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A Hybrid Pore-Network-Continuum Modeling Framework for Flow and Transport in 3D Digital Images of Porous Media

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Abstract

Understanding flow and transport in multiscale porous media is challenging due to the presence of a wide range of pore sizes. Recent imaging advances offer high-resolution characterization of the multiscale pore structures. However, simulating flow and transport in 3D digital images requires models to represent both the resolved and sub-resolution pore structures. Here, we develop a hybrid pore-network-continuum modeling framework. The hybrid framework treats the smaller pores below the image resolution as a continuum using the Darcy-scale formalism and explicitly represents the larger pores resolved in the images employing a pore network model. We validate the hybrid model against direct numerical simulations for single-phase flow and solute transport and further demonstrate its applicability for simulating two-component gas transport in a shale rock sample. The results indicate that the new hybrid model represents the flow and transport process in multiscale porous media while being much more computationally efficient than direct numerical simulation methods for the range of simulated conditions.

Keywords: Pore network model, Multiscale pore structure, 3D digital images, Hybrid modeling, Continuum

Highlights

- We develop a hybrid pore-network-continuum (h-PNM-continuum) modeling framework.
- It applies to pore-scale flow and transport in 3D multiscale digital images.
- Resolved and sub-resolution pores, and the mass transfer between them are modeled.
- h-PNM-continuum is computationally more efficient than standard direct numerical simulations.
- The hybrid modeling framework can be extended to more complex processes.

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1 1. Introduction

Pore-scale flow and transport processes in porous materials play a critical role in many geoscience and 2 engineering applications, such as hydrocarbon recovery, CO_2 sequestration, contaminant transport, nuclear 3 waste storage, fuel cells, and biomedical applications. Thanks to advances in imaging technologies and 4 computational modeling, pore-scale imaging and modeling (sometimes referred to as digital rock or soil physics) have become a standard technology to study fluid flow and transport in porous media at microscopic 6 scales (Blunt et al., 2013; Bultreys et al., 2016; Xiong et al., 2016). Advanced imaging and analysis techniques 7 including non-destructive X-ray computed tomography (e.g., micro-CT or nano-CT), focused ion beam-8 scanning electron microscopy (FIB-SEM), transmission electron microscopy (TEM), and focused ion beam-9 helium ion microscope (FIB-HIM), have been increasingly used to visualize and quantify 3D pore-scale 10 structures of porous media (e.g., Wildenschild & Sheppard, 2013; Cnudde & Boone, 2013; Elmorsy et al., 11 2023; Zahasky et al., 2023; Zheng et al., 2017). While imaging characterization can now achieve resolutions 12 down to a few nanometers, digital images often cannot resolve all the pores of porous media with a wide 13 range of pore sizes due to a trade-off between the spatial resolution and the size of the field of view. As a 14 result, "two-scale" digital images—large pores resolved in the image and much smaller sub-resolution pores 15 coexist—arise when applying imaging characterization to various porous media, e.g., shale rocks, carbonates, 16 tight sandstones, coalbed, and soils. These unresolved microporous subdomains have been demonstrated to 17 have a strong impact on the flow and transport processes for many applications (Mehmani & Prodanović, 18 2014; Bultreys et al., 2015; Scheibe et al., 2015; Soulaine et al., 2016; Guo et al., 2018; Mehmani et al., 2021; 19 Yang et al., 2021), which suggests that they should be accounted for in modeling approaches. 20

The need to represent flow and transport processes in both scales of these two-scale images poses sig-21 nificant challenges to pore-scale modeling. Pore-scale modeling based on high-resolution images can be 22 generally grouped into two categories (Meakin & Tartakovsky, 2009; Blunt et al., 2013; Xiong et al., 2016; 23 Mehmani et al., 2021; Yang et al., 2016; Gong et al., 2021): (1) direct numerical simulation (DNS) and 24 (2) pore network model (PNM). Both DNS and PNM approaches have been extended to model flow and 25 transport in "two-scale" digital images. Soulaine & Tchelepi (2016) proposed a micro-continuum modeling 26 framework based on the Darcy-Brinkman-Stokes equation (DBS), where the Stokes equation is solved in the 27 resolved pores and Darcy's law is utilized for the fluid flow in the sub-resolution microporous subdomain. 28 Similar DBS-based micro-continuum frameworks have been used and generalized to include various pro-29 cesses, including solute transport (Scheibe et al., 2015), mineral dissolution and reactive transport (Golfier 30 et al., 2002; Soulaine et al., 2017, 2018; Maes & Menke, 2021; You & Lee, 2021; Molins et al., 2021; Maes 31 et al., 2022), gas and multi-fluid transport in shale rocks (Guo et al., 2018; Soulaine et al., 2019; Guo et al., 32 2019), two-phase flow (Carrillo et al., 2020; Maes & Menke, 2021; Carrillo et al., 2022), fluid-solid interaction 33 and deformation (Carrillo & Bourg, 2019, 2021), and heat transfer (Maes & Menke, 2022; Xu et al., 2022). 34 While the DBS-based micro-continuum model provides a DNS modeling framework for flow and transport in 35 "two-scale" digital images, like any other DNS models, it requires high computation costs when simulating ³⁷ high-resolution digital images of porous materials.

PNM has also been extended in various forms to model the flow and transport processes in "two-scale" porous media. The extensions are often referred to as the dual pore network model (dual PNM). Bekri et al. 39 (2005) presented a dual PNM that simulated mass transfer in a two-scale porous medium, which employed 40 a large-scale network to represent vugs or fractures and a small-scale network for the homogeneous matrix. 41 Bauer et al. (2012) introduces a dual PNM where the large-scale pore network is extracted from images, and 42 micropores function as parallel throats alongside the resolved pore throats. Based on statistical information 43 from the images at different scales, Jiang et al. (2013) employed stochastic methods to create a small-scale 44 pore network, which was then combined with the larger-scale network, resulting in a two-scale network. 45 Mehmani & Prodanović (2014) developed a dual PNM considering different connectivities between the big 46 and small networks, and used the dual PNM to investigate the impact of microporosity on transport in 47 porous media. Prodanović et al. (2015) extended this approach by determining the location of the small-48 scale pore network based on digital images of porous media samples. Bultreys et al. (2015) developed another 49 dual PNM where the microporous regions are approximated as effective pore throats that can connect with 50 the large-scale pore network in parallel and in serial. de Vries et al. (2017) developed a dual PNM to 51 study the impact of aggregates on solute transport in dual porosity porous media. These methods have 52 provided critical insights into flow and transport processes in multiscale porous media, but they generally 53 require large computation costs due to a large number of network elements (if the pore structures in the 54 microporous regions are explicitly represented) or need to simplify the spatial interconnectivity between the 55 resolved pores and the microporous regions (when the microporous regions are approximated as effective 56 throats). 57

To address the above-discussed challenges, we present a new hybrid pore-network-continuum modeling 58 framework (h-PNM-continuum). This hybrid framework treats the unresolved smaller pores (sub-resolution 59 microporous regions) as a continuum described by the Darcy-scale formalism, while explicitly representing 60 the flow and transport processes in the larger pores (the resolved pores) using a PNM. The pore network 61 is extracted from the resolved pore subdomain using the watershed-segmentation-based SNOW algorithm 62 developed by Gostick (2017). Additionally, we develop a new algorithm to connect the pore network to the 63 microporous subdomain while extracting the pore network from the 3D image. Building upon the hybrid 64 framework, we present mathematical formulations for single-phase flow, solute transport, and gas transport 65 in nanoporous shale rocks that accounts for the confinement effects of gas flow in nanopores. 66

67 2. Hybrid pore-network-continuum modeling framework (h-PNM-continuum)

In this section, we present the key ingredients of the hybrid modeling framework, which include a modified network extraction algorithm that links the extracted pore network to the microporous subdomain, mathematical formulations for single-phase flow, solute transport, and two-component gas transport in shale rocks, and numerical methods used to discretize and solve the governing equations.



