

This preprint has been submitted to Freshwater Biology on the 24th of October 2019.



Macrophyte cover type and groundwater as the key drivers of the extremely high organic carbon concentration of soda pans

| | |
|-------------------------------|---|
| Journal: | <i>Freshwater Biology</i> |
| Manuscript ID | FWB-P-Jun-19-0305.R1 |
| Manuscript Type: | Standard Paper |
| Date Submitted by the Author: | n/a |
| Complete List of Authors: | Boros, Emil; Hungarian Academy of Sciences Centre for Ecological Research, GINOP Sustainable Ecosystems Group, Danube Research Institute, V-Balogh, Katalin; Hungarian Academy of Sciences Centre for Ecological Research, Balaton Limnological Institute, Vörös, Lajos; Hungarian Academy of Sciences Centre for Ecological Research, Balaton Limnological Institute Csitári, Bianka; Eötvös Loránd Tudományegyetem, Department of Microbiology; Uppsala University, Evolutionary Biology Centre, Limnology Szekely, Anna ; Uppsala University, Evolutionary Biology Centre, Limnology |
| Keywords: | Climate change < Applied Issues, Saline waters < Habitat, Chemical environment < Level of Organisation, Higher plants < Taxonomic Group / Assemblage, Greenhouse gases < Applied Issues, Ground water < Habitat, Temporary pools < Habitat, Nutrient cycling < Process / Approach / Methods |
| | |

1
2
3 **1** **Macrophyte cover type and groundwater as the key drivers of the**
4 **2** **extremely high organic carbon concentration of soda pans**
5
6 **3**

7
8 **4** Emil Boros^{1*}, Katalin V.-Balogh², Bianka Csitári^{3,4}, Lajos Vörös², Anna J. Székely⁴
9

10 **5**
11
12 **6** ¹ GINOP Sustainable Ecosystems Group, Danube Research Institute, MTA Centre for

13
14
15 **7** Ecological Research, Danube Research Institute, Klebelsberg Kuno str. 3. P.O. Box 35, H-

16
17 **8** 8237 Tihany, Hungary

18
19 **9** ² Balaton Limnological Institute, MTA Centre for Ecological Research, Klebelsberg Kuno str.

20
21
22 **10** 3. P.O. Box 35, H-8237 Tihany, Hungary

23
24 **11** ³ Department of Microbiology, ELTE Eötvös Loránd University, Pázmány Péter stny. 1/c., H-

25
26 **12** 1117 Budapest, Hungary

27
28
29 **13** ⁴ Evolutionary Biology Centre, Limnology, Uppsala University, Norbyvägen 18 D, SE-752

30
31 **14** 36 Uppsala, Sweden

32
33
34 **15**
35
36 **16** *Corresponding author; E-mail: drborose@gmail.com; boros.emil@okologia.mta.hu;
37
38 **17**
39
40 **18**

1
2
3 19 **Summary**
4
5
6 20
7
8 21

9
10 22
11
12 23
13
14 24
15
16 25
17 26
18
19 27
20
21 28
22
23 29
24
25 30
26
27 31
28
29 32
30
31 33
32
33 34
34
35 35
36
37 36
38
39 37
40
41 38
42
43 39
44
45 40
46
47 41
48
49 42
50
51 43
52
53 44
54
55 45
56
57 46
58
59 47
60

1. Endorheic soda pans are among the aquatic systems that have the highest dissolved organic carbon (DOC) content on the planet with concentrations reaching values close to 1 g L⁻¹. Considering the importance of inland waters in the global carbon cycle, the understanding of the drivers of such outstanding aquatic organic carbon pools is eminent. The soda pans of the Carpathian Basin present a wide variability of biotic and abiotic characteristics that provides an adequate system to assess the determinants of the extreme high DOC concentrations of soda pans. Here we demonstrate through a multisite comparison, a multiyear seasonal monitoring, and a laboratory experiment that the dissolved organic matter content of the highest DOC concentration soda pans is primarily of groundwater and emergent macrophyte origin.

2. More precisely, the multisite comparison of 14 soda pans revealed that the variation of colored dissolved organic matter (CDOM) content of the pans is partially explained by the CDOM content (21% of variation) and conductivity (14%) of the local groundwater indicating significant role of allochthonous terrestrial DOC sources. However, 46% of the variation in CDOM content of the studied soda pans could be accounted for the dominant type of emergent macrophyte with *Bolboshoenus maritimus* dominated macrophyte cover leading to higher CDOM content than *Phragmites australis*.

3. In line with the results of the multisite comparison, we demonstrated by a decomposition experiment that both *B. maritimus* and *P. australis* have the potential to release substantial amount of organic matter into soda pans. However, the organic matter release of *B. maritimus* is much more intensive than that of *P. australis* leading to twice as high DOC and 3.5-times higher CDOM concentrations. In general, considering previous organic matter release studies we concluded that *P. australis* is a relatively low

1
2
3 44 organic matter releaser emergent macrophyte, and therefore the species composition of
4
5 45 emergent macrophytes has to be considered in autochthonous plant-derived DOM
6
7 46 estimations.
8
9

10 47 4. Finally, the multi-year seasonal monitoring of two distinctive soda pans showed that the
11
12 48 high organic matter concentrations of the pans depends not only on their intrinsic
13
14 49 characteristics but also on interannual variability. More precisely, we demonstrated that
15
16 50 the highest CDOM and DOC concentrations occurred in a colored (i.e., brown, low
17
18 51 TSS) soda pan that had extensive (95%) macrophyte cover dominated by *B. maritimus*
19
20 52 in a period characterized by high pH due to low water levels, which were presumably
21
22 53 the consequence of increased evaporation due to decreased precipitation and above
23
24 54 average temperature. Considering the trends of climate change in Central-Europe (i.e.,
25
26 55 increased temperature and modified precipitation regimes), our results indicate that
27
28 56 extremely high organic matter concentrations in soda pans might become more frequent
29
30 57 in the near future.
31
32
33
34
35
36

37 59 *Keywords:* high pH, DOM, groundwater effect, emergent macrophytes, interannual variability
38
39
40 60
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

61 Introduction

62

63 As inland waters receive, process, store and emit carbon in globally significant amounts (Cole
64 *et al.*, 2007), mapping active carbon in even the smallest aquatic systems is necessary for
65 accurate global estimates of carbon cycling (Tranvik, Cole & Prairie, 2018). Dissolved organic
66 matter (DOM) is quantitatively the most significant pool of organic carbon in aquatic systems,
67 which is either derived from terrestrial sources (i.e., allochthonous organic matter) or from
68 biological material produced *in situ* by phytoplankton and macrophytes (i.e., autochthonous
69 OM) (Williamson *et al.*, 1999; Zhang *et al.*, 2013). Terrestrial DOM can reach surface waters
70 via surface inflow or runoff as well as via groundwater seepage (Wetzel, 2001; Grabs *et al.*,
71 2012; Einarsdottir, Wallin & Sobek, 2017). Therefore, the quantity and quality of DOM in
72 inland waters is not only influenced by internal processes but also by the characteristics of the
73 catchment area (e.g., vegetation, soil type, hydrology) (Wetzel, 2001; Kothawala *et al.*, 2014;
74 Sepp *et al.*, 2019).

75 In inland waters DOM is considered to be dominated (15–80%) by soluble humic substances
76 (i.e., fluvic acids and humic acids). While fulvic acids are soluble at any pH, the solubility of
77 humic acids depends on pH with no dissolution at lower pH ($\text{pH} < 2$) and complete solubility
78 at pH 13 (Aiken *et al.*, 1985). The pH of the water also affects the properties of the dissolved
79 humic substances with more aromatic and aliphatic humic acids being dissolved at higher pH
80 (Baglieri *et al.*, 2014). Humic acids also contribute to the fluorescence signal of colored
81 (chromophoric) DOM (CDOM) (Lapierre & Frenette, 2009). Generally, allochthonous DOM
82 is composed of more recalcitrant and highly colored humic substances than phytoplankton
83 produced autochthonous DOM (Tranvik L. J., 1988; Wetzel, 2001). However, the refractory
84 plant-derived CDOM can also originate from autochthonous sources such as littoral marshland
85 vegetation (i.e., emergent macrophytes) (Lapierre & Frenette, 2009). Humic substances are

1
2
3 86 considered to be recalcitrant due to their aromatic core, which is relatively resistant to microbial
4
5 87 degradation (Kellerman *et al.*, 2015). However, it has been shown that terrestrial carbon can be
6
7 88 substantial contributor to microbial biomass (Guillemette, McCallister & del Giorgio, 2015).
8
9
10 89 The global mean of dissolved organic carbon (DOC) concentration of lakes has been
11
12 90 determined to be 5.02 mg L⁻¹ (equivalent of 5.58 mg L⁻¹ total organic carbon) and influenced
13
14 91 by climatic factors (Chen *et al.*, 2015) as well as watershed characteristics (Sobek *et al.*, 2007).
15
16
17 92 Although the globally most abundant shallow lakes (Downing *et al.*, 2006; Verpoorter *et al.*,
18
19 93 2014) have been shown to have double mean DOC concentration than deep lakes (6.56 and
20
21 94 3.12 mg L⁻¹, respectively) (Chen *et al.*, 2015), the 90 mg L⁻¹ mean DOC concentration of soda
22
23 95 lakes can be considered as extreme even when compared to other highly productive aquatic
24
25 96 systems such as eutrophic lakes, marshes or bogs (Fig. 1a). In addition, when looking at single
26
27 97 measurements, soda pans are very likely the global record holders in DOC as occasionally DOC
28
29 98 concentrations close to 1 g L⁻¹ have been reported (Lake Nakuru: 980 mg L⁻¹ (Jirsa *et al.*, 2013);
30
31 99 Sósér: 988 mg L⁻¹ (Boros *et al.*, 2016) (Fig. 1b).
32
33
34
35 100 Soda lakes and pans can be found on all continents except Antarctica and represent the most
36
37 101 alkaline natural environments on Earth (Grant & Sorokin, 2011). Soda lakes are formed in
38
39 102 endorheic basins (i.e., limited drainage basins), where evaporation exceeds water outflow
40
41 103 (Warren, 2006) and the levels of calcium (Ca²⁺) and magnesium (Mg²⁺) are low, while sodium
42
43 104 (Na⁺) and carbonate species (HCO₃⁻ + CO₃²⁻) are high (Boros & Kolpakova, 2018). Within
44
45 105 Europe—to the best of our knowledge—soda pans are restricted to the Carpathian Basin and
46
47 106 are found only in Austria, Hungary and Serbia. The climatic conditions of the region (i.e.,
48
49 107 continental with influence of both oceanic and Mediterranean climate) in combination with the
50
51 108 shallowness of the pans cause high fluctuation in respect of water level and temperature. Water
52
53 109 level fluctuation is of particular interest as it results in intermittent hydroperiods for many pans
54
55 110 and affects the concentration of both organic and inorganic compounds. In the case of soda pans
56
57
58
59
60

1
2
3 111 of the Carpathian basin, groundwater inflow typically exceeds the surface-related watershed
4
5 112 inflow and precipitation (Boros, Ecsedi & Oláh, 2013). The analysis of 84 soda pans of the
6
7 113 region showed that many pans have extremely high DOC and CDOM concentrations (DOC:
8
9 114 median = 47; range: 20–664 mg L⁻¹; CDOM: median = 310; range: 20–7,100 mg Pt L⁻¹), which
10
11 115 indicates polyhumic (CDOM > 90 mg Pt L⁻¹) character (Boros *et al.*, 2017). Furthermore,
12
13 116 extreme high DOC concentrations have also been recorded (Fig. 1b) (Boros *et al.*, 2016).
14
15 117 Positive correlation between DOC and CDOM has been determined before for Carpathian soda
16
17 118 pans (V.-Balogh *et al.*, 2009; Boros *et al.*, 2013; 2016; 2017), while—despite the pH-dependent
18
19 119 solubility of humic substances—correlation between pH and DOC has not been demonstrated
20
21 120 (Boros *et al.*, 2017).

22
23
24
25
26 121 Although high organic matter content is an inherent property of most soda lakes and pans, the
27
28 122 regulators of CDOM variation across soda pans are understudied and the causes of extremely
29
30 123 high DOC and CDOM concentrations such as those measured in the pans of the Carpathian
31
32 124 basin, are not properly understood. Therefore, in this study we aim to identify and test the main
33
34 125 sources of CDOM and DOC in polyhumic soda pans. Namely, we hypothesize that CDOM
35
36 126 concentration of polyhumic soda pans is positively affected by the allochthonous CDOM
37
38 127 concentrations of local groundwater and the extent of emergent macrophyte cover.
39
40 128 Furthermore, based on our earlier observations (Boros *et al.*, 2017), we hypothesize that the
41
42 129 species composition of the autochthonous emergent macrophytes influences CDOM content
43
44 130 with cosmopolitan bulrush (*Bolboshoenus maritimus* (L.) Palla, Cyperaceae) dominated cover
45
46 131 leading to higher CDOM concentrations than common reed (*Phragmites australis* (Cav.) Trin.
47
48 132 *ex Steud.*, Poaceae) dominated cover. To disentangle the effect of different macrophyte species
49
50 133 on CDOM and DOC concentrations, experimental assessment of the amount and ratio of
51
52 134 CDOM and DOC released from *P. australis* and *B. maritimus* was performed. Finally, to
53
54 135 identify the environmental conditions that lead to extreme DOC and CDOM values in soda
55
56
57
58
59
60

1
2
3 136 pans, seasonal monitoring of a turbid and a colored polyhumic soda pans was conducted in two
4
5 137 separate years.
6

7
8 138

9 10 139 **Methods**

11 12 140 *Study sites and sampling*

13
14 141 The soda pans studied in this work are located in the central area of the Carpathian Basin, on
15
16 142 the interfluvial area of the Danube and Tisza rivers (Fig. 2). The water budget of the soda pans
17
18 143 in this area is highly influenced by evaporation, precipitation and groundwater influx, while the
19
20 144 surface water inflow from the watershed is negligible as usually no major watercourse enters
21
22 145 these systems (Boros *et al.*, 2013). The primary source of the high Na–HCO₃–Cl⁻ content of the
23
24 146 soda pans of the region is the discharge from upwelling deep saline groundwater, which is
25
26 147 enhanced in the pans by evaporation (Simon *et al.*, 2011). These soda pans can be categorized
27
28 148 into two groups based on their optical characteristics: the turbid type, where inorganic
29
30 149 suspended solids (ISS) are the main cause of turbidity or more precisely, the contribution of
31
32 150 ISS to light attenuation (K_d) exceeds 50%, and the colored (brown) type, where CDOM
33
34 151 contribution dominates K_d (> 50%) (Boros *et al.*, 2013). Submerged and floating macrophytes
35
36 152 are sparse or absent from the open water areas of both turbid and colored pans. However,
37
38 153 marshland vegetation (*Bolboschoeno-Phragmitetum*) characterized primarily by varying ratios
39
40 154 of emergent macrophyte species *B. maritimus* and *P. australis* can be found on their shoreline.
41
42 155 This study comprises of three parts: a multisite comparison of soda pans, a decomposition
43
44 156 experiment and a seasonal analysis of a turbid and a colored soda pan.
45
46

47
48
49
50
51 157 In the multisite comparison 14 natural pans were sampled between April and September of
52
53 158 2017 to assess the potential effect of groundwater and macrophytes on the organic matter
54
55 159 content. The pans were selected in order to cover a broad range in respect of turbidity, CDOM
56
57 160 and emergent macrophyte cover (Table 1). The coordinates of the location of the pans and the
58
59
60

1
2
3 161 groundwater wells are listed in Table S1. The chemical type of the pans was determined
4
5 162 following the guidelines of Boros & Kolpakova (2018). Two of the 14 pans were sampled in
6
7 163 the seasonal analysis: Zab-szék, a typical turbid pan (Table 1: Total Suspended Solids (TSS) =
8
9 164 1574 mg L⁻¹, CDOM = 364 mg Pt L⁻¹) and Sósér, a typical colored pan (TSS 83 mg L⁻¹,
10 165 CDOM = 2088 mg Pt L⁻¹) (Boros et al., 2013; 2016; 2017). The seasonal assessment of these
11
12 166 two pans was performed in 2014 and 2017 with samplings approximately every other week
13
14 167 from January to December in 2014 and from March to November in 2017. On two occasions in
15
16 168 2017 (11-Sep and 16-Oct) Zab-szék was completely dry, so no measurement was possible.
17
18 169 For all samplings water depth, conductivity and pH were determined on site in the open water
19
20 170 part of the pans using a centimeter-scale pole, and a WTW MultiLine P4 field instrument for
21
22 171 pH and conductivity, respectively. Samples were collected and transported to the lab for
23
24 172 CDOM, DOC, TSS, and fluorescence excitation emission matrix spectroscopy (EEMS)
25
26 173 measurement. The groundwater was assessed by sampling dug wells (mean depth: 3–5 meters)
27
28 174 located within 500 meters from the shoreline of each pan (Table 1). According to topographic
29
30 175 maps, all of the sampled wells were established decades ago. For the macrophyte DOM release
31
32 176 experiment, aboveground fraction of stems and leaves of *B. maritimus* and *P. australis*
33
34 177 specimens in their early senescing stage were collected from Sósér soda pan in October 2012.
35
36 178 After collection, the plant samples were washed with pan water and kept in clean plastic bags
37
38 179 and brought to the laboratory.
39
40
41
42
43
44
45
46
47
48

49 181 *Autochthonous emergent macrophyte cover assessment*

50
51 182 The cover of open water and marshland vegetation (*Bolboschoeno-Phragmitetum*) of pans and
52
53 183 the classification based on *P. australis* and *B. maritimus* ratio were estimated with remote-
54
55 184 sensing databases extracted from Google Satellite raster data via OpenLayers plugin and
56
57 185 complemented with local field observations in 2017. The GIS mapping procedure and spatial
58
59
60

1
2
3 186 calculations were carried out using ArcMap (Environmental Systems Research Institute 2013).
4
5 187 The proportion of emergent macrophyte cover (%) was calculated based on the ratio of open
6
7 188 water and macrophyte cover (ha) and the characteristic macrophyte type of each soda pan was
8
9 189 determined by estimation of the ratio of the most common species (*P. australis* and *B.*
10
11 190 *maritimus*). Pans with ratio of either species above 50% were classified as either *P. australis* or
12
13 191 *B. maritimus* dominated (Table 1).
14
15
16
17 192

