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31 associated processes, such as greenhouse gas generation by crops or biomass production (Keane et al.,
32 2018; Tan et al., 2018). The knowledge of PAR helps one to estimate the plant's primary production
33 (Mercado et al., 2009). Like the global horizontal irradiance (GHI), PAR can also be partitioned into its
34 diffuse ($PAR_{diffuse}$) and direct (PAR_{direct}) components. This separation is of particular interest to many
35 applications, especially for PAR estimation over land with complex topography, where the surrounding
36 features can block the direct PAR component in an intricate and time-varying way (Olseth, 1997; Wang et
37 al., 2006). Another application of this diffuse–direct separation of PAR is to study PAR distribution in plant
38 canopies, where the diffuse light penetrates to a greater depth within the canopies than does the direct light
39 (Mariscal et al., 2004). Furthermore, the light-use efficiency of plant canopies increases under cloudy
40 conditions, due to the enhancement of the PAR diffuse component (Gu et al., 2002; Kanniah et al., 2012;
41 Mercado et al., 2009). Li et al. (2020) studied the influence of diffuse PAR radiation in a desert steppe
42 ecosystem and concluded that the maximum canopy photosynthesis was reached under cloudy skies.

43 The implication of PAR separation becomes more profound in the field of agrivoltaic systems. Agrivoltaic
44 system is a novel concept, which combines solar photovoltaic and agricultural activities on the same land
45 area. The agrivoltaic technology is an efficient, effective, and innovative solution to tackling land use
46 competition (Adeh et al., 2019). Nonetheless, one important concern of using such systems is that, for the
47 coexistence of solar energy and agricultural farming, crop yield must not go below tolerable limits. It is
48 known that shading generally decreases crop yield, and different crops behave differently under shading
49 conditions (Barron-Gafford et al., 2019). In open-field agrivoltaic systems, the amount of PAR reaching
50 the agricultural land is not homogeneously distributed. The solar modules installed in the system produce
51 variable levels of shading directly on the crops throughout a day and over a year. In these shaded areas, the
52 diffuse component of PAR plays a dominant role. Therefore, knowing the amount of diffuse and direct
53 PAR incident to a specific crop area beneath the agrivoltaic system implies a more accurate crop yield
54 estimation. Noticeably, the study by Campana et al. (2021) was among the first works in agrivoltaic systems
55 that introduced the concept of PAR separation for calculating crop yield; the topic of concern is an
56 exceedingly recent one.

57 Despite the relevance of PAR on crop growth, the scarcity of PAR measurements and the lack of a
58 worldwide measurement network with standardized quality control protocols (Ferrera-Cobos et al., 2020;
59 Mizoguchi et al., 2010; Niu et al., 2019; Wang et al., 2016) directly explain the limited number of studies
60 about PAR thus far as compared to, for example, to more extensive studies of GHI or diffuse horizontal
61 irradiance (DHI). The lack of measurements is even more pronounced for the diffuse component of PAR.
62 Therefore, as a work-around, several authors have suggested a variety of models to estimate the different
63 components of PAR. PAR components can be estimated using atmospheric radiative transfer models

64 (ARTM), e.g., Bird and Riordan (1986), Gueymard (1995) or Emde et al. (2016) and methods derived from
65 these, e.g., Wandji et al. (2019) or Thomas et al. (2019). However, since ARTM is associated with high
66 complexity and using it demands much knowledge in atmospheric sciences, most of the models are
67 empirical. These empirical models can derive the global component of PAR, and a limited number can also
68 derive diffuse PAR (e.g., Weiss and Norman, 1985, Kathilankal et al., 2014), from parameters commonly
69 measured at weather stations (e.g., Alados et al., 1996, Hu et al., 2007), from spectral band measurement
70 (e.g., Trisolino et al., 2016), and from satellite data (e.g., Su et al., 2007, Janjai et al., 2011, Hao et al.,
71 2019). The exhaustive review by Nwokolo et al. (2018) offers an overview of empirical models to estimate
72 the global PAR (i.e., $PAR_{\text{global}} = PAR_{\text{diffuse}} + PAR_{\text{direct}}$). It is worth mentioning that the correlation
73 between PAR and meteorological parameters is location-dependent (García-Rodríguez et al., 2020).

74 Several works have focused on the ratio PAR/GHI and its behaviour in different climate zones. According
75 to the review by Noriega et al. (2020), the ratio is typically higher during summer and lower during winter,
76 though exceptions to this rule have been highlighted by Yu and Guo (2016) or and Ma Lu et al. (2022).
77 Analysis of the PAR/GHI ratio under cloudless conditions shows a clear dependence on air mass (González
78 and Calbó, 2002). However, under all-sky conditions, the dependence of the ratio is unclear. Yu et al.
79 (2015), Akitsu et al. (2015), and, Ferrera-Cobos et al. (2020) observed a decrease in the ratio when the
80 clearness index (i.e., $k_t = GHI/E_{\text{ext}}$) increases. In contrary, Lozano et al. (2022) found no significant
81 dependence of the ratio on k_t . Most research studies admit that the PAR/GHI ratio is location- and season-
82 dependent (Hu et al., 2007; Jacovides et al., 2003; Li et al., 2010; Proutsos et al., 2022), therefore pointing
83 out the need to further investigate the behaviour of the ratio at more sites with different climates around the
84 globe.

85 The PAR_{diffuse} component is generally analysed by the PAR diffuse fraction (i.e., $k_{\text{PAR}} =$
86 $PAR_{\text{diffuse}}/PAR_{\text{global}}$). Several models have been proposed to obtain k_{PAR} and most of them are inspired by
87 GHI separation models, which estimate DHI from GHI, and their clearness index dependence (Gu et al.,
88 1999; Jacovides et al., 2010; Kathilankal et al., 2014; Oliphant and Stoy, 2018; Ren et al., 2018). Since the
89 spectral range of PAR is a portion of that of GHI, it is logically attractive to use just GHI separation models
90 to partition PAR_{global} . Indeed, the recent work by Ma Lu et al. (2022) applied and compared several GHI
91 separation models for separating PAR_{global} .

92 Generally, empirical models based on simple mathematical expressions reported in the literature are
93 applicable when the local conditions are similar to those used for calibrating the models. However, a limited
94 number of studies investigate the transferability of the models to other locations around the globe. For
95 instance, de Blas et al. (2022) analysed the accuracy of 21 semi-empirical models of PAR_{global} in seven

96 locations of the SURFRAD network in the United States that the authors claimed to be representative of a
97 large variety of weather conditions. All 21 models use a combination of easily retrievable parameters (see
98 section 3.1 for further details). The results show that calibrating the model parameters according to the
99 studied locations can slightly improve the estimation of the PAR components. But since the global
100 calibrated models already offer very satisfactory results, they should be chosen considering the availability
101 of the input variables at each specific location. These findings, nevertheless, cannot necessarily be applied
102 to high latitudes ($>49^{\circ}\text{N}$), and to northern European countries where agrivoltaics research in these territories
103 has expanded during the latest decade. There exists an overall lack of knowledge on the transferability and
104 performance of PAR separation models in high-latitude environments.

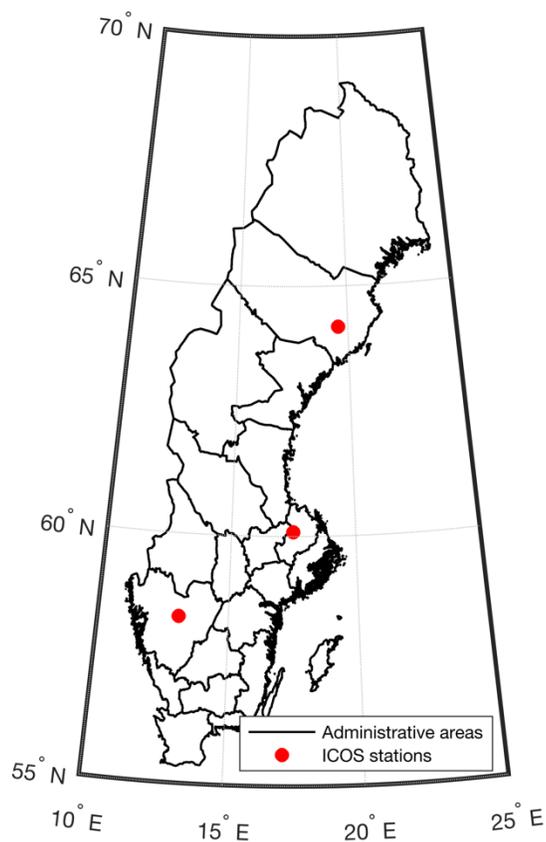
105 In this work, a new separation model to estimate $\text{PAR}_{\text{diffuse}}$ is proposed. It is derived from the original
106 YANG2 model (Yang and Boland, 2019), which is a GHI separation model, because of its high accuracy
107 demonstrated for both GHI and $\text{PAR}_{\text{global}}$ (Ma Lu et al., 2022). In addition, the newly proposed model is
108 based on atmospheric inputs conveniently retrievable from available databases, algorithms, and satellite-
109 derived data. The study is done for three locations in Sweden, considering an evident gap in PAR separation
110 model studies applied to northern latitudes exists. At the same time, an analysis of the seasonal trends and
111 variation of the different PAR components is provided for these colder climates. Additionally, the authors
112 are experimenting with agrivoltaic systems facilities based in Sweden. Hence, it is a priori opportune to
113 explore and be able to apply the developed model in situ in the upcoming future.

