

1 CEDAR-GPP: spatiotemporally upscaled estimates of gross primary
2 productivity incorporating CO₂ fertilization

3 Yanghui Kang^{1,2}, Max Gaber^{1,3}, Maoya Bassiouni^{1,2}, Xinchun Lu^{1,2}, Trevor F. Keenan^{1,2}

4 ¹ Department of Environmental Science, Policy, and Management, University of California,
5 Berkeley, Berkely, CA 94720, USA

6 ² Climate and Ecosystem Sciences Division, Lawrence Berkeley National Laboratory, Berkeley, CA
7 94720, USA

8 ³ Department of Geosciences and Natural Resource Management, University of Copenhagen,
9 Copenhagen, 1350, Denmark.

10 Correspondence: Yanghui Kang (yanghuikang@berkeley.edu)

11 Trevor Keenan (trevorkeenan@berkeley.edu)

12
13
14 This paper is a non-peer-reviewed preprint submitted to EarthArXiv. Subsequent
15 versions of this paper may contain revisions and updates.
16
17
18

19 **Abstract:** Gross primary productivity (GPP) is the largest carbon flux in the Earth system, playing a
20 crucial role in removing atmospheric carbon dioxide and providing the sugars and starches needed
21 for ecosystem metabolism. Despite the importance of GPP, however, existing estimates present
22 significant uncertainties and discrepancies. A key issue is the underrepresentation of the CO₂
23 fertilization effect, a major factor contributing to the increased terrestrial carbon sink over recent
24 decades. This omission could potentially bias our understanding of ecosystem responses to climate
25 change.

26 Here, we introduce CEDAR-GPP, the first global upscaled GPP product that incorporates
27 the direct CO₂ fertilization effect on photosynthesis. Our product is comprised of monthly GPP
28 estimates and their uncertainty at 0.05° resolution from 1982 to 2020, generated using a
29 comprehensive set of eddy covariance measurements, multi-source satellite observations, climate
30 variables, and machine learning models. Importantly, we used both theoretical and data-driven
31 approaches to incorporate the direct CO₂ effects. Our machine learning models effectively predicted
32 monthly GPP ($R^2 \sim 0.74$), the mean seasonal cycles ($R^2 \sim 0.79$), and spatial variabilities ($R^2 \sim 0.67$).
33 Incorporation of the direct CO₂ effects substantially improved the models' ability to estimate long-
34 term GPP trends across global flux sites. While the global patterns of annual mean GPP, seasonality,
35 and interannual variability generally aligned with existing satellite-based products, CEDAR-GPP
36 demonstrated higher long-term trends globally after incorporating CO₂ fertilization, particularly in
37 the tropics, reflecting a strong temperature control on direct CO₂ effects. CEDAR-GPP offers a
38 comprehensive representation of GPP temporal and spatial dynamics, providing valuable insights
39 into ecosystem-climate interactions.

40

41 1. Introduction

42 Terrestrial ecosystem photosynthesis, known as Gross Primary Productivity (GPP), is the
43 primary source of food and energy for the Earth system and human society. Through
44 photosynthesis, terrestrial ecosystems also mitigate climate change, by removing thirty percent of
45 anthropogenic carbon emissions from the atmosphere each year (Friedlingstein et al., 2023).
46 However, due to the lack of direct measurements at the global scale, our understanding of
47 photosynthesis and its spatiotemporal dynamics is limited, leading to considerable disagreements
48 among various GPP estimates (Anav et al., 2015; O’Sullivan et al., 2020; Smith et al., 2016; Yang et
49 al., 2022). Addressing these uncertainties is crucial for improving the predictability of ecosystem
50 dynamics under climate change (Friedlingstein et al., 2014).

51 Over the past three decades, global networks of eddy covariance flux towers collected *in situ*
52 carbon flux measurements that allow for accurate estimates of GPP, providing valuable insights into
53 photosynthesis dynamics under various environmental conditions (Baldocchi, 2020; Beer et al.,
54 2010). To quantify and understand GPP at scales and locations beyond the $\sim 1\text{km}^2$ flux tower
55 footprints, machine learning has been employed with gridded satellite and climate datasets to upscale
56 site-based measurements and produce wall-to-wall GPP maps (Yang et al., 2007; Xiao et al., 2008;
57 Jung et al., 2011; Tramontana et al., 2016; Joiner and Yoshida, 2020; Zeng et al., 2020; Dannenberg
58 et al., 2023). This approach provides important observational constraints of global carbon dynamics,
59 complementing process-based and semi-process-based modeling such as Terrestrial biosphere
60 models or the Light Use Efficiency (LUE) models (Beer et al., 2010; Jung et al., 2017; Schwalm et
61 al., 2017; Gampe et al., 2021).

62 Effective machine learning upscaling depends on a complete set of input predictors that fully
63 explain GPP dynamics. Upscaled datasets have primarily relied on satellite-observed greenness
64 indicators, such as vegetation indexes (VIs), Leaf Area Index (LAI), the fraction of absorbed
65 photosynthetically active radiation (fAPAR), which effectively capture canopy-level GPP dynamics
66 related to leaf area changes (Tramontana et al., 2016; Ryu et al., 2019; Joiner and Yoshida, 2020).
67 However, important aspects of leaf-level physiology, such as those controlled by climate factors, are
68 often omitted in major upscaled datasets, preventing accurate characterization of GPP responses to
69 climate change (Stocker et al., 2019; Bloomfield et al., 2023). In particular, none of the previous
70 upscaled datasets have considered the direct effect of atmospheric CO_2 on leaf-level photosynthesis,
71 which is a key factor contributing to at least half of the enhanced land carbon sink observed over the

72 past decades (Keenan et al., 2016; Keenan and Williams, 2018; Walker et al., 2021; Ruehr et al.,
73 2023). This omission can lead to incorrect inference regarding long term trends in various
74 components of the terrestrial carbon cycle (De Kauwe et al., 2016).

75 Multiple independent lines of evidence from atmospheric inversion (Wenzel et al., 2016),
76 atmospheric $^{13}\text{C}/^{12}\text{C}$ measurements (Keeling et al., 2017), ice core records of carbonylsulfide
77 (Campbell et al., 2017), glucose isotopomers (Ehlers et al., 2015), as well as free-air CO_2 enrichment
78 experiments (FACE) (Walker et al., 2021), suggest a widespread positive effect of elevated
79 atmospheric CO_2 on GPP from site to global scales. Increasing CO_2 *directly* stimulates the
80 biochemical rate of leaf-level photosynthesis, leading to an increase in net carbon assimilation and
81 leaf area, which enhances canopy-level GPP. Furthermore, high CO_2 concentration is expected to
82 reduce stomatal conductance and increase water use efficiency, indirectly enhancing photosynthesis
83 under water-limited conditions (De Kauwe et al., 2013; Keenan et al., 2013). The direct biochemical
84 effect has been found to dominate GPP responses to CO_2 , from both theoretical and observational
85 analyses (Haverd et al., 2020; Chen et al., 2022).

86 Satellite-based estimates have shown an increasing global GPP trend in the past few decades
87 largely attributable to CO_2 -induced increases in LAI (De Kauwe et al., 2016; Zhu et al., 2016; Chen
88 et al., 2019; Piao et al., 2020). However, previous upscaled GPP datasets, as well as most LUE
89 models such as the MODIS GPP product, have failed to consider the direct CO_2 effects on leaf-
90 level biochemical processes (Jung et al., 2020; Zheng et al., 2020). Consequently, these products
91 likely underestimated the long-term trend of global GPP, leading to large discrepancies when
92 compared to process-based models, which typically consider leaf-level CO_2 effects (Anav et al.,
93 2015; De Kauwe et al., 2016; O'Sullivan et al., 2020). Notably, recent improvements in LUE models
94 have included the CO_2 response and show improved long-term changes in GPP globally (Zheng et
95 al., 2020), yet, this important mechanism is still missing in GPP products upscaled from *in situ* eddy
96 covariance flux measurements.

97 To improve the quantification of GPP spatial and temporal dynamics and provide a robust
98 representation of long-term dynamics in global photosynthesis, we developed the CEDAR-GPP¹
99 data product. CEDAR-GPP was upscaled from global eddy covariance carbon flux measurements
100 using machine learning along with a broad range of multi-source satellite observations and climate
101 variables. In addition to incorporating direct CO_2 fertilization effects on photosynthesis, we also

¹ CEDAR stands for upsCaling Ecosystem Dynamics with ARTificial intelligence

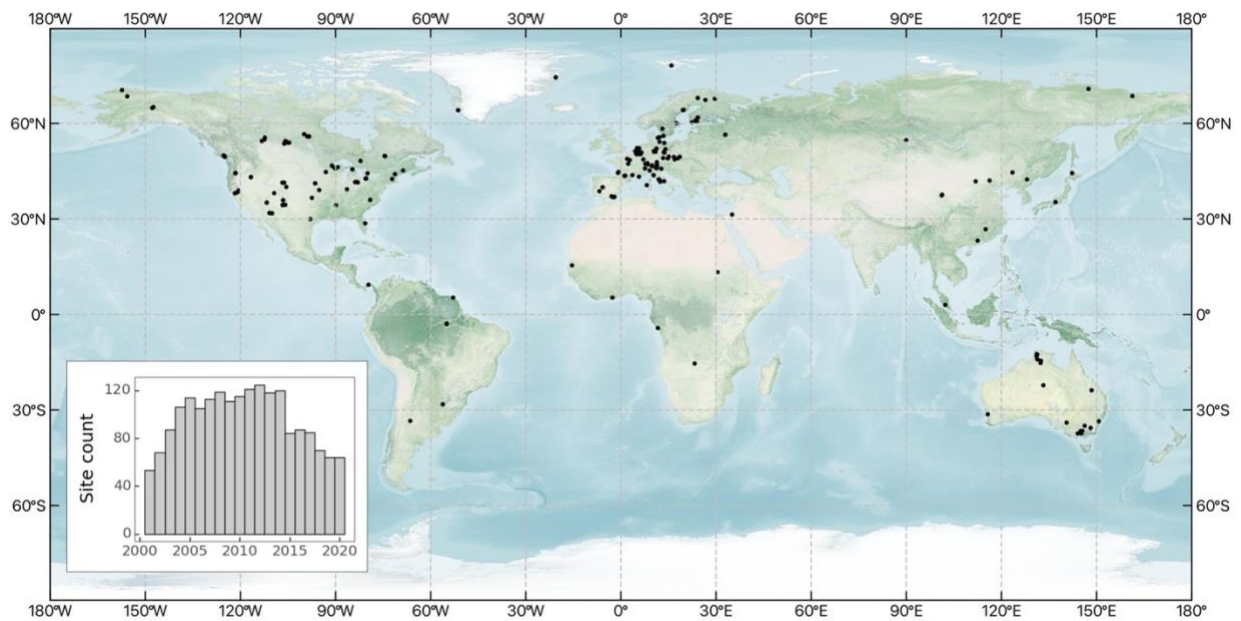
102 account for indirect effects via greenness indicators, and include novel satellite datasets such as
103 solar-induced fluorescence (SIF), Land Surface Temperature (LST) and soil moisture to explain
104 variability under environmental stresses. We provide monthly GPP estimations and associated
105 uncertainties at 0.05° resolution derived from ten model setups. These setups differ by the temporal
106 range depending on satellite data availability, the method for incorporating the direct CO₂
107 fertilization effects, and the partitioning approach used to derive GPP from eddy covariance
108 measurements. Short-term GPP datasets were from 2001 to 2020, primarily based on data derived
109 from MODIS satellites, and long-term datasets spanned 1982 to 2020 using combined Advanced
110 Very High Resolution Radiometer (AVHRR) and MODIS data. We used two approaches to
111 incorporate the direct CO₂ fertilization effects including direct prescription with eco-evolutionary
112 theory and machine learning inference from the eddy-covariance data. Additionally, we provided a
113 baseline configuration that did not incorporate the direct CO₂ effects. Uncertainties in GPP
114 estimation were quantified using bootstrapped model ensembles. We evaluate the machine learning
115 models' skills in predicting monthly GPP, seasonality, interannual variability, and trend against eddy
116 covariance measurements, and compare the CEDAR-GPP spatial and temporal variability to
117 existing satellite-based GPP estimates.

118 **2. Data and Methods**

119 **2.1 Eddy covariance data**

120 We obtained monthly eddy covariance GPP measurements from 2001 to 2020 from
121 FLUXNET2015 (Pastorello et al., 2020), AmeriFlux FLUXNET
122 (<https://ameriflux.lbl.gov/data/flux-data-products/>), and ICOS Warm Winter 2020 (Warm Winter
123 2020 Team and ICOS Ecosystem Thematic Centre., 2022) datasets. All data were processed with the
124 ONEFLUX pipeline (Pastorello et al., 2020). Following previous upscaling efforts (Tramontana et
125 al., 2016), we selected monthly GPP data that had at least 80% of high-quality hourly or half-hourly
126 data for temporal aggregation. We further excluded large negative GPP values, setting a cutoff of -1
127 gCm⁻²d⁻¹. We utilized GPP estimates from both the night-time (GPP_REF_NT_VUT) and day-time
128 (GPP_REF_DT_VUT) partitioning approaches and trained separate machine learning models for
129 each. We classified flux tower sites according to the primary C3 and C4 plant categories reported in
130 metadata and related publications when available and used a C4 plant percentage map (Still et al.,

131 2003) otherwise. Our analysis encompassed 233 sites, predominately located in North America,
 132 Western Europe, and Australia (Figure 1). In total, our dataset included roughly 18000 site-months.



133
 134 Figure 1. Global distribution of eddy covariance sites used to generate the CEDAR-
 135 GPP product. The inset displays the annual count of sites.

136 2.2 Global input datasets

137 We compiled an extensive set of covariates from gridded climate reanalysis data, multi-source
 138 satellite datasets including optical, thermal, and microwave observations, as well as categorical
 139 information on land cover, climate zone, and C3/C4 classification. The datasets that we compiled
 140 offer comprehensive information about GPP dynamics and its responses to climatic variabilities and
 141 stresses. Table 1 lists the inputs datasets and associated variables used to generate CEDAR-GPP.

142 Table 1. Datasets and input variables used to generate the CEDAR GPP product. For a list of
 143 selected variables used in different model setups, please refer to Table S1.

Category	Dataset	Temporal coverage	Spatial resolution	Temporal resolution	Variables	Reference
Climate	ERA5-Land Monthly Averaged data	1950 – present	0.1°	Monthly	Air temperature; vapor pressure deficit, Precipitation, Air and skin temperature, surface downwelling solar radiation, Potential evaporation	(Sabater, 2019)

	ESRA Global Monitoring Laboratory Atmospheric Carbon Dioxide	1976 – present	-	Monthly	Atmospheric CO ₂ concentration averaged from Mauna Loa, Hawaii, US and South Pole, Antarctica	(Thoning et al., 2021)
Satellite-based datasets	MODIS Nadir BRDF-adjusted reflectance (MCD43C4)	2000 – present	0.05°	Daily	Surface reflectance b1 – b7, Vegetation indices (NIRv, NDVI, kNDVI, EVI, GCI, NDWI), percent snow	(Schaaf and Wang, 2015)
	MODIS Terra and Aqua LAI/fPAR (MCD15A3H, MOD15A2H)	2000 – present	500m	4-day, 8-day	LAI, fPAR	(Myneni et al., 2015a, b)
	MODIS Terra and Aqua LST (MYD11A1, MOD11A1)	2000 – present	1 km	Daily	Daytime LST Nighttime LST	(Wan et al., 2015b, a)
	BESS_Rad	2000 – 2020	0.05°	Daily	PAR, diffuse PAR, downwelling solar radiation	(Ryu et al., 2018)
	Continuous-SIF (from OCO-2 and MODIS)	2000 – 2020	0.05°	4-day	all-sky daily average SIF	(Zhang, 2021)
	ESA CCI Soil Moisture Combined Passive and Active	1979 – 2021	0.25°	Daily	Surface soil moisture	(Gruber et al., 2019)
	GIMMS LAI4g	1982 – 2021	0.0833°	Half-month	LAI	(Cao et al., 2023)
	GIMMS NDVI4g	1982 – 2021	0.0833 °	Half-month	NDVI	(Li et al., 2023)
Static categorical datasets	MODIS Land Cover (MCD12Q1)	Average status used between 2001 and 2020	500m	-	Plant function types	(Friedl and Sulla-Menashe, 2019)
	Koppen-Geiger Climate Classification	present	1 km	-	Koppen-Geiger climate classes	(Beck et al., 2018)
	C4 percentage map	present	1°	-	Percentage of C4 plants	(Still et al., 2003, 2009)

145 2.2.1 Climate variables

146 We obtained air temperature, vapor pressure deficit, precipitation, potential
147 evapotranspiration, and skin temperature from the EAR5-Land reanalysis dataset (Sabater, 2019)
148 (Table 1; Table S1). We applied a three-month lag to precipitation, to reflect the memory of soil
149 moisture and represent the root zone water availability. Averaged monthly atmospheric CO₂
150 concentrations were calculated as an average of records from the Mauna Loa Observatory and South
151 Pole Observation stations, retrieved from NOAA's Earth System Research Laboratory (Thoning et
152 al., 2021).

153 2.2.2 Satellite datasets

154 We assembled a broad collection of satellite-based observations of vegetation greenness and
155 structure, LST, solar radiation, solar-induced fluorescence (SIF), and soil moisture (Table 1, Table
156 S1).

157 Three MODIS products were used: surface reflectance, LAI/fAPAR, and LST. Surface
158 reflectance from optical to infrared bands (band 1 to 7) was sourced from the MODIS Nadir
159 BRDF-adjusted reflectance (NBAR) daily dataset (MCD43C4) (Schaaf and Wang, 2015). From this,
160 we derived several vegetation indexes, including NIRv (Badgley et al., 2019), kNDVI (Camps-Valls
161 et al., 2021), NDVI, Enhanced Vegetation Index (EVI), Normalized Difference Water Index
162 (NDWI) (Gao, 1996), and a green chlorophyll index (CIgreen) (Gitelson, 2003). We also used snow
163 percentages from the NBAR dataset. We used the 4-day LAI and fPAR composite derived from
164 Terra and Aqua satellites (MCD15A3H) (Myneni et al., 2015a; Yan et al., 2016a, b) from July 2002
165 onwards and the MODIS 8-day LAI and fPAR dataset from Terra only (MOD15A2H) prior to July
166 2002 (Myneni et al., 2015b). We used day-time and night-time LST from the Aqua satellite
167 (MYD11A1) (Wan et al., 2015b), with the Terra-based LST product (MOD11A1) used after July
168 2002 (Wan et al., 2015a). Terra LST was bias-corrected with the differences in the mean seasonal
169 cycles between Aqua and Terra following Walther et al. (2021).

