

This manuscript has been submitted for publication to *Geochemical Perspectives Letters* on 19 March 2024. Please note that, despite having undergone one round of peer-review, the manuscript has not yet been accepted for publication (corrections recently returned to journal). Subsequent versions of this manuscript may have slightly different content. Feel free to reach out the corresponding author. We welcome constructive feedback.

1 Medieval and recent SO₂ budgets in the Reykjanes
2 Peninsula: implication for future hazard

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20 Abstract

21 Exposure to volcanic SO₂ can have adverse effects on human health, with severe respiratory disorders
22 documented on short- and long-term timescales. Here, we use melt inclusion and groundmass glass
23 data to calculate potential syn-eruptive SO₂ emissions during the medieval and the recent 2021-2024
24 eruptions across the Reykjanes peninsula, the most populated area of Iceland that has recently
25 undergone magmatic reactivation with the 2021-2024 eruptions at Fagradalsfjall and Svartsengi. We
26 target 16 individual eruptions from the medieval volcanic cycle at the RP, the 800-1240 AD Fires, along
27 with the 2021-23 Fagradalsfjall eruptions and the 2023-24 eruptions at Sundhnúksgígar. We calculate
28 potential SO₂ emissions across the RP to be in the range 0.004-6.3 Mt. These estimates correspond to
29 mean daily SO₂ emissions in the range 1000-111000 tons/day, higher than the mean SO₂
30 measurements of 5240 ± 2700 tons/day during the 2021 Fagradalsfjall eruption. By using pre-eruptive
31 sulfur values preserved in undegassed melt inclusions, we develop an empirical approach to calculate
32 best- and worst-case potential SO₂ emission scenarios of any past or ongoing RP eruption of known
33 effusion rate. We conclude that the potential sulfur emissions across the RP can be significantly higher
34 than observed during the 2021 Fagradalsfjall eruption, mainly because of the more evolved nature and
35 higher sulfur contents of magmas erupted during the medieval time. Based on dominant NW wind
36 directions on the RP, eruptions in Brennisteinsfjöll pose the greatest health hazard to the capital area.
37 Sulfate aerosol produced during long-term eruptions may impact visibility and air quality in the Keflavík
38 airport area. Our findings enable assessment of SO₂ emission scenarios of future eruptions across the
39 RP and can be used together with gas dispersal models to forecast SO₂ pollution at ground level and
40 its impact on human health.

41

42

43 Introduction

44 The release of volcanic gases and aerosols during volcanic eruptions can significantly impact the air
45 quality and climate (e.g. Ilyinskaya *et al.*, 2017), as well as the biodiversity (e.g., Weiser *et al.*, 2022).
46 Among volcanic gases, sulfur species (SO_2 , H_2S) and associated aerosols (SO_4 , H_2SO_4) are the most
47 critical airborne hazard to human health, with short and long-term impacts that have been recorded
48 at variable distances from eruptive vents (e.g., Horwell *et al.* 2023, Stewart *et al.* 2022; Ilyinskaya *et*
49 *al.*, 2017; Schmidt *et al.*, 2015). For example, several studies have associated cardiorespiratory issues
50 with volcanic sulfur emissions (e.g., Carlsen *et al.*, 2021; Heaviside *et al.*, 2021). Hence, a detailed
51 knowledge of potential sulfur releases of active volcanoes located in densely populated areas is critical
52 to understand air quality hazards of future volcanic eruptions. This is the case of the Reykjanes
53 Peninsula (RP) in southwest Iceland, an active spreading area segmented into five volcanic systems,
54 which from west to east are Reykjanes, Svartsengi, Fagradalsfjall, Krýsuvík and Brennisteinsfjöll. The
55 latest magmatic period in the RP occurred ~800 years ago (Sæmundsson *et al.*, 2020) but knowledge
56 about sulfur outputs during those eruptions has been lacking thus far. Each volcanic system on the RP
57 tends to activate during individual magmatic periods (Sæmundsson *et al.*, 2020) and the recent 2021-
58 2024 Fagradalsfjall and Svartsengi eruptions (Barsotti *et al.*, 2023; Sigmundsson *et al.* 2024) suggest
59 the potential initiation of a new eruptive period in an area that hosts ~70% of the Icelandic population.
60 Consequently, there is an increased societal need for a deeper understanding of sulfur emissions
61 across the RP, which is crucial for a comprehensive assessment of sulfur's impact during future
62 eruptions and its potential consequences for human health.

63 Here, we focus on magmatic units erupted in the volcanic systems of Reykjanes, Svartsengi, Krýsuvík
64 and Brennisteinsfjöll in the RP during the last medieval eruptive cycle, which occurred between the 8th
65 century and 1240 AD, hereafter referred to as the 800-1240 AD Fires (Caracciolo *et al.*, 2023; Peate *et*
66 *al.*, 2009). Additionally, we target the 2021-23 Fagradalsfjall eruptions and the December 2023,
67 January 2024 and February 2024 eruptions at Sundhnúksíggar in Svartsengi. We calculate syn-eruptive
68 sulfur release and potential sulfur emissions of 19 geologically and petrochemically well characterized

69 magmatic units (Caracciolo et al., 2023; Peate et al., 2009) and compare those with sulfur emissions
70 from the 2021 Fagradalsfjall eruption (Barsotti et al., 2023; Halldórsson et al., 2022). Also, we estimate
71 daily SO₂ emissions and develop an empirical approach to calculate worst- and best-case potential
72 sulfur emissions for any eruption of a given volume emplaced in the RP.