Figure 1: Schematic of the hybrid pore-network-continuum modeling framework. (a) A ternary segmented image (Image #1) with blue, pink, and grey colors denoting the resolved pores, sub-resolution microporous regions, and non-porous solids, respectively. (b1) Pore network extraction from the resolved pore subdomain (Image #2). (b2) Numerical discretization of the sub-resolution microporous subdomain (Image #3). (c) The coupled hybrid pore-network-continuum (h-PNM-continuum) representation of the original ternary image. (d) A zoom-in view to illustrate the coupling between the pore network and the discretized continuum subdomain. (e) An example of pores *i* and *j* from (d) that are coupled with the discretized continuum subdomain. (f) A zoom-in view of the interface between pore *i* and the adjacent numerical cells *m*, *n*, and *k* in the continuum subdomain.

⁷² 2.1. Network extraction and coupling between the pore network and the microporous subdomain

For a ternary segmented image (Image #1) that consists of three subdomains including resolved pores, 73 sub-resolution microporous regions, and solids (Figure 1(a)), we need to extract the pore network from the 74 resolved pore subdomain and then connect the pore bodies to the sub-resolution microporous subdomain 75 based on original connectivity between the two subdomains. We do so by augmenting the watershed-76 segmentation-based SNOW algorithm developed by Gostick (2017). The augmented algorithm involves two 77 steps: (1) decompose the ternary image to extract the pore network from the resolved pore subdomain 78 (Figure 1(b1)) and discretize the continuum subdomain (Figure 1(b2)); and (2) couple the extracted pore 79 network and the discretized microporous subdomain (Figure 1(c)). We present each step in the following 80 subsections. 81

⁸² 2.1.1. Network extraction and discretization

We decompose the ternary image (Image #1) into two binary images (Figure 1): Image #2 containing only the resolved pores and solid, and Image #3 containing only microporous region (i.e., continuum) and

solid. Then, we extract the pore network from Image #2 using the watershed-segmentation-based SNOW 85 algorithm developed by Gostick (2017) (Figure 1(b1)), where we approximate the pore bodies as spheres and 86 pore throats as cylindrical tubes. The SNOW algorithm employs several techniques, including prefiltering the 87 distance map, removing peaks on saddles and plateaus, and merging nearby peaks, to mitigate the peaks and 88 the spurious local maxima resulting from over-segmentation of the standard watershed-segmentation-based 89 algorithms (Rabbani et al., 2014; Gostick, 2017). When applying the SNOW algorithm, a multidimensional 90 Gaussian blur filter with a sigma value of 0.35 is used to smooth the distance map. The radius of the 91 structuring element in the maximum filter is set to 4. The network extraction algorithm provides the 92 volume, coordination number, diameter, and surface area of each pore body, and the diameter, length, and 93 connectivity of each pore throat. Finally, we discretize the microporous subdomain using structure grids 94 (Figure 1(b2)). In the present work, the discretized numerical cells have a one-to-one correspondence to the 95 voxels of the 3D digital image. 96

97 2.1.2. Coupling the pore network and the microporous subdomain

To couple the extracted pore network and the discretized microporous subdomain, we identify the voxels 98 of the microporous subdomain residing at the boundary of each pore and subsequently link them to the 99 corresponding extracted pore body. Taking pore i as an example, the coupling process involves the following 100 morphological operations: (1) Dilate the watershed segment of pore i with a spherical structuring element of 101 one-voxel radius. (2) Identify the interface voxels in the overlapping region of the dilated watershed segment 102 of pore i and the microporous region surrounding this pore (we refer to these voxels as the interface voxels 103 for this pore). For example, voxels labeled m, n, k in Figure 1(f) are the interface voxels for pore i. (3) 104 Count the number of side faces of each interface voxel shared with pore i. This is done by dilating the 105 interface voxel and counting the number of dilated voxels overlapped with pore i. (4) Append to pore i the 106 information of its interface voxels, including the total number of interface voxels, the coordinates, and the 107 number of shared faces of each interface voxel. Using the above method, we identify the interface voxels 108 adjacent to each specific pore and finally couple the PNM and the continuum model. The extracted pore 109 network and the coupling between the two submodels (Figure 1(c-f)) are then passed onto a computational 110 framework to simulate flow and transport in two-scale digital images of pore structures. 111

112 2.2. Mathematical formulations

The hybrid pore-network-continuum modeling framework can be used to model a wide variety of problems involving fluid flow and transport in two-scale porous media. In the present study, we focus on illustrating three example problems: (1) steady-state incompressible single-phase flow, (2) solute transport, and (3) twocomponent gas transport in shale. In the following subsections, we present the mathematical formulations for the three examples. Because of the small spatial scales involved, we do not consider gravity.

118 2.2.1. Steady-state incompressible single-phase flow

For steady-state incompressible single-phase flow, the mass balance equation in a pore body i of the resolved pore subdomain can be written as

$$\sum_{j=1}^{N_i} g_{ij} \left(p_i - p_j \right) + \sum_{m=1}^{M_i} q_{i,m} = 0, \tag{1}$$

where "j" indicates the pore body adjacent to pore body i, and "m" indicates the grid cell in the continuum 121 microporous subdomain that is connected to pore body i. "ij" indicates the pore throat between pore bodies 122 i and j; "i, m" indicates the connection between pore body i and the connected continuum cell m. N_i is 123 the number of pore bodies connected to pore body i. p_i and p_j are the pressure in pore bodies i and j124 (Pa), respectively. $g_{ij} = \pi d_{ij}^4 / (128\mu_{ij}l_{ij})$ is the conductance between pore bodies *i* and *j* (m⁴ · s/kg), which 125 is derived from Hagen-Poiseuille law for a cylindrical tube, where d_{ij} is the diameter of the pore throat ij126 (m), μ_{ij} is the gas viscosity (Pa · s) which is determined using the method from Lee et al. (1966), and l_{ij} 127 is the length of the pore throat ij (m). M_i is the number of continuum cells connected to pore body i. 128 $q_{i,m} = k_{i,m} A_{i,m} (p_i - p_m) / (\mu_{i,m} l_{i,m})$ is the volumetric flux from pore body *i* to continuum cell *m* (m³/s), 129 where $A_{i,m}$ is the area of their interface (m²), p_m is the pressure in continuum cell m (Pa), $l_{i,m}$ is the 130 distance between pore body i and continuum cell m (m), and $k_{i,m}$ and $\mu_{i,m}$ are the permeability (m²) and 131 fluid viscosity (Pa \cdot s) at the interface between pore body *i* and continuum cell *m*, respectively. 132

In the continuum microporous subdomain, the mass balance equation for a continuum cell m can be written as

$$\int_{\Omega_m} \nabla \cdot \left(-\frac{k_m}{\mu_m} \nabla p_m \right) dV + \sum_{i=1}^{W_i} q_{m,i} = 0,$$
⁽²⁾

where Ω_m denotes the integral domain, i.e., the control domain of cell m. The subscript "m, i" indicates the connection between continuum cell m and the connected pore body i. k_m is the the permeability in continuum cell m (m²). μ_m is gas viscosity in continuum cell m (Pa · s) computed from Lee et al. (1966). W_i is the number of pore bodies connected to continuum cell m. $q_{m,i} = k_{i,m}A_{i,m} (p_m - p_i)/(\mu_{i,m}l_{i,m})$ is the volumetric flux from continuum cell m to the adjacent pore body i (m³/s).

140 2.2.2. Solute transport

This section presents the governing equations for single-phase solute transport processes accounting for advection and dispersion.

