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1 **Network topology and rainfall controls on the**
2 **variability of combined sewer overflows and loads**

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17 **Key Points:**

- 18 • A parsimonious stochastic model is developed for CSO flows and solute fluxes.
19 • Uncalibrated stochastic model agrees with calibrated SWMM model.
20 • Network structure and rainfall control CSO load variability.

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Abstract

Water and pollutant fluxes from combined sewer overflows (CSO) have a significant impact on receiving waters. The random nature of rainfall forcing dominates the variability of sewer discharges, pollutant loads, and concentrations. An analytical model developed here, shows how sewer network topology and rainfall properties variously impact the stochasticity of CSO functioning. Probability distributions of sewer discharge and concentration compare well with the results from a calibrated Storm Water Management Model in an application to a sewershed located in Dresden, Germany. The model is determined by only four parameters, three of which can be predicted a priori, two from the rainfall record and one from the network topology using geomorphological flow recession theory, while the fourth can be estimated from a short discharge time series. The sensitivity of CSO and wastewater treatment loads to network structure suggests simple topologies may be more vulnerable to poor performance. The analytical model is useful for evaluating various CSO management strategies to reduce adverse impacts on receiving waters in a probabilistic setting.

1 Introduction

With a preference for human settlement next to rivers globally (Fang et al., 2018), wastewater discharges from urban areas have significant impacts on the health of riverine ecosystems and other human settlements downstream. A large number of cities globally have combined stormwater-sanitary sewer systems which discharge only mechanically treated sewage to aquatic and marine ecosystems during heavy rainfall. While urban wastewater treatment plants (UWWTPs) can take the majority of sewerage when present, combined-sewer overflow (CSO) discharges, rich in nitrogen, phosphorous, heavy metals, antibiotics, hormones and other sanitary pollutants, can have significant environmental impacts (David, Borchardt, von Tmpling, & Krebs, 2013; Phillips et al., 2012). Impacts on ecosystems arise from chemical (i.e. oxygen depletion, non-ionized ammonia peaks), and physical (i.e. frequently increased bed shear) stresses which depend to a large degree on local conditions (Borchardt & Sperling, 1997). Predicting the variability of CSO loads, concentrations and the frequency of events are key to understanding their impacts and for working towards resilient and sustainable urban drainage systems.

The variability of CSO functioning is a crucial component of its design. Key design criteria include: dilution rates in relation to dry weather flow; storage capacity in

53 relation to design storms; an acceptable number of overflows per year; a maximum tol-
 54 erable pollution load; and a maximum CSO discharge, among others (Riechel et al., 2016).
 55 Accounting for the stochastic nature of rainfall is one of the key challenges in CSO treat-
 56 ment design (Geiger, 1998). On the other hand, the sewer network controls the travel
 57 time distribution and also influences the flows and therefore the distribution of loads (Lhomme,
 58 Bouvier, & Perrin, 2004). In the following, the hypothesis that both rainfall variability
 59 and the sewer network topology are significant controls on the statistical properties of
 60 CSO functions are elaborated.

61 1.1 Rainfall Controls

62 Rainfall variability is a dominant control of CSO event timing, event loads, and con-
 63 centration variability (Coutu, Giudice, Rossi, & Barry, 2012; Geiger, 1998; Sandoval, Tor-
 64 res, Pawlowsky-Reusing, Riechel, & Caradot, 2013). Short intense rainfall can promote
 65 elevated loads in first flush events (Krebs, Holzer, Huisman, & Rauch, 1999). Long du-
 66 ration, low intensity events can lead to poorer efficiency at an UWWTP, reducing the
 67 relative contribution of CSOs to river pollution (Phillips et al., 2012). Rainfall event in-
 68 tensities correlate with CSO water quantity and pollutant loads, while event duration
 69 and rain depth predict CSO pollutant concentrations (Sandoval et al., 2013). Under the
 70 changing climatic conditions, the frequency of intense rainfall may increase, which brings
 71 concerns about an increasing frequency of CSO events (Semadeni-Davies, Hernebring,
 72 Svensson, & Gustafsson, 2008; Sterk, de Man, Schijven, de Nijs, & de Roda Husman, 2016).
 73 These aspects of CSO performance are suited to treatment as a stochastic process, specifically
 74 accounting for the statistical properties of the timing and magnitude of rainfall events
 75 on the hydrological response (Botter, Porporato, Rodriguez-Iturbe, & Rinaldo, 2009).

76 1.2 Network Topology and Discharge Variability

77 Taking a nonlinear relationship between storage and runoff, Q , the continuity equa-
 78 tion can be stated as (Botter et al., 2009):

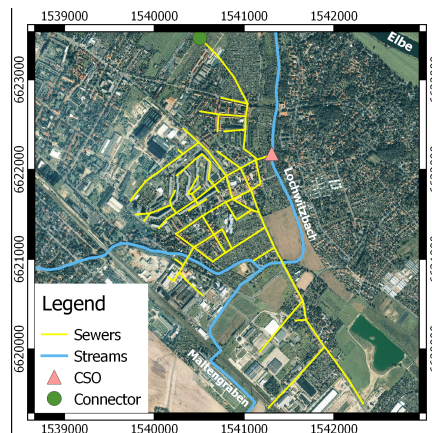
$$79 \frac{dQ}{dt} = -kQ^\alpha + \xi(t) \quad (1)$$

80 where k is related to the hydraulic residence time, $0 < \alpha$ is a flow recession exponent.
 81 The rainfall, ξ , is assumed to follow a marked Poisson process with exponentially dis-
 82 tributed times between events and event depths. From Eq. (1) the probability density

83 function (pdf) for long-term temporal variability of Q was previously derived and the
 84 shape of the pdf was shown to be strongly controlled by α (Botter et al., 2009).

85 The topological properties of river networks have also been shown to be related to
 86 α (Biswal & Marani, 2014). Through a decomposition of a river network into so-called
 87 independent links, a power-law relationship between the number of independent links
 88 $N(l)$ and the total lengths of those same links, $G(l)$, at a distance l was derived: i.e. $N(l) \propto$
 89 $G(l)^\alpha$ (Biswal & Marani, 2010). In rivers, at least, there appears to be an intrinsic re-
 90 lationship between the network structure, the hydrological response and the variability
 91 of discharge.

92 Sewers share many topological characteristics with rivers (Yang et al., 2017). Like
 93 rivers, sewers follow power laws in the area-distance relationship (Hack's Law) and in
 94 the probability distribution of contributing area, with exponent values similar to those
 95 found in rivers (Yang et al., 2017). A topological model also predicts runoff character-
 96 istics from sewers, as in rivers (Lhomme et al., 2004). We therefore hypothesize that the
 97 topological properties of gravity-driven sewer networks will influence the pdf of discharges,
 98 as well as pollutant loads and concentrations.



