

1 **Stratigraphic reservoir compartmentalization: causes, recognition, and implications for the**
2 **geological storage of carbon dioxide**

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8
9 **Abstract**

10 The impact of carbon capture and storage (CCS) in mitigating anthropogenic emissions of greenhouse
11 gases is potentially great, but its success is strongly dependent on identifying suitable geological
12 storage sites. One of the key uncertainties in this regard is the degree of compartmentalization of the
13 target storage horizon. Many studies have examined reservoir compartmentalization in oil and gas
14 fields and its implications to hydrocarbon production, but few have looked at compartmentalization
15 from the perspective of CCS storativity. Using case studies from the oil and gas industry, this paper
16 examines the predictability of reservoir compartmentalization and the techniques used to assess it. It
17 then focusses on stratigraphic compartmentalization in saline aquifers, a major category of potential
18 carbon storage sites, examining in particular the Triassic Bunter Sandstone Formation of the UK
19 North Sea. It is clear that the analysis of stratigraphic, seismic and fluid compositional data can
20 provide early evidence of reservoir compartmentalization. The available data indicate that the Bunter
21 Sandstone has a low risk of stratigraphic and structural compartmentalization. However, the presence
22 of compartments in oil and gas fields is often confirmed only once well production data are available.
23 As such, though the broad likelihood of compartmentalization in a CCS site can be forecasted,
24 understanding its exact nature may be dependent on injection data.

25
26 **Introduction**

27 Carbon capture and storage (CCS) is the extraction, transport and geological trapping of carbon
28 dioxide produced during the combustion of fossil fuels. It has been identified as a mechanism for
29 generating cleaner energy from coal, oil and gas, and reducing pollution emitted by major industrial
30 producers. The process typically involves the chemical removal of CO₂ from fossil fuel emissions, its
31 compression into a supercritical fluid state, transport by underground pipeline, and injection into a
32 permeable rock unit for permanent storage by mineralization.

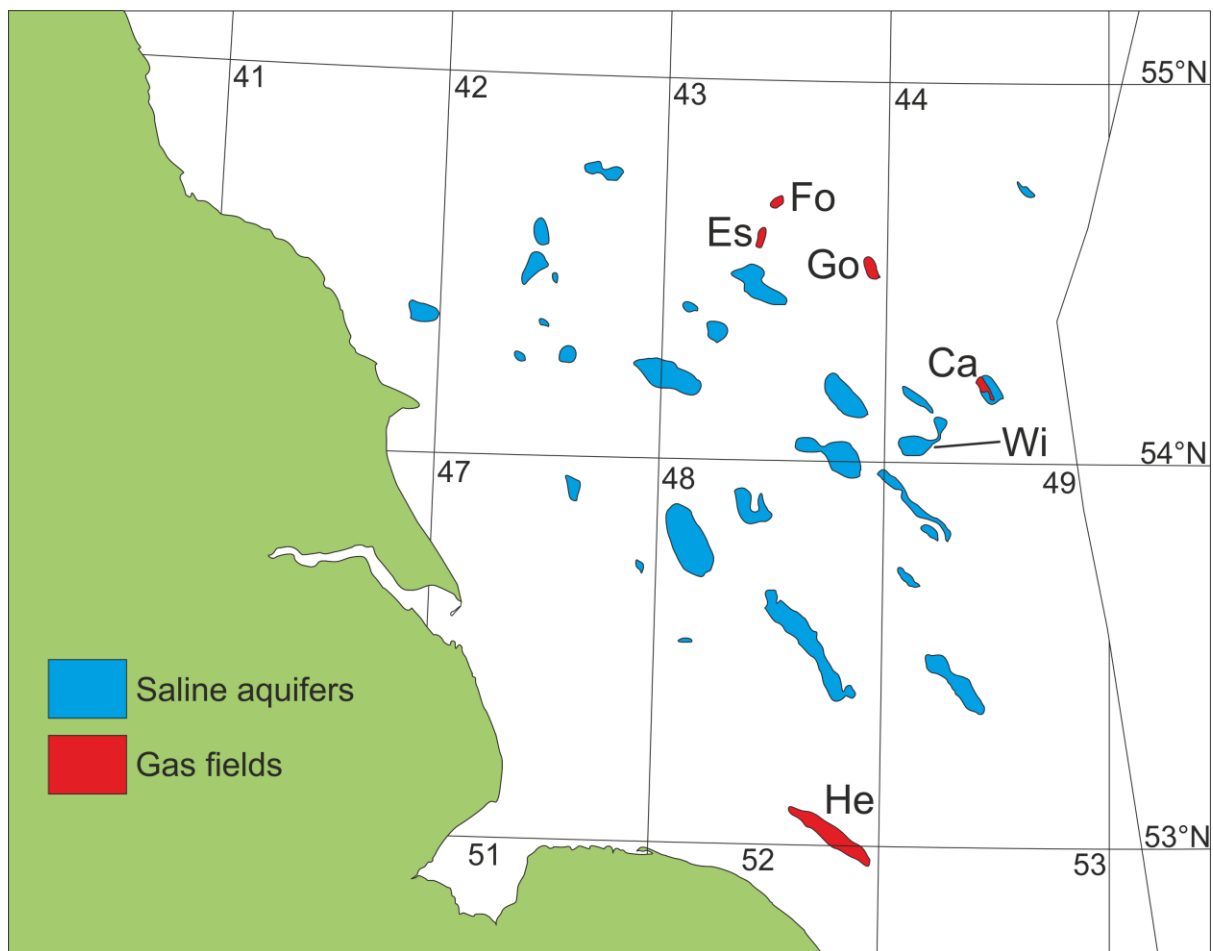
33 One of the fundamentals of CCS viability is the identification of suitable geological reservoir intervals
34 into which CO₂ can be injected and stored. It is critical that the target horizons are large, permeable,
35 and structured such that physical or geochemical barriers trap the carbon dioxide in the pore waters,
36 where it can mineralize. Saline aquifers – brine-filled permeable sedimentary rock formations – are of
37 particular interest in this regard.