19 193 *Experimental release of DOM from macrophytes*

21 194 In the laboratory the plant material was cut into 20-cm-long pieces, oven-dried to constant
22
23 195 weight at 35°C to avoid the destruction of the associated microbiome, and stored at room
24
25 196 temperature (23°C) until the experiment. Fifty grams of oven-dried plant material (stems and
26
27 197 leaves) were placed into 5-liter bottles. The bottles were filled with 3.5 l water (pH = 8.4)
28
29 198 collected from the well near Kelemen-szék (Table 1, N: 46.8012; E: 19.1717). The well-water
30
31 199 was filtered with pre-combusted GF-5 acid-washed glass fiber filter (pore size = 0.4 µm). The
32
33 200 DOC and CDOM concentration of the well water were determined as for other samples. The
34
35 201 incubation was performed in three replicates in the dark at room temperature (22-24 °C) for 29
36
37 202 days. The bottles were aerated with sterile-filtered atmospheric air and the dissolved oxygen
38
39 203 saturation ranged between 64% and 94% throughout the experiment. This aerobic treatment
40
41 204 was chosen, because an earlier study on *P. australis* showed that aerobic conditions are more
42
43 205 similar to the conditions in the field than anaerobic (V.-Balogh *et al.*, 2006). On days 0, 1, 4, 7,
44
45 206 11, 14, 18, 22, 25 and 29, 100 ml of water were sampled from the bottles for analyses. The
46
47 207 water volume of the samples was always replaced, consequently the water volume remained
48
49 208 constant during the experiment. All glassware used for sample collection and analytical
50
51 209 processes was acid-washed and Milli-Q water rinsed.
52
53
54
55
56
57
58 210
59
60

1
2
3 211 *CDOM, DOC, and TSS measurements*
4

5 212 The water samples were filtered as for the DOM release experiment and the concentration of
6
7 213 colored dissolved organic matter (CDOM) was expressed as Pt (platina) units (mg Pt L^{-1}) using
8
9
10 214 absorbance (440 nm) measurements with a Shimadzu UV 160A spectrophotometer (Cuthbert
11
12 215 & del Giorgio, 1992). For DOC analyses, the filtered samples were acidified (to pH 2 with HCl)
13
14 216 and bubbled to remove dissolved inorganic carbon (DIC), and DOC concentration was
15
16
17 217 measured by thermal catalysis at 1050 °C in an Elementar High TOC instrument equipped with
18
19 218 a platinum cartridge using synthetic air as carrier gas. The concentration of total suspended
20
21 219 solids (TSS) was measured by filtering water (100–2000 ml) through pre-dried and pre-weighed
22
23
24 220 cellulose acetate filters (pore size = 0.45 μm) followed by oven-drying at 105°C, and weighing
25
26 221 of dry filters using a CHYO YMC SM-200 analytical balance (accuracy 0.01 mg, precision
27
28 222 0.02 mg).
29

30
31 223

32
33 224 *Fluorescence Excitation Emission Matrix Spectroscopy*
34

35 225 During the 2017 seasonal comparison of Sósér and Zab-szék samples were collected for
36
37 226 dissolved organic matter (DOM) characterization using fluorescence excitation-emission
38
39
40 227 matrix spectroscopy (EEMS). The samples were filtered through a 0.1- μm pore size Millipore
41
42 228 Isopore Membrane Filters the same day and stored in combusted glassware at 4°C until
43
44 229 processing. EEMS profiles of the samples were determined as in Kothawala *et al.* (2014).
45
46
47 230 Briefly, excitation-emission matrices (EEM) were determined by measuring UV-visible
48
49 231 absorbance spectra using a Lambda 40 UV-visible spectrophotometer (Perkin Elmer, Waltham,
50
51 232 MA, United States) and measuring fluorescence emission using a fluorescence
52
53 233 spectrophotometer (SPEX FluoroMax-2, Horiba Jobin Yvon, Kyoto, Japan). Milli-Q water was
54
55
56 234 used as blank and its values were subtracted from the EEMs. Manufacturer supplied instrument
57
58 235 correction factors and the absorbance spectra were used for the correction of instrument and
59
60

1
2
3 236 filter biases, respectively, while fluorescence intensity was calibrated to the Raman area of the
4
5 237 blank water.

6
7
8 238

9
10 239 *Data analyses*

11
12 240 To assess the effect of macrophytes on the CDOM content of the pans, the amount of non-
13
14 241 groundwater related CDOM (i.e., CDOM_{diff}) was calculated by subtraction of the CDOM
15
16 242 concentration of the corresponding groundwater wells from the total CDOM concentration of
17
18
19 243 pans.

20
21 244 From the EEMS results three indicators have been calculated: fluorescence index (FI), freshness
22
23 245 index (FRESH), and humification index (HIX). The FI is used as an indicator of the source of
24
25 246 DOM: high FI ~1.8 indicates microbial and algal origin of DOM, while low FI ~1.2 suggests
26
27 247 terrestrial plant and soil derived DOM. The FRESH index is used as an estimator of how
28
29 248 recently the DOM has been produced ('freshness'). Finally, HIX indicates the humification of
30
31 249 DOM (i.e., amount of aromatic compounds) (Fellman, Hood & Spencer, 2010).

32
33 250 In the multisite comparison, the variables influencing CDOM concentration in the pans were
34
35 251 evaluated by a linear model. The relevant variables for the linear model were selected from
36
37 252 those in Table 1 (except the distance of the well to the corresponding pan) by backward
38
39 253 selection using Akaike information criterion (AIC) estimator (Table S2). The linear model was
40
41 254 checked for multicollinearity and the variables that had the highest generalized variance
42
43 255 inflation factor (GVIF) were stepwise removed until reaching a model where the squared degrees
44
45 256 of freedom corrected GVIFs ($(GVIF^{1/(2*Df)})^2$) were less than five (Table S3). Differences
46
47 257 of measured variables and linear models were tested by Welch's *t*-tests and Analyses of
48
49 258 Variance (ANOVA), respectively. Correlations between parameters were evaluated with
50
51 259 Pearson's correlation analyses. To meet the assumptions of normality of residuals of parametric
52
53 260 tests and the requirements of linear correlations data was log-transformed when needed.
54
55
56
57
58
59
60

1
2
3 261 Normality of residuals and variables was checked by Shapiro-Wilk test and visual analysis. In
4
5 262 the seasonal study the relationships over time of CDOM and DOC with TSS, water depth, pH,
6
7 263 conductivity and the EEMS indexes were assessed on monthly averaged data using cross-
8
9
10 264 correlation analyses (Zab-szék in 2017 was excluded from these analyses due to repeated
11
12 265 complete droughts). Prior analyses the autocorrelation of the variables was removed by
13
14 266 differencing the series with a lag of one. The successful removal of autocorrelation was
15
16
17 267 confirmed by the Ljung-Box test. The Spearman correlation, non-linear curve fitting for
18
19 268 Michaelis-Menten kinetic function, Mann-Whitney, and Kruskal-Wallis tests used to analyze
20
21 269 the DOM release experiment were performed using OriginPro 9 (OriginLab, Northampton,
22
23 270 MA), while the EEMs data were analyzed with MATLAB. All other analyses were performed
24
25
26 271 using the R environment for statistical computing (R Core Team, 2015).
27
28
29 272

30 273 **Results**

31 274 *Multisite comparison of the effect of groundwater and macrophytes*

32
33 275 The identity, optical category (i.e., turbid or colored), chemical type (i.e., category based on
34
35 276 ionic composition, Boros & Kolpakova, 2018), TSS as a measure of turbidity, emergent
36
37 277 macrophyte cover and type (i.e., *Bolboschoeno* or *Phragmitetum* dominated), and the shortest
38
39
40 278 distance between the sampled groundwater wells and the shoreline of the corresponding pans
41
42
43 279 are listed in Table 1. The CDOM concentration was almost always higher in the pan (median =
44
45 280 287.6 mg Pt L⁻¹) than in the corresponding groundwater well (median = 112.2 mg Pt L⁻¹; paired
46
47 281 t-test logCDOM: $t = 3.912$; $p = 0.002$; Fig 3a). The pH of each pan (median = 9.28) was also
48
49 282 significantly higher than the pH of the corresponding well (median = 8.60; paired t-test $t =$
50
51
52 283 3.896, $p = 0.002$), while the difference between the conductivity of the pans and the
53
54 284 corresponding wells was not significant (median_{pans} = 3.0 g L⁻¹; median_{wells} = 1.8 g L⁻¹; paired
55
56
57 285 t-test $t = 2.086$, $p = 0.057$) (Table 1).
58
59
60

1
2
3 286 In the pans dominated by *B. maritimus* CDOM was slightly but not significantly higher than in
4
5 287 those dominated by *P. australis* (median_{*B.maritimus*} = 1059 mg Pt L⁻¹; median_{*P.australis*} = 253 mg
6
7 288 Pt L⁻¹; t-test t = 2.269, p = 0.053; Fig. 3c). In the case of the soda pan that completely lacked
9
10 289 macrophyte cover the CDOM concentration of the pan was very similar to the CDOM of the
11
12 290 groundwater well (Table 1: Unnamed pan). For both the pans dominated by *B. maritimus* and
13
14 291 *P. australis* CDOM concentration of the soda pans showed positive relationship to the CDOM
15
16 292 concentration of the wells (Fig. 3c), although this correlation was only significant in the case
17
18 293 *P. australis* dominated pans (Pearson's *B. maritimus*: r=0.643; p=0.168; *P. australis*: r=0.803;
19
20 294 p=0.029). For the pans with *B. maritimus* dominance the non-groundwater related CDOM_{diff}
21
22 295 (median=653.5 mg Pt L⁻¹) was also higher (t-test t = 2.603, p = 0.044) than for those dominated
23
24 296 by *P. australis* (median=114.94 mg Pt L⁻¹; Fig. 3b).

25
26 297 For the linear model, backward variable selection retained all assessed variables but the water
27
28 298 depth and the optical type (Table S2), while the chemical type and the pH of the pans were
29
30 299 removed to avoid collinearity (Table S3). The linear model using CDOM and pH of the wells,
31
32 300 conductivity of both pans and wells, TSS of then pans, and emergent macrophyte cover and
33
34 301 type (Table 1) explained 92.4% of the variation of CDOM content of the pans, and the ANOVA
35
36 302 (Fig. 3d, Table S4) revealed that the macrophyte type of the pans significantly explained 45.8%
37
38 303 of the variation (p = 0.004), while the CDOM content and the conductivity of the groundwater
39
40 304 wells significantly explained 21.2% (p = 0.006) and 13.9% (p = 0.012), respectively. The 9.8%
41
42 305 variance explained by the conductivity of the pans, and the 1.2% explained by the extent of the
43
44 306 emergent macrophyte cover was not significant (p = 0.052 and p = 0.416, respectively), while
45
46 307 the pH of the pans and TSS explained less than 1% and was not significant (p > 0.1).

47
48
49
50
51
52
53
54 308

55
56 309 *Experimental release of DOM from macrophytes*
57
58
59
60

1
2
3 310 After the 29-days-long incubation of the DOM release experiment, the plant material of *P.*
4
5 311 *australis* lost $7.33\pm 0.15\%$ (3.66 ± 0.075 g) dry mass, while *B. maritimus* lost $10.28\pm 2.79\%$
6
7 312 (5.14 ± 1.40 g). The initial CDOM concentration in the experimental bottles was 39.23 ± 2.22 mg
8
9 313 Pt L⁻¹ and by the end of the experiment it increased in average to 1190 mg Pt L⁻¹ and to 3900
10
11 314 mg Pt L⁻¹ for *P. australis* and *B. maritimus*, respectively. The total CDOM release from the
12
13 315 degraded plant material was 1136 mg Pt g⁻¹ dry weight loss and 2675 mg Pt g⁻¹ dry weight loss
14
15 316 for *P. australis* and *B. maritimus* respectively. According to the Michaelis-Menten kinetics, at
16
17 317 the end of the experiment the CDOM concentration was 90.9% and 86.7% of the possible
18
19 318 maximum (V_{\max}) for *P. australis* and *B. maritimus*, respectively (Fig. 4a).

20
21
22 319 The initial DOC concentration was 13.40 ± 0.099 mg L⁻¹ and by the end of the experiment
23
24 320 increased in average to 82 mg L⁻¹ and to 183 mg L⁻¹ for *P. australis* and *B. maritimus*
25
26 321 incubations, respectively. The total DOC released from the degraded plant material was 78 mg
27
28 322 g⁻¹ dry weight loss and 125 mg g⁻¹ dry weight loss for *P. australis* and *B. maritimus*,
29
30 323 respectively. According to the Michaelis-Menten kinetics, at the end of the experiment the DOC
31
32 324 concentration was 100.9% and 96.8% of the possible maximum (V_{\max}) for *P. australis* and *B.*
33
34 325 *maritimus*, respectively (Fig. 4b). The CDOM/DOC ratio increased linearly with significant
35
36 326 parameters from 9 to 14.5 and from 16 to 21 for *P. australis* and *B. maritimus*, respectively
37
38 327 (Fig. 4c).

39
40
41
42
43
44 328

45 329 *Seasonal comparison of a turbid and a colored pan*

46
47 330 As expected, the turbid Zab-szék had substantially higher turbidity than the colored Sósér soda
48
49 331 pan throughout most of the study period except for two timepoints in summer 2014 (Table 2,
50
51 332 Fig 5a). Water depth showed high variation in both pans (Fig 5b), although it was slightly higher
52
53 333 in Sósér than in Zab-szék (Table 2). Water depth changes also indicated that the hydrology of
54
55 334 the two studied years also followed different patterns as for both pans the beginning of the year
56
57
58
59
60

1
2
3 335 in 2014 was characterized by low water depth, which increased from September-October until
4
5 336 the end of the year. Meanwhile, in 2017 both pans had higher water levels at the beginning of
6
7 337 the year until August, when the water levels decreased and remained low until the end of the
8
9 338 study period (Fig 5b). Throughout the two years the pans differed in respect of pH but not
10
11 339 conductivity (Table 2, Fig 5c,d). Both CDOM and DOC were higher in Sósér pan than in Zab-
12
13 340 szék (Table 2, Fig 5e,f). The concentration of CDOM and DOC also differed between the two
14
15 341 years with higher values for both pans in 2014 than in 2017 (Table 2, Fig 5e,f) More precisely,
16
17 342 the highest CDOM and DOC levels were all measured in Sósér in 2014 between January and
18
19 343 August, which corresponded to the lowest water level period of the given year (Fig. 5b). In this
20
21 344 period the mean concentration of CDOM was 6,649 mg Pt L⁻¹ and 563 mg L⁻¹ for DOC, while
22
23 345 for the rest of the year was 1,294 mg Pt L⁻¹ and 111 mg L⁻¹ for CDOM and DOC, respectively.
24
25 346 The two pans also differed in respect of the indicators calculated from the EEMS (Fig 6). Sósér
26
27 347 had higher humification index (HIX) throughout the study period, while the freshness index
28
29 348 (FRESH) was higher at every timepoint in Zab-szék than in Sósér (Table 2, Fig 6b,c). In
30
31 349 addition, both pans had generally higher FRESH values in summer and Zab-szék had a
32
33 350 prominent FRESH peak in the beginning of summer (29-May) (Fig 6b). The fluorescence index
34
35 351 (FI) did not differ between the two pans (Table 2), although it showed much higher variation in
36
37 352 Zab-szék with prominent peaks in summer and autumn (Fig. 6a). For both pans the mean FI
38
39 353 index was close to the value expected for DOM of terrestrial and soil origin (Sósér: FI_{median} =
40
41 354 1.235; Zab-szék: FI_{median} = 1.236).

42
43 355 The cross-correlation analyses revealed that both CDOM and DOC were most strongly
44
45 356 correlated with TSS, depth, pH and conductivity measured at the same time (lag = 0) except for
46
47 357 DOC concentration in Zab-szék in 2014 and its correlation with TSS, depth and pH (Table S5).
48
49 358 In the case of Sósér, the lag = 0 correlations between DOC concentration and TSS, depth, pH
50
51 359 and conductivity were always significant (positive correlation with TSS, pH and conductivity,
52
53
54
55
56
57
58
59
60

1
2
3 360 and negative with mean)(Table 3). In Zab-szék DOC showed significant (positive) correlation
4
5 361 at lag = 0 only with depth and conductivity (Table 3), while for TSS only values measured 2
6
7 362 months earlier (lag = 2) showed significant positive correlation with actual DOC, and for pH
8
9 363 only values measured 5 months later (lag = -5) were significantly correlated with actual DOC
10
11 364 (Table S5). The CDOM concentration did not significantly correlate with the TSS of the pans,
12
13 365 while mean depth, pH and conductivity significantly correlated with CDOM in all cases but pH
14
15 366 in Sósér in 2017 and conductivity in Zab-szék in 2014. Interestingly, the sign of the correlations
16
17 367 with CDOM depended on the pan: in Sósér CDOM positively correlated with pH and
18
19 368 conductivity, and negatively with depth, while in Zab-szék CDOM negatively correlated with
20
21 369 pH and conductivity, and positively with depth. In the case of the EEMS indexes significant
22
23 370 correlations were only found for DOC in Sósér: FRESH showed significant positive correlation
24
25 371 with DOC measured the same time and HIX showed significant negative correlation with DOC
26
27 372 measured three months earlier (Table 3, Table S5). In the case of HIX the highest (non-
28
29 373 significant) correlations with CDOM in both pans were also detected with CDOM measured
30
31 374 three months earlier (Table S5).
32
33
34
35
36
37
38
39

40 376 **Discussion**

41
42 377 The results of our study indicate—as we hypothesized—that the CDOM variation in the soda
43
44 378 pans of the Carpathian basin is highly influenced by groundwater and the dominant species of
45
46 379 the macrophyte cover but—opposite to our hypothesis—it is not significantly driven by the
47
48 380 extent of the cover. Furthermore, we demonstrated that the variation of DOC and CDOM
49
50 381 concentrations of these pans is related to the variations in pH, conductivity and water depths,
51
52 382 and it is presumably influenced by yearly variations in hydroperiods and intrinsic pan properties
53
54 383 such as turbidity.
55
56
57
58
59
60