99 **Figure 1.** The Lockwitzbach sewer network and CSO. Coordinates are UTM Zone 33 North.

100 1.3 A Utilitarian Perspective

101 Clearly the structure of the sewer network and rainfall properties are important
 102 factors, together with regulations and/or guidelines, impacting upon CSO design and

103 function. The manager of a sewer system might wonder what the use is to predict vari-
104 ability of a CSO system given that the rainfall properties cannot be controlled or that
105 only small changes to the structure of a sewer system can be changed at any one time.
106 Firstly, in response, many parts of the world face the challenge of constructing sewer sys-
107 tems to keep pace with rapid urbanization (Xu et al., 2019). As such there is a need for
108 general design tools to plan future urban infrastructure as distinct from comprehensive
109 hydrodynamic models solving the mass and energy balance equations for water and so-
110 lute transport. Secondly managers of established systems more and more need to be aware
111 of climate change impacts and to have a whole-of-catchment approach to managing sewer
112 performance. This necessitates a systems-scale understanding of the transformation of
113 rainfall variability into the variability of runoff production and sewer functioning.

114 Treating runoff as a stochastic process, has led to recent insights into how urban-
115 ization is changing the statistical properties of runoff as well as the variability of urban
116 wash-off (Daly, Bach, & Deletic, 2014; Mejía, Daly, Rossel, Jovanovic, & Gironás, 2014).
117 A stochastic approach was recently developed to evaluate the variability of water stor-
118 age within, and discharges from a CSO tank (Wang & Guo, 2018). The process descrip-
119 tions of storage and discharge used by Wang and Guo (2018) are identical to those used
120 to previously examine soil water storage (McGrath, Hinz, & Sivapalan, 2007; Milly, 1993)
121 and the temporal clustering of threshold flow events (Aquino, Aubeneau, McGrath, Bol-
122 ster, & Rao, 2017; Laio, Porporato, Ridolfi, & Rodriguez-Iturbe, 2001; McGrath et al.,
123 2007). Furthermore there have been recent advances in understanding how the network
124 structure of rivers influences the hydrodynamics of discharge (Biswal & Marani, 2010).
125 As a result, there is an opportunity to draw upon these new ideas in hydrology and ap-
126 ply them to improve the theoretical underpinnings of the practice of CSO management.

127 In this contribution we develop analytical expressions for the pdfs of CSO discharges,
128 loads and concentrations with parameters derived from rainfall and the structure of the
129 sewer network. The pdfs compare favourably with the results of a calibrated Storm Wa-
130 ter Management Model (SWMM). The model developed here allows a sewer system man-
131 ager/designer to easily assess how changing rainfall patterns (e.g. climate change sce-
132 narios) or urban growth (e.g., expansion and redesign of the sewer network) would im-
133 pact CSO functioning and the risks to urban rivers.

134 2 Stochastic Analytical CSO Network Yield (saCSOny) Model

135 2.1 Discharges, Concentrations and Loads

136 The combined flows (and loads), Q_c (L_c) at a CSO diversion are given by the sum
 137 of the sanitary flow (load), Q_s (L_s), and urban stormwater flow (load), Q_u (L_u):

$$138 \quad Q_c = Q_u + Q_s \quad (2)$$

$$139 \quad L_c = C_c Q_c = L_u + L_s = C_u Q_u + C_s Q_s \quad (3)$$

140 where Q_u is the stochastic stormwater runoff, Q_s is the sanitary discharge, and C_u is the
 141 solute concentration in stormwater, assumed to be constant and much smaller than the
 142 steady concentration in the sanitary flow, C_s . Implicitly we assume $L_u \ll L_s$ and that
 143 the above terms represent system averages and thus describe well-mixed conditions at
 144 the catchment-scale. Sanitary flows typically display strong diurnal and weekly variabil-
 145 ity while stormwater flows vary significantly at sub-hourly time scales during rainfall events.

146 While Q_s and C_s are initially assumed constant this assumption is later relaxed, such
 147 that fluctuations in the sanitary fluxes can be taken into account. The difference be-
 148 tween Q_c and a threshold discharge, Q_t , at a CSO diversion, determines the CSO dis-
 149 charge, $Q_{CSO} = Q_c - Q_t$, and the load during a CSO event, $L_{CSO} = C_c Q_{CSO}$. The
 150 overflow structure is typically a weir and when the water level in the upstream pipe reaches
 151 a certain height, the weir overflows into the CSO pipes. These structures are constructed
 152 such that the flow directed towards the wastewater treatment plant depends on the up-
 153 stream flow rate only to a minor extent. A simple threshold is therefore a good approx-
 154 imation to the hydrodynamics. The WWTP receives a flow, $Q_{WWTP} = Q_c - Q_{CSO}$,
 155 and a load, $L_{WWTP} = C_c Q_{WWTP}$. A stochastic model for Q_u is described next from
 156 which pdfs for the flows, loads and concentrations are derived.

157 2.2 Stormwater pdfs

158 Starting with Eq. (1), Botter et al. (2009) previously derived the pdf of discharges
 159 for rivers (the term Q_u here). In relation to Eq. (1), when $\alpha = 1$, the storage - discharge
 160 relationship is described as a linear reservoir, such that flows decrease exponentially with
 161 time during the recession phase. The pdf of Q_u in this case is given by Eq. (A.1). Flow
 162 recession in rivers, however, are often better described by power-laws (Wittenberg, 1999).
 163 When $0 < \alpha < 1$ the nonlinearity is termed concave; when $1 < \alpha < 2$, a range often
 164 observed in rivers, the nonlinearity is termed convex and finally, when $\alpha > 2$ the re-

165 lation is termed hyperbolic. For concave recession the pdf is given by Eq. (A.2) (Bot-
 166 ter et al., 2009). The pdfs of the convex and hyperbolic models have the same form as
 167 Eq. (A.2) without the Dirac delta term. For the variables of interest (Q_c , Q_{CSO} , Q_{WWTP} ,
 168 C_c , L_{CSO} , and L_{WWTP}) we can apply a change of variables to derive their pdfs from
 169 the pdfs for Q_u (see Appendix A).