170 We used the PKU GIMMS NDVI4g dataset (Li et al., 2023) and PKU GIMMS LAI4g (Cao
171 et al., 2023) datasets available from 1982 to 2020. PKU GIMMS NDVI4g is a harmonized time
172 series that includes AVHRR-based NDVI from 1982 to 2003 (with biases and corrections mitigated
173 through inter-calibration with Landsat surface reflectance images) and MODIS NDVI from 2004
174 onward. PKU GIMMS LAI4g consisted of AVHRR-based LAI from 1982 to 2003 (generated using

175 machine learning models trained with Landsat-based LAI data and NDVI4g) and MODIS BNU
176 LAI from 2004 onwards (Yuan et al., 2011).

177 We utilized photosynthetically active radiation (PAR), diffusive PAR, and shortwave
178 downwelling radiation from the BESS_Rad dataset (Ryu et al., 2018). We also obtained the
179 continuous-SIF (CSIF) dataset (Zhang et al., 2018; Zhang, 2021) produced by a machine learning
180 algorithm trained using OCO-2 SIF observations and MODIS surface reflectance. We also obtained
181 surface soil moisture from the ESA CCI soil moisture combined passive and active product (Dorigo
182 et al., 2017; Gruber et al., 2019).

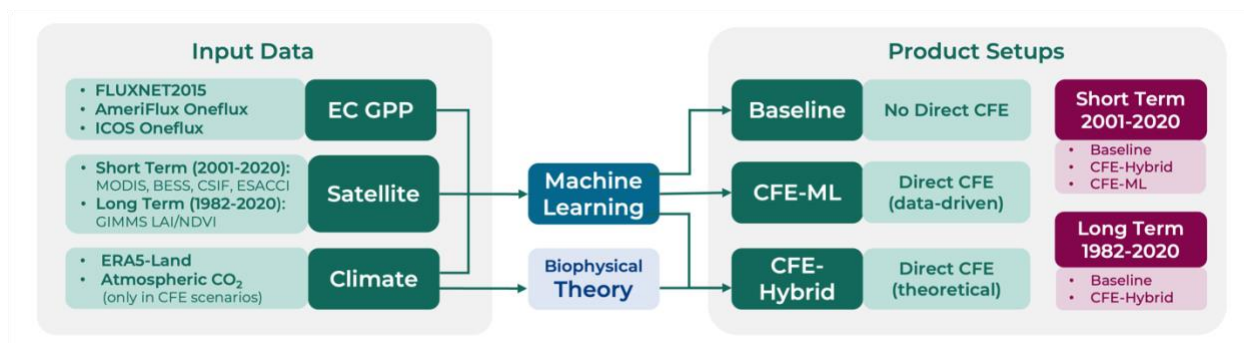
183 2.2.3 Other categorical datasets

184 We used plant functional type (PFT) information derived from the MODIS Land Cover
185 product (MCD12Q1) (Friedl and Sulla-Menashe, 2019). We followed the International Geosphere-
186 Biosphere Program classification scheme but merged several similar categories to maximize the
187 amount of eddy covariance sites/observations available for each category. Closed shrublands and
188 open shrublands are combined into a shrubland category. Woody savannas and savannas are
189 combined into savannas. We generated a static PFT map by taking the mode of the MODIS land
190 cover time series between 2001 – 2020 at each pixel to mitigate uncertainties from misclassification
191 in the MODIS dataset. Nevertheless, changes in vegetation structure induced by land use and land
192 cover change are reflected in the dynamics surface reflectance and LAI/fAPAR datasets we used.
193 We used the Koppen-Geiger main climate groups (tropical, arid, temperate, cold, and polar) (Beck et
194 al., 2018). We also utilized a C4 plant percentage map to account for different photosynthetic
195 pathways when incorporating CO₂ fertilization (Still et al., 2003, 2009).

196 2.2.4 Data preprocessing

197 We implemented a three-step preprocessing strategy for the satellite datasets: 1) quality
198 control, 2) gap-filling, and 3) spatial and temporal aggregation. In the first step, we selected high-
199 quality data based on the quality control flags of the satellite products when available. For the
200 MODIS NBAR dataset (MCD43C3), we used data with 75% or more high-resolution NBAR pixels
201 retrieved with full inversions for each band. For MODIS LST, we selected the best quality data from
202 the quality control bitmask as well as data where retrieved values had an average emissivity error of
203 no more than 0.02. For MODIS LAI/fAPAR, we used retrievals from the main algorithm with or
204 without saturation. We used all available data in ESA-CCI soil moisture due to the presence of
205 substantial data gaps. In the gap-filling step, missing values in satellite datasets were temporally filled

206 at the native temporal resolution, following a two-step protocol adapted from Walther et al (2021).
 207 Short temporal gaps were first filled with medians from a moving window, and the remaining gaps
 208 were filled with the mean seasonal cycle. For datasets with a high temporal resolution, including
 209 MODIS NBAR (daily), LAI/fPAR (4-day), BESS (4-day), CSIF (4-day), ESA-CCI (daily), temporal
 210 gaps no longer than 5 days (8 days for 4-day resolution products) were filled with medians of 15-day
 211 moving windows in the first step. An exception is MODIS LST (daily), for which we used a shorter
 212 moving window of 9 days due to rapid changes in surface temperature. GIMMS LAI4g and
 213 NDVI4g data were only filled with mean seasonal cycle due to their low temporal resolution (bi-
 214 monthly). In the last processing step, all the datasets were aggregated to a monthly time step and
 215 0.05-degree spatial resolution.



216
 217 Figure 2. Schematic overview of the CEDAR-GPP model setups.

218 2.3 Machine learning upscaling

219 2.3.1 CEDAR-GPP model setups

220 We trained machine learning model with eddy covariance GPP measurements as targets and
 221 climate/satellite variables as input features. We created ten model setups to produce ten different
 222 global monthly GPP datasets (Figure 2; Table 2). The model setups were characterized by the
 223 temporal range of input datasets used, the configuration of CO₂ fertilization effects, and the
 224 partitioning approach used to derive the GPP from eddy covariance measurements.

225 We provided a short-term (ST) configuration producing GPP estimates from 2001 to 2020,
 226 and a long-term (LT) configuration spanning 1982 to 2020. Each temporal configuration uses a
 227 different set of input variables depending on their temporal availability. Inputs for the short-term
 228 configuration included MODIS, CSIF, BESS PAR, ESA-CCI soil moisture, ERA5-Land, as well as
 229 PFT and Koppen Climate zone as categorical variables with one-hot encoding. The long-term used
 230 GIMMS NDVI4g and LAI4g data, ERA5-land, PFT and Koppen climate. ESA CCI soil moisture

231 datasets were excluded from the long-term model setups due to concerns about the product quality
 232 in the early years when the number and quality of microwave satellite data were limited (Dorigo et
 233 al., 2015). A detailed list of input features for each setup is provided in Table S1.

234 Regarding the direct CO₂ fertilization effects (CFE), we established a “Baseline” configuration
 235 that did not incorporate these effects, a “CFE-Hybrid” configuration that incorporated the effects
 236 via eco-evolutionary theory, and a “CFE-ML” configuration that inferred direct effects from eddy
 237 covariance data using machine learning. Detailed information about these approaches is provided in
 238 Section 2.4.2. Furthermore, separate models were trained for GPP target variables from the night-
 239 time (NT) and daytime (DT) partitioning approaches.

240 Table 2 lists the characteristics of ten model setups. Note due to the limited availability of
 241 eddy covariance observations before 2001, we did not apply the CFE-ML approach to the long-term
 242 setups, as the machine learning inferred CO₂ fertilization effects cannot be robustly extrapolate
 243 GPP back to 1982.

244 Table 2. Specifications of the CEDAR-GPP model setups.

Model Setup Name	Temporal range	Direct CO ₂ Fertilization Effects		GPP Partitioning Method
		Configuration	Method	
ST_Baseline_NT	Short-term (ST) 2001 – 2020	Baseline	Not incorporated	Night-time (NT)
ST_Baseline_DT				Day-time (DT)
ST_CFE-Hybrid_NT		CFE-Hybrid	Theoretical	NT
ST_CFE-Hybrid_DT				DT
ST_CFE-ML_NT		CFE-ML	Data-driven	NT
ST_CFE-ML_DT				DT
LT_Baseline_NT	Long-term (LT) 1982 – 2020	Baseline	Not incorporated	NT
LT_Baseline_DT				DT
LT_CFE-Hybrid_NT		CFE-Hybrid	Theoretical	NT
LT_CFE-Hybrid_DT				DT

245

246 2.3.2 CO₂ fertilization effect

247 We established three configurations considering the direct CO₂ fertilization effects on
 248 photosynthesis. In the baseline configuration, we trained machine learning models with eddy
 249 covariance GPP measurements, input climate and satellite features, but excluding CO₂
 250 concentration. As such, the models only include indirect CO₂ effects from the satellite-based proxies
 251 of vegetation greenness and structure and do not consider the direct effect of CO₂ on light use
 252 efficiency. Our baseline model is therefore directly comparable to other satellite-derived GPP
 253 products that only account for indirect CO₂ effects (ref FLUXCOM, FLUXSAT, MODIS).

254 In the CFE-ML configuration, we added trained monthly CO₂ concentration into the feature
255 set in addition to those incorporated in the baseline models. Thus, models inferred the functional
256 relationship between GPP and CO₂ from the eddy covariance data, encompassing both CO₂
257 fertilization pathways – direct effects on LUE and indirect effects from the satellite-based proxies of
258 vegetation greenness and structure.

259 In the CFE-Hybrid configuration, we applied biophysical theory to estimate the response of
260 LUE to elevated CO₂. First, we estimated a reference GPP, where LUE is not affected by any
261 increase in atmospheric CO₂, by applying the CFE-ML model with a constant atmospheric CO₂
262 concentration equal to the 2001 level while keeping all other variables temporally dynamic. Then, the
263 impacts of CO₂ on LUE were prescribed onto the reference GPP estimates using a theoretical CO₂
264 sensitivity function of LUE according to eco-evolutionary theories (Supplementary Text S1). The
265 theoretical CO₂ sensitivity function represents a CO₂ sensitivity that is equivalent to that of the
266 electron-transport-limited (light-limited) photosynthetic rate. When light is limited, elevated CO₂
267 suppresses photorespiration leading to increased photosynthesis at a lower rate than when
268 photosynthesis is limited by CO₂ (Lloyd and Farquhar, 1996; Smith and Keenan, 2020). Thus, the
269 CFE-Hybrid scenario provides a conservative estimation of the direct CO₂ effects on LUE. Note
270 that the theoretical sensitivity function describes the fractional change in LUE due to direct CO₂
271 effects relative to a reference period (i.e. 2001). Therefore, we used the CFE-ML model to establish
272 this reference GPP by fixing the CO₂ effects to the 2001 level, rather than simply using the GPP
273 from the Baseline model in which the direct CO₂ effects were not clearly represented.

274 For both CFE-ML and CFE-Hybrid scenarios, we made another conservative assumption that
275 C4 plants do not benefit from elevated CO₂, despite potential increases in photosynthesis during
276 water-limited conditions due to enhanced WUE. Data from flux tower sites dominated by C4 plants
277 were removed from our training set, so the machine learning models inferred CO₂ fertilization only
278 from flux tower sites dominated by C3 plants. When applying models globally, we assumed the
279 reference GPP values (with constant atmospheric CO₂ concentration equal to the 2001 level) to
280 represent C4 plants, and GPP estimates from CFE-ML or CFE-Hybrid models were applied in
281 proportion to the percentage of C3 plants in a grid cell.

282 2.3.3 Machine learning model training and validation

283 We employed the state-of-the-art XGBoost machine learning model, known for its high
284 accuracy in regression problems across various domains, including environmental and ecological

285 predictions (Chen and Guestrin, 2016; Kang et al., 2020; Berdugo et al., 2022). XGBoost is a
286 scalable and parallelized implementation of the gradient boosting technique that iteratively trains an
287 ensemble of decision trees, with each iteration targeting to minimize the residuals from the last
288 iteration. A notable merit of XGBoost is its ability to make prediction in the presence of missing
289 values, a common issue in remote sensing datasets.

290 We used five-fold cross-validation for model evaluation. Training data was randomly split into
291 five groups (folds), with each fold held out for testing while the rest four folds were used for model
292 training. We imposed two restrictions on fold splitting: each flux site was entirely assigned to a fold
293 to test model performance over unseen locations; the random sampling was stratified based on PFT
294 to ensure coverage of the full range of PFTs in both training and testing. Within each training set,
295 we performed a randomized search using three-fold cross-validation to determine the optimal
296 hyperparameter set, to reduce the risk of overfitting and improve the robustness of the evaluation.

297 We assessed the models' ability to capture the temporal and spatial characteristics of GPP,
298 including monthly variabilities, mean seasonal cycles, monthly anomalies, cross-site variability.
299 Model performance was assessed separately for each model setup (Table 2) and summarized by PFT
300 and Koppen climate zone. Mean seasonal cycles were calculated as the mean monthly GPP over the
301 site observation period, and monthly anomalies were the residuals of monthly GPP after subtracting
302 mean seasonal cycles. Monthly GPP averaged over years for each site was used to assess cross-site
303 variability. Goodness-of-fit metrics include RMSE, bias, and coefficient of determination (R^2 ,
304 equivalent to NSE Nash–Sutcliffe model efficiency coefficient).

305 To evaluate the models' ability to capture long-term GPP trends, we aggregated the monthly
306 GPP to annual values for sites with at least 5 years of observations following Chen et al. (2022).
307 GPP anomalies were computed by subtracting the multi-year mean GPP from the annual GPP for
308 each site. Anomalies were aggregated across site to achieve a single multi-site GPP anomaly per year.
309 We used the Sen slope and Mann-Kendall test to examine the GPP trends from 2002 to 2019,
310 excluding 2001 and 2020 due to the limited number of available sites.

311 2.3.4 Product generation and uncertainty quantification

312 In the CEDAR-GPP product, we generated a GPP dataset for each of the ten model setups,
313 by applying the model to global gridded datasets within the corresponding temporal range (Table 2).
314 GPP datasets were named after the corresponding model setups. For each model setup, we first
315 generated 30 sample set using bootstrapping, which were then used to train an ensemble of 30

316 XGBoost models. The bootstrapping was performed at the site level, and each bootstrapped sample
317 set contained around 140 to 150 unique sites, 17000 to 19000 site months covering all PFTs. The
318 relative composition of sites in each PFT was consistent with the full dataset. The 30 models trained
319 with bootstrapped samples generated an ensemble of 30 GPP values. We provided the ensemble
320 GPP mean and standard deviation from each of the ten model setups.

321 2.4 Product inter-comparison

322 We compared the global spatial and temporal patterns of CEDAR-GPP with other major
323 satellite-based GPP products including three machine learning upscaled and two LUE-based
324 datasets. We obtained two FLUXCOM products (Jung et al., 2020), the latest version of
325 FLUXCOM-RS (FLUXCOM-RSv006) available from 2001 to 2020 based on remote sensing
326 (MODIS collection 6) datasets only, as well as the FLUXCOM-RS+METEO ensemble available
327 between 1979 to 2018 and based on the climatology of remote sensing observations and ERA5
328 forcings (hereafter FLUXCOM-ERA5). We used FluxSat (Joiner and Yoshida, 2020), available from
329 2001 to 2019, which is an upscaled dataset based on MODIS NBAR surface reflectance and PAR
330 from Modern-Era Retrospective analysis for Research and Applications 2 (MERRA-2). Importantly,
331 FluxSat does not incorporate climate forcings. We used the MODIS GPP product (MOD17)
332 available since 2001, which was generated based on MODIS fAPAR and LUE as a function of air
333 temperature and vapor pressure deficit but not atmospheric CO₂ concentration (Running et al.,
334 2015). We also used the rEC-LUE products, available from 1982 to 2018 and based on a revised
335 LUE model that incorporated the effect of atmospheric CO₂ concentration and the fraction of
336 diffuse PAR on LUE in addition to air temperature and vapor pressure deficit (Zheng et al., 2020).
337 All datasets were resampled to 0.1 ° spatial resolution, and a common mask for the vegetated land
338 area was applied. We evaluated global mean annual GPP, mean seasonal cycle, interannual
339 variability, and trend among different datasets, comparing them over a common time period
340 determined by their data availability. Global total GPP was computed by scaling the global average
341 GPP flux with the global land area (122.4 million km²) following Jung et al. (2020). Mean seasonal
342 cycle was defined as above (Sec 2.3.3). We used the standard deviation of annual GPP to indicate the
343 magnitude of interannual variability, the Sen slope to indicate GPP annual trend and the Mann-
344 Kendall test for the statistical significance of trends.

345 3. Results

346 3.1 Evaluation of model performance

347 3.1.1 Overall performance

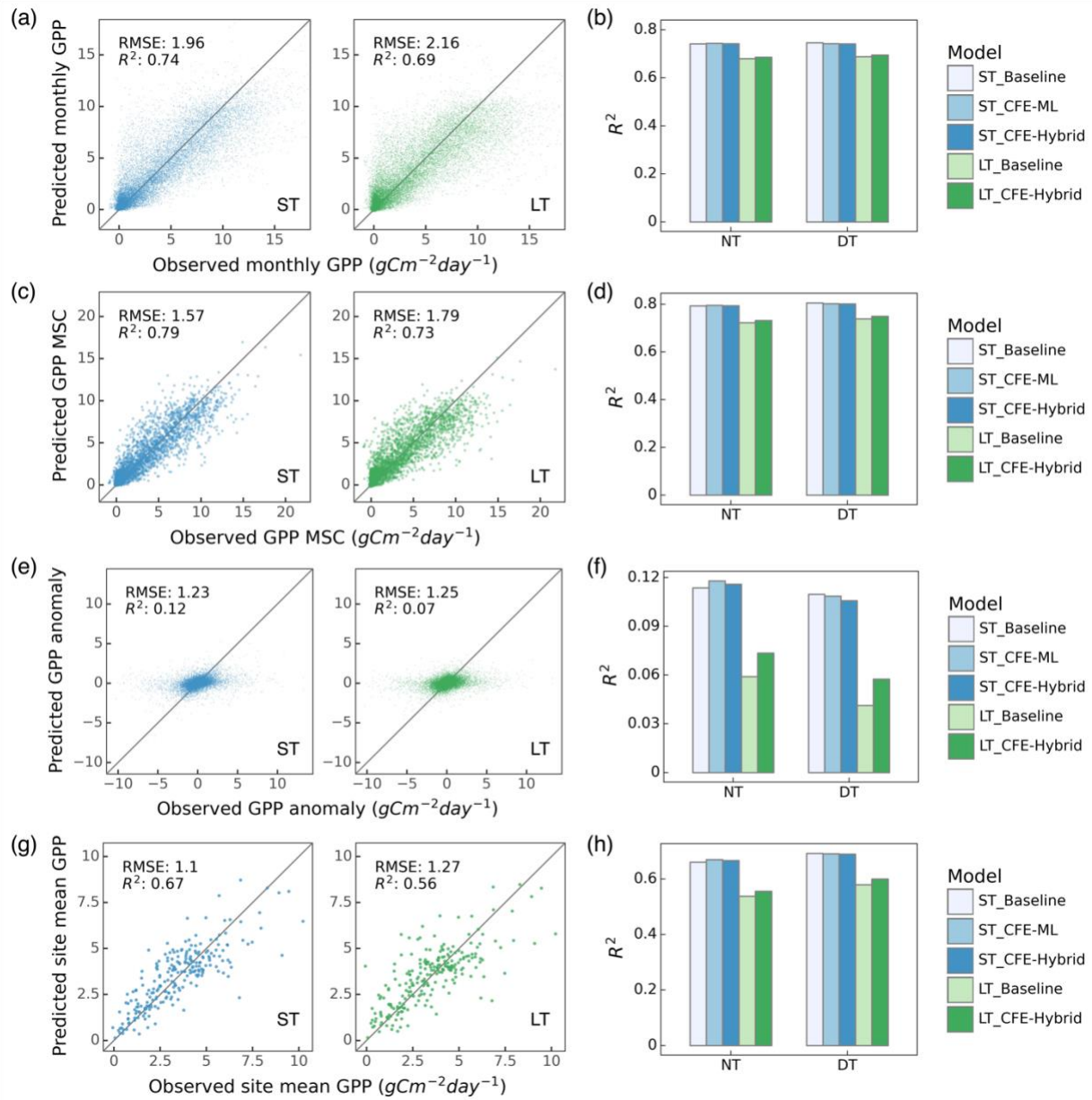
348 The short-term and long-term models explained approximately 74% and 68%, respectively, of
 349 the variation in eddy covariance estimated monthly GPP across global sites (Figure 3a). The long-
 350 term models consistently yielded lower performance than the short-term models, likely due to
 351 differences in the satellite remote sensing datasets used, as the short-term models benefited from a
 352 richer information including surface reflectance from individual bands, LST, CSIF, as well as soil
 353 moisture, while the long-term model only exploited NDVI and LAI. The models with different CFE
 354 scenarios and target GPP variables had similar performance on predicting monthly GPP (Figure 3b,
 355 Table 3, Table S2). All models exhibited minimal bias less than 0.15.