38 As geological reservoirs, saline aquifers may be ‘closed’ structures, in which geological barriers such
39 as faults or impermeable lithologies prevent the upward (and lateral) migration of injected CO₂ out of
40 the target storage interval. Examples of closed structures include four-way dip closures and fault-
41 bounded reservoirs. Their integrity is dependent on the seal being able to withstand the pressure
42 increase produced by the volume of carbon dioxide injected (see e.g. Mathias et al. 2009, 2011).
43 Saline aquifers can occur additionally in open, tilted structures, with flow dynamics and viscosity
44 controlling the retention of the CO₂ in the reservoir formation (Goater et al. 2013). Saline aquifers
45 within large-scale closed structures tend to have greater storage efficiency, where a higher volume of
46 is CO₂ is retained per unit of available pore space, but dipping open aquifers may offer greater total
47 storage volumes (Goater et al. 2013).

48 Modelling techniques have also been used to argue that closed aquifers are not feasible as storage
49 sites, due to <1% of the pore space being available for injected CO₂ storage before the reservoir
50 pressure is exceeded (Ehlig-Economides & Economides 2010). Consequently, Ehlig-Economides &
51 Economides (2010, p. 130) argued that reservoirs would either need to be “the size of a small US
52 state” or require “hundreds” of injection wells. This interpretation has been contested by Chadwick et
53 al. (2010, 2012) who questioned whether ‘closed’ reservoirs functioned as closed systems, and used
54 data from the Sleipner CCS project in the Norwegian North Sea to show that, over more than a decade
55 of operation, pressure increases were negligible. However, the reservoir permeability of Sleipner is
56 unusually high, at around 5000 mD, more than three times greater than that of any of the other
57 commercial CCS projects assessed by Hosa et al. (2011, fig. 4). Its injectivity (>1200 Darcy-metres)
58 is also exceptional, although there is unlikely to be a clear correlation between reservoir transmissivity
59 (formation thickness × permeability) and injection rate (Mathias et al. 2013). Non-geological factors
60 such as plant design and economics can also play a key role in the injectivity of a storage horizon
61 (Hosa et al. 2011; Mathias et al. 2013). For large-scale injection and storage of CO₂ (>1 Mt per annum
62 per well), therefore, considerable uncertainty remains about the optimal reservoir parameters.

63 In the UK, saline aquifers in the Bunter Sandstone Formation have been identified as some of the
64 most prospective CO₂ storage targets (Holloway & Savage 1993, Holloway et al. 2006, Heinemann et
65 al. 2012). Of Early Triassic age and sub-cropping much of the central and southern UK North Sea, the
66 formation comprises fluvial, alluvial, aeolian and lacustrine sediments deposited in a semi-arid
67 environment. Although their porosity can be reduced by halite cements (Laier & Nielsen 1989), the

68 sandstone-dominated intervals of the formation have suitably high porosity and permeability and
 69 trapping mechanisms to hold large volumes of carbon dioxide, with Heinemann *et al.* (2012)
 70 estimating the formation storage capacity as 3.8–7.8 Gt. The potential reservoirs of the UK North Sea
 71 (Fig. 1) typically occur as four-way dip closures above basement highs or salt domes and, although
 72 fault-bounded, have low degrees of internal faulting. As such, there is a low probability that they are
 73 structurally compartmentalized; the main barriers and baffles to the flow of injected CO₂ are more
 74 likely to be stratigraphic. Determining the quantity of CO₂ that can be stored in these and other saline
 75 aquifers, however, is dependent upon accurate assessment of the risk and nature of
 76 compartmentalization within the reservoirs.



77

78 **Figure 1.** Map of licence blocks in UK Southern North Sea, showing locations of major domal
 79 reservoir structures in the Bunter Sandstone Formation (after Holloway *et al.*, 2005). Saline aquifers
 80 shown in blue, gas fields in red. Abbreviations: Ca – Caister B; Es – Esmond; Fo – Forbes; Go –
 81 Gordon; He – Hewett; Wi – Williams *et al.* (2013) study area.

82

83 Compartmentalization is a concept derived from petroleum geology, in which a geological reservoir is
 84 subdivided into vertical or lateral segments, typically (but not exclusively) with different pressure or

85 fluid properties. Compartmentalization is recognized when there is no communication of fluid flow
86 between different parts of an oil or gas field over production time-scales. It is usually identified by a
87 drop in well-head pressure and oil production rate, and by a rapidly rising gas-oil ratio.

88 Compartmentalization can be controlled by structural features, such as faults acting as barriers or
89 baffles to flow; by stratigraphic variations in porosity and permeability; or by a combination of the
90 two. As it can strongly affect the producible volumes of oil and gas, understanding the degree and
91 nature of compartmentalization present within a reservoir is critical to a field's operational success.

92 A potential CO₂ storage horizon could be compromised if the reservoir interval is compartmentalized.
93 However, this aspect has been little-considered from a CCS perspective, with structural
94 compartmentalization in specific localities examined by only a small number of studies (Hovorka et
95 al. 2011, Castelletto et al. 2013, Meckel et al. 2013). There is also limited experience of injecting CO₂
96 into saline aquifers, and many potential sites have a paucity of existing data (see e.g. Hedley et al.
97 2013). A broad review, utilizing exploration and production data from compartmentalized oil and gas
98 fields, therefore has the potential to provide useful insights into the effects of reservoir
99 compartmentalization on the storage of carbon dioxide in saline aquifers.

100 This report begins by giving an overview of the types of reservoir compartmentalization and their
101 impacts upon oil and gas production. The recent review of reservoir compartmentalization in
102 petroleum geology by Jolley et al. (2010) provides many relevant case studies and discussions of
103 monitoring techniques. These show that for effective asset management, it is vital that
104 compartmentalization is identified early, requiring the integration of different datasets, primarily
105 seismic, sedimentological, and fluid data. Definitive proof of compartmentalization, however,
106 typically comes from data that are not available prior to the start of production (Smalley and
107 Muggeridge 2010). As such, the petroleum industry has developed various techniques and approaches
108 to mitigate against the risk of late discovery of compartmentalization. These are discussed below.

109 The review then explores the likely importance of reservoir compartmentalization to the geological
110 storage of CO₂. It examines stratigraphic compartmentalization in particular detail, with a focus on the
111 Bunter Sandstone saline aquifers of the UK North Sea and the potential impacts of stratigraphic
112 heterogeneities on its storativity. From this, the risks of stratigraphic compartmentalization in the
113 Bunter Sandstone are investigated, the importance to potential CO₂ storage sites are assessed and
114 recommendations are made.