170 2.3 Accounting for Sanitary Discharge Variability

171 To take into account the diurnal variation in sanitary flows (Q_s) and concentra-
 172 tions (C_s), they can be treated as random variables, independent of Q_u and C_u . Using
 173 the marginal distribution rule, the pdf of Q_c is related to the marginal distribution of
 174 Q_c , given Q_s and the pdf of Q_s i.e.:

$$175 \quad p_{q_c}(Q_c) = \int_0^{\infty} p_{q_c}(Q_c|Q_s) p_{q_s}(Q_s) dQ_s \quad (4)$$

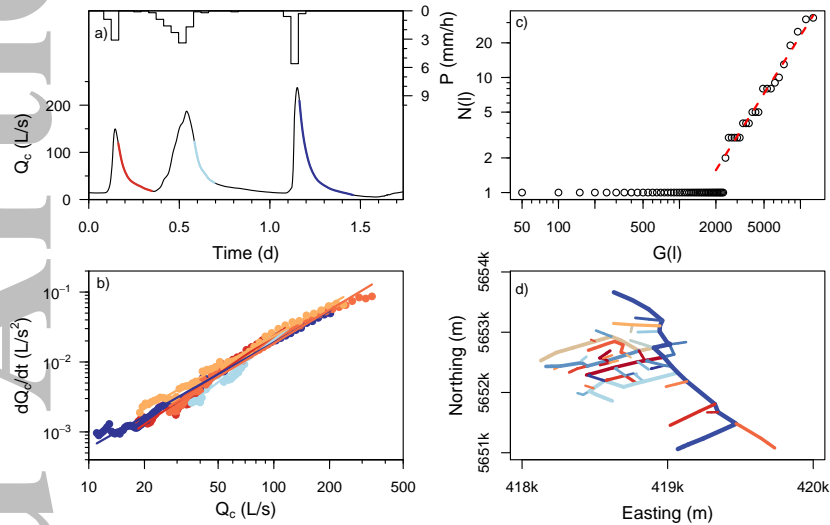
176 This is effectively a weighted average of the pdf of combined flows (Eq. A.4), where the
 177 weights are determined from the distribution of sanitary flows (p_{q_s}). A short period of
 178 observed dry-weather flows suffice to estimate p_{q_s} . The pdfs of discharges to the WWTP
 179 and from the CSO can be rescaled similarly. To derive the pdf of C_c the same approach
 180 can be used together with the distribution of sanitary loads, p_{L_s} , or alternatively the joint
 181 distribution of Q_s and C_s using the marginal distribution of C_c (Eq. A.10), and the joint
 182 pdf of Q_s and C_s i.e.:

$$183 \quad p_{c_c}(C_c) = \int_0^{\infty} \int_0^{\infty} p_{c_c}(C_c|Q_s, C_s) p_{q_s c_s}(Q_s, C_s) dQ_s dC_s \quad (5)$$

184 Practically this is achieved by sampling a short time series of dry-weather flows $Q_s(t_i)$
 185 and $C_s(t_i)$ at corresponding times then averaging the resulting ensemble of $p_{C_c}(C_c|Q_s(t_i), C_s(t_i))$
 186 over the set of samples, at each concentration, C_c , then normalizing the result to obtain
 187 a pdf. The complete set of equations are presented in Appendix A.

188 Next, the above model (defined by Eq 3-5 and A.1 - A.13, which we refer to as saC-
 189 SOny) is applied to a sewershed located in Dresden, Germany. The software R was used
 190 for data analysis (R Core Team, 2018). All code used in this paper is documented in Sup-
 191plementary Material) and the SWMM input and output needed to run the scripts can
 192 be found at <https://doi.org/10.26182/5bbbff6fadf94>. The R code includes scripts
 193 to numerically determine the pdfs and their corresponding cumulative distribution func-
 194 tions, analyse rainfall time series to determine the rainfall parameters, analyse SWMM

195 input files to calculate α from the network properties, analyse a discharge time series to
 196 conduct flow recession analysis, and reproduce all figures in the main text and supple-
 197 mentary material.

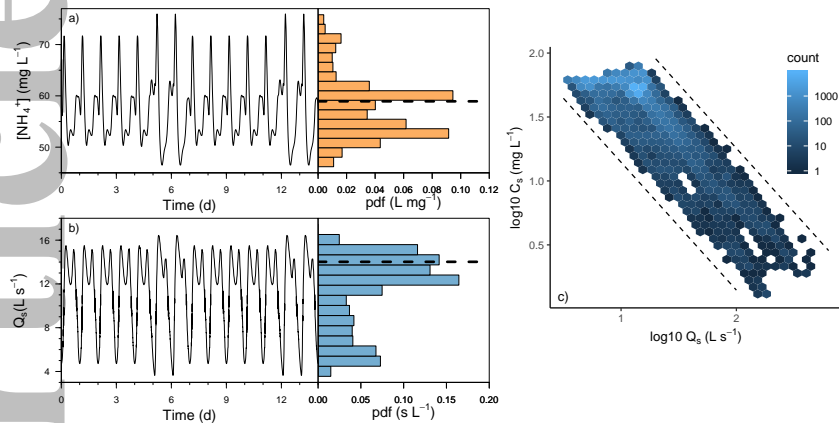


198 **Figure 2.** Empirical flow recession analysis: a) discharge, Q_c , and rainfall, P , time series for
 199 three of the five events shown in (b); b) Linear regression of the logarithms of the rate of change
 200 in discharge, and mean discharge, i.e. $\log(-dQ_c/dt) = \log(k) + \alpha \log(Q_c)$ where the mean $\alpha = 1.7$
 201 (Table S3); c) The power law relation found between length and number of independent links i.e.
 202 $G(l) \propto N(l)^{1.7}$; and d) The associated decomposed sewer network of independent links (color
 203 coded) (Biswal & Marani, 2014).

207 3 Application

208 3.1 The Lockwitzbach Sewershed

209 The Lockwitzbach sewer network, located in Dresden, Germany has a mean an-
 210 nual rainfall of 665 mm a⁻¹ (1981-2010), a potential evaporation rate of 605 mm a⁻¹ and
 211 a mean annual temperature of 9.4°C (Deutscher Wetterdienst, 2017). The sewershed has
 212 an area of 144.3 ha, with 36 ha of connected impervious surface. Wastewater from ap-
 213 proximately 7,630 inhabitants and stormwater from primarily suburban land use is col-
 214 lected by 12.83 km of pipes. Extraneous water does not impact upon this sewer network
 215 Karpf and Krebs (2011). The CSO structure operates as a sideflow weir with a flow thresh-



204 **Figure 3.** Characteristics of the sewer dynamics with: (a) Dry period discharges, Q_s (and pdf
 205 p_{q_s}); (b) concentrations, C_s ($[NH_4^+]$) (and pdf p_{c_s}); and (c) the $C_c - Q_c$ relationship, bounded by
 206 $C_c \propto Q_c^{-1}$.