356 Model performance in terms of the different temporal and spatial characteristics of monthly
 357 GPP was variable (Figure 3c-h). The models were most successful at predicting mean seasonal
 358 cycles, with the short-term and long-term models explaining around 79% and 72% of the variability,
 359 respectively (Figure 3c-d). The short-term and long-term models captured 66% and 54% ,
 360 respectively of the spatial variabilities of multi-year mean GPP across global sites (i.e., cross-site
 361 variability) (Figure 3g-h). However, all models predicted monthly anomalies across the sites, with R²
 362 values below 0.11 (Figure 3e-f). The CFE-ML and CFE-Hybrid models showed slightly higher
 363 accuracy than the Baseline model across all temporal and spatial characteristics.

364 Table 3. Machine learning model performance for five CEDAR-GPP setups based on NT GPP
 365 (Table 2). Results of DT setups can be found in Table S2.

Model Setup Name	Monthly			Mean seasonal cycles			Monthly anomalies			Cross-site		
	RMSE	Bias	R ²	RMSE	Bias	R ²	RMSE	Bias	R ²	RMSE	Bias	R ²
ST_Baseline_NT	1.96	-0.05	0.74	1.57	0.02	0.79	1.22	0.00	0.11	1.11	0.03	0.66
ST_CFE-ML_NT	1.95	-0.05	0.74	1.56	0.02	0.80	1.22	0.00	0.12	1.10	0.03	0.67
ST_CFE-Hybrid_NT	1.96	-0.05	0.74	1.57	0.03	0.79	1.23	0.00	0.12	1.10	0.04	0.67
LT_Baseline_NT	2.18	-0.10	0.68	1.82	0.01	0.72	1.26	0.00	0.06	1.29	0.03	0.54
LT_CFE-Hybrid_NT	2.16	-0.11	0.69	1.79	0.01	0.73	1.25	0.00	0.07	1.27	0.03	0.56

366



367

368

369

370

371

372

373

374

Figure 3. Machine learning model performance in predicting monthly GPP and its spatial and temporal variability. Scatter plots illustrated relationships between model predictions and observations for monthly GPP (a), mean seasonal cycles (MSC (c), monthly anomaly (e), and cross-site variability (g) for ST_CFE-Hybrid_NT (left, blue) and LT_CFE-Hybrid_NT (right, green) models. Corresponding bar plots show the R^2 values for all ten model setups in predicting monthly GPP (b), MSC (d), monthly anomaly (f), and cross-site variability (h).

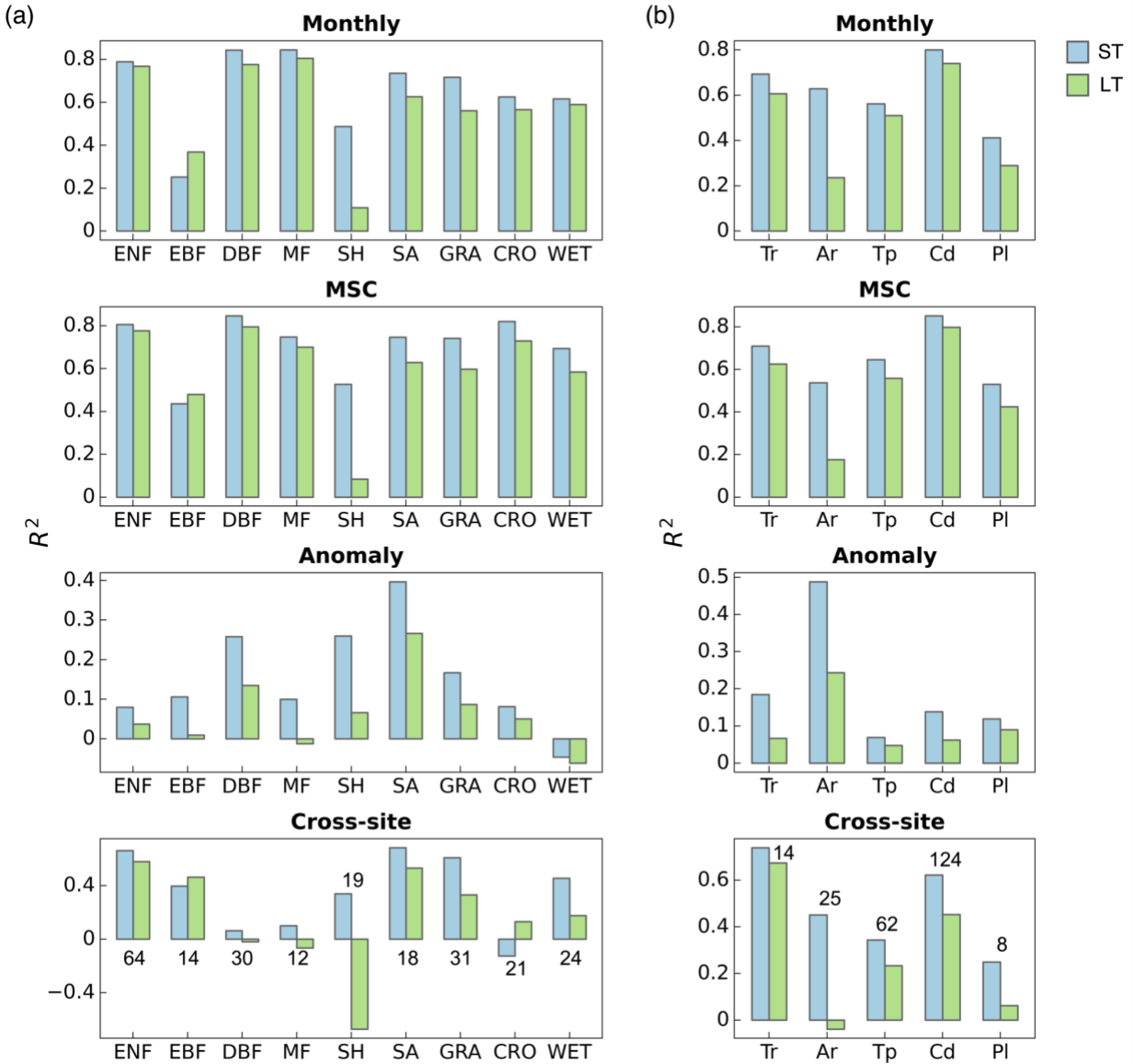
375 3.1.2 Performance by biome and climate zone

376 The predictive ability of our models varied across different PFTs and Koppen climate zones
377 (Figure 4). Here we present results from the CFE-Hybrid models and note that patterns for the
378 other CFE configurations were similar.

379 Model performance in terms of monthly GPP was highest for forests with distinct seasonality,
380 including deciduous broadleaf forests, mixed forests, and evergreen needleleaf forests, with R^2 values
381 above 0.78. Model accuracies were also high for savannas, and grasslands, followed by croplands and
382 wetlands, with R^2 values between 0.57 and 0.74. Model accuracies were lowest in evergreen
383 broadleaf forests and shrublands, with R^2 values as low as 0.14. Across climate zones, models
384 achieved the highest accuracy in predicting monthly GPP in cold and tropical climate zones with R^2
385 values between 0.64 and 0.80. The short-term models had lowest performance in polar regions with
386 an R^2 value around 0.42 and the long-term model had the lowest performance in arid regions with
387 an R^2 value of 0.25.

388 Model performance in terms of mean seasonal cycles across PFTs and climate zones followed
389 patterns for monthly GPP, while disparities emerged for performance in terms of GPP anomaly and
390 cross-site variability (Figure 4). The short-term model showed the highest predictive power in
391 explaining monthly anomalies in arid regions with an R^2 value of 0.49, where savanna and
392 shrublands sites are primarily located. Model performance in all other climate zones was significantly
393 lower with R^2 values below 0.2, and as low as 0.07 in temperate regions. Besides, the short-term
394 model demonstrated good performance in capturing anomalies in deciduous broadleaf forests. The
395 long-term model's relative performance between PFTs and climate zones was mostly consistent with
396 that of the short-term model, with lower accuracy in shrublands when compared to the short-term
397 model.

398 Model performance in terms of cross-site variability demonstrated highest accuracy in
399 savannas, grasslands, evergreen needleleaf forests, and evergreen broadleaf forests ($R^2 > 0.36$) and
400 lowest accuracy in deciduous broadleaf forests, mixed forests, and croplands ($R^2 < 0.20$). The short-
401 term model additionally showed good performance in shrublands and wetlands ($R^2 > 0.36$), whereas
402 the long-term model failed to capture any variability for shrublands. In terms of climate zones,
403 models were most successful at explaining the variabilities across tropical and cold climate zones
404 ($R^2 > 0.46$), the short-term model was least successful across polar regions, with a R^2 value of 0.29,
405 and the long-term model had low performance for both polar and arid regions with R^2 values below
406 0.15.



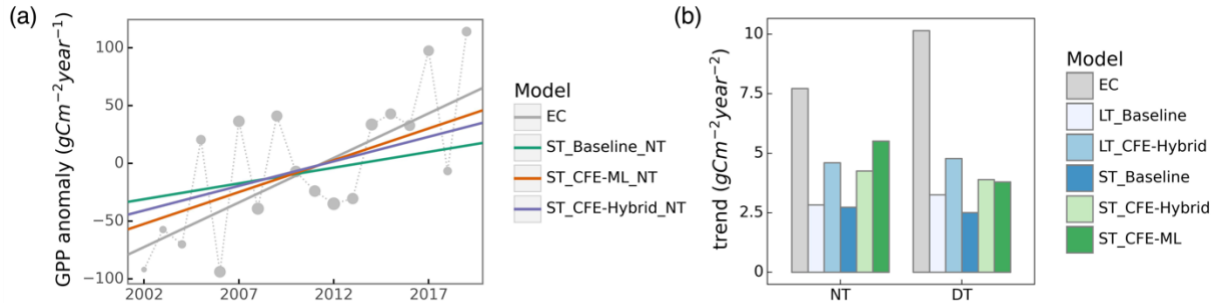
407
 408 Figure 4. Performance of the ST_Baseline_NT (blue) and LT_Baseline_NT (green)
 409 models on GPP spatiotemporal estimation by plant functional types (a) and climate
 410 zones (b). The cross-site panels included the number of sites within each category.
 411 ENF: evergreen needleleaf forest, EBF: evergreen broadleaf forest, DBF: deciduous
 412 broadleaf forest, MF: mixed forest, SH: shrubland, SA: savanna, GRA: grassland, CRO:
 413 cropland, WET: wetland. Tr: tropical, Ar: arid, Tp: temperate, Cd: cold, Pl: polar.

414 3.1.3 Prediction of long-term trends

415 Eddy covariance measured GPP presented a substantial increasing trend in GPP across flux
 416 sites between 2002 and 2019 (Figure 5a). The observed GPP from the night-time partitioning
 417 approach indicated an overall trend of $7.7 \text{ gCm}^{-2}\text{year}^{-2}$. In contrast, the ST_Baseline_NT model

418 predicted a more modest trend of $2.7 \text{ gCm}^{-2}\text{year}^{-2}$, primarily reflecting the indirect CO_2 effect
 419 manifested through the growth of LAI. Both the ST_CFE-ML_NT and ST_CFE-hybrid_NT
 420 models predicted much higher trends of 5.5 and $4.3 \text{ gCm}^{-2}\text{year}^{-2}$ respectively, aligning more closely
 421 to eddy covariance observations.

422 Across all model setups, the CFE-ML and CFE-hybrid models consistently outperformed the
 423 Baseline models in predicting GPP trends in global eddy covariance towers (Figure 6b) and all
 424 trends were statistically significant ($p < 0.05$). Notably, we found a considerably higher trend in eddy
 425 covariance GPP measurements derived from the day-time versus night-time partitioning approach.
 426 The predicted trends of different model setups between the partitioning approaches were similar
 427 despite a smaller trend predicted by the ST_CFE-ML_DT model compared to the corresponding
 428 NT model (Figure 6b).



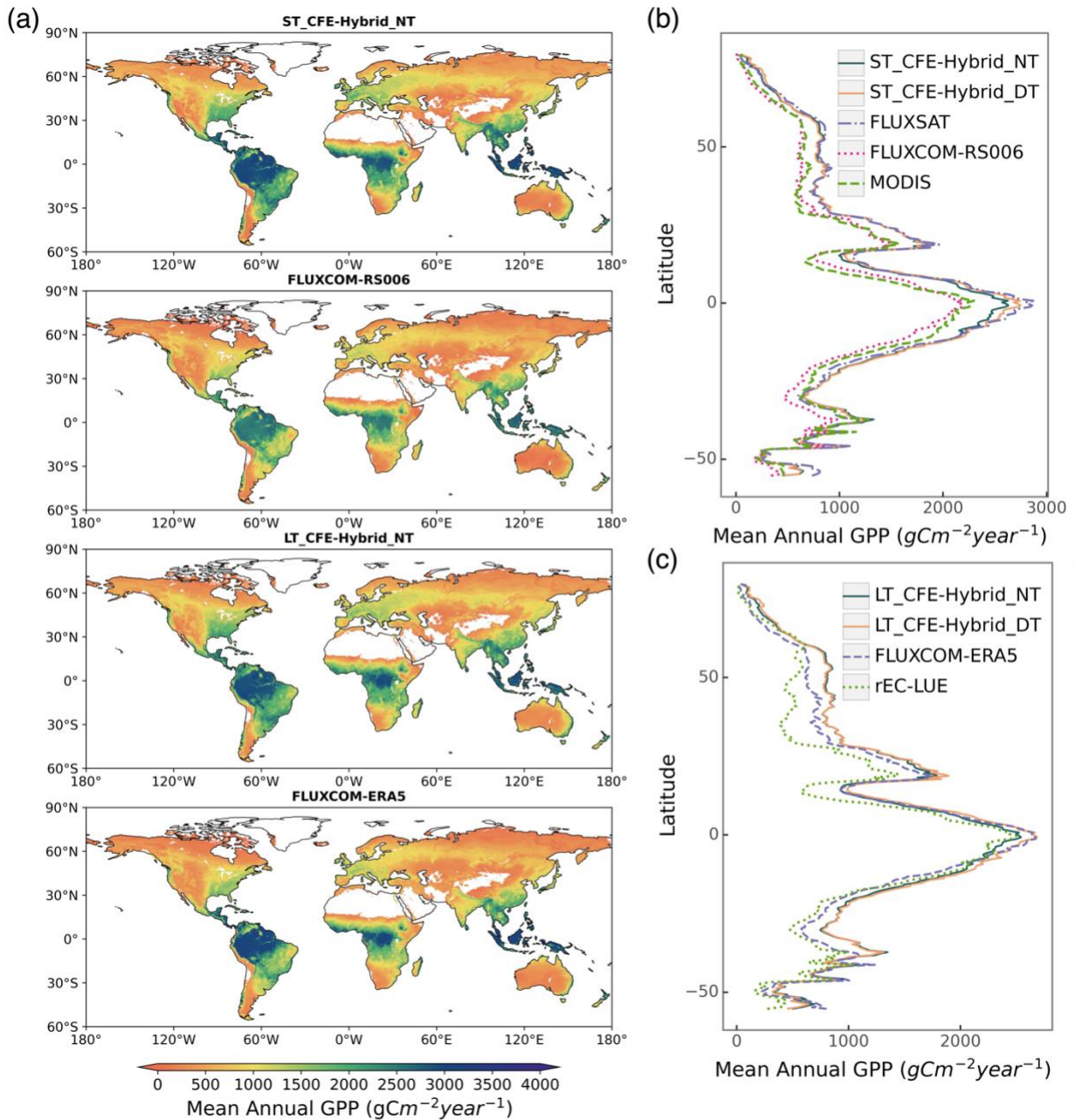
429
 430 Figure 5. Comparison of observed and predicted GPP trends across eddy covariance
 431 flux towers. (a) Aggregated annual GPP anomaly from 2002 to 2019 and trend lines
 432 from eddy covariance (EC) measurements, and three CFE model setups (short-term,
 433 night-time partitioning). Size of the grey circle markers is proportional to the number
 434 of sites. (b) Annual trends from eddy covariance measurements and ten CEDAR-GPP
 435 model setups.

436 3.2 Evaluation of GPP spatial and temporal dynamics

437 In this section, we present comparisons between CEDAR-GPP datasets and other upscaled or
 438 LUE-based datasets regarding the mean annual GPP (Section 3.2.1), GPP seasonality (Section 3.2.2),
 439 interannual variability (Section 3.2.3), and annual trends (Section 3.2.4). CEDAR-GPP model setups
 440 generally showed similar patterns in mean annual GPP, seasonality, and interannual variability,
 441 therefore, in corresponding sections, we present the CFE-Hybrid model setups as representative
 442 examples for comparisons with other independent datasets, unless otherwise stated. Supplementary
 443 figures include comparisons involving all CEDAR-GPP.