114 The remainder of the study is organized as follows: Section 2 presents the meteorological data used for
115 developing, calibrating, testing, and validating the model proposed in this study. Section 3 describes the
116 steps taken to develop the new separation model. Section 4 evaluates the performance of the proposed
117 model and discusses the results obtained for the selected sites. More specifically, an analysis of the
118 fluctuations in PAR components in these high-latitude locations is presented and discussed. Section 5
119 concludes the study.

120 2. Weather Data

121 The dataset used in this work for training and testing the proposed PAR separation model consists of
122 multiple-year measurements of $\text{PAR}_{\text{global}}$ and $\text{PAR}_{\text{diffuse}}$ among other variables from the Integrated Carbon
123 Observation System in Sweden (ICOS Sweden, 2022) network. Three locations in Sweden with available
124 measurements were selected, namely, Lanna, Degerö, and Norunda (Figure 1). The dataset spans three
125 years of data for each station with a time resolution of 30 min. Since the measurements of PAR from ICOS
126 stations are in units of flux density as a quantum process (PPFD), a conversion factor of $1 \text{ W/m}^2 \approx$
127 $4.6 \mu\text{mol/m}^2/\text{s}$ (Langhans et al., 1997) is applied whenever required. The data for each location is divided

128 into two subsets. On one hand, the training set consists of two years of data, which is used to fit the
 129 separation model parameters for the site. On the other hand, the validation (or testing) set consists of the
 130 remaining one year of data, which is used to test the fitted models with unseen data for the location of
 131 concern. The metadata of the sites considered in this study is tabulated in Table 1. A complete list of the
 132 available variables from these locations of the ICOS Sweden network is provided in section 3.2.



133
 134 Figure 1. Map of Sweden with the location of the ICOS Sweden network stations selected for the analysis. Map source: (GADM,
 135 2022).

136 Table 1. Information about the study locations and details of the data extracted from ICOS-Sweden network. The last column
 137 indicates the numbers of train/test samples (or data points) at each location after quality control (described in section 2.2).

Station	Latitude (°N)	Longitude (°E)	Elevation (m)	Data period	Samples training/ testing
Lanna	58°20'	13°06'	75	2016-2018	7062/ 3618
Degerö	64°18'	19°55'	270	2016- 2018	6993/ 2117
Norunda	60°05'	17°29'	46	2016-2018	5727/ 2676

138 2.1. Auxiliary Data

139 Besides PAR_{global} and PAR_{diffuse} , separation models often require as input several auxiliary variables,
140 which are often computable or can be accessed for general time periods and locations. These auxiliary
141 variables are described in this section. Firstly, the extraterrestrial radiation (E_{ext}) on a horizontal plane,
142 which is needed to compute k_t , is calculated as explained in Duffie & Beckman (2013). It is noted that the
143 computation of E_{ext} requires further a parameter known as the solar constant (SC), which is here in taken
144 to be $SC = 1361.1 \text{ W/m}^2$, following Gueymard (2018). Moreover, the Earth's orbit eccentricity correction
145 factor is used as per the definition by Spencer's equation (Spencer, J. W, 1971). Extraterrestrial PAR
146 (PAR_{ext}) is calculated analogously to E_{ext} , but with the approximated PAR solar constant, which is
147 $PAR_{\text{SC}} = 634.4 \text{ W/m}^2$ (Iqbal, 1983).

148 The solar zenith angle is calculated from the solar elevation and the latter is derived using the solar
149 positioning algorithm developed by Koblick (2021). Moreover, to account for the atmospheric refraction
150 effects, the model from the ESRL Global Monitoring Laboratory (US Department of Commerce, 2021) is
151 applied to correct the solar elevation angle. Both the clear-sky GHI (G_{cs}) and clear-sky PAR (PAR_{cs}) are
152 acquired from the Clouds and the Earth's Radiant Energy System (CERES) satellite-based observations
153 (Wielicki et al., 1996). Both satellite-derived diffuse fraction of GHI and the diffuse fraction of PAR are
154 obtained from the CERES SYN1deg Ed. 4.1 product (Doelling, 2017). CERES offers hourly satellite-
155 derived GHI, DHI, PAR total, and PAR diffuse from March 2000 till March 2022 with global coverage
156 with a $1^\circ \times 1^\circ$ resolution in both latitudes and longitudes. All satellite-derived data is downloaded to match
157 the spatial locations and temporal range of the measured ICOS data.

158 It should be noted that even though ICOS data has a temporal resolution of 30 min, due to the shortest time
159 step availability of CERES data, which has an hourly resolution at the midpoint, the remaining part of this
160 work (including both analysis and results) is performed with a 1-hour time step. In the present study, the
161 half-hourly time stamps of the 30-min data points from ICOS are taken (i.e., 9:30, 10:30, 11:30, and so
162 forth).

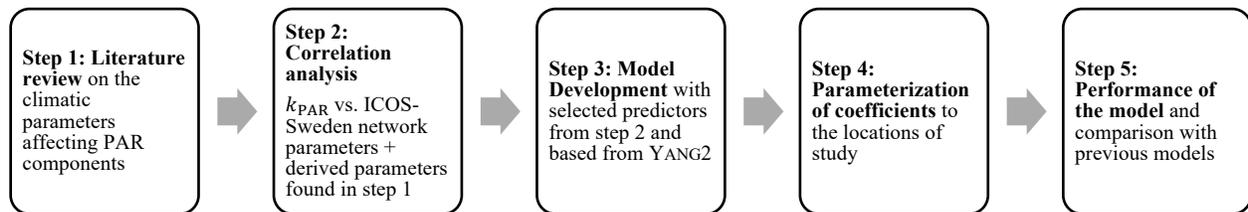
163 2.2. Quality control

164 Quality control (QC) constitutes an essential part of radiation modelling, with the goal of filtering and
165 eliminating spurious and erroneous data points. Since the observational data are to be used for the
166 determination of fitting parameters, validation, and performance comparison of the separation models, QC
167 must be applied to ensure that exclusively the highest-quality data points are selected. That said, there is no
168 ideal or universally accepted QC procedure for broadband irradiance data, not to mention PAR data. This
169 issue has been pointed out in the introduction section and in the previous work by Ma Lu et al. (2022). On
170 that account, the previously used QC procedure for PAR is adopted for this work as well. The reader is

171 referred to the previous publication for a detailed list of quality filters (Ma Lu et al., 2022). The only added
172 quality filter corresponds to albedo for which data with values greater than one was rejected.

173 3. Methodology

174 Before the development of the new PAR separation model is revealed, it is important to conduct a literature
175 review on studies that highlighted influencing atmospheric parameters to PAR components, and particularly
176 to $PAR_{diffuse}$. Subsequently, a correlation analysis is conducted between the diffuse fraction of PAR and
177 various meteorological variables available from the ICOS Sweden network to identify highly correlated
178 parameters. This correlation analysis was extended to include variables drawn from the first-step screening
179 and not provided by the ICOS Sweden network. Based on the correlation analysis, a selection of new
180 predictors results. The new separation model, which is based on the YANG2 model (Yang and Boland,
181 2019), is then developed in the third step, with consideration to those chosen predictors resulting from the
182 second step. Finally, the last two steps are to fit the model coefficients using the experimental data and to
183 evaluate the performance of the new model. A further comparison is made to two existing models, i.e.,
184 YANG2 and STARKE, which were previously found to be the most accurate. A graphical summary of the
185 methodology is presented in Figure 2.



186

187 Figure 2. Schematic diagram of the workflow applied in this work for the development of a new PAR separation model.

188 3.1 Literature review on the climatic parameters affecting PAR diffuse

189 An initial literature review has been performed to analyse which atmospheric variables are most influencing
190 the ecosystem production efficiency and, thus, the PAR components.

191 In the study by Li et al. (2020), in a desert steppe ecosystem, lower vapour pressure deficit ($VPD \leq 1$ kPa),
192 lower air temperature ($T_a < 20^\circ\text{C}$) and non-stressed water conditions were more favourable conditions
193 for enhanced ecosystem photosynthesis under cloudy skies ($k_t < 0.7$). $PAR_{diffuse}$ peaked when k_t was
194 around 0.5.