216 old of approximately 600 L s^{-1} . Excess water is discharged into the Lockwitbach, an ur-
 217 ban stream that drains into the Elbe River. A gate prevents backflow from the stream
 218 or downstream pipes. The northern outlet of the sewershed provides a connection for
 219 transport to the central Dresden WWTP.

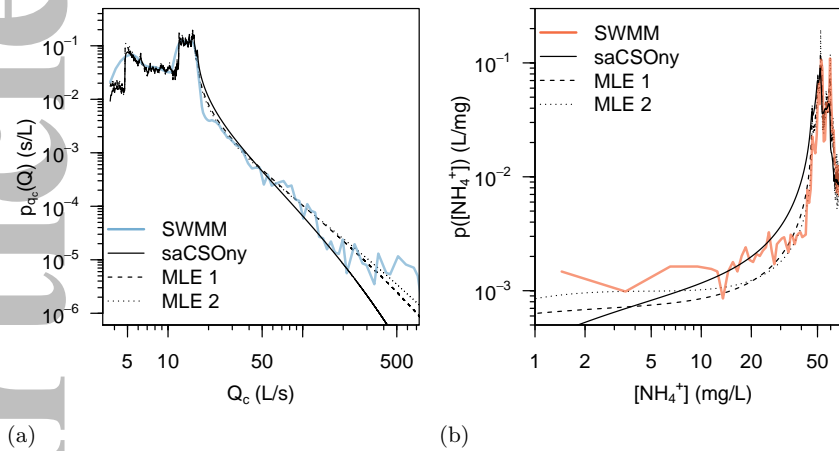
227 A monitoring program of the joint Urban Observatory Dresden of Dresden Uni-
 228 versity and the Helmholtz Centre for Environmental Research-UFZ under the Terres-
 229 trial Environmental Observation Initiative (TERENO) was established with the aims to
 230 analyze transport processes in sewer networks and the impacts of urban water manage-
 231 ment on river quality (Helm et al., 2015; Wollschlger et al., 2016). For hydraulic and wa-
 232 ter quality simulations the open source software EPA-SWMM v. 5.1.011 was previously
 233 calibrated to this data (Deb, Pratap, Agarwal, & Meyerivan, 2002; Kaeseberg et al., 2018;
 234 Rossmann, 2010; Steinberg, 2015). The calibrated model was run with a time step of ten
 235 minutes, using rainfall at a similar temporal resolution. A 17 year simulation was pro-
 236 duced providing modeled discharge and ammonia concentrations at the CSO junction
 237 with a 10 minute resolution. Further details can be found in Supplementary Material
 238 (Text S2, Figures S1 and S2).

220 **Table 1.** saCSOny model parameters. ^aJFM rainfall parameters are listed with other seasonal
 221 parameters listed in Table S2. Mean values for dry-spell sanitary parameters are tabulated. The
 222 95% confidence interval is denoted by \pm .

Parameter	Value	Estimation method
λ	0.30 d ⁻¹	Rainfall event analysis ^a
γ	0.54 mm ⁻¹	
k	$2 \pm 0.03 \text{ mm}^{1-\alpha} \text{ d}^{\alpha-2}$	Flow recession.
α	1.7 ± 0.2	Flow recession & topology.
Q_s	11.3 L s ⁻¹	Empirical pdf of dry spell flows.
C_s	57.2 mg L ⁻¹	
C_u	0 mg L ⁻¹	Assumed.

239 3.2 Parameter Estimation

240 The climate parameters, λ and γ , were determined from the precipitation time se-
 241 ries (Supplementary Text S2, Table S2, Figures S3 - S5). These parameters describe the
 242 exponential probability distributions of the time between rainfall events and the mag-
 243 nitude rain events (Rodriguez-Iturbe, Porporato, Ridolfi, Isham, & Coxi, 1999). A min-
 244 imum rainfall-free period of 5 h, selected as the threshold to delineate distinct rain events,
 245 was chosen based upon the flow recession characteristics which typically had returned
 246 to near pre-event flow rates within this time-frame. Due to the seasonality of rainfall the
 247 analysis was separated into annual quarters defined as January - March (JFM), April
 248 - June (AMJ), July - September (JAS) and October - December (OND). Precipitation
 249 totals for each event and the time between the start of events were determined and found
 250 to be approximately exponentially distributed for each quarter (Figures S3 and S4). The
 251 parameter λ was estimated by multiplying the frequency of actual rainfall by the long
 252 term runoff coefficient, 0.55. The parameters were estimated as the inverse of the mean
 253 of the time between rainfall events and the mean storm depth, respectively (Tables 2 and
 254 S2), equivalent to maximum-likelihood estimation. Potential for bias in the diurnal tim-
 255 ing of events was assessed, with JFM and OND events distributed indistinctly from uni-
 256 form distributions, indicating no bias in timing at a daily timescale (Figure S5). Events
 257 in AMJ and JAS were found to be significantly different from a uniform distribution by



223 **Figure 4.** Probability distributions of SWMM modeled and saCSOny predicted: (a) dis-
 224 charge, Q_c ; and (b) concentration, C_c . Parameters as in Table 1. A posteriori fits of the pdf by
 225 maximum likelihood are also shown (MLE 1 where k was estimated with fixed $\alpha = 1.7$; and MLE
 226 2 where both k and α were estimated, see Table S5).

258 the Kolmogorov-Smirnov (KS) test, with a preference for early to mid-morning events
 259 as compared to the late evening. While present, this bias had little impact on the esti-
 260 mated pdfs.

261 The point in the network chosen to represent combined flows was the junction im-
 262 mediately upstream of the pipe to the CSO structure (Figure 1). The parameters, k and
 263 α , were estimated from the mean of five flow recession events (Brutsaert & Nieber, 1977)
 264 (Figure 2, Table S3). Additional flow recession analyses were performed on 93 events
 265 (Table S4, Figure S6), selected with the criterion that the maximum discharge during
 266 the event was $>160 \text{ L s}^{-1}$ (i.e. approximately ten times the sanitary flow rate). Both anal-
 267 yses found a mean $\alpha = 1.7$ and mean $k = 2 \text{ mm}^{1-\alpha} \text{ d}^{\alpha-2}$ (Table 1). A log-linear rela-
 268 tionship (Figure S6c) between parameters was found between k and α . There was no ev-
 269 idence for a seasonal pattern in k or a normalized k (Dralle, Karst, & Thompson, 2015).
 270 The geomorphological approach of Biswal and Marani (2014) was applied to estimate
 271 α using the topology of the sewer network (Figure 2c,d). This independently resulted
 272 in the same value, $\alpha = 1.7$, as the mean measured recession exponent (Text S5). Sep-
 273 arately, maximum likelihood estimation (MLE) was applied to estimate k (MLE 1) and
 274 both k and α (MLE 2) using p_{Q_c} (Table S5).