73 Samples and methods

74 Scoria samples were collected from multiple vents within individual eruptive units of the 800-1240 AD
75 Fires (Caracciolo et al., 2023). Here, we present new sulfur (S) data for the same groundmass glass
76 (n=889) and melt inclusions (MIs) (n=416) dataset published in Caracciolo et al. (2023). Additionally,
77 we include new MI and groundmass glass data from the 2022 and 2023 Fagradalsfjall eruptions, as
78 well as data from the eruptions at Sundhnúsgíggar that occurred in the Svartsengi volcanic system in
79 December 2023, January 2024 and February 2024. S was analysed by electron microprobe analyser
80 (EMPA) at the University of Iceland using the same analytical settings as in Caracciolo *et al.* (2020) and
81 MI compositions have been corrected for post-entrapment processes (PEP) (Caracciolo et al. 2023).
82 Here, we use of the ‘petrological method’ (Devine et al., 1984) to calculate eruptive sulfur emissions
83 based on the difference between S concentrations in mineral-hosted MIs and S concentrations
84 measured in groundmass glass (ΔC_S). The idea behind this reconstruction method is that melt
85 inclusions with similar composition to erupted melts preserve the pre-eruptive volatile content, and
86 quenched groundmass glasses provide an estimate of the post-eruptive volatile content. For the
87 different magmatic units, the highest S concentration measured in PEP-corrected MIs ($C_{S\ MI}$) is
88 selected as the pre-eruptive S concentration, whereas the lowest S concentration in groundmass
89 glasses ($C_{S\ glass}$) is chosen as the post-eruptive S concentration. By combining the mass of erupted
90 magmas with the mass of S released, we can assess vent syn-eruptive SO₂ emissions (M_s) of individual
91 eruptions (eq. 1 and eq. 2 in the supplement) (e.g., Bali *et al.*, 2018; Thordarson *et al.*, 2003).
92 Furthermore, we calculate the magnitude of potential SO₂ emissions (potential M_s), which refers to
93 complete degassing of all pre-eruptive sulfur ($C_{S\ glass} = 0$) and reflects the maximum amount of SO₂

94 that a specific eruption could potentially have released, assuming that there is no degassing of
95 unerupted magma. This reconstruction method has been showed to have matched field-based volatile
96 measurements exceptionally well during the 2014-15 Holuhraun eruption (Bali et al., 2018; Pfeffer et
97 al., 2018) and the 2021 Fagradalsfjall eruption (this work, Table 1).

98

99 Sulfur concentrations in MI and groundmass glass

100 Sulfur concentration in MIs is in the range 200-1900 ppm, with a relatively large variability of S at a
101 given MI Mg#. Particularly, the most primitive MIs (Mg#>65), exclusively preserved in Reykjanes and
102 Krýsuvík, record S contents in the range 580-1070 ppm (Fig. 1). S concentration in PEP-corrected MI
103 compositions increases with decreasing MI Mg#, as expected for melt compositions controlled by
104 fractional crystallization. MIs from the 2023-24 eruptions at Sundhnúksíggar record pre-eruptive S
105 concentrations in the range 1400-1600 ppm, in agreement with MI data from the medieval eruptions
106 (Fig. 1b). MIs from the 2022-23 Fagradalsfjall eruptions closely match S concentrations measured in
107 the 2021 products (Fig. 1c). Groundmass glasses from Brennisteinsfjöll have mean S contents in the
108 range 150-280 ppm, lower than mean S contents measured in glasses from the other volcanic systems
109 (280-450 ppm) (Fig. 1, Table 1). For comparison, MIs from the 2021 Fagradalsfjall eruption contain
110 maximum S concentrations of 1200 ppm, whereas the groundmass glasses contain 20-200 ppm S.
111 Sulfide globules were not observed in the erupted samples.

112 Assessing sulfur variability and degassing during the 800-1240 AD

113 Fires

114 Considering that medieval and recent eruptions on the RP are likely sourced from mantle-derived
115 melts of diverse compositions (Caracciolo et al., 2023; Harðardóttir et al., 2022; Halldórsson et al.,
116 2022; Peate et al., 2009), including melts with variable S contents (Ranta *et al.*, 2022), we use our MI
117 record to estimate S contents of the local enriched and depleted end-member melt components. We

118 distinguish between these components from the K_2O/TiO_2 variability, a robust tracer of mantle
119 heterogeneities in Iceland (Halldórsson et al., 2022; Harðardóttir et al., 2022) (see supplement). Our
120 modelling, considering that S behaves as an incompatible element in basaltic magmas, shows that
121 most of the MI S variability can be explained by fractional crystallization (FC) and mixing of at least two
122 end-member melt compositions (Fig. 1a-d).

123 In order to evaluate S saturation during magma ascent and fractional crystallization through the crust,
124 we calculate sulfur content at sulfide saturation (SCSS) along a FC path, which reflects the amount of
125 S^{2-} present in a melt in equilibrium with a sulfide phase (Smythe et al., 2017) (see supplement). Our
126 modelling suggests that melts are sulfide undersaturated during most of magmatic fractionation across
127 the RP (Fig. 1 and Fig. S3-S4). Only magmas from Svartsengi and Brennisteinsfjöll have a high likelihood
128 to be sulfide saturated prior to eruptions. Furthermore, sulfide saturation is reached earlier during
129 magmatic differentiation of enriched mantle-derived melts than depleted melts (Fig. 1).

130 Modelling of S degassing with Sulfur_X (Ding et al. 2023) suggests that basaltic melts erupted during
131 the 800-1240 AD Reykjanes Fires are unlikely to degas significant amounts of S at known pre-eruptive
132 magma storage depths (Caracciolo et al. 2023) and that significant S degassing only takes place during
133 magma ascent in the last 0.2 kbar (<700 m) (Fig. S1).

134 Sulfur emissions across the RP

135 Sulfur release ranges between 1000-1770 ppm across the RP, a typical range for Icelandic rift basalts
136 (Ranta et al. 2024), with the largest ΔC_s found in lavas from Svartsengi and Brennisteinsfjöll (Table 1).