195 A work by Lu et al. (2022) using data from 40 sites around the globe has concluded that VPD and soil
196 moisture (SM) are significant variables in ecosystem production efficiency that should be fairly valued in
197 ecosystem modelling. For most of the studied sites, high VPD values cause positive changes in PAR while
198 low SM values cause negative changes in the fraction of PAR absorbed by the plants (fPAR). The study
199 underlines the influence of VPD on incident PAR in a multitude of locations. Yet none of those sites was
200 in northern latitudes.

201 A new method to estimate PAR values for clear-sky conditions used solar zenith angle, total column
202 contents of ozone (TOC) and water vapour (TWV), aerosol optical depth (AOD), vertical profiles of
203 temperature, pressure, density and volume mixing ratio of gases, elevation and ground albedo as inputs
204 (Wandji Nyamsi et al., 2019). The study emphasized that the errors in the suggested method were caused
205 by the overestimation of the input variables AOD and the assumption of constant PAR_{albedo} , suggesting
206 these two variables have a significant effect on the PAR under clear skies.

207 Recent work by de Blas et al. (2022), analysed PAR_{global} estimations at 1-min, hourly, and daily time steps
208 at seven sites from 21 models that use a combination of the following meteorological parameters: GHI,
209 clearness index, diffuse fraction, vapour pressure, relative optical air mass, precipitable water, solar zenith
210 angle, sky's brightness, and sky's clearness. The work further analysed the performance of the models for
211 different groups of sky conditions (clear to overcast) and found that for some models, the accuracy
212 worsened when applied to overcast skies.

213 Another recent work by Proutsos et al. (2022) studied the atmospheric factors affecting the PAR/GHI ratio
214 in a Mediterranean site. The authors concluded that the atmospheric water content (expressed by the degree
215 of cloudiness, actual water vapour, optical thickness, or dew point temperature) and the clearness index
216 were the most influential factors in the ratio. Air temperature and related meteorological variables (relative
217 humidity, vapour pressure deficit and saturation vapour pressure) were found to have no significant effect
218 on the ratio.

219 Regarding PAR diffuse estimations, the latest work by Lozano et al. (2022) found a clear dependence of
220 the k_{PAR} on the clearness index and total cloud cover (TCC) at a Mediterranean site. The authors proposed
221 a model to estimate k_{PAR} obtained through the first adaptation of the Boland-Ridley-Lauret (BRL) model
222 (Ridley et al., 2010) based on the clearness index, solar elevation angle, apparent solar time (AST), daily
223 clearness index and persistence index. When fitting the model to the studied site, the authors found that
224 AST and daily clearness index were insignificant and suggested these terms be removed from the model.

225 Kathilankal et al. (2014) developed a semi-parametric PAR separation model for the United States. It adapts
226 the BRL model using physically viable climate variables as predictors: relative humidity, PAR clearness

227 index, surface albedo and solar elevation angle. The proposed model takes a conditional approach, which
228 uses two logistic fits, one for clear-sky conditions and the other for cloudy conditions.

229 **3.2 Correlation Analysis**

230 The second step is to perform a correlation analysis between the observed diffuse fraction of PAR to each
231 of the meteorological variables available from the selected ICOS Sweden network stations (see Table 2),
232 and the derived potential variables that could benefit the separation model (see Table 3). The derived
233 variables are selected according to the literature review presented above. The correlation analysis of the
234 meteorological variables to k_{PAR} was used to rule out variables that are not important predictors of k_{PAR} ,
235 double-check variables considered already in well-known models, and to detect potential new significant
236 variables, either based on Pearson's correlation coefficient or visually from the scatterplot pattern
237 (Appendix A1). The analysis is performed with hourly data. Shape-preserving piecewise cubic interpolation
238 is used when the availability of the data has a larger timestep than hourly (e.g., AOD).

239
240

Table 2. List of variables available from the ICOS Sweden network Lanna, Degerö and Norunda stations for the period of 2016-2018 (“ICOS Sweden,” 2022).

Variable name	Variable description	Unit	Quantity kind
Swin_p	incoming shortwave radiation, pyranometer	W m-2	energy flux
Lwin	incoming long-wave radiation, net radiometer	W m-2	energy flux
Lwnet	net long-wave radiation, net radiometer	W m-2	energy flux
Lwout	outgoing long-wave radiation, net radiometer	W m-2	energy flux
NetRad	net radiation, net radiometer	W m-2	energy flux
PPFD_DIFF	photosynthetic photon flux density diffuse	µmol m-2 s-1	particle flux
PPFD_DIR	photosynthetic photon flux density direct	µmol m-2 s-1	particle flux
PPFD_IN	photosynthetic photon flux density incoming	µmol m-2 s-1	particle flux
PPFD_OUT	photosynthetic photon flux density outgoing	µmol m-2 s-1	particle flux
P	precipitation (total)	mm	length
Pa	air pressure	hPa	pressure
RH	relative humidity	%	portion
Sun	sunshine duration, sunshine sensor	1	portion
Swin_n	incoming shortwave radiation, net radiometer	W m-2	energy flux
Swnet_n	net shortwave radiation, net radiometer	W m-2	energy flux
Swout_n	outgoing shortwave radiation, net radiometer	W m-2	energy flux
T_canopy	target surface temperature	°C	temperature
Ta	air temperature	°C	temperature
H¹	sensible heat flux	W m-2	energy flux
LE¹	latent heat flux	W m-2	energy flux
Fc¹	carbon dioxide (CO2) flux	µmol m-2 s-1	particle flux
Fn2o^{1,2}	nitrous oxide (N2O) flux	µmol m-2 s-1	particle flux
Ustar¹	friction velocity	m s-1	velocity
WS¹	wind speed	m s-1	velocity
WD¹	wind direction	°	angle
NEE¹	net ecosystem exchange	µmol m-2 s-1	particle flux
LE_f¹	gap-filled latent heat flux	W m-2	energy flux
H_f¹	gap-filled sensible heat flux	W m-2	energy flux

¹Variables not available in Degerö ICOS station.
² Variable only available in Lanna ICOS station.

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242
243

Table 3. List of variables investigated through correlation analysis for the development of the new separation model of PAR based on the findings provided in the literature review.

Variable name	Variable description	Unit	Derived from	Source
Alb	surface albedo	-	$\text{Alb} = \frac{\text{Swout}_n}{\text{Swin}_n}$	
PAR_{albedo}	surface PAR albedo	-	$\text{PAR}_{\text{albedo}} = \frac{\text{PPFD}_{\text{OUT}}}{\text{PPFD}_{\text{IN}}}$	
AM	air mass	-	$\text{AM} = \frac{1}{\cos(Z) + 0.50572 * (6.07995 + (90 - Z))^{-1.6364}}$	(Kasten and Young, 1989)
e_a	actual vapor pressure	mbar	$e_a = e_s * \frac{\text{RH}}{100}$	(Technical Committee on Standardization of Reference Evapotranspiration, 2005)

e_s	saturation vapor pressure	mbar	$e_s = 6.1078 * \exp\left(\frac{17.27 * T_a}{T_a + 237.3}\right)$	(Technical Committee on Standardization of Reference Evapotranspiration, 2005)
VPD	vapor pressure deficit	mbar	$VPD = e_s - e_a$	
δ	optical thickness	-	$\delta = \ln\left(\frac{E_{ext}}{GHI}\right)$	(Proutsos et al., 2022)
δ_{PAR}	PAR optical thickness	-	$\delta_{PAR} = \ln\left(\frac{PAR_{ext}}{PAR_{global}}\right)$	
T_{dew}	dew point temperature	°C	$X = \frac{17.27 * T_a}{T_a + 237.3} + \ln\left(\frac{RH}{100}\right)$ $T_{dew} = \frac{237.3 * X}{17.27 - X}$	(Barenbrug, 1974)
AOD	total aerosol optical depth 550 nm	-	CAMS-AOD satellite-derived service provided by ECMWF. Time coverage from 2004-01-01 up to current day-2. Time step of 3 h.	(ECMWF, 2022)

244 From the analysis, the variables sunshine duration (Sun) and the PAR optical thickness (δ_{PAR}) were
245 identified as having a high degree of correlation to k_{PAR} . Despite sunshine duration exhibits the highest
246 correlation to k_{PAR} , the variable is not considered as a new predictor due to the difficulty in obtaining it. In
247 addition, even though exhibiting moderate to lower degrees of linear correlation (Table 4), the following
248 variables VPD, AOD, and PAR_{albedo} are selected. These variables have been shown to influence either
249 $PAR_{diffuse}$ or k_{PAR} based on the previous literature review (Li et al., 2020; Lu et al., 2022; Wandji Nyamsi
250 et al., 2019). The correlation coefficients of the selected variables are also shown for the datasets of Degerö
251 and Norunda sites (Table 4).