275 The sanitary discharge concentration has a characteristic diurnal and weekly pe-
276 riodicity (Figure 3a,b). A two-week long period of dry weather flows was used to deter-
277 mine Q_s and C_s , and from these their respective pdfs, p_{Q_c} and p_{C_c} (Text S6). Across
278 the entire time series the $C_c - Q_c$ relationship is bound by strong dilution (i.e. $C_c \propto Q_c^{-1}$)
279 with the variation of dry weather concentrations preserved over several orders of mag-
280 nitude of Q_c (Figure 3c), supporting the use of the well mixed assumption. Hysteresis
281 is also evident, indicating that mixing is not perfect during individual events and thus
282 apparently well mixed conditions emerge over the ensemble of flow events.

283 3.3 Predicted pdfs of CSO Function

284 The observed pdfs for JFM discharge and ammonia concentration [NH_4^+] agree well
285 with the pdfs predicted from a priori estimated parameters (Figures 4; Figures S7 - S8,
286 Text S7). A posteriori fits of the pdfs by MLE (Table S5) have very similar shapes. The
287 multiple modes stem from the diurnal variation in sanitary flows and the roughness of
288 the pdfs stem from the application of Eq. 5 using a fine discretization of the empirical
289 pdfs of Q_s and C_s . Importantly, the pdfs capture the long tails of both distributions which
290 is necessary to correctly capture the load distribution for CSO events.

291 Despite the similarity, one-sample KS tests reject the hypothesis that the empiri-
292 cal and model pdfs share the same distribution. As the KS tests develops statistics based
293 upon the maximum deviation between the distributions, it is a conservative test. The
294 failure of the test may stem from some clear differences between the two distributions.
295 For discharge in the range of flows close to the upper end of sanitary flows (20 L s^{-1})
296 the saCSOny model tends to slightly over predict the likelihood of discharges. Similarly
297 for concentrations near the lower end of sanitary concentrations (30 mg L^{-1}). The lat-
298 ter may be due to the over-prediction of discharges. Small to medium rainfall events of
299 long-duration and low intensity, not well described as Poisson shocks, may be another
300 contributing factor. Hydrodynamic processes such as storage, pipe friction, and hydro-
301 dynamic dispersion may also have an influence. Despite some minor deficiencies the sim-
302 ple model is able to capture significant features of the pdfs, from a just two weeks of ob-
303 served dry-weather flows and a handful of flow recessions.

4 Sensitivity Analysis

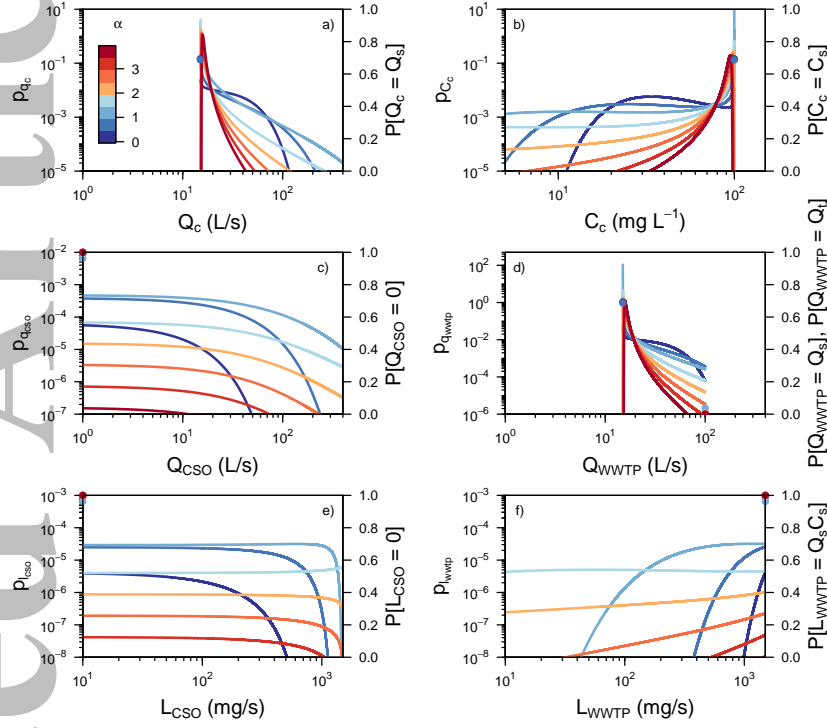
In the following subsections the effects of the four model parameters are illustrated (Figures 4 - 6 and S9- S11). The values listed in Table 1 form the base scenario and sensitivity analysis is conducted by systematically varying the others.

4.1 Network and Hydrodynamic Controls

Flow recession has a significant impact upon CSO functioning (Figure 5). As α decreases the mode of p_{Q_c} increases near Q_s , intermediate flows become more probable and larger flows less likely (Figure 5a). The pdf p_{C_c} is a mirror image of p_{Q_c} , with lower concentrations less likely with higher α . For small α there is the potential for the pdf to become bimodal (Figure 5b). Interestingly the probability of high Q_{CSO} increases with decreasing α until $\alpha = 1$, and then with further decreases the probability of high discharges declines (Figure 5c). The frequency of CSO events first increases as α increases then for $\alpha > 1$ event frequency decreases again (Figure 5c). The distribution of WWTP discharges resembles that of Q_c , albeit truncated at the acceptance threshold, Q_t (Figure 5d). For the parameters used, the pdfs of CSO load are relatively uniform for $\alpha > 1$ indicating a wide range of loads are equally probable (Figure 5e). The load probabilities decrease and increase in accord the the frequency of CSO discharge. The likelihood of smaller loads to the WWTP changes similarly with α , peaking at $\alpha \sim 1.5$ in this instance (Figure 5f).