1
2
3 385 *Groundwater as allochthonous source of CDOM*
4

5 386 Plant-derived colored DOM in inland waters originates either from macrophytes or from
6
7 387 terrestrial vegetation of the watershed via surface or groundwater inflow (i.e., allochthonous
8
9 388 source) (V.-Balogh *et al.*, 1998; Wetzel, 2001; Lapierre & Frenette, 2009; Einarsdottir *et al.*,
10
11 389 2017). As no watercourse enters the endorheic intermittent soda pans studied here, the role of
12
13 390 surface inflow in their water budget is negligible (Boros *et al.*, 2013), and therefore their plant-
14
15 391 derived DOM content is expected to be primarily derived from macrophytes and groundwater
16
17 392 inflow. The multisite comparison of polyhumic soda pans demonstrated that the CDOM
18
19 393 concentration of pans is related to the CDOM concentration of groundwater (Fig. 3).
20
21 394 Groundwaters store and release carbon to surface waters in amounts that are meaningful for
22
23 395 global budgets (Downing & Striegl, 2018) and import of terrestrial carbon into lakes via
24
25 396 groundwater seepage has been shown to be substantial even in the case of boreal lakes with
26
27 397 hydrology dominated by surface water inflow (Einarsdottir *et al.*, 2017). In areas of lower relief,
28
29 398 shallow basins can be the focus of local discharge and evaporation from regionally extensive
30
31 399 groundwater systems (Deocampo & Jones, 2014) and modern continental evaporates like soda
32
33 400 lakes and pans typically accumulate within groundwater discharge (Warren, 2006). Such
34
35 401 groundwater driven hydrological processes have been also identified behind the formation of
36
37 402 soda pans of the Danube-Tisza Interfluve (Mádl-Szőnyi & Tóth, 2009; Simon *et al.*, 2011). The
38
39 403 groundwaters analyzed in this study had relatively high pH (median: 8.60, range: 7.02-9.56,
40
41 404 Table 1.), which—considering the pH-dependent dissolution of humic substances—is very
42
43 405 likely one of the reasons of their high CDOM content. Considering the high CDOM content of
44
45 406 local groundwaters and the importance of groundwater in the hydrology of the soda pans of the
46
47 407 region (Simon *et al.*, 2011), it is not surprising that groundwater CDOM explained a substantial
48
49 408 part of the variation of the CDOM concentration of the studied soda pans. Finally, the mean FI
50
51 409 indexes measured by EEMS in the seasonal monitoring were close to 1.2, which is considered
52
53
54
55
56
57
58
59
60

1
2
3 410 to indicate plant and soil derived organic matter (Fellman *et al.*, 2010) further enforcing the
4
5 411 partly terrestrial origin of the DOM of the pans.
6
7
8 412

9
10 413 *Autochthonous emergent macrophyte cover type influences CDOM content*
11

12 414 However, the relationship between CDOM content of the pans and the corresponding
13
14 415 groundwaters depended on the dominant type of emergent macrophytes of the pans with pans
15
16 416 dominated by *B. maritimus* having higher CDOM content than those dominated by *P. australis*
17
18 417 (Fig. 3). While in the case of lakes macrophyte derived DOC is considered to contribute only
19
20 418 to 1-20% of the total DOC, for wetlands and small lakes macrophytes role in the carbon budget
21
22 419 can be substantial (Sobek *et al.*, 2006; Reitsema, Meire & Schoelynck, 2018). However, in our
23
24 420 linear model macrophyte species dominance was much more important in explaining the
25
26 421 variance in CDOM content among the studied pans than the extent of the macrophyte cover
27
28 422 (45.8 vs 1.2% of variance explained, respectively). A possible explanation to this seemingly
29
30 423 surprising result is that in the multisite comparison we focused on CDOM and not total DOC,
31
32 424 which depends both on total DOC and the color of DOM. It has been shown that qualitatively
33
34 425 macrophyte-derived DOM depends on the type of macrophyte (Qu *et al.*, 2013). Accordingly,
35
36 426 it is reasonable that in the given dataset the dominant species of macrophytes is more important
37
38 427 in explaining the pans' CDOM content than the extent of emergent macrophyte cover.
39
40
41
42
43

44 428 The importance of the type of vegetation was corroborated by the results of the DOM release
45
46 429 experiment, which also showed that more than triple CDOM and more than double the amount
47
48 430 of DOC is released from *B. maritimus* than from *P. australis*. While DOM release from *P.*
49
50 431 *australis* has been studied before, to the best of our knowledge this is the first report for *B.*
51
52 432 *maritimus*. The DOC release results for *P. australis* presented here (82 mg L⁻¹) compare well
53
54 433 to those from a previous experiment conducted under similar conditions (75 mg L⁻¹) (V.-Balogh
55
56 434 *et al.*, 2006) indicating high reproducibility of our experiment. When compared to other
57
58
59
60

1
2
3 435 emergent macrophytes *P. australis* also released less DOC than crofton weed (*Eupatorium*
4
5 436 *adenophorum* (Spreng.) R.M.King & H.Rob., Asteraceae), water oats (*Zizania latifolia*
6
7 437 (Griseb.) Turcz., Poaceae), oriental pepper (*Polygonum orientale* (L.) Spach., Polygonaceae)
8
9 438 (Qu *et al.*, 2013), or seepweed (*Suaeda salsa* (L.) Pall., Amaranthaceae) (Qi, Xue & Wang,
10
11 439 2017). All these indicates that common reed (*P. australis*) is a relatively low DOC releasing
12
13 440 emergent macrophyte and accordingly, it is essential to consider the composition of emergent
14
15 441 macrophytes when assessing macrophyte impact on the carbon balance of aquatic systems.

16
17 442 The total release of DOC within the time of the experiment reached values close to the
18
19 443 maximum expected for both species (Fig. 4b), while the total CDOM release did not reach the
20
21 444 maximum expected for either of the species (Fig. 4a). This contrasting dynamic between
22
23 445 CDOM and DOC resulted in continuously increasing ratio of CDOM and DOC (Fig. 4c). To
24
25 446 better mimic natural conditions our experiment was not conducted under sterile conditions and
26
27 447 oxygen for aerobic microbial degradation was provided by aeration, therefore, the increase in
28
29 448 the ratio of CDOM suggest microbial degradation of the labile carbon fraction and concomitant
30
31 449 accumulation of recalcitrant colored DOM. The importance of microbial degradation in DOC
32
33 450 release was demonstrated by the experiment of Qi, Xue & Wang (2017) were the amount of
34
35 451 DOC released from *P. australis* under sterile conditions was nine times higher than the quantity
36
37 452 released without inhibition of microbial degradation.

38
39 453 Considering the implications of our experiment to natural conditions, we have to emphasize
40
41 454 that the pH of the water used in the experiment was lower (pH 8.40) than the pH of almost all
42
43 455 pans analyzed in the multisite comparison (Table 1; pH 9.36). In a decomposition experiment
44
45 456 increasing pH has been shown to correlate with increasing dry mass loss from macrophyte litter
46
47 457 (Krachler *et al.*, 2010), which together with the pH dependent dissolution of humic substances
48
49 458 suggests even higher release of DOC and CDOM from macrophytes in the pans than in our
50
51 459 experiment. However, the CDOM/DOC ratio at the end of the experiment was 14 times higher
52
53
54
55
56
57
58
59
60

1
2
3 460 for *P. australis* and 21 times higher for *B. maritimus* than the average ratio of 6.6 measured for
4
5 461 84 soda pans of the Carpathian Basin (Boros *et al.*, 2017). A possible explanation for higher
6
7 462 proportion of colorless organic matter in natural habitats is photochemical degradation of
8
9 463 CDOM, which has been demonstrated to be a significant contributor to organic carbon
10
11 464 mineralization in boreal lakes, and which is expected to be particularly important in shallow
12
13 465 high DOC lakes (Koehler *et al.*, 2014). In summary, it is likely that the autochthonous plant-
14
15 466 derived DOM of soda pans depends on the species composition of the emergent macrophytes
16
17 467 and the interplay of microbial and photochemical degradation.
18
19
20
21
22

23 469 *Potential drivers of the seasonal variation of CDOM and DOC*

24
25
26 470 The seasonal comparison of two soda pans with significant differences in respect of optical
27
28 471 characteristics (i.e., colored Sósér and turbid Zab-szék) is reasonable particularly considering
29
30 472 the potential importance of photochemical reactions. Sósér and Zab-szék also differ in respect
31
32 473 of emergent macrophyte cover (95% and 16%, respectively) and dominant macrophyte species
33
34 474 (*B. maritimus* and *P. phragmites*, respectively) (Table 1). Although both pans had much higher
35
36 475 CDOM and DOC concentration than those characteristic for other aquatic systems (Fig. 1a), in
37
38 476 the case of the colored pan the concentration of both CDOM and DOC was extremely high.
39
40 477 Based on the findings of the multisite comparison and DOC release experiment the extensive
41
42 478 *B. maritimus* dominated macrophyte cover supposedly contributed to the extreme CDOM and
43
44 479 DOC concentrations of Sósér. Meanwhile, compared to Sósér Zab-szék had slightly higher pH,
45
46 480 lower water depth and consequently intermittent hydrology with several droughts in 2017. The
47
48 481 two pans also differed based on the qualitative EEMS analyses of DOM. The higher HIX index
49
50 482 of Sósér suggests higher ratio of plant-derived recalcitrant DOM, while the higher FRESH
51
52 483 index of Zab-szék suggests greater ratio of freshly produced autochthonous DOM (Fig 6).
53
54
55
56
57
58
59
60

1
2
3 484 In addition, substantial differences in the hydrology and overall CDOM and DOC
4
5 485 concentrations of the two years monitored were also observed. A potential explanation of the
6
7 486 contrasting patterns of hydrology of the two years might be in the differences in temperature
8
9 487 and precipitation regimes of the study years and the preceding periods. Although weather
10
11 488 analyses are beyond the scope of this study, regional records show that between July 2013 and
12
13 489 June 2014 the weather was dryer and much warmer than usual. Specifically, the precipitation
14
15 490 between July and December of 2013 was only 75% of the monthly mean measured for the
16
17 491 region between 1981 and 2010 and in the following period (Jan-Jun 2014) the temperature was
18
19 492 2.3°C higher than the monthly mean (Fig. S1). Meanwhile, for the same period in 2016-2017
20
21 493 both temperature and precipitation were close to average (Fig. S1). Considering the importance
22
23 494 of groundwater in the hydrology of the soda pans in this region (Mádl-Szónyi & Tóth, 2009),
24
25 495 the combination of decreased precipitation followed by high temperature in 2013 and 2014
26
27 496 might have caused increased evapotranspiration from soils, leading to decreased groundwater
28
29 497 levels, and consequently less groundwater influx to the pans, which explains the lower water
30
31 498 levels measured for both pans in the beginning of 2014 compared to 2017. Similarly, the
32
33 499 opposite water level trends in autumn (i.e., increase in 2014 and decrease in 2017; Fig. 5b)
34
35 500 could be explained with precipitation and temperature anomalies such as the unusually high
36
37 501 precipitations in the second half of 2014 (156% of average precipitation; Fig. S1).

38
39 502 In Sósér the lower water level period of 2014 coincided with the highest DOC and CDOM
40
41 503 concentrations measured (Fig. 5e,f). In this period the mean DOC concentration of the pan was
42
43 504 563 mg L⁻¹, which is more than an order of magnitude higher than the average of wetlands (Fig.
44
45 505 1a) and remarkable even among soda lakes and pans (e.g. Jirsa *et al.*, 2013; Boros *et al.*, 2017).
46
47 506 Both the CDOM and DOC concentrations of Sósér were negatively correlated with water depth
48
49 507 throughout the study period (Table 3), which suggests that evaporation driven concentration
50
51 508 contributed to the extremely high CDOM and DOC levels measured here. This is supported by
52
53
54
55
56
57
58
59
60

1
2
3 509 the positive correlation of CDOM and DOC to conductivity and pH as both are also expected
4
5 510 to increase with decreasing water levels due to the concentration of inorganic ions. Furthermore,
6
7 511 considering the pH dependence of the dry matter loss from macrophyte litter (Krachler *et al.*,
8
9 512 2010), the higher pH of the period could have further aggravated the OM release from the dense
10
11 513 *B. maritimus* dominated macrophyte cover of Sósér contributing to the measured record high
12
13 514 DOC levels. Turbidity positively correlated with DOC suggesting that in this otherwise low
14
15 515 turbidity pan (TSS in Zab-szék was almost 20 times higher than in Sósér; Fig. 5a), where
16
17 516 organic carbon is the main cause of turbidity opposite to clay minerals, which are typical for
18
19 517 turbid type pans such as Zab-szék (Boros *et al.*, 2013).

20
21
22
23 518 On the other hand, in Zab-szék the correlation between depth and both CDOM and DOC was
24
25 519 positive (Table 3) suggesting that in this groundwater fed pan, possible groundwater itself was
26
27 520 the primary source of CDOM and DOC. This explanation was also supported by the results of
28
29 521 the multisite comparison, where for this pan—exceptionally among the other studied pans—
30
31 522 the CDOM content of the nearby groundwater well was almost double of the CDOM content
32
33 523 of the pan (Table 1). Interestingly, in this pan CDOM and DOC correlated differently with pH
34
35 524 and conductivity: for CDOM the correlation was negative, while for DOC it was positive.
36
37 525 Considering the lower pH and conductivity of the nearby groundwater well, their negative
38
39 526 correlations with CDOM agree with its groundwater origin theory. However, the positive
40
41 527 correlations of DOC and conductivity are seemingly contradictory to this theory. As DOC
42
43 528 measures both labile and recalcitrant organic carbon, while CDOM reflects more recalcitrant
44
45 529 organic carbon, a potential explanation to the increasing DOC content at higher conductivity
46
47 530 could be the presence of salinity dependent biodegradation inhibition as salinity is one of the
48
49 531 strongest microbial inhibitors (Székely *et al.* 2013). However, it is also possible that the higher
50
51 532 DOC but lower CDOM concentrations at high conductivity reflect non-humic freshly produced
52
53 533 autochthonous DOM, which together with the lesser macrophyte coverage and the *P.*
54
55
56
57
58
59
60

1
2
3 534 *phragmites* dominance suggests that the influence of phytoplankton and microbial communities
4
5 535 on the DOM of this pan could also be prominent. This explanation is enforced by the contrasting
6
7 536 microbial community of the two pans described by Szabó *et al.* (2017) and by the peaks of FI
8
9 537 index in Zab-szék (Fig 6a), which are potential indicators of phytoplankton and microbial
10
11 538 blooms suggesting greater importance of microbial processes in the carbon dynamics of this
12
13
14 539 pan than in Sósér.

15
16
17 540 In a broader perspective, our results demonstrate that in the case of endorheic water bodies—
18
19 541 particularly those lacking surface inflow—groundwater can be an important source of organic
20
21 542 carbon that should be accounted for in carbon budget calculations. We also showed that
22
23 543 emergent macrophytes are essential sources of recalcitrant organic carbon. Although when
24
25 544 estimating macrophyte effect, species composition has to be also considered since common
26
27 545 reed (*P. australis*), one of the most common emergent macrophytes on a global scale has
28
29 546 relatively low organic carbon release compared to other species such as the cosmopolitan
30
31 547 bulrush (*B. maritimus*), for which this study comprises the first report of experimental
32
33 548 decomposition measurements. Finally, we demonstrated that the record high DOC values (0.5-1
34
35 549 g L⁻¹) measured in the soda pans of the Carpathian basin are the result of the interplay of
36
37 550 intrinsic soda pan characteristics such as *B. maritimus* dominated macrophyte cover and most
38
39 551 importantly persistent low water levels that occur in consequence of weather anomalies. More
40
41 552 precisely, we showed that high organic carbon content periods follow extreme warm and dry
42
43 553 seasons. Considering that such weather patterns might increase in frequency in the near future
44
45 554 due to the ongoing climate change, soda pans could become increasingly important hotspots of
46
47 555 terrestrial carbon processing urging further studies exploring the carbon biogeochemistry of
48
49 556 soda lakes.

50
51
52
53
54
55
56 557

57
58 558 **Acknowledgments**
59
60

1
2
3 559 The data collection, assessment and study was founded by European-financed Hungarian
4
5 560 Economic Development and Innovation Operative Programme (GINOP-2.3.2-15-2016-00019)
6
7
8 561 at the GINOP Sustainable Ecosystems Group of Hungarian Academy of Sciences (MTA)
9
10 562 Centre for Ecological Research, as well as by the Swedish Research Council Formas through a
11
12 563 grant to A.J.S. and the Nation's Young Talent Scholarship from the Ministry of Human
13
14 564 Capacities (NTP-NFTÖ-18-B-0217) to B.C. We deeply thank to Tamás Sápi, Csaba Pigniczki
15
16
17 565 for assistance in the field at the Kiskunság National Park (Hungary) and we are also grateful to
18
19 566 Karólína Einarisdóttir and Marloes Groeneveld for their help with sample and data processing.
20

21 567

22 23 24 568 **Data Availability Statement**

25
26 569 The seasonal monitoring data and fluorescence spectra used for the EEMS analyses are
27
28 570 available on the online repository at: <http://urn.kb.se/resolve?urn=urn:nbn:se:uu:diva-385197>.
29

30 571

31 32 33 572 **Conflict of Interest Statement**

34
35 573 We declare that there is no actual or potential conflict of interest.
36

37 574

38 39 40 575 **References**

- 41
42 576 Aiken, G.R., McKnight, D.M., Wershaw, R.L., and MacCarthy P. (1985). Humic Substances
43
44 577 in Soil, Sediments, and Water. In: *Humic substances in soil, sediment and water:*
45
46 578 *Geochemistry, isolation and characterization*. (Ed. R.L.W. and P.M. G. R. Aiken, D. M.
47
48 579 McKnight), pp. 1–9. John Wiley & Sons, Inc.
49
50
51 580 Baglieri A., Vindrola D., Gennari M. & Negre M. (2014). Chemical and spectroscopic
52
53 581 characterization of insoluble and soluble humic acid fractions at different pH values.
54
55 582 *Chemical and Biological Technologies in Agriculture* **1**, 1–11.
56
57
58 583 Boros E., Ecsedi Z. & Oláh J. (2013). *Ecology and management of soda pans in the*
59
60