252 Table 4. Pearson's correlation coefficients of the variables added to the new separation model of PAR for the data of the three
253 studied locations described in Table 1.

Variable	Pearson's correlation coefficient to k_{PAR}		
	Lanna	Degerö	Norunda
δ_{PAR}	0.8111	0.7907	0.8453
VPD	-0.5478	-0.5329	-0.4373
AOD	0.2613	0.2353	0.2873
PAR_{albedo}	0.1360	0.1200	0.0902

254

255 3.3 Model development

256 YANG2 (Yang and Boland, 2019), which is a logistic form model, has been selected as the starting point for
257 developing the new PAR separation model. The logistic form is chosen based on the agreement in the
258 literature as yielding higher accuracy for both for separation models of GHI and separation models of PAR

259 in comparison with other functional shapes. Previous work by Ma Lu et al. (2022) showed that YANG2 and
 260 STARKE (Starke et al., 2018) were among the best-performing models to obtain $PAR_{diffuse}$ from PAR_{global} .
 261 It should be noted that both YANG2 (Eq.1) and STARKE (Eq.4) were originally developed for decomposing
 262 GHI. For this reason, Ma Lu et al. (2022) have applied the Spitters relationship (Eq.6) (Spitters et al., 1986)
 263 to expand the applicability of these models to PAR separation.

$$264 \quad k^{YANG2} = C + \frac{1 - C}{1 + e^{\beta_0 + \beta_1 k_t + \beta_2 AST + \beta_3 Z + \beta_4 \Delta k_{tc} + \beta_6 k^{(s)}}} + \beta_5 k_{de}, \quad (1)$$

$$265 \quad \Delta k_{tc} = k_{tc} - k_t = \frac{G_{cs}}{E_{ext}} - k_t, \quad (2)$$

$$266 \quad k_{de} = \max\left(0, 1 - \frac{G_{cs}}{GHI}\right), \quad (3)$$

$$267 \quad k^{STARKE} = \begin{cases} \frac{1}{1 + e^{\beta_7 + \beta_8 k_t + \beta_9 AST + \beta_{10} Z + \beta_{11} K_T + \beta_{12} \psi + \frac{\beta_{13} G_{cs}}{277.78}}}, & k_{CSI} \geq 1.05 \text{ and } k_t > 0.65; \\ \frac{1}{1 + e^{\beta_0 + \beta_1 k_t + \beta_2 AST + \beta_3 Z + \beta_4 K_T + \beta_5 \psi + \frac{\beta_6 G_{cs}}{277.78}}}, & \text{otherwise} \end{cases} \quad (4)$$

$$268 \quad K_T = \frac{\sum_{n=1}^{24} GHI_n}{\sum_{n=1}^{24} E_{ext_n}}, \quad (5)$$

$$269 \quad k_{PAR}^{model} = \frac{PAR_{diffuse}}{PAR_{global}} = \frac{[1 + 0.3(1 - (k^{model})^2)]k^{model}}{1 + (1 - (k^{model})^2) \cos^2(90 - \beta) \cos^3 \beta} \quad (6)$$

270 Briefly, k_t is the clearness index, G_{cs} is the clear-sky GHI [W/m^2], Z is the solar zenith angle [$^\circ$], AST is the
 271 apparent solar time [h], E_{ext} is the extraterrestrial radiation [W/m^2], $k^{(s)}$ is the satellite-derived diffuse
 272 fraction, K_T is the daily clearness index, the ψ predictor is the three-point moving average of clearness
 273 index, k_{CSI} is the clear-sky index, and β is the solar elevation angle [$^\circ$].

274 In the present work, the model form of YANG2 is taken as a basis but with all the predictors adapted into
 275 PAR (i.e., k_t to k_{t_PAR} , Δk_{tc} to Δk_{tc_PAR}). The following model, hereafter called CLY (i.e., an abbreviation
 276 of the main developers' family names in alphabetical order), is proposed by including the four new relevant
 277 variables found in the previous subsection (see Table 4).

$$278 \quad k_{PAR}^{CLY} = C + \frac{1 - C}{1 + e^{\beta_0 + \beta_1 k_{t_PAR} + \beta_2 AST + \beta_3 Z + \beta_4 \Delta k_{tc_PAR} + \beta_5 PAR_{albedo} + \beta_6 \delta_{PAR} + \beta_7 AOD + \beta_8 VPD + \beta_{10} k_{PAR}^{(s)}}} + \beta_9 k_{de_PAR}, \quad (7)$$

279 where,

$$280 \quad k_{t_PAR} = \frac{PAR_{total}}{PAR_{ext}} \quad (8)$$

281
$$k_{de_PAR} = \max\left(0, 1 - \frac{PAR_{cs}}{PAR}\right) \quad (9)$$

282
$$\Delta k_{tc_PAR} = \frac{PAR_{cs}}{PAR_{ext}} - k_{t_PAR}, \quad (10)$$

283 Similarly, k_{t_PAR} is the PAR clearness index, PAR_{cs} is clear sky PAR [W/m^2], PAR_{ext} is the extra-terrestrial
 284 PAR [W/m^2] and $k_{PAR}^{(s)}$ is the satellite-based diffuse fraction of PAR. The use of satellite-derived predictors
 285 was introduced by Yang and Boland (2019). Satellite-based predictors are efficient in illustrating the low-
 286 frequency variability of the diffuse component since they are based on physical models.

287 The proposed model (CLY) is evaluated for the selected locations in Section 2 and the performance
 288 compared to the original YANG2 and STARKE along with all PAR version of YANG2 and STARKE, which
 289 are annotated with an asterisk, i.e., YANG2* (Eq.11) and STARKE* (Eq.12):

290
$$k_{PAR}^{YANG2*} = C + \frac{1 - C}{1 + e^{\beta_0 + \beta_1 k_{t_PAR} + \beta_2 AST + \beta_3 Z + \beta_4 \Delta k_{tc_PAR} + \beta_6 k_{PAR}^{(s)}}} + \beta_5 k_{de_PAR}, \quad (11)$$

291
$$k_{PAR}^{STARKE*} = \begin{cases} \frac{1}{1 + e^{\beta_7 + \beta_8 k_{t_PAR} + \beta_9 AST + \beta_{10} Z + \beta_{11} K_{T_PAR} + \beta_{12} \psi_{PAR} + \frac{\beta_{13} PAR_{cs}}{277.78}}}, & k_{CSL_PAR} \geq 1.05 \text{ and } k_{t_PAR} > 0.65; \\ \frac{1}{1 + e^{\beta_0 + \beta_1 k_{t_PAR} + \beta_2 AST + \beta_3 Z + \beta_4 K_{T_PAR} + \beta_5 \psi_{PAR} + \frac{\beta_6 PAR_{cs}}{277.78}}}, & \text{otherwise} \end{cases} \quad (12)$$

292 3.4 Statistical indicators for the assessment of the models

293 The performance of the proposed CLY model (Eq.7) is evaluated at the different sites introduced in Section
 294 2 using several popular error metrics. The results are then compared to the performances of the original and
 295 reparametrized YANG2 and STARKE models applied to PAR, as described in the work by Ma Lu et al.
 296 (2022). In addition, the proposed CLY model is compared to the reparametrized YANG2* and STARKE*
 297 models with all predictors adapted to PAR, as noted in section 3.3.

298 The error metrics selected in this work are the ones utilized by Ma Lu et al. (2022): the normalized mean
 299 bias error (nMBE), the normalized root mean square error (nRMSE), and the coefficient of determination
 300 (R^2). The observations of k_{PAR} are derived from the measurements of PAR_{global} and $PAR_{diffuse}$ at the
 301 studied ICOS stations. The predictions are the k_{PAR}^{model} calculated from the models.

302 3.5 Reparameterization of coefficients

303 The training datasets listed in Table 1 for the three locations under study are utilized to estimate locally
 304 fitted coefficients for each of the analysed models. To achieve this, a nonlinear optimization solver-based
 305 approach is employed, as detailed in Ma Lu et al. (2022). In this study, the root mean square error (RMSE)

306 of k_{PAR} is selected as the target function to be minimized. This choice aligns with the statistical concept of
 307 consistency (Gneiting, 2011), as one of the main evaluation metrics is the nRMSE (section 3.4). The concept
 308 of consistency has been emphasized in previous research for the calibration and evaluation of point forecasts
 309 (Yang et al., 2020; Yang and Kleissl, 2022).