137 ΔC_s values can be scaled by the mass of erupted material to estimate M_s of individual eruptions, using
138 published volumes of individual eruptive units, in the range 0.01 km³ to 0.72 km³ (Table 1). Using a
139 melt density of 2700 kg/m³ and assuming a bulk vesicularity of 15 vol%, we calculate M_s between
140 0.003-5.9 Mt (Fig. 2a). The most voluminous lavas found in Svartsengi and Brennisteinsfjöll released
141 the highest mass of SO₂ into the atmosphere during the medieval time. The syn-eruptive SO₂ released
142 by these latter voluminous lavas is approximately 2 to 6 times larger than syn-eruptive SO₂ emissions

143 during the 2021 Fagradalsfjall eruption, for which we estimated $M_s = 0.78 \text{ Mt}$ ($M_{s \text{ measured}} = 0.97 \pm 0.5$,
144 Barsotti et al. 2023). These are roughly between 20 to 70 % of the syn-eruptive SO_2 emissions
145 estimated for the 2014-15 Holuhraun eruption ($M_s = 10.5 \text{ Mt}$, Bali et al. 2018). We calculate SO_2 release
146 of 0.06-0.07 Mt for the 2022 and 2023 Fagradalsfjall eruptions, respectively. However, for a given mass
147 of melt, the 2021-23 Fagradalsfjall eruptions released a comparable mass of SO_2 (Table 1). Oppositely,
148 the 2023-24 eruptions at Sundhnúksíggar slightly exceeded SO_2 emissions during the 2021-23
149 Fagradalsfjall eruptions (Table 1). Similarly, we have calculated potential M_s , the maximum mass of SO_2
150 that could potentially have been released during each eruption. Potential M_s across the RP ranges
151 between 0.003-6.3 Mt and is only slightly higher than vent M_s , as most of the S is released into the
152 atmosphere during eruptions rather than staying dissolved in the lava (Table 1).

153

154 Evaluating end-member scenarios of SO_2 emissions and hazard 155 potential for future eruptions across the RP.

156 Based on the MI record of the 2021-23 Fagradalsfjall eruption (Halldórsson et al., 2022 and this work),
157 the 2023-24 eruptions at Sundhnúksíggar and the 800-1240 AD Fires (this work), we constrain
158 potential maximum (1900 ppm) and minimum (1170 ppm) pre-eruptive S concentrations and use
159 these to estimate potential M_s of future eruptions in the RP. With these constraints, we developed an
160 empirical approach to assess potential M_s for a given eruption of known lava volume, with important
161 applications for forecasting the worst- and best-case scenarios of potential M_s of future eruptive events
162 (Fig. 2b). For example, based on our approach, an eruption with an eruptive volume of 0.4 km^3 , could
163 release between 2.9 Mt and 4.1 Mt of SO_2 . This method also has an application when it comes
164 evaluating the long-term SO_2 impact of ongoing eruptions in the RP. If the mean magma effusion rate
165 is known and fixed, one can roughly estimate the volume of the lava flow and calculate potential M_s at
166 any given moment from the onset of the eruption. This provides a valuable tool to assess best- and
167 worst-case scenarios for SO_2 pollution during ongoing events.

168 Eruptive M_s calculations are strongly dependent on lava flow volumes. Hence, when it comes to
169 comparing the 800-1240 AD Fires with the 2021-24 eruptions, a more relevant parameter is the mean
170 daily SO_2 emissions, which also is an important parameter from a hazard perspective. We have
171 estimated daily SO_2 emissions for the 800-1240 AD Fires using mean output rates (MOR) calculated by
172 Oskarsson et al. (2024), in the range 3-119 m^3/s (Table 1, Eq. 3 supplement). Mean daily SO_2 emissions
173 during the medieval eruptions likely ranged between 1000 ton/day to 111000 ton/day (Fig. 2c). In
174 comparison, during the 2021, 2022 and 2023 Fagradalsfjall eruptions, we calculate average daily SO_2
175 emissions of 5000, 3780 and 3360 ton/day, respectively. The estimate for the 2021 Fagradalsfjall
176 eruption is in agreement with the majority of measured daily SO_2 emissions throughout the 2021
177 Fagradalsfjall eruption, in the range 1000-7600 ton/day (Esse et al., 2023), and with daily SO_2
178 emissions of 5240 ± 2700 ton/day, calculated assuming 0.97 ± 0.5 Mt of total mass of SO_2 (Barsotti et
179 al. 2023). Oppositely, the Dec. 2023 Sundhnúkar eruption released 32000 ton/day SO_2 (Table 1). Our
180 calculations highlight that future eruptions in the RP may have the potential to release significantly
181 more SO_2 on a daily basis than the 2021-24 eruptions.

182 SO_2 emissions during the 800-1240 AD Fires and the 2021-24 eruptions are small compared to those
183 during the 2014-15 Holuhraun basaltic eruption (9.2 Mt SO_2 , Pfeffer *et al.*, 2018). However, volcanic
184 eruptions in the RP are potentially considered to be more hazardous due to their proximity to
185 inhabited areas, to the international airport and to the large number of visitors expected at eruption
186 sites (Fig. 3) (Barsotti et al. 2023). To assess the health hazard for potential future eruptions, we built
187 seasonal wind roses, for the period 2012-2022, reflecting dominant wind speeds and directions in the
188 RP (Hersbach et al. 2023). We find that most of the time prevailing winds blow towards the NW-NE,
189 suggesting different SO_2 health hazards potential associated with eruptions within different volcanic
190 systems (Fig. 3). The prevalent NW wind blowing direction suggests that volcanic SO_2 emissions could
191 still be disruptive to the Keflavik airport area if there were a long-duration eruption. Even if eruptions
192 in the RP produce little ash, sulfate aerosol in the atmosphere could reduce visibility and air quality
193 (Pattantyus et al. 2018). Eruptions in Brennisteinsfjöll are the most hazardous for Reykjavík, especially

194 in spring and autumn seasons, as SO₂ is likely to be blown towards the capital area. Eruptions in
195 Reykjanes pose minimal hazard as winds tend to blow away from inhabited areas. During assessment
196 of possible eruptive scenarios in the RP, our estimates provide key input parameters to model the
197 release and dispersion of volcanic SO₂ into the atmosphere. Our results can be used to inform SO₂
198 pollution hazard assessments for potential eruptive scenarios and prompt action and mitigation plans
199 during ongoing volcanic crises in the RP.

200

201 **Acknowledgement**

202 This research was financially support by a NordVulk fellowship awarded to AC and by the Icelandic
203 Research Fund (grant 228933-052). We acknowledge support from the [Gosvá project](#), a research
204 programme on the assessment of volcanic hazard risks in Iceland led by the Icelandic Meteorological
205 Office (IMO). SAH acknowledges support from the Icelandic Research Fund (Grant #196139-051). We
206 thank Christoph Kern, two anonymous reviewers, and editor Ambre Luguët for their constructive
207 comments, which significantly improved the quality of the manuscript.