444 3.2.1 Mean annual GPP

445 Global patterns of mean annual GPP were generally consistent among CEDAR-GPP model
446 setups, FLUXCOM, FLUXSAT, MODIS, and rEC-LUE, with few noticeable regional differences
447 (Figure 6, Figure S1). Differences among CEDAR-GPP model setups were minimal and only
448 evident between the NT and DT setups in the tropics (Figure 6b-c, Figure S1). CEDAR-GPP short-
449 term datasets showed highest consistency with FLUXSAT in terms of mean annual GPP
450 magnitudes (2001 – 2018) and latitudinal variations, although FLUXSAT presented slightly higher
451 GPP values in the tropics compared to CEDAR-GPP (Figure 6b). Mean annual GPP magnitude for
452 FLUXCOM-RS006 and MODIS tended to be lower globally than CEDAR-GPP and FLUXSAT,
453 with the most pronounced differences observed in the tropical areas. Among the long-term datasets
454 (CEDAR-GPP LT, FLUXCOM-ERA5, and rEC-LUE), mean annual GPP (1982 – 2018) exhibited
455 greater disparities in the northern mid-latitudes than in the tropics and southern hemisphere (Figure
456 6c). CEDAR-GPP aligned more closely with FLUXCOM-ERA5 than with rEC-LUE, with the latter
457 showing lower annual mean GPP globally, particularly between 20°N to 50° N.



458

459

460

461

462

463

464

465

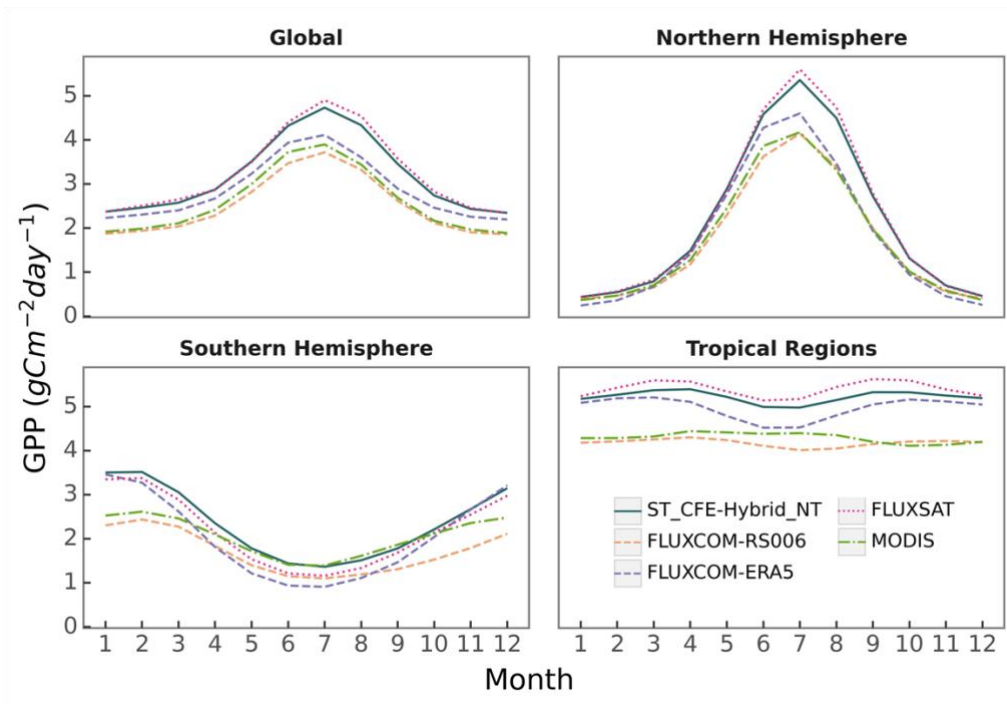
466

Figure 6. Global distributions of mean annual GPP from CEDAR-GPP and other machine learning upscaled and LUE-based reference datasets. (a) Global patterns of mean annual GPP from ST_CFE-Hybrid_NT, FLUXCOM-RS006, LT_CFE-Hybrid_NT, and FLUXCOM-ERA5. (b) Latitudinal distributions of mean annual GPP (2001 – 2018) from short-term datasets (ST_CFE-Hybrid_NT, ST_CFE-Hybrid_DT, FLUXSAT, FLUXCOM-RS006, and MODIS). (c) Latitudinal distributions of mean annual GPP (1982 – 2018) from long-term datasets (LT_CFE-Hybrid_NT, LT_CFE-Hybrid_DT, FLUXCOM-ERA5, and rEC-LUE).

467 3.2.2 Seasonal variability

468 CEDAR-GPP and other machine learning upscaled and LUE-based GPP datasets agreed on
469 seasonal variabilities (average between 2001 and 2018) at the global scale, characterized by a peak in
470 GPP in July and a nadir between December and January (Figure 7). At the global scale, CEDAR-
471 GPP was most closely aligned with FLUXSAT in GPP seasonal magnitude and amplitude, while
472 both FLUXCOM and MODIS displayed a relatively less pronounced magnitude.

473 In the northern hemisphere (20°N - 90°N), all GPP datasets agreed in seasonal GPP variation,
474 despite variances in the magnitude of peak GPP. In the southern hemisphere (20°S - 60°S), all
475 datasets exhibited their lowest GPP during June and July, and highest GPP from December to
476 January. However, the seasonal amplitude of GPP was greatest for FLUXCOM-ERA5, followed by
477 CEDAR-GPP and FLUXSAT, and substantially smaller for FLUXCOM-RS006 and MODIS GPP.
478 In the tropics (20°N - 20°S), differences between datasets were the strongest, where seasonal
479 variation is not as prominent compared to other regions. CEDAR-GPP, FLUXSAT, and
480 FLUXCOM-ERA5 each showed two GPP peaks, occurring in March-April and September-
481 October. Although FLUXCOM-RS006 had a similar seasonal pattern, its GPP magnitude was
482 markedly smaller. Interestingly, MODIS showed an inverse season pattern with a small peak from
483 June to August.

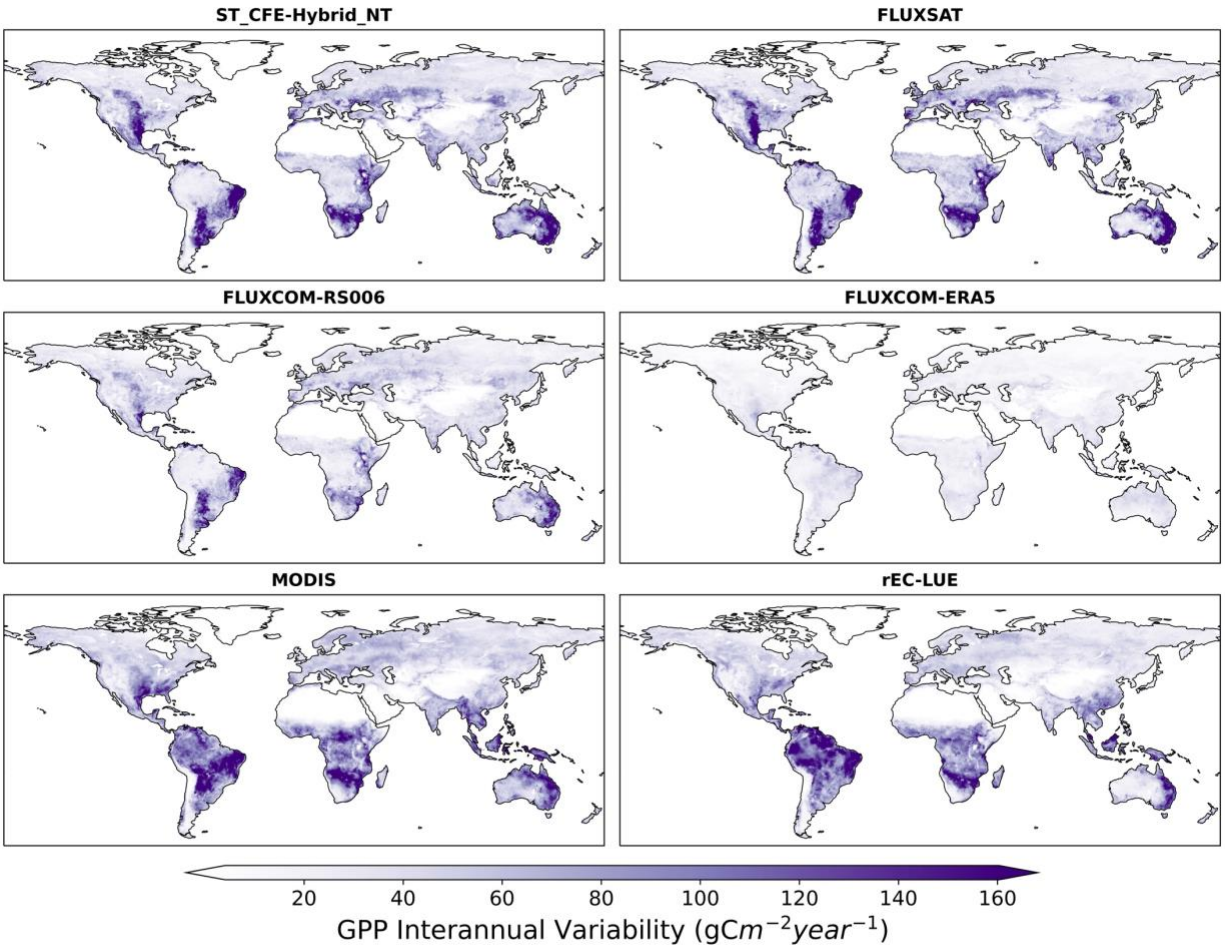


484
485 Figure 7. Comparison of GPP mean seasonal cycle between different datasets on a
486 global scale, and specifically with the Northern Hemisphere (20°N - 90°N), Southern

487 Hemisphere (20°S - 60°S), and Tropical regions (20°N - 20°S). Monthly means were
488 averaged from 2001 to 2018 for all datasets.

489 3.2.3 Interannual variability

490 We found distinct spatial patterns in GPP interannual variability between upscaled and LUE-
491 based datasets and a high level of agreement within each category, with the exception of
492 FLUXCOM-ERA5, which showed minimal interannual variability globally (Figure 8). All datasets
493 agreed on the presence of GPP interannual variability hotspots in eastern and southern South
494 America, central North America, southern Africa, and western Australia. These hotspots primarily
495 corresponded to arid and semi-arid areas characterized by grasslands, shrubs, and croplands (Figure
496 9). CEDAR-GPP was highly consistent with FLUXSAT, and both datasets also displayed relatively
497 high interannual variability in the dry subhumid areas of Europe, predominately covered by
498 croplands. FLUXCOM-RS006 mirrored the relative spatial patterns of CEDAR-GPP and
499 FLUXSAT, albeit at lower magnitudes. The LUE-based datasets (MODIS and rEC-LUE) predicted
500 a much higher interannual variability than the upscaled datasets in the tropical areas, particularly in
501 evergreen broadleaf forests and woody savannas (Figure 8, Figure 9). These datasets also depicted
502 slightly higher interannual variability for other types of forests, including evergreen needleleaf forests
503 and deciduous broadleaf forests, compared to the upscaled datasets.



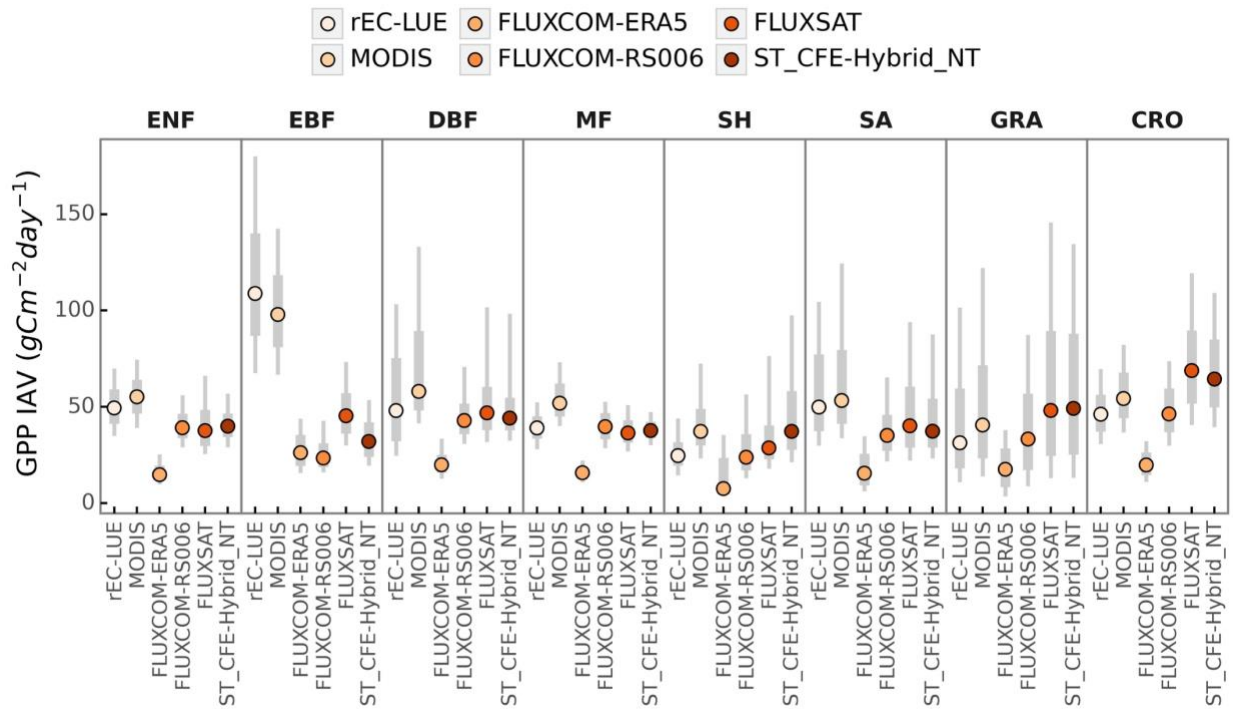
504

505

506

507

Figure 8. Spatial patterns of GPP interannual variability extracted over 2001 to 2018 for CEDAR-GPP (ST_CFE-Hybrid_NT), FLUXSAT, FLUXCOM-RS006, MODIS, FLUXCOM-ERA5, and rEC-LUE.



508

509 Figure 9. Comparison of GPP interannual variability (IAV across global datasets by
 510 PFT. Colored dots represent the median IAV, thicker grey bars indicate the 25% to
 511 75% percentiles of IAV distributions, and thinner grey bars show the 10% to 90%
 512 percentiles.

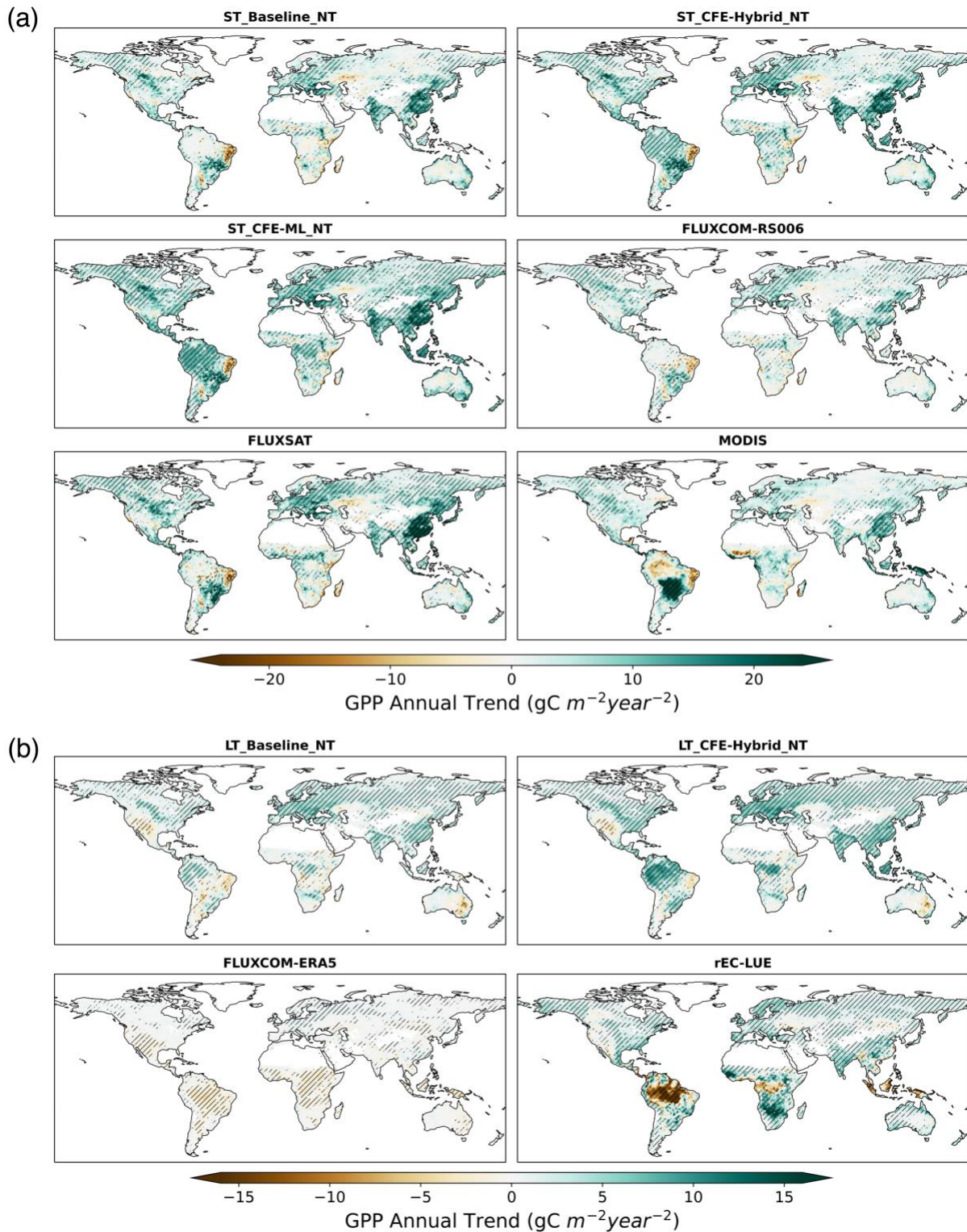
513 3.2.4 Trends

514 Differences in annual GPP trends among CEDAR-GPP model setups and other upscaled and
 515 LUE-based datasets mainly reflected the variability in the representation of CO₂ fertilization effects
 516 (Figure 10, Figure S4). From 2001 to 2018, the CEDAR-GPP Baseline model setups showed spatial
 517 variations in GPP trends consistent with the other upscaled datasets without direct CO₂ fertilization
 518 effects, including FLUXSAT and FLUXCOM-RS. In these datasets, substantial increases were seen
 519 in southeastern China and India, western Europe, and part of North and South America. These
 520 increases were largely associated with rising LAI due to land use changes and indirect CO₂
 521 fertilization effects, as identified by previous studies (Zhu et al., 2016; Chen et al., 2019). Although
 522 MODIS, which also does not include a direct CO₂ fertilization effect, generally agreed with these
 523 increasing trends, it also showed a declining GPP in the tropical Amazon and a stronger positive
 524 trend in central South America. After incorporating the direct CO₂ fertilization effects, both the
 525 CFE-Hybrid and CFE-ML setups predicted positive trends in tropical forests, an observation absent
 526 in all other datasets. Furthermore, the CFE-Hybrid and CFE-ML models also revealed increasing

527 GPP in temperate and boreal forests of North America and Eurasia. Notably, all datasets agreed on
528 a pronounced GPP decrease in eastern Brazil.