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117

118 **Reservoir compartmentalization**

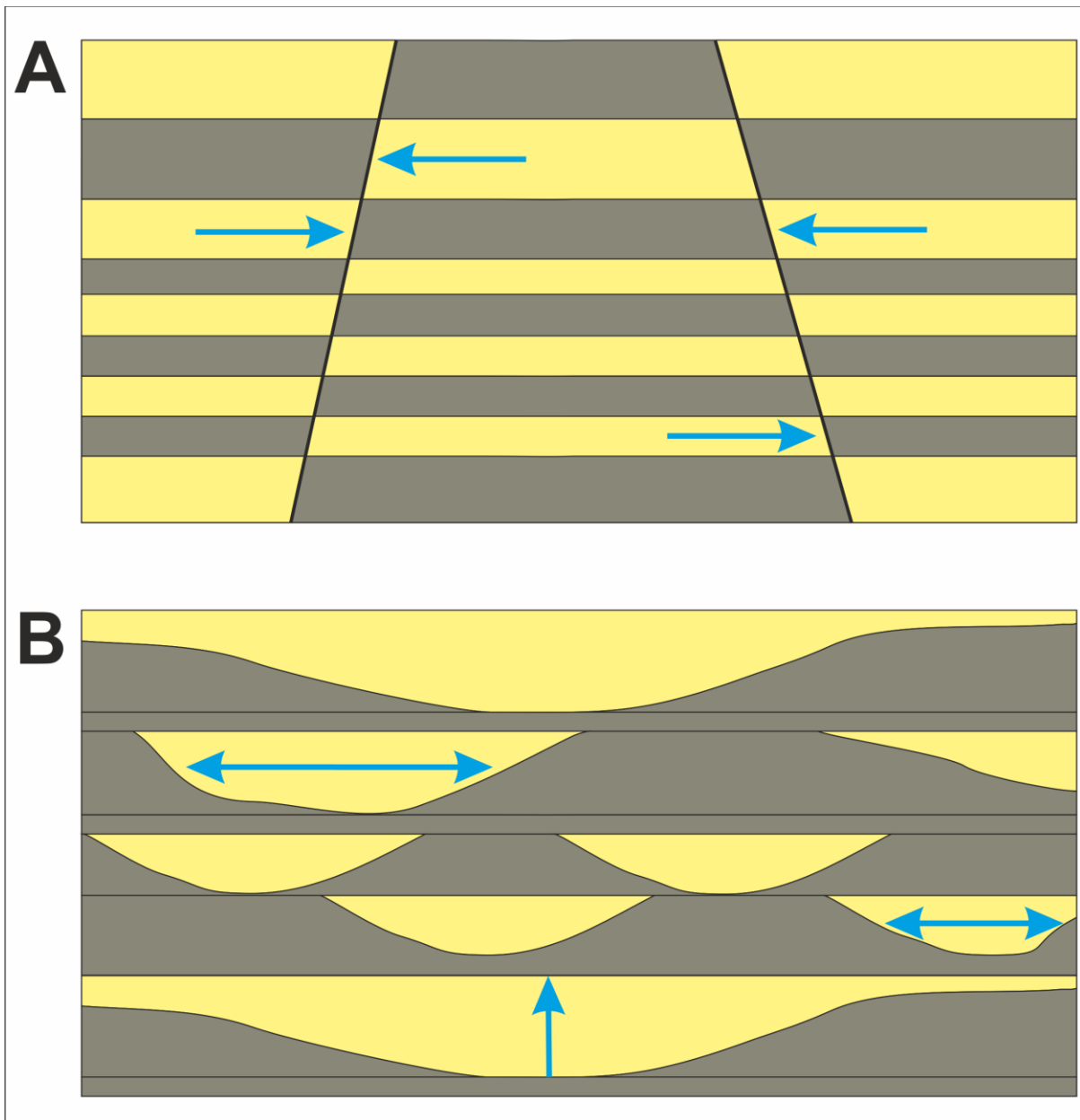
119 The connectivity of reservoir intervals is crucial to the success of a hydrocarbon field. If individual
120 components of a reservoir are not connected, hydrocarbons in an isolated segment will not flow to a
121 producing well in another segment during primary production or secondary, injection-driven recovery
122 (Hovadik and Larue 2010). Compartmentalization is produced by geological heterogeneities that
123 retard or prohibit the movement of fluids through a reservoir. A completely impermeable feature is
124 referred to as a static seal or barrier, whilst a dynamic seal or baffle is one across which partial flow is
125 possible (Jolley et al. 2010; Go et al. 2012). Three types of reservoir compartmentalization are
126 recognized: structural, stratigraphic, and mixed-mode. In structural compartmentalization, faults
127 partition the horizontal movement of fluids within the reservoir (Fig. 2A). Geometry and
128 transmissibility are the most important factors (Irving et al. 2010); faults may be static barriers across
129 which no flow is possible, or dynamic seals that act as baffles. The latter will still compartmentalize
130 the reservoir, however, if the speed of transmissibility across them is slower than the production time
131 scale.

132 Stratigraphic compartmentalization is generated by variations in the extent of impermeable
133 lithologies, such as mudstones and salts, and tends to produce barriers to vertical fluid movement
134 (Fig. 2B). Its magnitude is dependent on the net-to-gross (NTG) threshold and the reservoir
135 dimensions (Hovadik and Larue 2010), with the importance of these factors being both reservoir-
136 specific and time-dependent. Across different geological settings, NTG values of >60% are usually a
137 strong indicator of reservoir connectivity (Hovadik and Larue 2010, fig. 5). However, as discussed
138 below, this is dependent on the architecture of the permeable horizons and, as such, tends to be a
139 function of the depositional setting. Furthermore, the existence of static reservoir connectivity within
140 a field does not guarantee dynamic connectivity. As with faults, stratigraphic barriers can act as static
141 or dynamic seals, and numerical modelling is required to evaluate the dynamic connectivity (Hovadik
142 and Larue 2010).

143 Many reservoirs contain both faults and lateral and vertical variations in lithology, which combine to
144 produce mixed-mode compartmentalization. In many mixed-mode fields, though, one type of
145 compartmentalization will play a more dominant role than the other. In a case study of fields from
146 offshore north-western Australia, for example, Ainsworth (2006) showed that there was a critical
147 sealing fault density of 1.1 per km². Below this density, stratigraphic architecture controlled
148 compartmentalization, whilst above it structural compartmentalization was the major control.