For a pore body i, the mass conservation equation of a nonactive solute can be given as

$$V_{i}\frac{\partial(C_{i})}{\partial t} + \sum_{j=1}^{N_{i}} \left(g_{ij} \left(p_{i} - p_{j} \right) C_{ij} + D_{ij} A_{ij} \frac{C_{i} - C_{j}}{l_{ij}} \right) + \sum_{m=1}^{M_{i}} \left(q_{i,m} C_{i,m} + D_{i,m} A_{i,m} \frac{C_{i} - C_{m}}{l_{i,m}} \right) = 0, \quad (3)$$

where V_i is the volume of pore body i (m³). C_i and C_j are the concentrations in pore bodies i and j(kg/m³). C_m is the concentration in continuum cell m (kg/m³). C_{ij} is the concentration in pore throat

ij (kg/m³), which is assumed to be the upstream concentration between pore bodies i and j. Similarly, 146 $C_{i,m}$ is the upstream concentration at the interface of pore body i and continuum cell m (kg/m³). $D_{ij} =$ 147 $D_0(1+\xi Pe^2)$, is the dispersion coefficient in pore throat $ij \ (m^2/s)$, where D_0 is the molecular diffusion 148 coefficient (m²/s). The constant ξ depends on the shape of the flow channel (Taylor, 1953); for cylindrical 149 tubes, $\xi = 0.02083$. $Pe = u_{ij}d_{ij}/(2D_0)$ is the local Peclet number (-), where $u_{ij} = g_{ij}(p_i - p_j)/A_{ij}$ is the 150 fluid velocity in the pore throat (m/s) and $A_{ij} = \pi d_{ij}^2/4$ is the cross-sectional area of the pore throat (m^2) . 151 $q_{i,m} = k_{i,m} A_{i,m} (p_i - p_m) / (\mu_{i,m} l_{i,m})$ is the fluid flux from pore body *i* to the connected continuum cell *m* 152 (m³/s). $D_{i,m}$ is the dispersion coefficient at the interface between pore body *i* and continuum cell *m* (m²/s). 153 For each continuum microporous cell m, the mass conservation equation is given by 154

$$V_m \frac{\partial(\phi_m C_m)}{\partial t} + \int_{\Omega_m} \nabla \cdot \left(-\frac{k_m \nabla p_m}{\mu_m} C_m + D_m \nabla C_m \right) dV + \sum_{i=1}^{W_i} \left(q_{m,i} C_{i,m} + D_{i,m} A_{i,m} \frac{C_m - C_i}{l_{i,m}} \right) = 0, \quad (4)$$

where V_m is the volume of continuum cell m (m³). ϕ_m is the porosity (-). D_m is the dispersion coefficient in continuum subdomain (m²/s). $q_{m,i} = k_{i,m}A_{i,m}(p_m - p_i)/(\mu_{i,m}l_{i,m})$ is the fluid flux from continuum cell m to pore body i (m³/s).

158 2.2.3. Two-component gas transport in shale

We present the governing equations for CO_2 and CH_4 in a shale rock sample with two-scale pore struc-159 tures. The properties of the bulk gas (i.e., a mixture of CH_4 and CO_2) are functions of both pressure 160 and composition (i.e., the mass fractions of CO_2 and CH_4). Because the majority of the pore space in the 161 continuum subdomain of shale rock is on the order of nanometers, we need to account for the interactions 162 between gas molecules and the solid surface (e.g., pore wall) as well as diffusion of the molecules adsorbed 163 at the solid surface in the continuum models (e.g., Guo et al., 2018). We neglect the adsorption at the solid 164 surfaces and surface diffusion in the resolved pore subdomain due to a relatively small specific surface area 165 of the pore bodies (compared to the continuum microporous subdomain). We need two governing equations 166 for both the pore network and continuum models. One is for the bulk gas, and the other is for one of the 167 components (in this study, CO_2). We select two primary variables, the pressure of bulk gas (p) and the CO_2 168 mass fraction (ω^{CO_2}). 169

The mass balance equations for the bulk gas and CO_2 in a pore body *i* can be written as

$$V_i \frac{\partial \left(\rho_i^b\right)}{\partial t} + \sum_{j=1}^{N_i} \mathcal{J}_{ij}^b + \sum_{m=1}^{M_i} \mathcal{J}_{i,m}^b = 0,$$
(5)

$$V_{i} \frac{\partial(\rho_{i}^{b} \omega_{i}^{CO_{2}})}{\partial t} + \sum_{j=1}^{N_{i}} \left(\mathcal{J}_{ij}^{CO_{2},adv} + \mathcal{J}_{ij}^{CO_{2},disp} \right) + \sum_{m=1}^{M_{i}} \mathcal{J}_{i,m}^{CO_{2}} = 0,$$
(6)

where the superscript "b" denotes bulk gas (i.e., the mixture of CO₂ or CH₄). The superscripts "adv" and "disp" denote advection and dispersion, respectively. ρ_i^b is the bulk gas density in pore body $i (\text{kg/m}^3)$. \mathcal{J}_{ij}^b ¹⁷³ is the mass flux of bulk gas from pore body *i* to pore body *j* (kg/s). $\mathcal{J}_{i,m}^{b}$ is the mass flux of bulk gas from ¹⁷⁴ pore body *i* to continuum cell *m* (kg/s). $\omega_{i}^{CO_{2}}$ is the mass fraction of CO₂ in pore body *i* (-). $\mathcal{J}_{ij}^{CO_{2},adv}$ and ¹⁷⁵ $\mathcal{J}_{ij}^{CO_{2},disp}$ are the advective and dispersive mass fluxes of CO₂ from pore body *i* to *j* (kg/s), respectively. ¹⁷⁶ $\mathcal{J}_{i,m}^{CO_{2}}$ is the mass flux of CO₂ from pore body *i* to continuum cell *m* (kg/s), which includes both advective ¹⁷⁷ and dispersive fluxes.

¹⁷⁸ The bulk gas density in both pores and continuum cells can be expressed as

$$\rho^b = \frac{pM^b}{ZRT},\tag{7}$$

where p is the pressure of bulk gas (Pa). $M^b = M^{CO_2} \varpi^{CO_2} + M^{CH_4} \varpi^{CH_4}$ is molar mass of the bulk gas 179 (kg/mol), where M^{CO_2} and M^{CH_4} are molar mass of CH₄ and CO₂ (kg/mol), respectively; ϖ^{CO_2} and ϖ^{CH_4} 180 $(\varpi^{CO_2} + \varpi^{CH_4} = 1 \text{ in the system})$ are the mole fraction of CH₄ and CO₂ (-), respectively. T is temperature 181 (K). R is the universal gas constant $(R = 8.314 \text{ kg} \cdot \text{m}^2/\text{s}^2/\text{K/mol})$. $Z = 0.702e^{-2.5T_r}p_r^2 - 5.524e^{-2.5T_r}p_r + 10^{-2.5T_r}p_r^2$ 182 $0.044T_r^2 - 0.164T_r + 1.15$ is the compressibility factor (-), where $T_r = T/T_{crt}$ is the reduced temperature (-) 183 and T_{crt} is the critical temperature (K); $p_r = p/p_{crt}$ is the reduced pressure (-) and p_{crt} is the critical pressure 184 (Pa) (Mahmoud, 2014). In our study, $T_{crt} = T_{crt}^{CO_2} \varpi^{CO_2} + T_{crt}^{CH_4} \varpi^{CH_4}$ and $p_{crt} = p_{crt}^{CO_2} \varpi^{CO_2} + p_{crt}^{CH_4} \varpi^{CH_4}$, 185 where $T_{crt}^{CO_2}$ and $T_{crt}^{CH_4}$ are the critical temperature of CO₂ and CH₄ (K), respectively; $p_{crt}^{CO_2}$ and $p_{crt}^{CH_4}$ are 186 the critical pressure of CO_2 and CH_4 (Pa), respectively. 187

Here we use a flux model that describes gas flux through cylindrical nanotubes (Beskok & Karniadakis, 189 1999) to model the flux of the bulk gas through pore throats. The mass flux of the bulk gas and CO₂ can 190 be written as the Hagen-Poiseuille flux q_{H-P} multiplied by a correction factor,

$$\mathcal{J}_{ij}^{b} = f(Kn)q_{H-P} = f(Kn)g_{ij} (p_i - p_j)\rho_{ij}^{b},$$
(8)

$$\mathcal{J}_{ij}^{CO_2,adv} = J_{ij}^b \omega_{ij}^{CO_2},\tag{9}$$

where $f(Kn) = (1 + \alpha Kn) (1 + 4Kn/(1 + 4Kn))$ is the correction factor (-). $\alpha = 2\alpha_0/(\pi \tan(\alpha_1 Kn^\beta))$, with $\alpha_0 = 64/15\pi$, $\alpha_1 = 4$, and $\beta = 0.4$. $Kn = \lambda/L$ is the Knudsen number (-), where $\lambda = \mu \sqrt{\pi ZRT}/(p\sqrt{2M})$ is the mean free path (m), and L is a characteristic length scale for gas flow (m). In the present study, L is assumed to be equal to the length of the pore throat ij.