208

209 **Figure Captions**

210 **Table 1.** Eruptive units studied in this work and summary of main results. C_{S MI}= pre-eruptive S
211 concentration. C_{S glass} = post-eruptive S concentration. ΔC_S = sulfur emissions at the vent per unit mass
212 of melt, accounting for crystallinity. V_l = bulk lava volume. V_{DRE} = Vesicle-free lava volume. M_s = Syn-
213 eruptive SO₂ emissions at the vent. Potential M_s = potential SO₂ emissions. Lava volumes for the
214 medieval eruptions are from Einarsson et al. (1991), Jónsson (1978) and Sigurgeirsson (2004). ^a Lava
215 volume estimated by assuming a thickness of 5 m, consistent with average thicknesses of lava flows of
216 known volumes with a similar aerial extent. ^b MOR values (within brackets) and uncertainty ranges for
217 the medieval eruptions are from Óskarsson et al. (2024). ^c V and MOR from Pedersen et al. (2022). ^d V
218 and MOR from Pedersen et al. (2024). *Lava volumes for the 2024 eruptions at Sundhnúksígígar are
219 not available at the current stage.

220

221 **Fig. 1.** (a-d) Variation of S contents in groundmass glasses (filled circles) and PEP-corrected MIs (filled
222 triangles) as a function of Mg# [$Mg\# = 100 \cdot Mg / (Mg + Fe^{2+})$, $Fe^{2+} / Fe^{tot} = 0.9$] in samples from the 800-
223 1240 AD Fires, the 2021-2023 Fagradalsfjall eruptions and the 2023-24 eruptions at Sundhnúksíggar.
224 Data from the 2021 Fagradalsfjall eruption are from Halldórsson *et al.* (2022). Red and blue solid lines
225 indicate fractional crystallization paths calculated for a geochemically enriched and depleted initial
226 melt compositions, respectively (see supplement). The black dotted curve indicates SCSS along an
227 empirical fractional crystallization path calculated after Smythe *et al.* (2017), implemented in PySulfSat
228 (Wieser and Gleeson, 2022).

229

230 **Fig. 2.** (a) Variation of vent M_s (b) Magnitude of potential M_s as a function of eruption volume for the
231 800-1240 AD Fires, the 2021-23 Fagradalsfjall eruptions and the 2023 Sundhnúkar eruption. At a given
232 volume, straight lines allow to calculate potential M_s corresponding to maximum and minimum pre-
233 eruptive S concentrations measured across the RP. Inset plot show most common potential M_s across
234 the RP (c) Daily SO_2 emissions are calculated using MOR values and associated uncertainties from
235 Óskarsson *et al.* (2024). Blue histogram indicates measured SO_2 emissions during the 2021
236 Fagradalsfjall eruption (Esse *et al.*, 2023). Data are coloured according to the volcanic system and only
237 lavas with known volumes or MORs are included in the plots.

238

239 **Fig. 3.** Simplified geological map of the RP and lava flows emplaced during the 800 – 1240 AD Fires.
240 The map also illustrates the aerial extent of the 2021 (Pedersen *et al.*, 2022), 2022 (Gunnarson *et al.*,
241 2023) and 2023 (Belart *et al.* 2023) Fagradalsfjall lavas. Data from the 2023-24 eruptions at
242 Sundhnúksíggar are from the Landmælingar Íslands geoserver (gis.lmi.is/geoserver). When possible,
243 lava flows are coloured according to calculated syn-eruptive SO_2 emissions, ranging from 0.1 to 6 Mt.
244 Orange outlines show urban areas. Numbers reflect the different lava units as listed in table 1.
245 Seasonal wind roses reflect data at 900 mPa (~1000 m a.s.l.), which was the most common SO_2 injection

246 altitude during the 2021 Fagradalsfjall eruption (Esse et al. 2023). Spokes indicate the direction the
247 wind is blowing from, and the length of each spoke shows the frequency. Wind data were extracted
248 from ERA5 (Hersbach et al. 2023).