Longer mean residence times (smaller k) increase the probability of larger combined flows, lower concentrations in combined flows, higher flows from CSOs, higher flows to WWTPs and higher loads (Figure 6). Of the hydrological parameters α is a key influence on the probability that a CSO is discharging (Figures 7). The parameter k , which controls hydrodynamic response times, influences the probability of CSO discharge at smaller values. The discharge threshold is also significant with declining likelihood of CSO discharge the higher the threshold and the higher the threshold the more significant are seasonal differences in the rainfall (Figure S9). Some variability of the discharge threshold could be expected to occur due to the hydraulics of pipe flow. Some variability is also due to discharge measurement errors, as practical CSO construction often differs from principles of weir design for the purposes of flow measurement (Ahm, Thorndahl, Nielsen, & Rasmussen, 2016). The effect of degree of variation in Q_t on the frequency of CSO discharge can be inferred from Figure S9. In the case of Lochwitzbach 10 - 20% varia-

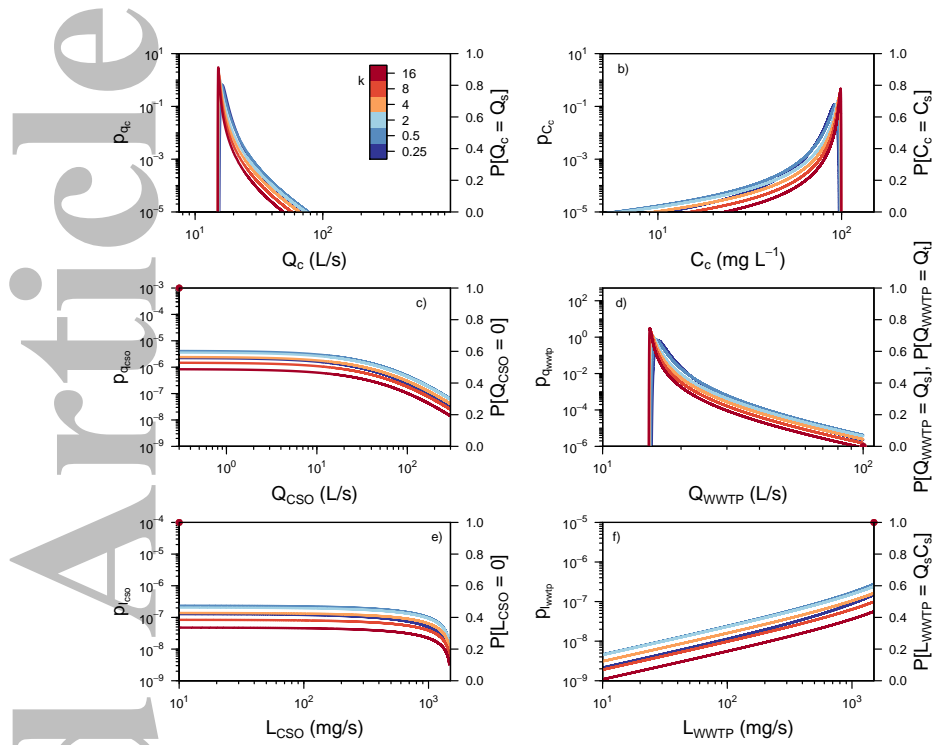
335 tion in Q_t would produce small differences in the order of magnitude estimate of the CSO
 336 frequency. Smaller thresholds however would result in larger absolute errors.



337 **Figure 5.** The impact of the flow recession parameter, α , on pdfs of: (a) Q_c ; (b) C_c ; (c)
 338 Q_{CSO} ; (d) Q_{WWTP} ; (e) L_{CSO} ; and (f) L_{WWTP} . Parameters used: $C_s = 100 \text{ mg L}^{-1}$, $Q_s = 15 \text{ L}$
 339 s^{-1} , $C_u = 0 \text{ mg L}^{-1}$, $k = 2 \text{ mm}^{1-\alpha} \text{ d}^{\alpha-2}$, $\gamma = 0.45 \text{ mm}^{-1}$, $\lambda = 0.3 \text{ d}^{-1}$; $Q_t = 100 \text{ L s}^{-1}$. Lines denote
 340 the continuous part of the pdf (left axes) while the circles denote the atom of probability (right
 341 axes). For L_{CSO} and Q_{CSO} the points correspond to $L_{CSO} = 0$ and $Q_{CSO} = 0$.

344 4.2 Climate Controls

345 Rainfall has a significant impact on function, as expected. Increasing rainfall fre-
 346 quency (also increasing total annual rainfall) shifts the pdfs of C_c such that lower con-
 347 centrations are more probable (Figure S10). This is in response to greater rainfall over-
 348 all. The effect of increasing mean rain event depth ($1/\gamma$) is similar (Figure S11). Increas-
 349 ing rainfall frequency and mean rain event depth increases the probability of CSO events
 350 and higher loads as both contribute to greater overall rainfall (Figures 7b, S10, S11). The



342 **Figure 6.** Sensitivity analysis to the flow recession parameter, k , with $\alpha = 1.7$. All other pa-
 343 rameters as in Figure 5.

351 impact of fewer but more intense rainfall can be seen in the frequency of CSO events (Fig-
 352 ure 7b). Climates of equal mean rainfall lie along lines with a slope of 1 in that figure,
 353 and it can be seen that a shift from a high frequency, low intensity rainfall to a low fre-
 354 quency higher intensity rainfall results in an increasing probability a CSO is discharg-
 355 ing.

356 5 Discussion

357 The saCSOny model quantified relative roles of climate and network parameters
 358 in controlling the statistics of CSO functions. The important role of climate is well known.
 359 Perhaps less well recognized is the significant effect that the network topology has upon
 360 the variability of CSO functioning.