310 4. Results and discussion

311 4.1 CLY separation model performance

312 The proposed CLY satellite-augmented model for estimating diffuse PAR is evaluated alongside four other
 313 models at the three studied locations, using hourly data. These include the original YANG2 and STARKE
 314 GHI decomposition models with Spitters amendment for PAR, as well as the modified versions of YANG2*
 315 (Eq.11) and STARKE* (Eq.12) presented in Section 3.3. Table 5 presents the models' performances.

316 Table 5. The nRMSE [%], nMBE [%] and R^2 in predicted hourly diffuse PAR from of the proposed PAR separation model, CLY,
 317 compared to the other 2 models with different versions, Yang2* and Starke* with PAR predictors, and the original Yang2 and
 318 Starke applied to PAR as presented in Ma Lu et al. (2022). Locally fitted coefficients (using training data over 2-years, period
 319 2016-2017) and validated (using testing data over 1-year, period 2018) at 3 ICOS-Sweden stations (Lanna, Degerö, Norunda). The
 320 errors are computed between the predicted and measured hourly PAR diffuse fraction values. Boldface denotes the best-performing
 321 model in a row.

Station	CLY	STARKE*	YANG2*	STARKE	YANG2
nRMSE [%]					
Lanna	12.86	15.29	13.92	15.00	13.71
Degerö	16.64	18.43	16.24	20.00	17.99
Norunda	14.89	16.03	15.12	17.18	15.96
nMBE [%]					
Lanna	0.8	-1.14	0.42	-1.87	-1.06
Degerö	-0.37	-1.89	0.14	-1.86	-0.01
Norunda	-1.85	-0.89	-1.13	-1.80	-1.74
R²					
Lanna	0.94	0.91	0.93	0.92	0.93
Degerö	0.91	0.89	0.91	0.87	0.90
Norunda	0.92	0.91	0.92	0.90	0.91

322 For two of the investigated locations with latitudes higher than 58° N, the CLY model's accuracy in terms
 323 of nRMSE and R^2 is superior to the other models. However, for Degerö, CLY performs slightly worse than
 324 YANG2 in terms of nRMSE. The added predictors to the YANG2 model, namely optical thickness, vapour
 325 pressure deficit, aerosol optical depth, and PAR albedo, can better represent the scattered processes in the
 326 atmosphere compared to the other models (Figure 3). Particularly, the CLY model outperforms other
 327 models (YANG2* and YANG2) in predicting clear-sky conditions when k_t values are between 0.7 and 0.8,
 328 and k_{PAR} values are lower than 0.2. An exception is observed in Degerö, where the behaviour is rather
 329 similar to YANG2. When compared to STARKE* and STARKE models, the CLY model estimates the shape

330 of the envelope and the larger spread of data during partly cloudy conditions ($0.3 < k_t < 0.7$) in a superior
 331 way. The CLY model coefficients are presented in Table 6.

332 Table 6. Model coefficients of the proposed CLY PAR separation model fitted to the 3 ICOS stations in Sweden with hourly time
 333 step (Lanna, Degerö, and Norunda) each with 2 years of data for the period 2016-2017.

Station	C	β_0	β_1	β_2	β_3	β_4	β_5	β_6	β_7	β_8	β_9	β_{10}
Lanna	0.1084	4.6754	-0.0111	-0.0683	0.0001	-2.8288	-1.4563	-2.5791	-1.2877	0.0245	0.2520	-2.1574
Norunda	0.0933	4.3594	-0.9755	-0.0590	0.0105	-2.0031	-0.4034	-3.2240	-1.0686	0.0427	0.2883	-1.8674
Degerö	0.1514	3.2011	1.5564	-0.0486	0.0079	-2.1034	-0.3636	-2.5949	-0.5795	0.0475	0.4814	-2.1945

334
 335
 336 Results show that predicting k_{PAR} is more accurate when using PAR-derived predictors (YANG2*) than
 337 GHI-derived predictors added to the Spitters relationship for YANG2. However, this trend is not observed
 338 for STARKE*, which does not seem to outperform STARKE. The reason could be due to the k_{CSI_PAR} and
 339 k_{t_PAR} constraint values not being recomputed for the PAR-derived predictors.

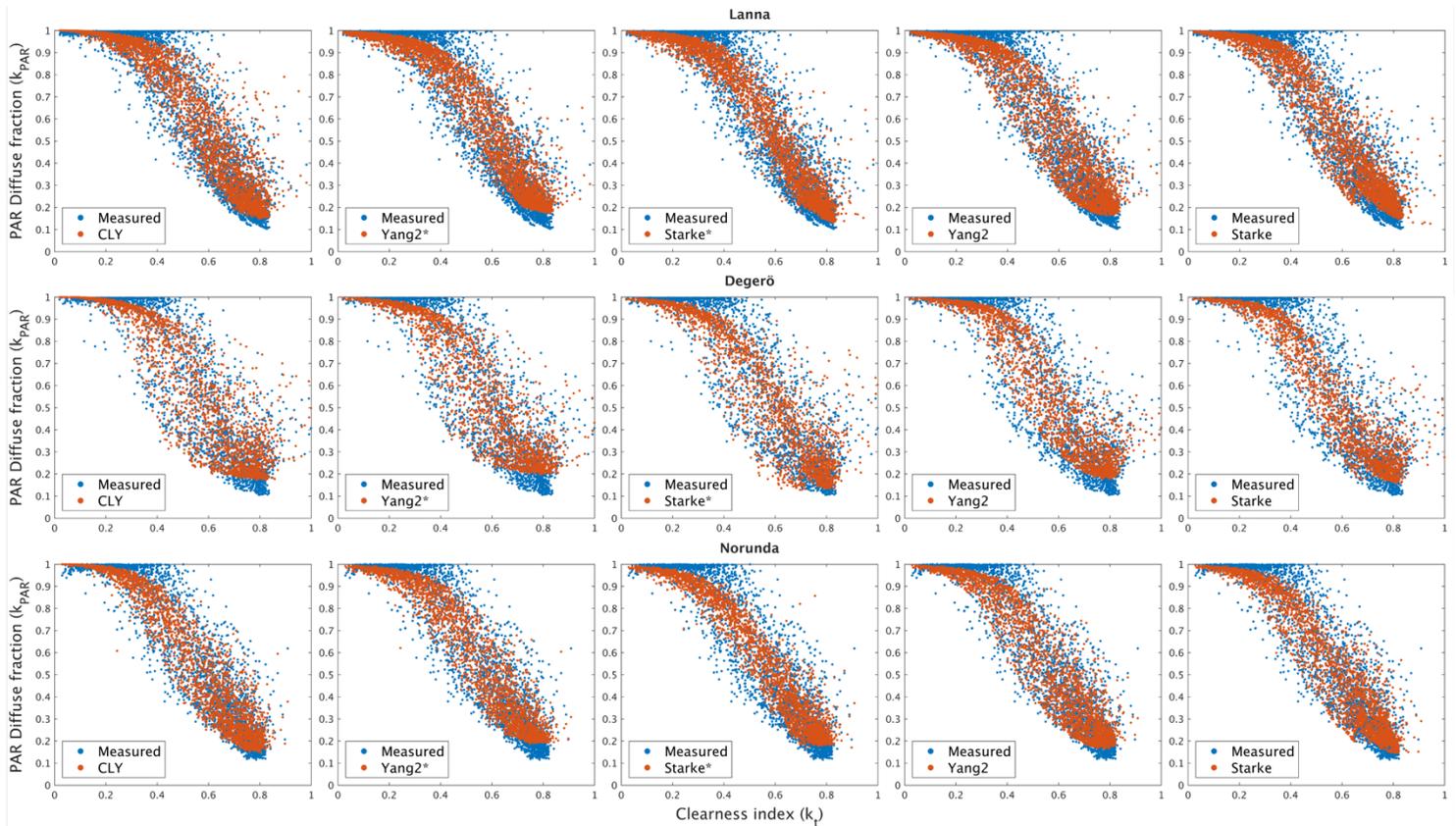


Figure 3. PAR diffuse fraction measured data plotted against the clearness index for the studied locations: Lanna (top row), Degerö (middle row), and Norunda (bottom row). The estimated results from the proposed PAR separation model CLY, YANG2* and STARKE* with PAR predictors, and the original YANG2 and STARKE applied to PAR are overlaid. The total number of data points in each plot refers to the testing data sample listed in Table 1.

341 Despite having several predictors, the proposed model is potentially widely applicable thanks to the
342 availability of satellite data. Complete data needed for the model can be easily retrieved from satellite-
343 derived data products or calculated using mathematical relationships from commonly available weather
344 data (Section 3).