529 From 2001 to 2018, a positive trend in global annual GPP was uniformly detected by all, albeit
530 with varying magnitudes (Figure 11a-b). The ST_Baseline_NT model predicted a GPP growth rate
531 of 0.35 Pg C per year, aligning with FLUXCOM-RS, but lower than FLUXSAT (0.51 Pg C yr⁻¹) and
532 MODIS (0.39Pg C yr⁻¹) (Figure 11b). The CFE-hybrid models estimated a notably faster GPP
533 growth at 0.58 Pg C per year. The CFE-ML models predicted the highest trends, up to 0.76 from
534 the ST_CFE-ML_NT model and 0.59 from the ST_CFE-ML_DT model. Also, a higher variance
535 was observed among ensemble members in the ST_CFE-ML setups compared to the ST_Baseline
536 and ST_CFE-Hybrid models.

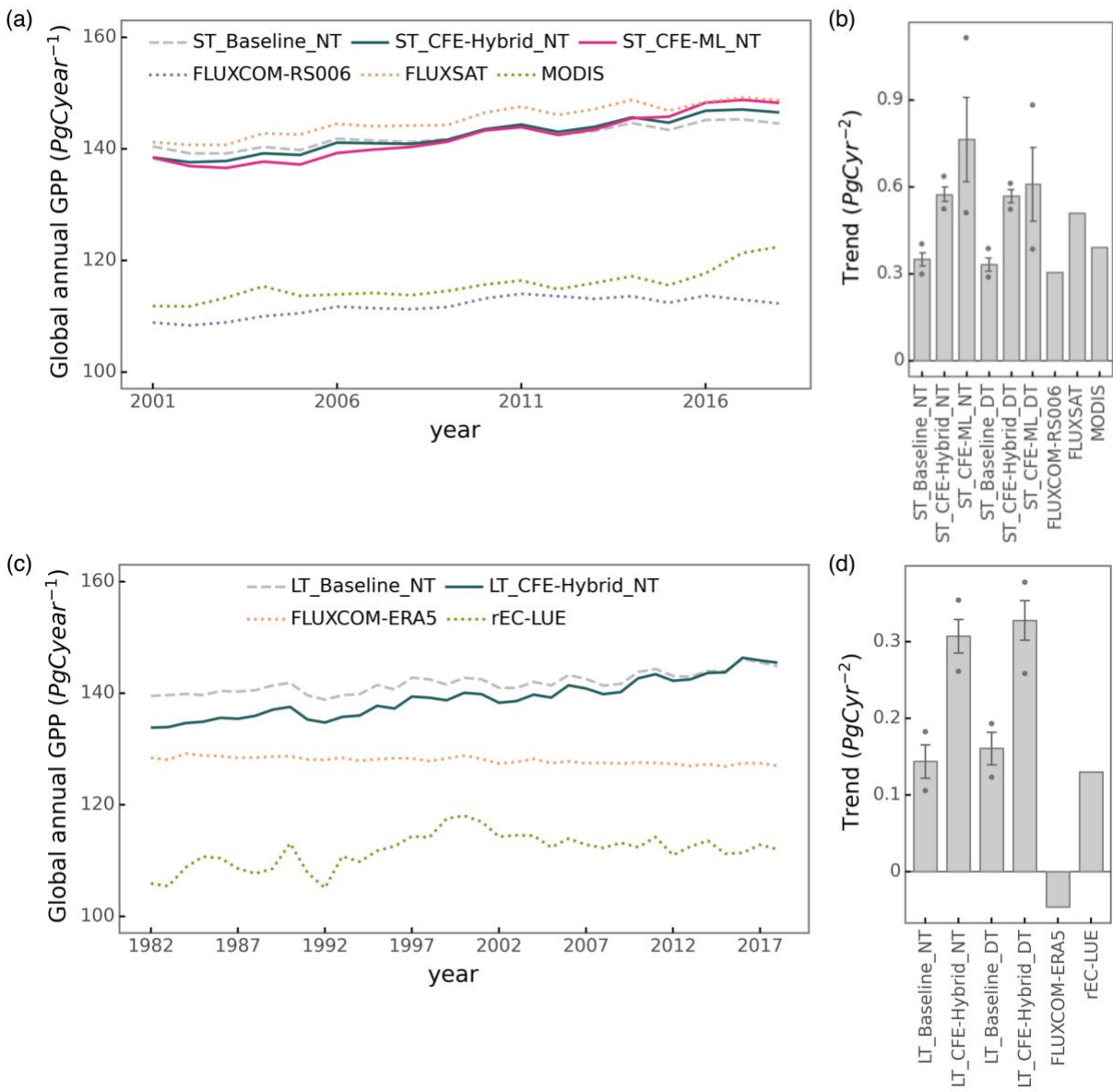
537 From 1982 to 2018, the LT_Baseline models identified increasing GPP trends in large areas of
538 Europe, East and South Asia, as well as Northern Amazon (Figure 10b). The patterns from
539 LT_CFE-Hybrid models aligned closely with the LT_Baseline models but exhibited a stronger
540 positive trend in global tropical areas as well as Eurasian boreal forests. In contrast, FLUXCOM-
541 ERA5 showed overall negative trends in the tropics, despite a small magnitude. Lastly, rEC-LUE
542 agreed with positive GPP trends identified in CEDAR-GPP in the extratropical areas, but predicted
543 a pronounced negative trend in the tropics. At the global scale, all the CEDAR-GPP long-term
544 models predicted an increasing global GPP trend (Figure 11d). The LT_Baseline models showed a
545 trend of 0.13 to 0.15 Pg C year⁻², while the LT_CFE-Hybrid setups doubled that rate. rEC-LUE
546 showed a two-phased pattern with a strong increase in GPP from 1982 to 2000 (0.54 Pg C year⁻²),
547 followed by a decreasing trend after 2001 (-0.20 Pg C year⁻²) (Figure S5). This resulted in an overall
548 positive change at a rate comparable to that of the Baseline model. FLUXCOM-ERA5 exhibited a
549 small negative trend.



550
551
552
553

Figure 10. Annual GPP trend over 2001 – 2018 for short-term CEDAR-GPP, FLUXCOM-RS006, FLUXSAT, and MODIS datasets (a) and over 1982 – 2018 for long-term CEDAR-GPP, FLUXCOM-ERA5 and rEC-LUE datasets (b). Hatched

554 areas indicate the GPP trend that is statistically significant at $p < 0.05$ level under the
 555 Mann-Kendal test.

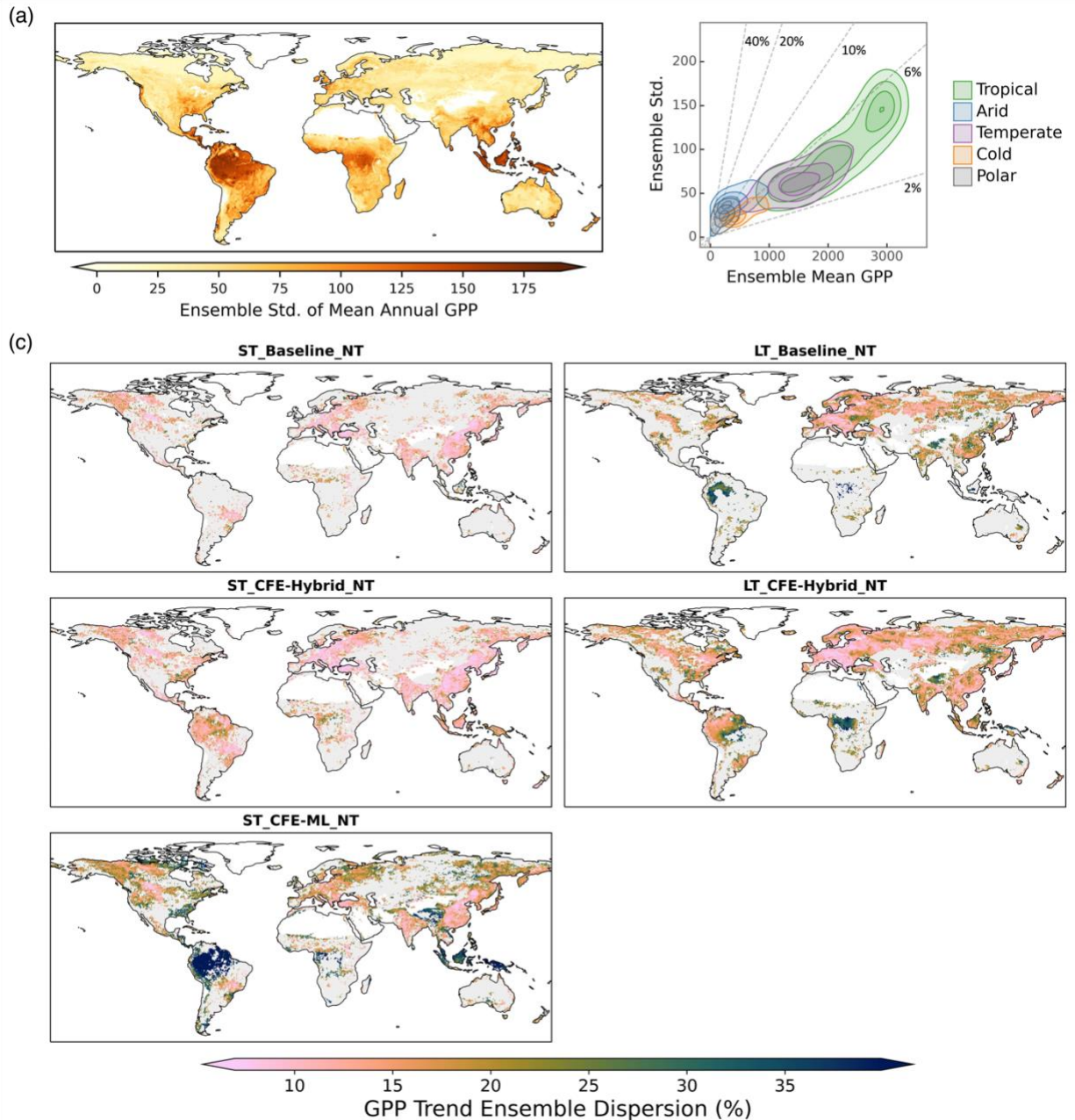


556
 557 Figure 11. Global annual GPP variations (a) and trends (b) over 2001 to 2018 for
 558 short-term CEDAR-GPP, FLUXCOM-RS006, FLUXSAT, and MODIS datasets
 559 Global annual GPP variations (c) and trends (d) over 1982 to 2018 for long-term
 560 long-term CEDAR-GPP, FLUXCOM-ERA5, and rEC-LUE datasets. Error bars in
 561 (b) and (d) represent the 25% to 75% percentile from the model ensembles of
 562 CEDAR-GPP. Dots in (b) and (d) indicate the minimum and maximum from the
 563 model ensembles of CEDAR-GPP.

564

565 3.3 GPP estimation uncertainties

566 We analyzed the spread between the 30 model ensemble members in CEDAR-GPP as an
567 indicator of uncertainties in GPP estimations. The spatial pattern of uncertainty in estimating annual
568 mean GPP largely resembled that of the mean map (Figure 12, Figure 6a). The largest model spread
569 was found in highly productive tropical forests, and this uncertainty decreased in temperate and cold
570 areas (Figure 12a). Tropical ecosystems, with a mean annual GPP between 1000 to 3500 PgCyr⁻¹,
571 only exhibited a 2% and 6% variation within the model ensemble (Figure 12b). Ecosystems in the
572 temperate and cold climates had a smaller annual GPP and proportionally small uncertainties up to
573 6%. However, ecosystems in Arid and Polar climates, despite their similarly low GPP, showed
574 higher model uncertainty, reaching 10% to 40% of the ensemble mean. The estimation uncertainty
575 of GPP trends was generally below 15% to 20% in the CEDAR-GPP datasets under the
576 ST_Baseline and ST_CFE-Hybrid setups (Figure 12c). However, in the ST_CFE-ML setup, the
577 estimation increased substantially, with model spread reaching up to 40% in tropical areas. Notably,
578 the long-term models showed a higher uncertainty compared to the short-term models.



579

580

581

582

583

584

585

586

587

588

589

Figure 12. CEDAR-GPP estimation uncertainty derived from ensemble spread (standard deviation of 30 model predictions). (a) Spatial patterns of the absolute standard deviation from ensemble members in estimating the mean annual GPP from 2001 to 2018, using data from the ST_CFE-Hybrid_NT setup. (b) Relationships between ensemble standard deviation and ensemble mean in mean annual GPP. Colored contours denote clusters of Koppen climate zones. Dashed lines indicate the ratio between the ensemble standard deviation and the ensemble mean with values shown in percentage. (c) Spatial patterns of model uncertainty in GPP long-term trend estimation. Only areas where 90% of the ensemble members showed a statistically significant trend ($p < 0.05$) are shown in the maps. The trend for the short-term datasets

590 (left column) was computed between 2001 to 2018. The trend for the long-term
591 datasets (right column) was computed between 1982 to 2018.

592 **4. Discussion**

593 **4.1 Uncertainties in GPP upscaling**

594 Uncertainties in CEDAR-GPP estimations primarily stem from three components of the
595 machine learning upscaling framework: eddy covariance GPP measurements, input satellite and
596 climate datasets, and the machine learning model itself. Here we examine the three sources of
597 uncertainties, discuss our strategies to reduce their impact, and assess the ability of our model
598 ensemble spread indicator to indicate these uncertainties.

599 **4.1.1 Eddy covariance data**

600 Measurement and modeling errors inherent in eddy covariance GPP can propagate through
601 the upscaling process. While random measurement errors in monthly GPP are typically very small
602 due to error cancellation through temporal aggregation from the half-hourly data, systematic biases
603 may be introduced by the partitioning approach used to derive GPP from the Net Ecosystem
604 Exchange (NEE) measurements, which could be specific to certain environmental conditions
605 (Keenan et al., 2019; Pastorello et al., 2020). This uncertainty was evidenced by a discrepancy
606 between the CEDAR-GPP NT and DT setups. The mean annual GPP from the DT setup was
607 slightly higher than that from the NT setup (Figure 6), and the DT setup also predicted a higher
608 GPP trend in the long-term dataset (Figure 11). Notably, these discrepancies were relatively small
609 compared to the predominant spatiotemporal patterns. Nevertheless, separate DT and NT setups in
610 CEDAR-GPP offered an interesting quantification of the associated uncertainties over space and
611 time, providing insights for future improvements of GPP partitioning approaches.

612 The heavily skewed representativeness of the eddy covariance data remains a key challenge
613 in upscaling, a limitation well-documented in previous studies (Jung et al., 2020). Effective
614 generalization in machine learning models requires a substantial volume of training data that
615 adequately represents and balances unseen conditions. However, the geographical and temporal
616 limitations of eddy covariance data introduce both systematic and random uncertainties in upscaled
617 GPP estimations. Flux sites are primarily clustered in North America and Western Europe, with
618 sparse availability in critical carbon exchange hotspots such as tropical, subtropical, and boreal

619 regions (Figure 1). Consequently, estimated GPP from upscaling was commonly believed to carry
620 high uncertainties in these areas. However, previous studies have found that humid tropical regions
621 were associated with low extrapolation uncertainty despite the limited amount of flux sites (Jung et
622 al., 2020). A pioneering study employing synthetic datasets from a deterministic terrestrial biosphere
623 model assessed the generalizability of the upscaling approach (Jung et al., 2009). Results indicated
624 that a machine learning model trained with data from the FLUXNET sites' location and times could
625 account for 92% variation of GPP globally. Therefore, to fully understand the upscaling uncertainty,
626 it is essential to carefully evaluate the generalization or extrapolation errors within the predictor
627 space, which indicates the actual environmental controls and mechanisms of the ecosystem carbon
628 fluxes (van der Horst et al., 2019; Villarreal and Vargas, 2021).

629 We compiled a large (~18000 site-months) set of high-quality eddy covariance data in
630 generating CEDAR-GPP to potentially mitigate uncertainties associated with eddy covariance
631 measurements. The FLUXNET2015 dataset was complemented by two regional networks
632 processed under the same ONEFLUX protocol with additional sites and longer records. Data were
633 screened with stringent quality control criteria to reduce measurement uncertainties. Nevertheless,
634 uncertainties in GPP upscaling due to limited eddy covariance data representation remain
635 substantial. Additionally, our analysis suggested that the estimated global GPP magnitudes were
636 related to the eddy covariance GPP data used in the model. Global GPP magnitudes derived from
637 CEDAR-GPP and FLUXSAT aligned closely with that from Terrestrial Biosphere Models (Anav et
638 al., 2015), while FLUXCOM's estimation was much lower (Figure 6, Figure 11). FLUXSAT used
639 FLUXNET2015 as the training set, which largely overlapped with that included in CEDAR-GPP
640 (Joiner and Yoshida, 2020). FLUXCOM utilized the FLUXNET La Thuile set combined with
641 CarboAfrica network, which consisted of a very different set of sites, with flux data processed with a
642 different pipeline (Tramontana et al., 2016). The influence from the predictor datasets was minimal
643 since all three datasets relied on MODIS-derived products. For a comprehensive evaluation of the
644 impacts of flux site representativeness on upscaling, future research directions could include
645 conducting synthetic experiments with simulations of ensembles of terrestrial biosphere models.

646 4.1.2 Input predictors and controlling factors

647 Input predictors, including satellite and climate datasets, bring additional uncertainties to
648 upscaled GPP. First, satellite remote sensing data contains noises resulting from atmospheric
649 conditions, sun-earth geometry, soil background, and geolocation inaccuracies. The models or

650 algorithms used for variable estimation, such as those for retrieving LAI, fAPAR, LST, and soil
651 moisture, also contain random errors and systematic biases specific to certain regions, biome types,
652 or climatic conditions (Yan et al., 2016b; Fang et al., 2019; Ma et al., 2019). Additionally, satellite
653 observations frequently contain missing values due to clouds, aerosols, snow, and satellite revisit
654 cycles, leading to both systematic and random uncertainties. In producing CEDAR-GPP, we aimed
655 to mitigate these uncertainties through comprehensive preprocessing procedures. Our temporal gap-
656 filling strategy improved upon previous upscaling efforts by exploiting both the temporal
657 dependency of vegetation status and long-term climatology, to reduce biases from missing values.
658 Temporal and spatial aggregation further diminished the remaining data gaps and random noises.
659 Nevertheless, it was still likely that considerable uncertainties remained in satellite datasets and
660 propagated into the upscaled estimations.