- 1
2
3 584 *Carpathian basin*. (Ed. Hortobágy Environmental Association), Balmazújváros.
- 4
5 585 Boros E., Pigniczki C., Sápi T., V.-Balogh K., Vörös L. & Somogyi B. (2016). Waterbird-
6
7 586 Mediated Productivity of Two Soda Pans in the Carpathian Basin in Central Europe.
8
9
10 587 *Waterbirds* **39**, 388–401.
- 11
12 588 Boros E., V.-Balogh K., Vörös L. & Horváth Z. (2017). Multiple extreme environmental
13
14 589 conditions of intermittent soda pans in the Carpathian Basin (Central Europe).
15
16 590 *Limnologica* **62**, 38–46.
- 17
18
19 591 Boros E. & Kolpakova M. (2018). A review of the defining chemical properties of soda lakes
20
21 592 and pans : An assessment on a large geographic scale of Eurasian inland saline surface
22
23 593 waters. 1–20. PLoS ONE 13(8): e0202205. <https://doi.org/10.1371/journal>.
- 24
25
26 594 Chen M., Zeng G., Zhang J., Xu P., Chen A. & Lu L. (2015). Global Landscape of Total
27
28 595 Organic Carbon, Nitrogen and Phosphorus in Lake Water. *Scientific Reports* **5**, 1–7.
- 29
30 596 Cole J.J., Prairie Y.T., Caraco N.F., McDowell W.H., Tranvik L.J., Striegl R.G., *et al.* (2007).
31
32 597 Plumbing the Global Carbon Cycle: Integrating Inland Waters into the Terrestrial Carbon
33
34 598 Budget. *Ecosystems* **10**, 172–185.
- 35
36
37 599 Cuthbert I.D. & del Giorgio P. (1992). Toward a standard method of measuring color in
38
39 600 freshwater. *Limnology and Oceanography* **37**, 1319–1326.
- 40
41
42 601 Deocampo D.M. & Jones B.F. (2014). Geochemistry of Saline Lakes Geochemistry of Saline
43
44 602 Lakes. In: *Treatise on Geochemistry*, 2nd edn. pp. 437–469. Elsevier Ltd.
- 45
46
47 603 Downing, J. A., Prairie, Y. T., Cole, J. J., Duarte, C. M., Tranvik, L. J., Striegl, R. G., ...
48
49 604 Middelburg, J. J. (2006). The global abundance and size distribution of lakes, ponds, and
50
51 605 impoundments. *Limnology and Oceanography*, 51, 2388–2397.
- 52
53
54 606 Downing J.A. & Striegl R.G. (2018). Size, age, renewal, and discharge of groundwater
55
56 607 carbon. *Inland Waters* **8**, 122–127.
- 57
58 608 Einarsdottir K., Wallin M.B. & Sobek S. (2017). High terrestrial carbon load via groundwater
59
60

- 1
2
3 609 to a boreal lake dominated by surface water inflow. *Journal of Geophysical Research:*
4
5 610 *Biogeosciences* **122**, 15–29.
6
7 611 Fellman J.B., Hood E. & Spencer R.G.M. (2010). Fluorescence spectroscopy opens new
8
9 612 windows into dissolved organic matter dynamics in freshwater ecosystems: A review.
10
11 613 *Limnology and Oceanography* **55**, 2452–2462.
12
13 614 Grabs T., Bishop K., Laudon H., Lyon S.W. & Seibert J. (2012). Riparian zone hydrology and
14
15 615 soil water total organic carbon (TOC): Implications for spatial variability and upscaling
16
17 616 of lateral riparian TOC exports. *Biogeosciences* **9**, 3901–3916.
18
19 617 Grant W. & Sorokin D.Y. (2011). Distribution and Diversity of Soda Lake Alkaliphiles. In:
20
21 618 *Extremophiles Handbook*. (Eds K. Horikoshi, G. Antranikian, A.T. Bull, F.T. Robb &
22
23 619 K.O. Stetter), pp. 27–54.
24
25 620 Guillemette F., McCallister S.L. & del Giorgio P.A. (2015). Selective consumption and
26
27 621 metabolic allocation of terrestrial and algal carbon determine allochthony in lakes. *The*
28
29 622 *ISME journal*, 1–10.
30
31 623 Jirsa F., Gruber M., Stojanovic A., Omondi S.O., Mader D., Körner W., *et al.* (2013). Major
32
33 624 and trace element geochemistry of Lake Bogoria and Lake Nakuru, Kenya, during
34
35 625 extreme draught. *Chemie der Erde* **73**, 275–282.
36
37 626 Kayranli B., Scholz M., Mustafa A. & Hedmark Å. (2010). Carbon storage and fluxes within
38
39 627 freshwater wetlands: A critical review. *Wetlands* **30**, 111–124.
40
41 628 Kellerman A.M., Kothawala D.N., Dittmar T. & Tranvik L.J. (2015). Persistence of dissolved
42
43 629 organic matter in lakes related to its molecular characteristics. *Nature Geoscience* **8**,
44
45 630 454–457.
46
47 631 Koehler B., Landelius T., Weyhenmeyer G.A., Machida N. & Tranvik L.J. (2014). Sunlight-
48
49 632 induced carbon dioxide emissions from inland waters Birgit. *Global Biogeochemical*
50
51 633 *Cycles*, 696–711.
52
53
54
55
56
57
58
59
60

- 1
2
3 634 Kothawala D.N., Stedmon C.A., Müller R.A., Weyhenmeyer G.A., Köhler S.J. & Tranvik L.J.
4
5 635 (2014). Controls of dissolved organic matter quality: evidence from a large-scale boreal
6
7 636 lake survey. *Global change biology* **20**, 1101–14.
- 8
9
10 637 Krachler R.F., Krachler R., Stojanovic A., Wielander B. & Herzig A. (2010). Effects of pH on
11
12 638 aquatic biodegradation processes. *Biogeosciences Discussions* **6**, 491–514.
- 13
14 639 Lapiere J.F. & Frenette J.J. (2009). Effects of macrophytes and terrestrial inputs on
15
16 640 fluorescent dissolved organic matter in a large river system. *Aquatic Sciences* **71**, 15–24.
- 17
18 641 Mádl-Szőnyi J. & Tóth J. (2009). A hydrogeological type section for the Duna-Tisza
19
20 642 Interfluve, Hungary. 961–980.
- 21
22 643 Qi Y., Xue Y. & Wang X. (2017). Release and Microbial Degradation of Dissolved Organic
23
24 644 Carbon and Nitrogen from *Phragmites australis* and *Suaeda salsa* in the Wetland of the
25
26 645 Yellow River Estuary. *Journal of Oceanography and Marine Research* **05**.
- 27
28 646 Qu X., Xie L., Lin Y., Bai Y., Zhu Y., Xie F., *et al.* (2013). Quantitative and qualitative
29
30 647 characteristics of dissolved organic matter from eight dominant aquatic macrophytes.
31
32 648 *Environmental science and pollution research international* **20**, 7413–7423.
- 33
34 649 Reitsema R.E., Meire P. & Schoelynck J. (2018). The Future of Freshwater Macrophytes in a
35
36 650 Changing World: Dissolved Organic Carbon Quantity and Quality and Its Interactions
37
38 651 With Macrophytes. *Frontiers in Plant Science* **9**, 1–15.
- 39
40 652 Saidy A.R., Smernik R.J., Baldock J.A., Kaiser K. & Sanderman J. (2013). Geoderma The
41
42 653 sorption of organic carbon onto differing clay minerals in the presence and absence of
43
44 654 hydrous iron oxide. *Geoderma* **209–210**, 15–21.
- 45
46 655 Sepp M., Kõiv T., Nõges P. & Nõges T. (2019). The role of catchment soils and land cover on
47
48 656 dissolved organic matter (DOM) properties in temperate lakes. *Journal of Hydrology*
49
50 657 **570**, 281–291.
- 51
52 658 Simon Sz., Mádl-Szőnyi J., Müller I., & Pogácsás Gy. (2011). Conceptual model for surface
53
54
55
56
57
58
59
60

- 1
2
3 659 salinization in an overpressured and a superimposed gravity-flow field, Lake
4
5 660 Kelemenszék area, Hungary. *Hydrogeology Journal* **19**, 701–717.
6
7 661 Sobek A.S., Söderbäck B., Karlsson S., Andersson E., Karlsson S., Andersson E., *et al.*
8
9 (2006). A Carbon Budget of a Small Humic Lake : An Example of the Importance of
10 662 Lakes for Organic Matter Cycling in Boreal Catchments A Carbon Budget of a Small
11 663 Humic Lake : An Example of the Importance of Lakes for Organic Matter Cycling in
12 664 Boreal Catchments. *AMBIO: A Journal of the Human Environment* **35**, 469–475
13
14 665
15 666 Sobek S., Tranvik L.J., Prairie Y.T., Kortelainen P. & Cole J.J. (2007). Patterns and
16 667 regulation of dissolved organic carbon: An analysis of 7,500 widely distributed lakes.
17 668 *Limnology and Oceanography* **52**, 1208–1219.
18
19 669 Szabó A., Korponai K., Kerepesi C., Somogyi B., Vörös L., Bartha D., *et al.* (2017). Soda
20 670 pans of the Pannonian steppe harbor unique bacterial communities adapted to multiple
21 671 extreme conditions. *Extremophiles* **21**, 1–11.
22
23 672 Székely, A. J., M. Berga, and S. Langenheder. 2013. Mechanisms determining the fate of
24 673 dispersed bacterial communities in new environments. *ISME J.* **7**: 61–71.
25
26 674 Thurman E.M. (1985). *Organic Geochemistry of Natural Waters*. Martinus Nijhoff/Dr. W.
27 675 Junk Publishers, Dordrecht, the Netherlands.
28
29 676 Tranvik L. J. (1988). Availability of dissolved organic carbon for planktonic bacteria in
30 677 oligotrophic lakes of differing humic content. *Microbial Ecology* **16**, 311–322
31
32 678 Tranvik L.J., Cole J.J. & Prairie Y.T. (2018). The study of carbon in inland waters-from
33 679 isolated ecosystems to players in the global carbon cycle. *Limnology and Oceanography*
34 680 *Letters* **3**, 41–48.
35
36 681 Verpoorter, C., T. Kutser, D. A. Seekell, and L. J. Tranvik. 2014. A global inventory of lakes
37 682 based on high-resolution satellite imagery. *Geophys. Res. Lett.* **41**: 6396–6402.
38
39 683 V.-Balogh K., Presing M., Hiripi L. & Vörös L. (1998). Stable carbon and nitrogen isotope
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

- 1
2
3 684 ratios of dissolved humic substances in a shallow reservoir covered by macrophytes.
4
5 685 *Internat. Rev. Hydrobiol.* **83**, 203–206
6
7
8 686 V.-Balogh K., Présing M., Vörös L. & Tóth N. (2006). A study of the decomposition of reed
9
10 687 (*Phragmites australis*) as a possible source of aquatic humic substances by measuring the
11
12 688 natural abundance of stable carbon isotopes. *International Review of Hydrobiology* **91**,
13
14 689 15–28.
15
16
17 690 V.-Balogh, K., Németh, B. & Vörös, L. (2009). Specific attenuation coefficients of optically
18
19 691 active substances and their contribution to the underwater ultraviolet and visible light
20
21 692 climate in shallow lakes and ponds. *Hydrobiologia* **632**, 91–105.
22
23
24 693 Warren J.K. (2006). Depositional chemistry and hydrology, *Evaporites: Sediments, Resources*
25
26 694 *and Hydrocarbons*. pp. 59–138. Springer-Verlag, Berlin, Heidelberg.
27
28
29 695 Wetzel R. (2001). *Limnology: Lake and River Ecosystems.*, Third. Academic Press, San
30
31 696 Diego.
32
33 697 Williamson C.E., Morris D.P., Pace M.L. & Olson O.G. (1999). Dissolved organic carbon and
34
35 698 nutrients as regulators of lake ecosystems: Resurrection of a more integrated paradigm.
36
37 699 *Limnology and Oceanography* **44**, 795–803.
38
39
40 700 Zhang Y., Liu X., Wang M. & Qin B. (2013). Compositional differences of chromophoric
41
42 701 dissolved organic matter derived from phytoplankton and macrophytes. *Organic*
43
44 702 *Geochemistry* **55**, 26–37.
45
46
47 703
48
49
50
51
52
53
54
55
56
57
58
59
60

1
2
3 704 *Figure legends*
4

5 705 **Fig. 1** Global mean (a) and maximum (b) concentrations of dissolved organic carbon (DOC) in
6 inland waters. Data sources: (Wetzel, 2001) for surface water; (Chen *et al.*, 2015) for lakes;
7 706 (Thurman, 1985) for lakes, swamps, marshes, bogs; Suhett et al (2004) for coastal lagoons;
8 707 (Boros *et al.*, 2013, 2016, 2017) for soda pans.
9
10
11
12
13

14 709
15
16 710 **Fig. 2** Location of studied soda pans within the Carpathian Basin based on Boros et al., 2017.
17
18

19 711
20
21 712 **Fig. 3** Multisite comparison of the factors related to colored dissolved organic matter (CDOM)
22 in 14 soda pans of the Carpathian Basin. (a) Comparison of CDOM concentration in soda pans
23 713 and nearby groundwater; (b) Difference between CDOM concentration of soda pans and
24 714 corresponding groundwater well by dominant macrophyte type of the pans (*Bolboschoenus*
25 715 *maritimus* or *Phragmites australis*); (c) Correlation of groundwater and soda pan CDOM. Filled
26 716 circles indicate *B. maritimus* and crosses *P. australis* dominated pans, while the empty circle
27 717 indicates no macrophyte cover. Dashed line depicts non-significant ($r = 0.643$; $p = 0.168$) and
28 718 solid line depicts significant ($r = 0.804$; $p = 0.029$) correlation; (d) Variance of CDOM of the
29 719 14 pans explained by the different parameters based on linear model (Table S4). Significance
30 720 of parameters according to ANOVA indicated by **: $p < 0.01$; *: $p < 0.05$; $p < 0.1$.
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45

46 722
47 723 **Fig. 4** Experimental organic matter release from *Bolboschoenus maritimus* and *Phragmites*
48 724 *australis* plant material. (a) Colored dissolved organic matter (CDOM) release. Non-linear
49 725 Michaelis-Menten kinetics-based curve fitting statistics for *B. maritimus*: $N=28$; $df=25$; $r^2=$
50 726 0.7881 ; $V_{max}=4495.6372$ $p \ll 0.0001$; $K_m=5.2482$ $p=0.0023$, and for *P. australis*: $N=28$; $df=26$;
51 727 $r^2= 0.8403$; $V_{max}=1308.8554$ $p \ll 0.0001$; $K_m=5.2816$ $p=0.0002$. (b) Dissolved organic matter
52 728 (DOC) release. Non-linear Michaelis-Menten kinetics-based curve fitting statistics for *B.*
53
54
55
56
57
58
59
60

1
2
3 729 *maritimus*: N=28; df=26; $r^2=0.9197$; $V_{\max}=189.4330$ $p \ll 0.0001$; $K_m=1.9014$ $p \ll 0.0001$, and
4
5 730 for *P. australis*: N=28; df=26; $r^2=0.8596$; $V_{\max}=81.2560$ $p \ll 0.0001$; $K_m=1.1827$ $p \ll 0.0001$.
6
7 731 (c) Changes CDOM/DOC. Linear curve fitting statistics for *B. maritimus*: N=28; df=26; $r^2=$
8
9 732 0.5630 ; intercept= 15.7668 $p \ll 0.0001$; slope= 0.2204 $p=0.0028$ and for *P. australis*: N=28;
10
11 733 df=26; $r^2=0.7529$; intercept= 9.3597 $p \ll 0.0001$; slope= 0.1896 $p \ll 0.0001$.
12
13
14
15 734

16
17 735 **Fig. 5** Multiyear comparison of the temporal changes of environmental parameters in Sósér and
18
19 736 Zab-szék soda pans in 2014 and 2017. (a) Total suspended solids (TSS); (b) Mean water depth;
20
21 737 (c) pH; (d) Electrical conductivity (EC); (e) Colored dissolved organic matter (CDOM); (f)
22
23 738 dissolved organic carbon (DOC).
24
25

26 739
27
28 740 **Fig. 6** Temporal changes of dissolved organic matter (DOM) characterizing indices derived
29
30 741 from fluorescence excitation emission matrix spectroscopy (EEMS) in Sósér and Zab-szék soda
31
32 742 pans in 2017. (a) Fluorescence index (FI) indicating the source of DOM with high values (FI \approx
33
34 743 1.8) suggesting DOM derived from extracellular release and leachate from bacteria and algae,
35
36 744 and low values (FI ≈ 1.2) indicating plant and soil derived organic matter. (b) Freshness index
37
38 745 (FRESH) indicating contribution of recently produced DOM. (c) Humification index (HIX)
39
40 746 indicating humic substance content or extent of humification of DOM (Fellman *et al.*, 2010).
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

747 Table 1. Environmental parameters determined for the soda pans and corresponding groundwater wells in the multisite comparison

| Name of pan | Soda pan | | | | | | | | Groundwater well | | | | |
|--------------|----------------|---------------|-------------------------------|---------------------------------------|---------------------------|--|-------|-------------------------------|------------------|--|------|-------------------------------|---|
| | Chemical type* | Optical type† | Emergent macrophyte cover (%) | Macrophyte type (species cover > 50%) | TSS (mg L ⁻¹) | Electrical conductivity (μS cm ⁻¹) | pH | CDOM (Pt mg L ⁻¹) | Water depth (cm) | Electrical conductivity (μS cm ⁻¹) | pH | CDOM (Pt mg L ⁻¹) | Distance to closest shore of soda pan (m) |
| Bogárczó | Soda | Turbid | 43 | <i>B. maritimus</i> | 807 | 4970 | 9.28 | 1424 | 30 | 5460 | 8.55 | 107 | 34 |
| Böddi-szék 1 | Soda-Saline | Turbid | 15 | <i>P. australis</i> | 745 | 6420 | 9.80 | 239 | 9 | 2810 | 8.83 | 74 | 24 |
| Böddi-szék 2 | Soda-Saline | Turbid | 5 | <i>P. australis</i> | 1100 | 1710 | 9.69 | 253 | 2 | 5100 | 9.56 | 117 | 16 |
| Böddi-szék 3 | Soda-Saline | Colored | 92 | <i>P. australis</i> | 8 | 1847 | 8.39 | 293 | 40 | 3400 | 9.29 | 303 | 239 |
| Büdös-szék | Soda | Turbid | 62 | <i>B. maritimus</i> | 384 | 5230 | 9.27 | 174 | 12 | 2200 | 7.73 | 28 | 190 |
| Csaba-szék | Soda | Colored | 52 | <i>B. maritimus</i> | 50 | 2620 | 8.57 | 1659 | 20 | 1863 | 8.34 | 841 | 25 |
| Dongér | Soda-Saline | Turbid | 11 | <i>P. australis</i> | 67 | 1625 | 10.57 | 99 | 20 | 757 | 8.64 | 67 | 31 |
| Fehér-szék | Soda | Turbid | 91 | <i>P. australis</i> | 585 | 3300 | 8.86 | 684 | 45 | 2310 | 8.74 | 437 | 155 |
| Fülöp-szék | Soda | Colored | 75 | <i>P. australis</i> | 350 | 4090 | 9.12 | 143 | 8 | 4010 | 7.02 | 28 | 198 |
| Kelemen-szék | Soda | Turbid | 30 | <i>B. maritimus</i> | 1850 | 3010 | 9.13 | 693 | 20 | 1146 | 8.15 | 205 | 201 |
| Sósér | Soda | Colored | 95 | <i>B. maritimus</i> | 83 | 8680 | 9.89 | 2088 | 3 | 3540 | 9.28 | 69 | 47 |
| Unnamed | Soda | Turbid | 0 | none | 692 | 4840 | 9.43 | 203 | 20 | 1609 | 8.73 | 146 | 85 |
| Vesszős-szék | Soda | Turbid | 72 | <i>B. maritimus</i> | 80 | 2050 | 9.73 | 282 | 20 | 1678 | 8.17 | 41 | 467 |
| Zab-szék | Soda | Turbid | 16 | <i>P. australis</i> | 1574 | 4270 | 9.28 | 364 | 3 | 1051 | 8.30 | 667 | 45 |

748 * Based on Boros & Kolpakova (2018)

749 † Based on Boros *et al.*, (2013, 2017)

750

751 Table 2. Mean and standard deviation of the measured parameters of the turbid and colored
 752 soda pan in the two years of seasonal monitoring

| Parameter | Sósér | | | | Zab-szék | | | |
|-----------|--------|--------|--------|-------|----------|--------|--------|--------|
| | 2014 | | 2017 | | 2014 | | 2017 | |
| | Mean | SD | Mean | SD | Mean | SD | Mean | SD |
| TSS | 180.8 | 293.0 | 68.7 | 71.1 | 2325.9 | 2201.9 | 2100.9 | 1618.3 |
| Depth | 18.6 | 15.3 | 26.8 | 14.3 | 13.1 | 12.1 | 11.3 | 9.6 |
| pH | 9.26 | 0.54 | 9.23 | 0.40 | 9.60 | 0.28 | 9.61 | 0.21 |
| EC | 5.72 | 3.45 | 6.49 | 3.79 | 5.64 | 3.66 | 7.05 | 6.68 |
| CDOM | 5888.7 | 2908.6 | 1628.2 | 508.7 | 470.4 | 200.8 | 279.1 | 127.3 |
| DOC | 437.3 | 268.5 | 214.6 | 108.4 | 69.1 | 48.0 | 49.4 | 10.8 |
| FI | - | - | 1.228 | 0.028 | - | - | 1.234 | 0.049 |
| FRESH | - | - | 0.608 | 0.030 | - | - | 0.736 | 0.091 |
| HIX | - | - | 17.47 | 1.689 | - | - | 7.143 | 1.436 |

753

754

755 Table 3. Results of cross-correlation at lag = 0 of colored dissolved organic matter (CDOM)
 756 and dissolved organic carbon (DOC) and environmental parameters of the turbid and colored
 757 soda pans. Significant correlations ($p < 0.05$) are in bold.