345 This work introduces a new PAR separation model developed and applied for northern latitudes to support
346 the development of the agrivoltaic sector. In particular, to accurately determine the $PAR_{diffuse}$ reaching the
347 crops beneath an agrivoltaic system from PAR_{global} measurements. The proposed model could be applied to
348 other latitudes and climates to evaluate its worldwide performance, although this is beyond the scope of the
349 present study.

350

351 **4.2 PAR_{global} , PAR_{direct} , and $PAR_{diffuse}$ variation at northern latitudes**

352 As highlighted in the introduction, studies on the behavior of PAR components for high-latitude regions
353 are lacking. The annual evolution for PAR_{global} , PAR_{direct} and $PAR_{diffuse}$ measured at the ICOS sites at the
354 three study locations is depicted in Figure 4. The monthly distribution of PAR_{global} shows a clear cycle,
355 with maximum mean and median values around May and July for all locations, and the lowest values during
356 winter. This seasonality trend is similarly observed in other studies for the northern hemisphere, such as the
357 study by Lozano et al. (2022) in Granada, Spain (37.16° N, 3.61° W). However, the magnitude of PAR_{global}
358 differs. In the Mediterranean location, the PAR_{global} during the warmest months exhibited values higher
359 than 250 W/m², while the maximum in the Scandinavian sites was around 150 W/m² (with the exception of
360 2018, which reached average values slightly below 200 W/m²). Moreover, the Lanna station, located at the
361 southernmost latitude, received on average 30.64% more annual PAR_{global} radiation than Degerö, located
362 6° further north, for the period 2016-2017.

363 The seasonal pattern of the PAR_{direct} component exhibits the highest variation and distribution. The direct
364 component is clearly influenced by the Sun's position and the intensity of the incoming light. It is worth
365 noting that 2016 and 2017 present similar distributions, while 2018 shows a significantly different
366 distribution. The atypical behaviour is aligned with the drought that occurred in Sweden in 2018. The country
367 experienced an earlier onset of summer at the start of May, which lasted throughout the summer months,
368 with short interruptions mainly in June (Wilcke et al., 2020). For the three locations investigated, the
369 average PAR_{direct} value was 57.48% higher in May 2018 than in the previous two years. The increased
370 solar irradiance in 2018 was caused by the anomalous presence of clear sky conditions (Räisänen, 2019;
371 Sinclair et al., 2019).

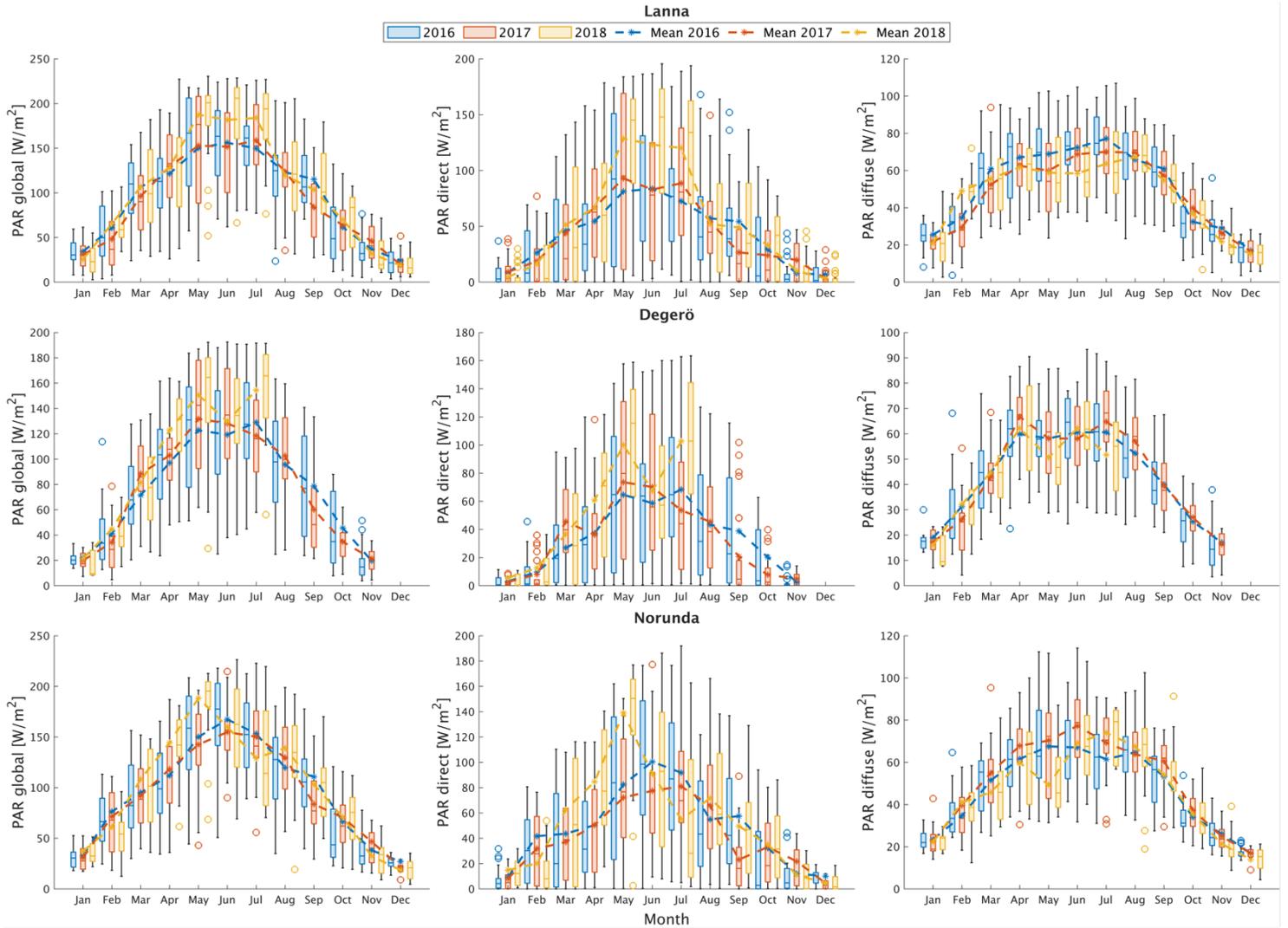


Figure 4. Monthly variation statistics for PAR_{global} , PAR_{direct} and $PAR_{diffuse}$ during the period 2016 – 2018 at the studied ICOS Sweden network stations: Lanna (top), Degerö (middle), and Norunda (bottom). For each box, central lines are the median, and upper and lower limits represent the percentiles 75th and 25th respectively. The limits of the segments represent the minimum and the maximum daily average values. The stars are the mean monthly values.

372 The monthly variation observed in Figure 4 for $PAR_{diffuse}$ is less pronounced than for PAR_{direct} or
373 PAR_{global} . The main reason is the high complexity of the scattering processes involved in the diffuse
374 component, affected by the presence of clouds, aerosols, surface albedo, and altitude. For the investigated
375 sites, the trend is similar for all the years with a slight alteration in 2018 due to decreased amount of clouds,
376 which brought overall lower values of $PAR_{diffuse}$. The annual mean $PAR_{diffuse}$ value for the locations
377 studied was 46.65 W/m^2 , marginally lower (59 W/m^2) than the one reported by Lozano et al. (2022) in
378 Granada (Spain) 2008-2018 and higher (35 W/m^2) than the one reported by Trisolino et al. (2018) in
379 Lampedusa (Italy) 2002-2016. Since there are scarce studies about PAR trends, the comparison is made to
380 available studies in these Southern European locations. It is interesting to observe that the $PAR_{diffuse}$ is
381 rather similar regardless of whether it is in the north or south of Europe.