149 Across this variety of contexts, the meaning of compartment can be somewhat ambiguous. The term
150 *drainage cell* was therefore used by Shepherd (2009) to describe situations where an orthogonal,
151 intersecting fault system combines with stratigraphic barriers to create box-shaped reservoir intervals

152 isolated by permeability barriers. However, Gill et al (2010) subsequently also applied the term to
153 non-connected units within stratigraphically compartmentalized reservoirs.



154
155 **Figure 2.** Schematic examples of reservoir compartmentalization types. A) structural
156 compartmentalization of horizontally bedded succession of reservoir sandstones (yellow) and non-
157 reservoir mudstones (grey), with normal faults creating barriers to horizontal fluid flow (blue arrows).
158 B) stratigraphic compartmentalization of stacked succession of non-reservoir mudstones (grey) and
159 channel sandstones (yellow), with channelization of reservoir sandstones into isolated bodies creating
160 barriers to horizontal and vertical fluid flow (blue arrows).

161

162 Impacts of compartmentalization to the oil and gas industry

163 Recognizing compartmentalization within a reservoir is crucial to the economic viability of an oil or
164 gas field; accurate estimates of reserve volumes dictate the field's value (Gluyas and Swarbrick 2004).
165 The late recognition of compartmentalization will lead to increased costs, through having to spend
166 longer collecting data before going into production, and reduced profitability from the realization that
167 fewer hydrocarbons are retrievable (Jolley et al. 2010). If compartmentalization can be identified
168 before a field goes into production, therefore, the economic impacts can be minimized.

169 In many cases, however, it is very difficult to confirm whether a reservoir is compartmentalized until
170 a field begins producing, and dynamic pressure-time data can be analysed. In the UK North Sea, for
171 example, numerous case studies discussed below show that reservoir compartmentalization has been
172 'systematically underestimated' (Smith 2008). As a consequence, various approaches have been
173 developed to better quantify the risk and nature of reservoir compartmentalization.

174

175 Mitigation technologies and industry methods for estimating compartmentalization risk

176 Reservoir compartmentalization is assessed using a variety of techniques. These include the analysis
177 of fluid properties, pressure-volume-temperature (PVT) data, and 3D or 4D seismic monitoring. For
178 field-scale reservoir characterization, 3D seismic surveys use acoustic reflectivity to analyse
179 subsurface geology and identify features indicative of compartmentalization. These include
180 lithological changes, fault distributions, and fluid variations. By repeating such surveys over time, 4D
181 seismic models can then be generated, enabling fluid movement and reservoir drainage to be assessed.
182 Such analyses are key to confirming the existence and nature of reservoir compartmentalization, but
183 are expensive and may therefore not be employed.

184 Well test data analysis, where the production behaviour of a well is examined, is a key method in
185 evaluating potential compartmentalization. This approach focusses in particular on the flow rates of
186 gas, oil and water over time as field production develops. Extended well testing over a duration of
187 weeks to months is ideal, to obtain maximal information on hydrocarbon properties, reservoir
188 structure and field behaviour. However, this can be economically or environmentally unfeasible
189 (Gluyas and Swarbrick 2004). Regardless of the duration of the test, evaluating the subsequent
190 behaviour of the reservoir is especially important. By the employment of pressure meters in the well,
191 the post-test recovery of pressure in the well is analysed. These pressure build-up tests enable the
192 natural restoration of pressure in the reservoir to be assessed, and potential flow barriers and baffles to
193 be identified (Gluyas and Swarbrick 2004).

194 Within a single well, test data analysis may reveal indications that compartmentalization is present,
195 but not its extent. To obtain more detailed information, an interference test may be carried out. In such

196 analyses, a single well is placed into production or injection, whilst adjacent wells are shut in with
197 pressure gauges installed (Gluyas and Swarbrick 2004). When production or injection commences in
198 the test well, the subsequent pressure changes and durations can be determined in the surrounding
199 wells, providing a much clearer picture of the reservoir structure and behaviour.

200 At the commencement of drilling, in order to provide the baseline field data on fluid characteristics
201 and behaviour, pressure-volume-temperature (PVT) analysis must be carried out. Variations in PVT
202 values from different wells within a field are a strong indicator of reservoir compartmentalization.
203 However, Paez et al. (2010) showed that if the fluids are in a critical or near-critical state, they can be
204 gravity segregated, such that wells within a connected reservoir might give different PVT values
205 suggestive of compartmentalization.

206 A variety of other fluid properties can be analysed to examine potential compartmentalization,
207 including chemistry, density, and viscosity. Though there is always a risk of water contamination,
208 careful interrogation of produced fluid data is cheap and high-impact (Jolley et al. 2010). As with
209 PVT data, if different wells or reservoir intervals show different fluid properties, the reservoir may be
210 interpreted as compartmentalized. However, as shown by Smalley & Muggeridge (2010), it is vital to
211 know the time required for the fluid property in question to equilibrate, along with the size and age of
212 the reservoir.

213 The isotopic composition of reservoir pore fluids, for example, will equilibrate slowly across a
214 reservoir. For a two million year-old reservoir, Smalley & Muggeridge (2010) calculated that
215 hydrocarbon mixing by molecular diffusion would have occurred over a distance of less than 1 km.
216 Even if clear differences in isotopic composition were measured within or between wells, this would
217 not indicate the reservoir is compartmentalized. Conversely, aquifer pressure differences will
218 equilibrate over 10 kilometres within 10 years (Smalley & Muggeridge 2010), so if differences in
219 fluid pressure are detected within a new field, it is much more likely that the reservoir is
220 compartmentalized.

221 Vertical and lateral variations in the chloride ions of produced water enabled Gill et al. (2010) to
222 identify nine separate drainage cells in the Nelson field of the UK Central North Sea. These were
223 controlled by stratigraphic compartmentalization by turbidite channel heterogeneity. Fluid analysis
224 can also be employed to assess changes in reservoir compartmentalization. Time-lapse geochemistry –
225 monitoring the composition of fluids during appraisal, development and production – was used in the
226 Auger oil field in the Gulf of Mexico (Chuparova et al. 2010). The acquisition of pre-production fluid
227 data provided the baseline by which changes in fluid composition could be measured. Collecting
228 produced fluid samples was then a cheap means of detecting dynamic changes in reservoir
229 communication. By assessing the ratio of oil to gas condensate it was shown that, after six years of
230 production, an oil and a gas reservoir that static pressure data indicated were separate, were now in

231 dynamic communication (Chuparova et al. 2010). Such early detection of changing reservoir
232 behaviour can yield a significant benefit to the economics of a field.