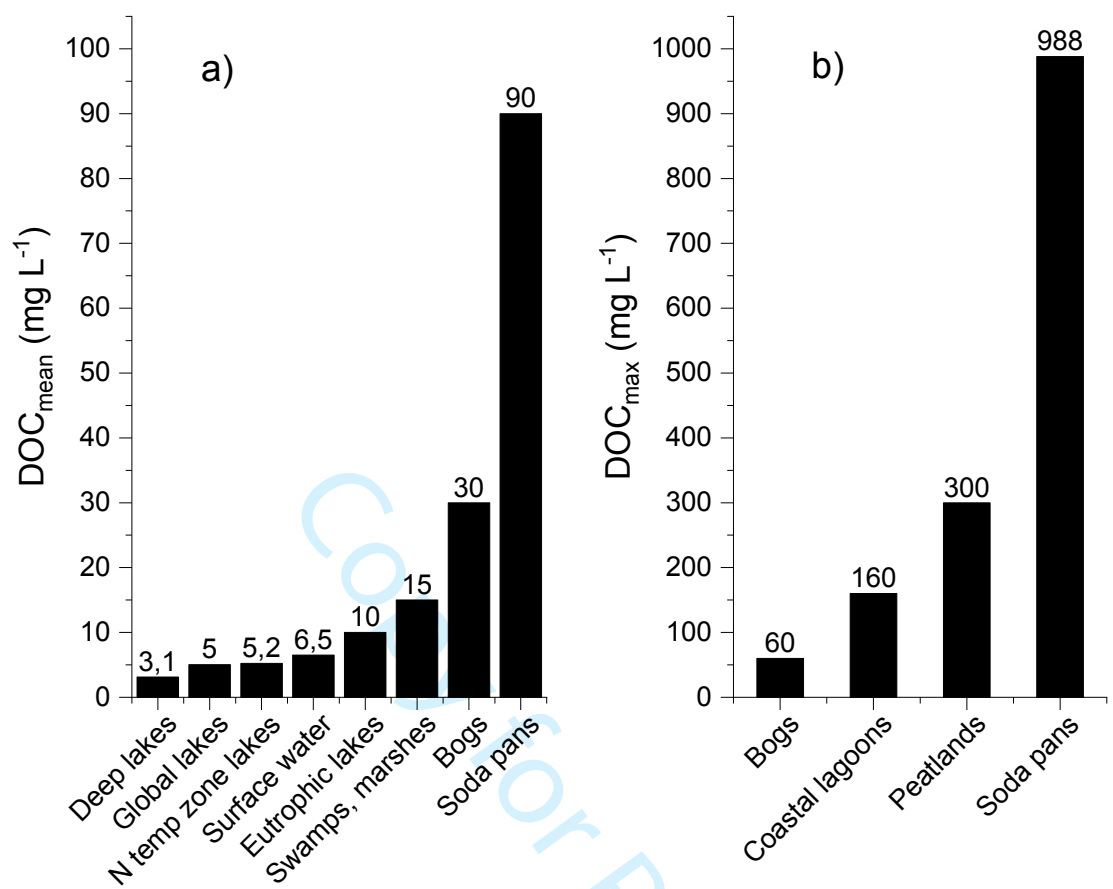
| | CDOM | | | DOC | | |
|--------------|---------------|---------------|---------------|---------------|---------------|--------------|
| | Sósér | | Zab-szék* | Sósér | | Zab-szék* |
| | 2014 | 2017 | 2014 | 2014 | 2017 | 2014 |
| TSS | 0.316 | 0.598 | 0.355 | 0.737 | 0.745 | 0.122 |
| Depth | -0.750 | -0.829 | 0.809 | -0.758 | -0.778 | 0.622 |
| pH | 0.778 | 0.563 | -0.792 | 0.930 | 0.839 | 0.414 |
| Conductivity | 0.802 | 0.660 | -0.447 | 0.994 | 0.870 | 0.951 |
| FI | | -0.297 | | | -0.568 | |
| FRESH | | 0.181 | | | 0.809 | |
| HIX | | -0.225 | | | -0.125 | |

758 * Measurements from 2017 for Zab-szék were omitted from these analyses because of missing
 759 data due to drought of the pan.

760

761

762 Figure 1

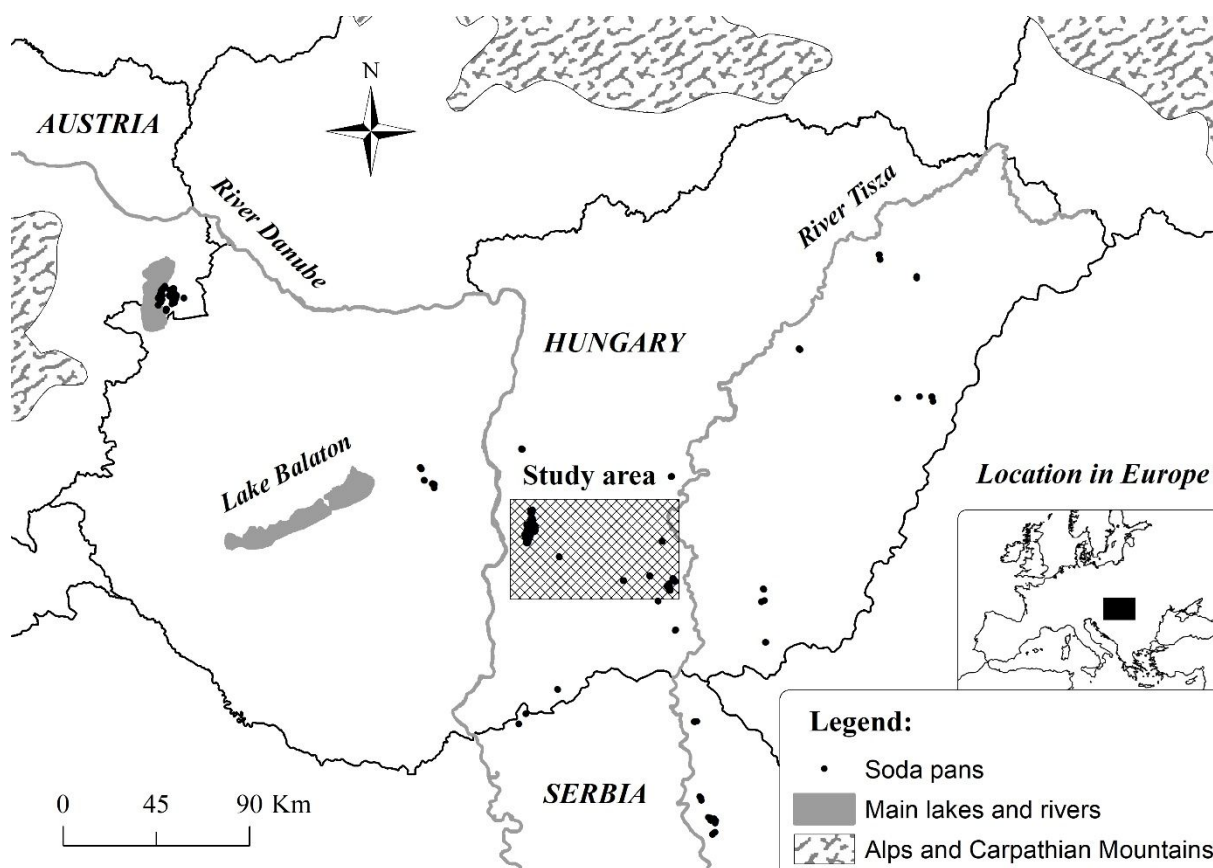


763

764

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

765 Figure 2



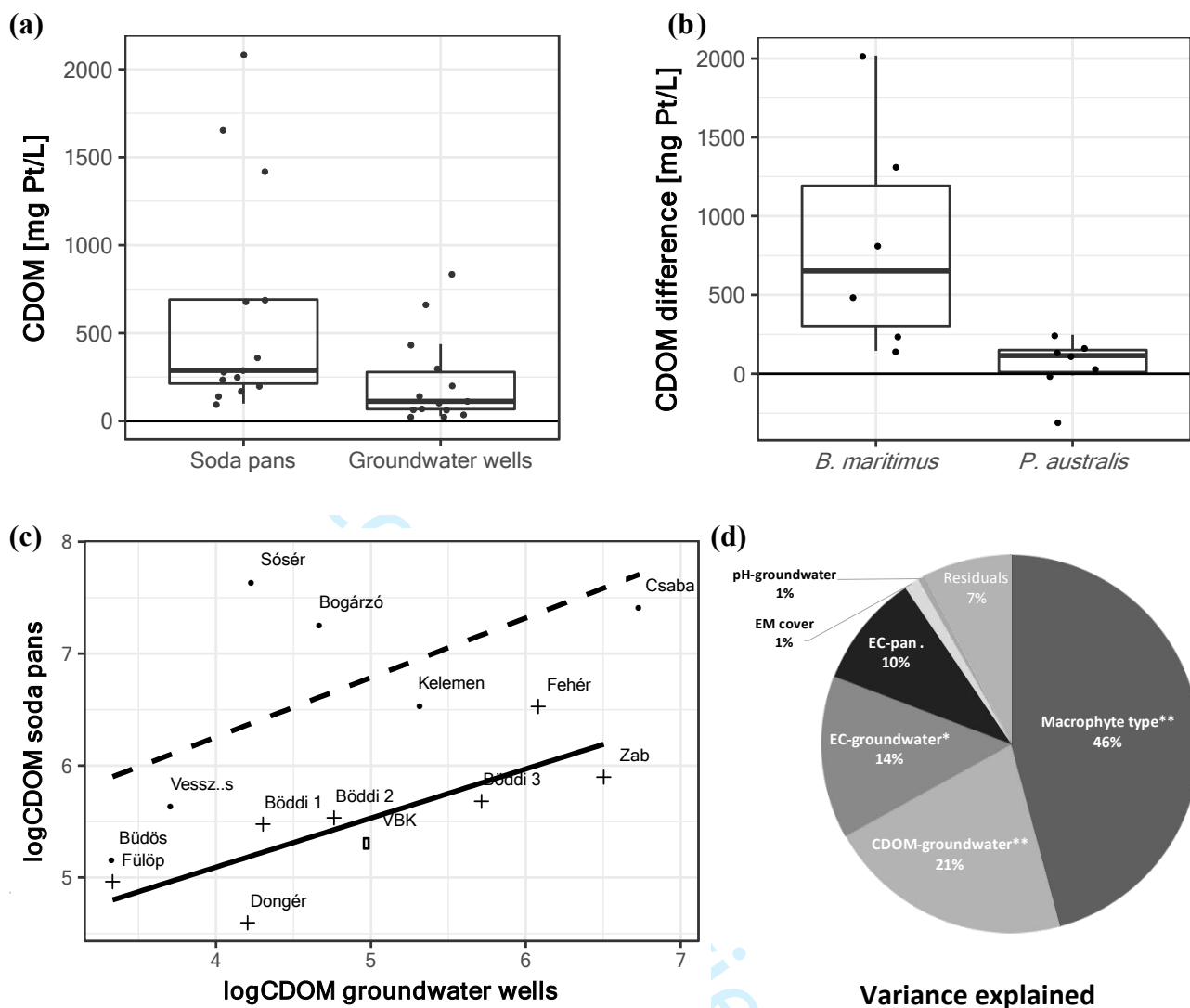
766

767

768

Review

769 Figure 3

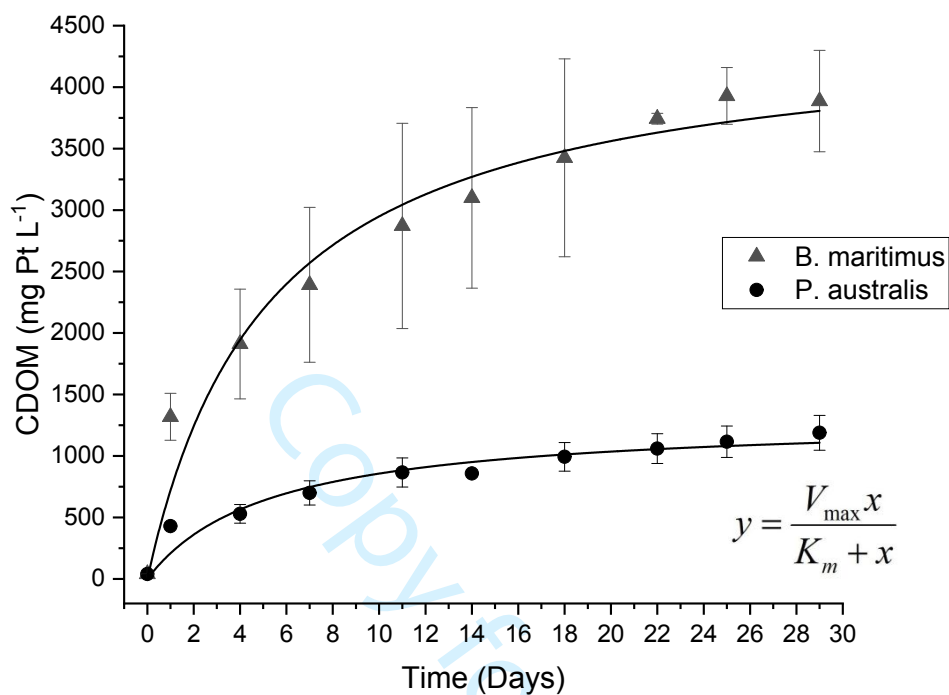


view

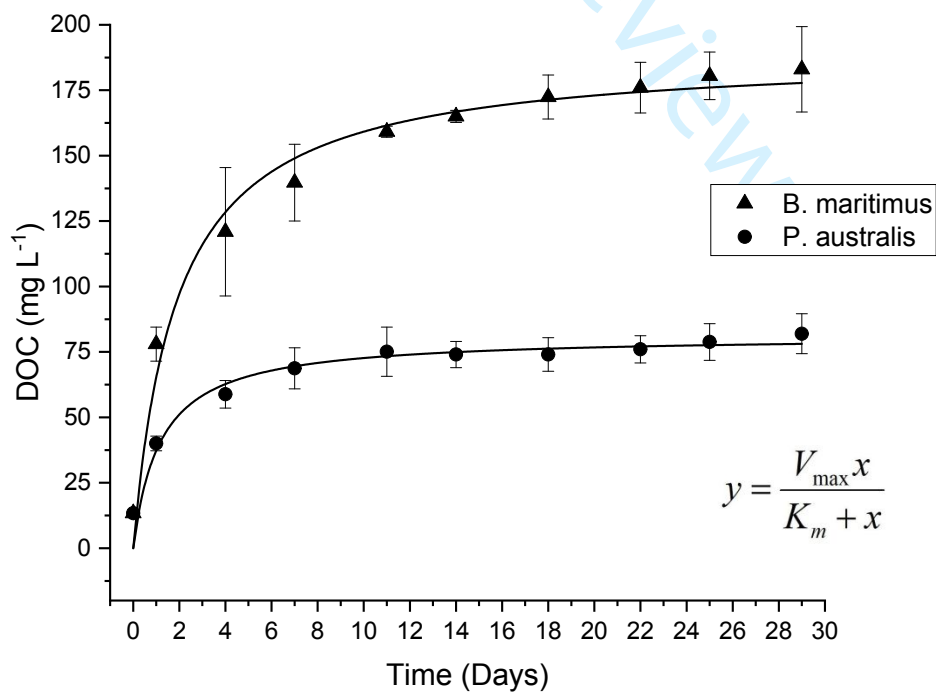
1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

793 Figure 4

794 (a)

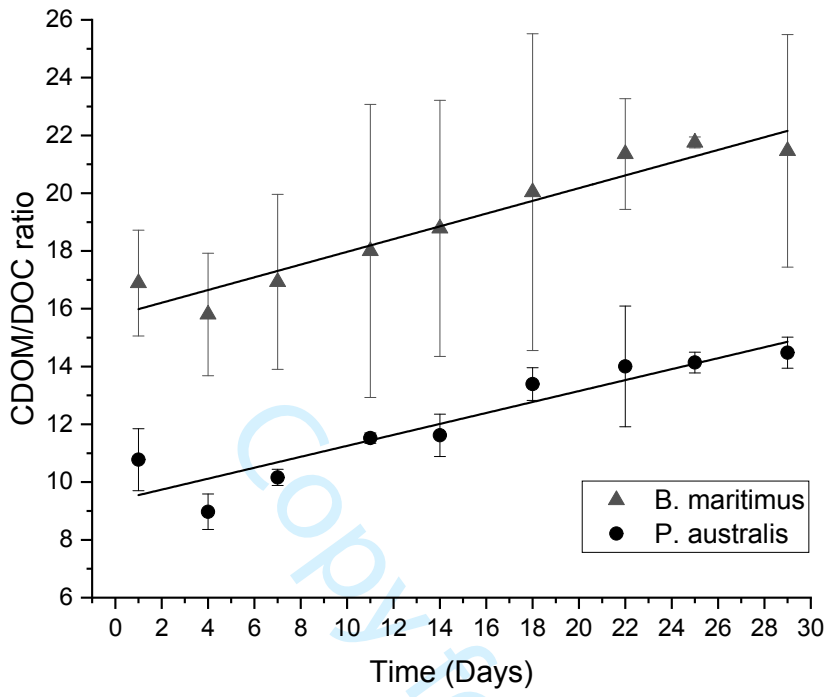


795
796
797 (b)