¹⁹⁵ The dispersive mass flux of CO_2 can be expressed as

$$\mathcal{J}_{ij}^{CO_2,disp} = \rho_{ij}^b D_{ij}^{CO_2,disp} A_{ij} \frac{\omega_i^{CO_2} - \omega_j^{CO_2}}{l_{ij}},\tag{10}$$

where ρ_{ij}^b is the bulk gas density in the pore throat ij (kg/m³). $D_{ij}^{CO_2,disp}$ is the dispersion coefficient of CO₂ in the pore throat ij in the resolved pore subdomain (m²/s).

The mass fluxes of bulk gas and CO_2 between pore body *i* and continuum cell *m* are

$$\mathcal{J}_{i,m}^{b} = \rho_{i,m}^{b} A_{i,m} \frac{k_{i,m}}{\mu_{i,m}} \frac{p_{i} - p_{m}}{l_{i,m}},$$
(11)

$$\mathcal{J}_{i,m}^{CO_2} = J_{i,m}^b \omega_{i,m}^{CO_2} + \rho_{i,m}^b D_{i,m}^{CO_2,disp} A_{i,m} \frac{\omega_i^{CO_2} - \omega_m^{CO_2}}{l_{i,m}},\tag{12}$$

where $\rho_{i,m}^{b}$ is the bulk gas density at the interface between pore body *i* and continuum cell *m* (kg/m³). $D_{i,m}^{CO_2,disp}$ is the dispersion coefficient of CO₂ at the interface between between pore body *i* and continuum cell *m* (m²/s), which is approximated using the smaller dispersion coefficients at either side of the interface. In the continuum microporous subdomain, the mass balance equations for the bulk gas and CO₂ in a continuum cell *m* can be written as

$$V_m \frac{\partial \left(\phi_m \rho_m^b + \Gamma_m^{CO_2, ex} + \Gamma_m^{CH_4, ex}\right)}{\partial t} + \int_{\Omega_m} \nabla \cdot \left(J_m^b + J_m^{CO_2, sf} + J_m^{CH_4, sf}\right) dV + \sum_{i=1}^{W_i} \mathcal{J}_{m,i}^b = 0, \quad (13)$$

$$V_m \frac{\partial \left(\phi_m \rho_m^b \cdot \omega_m^{CO_2} + \Gamma_m^{CO_2, ex}\right)}{\partial t} + \int_{\Omega_m} \nabla \cdot \left(J_m^{CO_2, adv} + J_m^{CO_2, sf} + J_m^{CO_2, disp}\right) dV + \sum_{i=1}^{W_i} \mathcal{J}_{m,i}^{CO_2} = 0, \quad (14)$$

where "ex" indicates excess adsorption, and "sf" indicates surface diffusion. ρ_m^b is the bulk gas density in 204 continuum cell m (kg/m³). $\Gamma_m^{CO_2,ex}$ and $\Gamma_m^{CH_4,ex}$ are the excess adsorption of CO₂ and CH₄ per unit volume of 205 bulk porous media (kg/m³), respectively. J_m^b is the mass flux of bulk gas at the interfaces between continuum 206 cell m and adjacent cells (kg/s/m²). $J_m^{CO_2,sf}$ is the mass flux of CO₂ by surface diffusion (kg/s/m²). $J_m^{CH_4,sf}$ 207 is the mass flux of CH₄ (kg/s/m²) by surface diffusion. $\mathcal{J}_{m,i}^{b}$ is the mass flux of bulk gas from continuum 208 cell m to pore body i (kg/s). $\omega_m^{CO_2}$ is the mass fraction of CO₂ in continuum cell m (–). $J_m^{CO_2,adv}$ and 209 $J_m^{CO_2,disp}$ are the advective and dispersive mass fluxes of CO₂ at the interfaces between continuum cell m 210 and its adjacent cells (kg/s/m²), respectively. $\mathcal{J}_{m,i}^{CO_2}$ is the mass flux of CO₂ from continuum cell m to pore 211 body *i* (kg/s). Other variables and parameters have been defined earlier. Note that $\mathcal{J}_{m,i}^{b}$ and $\mathcal{J}_{m,i}^{CO_2}$ are only 212 non-zero for continuum cells at the interface between the resolved pore and microporous subdomains. 213

Adsorption of CO_2 and CH_4 in shales may be described by the Langmuir isotherm (e.g., Zhang et al., 215 2012; Heller & Zoback, 2014). For a two-component mixture of CO_2 and CH_4 , we use the extended multi-216 component Langmuir isotherm to represent the impact of competitive adsorption. The excess adsorption of 217 CO_2 and CH_4 can be expressed as

$$\Gamma_m^{CO_2,ex} = \Gamma_m^{CO_2,ads} \left(1 - \frac{\omega_m^{CO_2} \rho_m^b}{\omega_m^{CO_2,ads} \rho_m^{ads}} \right),\tag{15}$$

$$\Gamma_m^{CH_4,ex} = \Gamma_m^{CH_4,ads} \left(1 - \frac{\omega_m^{CH_4} \rho_m^b}{\omega_m^{CH_4,ads} \rho_m^{ads}} \right),\tag{16}$$

$$\Gamma_m^{CO_2,ads} = \Gamma_{m,max}^{CO_2} \frac{K_m^{CO_2} p_m \varpi_m^{CO_2}}{1 + K_m^{CO_2} p_m \varpi_m^{CO_2} + K_m^{CH_4} p_m \varpi_m^{CH_4}},$$
(17)

$$\Gamma_m^{CH_4,ads} = \Gamma_{m,max}^{CH_4} \frac{K_m^{CH_4} p_m \varpi_m^{CH_4}}{1 + K_m^{CO_2} p_m \varpi_m^{CO_2} + K_m^{CH_4} p_m \varpi_m^{CH_4}},$$
(18)

where "ads" indicates adsorption. $\Gamma_{m,max}^{CO_2}$ and $\Gamma_{m,max}^{CH_4}$ are the maximum adsorption of CO₂ and CH₄ (kg/m³), respectively. $K_m^{CO_2}$ and $K_m^{CH_4}$ are the Langmuir coefficients of CO₂ and CH₄ (m · s²/kg), respectively. $\omega_m^{CH_4} = 1 - \omega_m^{CO_2}$ is the mass fraction of CH₄ in continuum cell m (–). $\omega_m^{CO_2,ads} = \Gamma_m^{CO_2,ads} / (\Gamma_m^{CO_2,ads} + \Gamma_m^{CH_4,ads})$ and $\omega_m^{CH_4,ads} = \Gamma_m^{CH_4,ads} / (\Gamma_m^{CO_2,ads} + \Gamma_m^{CH_4,ads})$ are the mass fractions of CO₂ and CH₄ in the adsorbed phase (–), respectively. ρ_m^{ads} is the adsorbed gas density, which is assumed to be constant (1,000 kg/m³) (Zhang et al., 2020).

The mass fluxes of the bulk gas and CO_2 in the continuum cell m can be expressed as

$$J_m^b = -\rho_m f(Kn) \frac{k_m}{\mu_m^b} \nabla p_m, \tag{19}$$

$$J_m^{CO_2,adv} = J_m^b \omega_m^{CO_2}.$$
 (20)

The adsorbed gas can migrate along the pore wall, which causes the surface diffusion process (Ruthven, 1984; Medved and Cerny, 2011). The surface diffusive fluxes of the bulk gas and CO₂ can be expressed as

$$J_m^{CO_2,sf} = -D_m^{CO_2,sf} \nabla \Gamma_m^{CO_2,ads},\tag{21}$$

$$J_m^{CH_4,sf} = -D_m^{CH_4,sf} \nabla \Gamma_m^{CH_4,ads},$$
(22)

where $D_m^{CO_2,sf} = D_0^{CO_2,sf}/\tau$ and $D_m^{CH_4,sf} = D_0^{CH_4,sf}/\tau$ are the surface diffusion coefficients of CO₂ and CH₄ (m²/s) in the continuum microporous subdomain, where $D_0^{CO_2,sf}$ and $D_0^{CH_4,sf}$ are surface diffusion coefficients of CO₂ and CH₄ on flat solid surfaces (m²/s), respectively.