233 An approach outlined by Richards et al. (2010) is Reservoir Connectivity Analysis (RCA), in which
234 reservoir compartments are defined from seismic, pressure and fluid data, and the mechanical and
235 capillary seal mechanisms between the compartments then evaluated. From this, the compartments
236 and connections are mapped schematically to produce connectivity diagrams, enabling the
237 compartmentalization of the reservoir to be assessed at different scales. Once production is underway,
238 continuous data surveillance enables the models to be tested and confirmed as valid. As Jolley et al.
239 (2010) state, compartmentalization may be suggested by various individual sources of data, but it is
240 most reliably confirmed by early integration of diverse datasets and their subsequent monitoring.

241

242 **Stratigraphic compartmentalization**

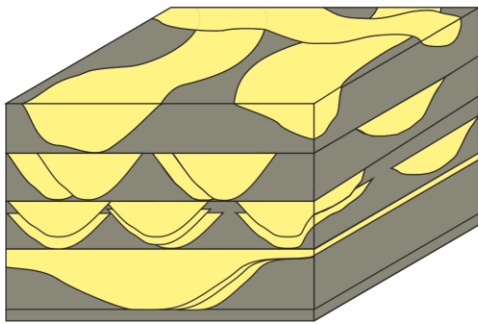
243 The most important factor in stratigraphically compartmentalized reservoirs is the NTG threshold.
244 The higher the NTG ratio of a field, the more likely it is that reservoir intervals are connected. Based
245 on a variety of reservoir analyses, Hovadik and Larue (2010, p. 241) defined 60% as the level above
246 which connectivity is ‘unlikely to be a problem’. As shown in Figure 3, however, the stratigraphic
247 architecture is important. In stacked, amalgamated channelized reservoir systems, NTG values of as
248 low as 30% could still indicate good connectivity (Figure 3A; Hovadik and Larue 2010). In straight,
249 parallel, channel systems (Figure 3B), NTG values of 60% are typically required, whilst in sheeted
250 reservoirs (Fig. 3C), geobody connectivity will typically be low even when NTG values are >60%
251 (Hovadik and Larue 2010). Other key elements to consider include reservoir dimensionality (e.g. the
252 size and continuity of impermeable horizons) and the size of geobodies (connected reservoir units)
253 when compared with that of the reservoir as a whole (Hovadik and Larue 2010). Mainly the
254 environment the reservoir rocks were deposited in and their subsequent diagenetic history themselves
255 control these factors.

256

257 Stratigraphic compartmentalization by depositional setting

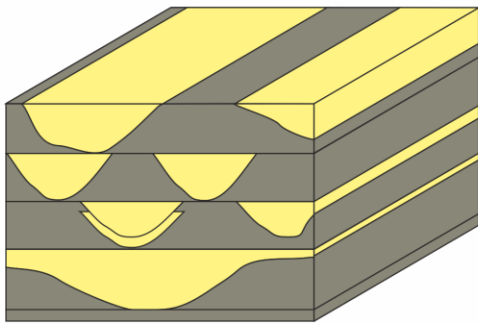
258 The size, shape and properties of reservoirs within a target formation are dictated by the depositional
259 conditions under which the succession formed. From terrestrial settings to deep marine basins,
260 different sedimentary environments produce very different types of geobodies. As such, to predict the
261 type of stratigraphic compartmentalization that might be present in a reservoir, the depositional
262 environment must be evaluated.

A



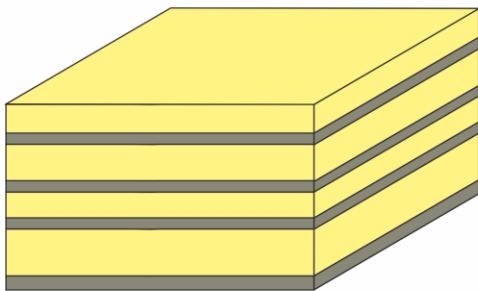
Amalgamated channels:
Geobody connectivity
requires NTG of ~ 0.3

B



Straight/parallel channels:
Geobody connectivity
requires NTG of ~ 0.6

C



Sheet deposits:
Geobody connectivity
requires NTG of ~ 1

263

264 **Figure 3.** Net-to-gross (NTG) versus connectivity in reservoirs of different architecture (after
265 Hovadik & Larue 2010, fig. 5). A, Stacked, amalgamated channel systems have good reservoir
266 connectivity even with low NTG values. B, Straight, parallel channel systems have good reservoir
267 connectivity only when NTG values are moderate–high. C, Sheet deposits typically have poor
268 connectivity, even with high NTG values.