249

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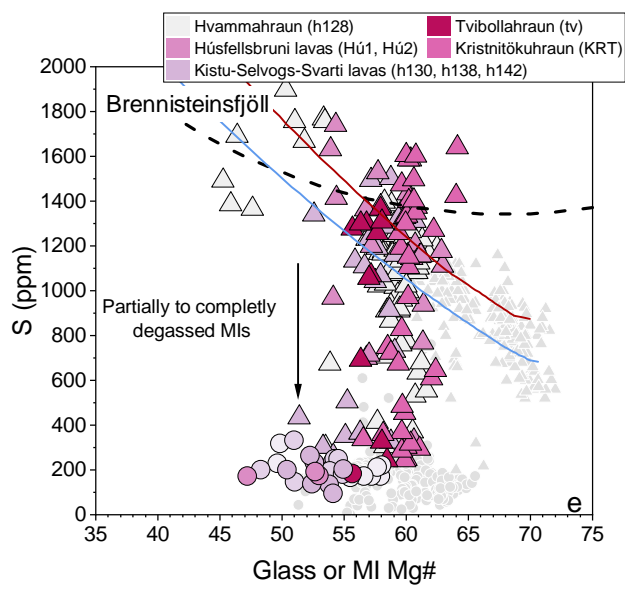
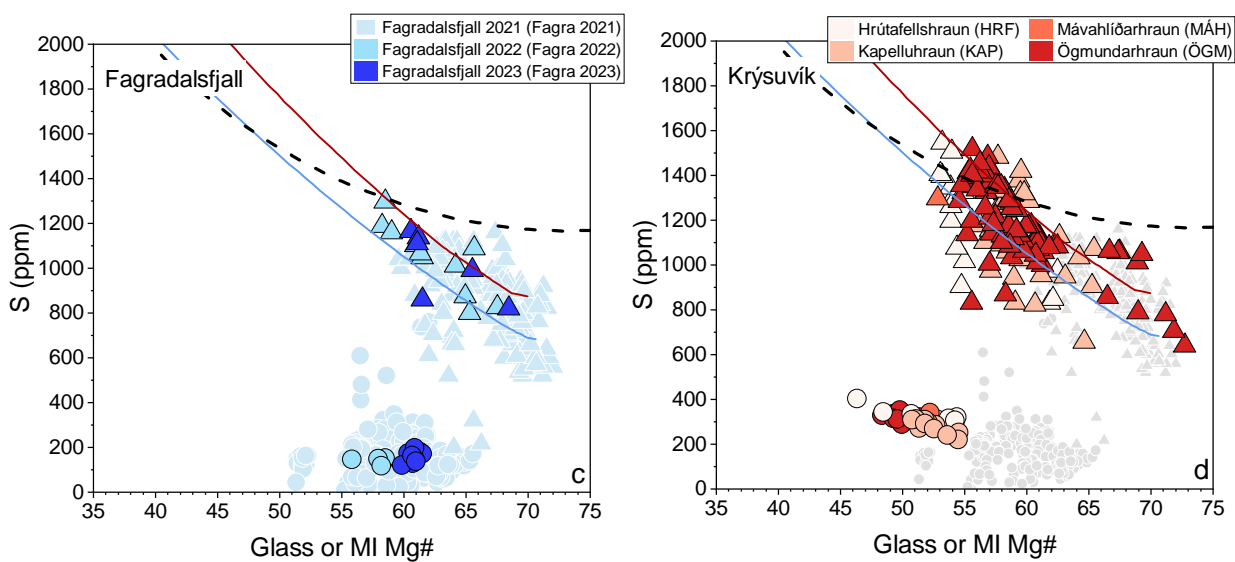
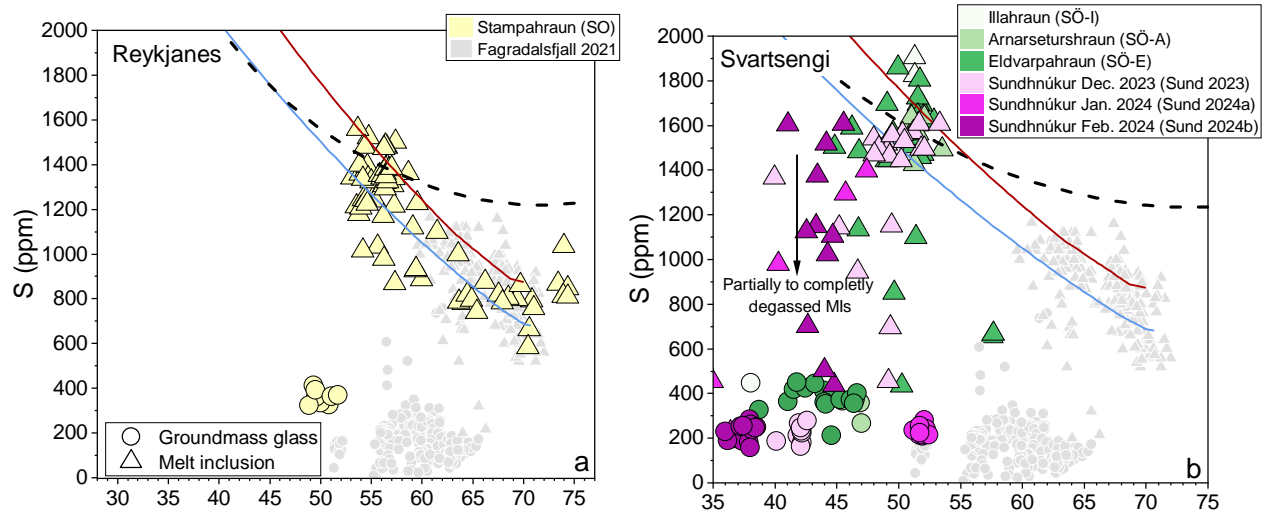