661 The mismatch between the footprint of the eddy covariance measurements and the coarse
662 resolution of satellite observations presents another significant, and potentially more impactful,
663 source of uncertainties. Flux towers typically have a footprint of around $\sim 1 \text{ km}^2$ (Chu et al., 2021),
664 whereas satellite observations employed in CEDAR-GPP and most other upscaled datasets were at
665 5 km or lower resolution. Systematic and random errors could be introduced due to this mismatch,
666 particularly in heterogenous biomes and areas with a mixture of vegetation and non-vegetated land
667 covers. Gaber et al. (in prep) discovered that machine learning models trained with MODIS datasets
668 at their native 500m resolution substantially outperformed those trained with aggregated datasets at
669 0.05-degree resolution, especially in capturing interannual variabilities. One approach to mitigate this
670 issue is by generating upscaled datasets at a higher spatial resolution (e.g. 500m). Alternatively,
671 models could be trained at a high resolution and applied to the coarse resolution to reduce
672 computation and storage requirements (Dannenberg et al., 2023). However, this approach does not
673 address inherent scaling errors in coarse-resolution satellite images (Yan et al., 2016a; Dong et al.,
674 2023).

675 Besides the quality of predictors, successful machine learning upscaling also depends on a
676 comprehensive set of features representing all controlling factors. Satellite-derived vegetation
677 structural indicators have been found to be the most influential features in GPP upscaling
678 (Tramontana et al., 2015); consequently, trends and interannual variabilities in upscaled GPP
679 reflected that from the underlying vegetation structural datasets. For example, the FLUXCOM-
680 RSv006 dataset was generated based on the MODIS collection 6 dataset, showing similar temporal
681 dynamics globally as CEDAR-GPP and FLUXSAT. However, the previous version that utilized

682 MODIS collection 5 products exhibited minimal interannual variability and insignificant trend (Ryu
683 et al., 2019).

684 Furthermore, information about resource limitations and stress factors can be crucial for
685 certain biomes and/or under specific conditions (Stocker et al., 2018). For example, Dannenberg et
686 al. (2023) found that incorporating LST from MODIS and soil moisture from the SMAP satellite
687 datasets improved the machine learning estimation accuracy of GPP in drylands from $R^2 \sim 0.4$ to
688 0.7 for dryland sites in North America. CEDAR-GPP integrated multi-source satellite observations
689 (optical, thermal, microwave) as well as climate variables to obtain comprehensive information about
690 GPP dynamics. Nevertheless, as was commonly observed in machine learning upscaling, the models
691 failed to capture interannual anomalies. This could imply that vital information for GPP estimation
692 remained missing or inadequately represented in existing datasets. Examples include factors related
693 to agricultural management practices (crop type, cultivar, irrigation, fertilization), plant hydraulic and
694 physiological properties, as well as root and soil characteristics.

695 4.1.3 Machine learning models and uncertainty quantification

696 The choice of machine learning models and their parameterization has been found to have a
697 relatively minor impact on the upscaling of GPP (Tramontana et al., 2015). CEDAR used the state-
698 of-the-art boosting algorithm, XGBoost, which provided near-optimal performance considering the
699 current data availability. Further reduction of model uncertainty will likely rely on additional
700 information, such as increasing the number of training samples or incorporating more high-quality
701 predictors. Additionally, temporal dependency of carbon fluxes responses to atmospheric controls
702 may also be exploited with specialized deep neural networks such as recurrent neural networks or
703 transformers (Besnard et al., 2019; Ma and Liang, 2022).

704 A key challenge, however, is the quantification of uncertainties in machine learning upscaling
705 (Reichstein et al., 2019). The limited availability of eddy covariance data hinders a comprehensive
706 assessment of the extrapolation errors; consequently, metrics of predictive performance from cross-
707 validation are inherently biased. CEDAR derived estimation uncertainty for each GPP prediction
708 using bootstrapping model ensemble, which naturally mimics the biased sampling of flux tower
709 locations. Notably, the choice of input climate reanalysis datasets could also induce systematic
710 differences in GPP spatial and temporal patterns (Tramontana et al., 2015). The FLUXCOM
711 product used model ensembles based on different reanalysis datasets to capture these uncertainties.
712 Additionally, different satellite datasets of vegetation structural proxies, such as LAI, also exhibited

713 significant discrepancies (Jiang et al., 2017). Thus, an ensemble approach combining site-level
714 bootstrapping with multiple sources of input predictors could potentially provide a more
715 comprehensive quantification of uncertainties. Future work may also explore Bayesian neural
716 networks, which provide uncertainty along with predictions and, at the same time, present high
717 predictive power comparable to ensemble tree-based algorithms (Ma et al., 2021).

718 4.2 Long-term GPP changes and CO₂ fertilization effect

719 CEDAR-GPP was constructed using a comprehensive set of climate variables and multi-
720 source satellite observations, thus, encapsulating long-term GPP dynamics from both direct and
721 indirect effects of climate controls. Particularly, CEDAR-GPP included the direct CO₂ fertilization
722 effect, which has been shown to dominate the increasing trend of global GPP (Chen et al., 2022).
723 Incorporating these effects substantially improved long-term trends of GPP from site to global
724 scales (Figure 5, 10, 11). CEDAR's CFE-Hybrid setup offered a conservative estimation of the
725 direct CO₂ effects by simulating the light-limited sensitivity on LUE for C3 plants (Walker et al.,
726 2021). Nevertheless, the model did not account for the impacts of nutrient availability, which could
727 potentially constrain CO₂ fertilization (Reich et al., 2014; Peñuelas et al., 2017; Terrer et al., 2019).
728 Furthermore, the sensitivity of light-limited photosynthesis is a function of temperature, resulting in
729 the most pronounced increasing trend in the tropics (Figure 10). Yet the model assumed a fixed
730 ratio of leaf-internal to ambient CO₂, and thus did not include any responses to vapor pressure
731 deficit.

732 The CFE-ML model adopted a data-driven approach to infer CO₂ effects directly from eddy
733 covariance data. This strategy allowed the model to capture any physiological pathways of the CO₂
734 impact evidenced in the eddy covariance measurements, including the increases of the biochemical
735 rates as well as enhancements in water use efficiency (Keenan et al., 2013). The model successfully
736 detected a strong positive effect of CO₂ on eddy covariance measured GPP, consistent with
737 previous studies based on process-based and statistical models (Fernández-Martínez et al., 2017;
738 Ueyama et al., 2020; Chen et al., 2022). Notably, the CFE-ML model could have included the
739 impacts of other factors that exhibit a strong temporal correlation with CO₂. For example,
740 industrialization-induced increases in nitrogen deposition could synergistically boost GPP alongside
741 CO₂ (O'Sullivan et al., 2019). Technological and management improvements in agriculture that
742 contribute to a global boost of crop photosynthesis (Zeng et al., 2014), might also be indirectly
743 reflected in the model estimates. As a result, the CFE-ML predicted a GPP trend that more closely

744 aligned with eddy covariance observations, and the upscaled dataset also showed a globally higher
745 trend than CFE-Hybrid (Figure 5; Figure 10). Nevertheless, spatial patterns of GPP trends from the
746 CFE-ML aligned with that from CFE-Hybrid, reflecting a strong temperature dependency, implying
747 that the effects of CO₂ likely remained the most significant factor. Additionally, the considerable
748 ensemble spread in the CO₂ trends from the CFE-ML model (Figure 11, Figure 13) underscored a
749 high level of uncertainty in the machine learning quantified CO₂ effects. Future work may exploit
750 explainable machine learning and causal inference to investigate the underlying mechanisms of CO₂
751 effects.

752 **5. Data availability**

753 The CEDAR-GPP product, comprising ten GPP datasets, is available at zenodo
754 (<https://doi.org/10.5281/zenodo.8212707>). These datasets were generated at a spatial resolution of
755 0.05° and monthly time steps. Each dataset includes an ensemble mean GPP and an ensemble
756 standard deviation.

757 **6. Summary and conclusions**

758 We present the CEDAR-GPP product generated by upscaling global eddy covariance
759 measurements with machine learning and a broad range of satellite and climate variables. CEDAR-
760 GPP comprises four long-term datasets from 1982 to 2020 and six short-term datasets from 2001 to
761 2020. These datasets encompass three different configurations of the incorporation of direct CO₂
762 fertilization effects and two partitioning approaches to derive GPP from eddy covariance data. The
763 machine learning models of CEDAR-GPP achieved high capability in predicting monthly GPP, its
764 seasonal cycles, and spatial variability within the globally distributed eddy covariance sites, with
765 cross-validated R² between 0.56 to 0.79. Short-term model setups consistently outperformed long-
766 term models due to considerably more and higher-quality information from multi-source satellite
767 observations.

768 CEDAR-GPP advances satellite-based GPP estimations, as the first upscaled dataset that
769 considered the direct biochemical effects of elevated atmospheric CO₂ on photosynthesis, which is
770 responsible for an increasing land carbon sink over the past decades. We showed that incorporating
771 this effect in our CFE-ML and CFE-Hybrid models substantially improved the estimation of GPP

772 trends at eddy covariance sites. Global patterns of long-term GPP trends in the CFE-ML setups
773 showed a strong temperature dependency consistent with biophysical theories. Aside from the trend,
774 global spatial and temporal GPP patterns from CEDAR generally aligned with other satellite-based
775 GPP datasets.

776 In conclusion, CEDAR-GPP, informed by global eddy covariance measurements (~18000
777 site-months) and a broad range of multi-source remote sensing observations and climatic variables,
778 offered a comprehensive representation of global GPP spatial and temporal dynamics over the past
779 four decades. The different CO₂ fertilization configurations integrated in CEDAR-GPP offer new
780 opportunities for understanding global ecosystem photosynthesis's response to increases in
781 atmospheric CO₂ along different pathways over space and time. CEDAR-GPP is expected to serve
782 as a valuable tool for benchmarking process-based modeling and constraining the global carbon
783 cycle.

784 **Supplement**

785 The supplement related to this article is available online.

786 **Author contributions**

787 T. K. and Y. K. conceptualized the study. Y. K. performed the formal analysis and generated
788 the final product. Y. K., T. K., M. B., and M. G. contributed to the development and investigation
789 of the research. Y. K., M. G., and X. L. contributed to data curation and processing. Y. K. prepared
790 the manuscript with contributions from all co-authors. T. K. supervised the project.

791 **Acknowledgments**

792 We are grateful to Dr. Youngryel Ryu for providing the BESS_Rad dataset and Dr. Martin
793 Jung for sharing the FLUXCOM-RS006 dataset. We also thank Dr. Muye Li for sharing early
794 versions of the PKU GIMMS NDVI4g and LAI4g datasets with us.

795 **Financial support**

796 This research was supported by the U.S. Department of Energy Office of Science Early
797 Career Research Program award #DE-SC0021023 and a NASA Award 80NSSC21K1705. TFK
798 acknowledges additional support from the LEMONTREE (Land Ecosystem Models based On New
799 Theory, observations and Experiments) project, funded through the generosity of Eric and Wendy

800 Schmidt by recommendation of the Schmidt Futures programme and NASA award

801 #80NSSC20K1801.

802 **Reference**

- 803 Anav, A., Friedlingstein, P., Beer, C., Ciais, P., Harper, A., Jones, C., Murray-Tortarolo, G., Papale,
804 D., Parazoo, N. C., Peylin, P., Piao, S., Sitch, S., Viovy, N., Wiltshire, A., and Zhao, M.:
805 Spatiotemporal patterns of terrestrial gross primary production: A review, *Reviews of*
806 *Geophysics*, 1–34, <https://doi.org/10.1002/2015RG000483>. Received, 2015.
- 807 Badgley, G., Anderegg, L. D. L., Berry, J. A., and Field, C. B.: Terrestrial gross primary production:
808 Using NIRV to scale from site to globe, *Global Change Biology*, 25, 3731–3740,
809 <https://doi.org/10.1111/gcb.14729>, 2019.
- 810 Baldocchi, D. D.: How eddy covariance flux measurements have contributed to our understanding
811 of Global Change Biology, *Global Change Biology*, 26, 242–260,
812 <https://doi.org/10.1111/gcb.14807>, 2020.
- 813 Beck, H. E., Zimmermann, N. E., McVicar, T. R., Vergopolan, N., Berg, A., and Wood, E. F.:
814 Present and future köppen-geiger climate classification maps at 1-km resolution, *Scientific Data*,
815 5, 1–12, <https://doi.org/10.1038/sdata.2018.214>, 2018.
- 816 Beer, C., Reichstein, M., Tomelleri, E., Ciais, P., Jung, M., Carvalhais, N., Rödenbeck, C., Arain, M.
817 A., Baldocchi, D., Bonan, G. B., Bondeau, A., Cescatti, A., Lasslop, G., Lindroth, A., Lomas, M.,
818 Luysaert, S., Margolis, H., Oleson, K. W., Rouspard, O., Veenendaal, E., Viovy, N., Williams, C.,
819 Woodward, F. I., and Papale, D.: Terrestrial gross carbon dioxide uptake: Global distribution and
820 covariation with climate, *Science*, 329, 834–838, <https://doi.org/10.1126/science.1184984>, 2010.
- 821 Berdugo, M., Gaitán, J. J., Delgado-Baquerizo, M., Crowther, T. W., and Dakos, V.: Prevalence and
822 drivers of abrupt vegetation shifts in global drylands, *Proceedings of the National Academy of*
823 *Sciences*, 119, e2123393119, <https://doi.org/10.1073/pnas.2123393119>, 2022.
- 824 Besnard, S., Carvalhais, N., Altaf Arain, M., Black, A., Brede, B., Buchmann, N., Chen, J., Clevers, J.
825 G. P. W., Dutrieux, L. P., Gans, F., Herold, M., Jung, M., Kosugi, Y., Knohl, A., Law, B. E.,
826 Paul-Limoges, E., Lohila, A., Merbold, L., Rouspard, O., Valentini, R., Wolf, S., Zhang, X., and
827 Reichstein, M.: Memory effects of climate and vegetation affecting net ecosystem CO₂ fluxes in
828 global forests, *PLoS ONE*, 14, 1–22, <https://doi.org/10.1371/journal.pone.0211510>, 2019.
- 829 Bloomfield, K. J., Stocker, B. D., Keenan, T. F., and Prentice, I. C.: Environmental controls on the
830 light use efficiency of terrestrial gross primary production, *Global Change Biology*, 29, 1037–
831 1053, <https://doi.org/10.1111/gcb.16511>, 2023.
- 832 Campbell, J. E., Berry, J. A., Seibt, U., Smith, S. J., Montzka, S. A., Launois, T., Belviso, S., Bopp, L.,
833 and Laine, M.: Large historical growth in global terrestrial gross primary production, *Nature*, 544,
834 84–87, <https://doi.org/10.1038/nature22030>, 2017.
- 835 Camps-Valls, G., Campos-Taberner, M., Moreno-Martínez, Á., Walther, S., Duveiller, G., Cescatti,
836 A., Mahecha, M. D., Muñoz-Marí, J., García-Haro, F. J., Guanter, L., Jung, M., Gamon, J. A.,
837 Reichstein, M., and Running, S. W.: A unified vegetation index for quantifying the terrestrial
838 biosphere, *Science Advances*, 7, eabc7447, <https://doi.org/10.1126/sciadv.abc7447>, 2021.
- 839 Cao, S., Li, M., Zhu, Z., Zha, J., Zhao, W., Duanmu, Z., Chen, J., Zheng, Y., and Chen, Y.:
840 Spatiotemporally consistent global dataset of the GIMMS Leaf Area Index (GIMMS LAI4g)
841 from 1982 to 2020, *Earth System Science Data Discussions*, 1–31, [https://doi.org/10.5194/essd-](https://doi.org/10.5194/essd-2023-68)
842 [2023-68](https://doi.org/10.5194/essd-2023-68), 2023.
- 843 Chen, C., Park, T., Wang, X., Piao, S., Xu, B., Chaturvedi, R. K., Fuchs, R., Brovkin, V., Ciais, P.,
844 Fensholt, R., Tømmervik, H., Bala, G., Zhu, Z., Nemani, R. R., and Myneni, R. B.: China and

845 India lead in greening of the world through land-use management, *Nature Sustainability*, 2, 122–
846 129, <https://doi.org/10.1038/s41893-019-0220-7>, 2019.

847 Chen, C., Riley, W. J., Prentice, I. C., and Keenan, T. F.: CO₂ fertilization of terrestrial
848 photosynthesis inferred from site to global scales, *Proceedings of the National Academy of*
849 *Sciences*, 119, 1–8, <https://doi.org/10.1073/pnas.2115627119/-/DCSupplemental>. Published,
850 2022.

851 Chen, T. and Guestrin, C.: XGBoost: a scalable tree boosting system, in: *Proceedings of the 22nd*
852 *ACM SIGKDD International Conference on Knowledge Discovery and Data Mining -*
853 *KDD '16*, 785–794, <https://doi.org/10.1145/2939672.2939785>, 2016.

854 Chu, H., Luo, X., Ouyang, Z., Chan, W. S., Dengel, S., Biraud, S. C., Torn, M. S., Metzger, S.,
855 Kumar, J., Arain, M. A., Arkebauer, T. J., Baldocchi, D., Bernacchi, C., Billesbach, D., Black, T.
856 A., Blanken, P. D., Bohrer, G., Bracho, R., Brown, S., Brunsell, N. A., Chen, J., Chen, X., Clark,
857 K., Desai, A. R., Duman, T., Durden, D., Fares, S., Forbrich, I., Gamon, J. A., Gough, C. M.,
858 Griffis, T., Helbig, M., Hollinger, D., Humphreys, E., Ikawa, H., Iwata, H., Ju, Y., Knowles, J. F.,
859 Knox, S. H., Kobayashi, H., Kolb, T., Law, B., Lee, X., Litvak, M., Liu, H., Munger, J. W.,
860 Noormets, A., Novick, K., Oberbauer, S. F., Oechel, W., Oikawa, P., Papuga, S. A., Pendall, E.,
861 Prajapati, P., Prueger, J., Quinton, W. L., Richardson, A. D., Russell, E. S., Scott, R. L., Starr, G.,
862 Staebler, R., Stoy, P. C., Stuart-Haëntjens, E., Sonnentag, O., Sullivan, R. C., Suyker, A., Ueyama,
863 M., Vargas, R., Wood, J. D., and Zona, D.: Representativeness of Eddy-Covariance flux
864 footprints for areas surrounding AmeriFlux sites, *Agricultural and Forest Meteorology*, 301–302,
865 <https://doi.org/10.1016/j.agrformet.2021.108350>, 2021.

866 Dannenberg, M. P., Barnes, M. L., Smith, W. K., Johnston, M. R., Meerdink, S. K., Wang, X., Scott,
867 R. L., and Biederman, J. A.: Upscaling dryland carbon and water fluxes with artificial neural
868 networks of optical, thermal, and microwave satellite remote sensing, *Biogeosciences*, 20, 383–
869 404, <https://doi.org/10.5194/bg-20-383-2023>, 2023.