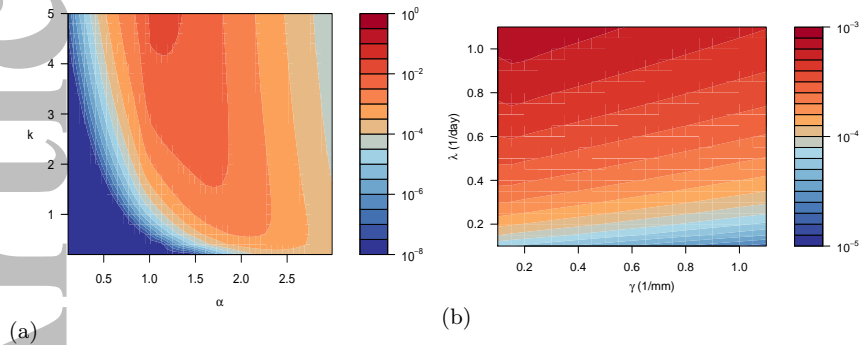
5.1 Network Controls

For Lockwitzbach at least it was demonstrated that the topology of the sewers could predict α . More work needs to be done to establish the extent to which this is more generally applicable to sewers. The empirical studies linking flow recession to topology have all been conducted on rivers to date (Biswal & Marani, 2014). With sewersheds evolving from simple linear features at early stages of development, towards fractal objects with topological properties of rivers, we expect α to change as they grow (Yang et al., 2017). Biswal and Marani (2014) suggested $\alpha \sim 1/(1 - H)$, where $H \sim 0.6$ is Hack's exponent. For sewers it has been shown H decreases from ~ 1 to 0.6 as they matured (Yang et al., 2017), which suggests α decreasing from ∞ to 2.5 during growth. While the relation suggested by Biswal and Marani (2010) may be valid for mature river networks, this suggests it may not be relevant for growing sewers. Intuition suggests that early on flow resembles a simple linear reservoir (i.e. $\alpha = 1$) and as the complexity of the network develops α likely increases. If this were the case the sensitivity analysis suggests that for the Lockwitzbach at least, high CSO loads and discharges tend to be more probable when $\alpha \sim 1$ (see Figure 4), thus poor performance of the CSO is more likely.

For k the expected changes seem to be clearer, as it is expected to decrease as the length of the pipe network and as the total area of connected impervious surface expands. The parameter k can be impacted by numerous factors. Longitudinal growth of the network would lengthen mean travel times of water and reduce k . Green infrastructure may also delay and lengthen travel times as a design goal. The results for Lockwitzbach suggests that the frequency of CSO events would decline further were k to decrease.

A take-home message for a sewer manager is that alternative network structures will have varying flow recession exponents and, as a result, varying water quality outcomes. Designing the right structure, from a network perspective, has the potential to lower the costs and reduce the constraints to mitigate CSO impacts on receiving waters. The Lockwitzbach was the first sewer system in which α was predicted from the topology, so much more work needs to be done to evaluate this approach and identify its limitations in application to other sewersheds. Additionally, to design for growing infrastructure sewer managers would be further supported by providing them with knowledge as to how to design a network to achieve a set of hydrodynamic parameters as well as

392 to predict how these properties age as the sewersheds self-organise over time (Semadeni-
 393 Davies et al., 2008).



394 **Figure 7.** Proportion of time (log10) a CSO discharges as a function of: (a) the topol-
 395 ogy/hydrodynamic parameters; and (b) the rainfall parameters. Parameters used include
 396 $Q_t = 100 \text{ L s}^{-1}$, an impervious catchment area of 36 ha and for: (a) $\lambda = 0.3 \text{ d}^{-1}$, $\gamma = 0.45$
 397 mm^{-1} ; and (b) $\alpha = 1.7$, $k = 2 \text{ mm}^{-0.7} \text{ d}^{-0.3}$.

398 5.2 Climate Controls

399 Regional, seasonal and inter-annual variations in rainfall properties vary significantly
 400 and may explain large differences in CSO performance. We see (Figure 7) that increas-
 401 ing the likelihood of large rainfall events (smaller γ) leads to increased frequencies of CSO
 402 events Sterk et al. (2016). As the model assumes exponential distributions of rainfall depth
 403 and inter-even times it is best suited to describing what happens during typical condi-
 404 tions and may not be best at describing very rare events.

405 Catchment managers can't be expected to control the rainfall, as one reviewer pointed
 406 out, but it should be remembered that λ is an effective rainfall event rate, incorporat-
 407 ing the filtering of smaller, non-productive events, and thus the runoff coefficient. Catch-
 408 ment managers can therefore directly influence the course of λ by supporting green in-
 409 frastructure, pervious paving, and managing the connectivity of impervious area, amongst
 410 others actions. For example, green infrastructure can increase infiltration, increase de-
 411 tention storage, and reduce the peak flows of urban runoff, thereby reducing CSO loads
 412 (Riechel et al., 2016). Increased detention storage would decrease λ though not significantly
 413 impact γ (Rodriguez-Iturbe et al., 1999). A λ for green infrastructure, λ_g , can be esti-

414 mated as: $\lambda_g = \lambda \exp(-\gamma s)$, where s is the effective catchment-scale detention stor-
415 age added. In the case of Lockwitzbach the effect of adding an extra 1 mm of detention
416 storage as green infrastructure would reduce the JFM λ from 0.3 d^{-1} to 0.17 d^{-1} . Assum-
417 ing that α and k remain unchanged the frequency of CSO discharges would be expected
418 to decrease approximately three-fold (Figures 7b, and S10). Natural multi-decadal vari-
419 ability as well as climate change related impacts on rainfall patterns have the potential
420 to impact water quality outcomes (Mellander et al., 2018; Semadeni-Davies et al., 2008;
421 Sterk et al., 2016). The saCSO_{ny} model offers the potential for sewer system managers
422 to better plan for a mitigate these impacts.

423 5.3 Mixing Assumptions

424 The $C \propto Q^{-1}$ relationship, bounding the SWMM-simulated values (Figure 3), may
425 be partly the result of the assumption in SWMM that individual pipes are completely
426 mixed, high-dispersion reactors (Rossmann, 2010). This need not necessarily be the case
427 at the scale of a sewershed, however in the case of the entire Lochwitzbach sewershed
428 well-mixed conditions remain a reasonable approximation. Contrasting spatial distribu-
429 tions of stormwater and sanitary inflows likely determine to what extent complete mix-
430 ing is a reasonable approximation (Krebs et al., 1999). Power law $C-Q$ relationships,
431 $C = dQ^{-h}$ with $h < 1$, may be evidence of such incomplete mixing. Partial mixing
432 could be introduced into Eq. 3 and distributions derived in a similar way (Text S1). Cur-
433 rently this would rely on an empirical $C - Q$ relationship to establish the mixing pa-
434 rameters which is somewhat unsatisfactory. As the flow recession exponent is estimated
435 from the network topology it seems plausible that in the future related methods might
436 be developed to predict d and h a priori, in a similar manner as has been done for flow
437 recession (Biswal & Marani, 2014).

438 5.4 Assessing impacts on receiving waters

439 The pdf of CSO loads can be used to estimate impacts upon receiving waters. Where
440 guidelines specify CSO loads with respect to dry flow rates in a river (Holzer & Krebs,
441 1998), then the pdfs of CSO load can be integrated to estimate the probability of not
442 meeting a dilution threshold. Alternatively, where the river responds on much longer time
443 scales, say several days to rise and fall from a single rainfall event, then the pdf of a dy-
444 namic load threshold can be estimated assuming load and river discharge are indepen-

445 dent random variables in a manner similar to Eq. (4). In the case of the small Lockwitzbach
 446 stream, the discharges would be strongly correlated with the sewer flows at sub-daily
 447 time scales. In this case consideration of the covariance between stream and CSO dis-
 448 charges would be required. The size of the sewershed in relation to the receiving water
 449 should also be a consideration in assessing the applicability of the saCSO_{ny} model. It
 450 is expected small to medium, gravity-driven sewersheds, with a small number of outlets
 451 would be most suitable, however additional research comparing saCSO_{ny} predictions with
 452 sewer performance would help clarify the situations where the model is and is not suit-
 453 able.