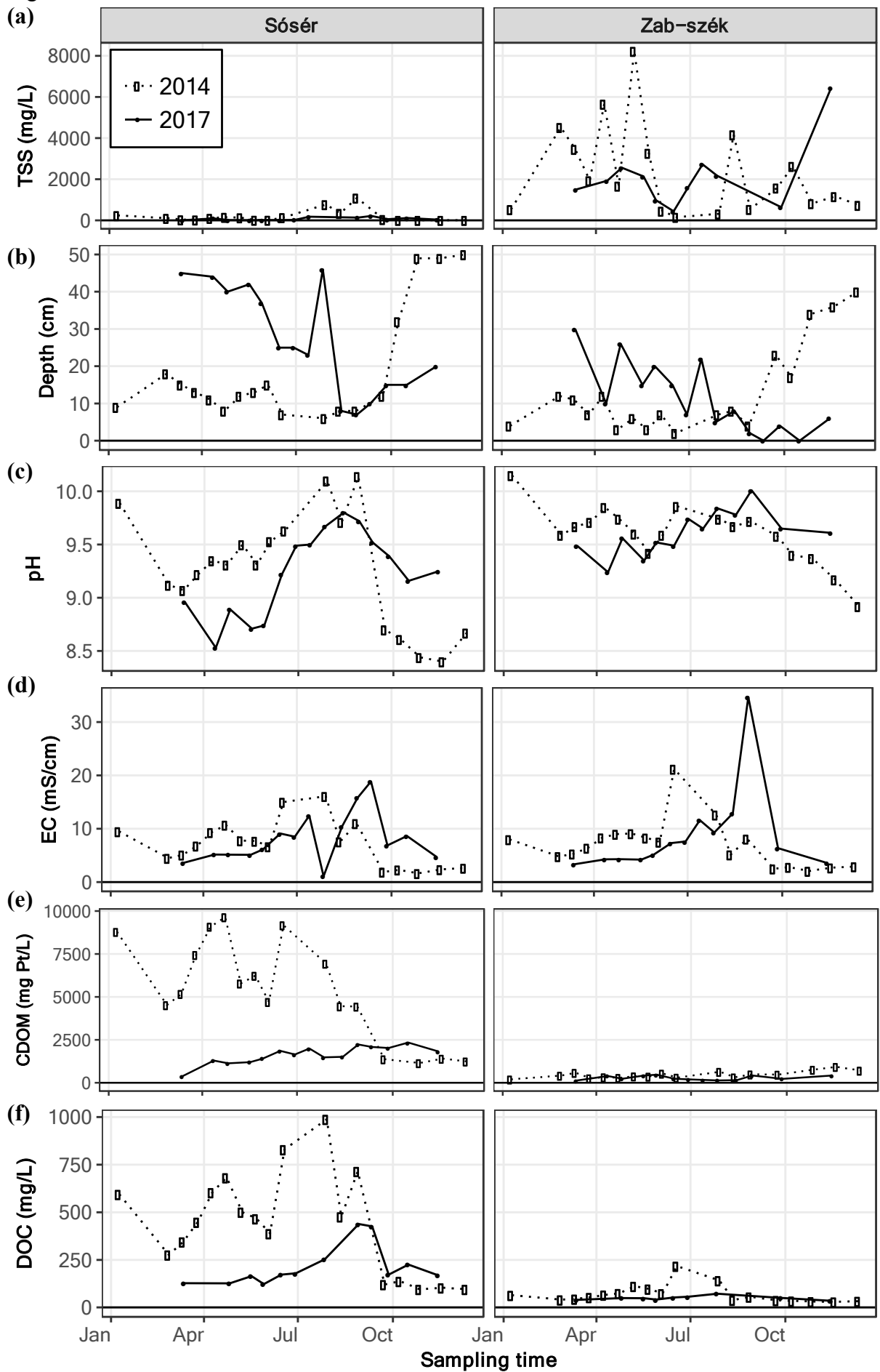
798
799

1
2
3 800
4 801 (c)
5
6
7



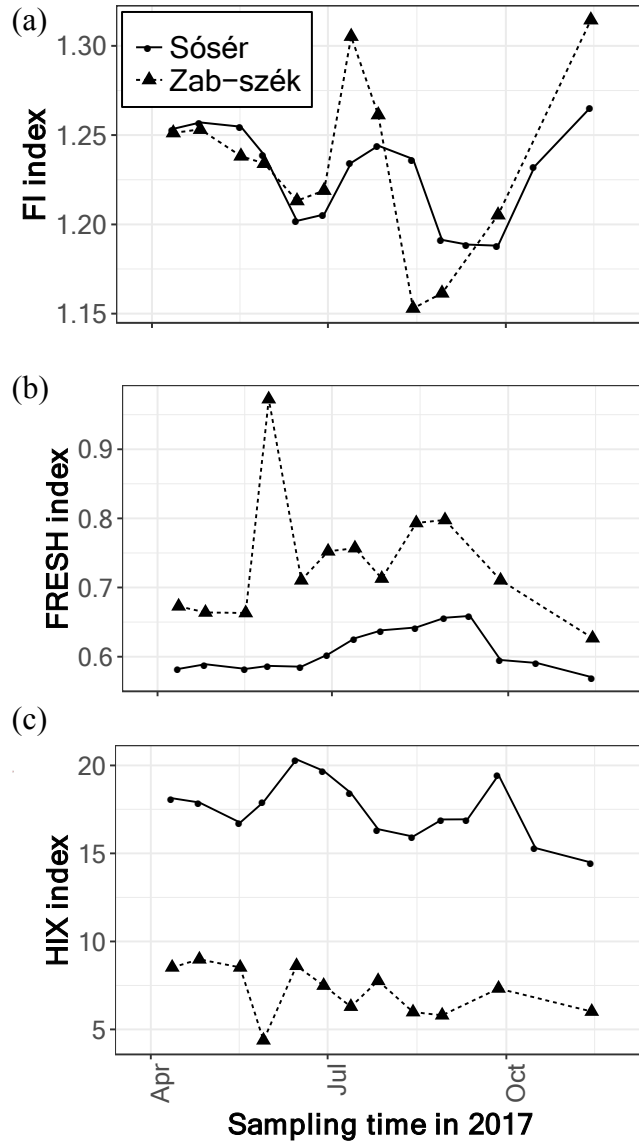
28
29
30 802
31 803
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

Figure 5



1
2
3 804
4 805
5 806
6 807
7 808
8 809
9 810
10 811
11 812
12 813
13 814
14 815
15 816
16 817
17 818
18 819
19 820
20 821
21 822
22 823
23 824
24 825
25 826
26 827
27 828
28 829
29 830
30 831
31 832
32 833
33 834
34 835
35 836
36 837
37 838
38 839
39 840
40 841
41 842
42 843
43 844
44 845
45 846
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

Figure 6



Review

Table S1. Coordinates of soda pans and groundwater wells studied I the cross-sectional analyses.

| Name of pan | Soda pans | | Groundwater wells | |
|--------------|-----------------|------------------|-------------------|------------------|
| | Latitude (N) | Longitude (E) | Latitude (N) | Longitude (E) |
| Bogárzó | 46.8054 | 19.1412 | 46.8045 | 19.1403 |
| Böddi-szék 1 | 46.7666 | 19.1500 | 46.7688 | 19.1500 |
| Böddi-szék 2 | 46.7603 | 19.1476 | 46.7604 | 19.1463 |
| Böddi-szék 3 | 46.7692 | 19.1229 | 46.7684 | 19.1246 |
| Büdös-szék | 46.5467 | 20.0298 | 46.5489 | 20.0259 |
| Csaba-szék | 46.8182 | 19.1888 | 46.8179 | 19.1875 |
| Dongér | 46.5725 | 20.0593 | 46.5725 | 20.0583 |
| Fehér-szék | 46.8083 | 19.1867 | 46.8106 | 19.1845 |
| Fülöp-szék | 46.5787 | 19.9877 | 46.5775 | 19.9889 |
| Kelemen-szék | 46.7973 | 19.1743 | 46.8012 | 19.1717 |
| Sósér | 46.7877 | 19.1350 | 46.7866 | 19.1355 |
| Unknown | 46.7636 | 19.1807 | 46.7642 | 19.1807 |
| Vesszős-szék | 46.5249 | 20.0373 | 46.5249 | 20.0310 |
| Zab-szék | 46.8342 | 19.1748 | 46.8352 | 19.1747 |

Table S2. Summary of backward selection of parameter for AIC minimalization

Initial model:

$$\log(\text{CDOM}_{\text{soda pans}}) \sim \log(\text{CDOM}_{\text{groundwater wells}}) + \text{Conductivity}_{\text{soda pans}} + \text{Conductivity}_{\text{groundwater wells}} + \text{pH}_{\text{soda pans}} + \text{pH}_{\text{groundwater wells}} + \text{Emerged macrophyte type} + \text{Emerged macrophyte cover} + \text{Depth} + \text{TSS} + \text{Chemical type} + \text{Optical type}$$

Model after backward selection:

$$\log(\text{CDOM}_{\text{soda pans}}) \sim \log(\text{CDOM}_{\text{groundwater wells}}) + \text{Conductivity}_{\text{soda pans}} + \text{Conductivity}_{\text{groundwater wells}} + \text{pH}_{\text{soda pans}} + \text{pH}_{\text{groundwater wells}} + \text{Emerged macrophyte type} + \text{Emerged macrophyte cover} + \text{TSS} + \text{Chemical type}$$

| Step | Removed parameter | Df | Deviance | Residuals Df | Residuals Deviance | AIC |
|------|-------------------|----|------------|--------------|--------------------|----------|
| 1 | | | | 1 | 0.2145496 | 32.49581 |
| 2 | Depth | 1 | 0.00065744 | 2 | 0.2152070 | 34.45297 |
| 3 | Optical type | 1 | 0.00077518 | 3 | 0.2159822 | 36.40263 |

Table S3. Summary of GVIF (Generalized Variance Inflation Factors) analyses before and after removal of inflated parameters ($GVIF^{1/(2 \cdot Df)} > 5$). Analyses were performed using *vif* function of *car* package in R.

Final model after backward selection:

$\log(CDOM_{\text{soda pans}}) \sim \log(CDOM_{\text{groundwater wells}}) + \text{Conductivity}_{\text{soda pans}} + \text{Conductivity}_{\text{groundwater wells}} + \text{pH}_{\text{soda pans}} + \text{pH}_{\text{groundwater wells}} + \text{Emerged macrophyte type} + \text{Emerged macrophyte cover} + \text{TSS} + \text{Chemical type}$

| | GVIF | Df | GVIF^{1/(2*Df)} |
|--|-------------|-----------|--------------------------------|
| $\log(CDOM_{\text{groundwater wells}})$ | 21.948539 | 1 | 4.684927 |
| $\text{Conductivity}_{\text{groundwater wells}}$ | 4.796074 | 1 | 2.189994 |
| $\text{Conductivity}_{\text{soda pans}}$ | 1.581812 | 1 | 1.257701 |
| $\text{pH}_{\text{groundwater wells}}$ | 20.822963 | 1 | 4.563218 |
| $\text{pH}_{\text{soda pans}}$ | 18.047921 | 1 | 4.248285 |
| Emerged macrophyte type | 20.992869 | 2 | 2.140513 |
| Emerged macrophyte cover | 21.216754 | 1 | 4.606165 |
| TSS | 5.424868 | 1 | 2.329135 |
| Chemical type | 24.508659 | 1 | 4.950622 |

Model after removal of highest GVIF parameter (Chemical type):

$\log(CDOM_{\text{soda pans}}) \sim \log(CDOM_{\text{groundwater wells}}) + \text{Conductivity}_{\text{soda pans}} + \text{Conductivity}_{\text{groundwater wells}} + \text{pH}_{\text{soda pans}} + \text{pH}_{\text{groundwater wells}} + \text{Emerged macrophyte type} + \text{Emerged macrophyte cover} + \text{TSS}$

| | GVIF | Df | GVIF^{1/(2*Df)} |
|--|-------------|-----------|--------------------------------|
| $\log(CDOM_{\text{groundwater wells}})$ | 4.707523 | 1 | 2.169683 |
| $\text{Conductivity}_{\text{groundwater wells}}$ | 2.288839 | 1 | 1.512891 |
| $\text{Conductivity}_{\text{soda pans}}$ | 1.551452 | 1 | 1.245573 |
| $\text{pH}_{\text{groundwater wells}}$ | 2.555346 | 1 | 1.598545 |
| $\text{pH}_{\text{soda pans}}$ | 5.345451 | 1 | 2.312023 |
| Emerged macrophyte type | 2.216132 | 2 | 1.22011 |
| Emerged macrophyte cover | 4.035098 | 1 | 2.008755 |
| TSS | 2.057321 | 1 | 1.434337 |

Final model after removal of highest GVIF parameter ($\text{pH}_{\text{soda pans}}$):

$\log(CDOM_{\text{soda pans}}) \sim \log(CDOM_{\text{groundwater wells}}) + \text{Conductivity}_{\text{soda pans}} + \text{Conductivity}_{\text{groundwater wells}} + \text{pH}_{\text{soda pans}} + \text{pH}_{\text{groundwater wells}} + \text{Emerged macrophyte type} + \text{Emerged macrophyte cover} + \text{TSS}$

| | GVIF | Df | GVIF^{1/(2*Df)} |
|--|-------------|-----------|--------------------------------|
| $\log(CDOM_{\text{groundwater wells}})$ | 1.740759 | 1 | 1.319378 |
| $\text{Conductivity}_{\text{groundwater wells}}$ | 1.41167 | 1 | 1.188137 |
| $\text{Conductivity}_{\text{soda pans}}$ | 1.337339 | 1 | 1.156434 |
| $\text{pH}_{\text{groundwater wells}}$ | 1.533761 | 1 | 1.238451 |
| Emerged macrophyte type | 1.615174 | 2 | 1.12734 |

| | | | |
|--------------------------|----------|---|----------|
| Emerged macrophyte cover | 2.270607 | 1 | 1.506853 |
| TSS | 1.955604 | 1 | 1.398429 |

Table S4. ANOVA table of the final linear model from table S3 ordered by variance explained. Significant correlations ($p < 0.05$) are in bold. Significance codes: $p < 0.001$: '***'; $p < 0.01$: '**'; $p < 0.05$: '*'; $p < 0.1$: '.'; $p > 0.1$: ''

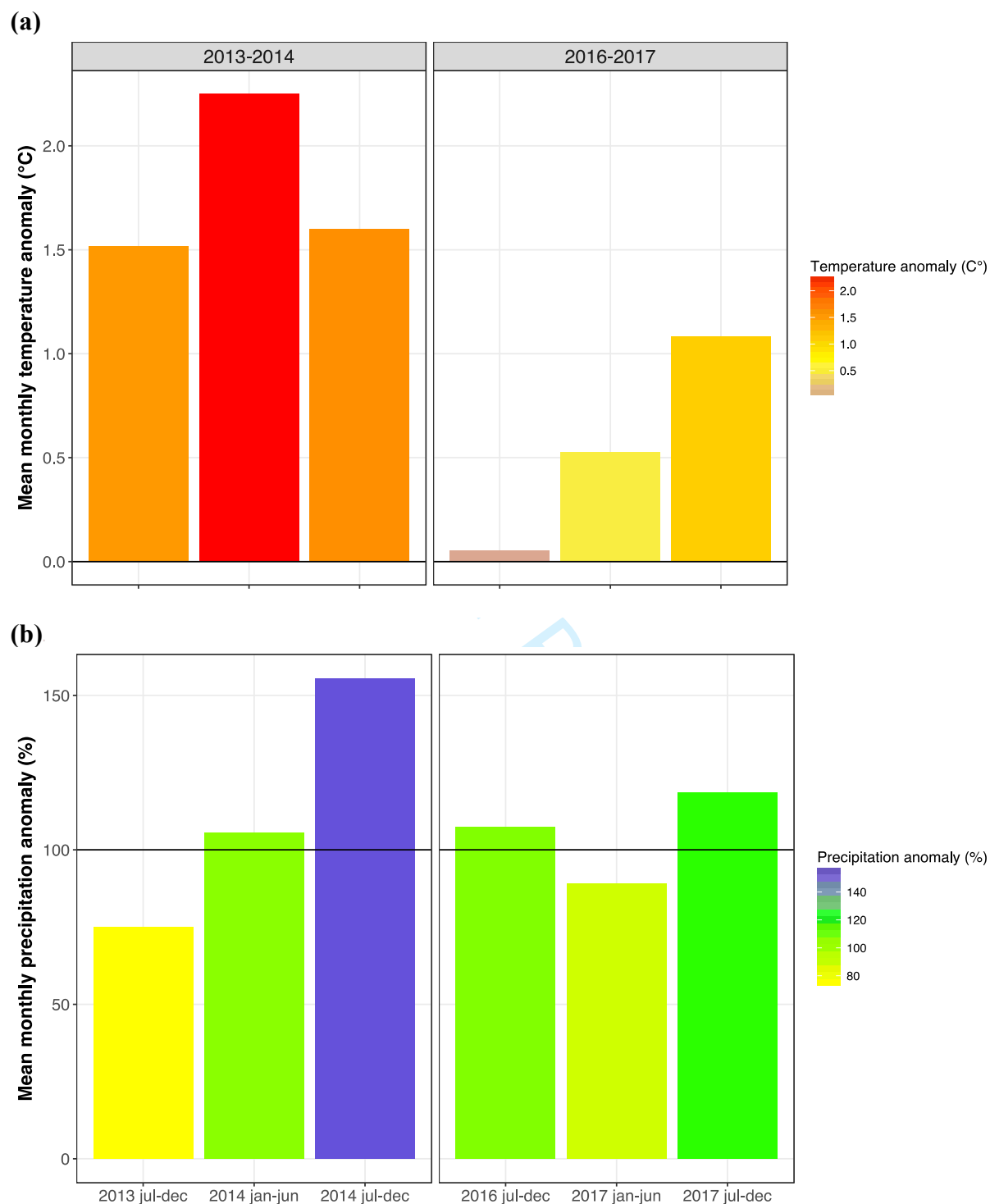
| | Df | Sum Sq | Mean Sq | F value | Pr (>F) | Variance explained (%) | Significance code |
|---|----|----------|----------|-----------|----------|------------------------|-------------------|
| Emerged macrophyte type | 2 | 5.409892 | 2.704946 | 15.099960 | 0.007605 | 45.783090 | ** |
| log(CDOM _{groundwater wells}) | 1 | 2.499060 | 2.499060 | 13.950630 | 0.013500 | 21.149160 | ** |
| Conductivity _{groundwater wells} | 1 | 1.646254 | 1.646254 | 9.189968 | 0.029032 | 13.932000 | * |
| Conductivity _{soda pans} | 1 | 1.152705 | 1.152705 | 6.434803 | 0.052103 | 9.755164 | . |
| Emerged macrophyte cover | 1 | 0.140855 | 0.140855 | 0.786304 | 0.415828 | 1.192038 | |
| pH _{groundwater wells} | 1 | 0.071898 | 0.071898 | 0.401362 | 0.554230 | 0.608466 | |
| TSS | 1 | 0.000009 | 0.000009 | 0.000051 | 0.994592 | 0.000077 | |
| Residuals | 5 | 0.895680 | 0.179136 | NA | NA | 7.580003 | |