Similarly, we consider component dispersion in the continuum subdomain. The dispersive flux of CO_2 can be expressed as

$$J_m^{CO_2,disp} = -\phi_m \rho_m^b D_m^{CO_2,disp} \nabla \omega_m^{CO_2},$$
(23)

where $D_m^{CO_2,disp}$ is the dispersion coefficient of CO₂ in continuum cell m (m²/s).

The flux of the bulk gas and CO_2 at the interface between pore *i* and continuum cell *m* can be expressed as

$$\mathcal{J}_{m,i}^{b} = \rho_{i,m}^{b} A_{i,m} \frac{k_{i,m}}{\mu_{i,m}} \frac{p_m - p_i}{l_{i,m}},$$
(24)

$$\mathcal{J}_{m,i}^{CO_2} = \mathcal{J}_{m,i}^b \omega^{CO_2} + \phi_m \rho_{i,m}^b D_{i,m}^{CO_2, disp} A_{i,m} \frac{\omega_m - \omega_i}{l_{i,m}}.$$
(25)

Note that we do not consider the surface diffusion at the interface between pore i and continuum cell m.

236 2.3. Numerical methods

We apply a fully implicit numerical framework to solve the mathematical formulations of the h-PNMcontinuum models presented in Section 2.2. The continuum porous domains are discretized in space using the cell-centered finite volume method. The advection term in the solute transport equations (Eqs. 3 & 4) is discretized using a first-order upwinding scheme. The nonlinear equations resulting from the fully implicit discretization are solved using the Newton-Raphson iteration method. Our implementation of the numerical methods employs the automatic differentiation and gridding utilities in the MATLAB Reservoir Simulation Toolbox (Lie, 2019).

We conduct simulations using existing DNS models implemented in the open-source code OpenFOAM to evaluate the accuracy of the first two mathematical formulations of the h-PNM-continuum models. The DNS results for the steady-state single-phase flow are obtained using the simpleFoam solver based on the SIMPLE (Semi-Implicit Method for Pressure Linked Equations) algorithm. The single-phase solute transport DNS results are conducted using the scalarTransportFoam solver.

The initial and boundary conditions for both the h-PNM-continuum and DNS models of each specific test case are presented in Sections 3.1–3.3.

251 3. Results

We conduct numerical experiments to evaluate the h-PNM-continuum modeling framework for each of the mathematical formulations: (1) steady-state single-phase flow, (2) solute transport process, and (3) compositional single-phase flow in shale. The following three subsections present the details of the problem setup and simulation results.

256 3.1. Steady-state single-phase flow

We simulate steady-state single-phase flow in a 3D digital image of sintered glass beads and compare the simulations to DNS. To analyze the contribution of the continuum subdomain to the flow behavior in the entire domain, we consider two cases: (1) flow in the whole domain including both the resolved pore and the sub-resolution microporous subdomains, and (2) flow only in the resolved pore subdomain.

The physical size of the domain is 0.0025 m \times 0.0025 m \times 0.01 m in x, y, and z with 100 \times 100 \times 261 400 voxels, respectively. We apply the SNOW algorithm to decompose the solid regions within the digital 262 images, and then we label one out of five as the continuum microporous region, creating a ternary digital 263 image. The unresolved pore space in the continuum microporous subdomain accounts for 8.8% of the total 264 pore space in the ternary image. The continuum subdomain is assumed to be homogeneous and isotropic 265 with a porosity of $\phi_m = 0.2$ and a permeability of $k_m = 3$ Darcy (1 Darcy $\approx 9.869 \times 10^{-13} \text{ m}^2$). For both 266 the h-PNM-continuum and DNS models, the porosity of the resolved pore subdomain is 0.278 and the total 267 porosity of the entire domain is 0.304. The fluid is assumed incompressible and is described by Eqs. (1–2). 268 The initial pressure of the domain is set to 10 Pa. The inlet and outlet boundaries have a fixed pressure of 269

Table 1: Computed permeability of the entire domain and the resolved pore subdomain by the h-PNM-continuum model and DNS method.

Model	Permeability (m^2)	
h-PNM-continuum model	k^{pore}	6.15×10^{-11}
	k	6.61×10^{-11}
DNS method	k^{pore}	5.39×10^{-11}
	k	5.98×10^{-11}

 $p_{inlet} = 10$ Pa and $p_{outlet} = 0$, respectively, imposing a gradient pressure from the inlet to the outlet. The other four faces of the domain are closed and set as no flux boundaries. For the hybrid-PNM-continuum model, an absolute tolerance of 10^{-5} for the numerical residual is used as a convergence criterion for the nonlinear Newton-Raphson iterations. For the DNS, the absolute tolerance is 10^{-10} for the SIMPLE solver. The overall permeability of the entire domain can be computed using Darcy's law as

$$k = \frac{Q\mu L}{A \triangle p} \tag{26}$$

where Q is the volumetric flow rate at the inlet and outlet of the domain (m³/s) at steady state. $L = L_z$ is the length of the domain along the flow direction (m). $A = L_x \times L_y$ is the cross-sectional area of the domain m². $\Delta p = p_{inlet} - p_{outlet}$ is the pressure drop between inlet and outlet.

The computed overall permeability of the entire domain is 6.61×10^{-11} m² from the h-PNM-continuum 278 model and 5.98×10^{-11} m² from the DNS model (Table 1). The computed permeability of the resolved 279 pore subdomain is 6.15×10^{-11} m² and 5.39×10^{-11} m² from the PNM and DNS models, respectively. The 280 permeability calculated from the two models is generally in good agreement. Meanwhile, some discrepancies 281 between the two models are observed, which can be primarily attributed to the differences in the resolved pore 282 subdomain. The higher permeability from the h-PNM-continuum model may result from the simplification 283 of the extracted pore network from the digital images by approximating the pore space as cylindrical pore 284 throats and spherical pore bodies. This is a known issue of image-based pore-network modeling. Multiple 285 more advanced approaches have been developed in the literature to improve the accuracy of PNM for image-286 based pore-scale simulations, including Zhao et al. (2020), and Raeini et al. (2017, 2018). These enhanced 287 PNM approaches can be incorporated into our h-PNM-continuum framework to improve accuracy. 288

The spatial distributions of pressure from the h-PNM-continuum model and DNS at steady state are presented in Figure 2. The results show that the pressure fields of both the resolved pore subdomain and unresolved microporous subdomain from the h-PNM-continuum model agree well with the DNS results (columns 1 and 2), which further demonstrates the accuracy of the h-PNM-continuum model for simulating steady-state flow.



Figure 2: Simulated pressure distributions in the sintered glass beads by the h-PNM-continuum model (row 1) and DNS method (row 2) assuming steady-state. The plots on the left included sub-resolution microporous subdomain, while the plots on the right did not.

294 3.2. Solute transport

We conduct numerical experiments to simulate the transport of a nonreactive solute under steady-state water flow using the same digital pore-scale image from Section 3.1. To verify the developed model under different flow conditions, we conduct simulations under both low and high Peclet numbers (i.e., Pe = 0.1and 100) and compare the results with that of the DNS method. The Peclet number is defined as

$$Pe = \frac{L_c u}{\phi_e D_0},\tag{27}$$

where L_c is a characteristic length (m), which is set to L_z . $u = k \Delta p/(\mu L_z)$ is the average water flow velocity (m/s). ϕ_e is the effective porosity of the entire domain. D_0 is the molecular diffusion coefficient in the resolved pore subdomain and is set to $2.295 \times 10^{-9} \text{ m}^2/s$ (which is equal to the self-diffusion coefficient of neat water at 298.15 K).

The pressure drop between the inlet and the outlet Δp is set to 0.1057 Pa and 105.7 Pa to generate the low and high Peclect number scenarios. For solute transport, the initial concentration of the solute is zero. At time t = 0, continuous injection of the solute at a concentration of 10 (mol/L) is applied to the inlet. At the outlet boundary, the normal gradient of the solute concentration is set to zero.