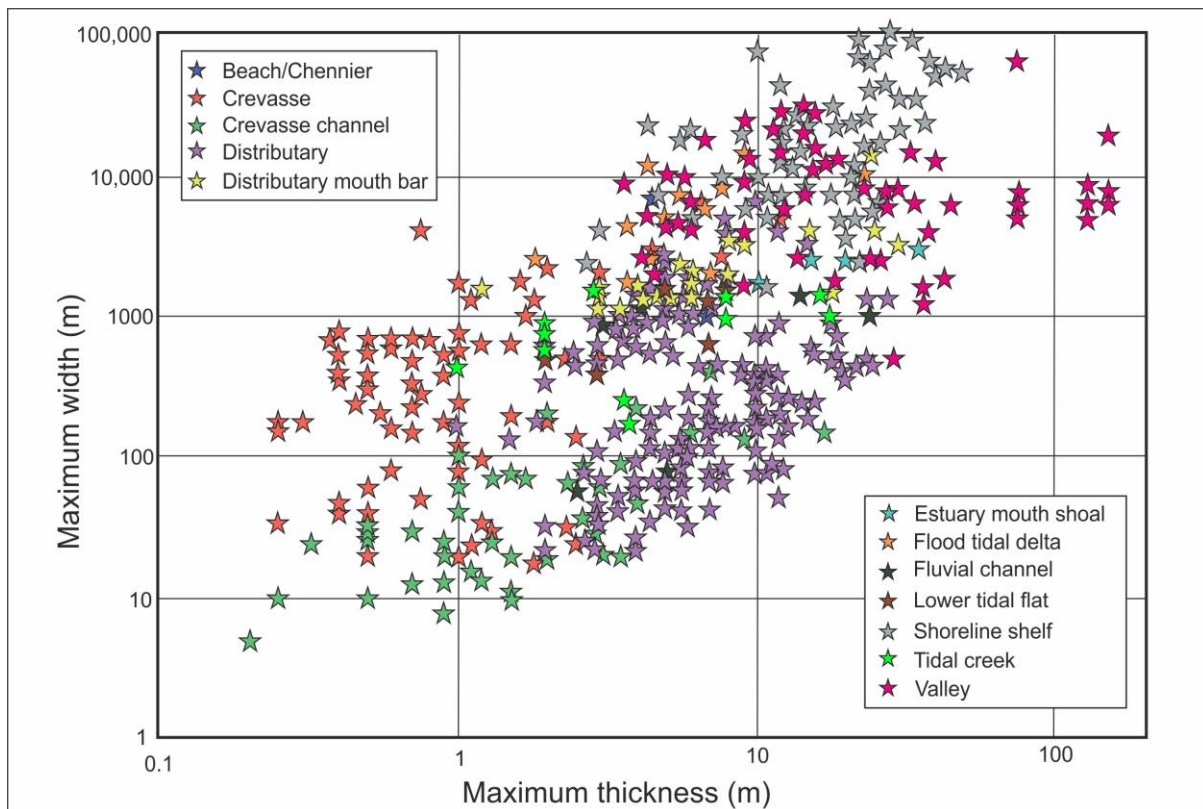
269

270 Turbidite successions are the main type of deep marine reservoir, formed where high-energy currents
271 carried large volumes of sand down the continental shelf into a marine basin. A number of such fields
272 occur in the UK North Sea and stratigraphic compartmentalization within them is controlled by the
273 geometry and connectivity of the sandstones. Since turbidite sands are typically transported in

274 channels or down canyons, the lateral and vertical variability of these features is key to determining
275 how compartmentalized a turbidite reservoir is.

276 In the Pierce field of the North Sea, Scott et al. (2010) argued that compartmentalization was
277 controlled by stratigraphic architecture, with early Cenozoic turbidite sandstones being channelized
278 around a pair of salt diapirs. The contemporaneous turbidite succession of the Nelson Field does not
279 have salt diapirs present, but channelization of the sandstones still controls the compartmentalization,
280 creating nine stratigraphic drainage cells (Gill et al. 2010).

281 There are many different types of marginal marine (paralic) successions and the dimensions of the
282 reservoir intervals within them are accordingly variable (Fig. 4; after Reynolds 1999). The largest
283 sand bodies by area occur in shoreline successions deposited during periods of high sea-level, and the
284 thickest in lowstand valley sandstone successions, whilst crevasse and crevasse channel sandstones are
285 the narrowest and thinnest (Fig. 4; after Reynolds 1999, figs 5, 6).



287 **Figure 4.** Sandbody dimensions of different reservoir types (after Reynolds 1999, figs 5, 6). Crevasse
288 and crevasse channel successions are typically thinnest and narrowest, whilst shoreline shelf and
289 valley successions are thickest and widest.

290

291 In general, though, NTG ratios in shallow marine reservoirs tend to be higher than deep marine
292 successions due to their elevated sand content. Stratigraphic compartmentalization of marginal marine
293 successions then occurs at three scales: intra-sandbody, inter-sandbody, and inter-parasequence scale
294 (Ainsworth 2010). All are the result of mudstones acting as baffles or barriers. Intra- and inter-
295 sandbody scale compartmentalization occurs in heterogeneous sequences deposited during periods of
296 high sedimentation rate. In contrast, inter-parasequence scale compartmentalization is created by
297 flooding surface shales, produced during major marine transgressions. These mudstones will have a
298 much greater extent and create much more laterally persistent barriers to reservoir communication. In
299 general, though, with a high sand component, reservoirs deposited in wave-dominated shallow marine
300 systems show low levels of stratigraphic compartmentalization (Ainsworth 2006).

301 Channelization of sandstones also creates stratigraphic compartmentalization in marginal marine
302 successions. In the Northwest Hutton field, the reservoir intervals comprise Middle Jurassic sheet or
303 channel sandstones deposited in coastal plain, marginal marine or shallow marine environments. Flint
304 et al. (1998) showed that connectivity was generally good, but that stratigraphic compartmentalization
305 was present and controlled by two main factors. The first was the laterally persistent mudstones of the
306 lower Ness Formation, produced during a marine transgression and creating a widespread, low
307 permeability barrier to vertical fluid flow. The second was the variable dimensions of channel
308 sandstones, particularly in the fluvially dominated Etive and upper Ness formations, with fine-grained
309 inter-channel sediments acting as barriers to horizontal fluid flow.

310 This type of horizontal stratigraphic compartmentalization is commonly also observed in reservoirs
311 deposited in fluvial or aeolian environments. Major barriers to vertical fluid flow are most prevalent
312 in terminal fluvial systems, where the greatest quantities of mudstone accumulate. This was described
313 in the Triassic Heron Cluster of the North Sea by McKie et al (2010), who noted that thin shale
314 intervals could be breached by subsequent channel sands, providing connectivity. Thicker shales
315 deposited in terminal, flood basin regions, however, were rarely incised by sandstones and therefore
316 created barriers over tens of kilometres. A similar scenario has been predicted by Archer et al. (2010)
317 for the Jasmine Field, with the laterally extensive Julius and Jonathan mudstone intervals being likely
318 to create stratigraphic compartmentalization within the Joanne Sandstone over production timescales.

319 For terrestrial successions deposited in semi-arid, alluvial and braided fluvial environments, the
320 reservoir intervals occur within channel and sheet sandstones, the shifting nature of braided channel
321 systems providing good lateral and vertical connectivity of sandbodies. Stratigraphic
322 compartmentalization is caused primarily by overbank (floodplain) mudstones, and playa lake
323 mudstone and evaporite (salt) deposits, particularly in distal settings. This is discussed further below
324 in the context of the stratigraphy and compartmentalization of the Bunter Sandstone Formation.