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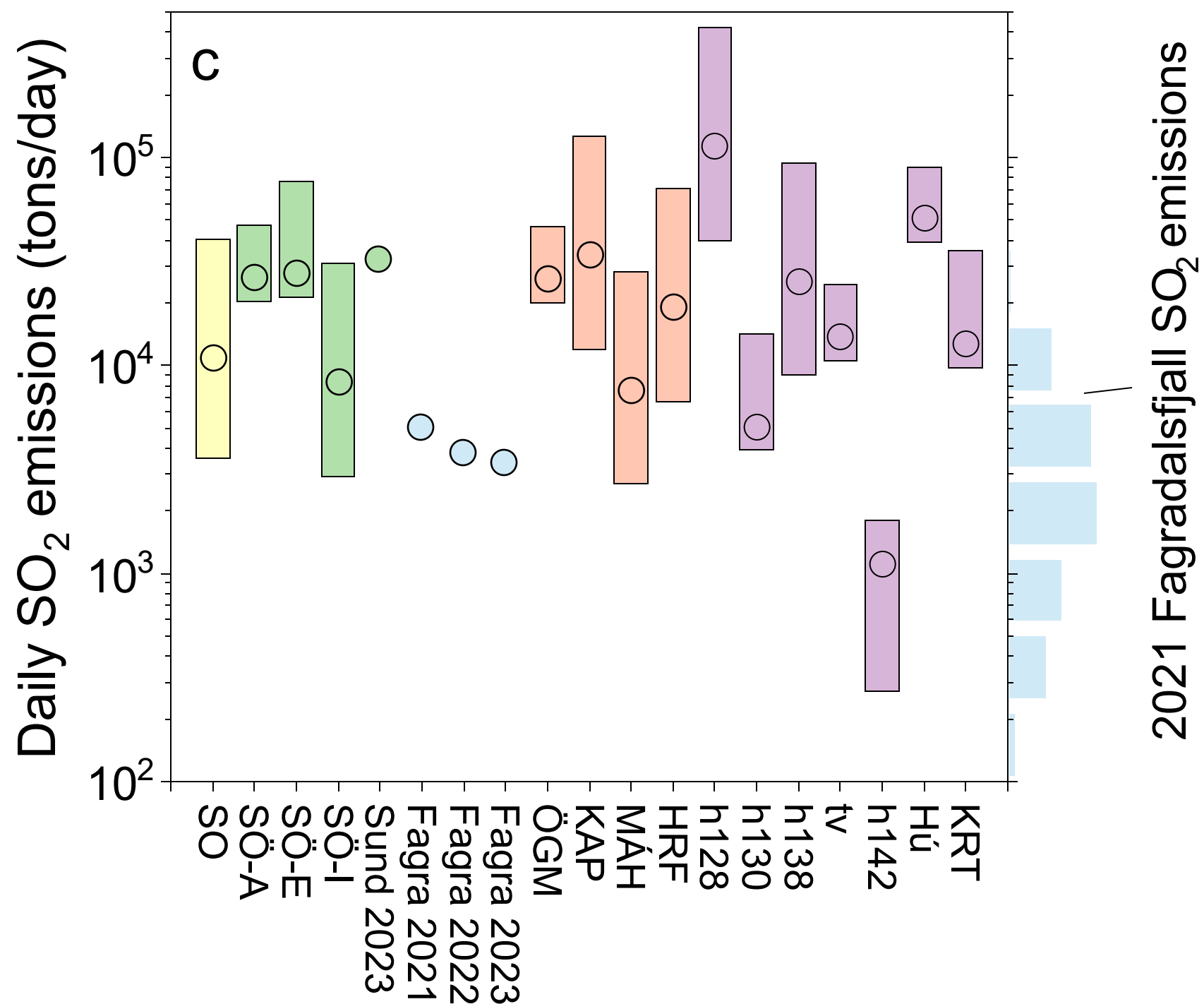
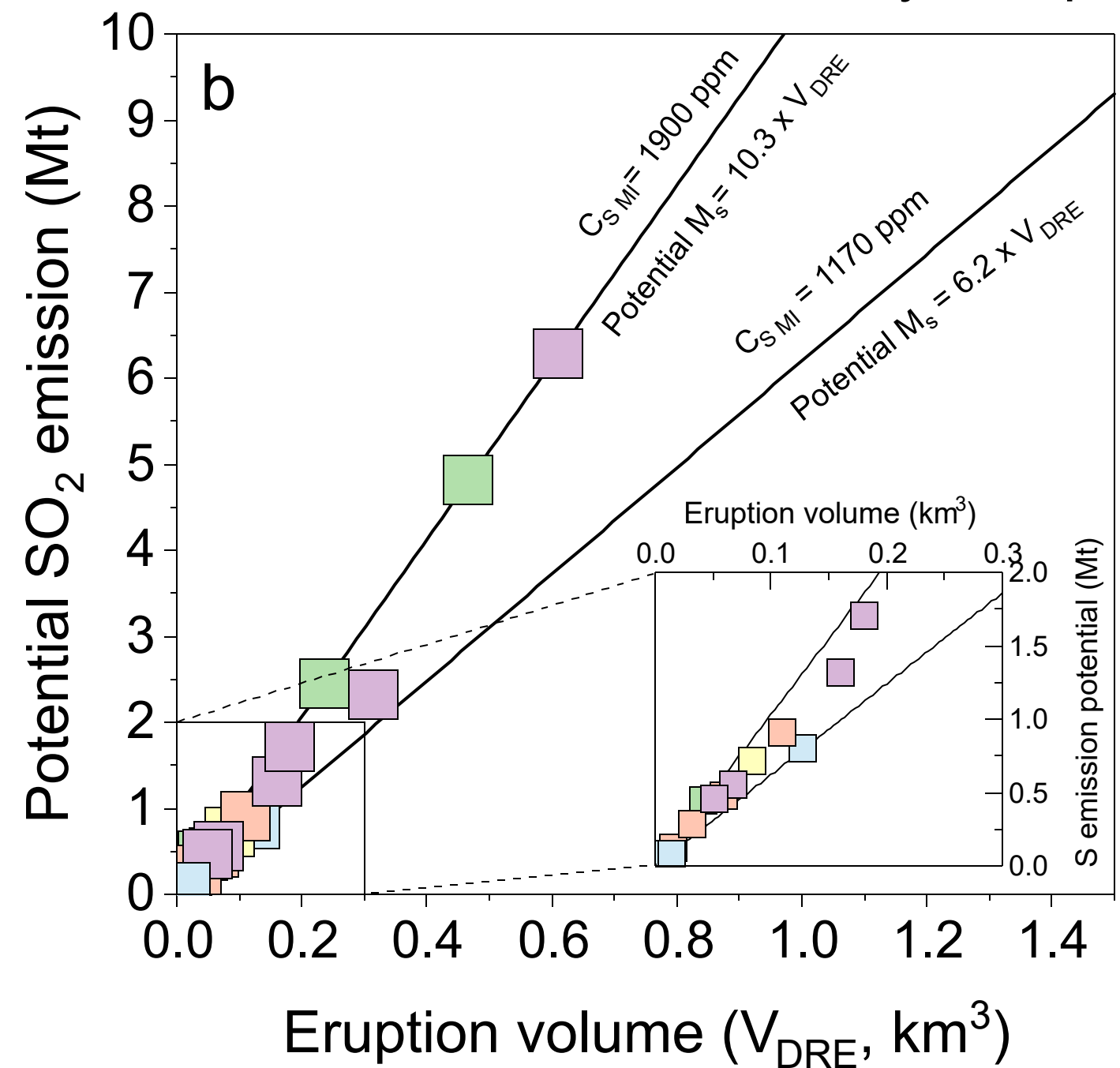
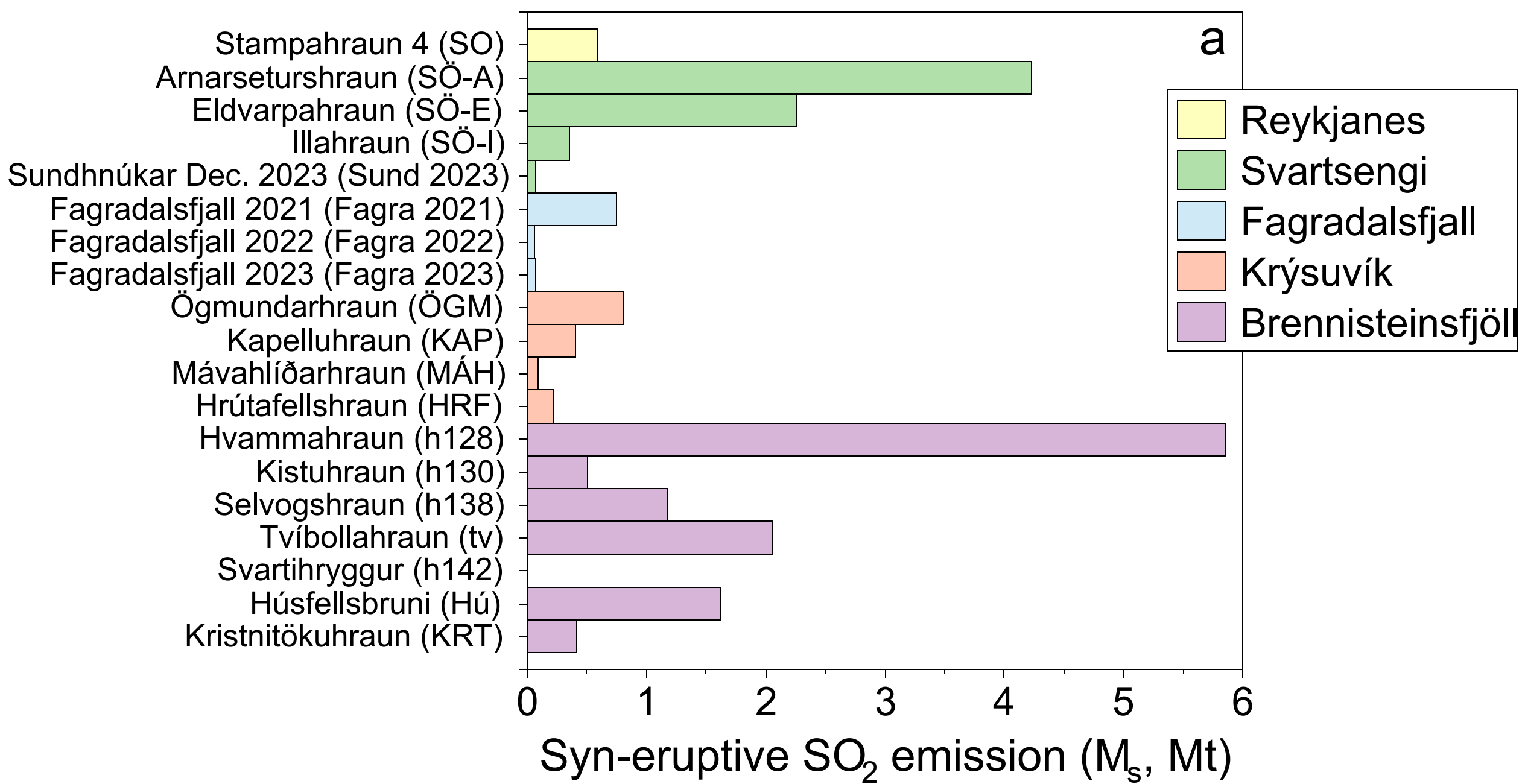