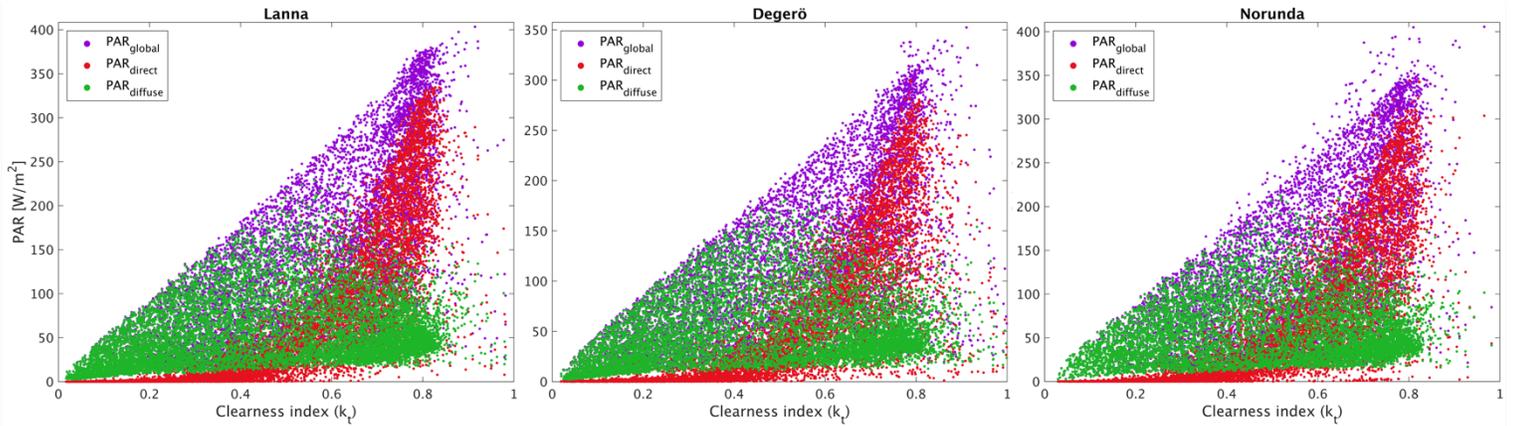


Figure 5. Scatterplots between the clearness index and PAR_{global} , PAR_{direct} and $PAR_{diffuse}$ for the period 2016 – 2018 at the studied ICOS Sweden network stations: Lanna (left), Degerö (middle), and Norunda (right). Hourly values at midpoint are used.

382 Figure 5 presents the effect of cloudiness on PAR_{global} , PAR_{direct} and $PAR_{diffuse}$ measurements for the
 383 investigated sites during the studied period. The upper envelope of PAR_{global} increases linearly with the
 384 clearness index. When the clearness index is low, $k_t < 0.3$, corresponding to thick cloud conditions (Chen
 385 et al., 2009), $PAR_{diffuse}$ makes the primary contribution to PAR_{global} . $PAR_{diffuse}$ increases with increasing
 386 k_t , peaking at values of k_t around 0.5 under thin cloud conditions ($0.3 \leq k_t < 0.7$), and then decreases
 387 towards clear-sky conditions, at high values of k_t . PAR_{direct} increases exponentially when the sky starts
 388 having clearer conditions ($k_t > 0.3$), and rapidly increases after the $PAR_{diffuse}$ decreases ($k_t > 0.7$). At
 389 high values of k_t , PAR_{direct} significantly contributes to the PAR_{global} . These trends are consistent across
 390 the three studied sites and align with Li et al.'s (2020) findings in a desert environment in the northern
 391 hemisphere. However, the magnitude of the PAR_{global} in this study are halved due to the climate and
 392 latitude characteristics.

393 The analysis demonstrates that the seasonality variation of PAR components and the relationship with
 394 cloudiness in high latitudes is similar to mid-latitudes in the northern hemisphere. However, the magnitude
 395 of the PAR components decreases as the location moves further north. This decrease is particularly
 396 noticeable for the PAR_{direct} component due to the distinct course of the solar zenith angle throughout the
 397 year resulting in reduced solar radiation. The $PAR_{diffuse}$ component, on the other hand, appears to have
 398 minor variability across seasons and locations, indicating that it is less influenced by incoming solar
 399 irradiance and more likely to be affected by sky conditions and atmospheric aerosols content.

400

401 5. Conclusions

402 The issue of conflicting land use between agricultural activities and ground-mounted solar photovoltaic
 403 power plants has become increasingly prevalent in recent years, and agrivoltaic systems offer a potential

404 solution to this problem. Accurately estimating $PAR_{diffuse}$ is crucial for analysing agrivoltaic systems, as
405 crops situated underneath do not receive PAR_{global} in a uniform manner, as is the case in open-field
406 conditions. Instead, they receive a non-uniform combination of $PAR_{diffuse}$ and PAR_{direct} due to the shading
407 produced by the PV system, with shaded areas receiving a greater proportion of $PAR_{diffuse}$. This shading
408 typically reduces crop yields, making accurate calculation of $PAR_{diffuse}$ essential for more precise crop
409 yield predictions.

410 To this end, the present study proposes a new separation model called CLY, which calculates $PAR_{diffuse}$
411 from PAR_{global} using the YANG2 decomposition model for GHI (Yang and Boland, 2019) as a basis. The
412 CLY model leverages atmospheric data from satellites, which are widely available worldwide, and utilizes
413 predictors selected through correlation analysis and previous literature findings.

414 The accuracy of the model has been compared to that of two previously identified best GHI separation
415 models for PAR (Ma Lu et al., 2022), namely YANG2 and STARKE, across different locations in Sweden.
416 Results show that the CLY model outperforms both the YANG2 and STARKE models in two of the three
417 locations studied. Across all locations, the model achieves R^2 values above 0.91, with an improvement of
418 up to 1% in both R^2 and nRMSE compared to the previously identified most accurate model, YANG2.
419 Although the CLY model has only been validated in three locations at high northern latitude ($>58^\circ N$),
420 primarily chosen because of the lack of studies in these regions, it could be subject to further studies to
421 investigate its applicability and performance in other climates and at other temporal resolutions.

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435 CRediT author statement

436 **Silvia Ma Lu**: Conceptualization, Methodology, Data Curation, Formal analysis, Validation, Visualization,
437 Writing - Original Draft, Writing - Review & Editing; **Dazhi Yang**: Writing - Review & Editing; **Martha**
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439 Writing - Review & Editing; **Anders Avelin**: Writing - Review & Editing; **Pietro Elia Campana**: Funding
440 acquisition, Conceptualization, Methodology, Writing - Review & Editing.

441 Appendix

442 A1. Correlation analysis

443 Table A - 1 displays the Pearson correlation coefficient (r) between the variables presented in Tables 2 and
444 3 and the diffuse fraction of PAR. The data used for the correlation analysis and the results in Section 4
445 were retrieved from the ICOS Sweden Lanna, Degerö and Norunda stations for the years 2016, 2017, and
446 2018. The Pearson correlation method is the most common way of measuring linear correlations. It assigns
447 a value between -1 and 1, where 0 is no correlation, 1 is total positive correlation, and -1 is total negative
448 correlation (Nettleton, 2014).

449 The variables that showed high and medium degrees of correlation but were not considered as new
450 predictors for the proposed model could be associated with various reasons. Firstly, these variables were
451 already accounted for or implicitly accounted for in existent significant variables in the YANG2 model (i.e.,
452 Swin, PPF_D_DIR). Secondly, these variables were accounted for in the calculation of another variable that
453 was considered as a new predictor (e.g., PPF_D_IN and PPF_D_OUT were used to calculate PAR_{albedo}.
454 Thirdly, these variables are difficult to obtain if they are not measured on-site (e.g., sunshine duration, H,
455 LE). Lastly, these variables showed minor or scarce influence on PAR (e.g., RH).

456

458 Table A - 1. Pearson's correlation coefficient values of the variables investigated to the diffuse fraction of PAR under the studied
 459 locations. Highlighted in bold are the variables chosen as predictors for the proposed model.

Variable	Pearson's correlation coefficient (r) to k_{PAR}		
	Lanna	Degerö	Norunda
Swin_p	-0.7617	-0.7233	-0.6872
Lwin	0.2899	0.3843	0.3340
Lwnet	0.8972	0.8767	0.8661
Lwout	-0.4335	-0.3066	-0.2315
NetRad	-0.6602	-0.5460	-0.6111
PPFD_DIFF	-0.0301	0.0650	0.1118
PPFD_DIR	-0.7510	-0.7155	-0.6838
PPFD_IN	-0.7408	-0.7047	-0.6741
PPFD_OUT	-0.3616	-0.2312	-0.6310
P	0.0942	0.1244	0.0869
Pa	-0.1793	-0.2304	-0.1282
RH	0.6266	0.6043	0.4886
Sun	-0.9513	-0.9466	-0.9382
Swin_n	-0.7596	-0.7173	-0.6964
Swnet_n	-0.7475	-0.6486	-0.6905
Swout_n	-0.7015	-0.4157	-0.7419
T_canopy	-0.3718	-0.2865	-0.2294
Ta	-0.3506	-0.2988	-0.2192
H	-0.5964	-	-0.5267
LE	-0.5565	-	-0.3376
Fc	0.3215	-	0.2193
Fn2o	0.1408	-	-
Ustar	-0.1389	-	-0.1520
WS	-0.0177	-	-0.0627
WD	-0.0990	-	-0.0785
NEE	0.3295	-	0.2540
LE_f	-0.5887	-	-0.3575
H_f	-0.6045	-	-0.5378
Alb	0.0437	0.0540	-0.1926
PAR_{albedo}	0.1360	0.1200	0.0902
AM	0.3242	0.2390	0.1960
e_a	0.0524	0.0873	0.1410
e_s	-0.3755	-0.3407	-0.2448
VPD	-0.5478	-0.5329	-0.4373
δ	0.7984	0.7732	0.8153
δ_{PAR}	0.8111	0.7907	0.8453
T_{dew}	0.0300	0.0745	0.1160
AOD	0.2613	0.2353	0.2873

461 In addition, to aid readers with visualization, the scatterplots of the analysed variables are presented in
 462 figures A-1, A-2. The scatterplots provide a graphical representation of the relationships between the
 463 variables, allowing readers to observe any trends or patterns that may exist besides linear relationships.