325

326 Recognition of stratigraphic compartmentalization:

327 One of the key approaches for identifying stratigraphic compartmentalization early is the use of
328 sequence stratigraphy. Understanding relative changes in sea-level during the deposition of a reservoir
329 succession enables a field-scale stratigraphic framework to be produced. As they have the greatest
330 lateral extent, parasequence-bounding shales (mudstones deposited during periods of high sea-level)
331 are stratigraphic compartmentalizers of the highest importance. Similarly, sequence boundaries
332 (erosive surfaces produced during sea-level fall) can produce widespread changes in permeability that
333 compartmentalize a reservoir.

334 In the Upper Jurassic turbidites of the Buzzard Field, North Sea, stratigraphic compartmentalization
335 was identified shortly after production began by pressure-time analyses of pressure monitoring wells
336 adjacent to a production well (Ray et al. 2010). These data showed that, after water injection began,
337 pressure declined in three wells but increased in a fourth. Re-examination of seismic and stratigraphic
338 data revealed that a previously unrecognized mudstone interval was partitioning this region of the
339 field into vertical compartments.

340 The turbidite reservoirs of the Schiehallion Field, meanwhile, were initially thought to be well-
341 connected. However, once production began, it transpired that they were more stratigraphically
342 compartmentalized than predicted, due to a greater number of localized mudstones acting as barriers
343 to flow (Gainski et al. 2010). As a consequence, more than double the number of production wells had
344 to be drilled; 4D seismic and downhole pressure data have been used subsequently to monitor
345 evolution of the field.

346 In block 21/28a of the UK central North Sea, six wells were drilled between 1971 and 1989,
347 encountering oil and gas in a series of Eocene sandstone reservoirs. The discoveries were named Fyne
348 and Dandy (Gluyas and Swarbrick 2004). The fields never went into production, however, as RFT
349 pressure analyses indicated that the reservoirs were not in communication, but for unclear reasons. In
350 the late 1990s, higher resolution seismic surveys and palaeoenvironmental analyses were carried out
351 (Gluyas and Swarbrick 2004), showing that there were four separate hydrocarbon pools in a complex
352 submarine fan succession. Though the individual sandstone reservoirs were of good porosity and
353 permeability, they comprised elements of a meandering channel system separated by numerous
354 mudstones, and the complexity of this meant the fields did not go into production.

355

356 **Stratigraphic compartmentalization and CCS in saline aquifers**

357 Unidentified stratigraphic compartmentalization has the potential to significantly reduce the viability
358 of a site for carbon storage. The Snøhvit project in northern Norway was predicted to be able to store

359 23 Mt of CO₂ in a saline aquifer within the Tubaen Formation, 2400 m below the sea floor (Hosa et
360 al. 2011). However, unexpected compartmentalization encountered during the early phase of injection
361 led to a rapid pressure increase and a greatly reduced storage volume (Eiken et al. 2011). One interval
362 – comprising only around one-sixth of the reservoir interval – received most of the injected CO₂, with
363 the compartmentalization being primarily stratigraphic. A previously unrecognized shale interval is
364 thought to act as a barrier to vertical flow, with possible additional compartmentalization caused by
365 channel sandstones (Eiken et al. 2011).

366 Brine-filled domes in the Bunter Sandstone Formation of the UK southern North Sea (Fig. 1) have
367 been targeted as promising potential CO₂ storage sites. Due to the more limited lateral and temporal
368 extent of the sandbodies, fluvio-aeolian reservoirs such as the Bunter are often more stratigraphically
369 compartmentalized than shallow marine systems, though as Reynolds (1999) discussed, there is wide
370 variability in paralic (marginal marine) successions. Fluvio-aeolian reservoirs are also prone to
371 vertical changes in lithology (from permeable sandstone to impermeable shale or salt) and horizontal
372 changes in the dimensions of lithological units (most notably the channelization of permeable
373 sandstones). To determine the migration pathway of any injected CO₂, therefore, the distribution and
374 continuity of low permeability layers in the Bunter Sandstone needs to be assessed (Williams et al.
375 2013).

376 Few direct data are available for the saline water-bearing stratigraphic traps, but a small number of
377 Bunter domes in the North Sea – Caister B, Esmond, Forbes, Gordon and Hewett – are productive gas
378 fields (Fig. 1). Potential carbon storage sites lie within this region, so published data on the geology of
379 the gas fields provide a source of information on potential stratigraphic compartmentalization within
380 the Bunter Sandstone Formation.

381

382 *Bunter Sandstone gas fields*

383 In the Hewett Gas Field (Southern North Sea), the Bunter Sandstone is around 600-800 feet (180-260
384 m) in thickness, and comprises thick, fluvial channel and sheetflood sandstones with an average NTG
385 of 0.96 (Cooke-Yarborough 1991, Cooke-Yarborough & Smith 2003). A second sandstone reservoir
386 interval is also present: the Lower Bunter (Hewett) Sandstone, a fault-controlled, laterally restricted
387 sandbody of late Permian–early Triassic age. This occurs only in the Hewett Field and a small
388 adjacent region of the Dutch North Sea (Geluk 1999). It has an average NTG of 0.86, and is up to 200
389 feet (61 m) thick in the main Hewett Field, but thins rapidly northwards to a thickness of only 20 feet
390 (6 m) (Cooke-Yarborough 1991; see also Clarke 2014). The two sandstone reservoir intervals are
391 separated by anhydritic floodplain mudstones of the Bunter Shale, which are up to ~800 feet in

392 thickness. The Bunter Shale and the younger, dolomitic Dowsing Formation act as both lateral and
393 vertical seals across the field.

394 The stratigraphy of the Bunter Sandstone in the Hewett Field was not discussed by Cooke-
395 Yarborough (1991) or Cooke-Yarborough & Smith (2003), but further data can be obtained from its
396 stratigraphy in nearby gas fields. In the Esmond Complex (the Esmond, Forbes and Gordon fields),
397 the formation is 400-500 feet (120-150 m) thick (Ketter 1991). The sandstones are predominantly
398 sheet-like, but with three internal hierarchies of heterogeneity: cross-sets, channels, and channel
399 complexes. Five lithofacies were described, with the highest porosity-permeability values occurring in
400 sheetflood sandstones and tabular-planar upper channel fill sandstones.