454 6 Conclusions

455 A four-parameter analytical model has been developed here to explore hydrolog-
 456 ical and climate factors influencing the functioning of a simple combined sewer overflow
 457 system. We demonstrated that three of the parameters of the model can be estimated
 458 readily a priori from the climate and the structure of the sewer network and one param-
 459 eter from a short time series of observed discharge by flow recession analysis. A significant
 460 finding is that the flow recession exponent may be estimated from the sewer topology,
 461 and it significantly impacts variability of CSO function. This suggests that the statis-
 462 tical properties can be estimated from the design and a minimum of data without the
 463 need for solution of the full de Saint-Venant equations. Furthermore, relative contribu-
 464 tions to variability from rainfall and the hydrodynamics/sewer structure can be disen-
 465 tangled. The equations derived here offer new approaches to rapidly assess options to
 466 mitigate CSO impacts on urban rivers. Future work is required to test the saCSO_{ny} model
 467 across diverse urban settings.

468 A The saCSO_{ny} Model

469 The pdf for stormwater discharge in the case of linear case is (Botter et al., 2009):

$$470 \quad p_{q_u}(Q_u) = \frac{\gamma^{\frac{\lambda}{k}} Q_u^{\frac{\lambda}{k}-1}}{\Gamma\left(\frac{\lambda}{k}\right)} \exp[-\gamma Q_u] \quad (\text{A.1})$$

471 and for the nonlinear case (Botter et al., 2009):

$$472 \quad p_{q_u}(Q_u) = \frac{K}{Q_u^\alpha} \exp\left[-\frac{\gamma}{k} \frac{Q_u^{2-\alpha}}{2-\alpha} + \frac{\lambda}{k} \frac{Q_u^{1-\alpha}}{1-\alpha}\right] + K \frac{k}{\lambda} \delta(Q_u) \quad (\text{A.2})$$

473 Using Eq. A.2 the remaining pdfs for flows, loads and concentrations for the nonlinear
 474 case (the linear case is omitted for space as it can be derived similarly) can be derived

475 using a change the variables, i.e.:

$$476 \quad p_y(Y) = p_x(f^{-1}(Y)) \left| \frac{\partial f^{-1}}{\partial Y} \right| \quad (\text{A.3})$$

477 where $f^{-1}(Y)$ is the inverse of a function $Y = f(X)$ of a random variable, X , with prob-
 478 ability density, $p_x(X)$, and $p_y(Y)$ is the pdf of Y . Applying a change of variables in the
 479 case of the combined flows i.e. $Q_c = Q_s + Q_u$, gives the pdf of Q_c , as:

$$480 \quad p_{q_c}(Q_c) = KG(Q_c - Q_s) + K \frac{k}{\lambda} \delta(Q_u) \quad (\text{A.4})$$

481 where:

$$482 \quad G(x) = \frac{1}{x^\alpha} \exp \left[-\frac{\gamma}{k} \frac{x^{2-\alpha}}{2-\alpha} + \frac{\lambda}{k} \frac{x^{1-\alpha}}{1-\alpha} \right] \quad (\text{A.5})$$

483 The remaining pdfs are derived similarly. With a CSO event triggered when $Q_c > Q_t$
 484 then the pdf of its discharge, Q_{CSO} , can be determined to be:

$$485 \quad p_{q_{CSO}}(Q_{CSO}) = KG(Q_{CSO} + Q_t - Q_s) + P[Q_c < Q_t] \delta(Q_{CSO}) \quad (\text{A.6})$$

486 where:

$$487 \quad P[Q_c < Q_t] = \int_0^{Q_t} p_{q_c}(Q_c) dQ_c \quad (\text{A.7})$$

488 The pdf for Q_{WWTP} is:

$$489 \quad p_{q_{WWTP}}(Q_{WWTP}) = KG(Q_{WWTP} - Q_s) + P[Q_t < Q_c] \delta(Q_{WWTP} - Q_t) + \\ 490 \quad P[Q_u = 0] \delta(Q_u) \quad (\text{A.8})$$

491 where:

$$492 \quad P[Q_t < Q_c] = \int_{Q_t}^{\infty} p_{q_c}(Q_c) dQ_c \quad (\text{A.9})$$

493 and $P[Q_u = 0]$ is given by the last term in Eq. (A.2). The pdf of the concentration of
 494 effluent is:

$$495 \quad p_{C_c|Q_s, C_s}(C_c) = K \frac{|C_u - C_s|}{(C_u - C_c)^2} G \left(Q_s \left(\frac{C_s - C_c}{C_c - C_u} \right) \right) + P[Q_u = 0] \delta(C_s - C_c) \quad (\text{A.10})$$

496 where we have written the pdf as a marginal distribution so as to recognize the possi-
 497 bility that Q_s and C_s may themselves display a degree of variability. While it is possi-
 498 ble to derive the full pdf of CSO loads, for the sake of space and simplicity the case when
 499 the stormwater concentrations are negligible, i.e. $C_u \ll C_s$, is shown:

$$500 \quad p_{l_{CSO}}(L_{CSO}) = K \frac{L_s}{(L_{CSO} - L_s)^2} G \left(\frac{L_s}{(L_s - L_{CSO})} Q_t - Q_s \right) + \\ 501 \quad P[Q_u \leq Q_t] \delta(L_{CSO}) \quad (\text{A.11})$$

502 where the sanitary load, $L_s = Q_s C_s$ has been substituted. The pdf of WWTP loads
 503 is:

$$504 \quad p_{L_{WWTP}}(L_{WWTP}) = K \frac{Q_t L_s}{L_{WWTP}^2} G \left(\frac{L_s}{L_{WWTP}} Q_t - Q_s \right) + \quad (\text{A.12})$$

$$505 \quad P[Q_u \leq Q_t] \delta(L_s) \quad (\text{A.13})$$

506 Acronyms

507 **CSO** Combined Sewer Overflow

508 **UWWTP** Urban Waste Water Treatment Plant

509 **SWMM** Stormwater Management Model

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