Table S5. Results of cross-correlation analyses between CDOM and DOC concentrations and TSS, mean depth, pH, EC, and EEMS indexes (FI, FRESH and HIX) according to $x_{t+k} \sim y_t$, where x_{t+k} is CDOM or DOM at timepoint $t+k$, and y_t is the tested parameter at timepoint t , $k = 1$ corresponds to 1 month lag. Significant correlations ($p < 0.05$) are in bold. Only correlations between $k = -5$ and $k = 5$ are presented as no greater lag cross-correlations were significant.

| TSS | | | k | | | | | | | | | | |
|----------|----------|------|---------------|--------|--------|---------------|---------------|---------------|--------------|--------------|--------|--------|--------|
| Variable | Pan | Year | -5 | -4 | -3 | -2 | -1 | 0 | 1 | 2 | 3 | 4 | 5 |
| CDOM | Sósér | 2014 | -0.159 | 0.227 | 0.331 | 0.181 | 0.37 | 0.316 | -0.188 | -0.412 | -0.378 | -0.4 | -0.126 |
| CDOM | Sósér | 2017 | -0.376 | -0.566 | 0.049 | 0.265 | 0.314 | 0.598 | 0.407 | 0.323 | 0.108 | -0.218 | -0.296 |
| CDOM | Zab-szék | 2014 | 0.182 | -0.183 | -0.049 | -0.097 | -0.479 | -0.355 | -0.311 | -0.016 | -0.156 | -0.333 | 0.045 |
| DOC | Sósér | 2014 | -0.346 | -0.087 | 0.082 | 0.113 | 0.63 | 0.737 | -0.033 | -0.426 | -0.405 | -0.473 | -0.197 |
| DOC | Sósér | 2017 | -0.219 | -0.416 | -0.356 | 0.143 | 0.482 | 0.745 | 0.666 | -0.058 | -0.363 | -0.343 | -0.282 |
| DOC | Zab-szék | 2014 | -0.187 | -0.031 | -0.091 | -0.147 | -0.151 | -0.122 | 0.371 | 0.702 | 0.374 | 0.195 | 0.008 |
| Depth | | | | | | | | | | | | | |
| CDOM | Sósér | 2014 | 0.115 | -0.035 | -0.404 | -0.645 | -0.735 | -0.75 | -0.315 | -0.033 | 0.132 | 0.234 | 0.225 |
| CDOM | Sósér | 2017 | 0.499 | -0.001 | -0.006 | -0.438 | -0.603 | -0.829 | -0.466 | -0.248 | 0.137 | 0.194 | 0.321 |
| CDOM | Zab-szék | 2014 | 0.106 | 0.162 | 0.156 | 0.44 | 0.663 | 0.809 | 0.495 | 0.2 | -0.025 | -0.198 | -0.245 |
| DOC | Sósér | 2014 | 0.447 | 0.446 | 0.037 | -0.453 | -0.718 | -0.758 | -0.377 | -0.093 | 0.043 | 0.134 | 0.165 |
| DOC | Sósér | 2017 | 0.37 | 0.209 | -0.039 | -0.39 | -0.669 | -0.778 | -0.246 | -0.066 | 0.305 | 0.411 | 0.482 |
| DOC | Zab-szék | 2014 | 0.738 | 0.449 | 0.068 | -0.287 | -0.598 | -0.622 | -0.429 | -0.227 | -0.06 | 0.048 | 0.041 |
| pH | | | | | | | | | | | | | |
| CDOM | Sósér | 2014 | -0.212 | 0.072 | 0.277 | 0.425 | 0.626 | 0.778 | 0.267 | -0.024 | -0.113 | -0.376 | -0.22 |
| CDOM | Sósér | 2017 | -0.437 | -0.44 | -0.264 | 0.324 | 0.57 | 0.563 | 0.418 | 0.349 | 0.126 | -0.124 | -0.372 |
| CDOM | Zab-szék | 2014 | -0.133 | 0 | -0.116 | -0.34 | -0.564 | -0.792 | -0.358 | -0.165 | -0.001 | 0.065 | 0.014 |
| DOC | Sósér | 2014 | -0.403 | -0.35 | -0.145 | 0.124 | 0.646 | 0.93 | 0.406 | 0.035 | -0.113 | -0.371 | -0.237 |
| DOC | Sósér | 2017 | -0.226 | -0.324 | -0.304 | -0.066 | 0.444 | 0.839 | 0.726 | 0.254 | -0.431 | -0.598 | -0.365 |
| DOC | Zab-szék | 2014 | -0.664 | -0.275 | 0.111 | 0.339 | 0.441 | 0.414 | 0.096 | 0.009 | 0.103 | 0.099 | 0.201 |

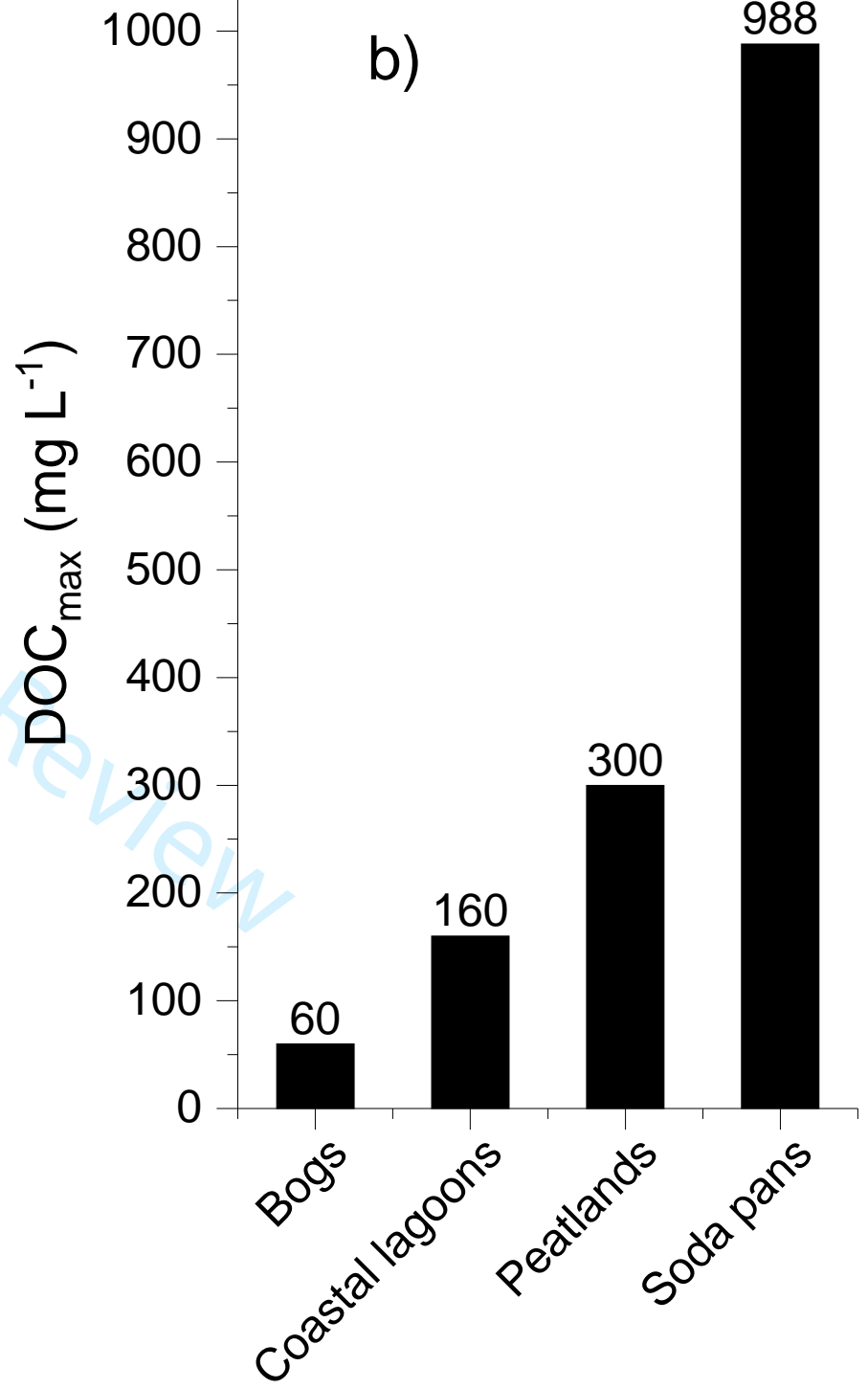
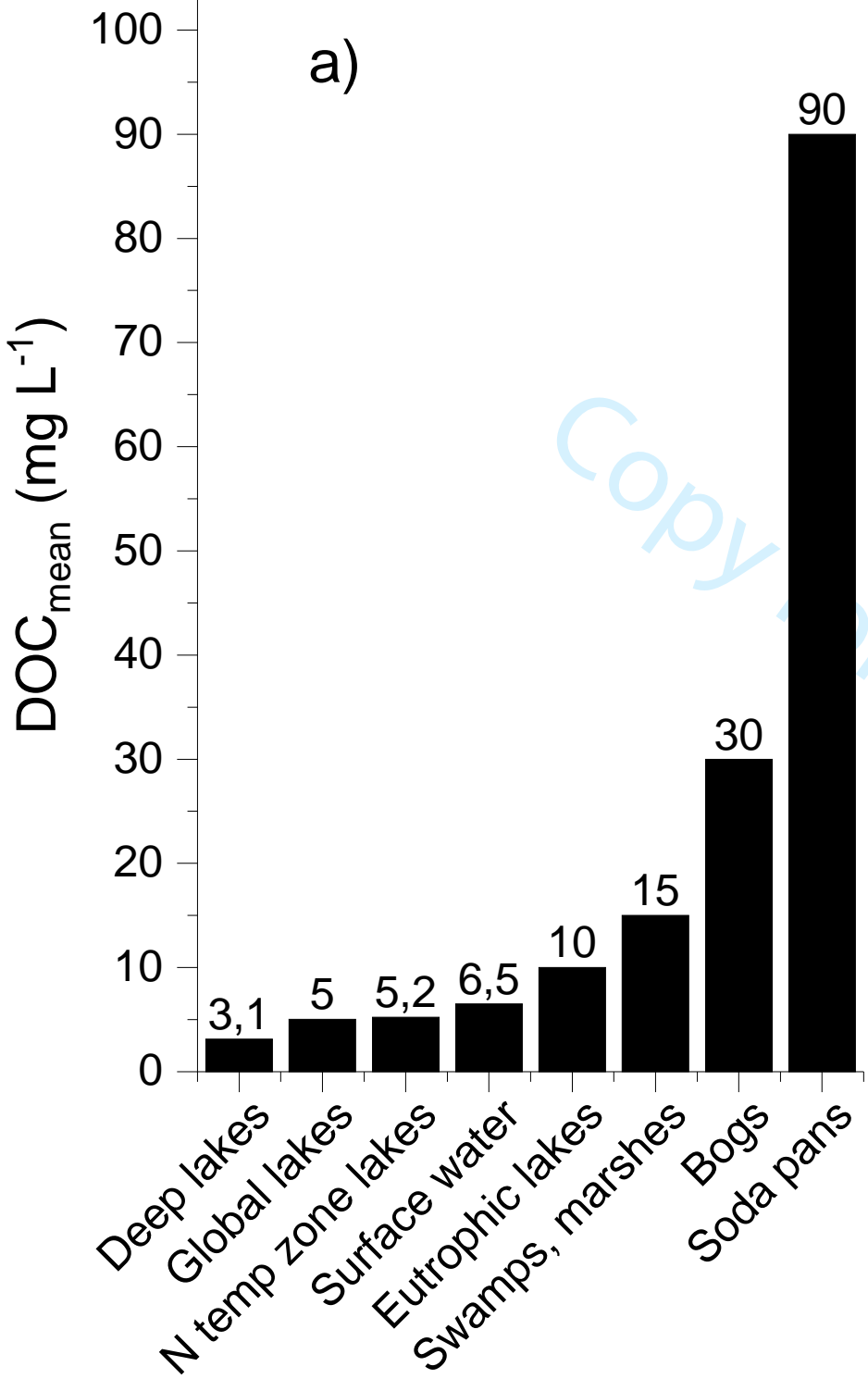
| EC | | | k | | | | | | | | | | |
|--------------|----------|------|--------|--------|---------------|--------|--------------|--------------|-------------|--------|--------|--------|--------|
| Variable | Pan | Year | -5 | -4 | -3 | -2 | -1 | 0 | 1 | 2 | 3 | 4 | 5 |
| CDOM | Sósér | 2014 | -0.209 | 0.017 | 0.394 | 0.36 | 0.535 | 0.802 | 0.304 | -0.051 | -0.171 | -0.432 | -0.41 |
| CDOM | Sósér | 2017 | -0.593 | -0.179 | -0.161 | 0.173 | 0.236 | 0.66 | 0.529 | 0.349 | -0.032 | -0.096 | -0.22 |
| CDOM | Zab-szék | 2014 | -0.179 | -0.108 | -0.246 | -0.369 | -0.384 | -0.447 | -0.377 | -0.338 | 0.061 | 0.444 | 0.502 |
| DOC | Sósér | 2014 | -0.39 | -0.337 | -0.053 | 0.015 | 0.469 | 0.994 | 0.542 | 0.059 | -0.066 | -0.389 | -0.464 |
| DOC | Sósér | 2017 | -0.322 | -0.334 | -0.448 | 0.027 | 0.627 | 0.87 | 0.326 | 0.104 | -0.199 | -0.291 | -0.42 |
| DOC | Zab-szék | 2014 | -0.515 | -0.563 | -0.349 | 0.093 | 0.673 | 0.951 | 0.62 | 0.125 | -0.164 | -0.401 | -0.344 |
| FI | | | | | | | | | | | | | |
| CDOM | Sósér | 2017 | 0.344 | 0.114 | 0.314 | -0.029 | -0.115 | -0.297 | -0.413 | -0.175 | -0.017 | -0.04 | 0.158 |
| CDOM | Zab-szék | 2017 | 0.143 | -0.598 | -0.143 | 0.296 | 0.028 | 0.207 | 0.303 | -0.211 | -0.344 | 0.188 | -0.104 |
| DOC | Sósér | 2017 | 0.176 | 0.235 | 0.578 | 0.132 | -0.538 | -0.568 | -0.175 | -0.197 | 0.163 | 0.32 | 0.181 |
| FRESH | | | | | | | | | | | | | |
| CDOM | Sósér | 2017 | -0.508 | -0.513 | -0.159 | 0.079 | 0.144 | 0.181 | 0.377 | 0.373 | 0.159 | -0.132 | -0.218 |
| CDOM | Zab-szék | 2017 | -0.092 | 0.208 | 0.207 | -0.169 | 0.081 | -0.02 | -0.468 | -0.225 | 0.387 | -0.077 | 0.071 |
| DOC | Sósér | 2017 | -0.164 | -0.396 | -0.617 | -0.275 | 0.365 | 0.89 | 0.624 | -0.027 | -0.387 | -0.356 | -0.059 |
| HIX | | | | | | | | | | | | | |
| CDOM | Sósér | 2017 | 0.079 | -0.123 | -0.529 | -0.389 | -0.575 | -0.225 | 0.038 | 0.103 | 0.21 | 0.349 | 0.16 |
| CDOM | Zab-szék | 2017 | 0.284 | 0.155 | -0.44 | 0.244 | 0.067 | -0.362 | 0.4 | 0.042 | -0.578 | 0.058 | 0.09 |
| DOC | Sósér | 2017 | 0.153 | -0.039 | -0.661 | -0.609 | -0.201 | -0.125 | 0.103 | 0.454 | 0.274 | 0.173 | -0.055 |

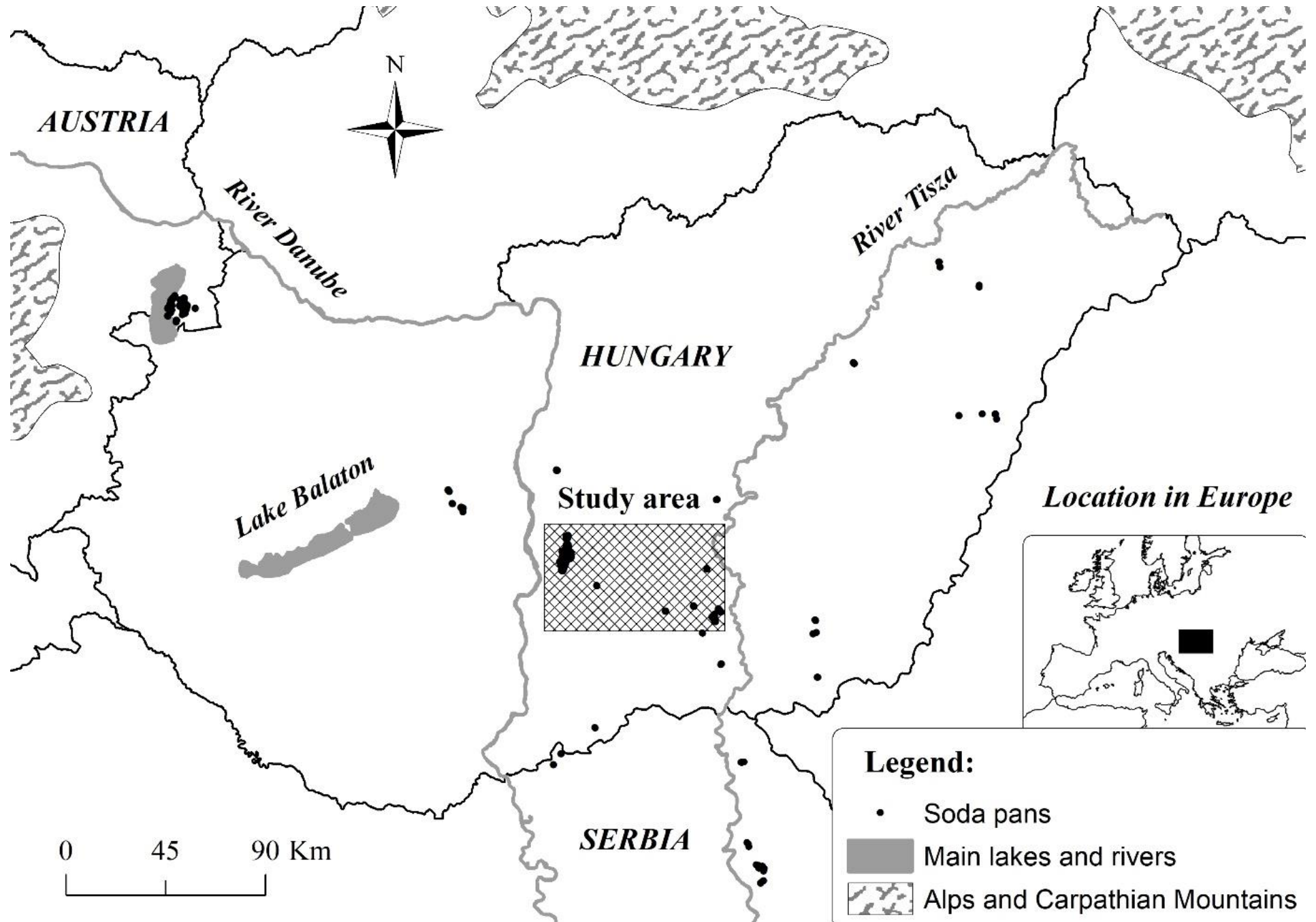
Fig. S1. Temperature and precipitation anomalies before and during the sampling periods of the seasonal analysis as six-months mean of difference from mean monthly average measured between 1981-2010. (a) Temperature and (b) precipitation. Based on data from Hungarian Meteorological Service (OMSZ).