870 De Kauwe, M. G., Medlyn, B. E., Zaehle, S., Walker, A. P., Dietze, M. C., Hickler, T., Jain, A. K.,
871 Luo, Y., Parton, W. J., Prentice, I. C., Smith, B., Thornton, P. E., Wang, S., Wang, Y.-P., Wårlind,
872 D., Weng, E., Crous, K. Y., Ellsworth, D. S., Hanson, P. J., Seok Kim, H.-, Warren, J. M., Oren,
873 R., and Norby, R. J.: Forest water use and water use efficiency at elevated CO₂: a model-data
874 intercomparison at two contrasting temperate forest FACE sites, *Global Change Biology*, 19,
875 1759–1779, <https://doi.org/10.1111/gcb.12164>, 2013.

876 De Kauwe, M. G., Keenan, T. F., Medlyn, B. E., Prentice, I. C., and Terrer, C.: Satellite based
877 estimates underestimate the effect of CO₂ fertilization on net primary productivity, *Nature*
878 *Climate Change*, 6, 892–893, <https://doi.org/10.1038/nclimate3105>, 2016.

879 Dong, Y., Li, J., Jiao, Z., Liu, Q., Zhao, J., Xu, B., Zhang, H., Zhang, Z., Liu, C., Knyazikhin, Y., and
880 Myneni, R. B.: A Method for Retrieving Coarse-Resolution Leaf Area Index for Mixed Biomes
881 Using a Mixed-Pixel Correction Factor, *IEEE Transactions on Geoscience and Remote Sensing*,
882 61, 1–17, <https://doi.org/10.1109/TGRS.2023.3235949>, 2023.

883 Dorigo, W., Wagner, W., Albergel, C., Albrecht, F., Balsamo, G., Brocca, L., Chung, D., Ertl, M.,
884 Forkel, M., Gruber, A., Haas, E., Hamer, P. D., Hirschi, M., Ikonen, J., de Jeu, R., Kidd, R.,
885 Lahoz, W., Liu, Y. Y., Miralles, D., Mistelbauer, T., Nicolai-Shaw, N., Parinussa, R., Pratola, C.,
886 Reimer, C., van der Schalie, R., Seneviratne, S. I., Smolander, T., and Lecomte, P.: ESA CCI Soil
887 Moisture for improved Earth system understanding: State-of-the art and future directions,
888 *Remote Sensing of Environment*, 203, 185–215, <https://doi.org/10.1016/j.rse.2017.07.001>,
889 2017.

890 Dorigo, W. A., Gruber, A., De Jeu, R. A. M., Wagner, W., Stacke, T., Loew, A., Albergel, C., Brocca,
891 L., Chung, D., Parinussa, R. M., and Kidd, R.: Evaluation of the ESA CCI soil moisture product

892 using ground-based observations, *Remote Sensing of Environment*, 162,
893 <https://doi.org/10.1016/j.rse.2014.07.023>, 2015.

894 Ehlers, I., Augusti, A., Betson, T. R., Nilsson, M. B., Marshall, J. D., and Schleucher, J.: Detecting
895 long-term metabolic shifts using isotopomers: CO₂-driven suppression of photorespiration in C₃
896 plants over the 20th century, *Proceedings of the National Academy of Sciences*, 112, 15585–
897 15590, <https://doi.org/10.1073/pnas.1504493112>, 2015.

898 Fang, H., Baret, F., Plummer, S., and Schaepman-Strub, G.: An overview of global leaf area index
899 (LAI): Methods, products, validation, and applications, *Reviews of Geophysics*, 2018RG000608,
900 <https://doi.org/10.1029/2018RG000608>, 2019.

901 Fernández-Martínez, M., Vicca, S., Janssens, I. A., Ciais, P., Obersteiner, M., Bartrons, M., Sardans,
902 J., Verger, A., Canadell, J. G., Chevallier, F., Wang, X., Bernhofer, C., Curtis, P. S., Gianelle, D.,
903 Grünwald, T., Heinesch, B., Ibrom, A., Knohl, A., Laurila, T., Law, B. E., Limousin, J. M.,
904 Longdoz, B., Loustau, D., Mammarella, I., Matteucci, G., Monson, R. K., Montagnani, L., Moors,
905 E. J., Munger, J. W., Papale, D., Piao, S. L., and Peñuelas, J.: Atmospheric deposition, CO₂, and
906 change in the land carbon sink, *Sci Rep*, 7, 9632, <https://doi.org/10.1038/s41598-017-08755-8>,
907 2017.

908 Friedl, M. and Sulla-Menashe, D.: MCD12Q1 MODIS/Terra+Aqua Land Cover Type Yearly L3
909 Global 500m SIN Grid V006 [Data set], NASA EOSDIS Land Processes DAAC,
910 <https://doi.org/10.5067/MODIS/MCD12Q1.006>, 2019.

911 Friedlingstein, P., Meinshausen, M., Arora, V. K., Jones, C. D., Anav, A., Liddicoat, S. K., and
912 Knutti, R.: Uncertainties in CMIP5 climate projections due to carbon cycle feedbacks, *Journal of*
913 *Climate*, 27, 511–526, <https://doi.org/10.1175/JCLI-D-12-00579.1>, 2014.

914 Friedlingstein, P., O’Sullivan, M., Jones, M. W., Andrew, R. M., Gregor, L., Hauck, J., Le Quéré, C.,
915 Luijkx, I. T., Olsen, A., Peters, G. P., Peters, W., Pongratz, J., Schwingshackl, C., Sitch, S.,
916 Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., Alkama, R., Arneeth, A., Arora, V. K., Bates,
917 N. R., Becker, M., Bellouin, N., Bittig, H. C., Bopp, L., Chevallier, F., Chini, L. P., Cronin, M.,
918 Evans, W., Falk, S., Feely, R. A., Gasser, T., Gehlen, M., Gkritzalis, T., Gloege, L., Grassi, G.,
919 Gruber, N., Gürses, Ö., Harris, I., Hefner, M., Houghton, R. A., Hurtt, G. C., Iida, Y., Ilyina, T.,
920 Jain, A. K., Jersild, A., Kadono, K., Kato, E., Kennedy, D., Klein Goldewijk, K., Knauer, J.,
921 Korsbakken, J. I., Landschützer, P., Lefèvre, N., Lindsay, K., Liu, J., Liu, Z., Marland, G., Mayot,
922 N., McGrath, M. J., Metzl, N., Monacci, N. M., Munro, D. R., Nakaoka, S.-I., Niwa, Y., O’Brien,
923 K., Ono, T., Palmer, P. I., Pan, N., Pierrot, D., Pockock, K., Poulter, B., Resplandy, L., Robertson,
924 E., Rödenbeck, C., Rodriguez, C., Rosan, T. M., Schwinger, J., Séférian, R., Shutler, J. D.,
925 Skjelvan, I., Steinhoff, T., Sun, Q., Sutton, A. J., Sweeney, C., Takao, S., Tanhua, T., Tans, P. P.,
926 Tian, X., Tian, H., Tilbrook, B., Tsujino, H., Tubiello, F., van der Werf, G. R., Walker, A. P.,
927 Wanninkhof, R., Whitehead, C., Willstrand Wranne, A., et al.: Global Carbon Budget 2022, *Earth*
928 *System Science Data*, 14, 4811–4900, 2023.

929 Gampe, D., Zscheischler, J., Reichstein, M., Sullivan, M. O., Smith, W. K., Sitch, S., and Buermann,
930 W.: Increasing impact of warm droughts on northern ecosystem productivity over recent
931 decades, *Nature Climate Change*, <https://doi.org/10.1038/s41558-021-01112-8>, 2021.

932 Gao, B. C.: NDWI - A normalized difference water index for remote sensing of vegetation liquid
933 water from space, *Remote Sensing of Environment*, 58, 257–266,
934 [https://doi.org/10.1016/S0034-4257\(96\)00067-3](https://doi.org/10.1016/S0034-4257(96)00067-3), 1996.

935 Gitelson, A. A.: Remote estimation of leaf area index and green leaf biomass in maize canopies,
936 *Geophysical Research Letters*, 30, 1248, <https://doi.org/10.1029/2002GL016450>, 2003.

937 Gruber, A., Scanlon, T., Van Der Schalie, R., Wagner, W., and Dorigo, W.: Evolution of the ESA
938 CCI Soil Moisture climate data records and their underlying merging methodology, *Earth System*
939 *Science Data*, 11, 717–739, <https://doi.org/10.5194/essd-11-717-2019>, 2019.

940 Haverd, V., Smith, B., Canadell, J. G., Cuntz, M., Mikaloff-Fletcher, S., Farquhar, G., Woodgate, W.,
941 Briggs, P. R., and Trudinger, C. M.: Higher than expected CO₂ fertilization inferred from leaf to
942 global observations, *Global Change Biology*, 26, 2390–2402, <https://doi.org/10.1111/gcb.14950>,
943 2020.

944 van der Horst, S. V. J., Pitman, A. J., De Kauwe, M. G., Ukkola, A., Abramowitz, G., and Isaac, P.:
945 How representative are FLUXNET measurements of surface fluxes during temperature
946 extremes?, *Biogeosciences*, 16, 1829–1844, <https://doi.org/10.5194/bg-16-1829-2019>, 2019.

947 Jiang, C., Ryu, Y., Fang, H., Myneni, R., Claverie, M., and Zhu, Z.: Inconsistencies of interannual
948 variability and trends in long-term satellite leaf area index products, *Global Change Biology*, 23,
949 4133–4146, <https://doi.org/10.1111/gcb.13787>, 2017.

950 Joiner, J. and Yoshida, Y.: Satellite-based reflectances capture large fraction of variability in global
951 gross primary production (GPP) at weekly time scales, *Agricultural and Forest Meteorology*, 291,
952 108092, <https://doi.org/10.1016/j.agrformet.2020.108092>, 2020.

953 Jung, M., Reichstein, M., and Bondeau, A.: Towards global empirical upscaling of FLUXNET eddy
954 covariance observations: Validation of a model tree ensemble approach using a biosphere model,
955 *Biogeosciences*, 6, 2001–2013, <https://doi.org/10.5194/bg-6-2001-2009>, 2009.

956 Jung, M., Reichstein, M., Margolis, H. A., Cescatti, A., Richardson, A. D., Arain, M. A., Arneth, A.,
957 Bernhofer, C., Bonal, D., Chen, J., Gianelle, D., Gobron, N., Kiely, G., Kutsch, W., Lasslop, G.,
958 Law, B. E., Lindroth, A., Merbold, L., Montagnani, L., Moors, E. J., Papale, D., Sottocornola, M.,
959 Vaccari, F., and Williams, C.: Global patterns of land-atmosphere fluxes of carbon dioxide , latent
960 heat , and sensible heat derived from eddy covariance , satellite , and meteorological observations,
961 *Journal of Geophysical Research: Biogeosciences*, 116, 1–16,
962 <https://doi.org/10.1029/2010JG001566>, 2011.

963 Jung, M., Reichstein, M., Schwalm, C. R., Huntingford, C., Sitch, S., Ahlström, A., Arneth, A.,
964 Camps-Valls, G., Ciais, P., Friedlingstein, P., Gans, F., Ichii, K., Jain, A. K., Kato, E., Papale, D.,
965 Poulter, B., Raduly, B., Rödenbeck, C., Tramontana, G., Viovy, N., Wang, Y. P., Weber, U.,
966 Zaehle, S., and Zeng, N.: Compensatory water effects link yearly global land CO₂ sink changes
967 to temperature, *Nature*, 541, 516–520, <https://doi.org/10.1038/nature20780>, 2017.

968 Jung, M., Schwalm, C., Migliavacca, M., Walther, S., Camps-Valls, G., Koirala, S., Anthoni, P.,
969 Besnard, S., Bodesheim, P., Carvalhais, N., Chevallier, F., Gans, F., S Goll, D., Haverd, V.,
970 Köhler, P., Ichii, K., K Jain, A., Liu, J., Lombardozzi, D., E M S Nabel, J., A Nelson, J.,
971 O’Sullivan, M., Pallandt, M., Papale, D., Peters, W., Pongratz, J., Rödenbeck, C., Sitch, S.,
972 Tramontana, G., Walker, A., Weber, U., and Reichstein, M.: Scaling carbon fluxes from eddy
973 covariance sites to globe: Synthesis and evaluation of the FLUXCOM approach, *Biogeosciences*,
974 17, 1343–1365, <https://doi.org/10.5194/bg-17-1343-2020>, 2020.

975 Kang, Y., Ozdogan, M., Zhu, X., Ye, Z., Hain, C., and Anderson, M.: Comparative assessment of
976 environmental variables and machine learning algorithms for maize yield prediction in the US
977 Midwest, *Environmental Research Letters*, 15, <https://doi.org/10.1088/1748-9326/ab7df9>,
978 2020.

979 Keeling, R. F., Graven, H. D., Welp, L. R., Resplandy, L., Bi, J., Piper, S. C., Sun, Y., Bollenbacher,
980 A., and Meijer, H. A. J.: Atmospheric evidence for a global secular increase in carbon isotopic
981 discrimination of land photosynthesis, *Proceedings of the National Academy of Sciences of the*
982 *United States of America*, 114, 10361–10366, <https://doi.org/10.1073/pnas.1619240114>, 2017.

983 Keenan, T. F. and Williams, C. A.: The terrestrial carbon sink, *Annual Review of Environment and*
984 *Resources*, 43, 219–243, <https://doi.org/10.1146/annurev-environ-102017-030204>, 2018.

985 Keenan, T. F., Hollinger, D. Y., Bohrer, G., Dragoni, D., Munger, J. W., Schmid, H. P., and
986 Richardson, A. D.: Increase in forest water-use efficiency as atmospheric carbon dioxide
987 concentrations rise, *Nature*, 499, 324–327, <https://doi.org/10.1038/nature12291>, 2013.

988 Keenan, T. F., Prentice, I. C., Canadell, J. G., Williams, C. A., Wang, H., Raupach, M., and Collatz,
989 G. J.: Recent pause in the growth rate of atmospheric CO₂ due to enhanced terrestrial carbon
990 uptake, *Nature Communications*, 7, 1–9, <https://doi.org/10.1038/ncomms13428>, 2016.

991 Keenan, T. F., Migliavacca, M., Papale, D., Baldocchi, D., Reichstein, M., Torn, M., and Wutzler, T.:
992 Widespread inhibition of daytime ecosystem respiration, *Nature Ecology and Evolution*, 3, 407–
993 415, <https://doi.org/10.1038/s41559-019-0809-2>, 2019.

994 Li, M., Cao, S., and Zhu, Z.: Spatiotemporally consistent global dataset of the GIMMS Normalized
995 Difference Vegetation Index (PKU GIMMS NDVI) from 1982 to 2020, *Earth System Science*
996 *Data Discussions*, 1–31, <https://doi.org/10.5194/essd-2023-1>, 2023.

997 Ma, H. and Liang, S.: Development of the GLASS 250-m leaf area index product (version 6) from
998 MODIS data using the bidirectional LSTM deep learning model, *Remote Sensing of*
999 *Environment*, 273, 112985, <https://doi.org/10.1016/j.rse.2022.112985>, 2022.

1000 Ma, H., Zeng, J., Chen, N., Zhang, X., Cosh, M. H., and Wang, W.: Satellite surface soil moisture
1001 from SMAP, SMOS, AMSR2 and ESA CCI: A comprehensive assessment using global ground-
1002 based observations, *Remote Sensing of Environment*, 231, 111215,
1003 <https://doi.org/10.1016/j.rse.2019.111215>, 2019.

1004 Ma, Y., Zhang, Z., Kang, Y., and Özdoğan, M.: Corn yield prediction and uncertainty analysis based
1005 on remotely sensed variables using a Bayesian neural network approach, *Remote Sensing of*
1006 *Environment*, 259, 112408, <https://doi.org/10.1016/j.rse.2021.112408>, 2021.

1007 Myneni, R., Knyazikhin, Y., and Park, T.: MCD15A3H MODIS/Terra+Aqua Leaf Area
1008 Index/FPAR 4-day L4 Global 500m SIN Grid V006 [Data set], NASA EOSDIS Land Processes
1009 DAAC., <https://doi.org/10.5067/MODIS/MCD15A3H.006>, 2015a.

1010 Myneni, R., Knyazikhin, Y., and Park, T.: MOD15A2H MODIS/Terra Leaf Area Index/FPAR 8-
1011 Day L4 Global 500m SIN Grid V006 [Data set], NASA EOSDIS Land Processes DAAC,
1012 <https://doi.org/10.5067/MODIS/MOD15A2H.006>, 2015b.

1013 O’Sullivan, M., Spracklen, D. V., Batterman, S. A., Arnold, S. R., Gloor, M., and Buermann, W.:
1014 Have Synergies Between Nitrogen Deposition and Atmospheric CO₂ Driven the Recent
1015 Enhancement of the Terrestrial Carbon Sink?, *Global Biogeochemical Cycles*, 33, 163–180,
1016 <https://doi.org/10.1029/2018GB005922>, 2019.

1017 O’Sullivan, M., Smith, W. K., Sitch, S., Friedlingstein, P., Arora, V. K., Haverd, V., Jain, A. K., Kato,
1018 E., Kautz, M., Lombardozzi, D., Nabel, J. E. M. S., Tian, H., Vuichard, N., Wiltshire, A., Zhu,
1019 D., and Buermann, W.: Climate-Driven Variability and Trends in Plant Productivity Over Recent
1020 Decades Based on Three Global Products, *Global Biogeochemical Cycles*, 34,
1021 <https://doi.org/10.1029/2020GB006613>, 2020.

1022 Pastorello, G., Trotta, C., Canfora, E., Chu, H., Christianson, D., Cheah, Y. W., Poindexter, C.,
1023 Chen, J., Elbashandy, A., Humphrey, M., Isaac, P., Polidori, D., Ribeca, A., van Ingen, C., Zhang,
1024 L., Amiro, B., Ammann, C., Arain, M. A., Ardö, J., Arkebauer, T., Arndt, S. K., Arriga, N.,
1025 Aubinet, M., Aurela, M., Baldocchi, D., Barr, A., Beamesderfer, E., Marchesini, L. B., Bergeron,
1026 O., Beringer, J., Bernhofer, C., Berveiller, D., Billesbach, D., Black, T. A., Blanken, P. D., Bohrer,
1027 G., Boike, J., Bolstad, P. V., Bonal, D., Bonnefond, J. M., Bowling, D. R., Bracho, R., Brodeur, J.,
1028 Brümmer, C., Buchmann, N., Burban, B., Burns, S. P., Buysse, P., Cale, P., Cavagna, M., Cellier,
1029 P., Chen, S., Chini, I., Christensen, T. R., Cleverly, J., Collalti, A., Consalvo, C., Cook, B. D.,
1030 Cook, D., Coursolle, C., Cremonese, E., Curtis, P. S., D’Andrea, E., da Rocha, H., Dai, X., Davis,

1031 K. J., De Cinti, B., de Grandcourt, A., De Ligne, A., De Oliveira, R. C., Delpierre, N., Desai, A.
1032 R., Di Bella, C. M., di Tommasi, P., Dolman, H., Domingo, F., Dong, G., Dore, S., Duce, P.,
1033 Dufre ne, E., Dunn, A., Du sek, J., Eamus, D., Eichelmann, U., ElKhidir, H. A. M., Eugster, W.,
1034 Ewenz, C. M., Ewers, B., Famulari, D., Fares, S., Feigenwinter, I., Feitz, A., Fensholt, R., Filippa,
1035 G., Fischer, M., Frank, J., Galvagno, M., Gharun, M., Gianelle, D., et al.: The FLUXNET2015
1036 dataset and the ONEFlux processing pipeline for eddy covariance data, *Scientific data*, 7, 225,
1037 <https://doi.org/10.1038/s41597-020-0534-3>, 2020.