³⁰⁷ For convenience, we define a dimensionless time as

$$T = \frac{Qt}{\phi_e L_x L_y L_z},\tag{28}$$

where the definition of Q, ϕ_e , L_x , L_y , and L_z are consistent with the above definitions.

In Figure 3, we present the 3D concentration distribution of the entire domain for snapshots in time from two models at low and high Peclet numbers (Pe = 0.1 vs. Pe = 100). The breakthrough curves (BTCs) of the entire domain for the two Peclet numbers are presented in Figure 4. The spatial concentration distribution and BTCs results show that the h-PNM-continuum model matches well with the DNS model for both the low and high Peclet numbers cases.



Figure 3: Comparison of the spatial distribution of the solute concentration in the sintered glass beads simulated by the h-PNM-continuum model (rows 1 & 3) and DNS method (rows 2 & 4) at low and high Peclet numbers (i.e., Pe = 0.1 and 100, respectively).

For the low Peclet number case, the solute transport is dominated by diffusion and dispersion. Under 314 this condition, the assumption of well-mixed solute in each pore body employed by the PNM is reasonable. 315 As a result, excellent agreement between the h-PNM-continuum model and the DNS method is observed for 316 both concentration distributions at different times (T = 0.08, 0.13, and 0.23) and the BTCs at the outlet. 317 For the high Peclet number case, advection becomes dominant. From the spatial distribution of the solute 318 concentration, we can see that the plume from the developed h-PNM-continuum model transports slightly 319 faster than the corresponding DNS method. This is because the pore network model assumes a well-mixed 320 condition in each pore body and it has a greater permeability in the resolved pore structure than that 321 computed by the DNS model, both of which lead to faster transport of the solute in the domain. However, 322 even for a high Peclet number regime of Pe=100, the h-PNM-continuum model still appears to provide 323 reliable results compared to direct numerical simulations. 324

325 3.3. Compositional single-phase flow

In the third set of numerical experiments, we simulate the injection of CO_2 into a shale sample reconstructed from images obtained using FIB-SEM. The sample is 250 nm \times 250 nm \times 1,000 nm in three dimensions, and the voxel size is 5 nm \times 5 nm \times 5 nm. Voxels in the digital 3D image are segmented into four constituents: resolved pore space, organic matter, clay, and granular minerals. The organic matter and clay regions are represented by the continuum (i.e., the unresolved microporous region), while the granular minerals are assumed non-porous solid impermeable to flow. We use the methods described in Section 2.1 to extract and couple the pore network and the continuum grid cells. We simulate two scenarios to examine the



Figure 4: Comparisons of outflow concentration simulated using the h-PNM-continuum model (red curves) and the DNS method (blue curves) at low and high Peclet numbers (Pe = 0.1 and 100, respectively).

contribution of the continuum subdomain to CH₄ production. One considers the entire domain, including the unresolved microporous regions and the other focuses on the resolved pore subdomain only.

For the numerical simulations, the h-PNM-continuum model solves the governing equations described in Section 2.2.3 (Eqs. 5, 6, 13 & 14). We assume that all of the pore space in the shale sample is initially occupied by CH₄ (i.e., $\omega^{CO_2} = 0$) at a gas pressure of 10 MPa and a temperature of 400 K. At the inlet boundary (where z = 0), we continuously inject CO₂ (i.e., $\omega^{CO_2} = 1$) at a fixed pressure of 10.1 MPa. At the outlet boundary (where z = 1,000 nm), the pressure is fixed to 10 MPa while the CO₂ mass fraction is set as zero-gradient (i.e., $\partial \omega^{CO_2} / \partial z|_{\text{outlet}} = 0$). Other parameters used in the simulations are presented in Table 2.

We define a characteristic time scale t_c and use it to define a dimensionless time $T = t/t_c$. t_c is the time 342 when half of the total CH_4 in the entire domain is displaced by CO_2 . The spatial distribution of the CO_2 343 mass fraction and cross-section-averaged mass fraction of CO₂ and CH₄ along the z-direction are presented 344 in Figure 5. The gas pressure in the domain reaches equilibrium relatively quickly (at approximately T =345 0.0012). However, the transport of the CO₂ and CH₄ components lasts much longer. It is interesting that 346 the transport of CO_2 is faster in the presence of the sub-resolution microporous subdomain (Figure 5). 347 This is because the microporous subdomain increases the connectivity of the resolved pores, resulting in 348 improved effective connectivity across the entire domain. Comparison of Figure 5 rows 1 & 2, we can see 349 that the isolated pores are connected by the sub-resolution microporous regions within the whole domain. 350 For real shale rocks, the connectivity of resolved pores is usually limited, while sub-resolution micropores in 351 organic matter and clay minerals have the potential to provide storage spaces and migration pathways for 352 hydrocarbons. 353

The displacement of CH_4 by the injected CO_2 can also be measured by the mass flow rate of the two components at the outlet boundary. In Figure 6(a), we observe a significantly advanced arrival of CO_2 when microporous subdomain is coupled with the pore network (simulated using the h-PNM-continuum

Parameter	Value
Permeability of the continuum subdomain (k_m)	0.5 nanoDarcy
Porosity of the continuum subdomain (ϕ_m)	0.1
Langmuir adsorption coefficient for CO ₂ $(K_m^{CO_2})$ ^[1]	$1\times 10^{-7}~{\rm m\cdot s^2/kg}$
Langmuir adsorption coefficient for CH_4 $(K_m^{CH_4})$ ^[1]	$1.7\times 10^{-7}~{\rm m\cdot s^2/kg}$
Maximum adsorption of CO ₂ $(\Gamma_{m,max}^{CO_2})^{[1]}$	45 kg/m^3
Maximum adsorption of CH ₄ $(\Gamma_{m,max}^{CH_4})^{[1]}$	5 kg/m^3
Surface diffusion coefficient for CO ₂ on flat solid surface $(D_0^{CO_2,sf})$ ^[2]	$3.135 \times 10^{-8} \text{ m}^2/\text{s}$
Surface diffusion coefficient for CH_4 on flat solid surface $(D_0^{CH_4,sf})$ ^[2]	$1.14 \times 10^{-8} \text{ m}^2/\text{s}$
Dispersion coefficient of CO ₂ in the continuum subdomain $(D_m^{CO_2,disp})$ ^[1]	$1\times 10^{-9}~{\rm m^2/s}$
Dispersion coefficient of CO ₂ in the pore throat $(D_{ij}^{CO_2,disp})$ ^[3]	$2\times 10^{-8}~{\rm m^2/s}$
Critical temperature of CO ₂ $(T_{crt}^{CO_2})$	304.13 K
Critical temperature of CH_4 $(T_{crt}^{CH_4})$	$190.564 { m K}$
Critical pressure of CO ₂ $(p_{crt}^{CO_2})$	7.3773×10^6 Pa
Critical pressure of CH ₄ $(p_{crt}^{CH_4})$	4.599×10^6 Pa
Tortuosity (τ)	5

Table 2: Parameters used for the simulations of CO_2 injection into a shale sample.

Note: [1] From Edwards et al. (2015); [2] $D_0^{CH_4,sf}$ is from Guo et al. (2018), while $D_0^{CO_2,sf}$ is scaled from $D_0^{CH_4,sf}$ using molecular weight; [3] From Guevara-Carrion et al. (2019).



Figure 5: Simulated spatial distributions of CO₂ mass fraction at T = 0.13, 0.33 and 0.84 during CO₂ injection with (Subplots a1-a3) and without (Subplots b1-b3) including the sub-resolution microporous subdomain. Subplots c1-c3 show the cross-sectional average mass fractions along the z axis for CO₂ and CH₄ with and without accounting for the sub-resolution microporous subdomain. The simulations were conducted using the h-PNM-continuum model.