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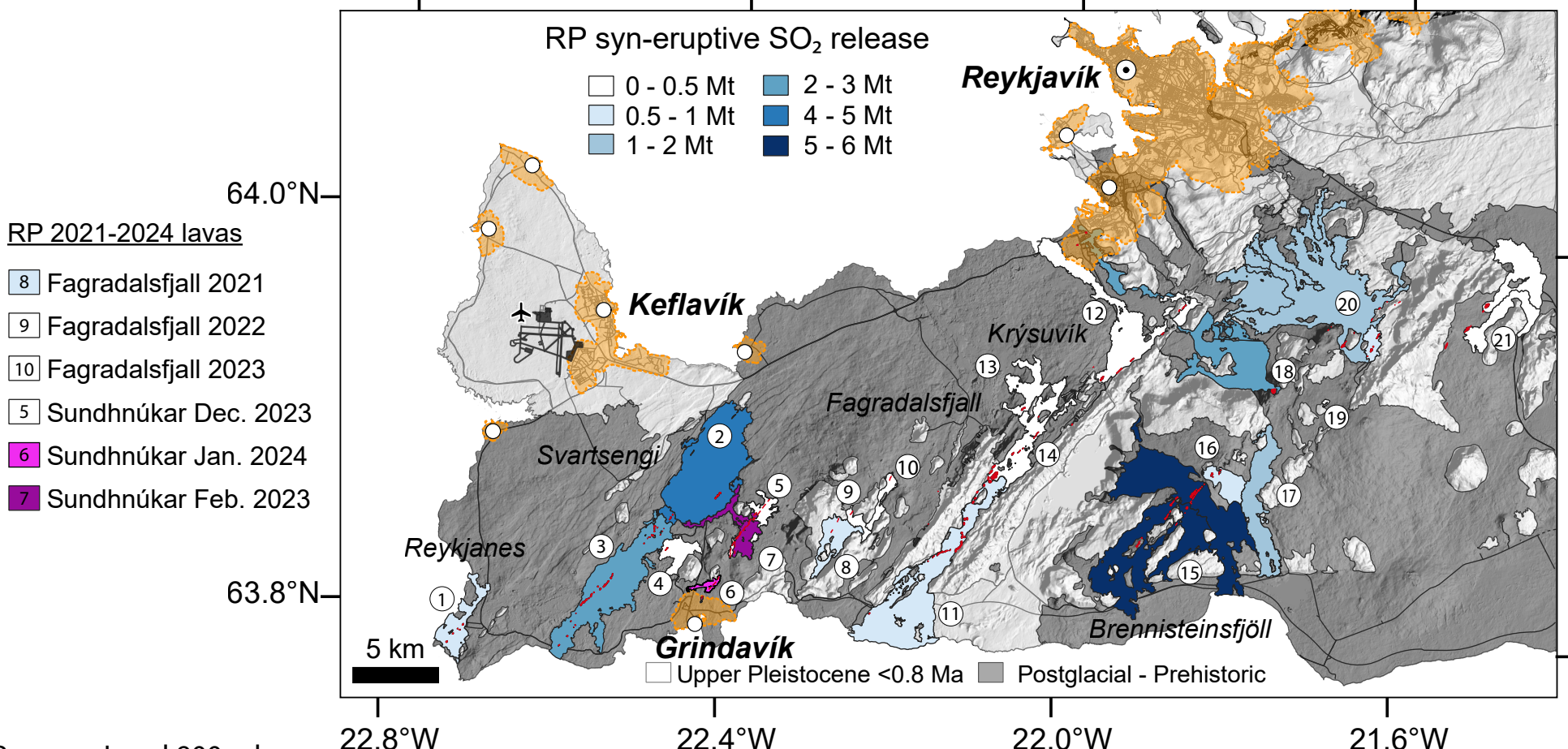
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385