401 Reservoir analysis by Ketter (1991) also identified seven discrete zones within the formation: an
402 upper sandstone (Zone I), a deep red silty mudstone acting as a seal to the units below (Zone II), a
403 thick, uniform sandstone (Zone III), an interval of mixed sandstone and mudstone (Zone IV), two
404 intervals of fluvial channels and sheetflood sandstones (Zones V and VI), and a lower interval of
405 increasingly silty sandstone (Zone VII). Zone III is the primary reservoir interval. More recent work
406 on the Esmond Complex (Centrica 2005) has shown that there was low to very low aquifer support in
407 most reservoir intervals, but that active aquifer support was present in Zone III of the Forbes field.

408 In the Caister B gas field, where the Bunter is around 150m thick, Ritchie and Pratsides (1993)
409 described a sheetflood sandstone-dominated succession, which they divided into five stratigraphic
410 intervals of varying NTG values. Using wireline log data, Williams et al. (2013) also identified five
411 stratigraphic zones within a Bunter Sandstone dome in block 44/26 of the southern North Sea (Fig. 1).
412 These comprise a thin, shale-dominated Zone 1, a thin Zone 2 of shales and cemented sandstones, a
413 thick Zone 3 of sandstones with thin shales, a thick sandstone-dominated Zone 4 with discontinuous
414 shale layers, capped by a 1–2 metre-thick cemented sandstone, and a thin Zone 5 of sandstones and
415 shales. The non-cemented sandstones have 5–35 % porosity, but very limited permeability data were
416 available.

417 During the deposition of the Bunter Sandstone Formation, sediment transport was in a broadly north-
418 easterly direction (Warrington & Ivimey-Cook 1992), so the axial orientation of higher permeability
419 channel sandstones is likely to follow that trend. There are also differing degrees of cementation
420 within the formation. The cemented sandstone at the top of Zone 4 is widespread, extending into
421 Caister B (Williams et al. 2013), but its lateral continuity is unresolved.

422 In the Esmond Complex, Ketter (1991) described the communication between reservoir intervals as
423 ‘good though locally tortuous’, with fluid flow paths being better along the axes of the fluvial
424 channels (broadly N-S) than across them (broadly E-W). Cooke-Yarborough & Smith (2003)
425 described “remarkable” communication between the different reservoir intervals in the Hewett Field,

426 even across faults. This has been developed further by Clarke (2014), whose fault connectivity
427 analysis indicates that reservoir communication between the Hewett and Little Dotty fields – across
428 the North Hewett fault of the Dowsing Fault system – is driven by the juxtaposition of reservoir
429 sandstones in the southern region of the fault.

430 Despite the heterogeneities, the very high NTG values recorded for the Bunter Sandstone in the
431 Hewett, Esmond Complex and Caister B fields indicate that connectivity should be very good (cf.
432 Hovadik & Larue 2010), and there is no indication of compartmentalization in the Bunter gas fields
433 (Holloway et al. 2005). Furthermore, variations in reservoir architecture within channelized sandstone
434 successions may not matter if those variations do not affect reservoir connectivity or tortuosity (Larue
435 and Hovadik 2006). Williams et al. (2013) stated that most of the low permeability shales and
436 cemented sandstones in the Bunter are not laterally continuous, and would therefore have little effect
437 on CO₂ storage potential. This is supported by Clarke (2014), who showed that a widespread shale
438 interval predicted to act as a barrier to vertical fluid flow in the Upper Bunter of the Hewett Field did
439 not actually act as such.

440 Gas production from the Bunter Sandstone Formation of the Hewett Field was thought to be by
441 depletion drive, but recent work by Clarke (2014) shows that, whilst this is true of the Lower Bunter
442 reservoir interval, the Upper Bunter reservoir experienced aquifer drive. In the Esmond Complex,
443 meanwhile, aquifer influx coefficients indicate that the depletion-driven Esmond and Gordon fields
444 will not have undergone significant repressurization since they were abandoned (Centrica 2005). In
445 the Forbes field, however, aquifer drive in the main reservoir interval (Zone III) means that
446 repressurization since abandonment is likely. This potential drive mechanism variability within
447 Bunter reservoirs needs to be taken into account when planning CO₂ injection strategies.

448 Further information could be obtained from analysis of fluid properties. In the Hewett Field, the
449 trapping structure formed between 50 and 150 Ma (Cooke-Yarborough 1991), an age range that is
450 likely to apply to other Bunter Sandstone reservoirs. Based on the work of Smalley & Muggeridge
451 (2010), all fluid properties ought to have equilibrated over that time scale, even by molecular
452 diffusion, so if analysis of fluid composition was carried out, any variations in fluid composition
453 would suggest the presence of reservoir compartments.

454

455 **Summary and conclusions**

456 The impacts of reservoir compartmentalization on the geological storage of carbon dioxide have been
457 little-considered previously, particularly in terms of stratigraphic variation. This review shows that

458 techniques used to detect compartmentalization in oil and gas fields are of relevance to understanding
459 the prospectivity of saline aquifers as sites for the geological storage of carbon dioxide.

460 As a particularly prospective CCS reservoir interval, the Bunter Sandstone Formation must be
461 carefully assessed for compartmentalization before the injection of CO₂ commences. The case studies
462 of Bunter gas fields from the UK North Sea described above show that stratigraphic heterogeneities
463 are present in the formation at various scales. Typically, five to seven main stratigraphic units have
464 been recognized within the formation, though in the Hewett Field, a distinct, lower reservoir unit also
465 occurs. There is potential therefore for lateral and vertical variation in the reservoir properties of the
466 formation across the UK Southern North Sea. Nonetheless, the high NTG values documented, and the
467 limited lateral extent of stratigraphic barriers mean that vertical and horizontal connectivity within the
468 formation is likely to be good, and the risk of significant compartmentalization low. Exploration data
469 on the geochemistry of produced fluids would help clarify this.

470 Numerous case studies from oil and gas fields show that, nonetheless, the precise nature of the
471 compartmentalization of a reservoir is often determinable only after production has begun. For saline
472 aquifers thought to be prospective for the storage of carbon dioxide, the collection of detailed
473 stratigraphic information, seismic analysis, and assessment of fluid properties within the formation
474 will provide a general indication of the presence of reservoir compartmentalization. Experimental
475 drilling and injection, however, is likely to be the key means of clarifying the distribution, orientation
476 and connectivity of any compartments – and therefore the true storage potential – of such units.

477

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