Figure 6: (a) The simulated mass flow rate of CO_2 and CH_4 with and without accounting for the sub-resolution microporous subdomain at the outlet boundary. (b) Simulated cumulative CH_4 production with and without accounting for the sub-resolution microporous subdomain. The simulations were conducted using the h-PNM-continuum model.

model), which indicates a quick invasion of CO₂ because of the improved effective connectivity through the 357 entire domain. Besides, the mass flow rate of CH₄ in the h-PNM-continuum model is much higher than 358 that simulated by the PNM model and declines much slower than the latter. Meanwhile, as is shown in 359 Figure 6(b), the cumulative mass production of CH_4 from the whole domain is higher than the resolved 360 pore subdomain. It indicates that the sub-resolution microporous subdomain can continuously serve as a 361 significant source for CH_4 production—the CH_4 production can be as high as 19% of the whole domain. 362 Our modeling results delineate the significant contribution of the continuum subdomain to the overall gas 363 production. 364

365 4. Discussion

We have developed a new hybrid pore-network-continuum modeling framework (h-PNM-continuum) for 366 simulating flow and transport in 3D multiscale pore structures. Our numerical experiments, spanning from 367 steady-state single-phase flow, single-phase solute transport, and single-phase two-component gas transport, 368 demonstrate that the h-PNM-continuum model can accurately simulate flow and transport processes and the 369 mass transfer between the resolved pore and continuum subdomains, while being much more computationally 370 efficient than the DNS method. Below, we compare the h-PNM-continuum model with other two-scale pore 371 network models, comment on the computational efficiency and numerical accuracy of h-PNM-continuum 372 model, and discuss how the h-PNM-continuum modeling framework may be extended to model two-phase 373 flow in 3D multiscale pore structures. 374

375 4.1. Comparison with other "two-scale" models

We comment on how the h-PNM-continuum model relates to other "two-scale" network models reported in the literature (Bauer et al., 2012; Jiang et al., 2013; Mehmani & Prodanović, 2014; Bultreys et al., 2015; Prodanović et al., 2015), which we refer to as dual PNMs in the present study. The dual PNMs represent both the resolved pores and the sub-resolution microporous regions of a "two-scale" porous medium. The difference between the different dual PNMs lies in the specific approach for representing the microporous regions and their connectivity with the resolved pore space.

One common approach employed by the dual PNMs is to represent the sub-resolution microporous 382 regions by smaller-scale pore networks. The smaller-scale pore networks can be stochastically generated or 383 based on higher-resolution images that resolve the microporous regions. In Jiang et al. (2013), a small-scale 384 pore network is stochastically generated and then combined with the larger-scale pore network. The dual 385 PNM involves the following steps. (1) Two networks (i.e., the large-scale and small-scale networks) were 386 firstly extracted from micro-CT images at different length scales using the Pore Analysis Tools (Jiang et al., 387 2007; Jiang, 2008). Then, (2) the statistic information from each pore network is used to reconstruct a 388 stochastic network that is representative of the original one (Jiang et al., 2012). Finally, (3) the large-scale 389 and small-scale networks are integrated into a single two-scale pore network by connecting the two networks 390 using added small-scale bonds. Micropores were uniformly distributed in the remaining space excluding the 391 regions occupied by the large-scale network, and the spatial correlations were not considered. In addition, 392 the coupling of the two networks was based on a stochastic algorithm, which does not follow the specific 393 connectivities in the 3D digital image. For the dual PNM by Prodanović et al. (2015), the smaller-scale pore 394 network is constructed from digital images of microporous subdomains as follows: (1) Extract the large-scale 395 pore network (i.e., macro-network) from the 3D digital images; (2) Identify the microporous regions in the 396 images; (3) Rescale the small-scale pore network (i.e., micro-network) by a scaling factor to fit into the 397 microporous regions; (4) Finally, connect the small-scale network to the larger-scale network to obtain a 398 two-scale network model. In their study, a network element (e.g., a pore) in the network does not necessarily 300

⁴⁰⁰ correspond to a single pore space; rather, it can encapsulate the collective characteristics of a small-scale ⁴⁰¹ network within the microporous subregion. The dual PNMs discussed above can have a large number of ⁴⁰² network elements due to the presence of many sub-resolution small pores, which can lead to prohibitive ⁴⁰³ computational costs even for a relatively small shale rock sample on the order of a few micrometers in each ⁴⁰⁴ dimension.

Another approach that has been used by dual PNMs is to approximate the microporous regions as 405 effective "pore throats" without resolving the pore structures of the microporous regions. Bauer et al. 406 (2012) extracted the resolved pore network (i.e., coarse-scale pore network) from the digital images and 407 assigned some portion of the resolved pores to be connected by micropores in parallel to the resolved pore 408 throats. Bultreys et al. (2015) developed another image-based dual PNM, in which a resolved pore network 409 is extracted from micro-CT images, and micropores are allowed to act both in parallel and in serial to the 410 resolved pore network. To determine the connection between micropores and resolved pores, a special link 411 (i.e., referred to as "micro-link" by the authors) is added if two pores in the resolved pore network touch 412 the same microporous region. Otherwise, the micro-link does not exist, or it can be considered as solid. 413 While the dual PNMs proposed by Bauer et al. (2012) and Bultreys et al. (2015) significantly reduce the 414 computational cost relative to the dual PNMs discussed in the previous paragraph, they do not represent 415 the actual location of the microporous regions and their specific connections with resolved pores. 416

The h-PNM-continuum model developed in the present study treats the micropores as a continuum using the Darcy-scale formalism and explicitly represents the flow and transport processes in the resolved pores using a pore network model. In the hybrid framework, the multiple scales originate directly from flow physics occurring at different scales, for example, the gas transport in micropores vs. resolved pores as described in Section 2.2.3. This hybrid approach significantly reduces the computational cost by representing the resolved pore space using PNM and modeling the sub-resolution microporous subdomain at the continuum scale.

423 4.2. Computational efficiency and numerical accuracy

The h-PNM-continuum model significantly reduces the computational cost relative to a DNS model. To 424 evaluate the computational efficiency, we compare the number of numerical cells in the two models. The 425 numerical cell numbers for a synthetic porous medium (Sections 3.1 & 3.2) and a shale sample (Section 426 3.3) are presented in Table 3. The number of numerical cells for the h-PNM-continuum model is over 427 99.9% smaller than that for the DNS model across all three test cases. While h-PNM-continuum still has 428 a large number of numerical cells for the continuum subdomain (i.e., the same as that of the DNS model), 429 the governing equation there is much simpler relative to that in the resolved pore subdomain for the DNS 430 model. Numerical coarsening strategies may also be developed to reduce the number of numerical cells in 431 the continuum subdomains, given that the flow and transport dynamics are anticipated to be much slower 432 than that in the resolved pore subdomain. 433

While the h-PNM-continuum is much more computationally efficient than the DNS model, the approximation of representing the resolved pore space by a pore network introduces additional errors just like any

Model		Synthetic porous medium	Typical shale sample
Resolved pore subdomain	h-PNM-continuum	509	76
	DNS method	$1,\!110,\!198$	153,095
Continuum subdomain	h-PNM-continuum	535,981	345,905
	DNS method		

Table 3: Comparisons of the number of numerical cells used in the hybrid pore-network-continuum model and the DNS method.

other pore network models. In the present study, we employ a basic version of PNM to demonstrate the
image-based hybrid framework for coupling a PNM with a continuum scale model. More advanced PNMs reported in the literature have significantly improved their predictive capabilities (Mehmani & Tchelepi, 2017;
Raeini et al., 2017, 2018; Zhao et al., 2020). As discussed in Section 3.1, These enhanced PNM approaches
can be incorporated into our h-PNM-continuum framework to further improve the accuracy.