Number	Eruptive unit	Acronym	Volcanic system	Age	C _{S MI}	C _{glass}	ΔC _S	V	V _{DRE}	Mass	M _S	Potential M _S	MOR ^b	Time of lava emplacement	Daily SO ₂ emissions
				A.D.	ppm	ppm	ppm	km ³	km ³	Kg	Mt	Mt	m ³ /s	days	tons/day
1	Stampahraun 4	SO	Reykjanes	1210-1240	1559	258	1275	0.10	0.09	2.3E+11	0.58	0.71	6.4 - 67.9 (17.9)	17 - 193 (65)	3570 - 40450 (10620)
2	Arnarseturshraun	SÓ-A	Svartsengi	1210-1240	1907	196	1667	0.55	0.47	1.3E+12	4.23	4.81	26.1 - 60.1 (33.5)	106 - 244 (190)	20420 - 47000 (26200)
3	Eldvarpahraun	SÓ-E	Svartsengi	1210-1240	1907	112	1759	0.28	0.24	6.4E+11	2.26	2.45	26 - 93.7 (33.3)	35 - 125 (97)	21340 - 76900 (27300)
4	Illahraun	SÓ-I	Svartsengi	1210-1240	1907	312	1563	0.05	0.04	1.1E+11	0.36	0.44	4 - 42.7 (11.2)	14 - 145 (52)	2920 - 31140 (8180)
5	Sundhnúkar Dec. 2023	Sund 2023	Svartsengi	2023	1610	210	1372	0.011 ^d	0.01	2.5E+10	0.07	0.08	50 ^d	2.5	32000
6	Sundhnúkar Jan. 2024*	Sund 2024a	Svartsengi	2024	1400	160	1215	-	-	-	-	-	-	<2	-
7	Sundhnúkar Feb. 2024*	Sund 2024b	Svartsengi	2024	1607	164	1414	-	-	-	-	-	-	<2	-
8	Fagradalsfjall 2021	Fagra 2021	Fagradalsfjall	2021	1170	20	1127	0.15 ^c	0.13	4.1E+11	0.78	0.80	9.5 ^c	185	5000
9	Fagradalsfjall 2022	Fagra 2022	Fagradalsfjall	2022	1300	120	1180	0.011 ^d	0.01	2.5E+10	0.06	0.07	7.0 ^d	18	3780
10	Fagradalsfjall 2023	Fagra 2023	Fagradalsfjall	2023	1170	120	1050	0.015 ^d	0.013	3.4E+10	0.07	0.08	7.0 ^d	26	3360
11	Ögmundarhraun	ÖGM	Krýsuvík	1151-1188	1517	138	1351	0.13	0.11	3.0E+11	0.81	0.90	31.9 - 73.3 (40.8)	21 - 47 (37)	20110 - 46220 (25750)
12	Kapelluhraun	KAP	Krýsuvík	1151-1188	1482	183	1273	0.07	0.06	1.6E+11	0.41	0.48	20.2 - 213 (56)	4 - 40 (14)	12000 - 126500 (33240)
13	Mávahlíðarhraun	MÁH	Krýsuvík	1151-1188	1297	280	997	0.02	0.02	4.6E+10	0.09	0.12	5.8 - 61 (16)	4 - 40 (14)	2700 - 28370 (7450)
14	Hrútafellshraun ^a	HRF	Krýsuvík	8 th - 9 th century	1546	252	1268	0.04 ^a	0.03	9.0E+10	0.23	0.28	11.4 - 120 (31.5)	4 - 40 (14)	6750 - 71000 (18660)
15	Hvammahraun	h128	Brennisteinsfjöll	8 th - 9 th century	1900	90	1810	0.72	0.61	1.7E+12	5.86	6.27	48.3 - 510.6 (134.1)	16 - 173 (62)	39970 - 422560 (111000)
16	Kistuhraun	h130	Brennisteinsfjöll	900-1100	1494	94	1372	0.08	0.07	1.8E+11	0.50	0.55	6.1 - 22 (7.8)	42 - 152 (119)	3900 - 14000 (5000)
17	Selvogshraun	h138	Brennisteinsfjöll	10 th - 11 th century	1504	126	1350	0.19	0.16	4.4E+11	1.18	1.31	14.2 - 149.9 (39.4)	15- 155 (56)	8950 - 94450 (24810)
18	Tvíbollhraun	tv	Brennisteinsfjöll	950	1367	129	1213	0.37	0.31	8.5E+11	2.06	2.32	18.7 - 43 (24)	100 - 229 (178)	10580 - 24340 (13580)
19	Svartihryggur ^a	h142	Brennisteinsfjöll	900-1200	1341	147	1170	0.0005 ^a	0.0004	1.3E+09	0.003	0.003	0.5 - 3.3 (2)	2 - 13 (3)	270 - 1800 (1100)
20	Húsfellsbruni ^a	Hú1 & Hú2	Brennisteinsfjöll	9 - 13 th century	1739	55	1650	0.20 ^a	0.17	4.9E+11	1.62	1.70	50.6 - 116.4 (64.9)	21 - 49 (38)	38960 - 89620 (50000)
21	Kristnitökuhraun ^a	KRT	Brennisteinsfjöll	1000	1640	118	1492	0.06 ^a	0.05	1.4E+11	0.41	0.45	14.1 - 50.9 (18)	14 - 50 (39)	9800 - 35420 (12570)







Pressure Level 900 mbar
Time period 2012-2022