464

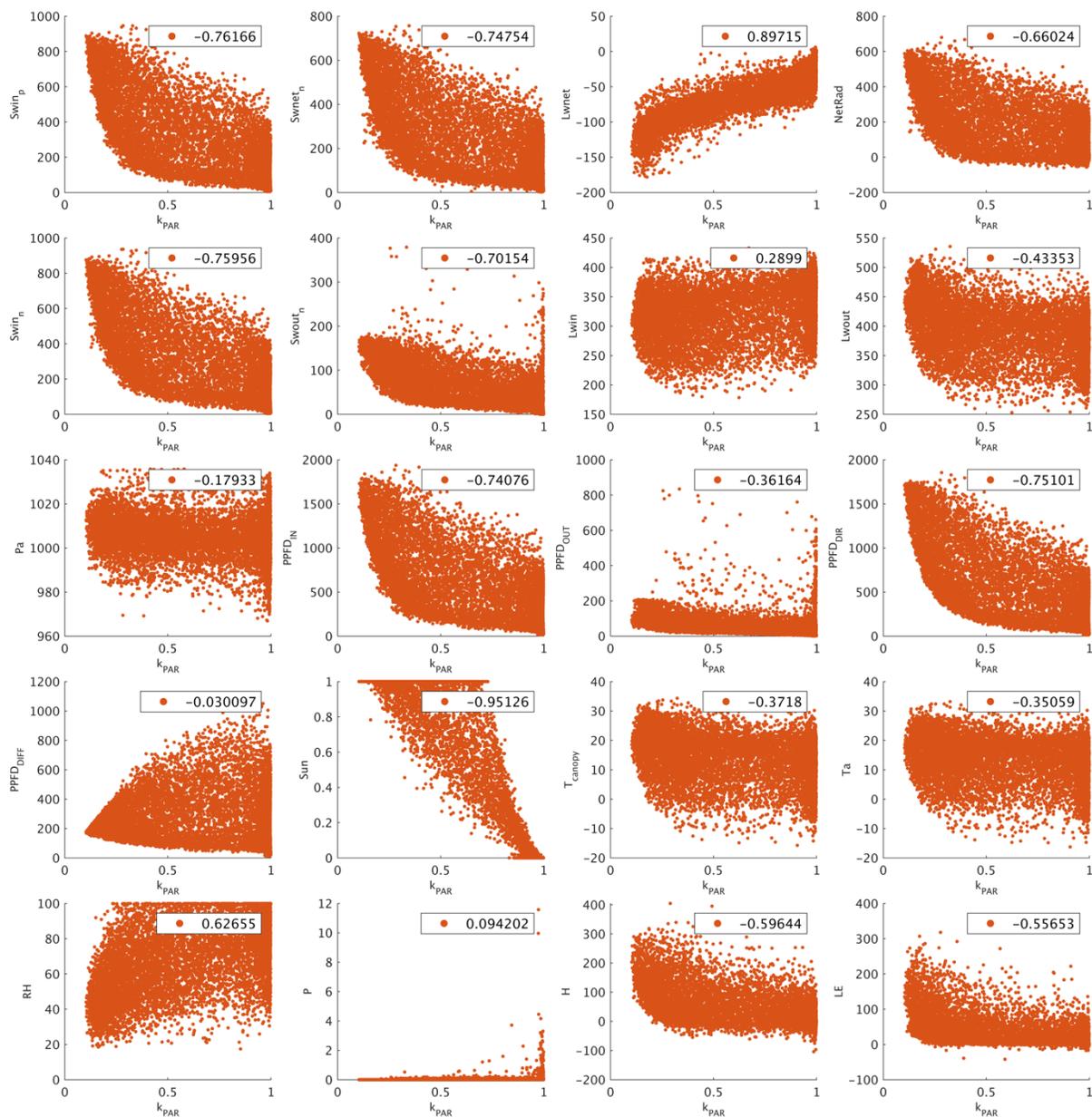
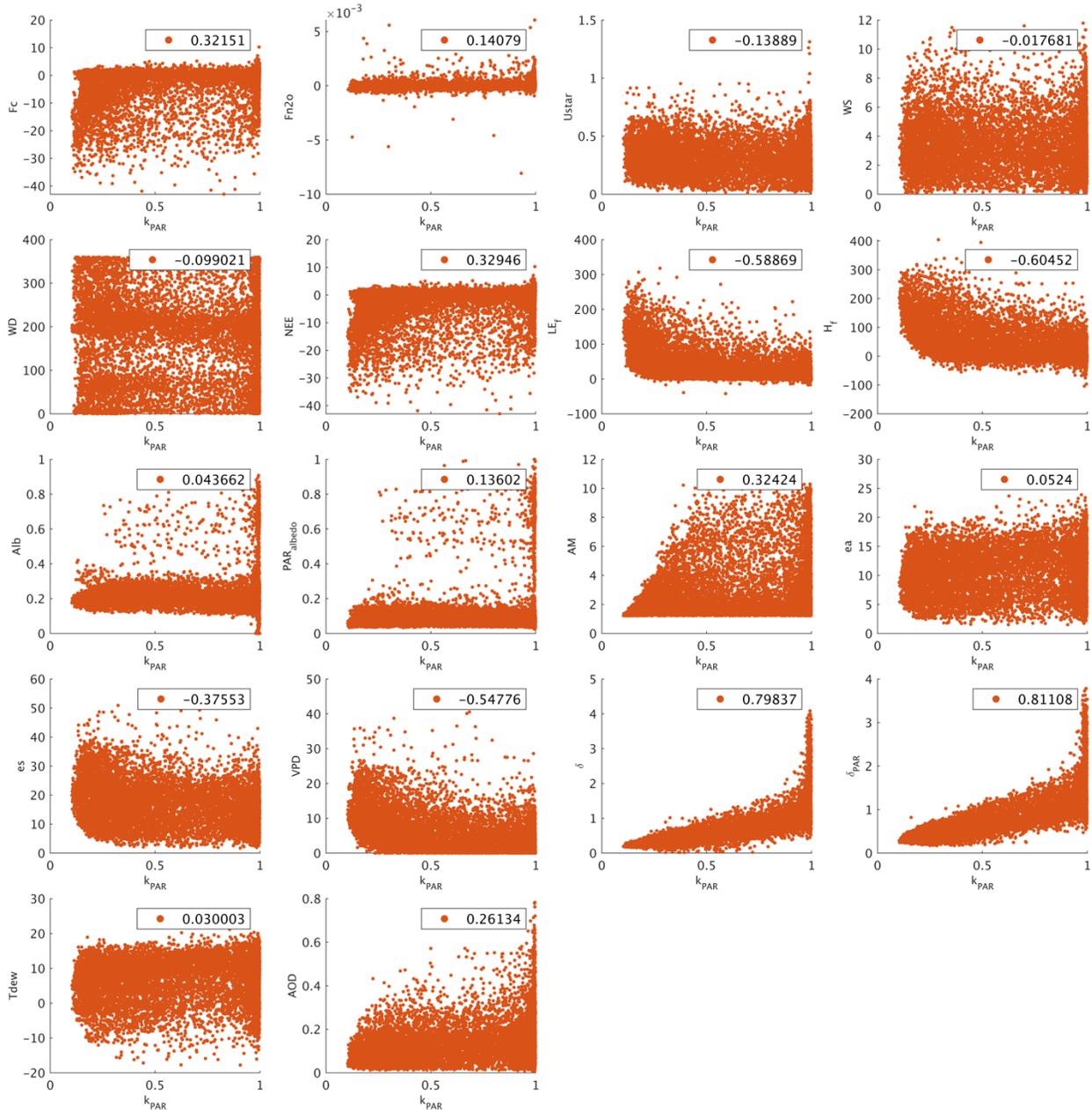


Figure A- 1. Scatterplots of the diffuse fraction of PAR (x-axis) to the variables from Table 2 (y-axis) from Lanna ICOS-Sweden network station for the period 2016-2018. Pearson's correlation coefficient value is displayed for each plot.

465



466

Figure A- 2. (Continuation) Scatterplots of the diffuse fraction of PAR (x-axis) to the variables from Tables 2 and 3 (y-axis) from Lanna ICOS-Sweden network station for the period 2016–2018. Pearson’s correlation coefficient value is displayed for each plot.

467

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