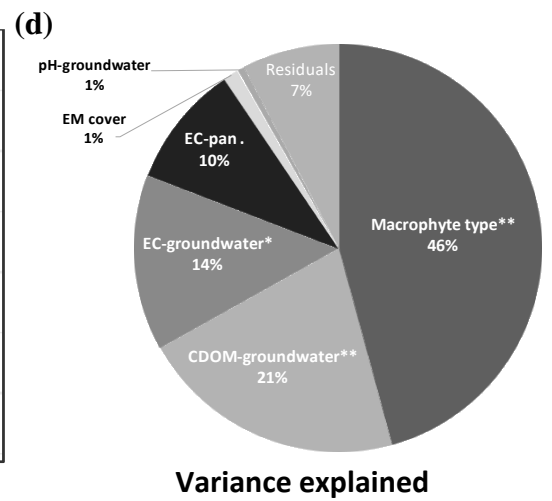
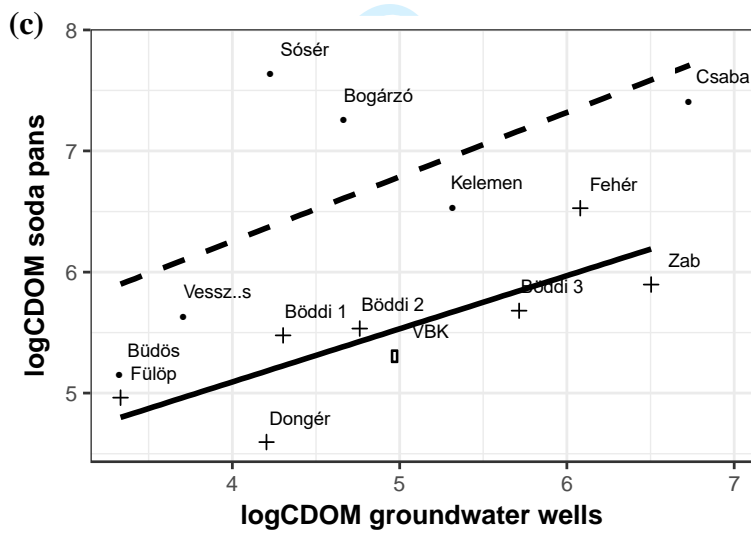
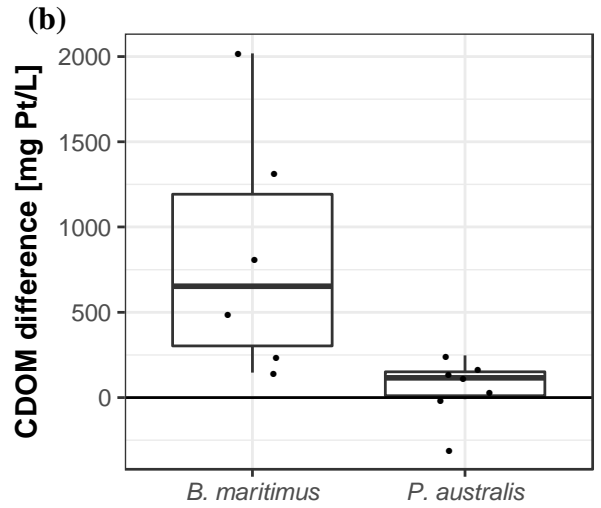
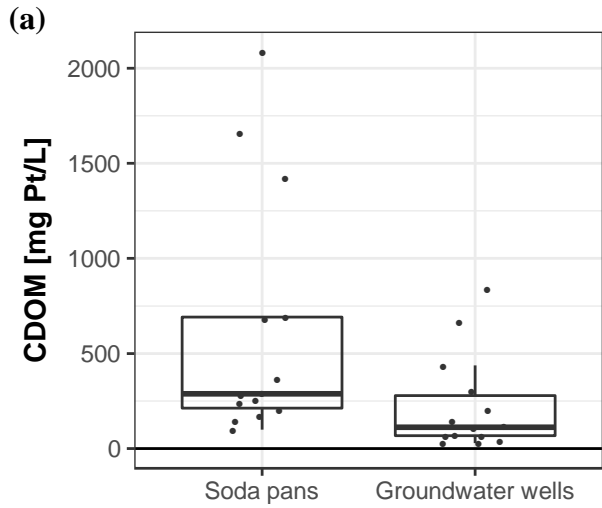
1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46





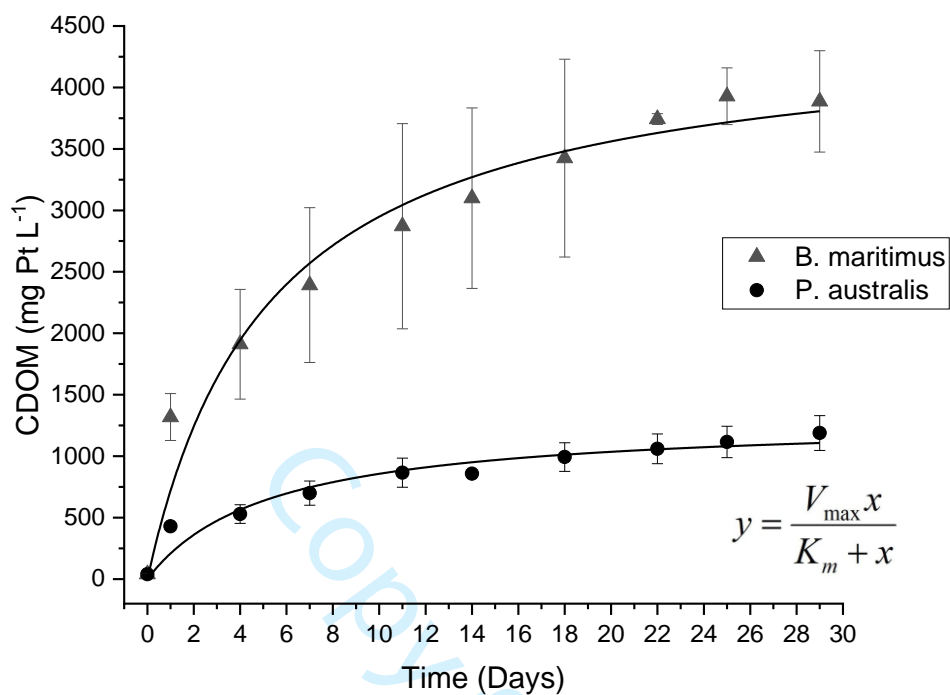
1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46



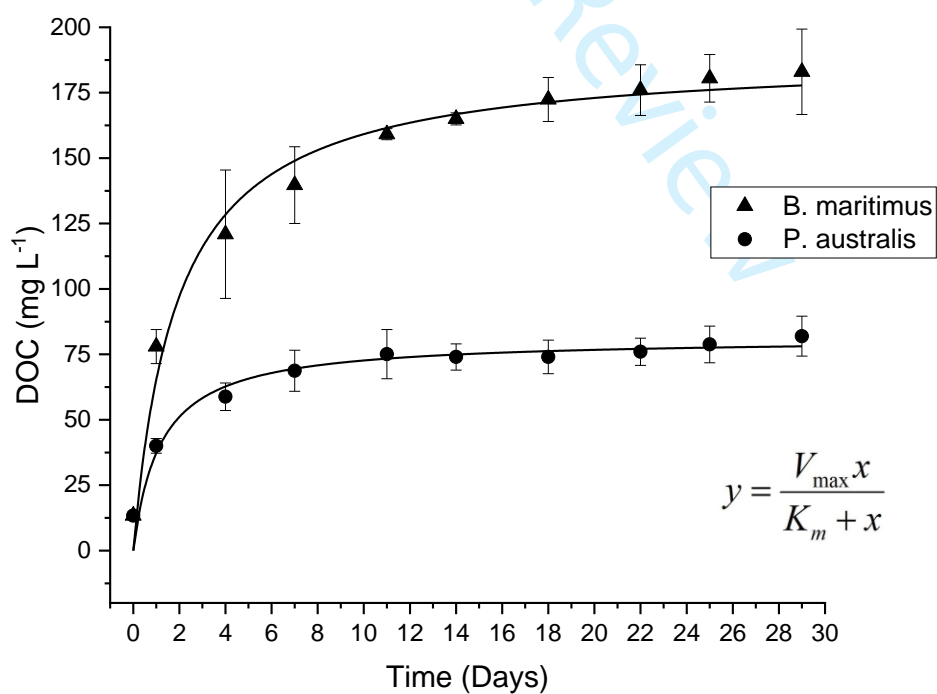
view

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

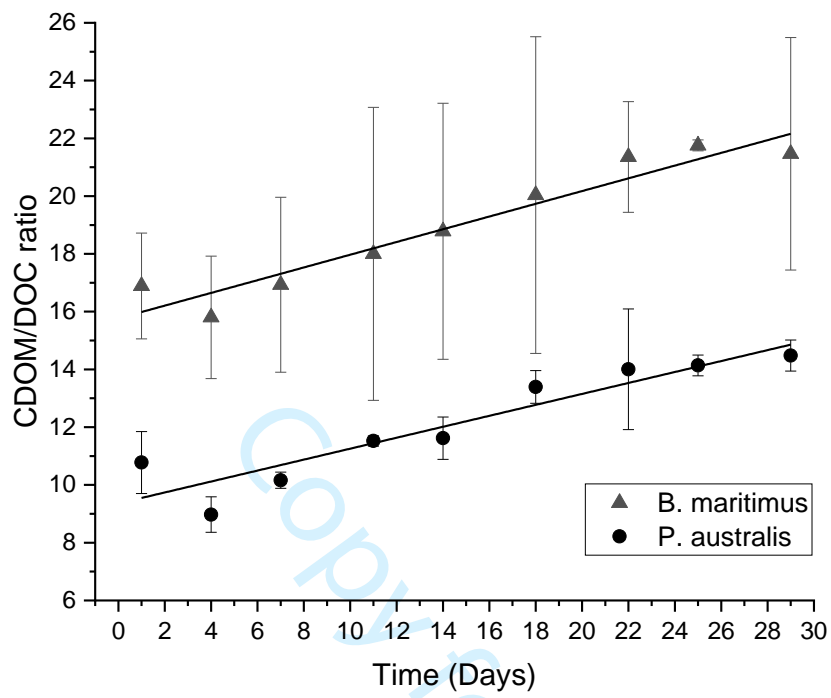
(a)

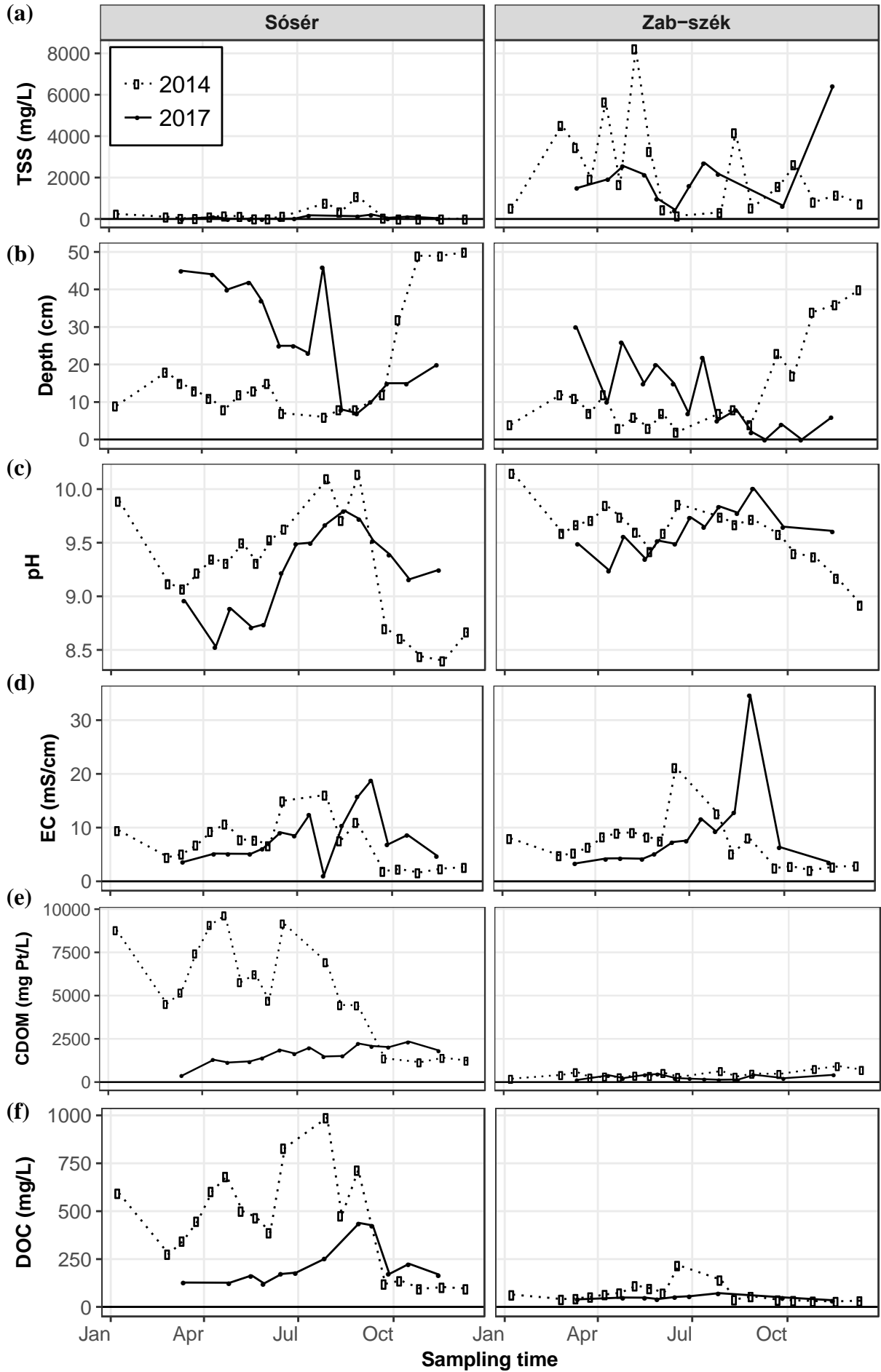


(b)

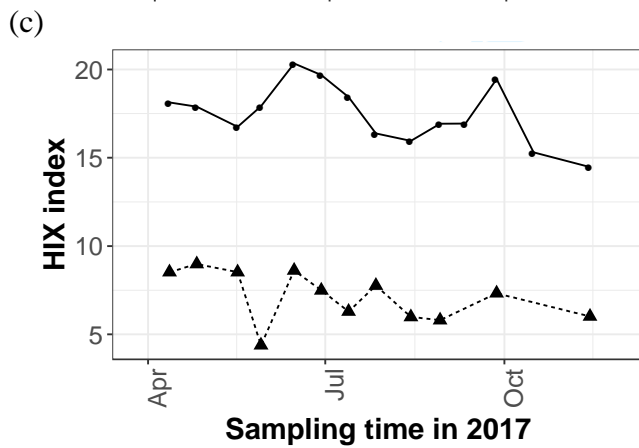
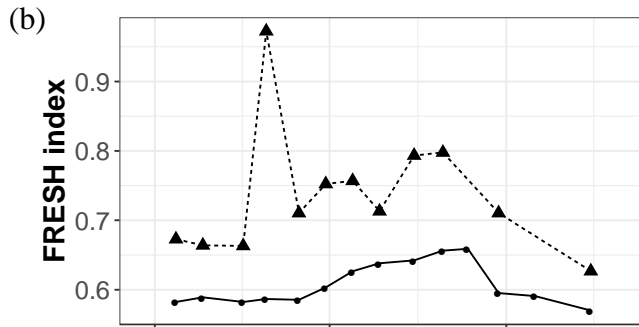
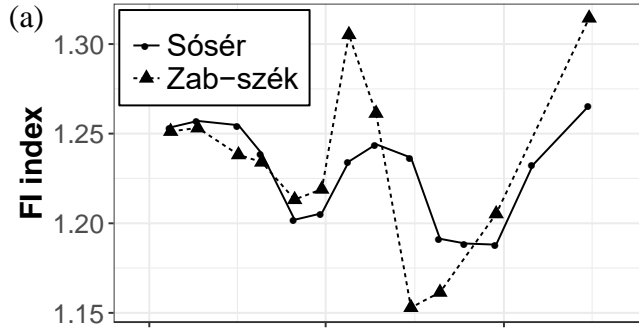


(c)





1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60



Review