1038 Pe uelas, J., Ciais, P., Canadell, J. G., Janssens, I. A., Fern andez-Mart nez, M., Carnicer, J.,
1039 Obersteiner, M., Piao, S., Vautard, R., and Sardans, J.: Shifting from a fertilization-dominated to a
1040 warming-dominated period, *Nature Ecology and Evolution*, 1, 1438–1445,
1041 <https://doi.org/10.1038/s41559-017-0274-8>, 2017.

1042 Piao, S., Wang, X., Park, T., Chen, C., Lian, X., He, Y., Bjerke, J. W., Chen, A., Ciais, P.,
1043 T mmervik, H., Nemani, R. R., and Myneni, R. B.: Characteristics, drivers and feedbacks of
1044 global greening, *Nature Reviews Earth and Environment*, 1, 14–27,
1045 <https://doi.org/10.1038/s43017-019-0001-x>, 2020.

1046 Reich, P. B., Hobbie, S. E., and Lee, T. D.: Plant growth enhancement by elevated CO2 eliminated
1047 by joint water and nitrogen limitation, *Nature Geoscience*, 7, 920–924,
1048 <https://doi.org/10.1038/ngeo2284>, 2014.

1049 Reichstein, M., Camps-Valls, G., Stevens, B., Jung, M., Denzler, J., Carvalhais, N., and Prabhat:
1050 Deep learning and process understanding for data-driven Earth system science, *Nature*, 566,
1051 195–204, <https://doi.org/10.1038/s41586-019-0912-1>, 2019.

1052 Ruehr, S., Keenan, T. F., Williams, C., Zhou, Y., Lu, X., Bastos, A., Canadell, J. G., Prentice, I. C.,
1053 Sitch, S., and Terrer, C.: Evidence and attribution of the enhanced land carbon sink, *Nat Rev*
1054 *Earth Environ*, 1–17, <https://doi.org/10.1038/s43017-023-00456-3>, 2023.

1055 Running, S., Mu, Q., and Zhao, M.: MOD17A2H MODIS/Terra Gross Primary Productivity 8-Day
1056 L4 Global 500m SIN Grid V006, <https://doi.org/10.5067/MODIS/MOD17A2H.006>, 2015.

1057 Ryu, Y., Jiang, C., Kobayashi, H., and Detto, M.: MODIS-derived global land products of shortwave
1058 radiation and diffuse and total photosynthetically active radiation at 5 km resolution from 2000,
1059 *Remote Sensing of Environment*, 204, 812–825, <https://doi.org/10.1016/j.rse.2017.09.021>,
1060 2018.

1061 Ryu, Y., Berry, J. A., and Baldocchi, D. D.: What is global photosynthesis? History, uncertainties and
1062 opportunities, *Remote Sensing of Environment*, 223, 95–114,
1063 <https://doi.org/10.1016/j.rse.2019.01.016>, 2019.

1064 Sabater, J. M.: ERA5-Land monthly averaged data from 1981 to present., Copernicus Climate
1065 Change Service (C3S) Climate Data Store (CDS), <https://doi.org/doi:10.24381/cds.68d2bb30>,
1066 2019.

1067 Schaaf, C. and Wang, Z.: MCD43C4 MODIS/Terra+Aqua BRDF/Albedo Nadir BRDF-Adjusted
1068 Ref Daily L3 Global 0.05Deg CMG V006 [Data set], NASA EOSDIS Land Processes DAAC.,
1069 <https://doi.org/10.5067/MODIS/MCD43C4.006>, 2015.

1070 Schwalm, C. R., Anderegg, W. R. L., Michalak, A. M., Fisher, J. B., Biondi, F., Koch, G., Litvak, M.,
1071 Ogle, K., Shaw, J. D., Wolf, A., Huntzinger, D. N., Schaefer, K., Cook, R., Wei, Y., Fang, Y.,
1072 Hayes, D., Huang, M., Jain, A., and Tian, H.: Global patterns of drought recovery, *Nature*, 548,
1073 202–205, <https://doi.org/10.1038/nature23021>, 2017.

1074 Smith, W. K., Reed, S. C., Cleveland, C. C., Ballantyne, A. P., Anderegg, W. R. L., Wieder, W. R.,
1075 Liu, Y. Y., and Running, S. W.: Large divergence of satellite and Earth system model estimates of
1076 global terrestrial CO2 fertilization, *Nature Climate Change*, 6, 306–310,
1077 <https://doi.org/10.1038/nclimate2879>, 2016.

1078 Still, C. J., Berry, J. A., Collatz, G. J., and DeFries, R. S.: Global distribution of C3 and C4
1079 vegetation: Carbon cycle implications, *Global Biogeochemical Cycles*, 17,
1080 <https://doi.org/10.1029/2001gb001807>, 2003.

1081 Still, C. J., Berry, J. A., Collatz, G. J., and DeFries, R. S.: ISLSCP II C4 Vegetation Percentage, in:
1082 Hall, Forrest G., G. Collatz, B. Meeson, S. Los, E. Brown de Colstoun, and D. Landis (eds.).
1083 ISLSCP Initiative II Collection. Data set., <http://dx.doi.org/10.3334/ORNLDAAC/932>, 2009.

1084 Stocker, B. D., Zscheischler, J., Keenan, T. F., Prentice, I. C., Peñuelas, J., and Seneviratne, S. I.:
1085 Quantifying soil moisture impacts on light use efficiency across biomes, *New Phytologist*, 218,
1086 1430–1449, <https://doi.org/10.1111/nph.15123>, 2018.

1087 Stocker, B. D., Zscheischler, J., Keenan, T. F., Prentice, I. C., Seneviratne, S. I., and Peñuelas, J.:
1088 Drought impacts on terrestrial primary production underestimated by satellite monitoring,
1089 *Nature Geoscience*, 12, 264–270, <https://doi.org/10.1038/s41561-019-0318-6>, 2019.

1090 Terrer, C., Jackson, R. B., Prentice, I. C., Keenan, T. F., Kaiser, C., Vicca, S., Fisher, J. B., Reich, P.
1091 B., Stocker, B. D., Hungate, B. A., Peñuelas, J., McCallum, I., Soudzilovskaia, N. A., Cernusak, L.
1092 A., Talhelm, A. F., Van Sundert, K., Piao, S., Newton, P. C. D., Hovenden, M. J., Blumenthal, D.
1093 M., Liu, Y. Y., Müller, C., Winter, K., Field, C. B., Viechtbauer, W., Van Lissa, C. J., Hoosbeek,
1094 M. R., Watanabe, M., Koike, T., Leshyk, V. O., Polley, H. W., and Franklin, O.: Nitrogen and
1095 phosphorus constrain the CO2 fertilization of global plant biomass, *Nature Climate Change*, 9,
1096 684–689, <https://doi.org/10.1038/s41558-019-0545-2>, 2019.

1097 Thoning, K. W., Crotwell, A. M., and Mund, J. W.: Atmospheric Carbon Dioxide Dry Air Mole
1098 Fractions from continuous measurements at Mauna Loa, Hawaii, Barrow, Alaska, American
1099 Samoa and South Pole. 1973–2020, Version 2021-08-09, National Oceanic and Atmospheric
1100 Administration (NOAA), Global Monitoring Laboratory (GML), Boulder, Colorado, USA,
1101 <https://doi.org/10.15138/yaf1-bk21>, 2021.

1102 Tramontana, G., Ichii, K., Camps-Valls, G., Tomelleri, E., and Papale, D.: Uncertainty analysis of
1103 gross primary production upscaling using Random Forests, remote sensing and eddy covariance
1104 data, *Remote Sensing of Environment*, 168, 360–373, <https://doi.org/10.1016/j.rse.2015.07.015>,
1105 2015.

1106 Tramontana, G., Jung, M., Schwalm, C. R., Ichii, K., Camps-Valls, G., Ráduly, B., Reichstein, M.,
1107 Arain, M. A., Cescatti, A., Kiely, G., Merbold, L., Serrano-Ortiz, P., Sickert, S., Wolf, S., and
1108 Papale, D.: Predicting carbon dioxide and energy fluxes across global FLUXNET sites with
1109 regression algorithms, *Biogeosciences*, 13, 4291–4313, [https://doi.org/10.5194/bg-13-4291-](https://doi.org/10.5194/bg-13-4291-2016)
1110 2016, 2016.

1111 Ueyama, M., Ichii, K., Kobayashi, H., Kumagai, T., Beringer, J., Merbold, L., Euskirchen, E. S.,
1112 Hirano, T., Marchesini, L. B., Baldocchi, D., Saitoh, T. M., Mizoguchi, Y., Ono, K., Kim, J.,
1113 Varlagin, A., Kang, M., Shimizu, T., Kosugi, Y., Bret-Harte, M. S., Machimura, T., Matsuura, Y.,
1114 Ohta, T., Takagi, K., Takanashi, S., and Yasuda, Y.: Inferring CO2 fertilization effect based on
1115 global monitoring land-atmosphere exchange with a theoretical model, *Environmental Research*
1116 *Letters*, 15, 84009, <https://doi.org/10.1088/1748-9326/ab79e5>, 2020.

1117 Villarreal, S. and Vargas, R.: Representativeness of FLUXNET Sites Across Latin America, *Journal*
1118 *of Geophysical Research: Biogeosciences*, 126, e2020JG006090,
1119 <https://doi.org/10.1029/2020JG006090>, 2021.

1120 Walker, A. P., De Kauwe, M. G., Bastos, A., Belmecheri, S., Georgiou, K., Keeling, R. F.,
1121 McMahan, S. M., Medlyn, B. E., Moore, D. J. P., Norby, R. J., Zaehle, S., Anderson-Teixeira, K.
1122 J., Battipaglia, G., Brienen, R. J. W., Cabugao, K. G., Cailleret, M., Campbell, E., Canadell, J. G.,
1123 Ciais, P., Craig, M. E., Ellsworth, D. S., Farquhar, G. D., Fatichi, S., Fisher, J. B., Frank, D. C.,
1124 Graven, H., Gu, L., Haverd, V., Heilman, K., Heimann, M., Hungate, B. A., Iversen, C. M., Joos,
1125 F., Jiang, M., Keenan, T. F., Knauer, J., Körner, C., Leshyk, V. O., Leuzinger, S., Liu, Y.,

1126 MacBean, N., Malhi, Y., McVicar, T. R., Penuelas, J., Pongratz, J., Powell, A. S., Riutta, T., Sabot,
 1127 M. E. B., Schleucher, J., Sitch, S., Smith, W. K., Sulman, B., Taylor, B., Terrer, C., Torn, M. S.,
 1128 Treseder, K. K., Trugman, A. T., Trumbore, S. E., van Mantgem, P. J., Voelker, S. L., Whelan, M.
 1129 E., and Zuidema, P. A.: Integrating the evidence for a terrestrial carbon sink caused by increasing
 1130 atmospheric CO₂, *New Phytologist*, 229, 2413–2445, <https://doi.org/10.1111/nph.16866>, 2021.
 1131 Walther, S., Besnard, S., Nelson, J. A., El-madany, T. S., Migliavacca, M., Weber, U., Ermida, S. L.,
 1132 Brümmer, C., Schrader, F., Prokushkin, A. S., Panov, A. V., and Jung, M.: Technical note : A
 1133 view from space on global flux towers by MODIS and Landsat : The FluxnetEO dataset,
 1134 *Biogeosciences Discussions*, 2021.
 1135 Wan, Z., Hook, S., and Hulley, G.: MOD11A1 MODIS/Terra Land Surface
 1136 Temperature/Emissivity Daily L3 Global 1km SIN Grid V006 [Data set], NASA EOSDIS Land
 1137 Processes DAAC, <https://doi.org/10.5067/MODIS/MOD11A1.006>, 2015a.
 1138 Wan, Z., Hook, S., and Hulley, G.: MYD11A1 MODIS/Aqua Land Surface
 1139 Temperature/Emissivity Daily L3 Global 1km SIN Grid V006 [Data set], NASA EOSDIS Land
 1140 Processes DAAC, <https://doi.org/10.5067/MODIS/MYD11A1.006>, 2015b.
 1141 Warm Winter 2020 Team and ICOS Ecosystem Thematic Centre.: Warm Winter 2020 ecosystem
 1142 eddy covariance flux product for 73 stations in FLUXNET-Archive format—release 2022-1
 1143 (Version 1.0), <https://doi.org/10.18160/2G60-ZHAK>, 2022.
 1144 Wenzel, S., Cox, P. M., Eyring, V., and Friedlingstein, P.: Projected land photosynthesis constrained
 1145 by changes in the seasonal cycle of atmospheric CO₂, *Nature*, 538, 499–501,
 1146 <https://doi.org/10.1038/nature19772>, 2016.
 1147 Xiao, J., Zhuang, Q., Baldocchi, D. D., Law, B. E., Richardson, A. D., Chen, J., Oren, R., Starr, G.,
 1148 Noormets, A., Ma, S., Verma, S. B., Wharton, S., Wofsy, S. C., Bolstad, P. V., Burns, S. P., Cook,
 1149 D. R., Curtis, P. S., Drake, B. G., Falk, M., Fischer, M. L., Foster, D. R., Gu, L., Hadley, J. L.,
 1150 Hollinger, D. Y., Katul, G. G., Litvak, M., Martin, T. A., Matamala, R., McNulty, S., Meyers, T.
 1151 P., Monson, R. K., Munger, J. W., Oechel, W. C., Paw U, K. T., Schmid, H. P., Scott, R. L., Sun,
 1152 G., Suyker, A. E., and Torn, M. S.: Estimation of net ecosystem carbon exchange for the
 1153 conterminous United States by combining MODIS and AmeriFlux data, *Agricultural and Forest
 1154 Meteorology*, 148, 1827–1847, <https://doi.org/10.1016/j.agrformet.2008.06.015>, 2008.
 1155 Yan, K., Park, T., Yan, G., Chen, C., Yang, B., Liu, Z., Nemani, R. R., Knyazikhin, Y., and Myneni,
 1156 R. B.: Evaluation of MODIS LAI/FPAR product collection 6. Part 1: Consistency and
 1157 improvements, *Remote Sensing*, 8, 1–16, <https://doi.org/10.3390/rs8050359>, 2016a.
 1158 Yan, K., Park, T., Yan, G., Liu, Z., Yang, B., Chen, C., Nemani, R. R., Knyazikhin, Y., and Myneni,
 1159 R. B.: Evaluation of MODIS LAI/FPAR product collection 6. Part 2: Validation and
 1160 intercomparison, *Remote Sensing*, 8, 460, <https://doi.org/10.3390/rs8060460>, 2016b.
 1161 Yang, F., Ichii, K., White, M. A., Hashimoto, H., Michaelis, A. R., Votava, P., Zhu, A. X., Huete, A.,
 1162 Running, S. W., and Nemani, R. R.: Developing a continental-scale measure of gross primary
 1163 production by combining MODIS and AmeriFlux data through Support Vector Machine
 1164 approach, *Remote Sensing of Environment*, 110, 109–122,
 1165 <https://doi.org/10.1016/j.rse.2007.02.016>, 2007.
 1166 Yang, R., Wang, J., Zeng, N., Sitch, S., Tang, W., McGrath, M. J., Cai, Q., Liu, D., Lombardozzi, D.,
 1167 Tian, H., Jain, A. K., and Han, P.: Divergent historical GPP trends among state-of-the-art multi-
 1168 model simulations and satellite-based products, *Earth System Dynamics*, 13, 833–849,
 1169 <https://doi.org/10.5194/esd-13-833-2022>, 2022.
 1170 Yuan, H., Dai, Y., Xiao, Z., Ji, D., and Shangguan, W.: Reprocessing the MODIS Leaf Area Index
 1171 products for land surface and climate modelling, *Remote Sensing of Environment*, 115, 1171–
 1172 1187, <https://doi.org/10.1016/j.rse.2011.01.001>, 2011.

1173 Zeng, J., Matsunaga, T., Tan, Z.-H., Saigusa, N., Shirai, T., Tang, Y., Peng, S., and Fukuda, Y.:
1174 Global terrestrial carbon fluxes of 1999–2019 estimated by upscaling eddy covariance data with a
1175 random forest, *Sci Data*, 7, 313, <https://doi.org/10.1038/s41597-020-00653-5>, 2020.

1176 Zeng, N., Zhao, F., Collatz, G. J., Kalnay, E., Salawitch, R. J., West, T. O., and Guanter, L.:
1177 Agricultural Green Revolution as a driver of increasing atmospheric CO₂ seasonal amplitude,
1178 *Nature*, 515, 394–397, <https://doi.org/10.1038/nature13893>, 2014.

1179 Zhang, Y.: A global spatially contiguous solar-induced fluorescence (CSIF) dataset using neural
1180 networks (2000-2020), National Tibetan Plateau Data Center,
1181 <https://doi.org/10.11888/Ecolo.tpd.c.271751.>, 2021.

1182 Zhang, Y., Joiner, J., Hamed Alemohammad, S., Zhou, S., and Gentine, P.: A global spatially
1183 contiguous solar-induced fluorescence (CSIF) dataset using neural networks, *Biogeosciences*, 15,
1184 5779–5800, <https://doi.org/10.5194/bg-15-5779-2018>, 2018.

1185 Zheng, Y., Shen, R., Wang, Y., Li, X., Liu, S., Liang, S., Chen, J. M., Ju, W., Zhang, L., and Yuan,
1186 W.: Improved estimate of global gross primary production for reproducing its long-Term
1187 variation, 1982-2017, *Earth System Science Data*, 12, 2725–2746, [https://doi.org/10.5194/essd-](https://doi.org/10.5194/essd-12-2725-2020)
1188 [12-2725-2020](https://doi.org/10.5194/essd-12-2725-2020), 2020.

1189 Zhu, Z., Piao, S., Myneni, R. B., Huang, M., Zeng, Z., Canadell, J. G., Ciais, P., Sitch, S.,
1190 Friedlingstein, P., Arneth, A., Cao, C., Cheng, L., Kato, E., Koven, C., Li, Y., Lian, X., Liu, Y.,
1191 Liu, R., Mao, J., Pan, Y., Peng, S., Peuelas, J., Poulter, B., Pugh, T. A. M., Stocker, B. D., Viovy,
1192 N., Wang, X., Wang, Y., Xiao, Z., Yang, H., Zaehle, S., and Zeng, N.: Greening of the Earth and
1193 its drivers, *Nature Climate Change*, 6, 791–795, <https://doi.org/10.1038/nclimate3004>, 2016.

1194
1195

1196 **References**

1197 Automatic citation updates are disabled. To see the bibliography, click Refresh in the Zotero tab.