441 4.3. Extension of the hybrid framework for other applications

In addition to the examples demonstrated in the present study, the h-PNM-continuum modeling frame-442 work can also be potentially generalized to model more complex processes in two-scale porous media, such as 443 two-phase flow or reactive transport. Extending the h-PNM-continuum to model two-phase flow in two-scale 444 porous media can be particularly attractive given that they are computationally expensive when simulated 445 by DNS models. Here, we outline the conceptual steps for extending the h-PNM-continuum model for mod-446 eling two-phase flow in a two-scale porous medium. For the resolved pore subdomain, we may employ a 447 dynamic PNM or a quasi-static PNM to represent the two-phase flow, depending on the question of interest. 448 For instance, for the two-scale porous media where there is a significant contrast between the pore sizes of 449 the resolve pore subdomain and the sub-resolution continuum subdomain, the two-phase flow in the resolved 450 pore subdomain is expected to be much slower than that in the microporous region. In this case, it may be 451 sufficient to use a quasi-static PNM for the resolved pore subdomain. The quasi-static distribution of the two 452 fluid phases can be updated based on the fluxes at the interfaces between the resolved pore subdomain and 453 the continuum subdomain. For the continuum subdomain, we can employ the standard two-phase flow model 454 at the continuum scale. The two-phase flow properties of the microporous region (i.e., capillary pressure as 455 a function of phase saturation and relative permeability function) may be estimated from higher-resolution 456 images or pore-size distribution information (e.g., obtained by mercury intrusion or nitrogen adsorption). 457

458 5. Conclusion

We have developed an image-based hybrid pore-network-continuum (h-PNM-continuum) modeling framework for the flow and transport processes in two-scale porous media. The hybrid framework treats the unresolved microporous region as a continuum and models the flow and transport at the Darcy scale, while explicitly representing the flow and transport processes in the larger pores through a computationally efficient pore network model. The h-PNM-continuum modeling framework couples the pore network with the
microporous subdomain based on the original connectivity between the two subdomains from 3D digital
images.

We have validated the modeling framework by comparing it to DNS methods in a set of numerical 466 experiments including steady-state incompressible single-phase flow and solute transport. The results suggest 467 that the h-PNM-continuum model can accurately predict the overall flow and transport process and the mass 468 transfer between the resolved pores and the sub-resolution microporous subdomains while being much more 469 computationally efficient than the DNS methods. We have then demonstrated the applicability of the hybrid 470 model for simulating two-component gas transport in shale rocks using the displacement of CH_4 by CO_2 as 471 an example. The simulations illustrate the significant contribution of the continuum subdomain to the overall 472 gas production. While the present study focuses on single-phase flow and transport, the h-PNM-continuum 473 modeling framework may also be further extended to simulate more complex processes that involve multi-474 phase flow, reactive transport, and non-Newtonian displacement. If desired, more advanced PNM approaches 475 can also be incorporated into the h-PNM-continuum modeling framework to further improve the accuracy 476 of representing flow and transport processes in the resolved pore subdomain. 477

List of variables

Name De	efinition
Subscript i or j Po	ore body labeled i or j .
Subscript <i>m</i> Th	he grid cell in the continuum subdomain.
Subscript <i>ij</i> Po	ore throat ij between pore bodies i and j .
Subscript i, m The second s	he connection between pore body i and the connected continuum cell m .
Subscript m, i The second s	he connection between continuum cell m and the connected pore body i .
Superscript b Bu	ulk gas (i.e., the mixture of CO_2 or CH_4).
Superscript <i>ads</i> Ac	dsorption.
Superscript ex Ex	xcess adsorption.
Superscript adv Ac	dvection.
Superscript <i>disp</i> Di	ispersion.
Superscript sf Su	urface diffusion.
A_{ij} The second se	he cross-sectional area of the pore throat ij (m ²).
$A_{i,m}$ An	rea of the interface between pore body i and continuum cell m (m ²).
C_i Co	oncentration in pore body i (kg/m ³).
C _{ij} Co	oncentration in pore throat ij (kg/m ³).
C_j Co	oncentration in pore body j (kg/m ³).

Name	Definition
D_0	The molecular diffusion coefficient (m^2/s) .
$D_0^{CO_2,sf}$	Surface diffusion coefficient of CO_2 on flat solid surface (m^2/s) .
$D_0^{CH_4,sf}$	Surface diffusion coefficient of CH_4 on flat solid surface (m^2/s) .
$D_m^{CH_4,sf}$	Surface diffusion coefficient of CH_4 in the continuum subdomain (m ² /s).
$D_m^{CO_2,sf}$	Surface diffusion coefficient of CO_2 in the continuum subdomain (m ² /s).
$D_{ij}^{CO_2,disp}$	The dispersion coefficient of CO_2 in the pore throat ij (m ² /s).
$D_m^{CO_2,disp}$	The dispersion coefficient of CO_2 in the continuum subdomain (m^2/s) .
$D_{i,m}^{CO_2,disp}$	The molecular diffusion coefficient of CO_2 at the interface between pore body
	i and continuum cell m (m ² /s).
d_{ij}	Diameter of the pore throat ij (m).
ϕ_m	Porosity (–).
f(Kn)	The correction factor $(-)$.
$\Gamma_m^{CH_4,ex}$	Excess adsorption of CH_4 per unit volume of porous material (kg/m ³).
$\Gamma_m^{CO_2,ex}$	Excess adsorption of CO_2 per unit volume of porous material (kg/m ³).
$\Gamma^{CH_4}_{m,max}$	Maximum adsorption of CH_4 (kg/m ³).
$\Gamma^{CO_2}_{m,max}$	Maximum adsorption of CO_2 (kg/m ³).
g_{ij}	The conductance between pore bodies i and j (m ⁴ · s/kg).
\mathcal{J}^b_{ij}	Mass flux of bulk gas from pore body i to pore body j (kg/s).
$\mathcal{J}^b_{i,m}$	Mass flux of bulk gas from pore body i to continuum cell m (kg/s).
$\mathcal{J}^b_{m,i}$	Mass flux of bulk gas from continuum cell m to pore body i (kg/s).
$\mathcal{J}_{ij}^{CO_2,adv}$	Advective mass flux of CO_2 from pore body <i>i</i> to pore body <i>j</i> (kg/s).
$\mathcal{J}_{ij}^{CO_2,disp}$	Dispersive mass flux of CO_2 from pore body <i>i</i> to pore body <i>j</i> (kg/s).
$\mathcal{J}^{CO_2}_{i,m}$	Mass flux of CO_2 from pore body <i>i</i> to continuum cell <i>m</i> (kg/s).
$\mathcal{J}_{m,i}^{CO_2}$	Mass flux of CO_2 from continuum cell m to pore body i (kg/s).
J_m^b	Mass flux of bulk gas at the interfaces between continuum cell m and adjacent
	cells $(kg/s/m^2)$.
$J_m^{CO_2,adv}$	Advective mass fluxes of CO_2 at the interfaces between continuum cell m and
	adjacent cells $(kg/s/m^2)$.
$J_m^{CO_2,disp}$	Dispersive mass fluxes of CO_2 at the interfaces between continuum cell m and
	adjacent cells $(kg/s/m^2)$.
$J_m^{CH_4,sf}$	Mass flux of CH_4 by surface diffusion (kg/s/m ²).
$J_m^{CO_2,sf}$	Mass flux of CO_2 by surface diffusion (kg/s/m ²).
$K_m^{CH_4}$	The Langmuir coefficient of CH_4 (m \cdot s ² /kg).
$K_m^{CO_2}$	The Langmuir coefficient of CO_2 (m \cdot s ² /kg).
k_m	Permeability (m^2) .

Definition
The mean free path (m).
A characteristic length scale for gas flow (m).
Length of the pore throat ij (m).
Distance between pore body i and continuum cell m (m).
Fluid viscosity $(Pa \cdot s)$.
Number of continuum cells connected to pore body i (–).
Molar mass of the bulk gas (kg/mol).
Molar mass of CH_4 (kg/mol).
Molar mass of CO_2 (kg/mol).
Number of pore bodies connected to pore body i (–).
The domain of integration.
Mass fraction of CH_4 (–).
Mass fraction of CO_2 (–).
Mass fraction of CH_4 in the adsorbed phase (-).
Mass fraction of CO_2 in the adsorbed phase (-).
Pressure (Pa).
Local Peclet number (–).
The reduced pressure $(-)$.
The critical pressure (Pa).
The critical pressure of CO_2 (Pa).
The critical pressure of CH_4 (Pa).
Mole fraction of CO_2 .
Mole fraction of CH_4 .
The volumetric flux between pore body i and continuum cell m (m ³ /s).
The volumetric flux between continuum cell m and pore body i (m ³ /s).
Adsorbed gas density kg/m^3).
Bulk gas density (kg/m^3) .
The universal gas constant $(kg \cdot m^2/s^2/K/mol)$.
Temperature (K).
The reduced temperature $(-)$.
The critical temperature (K).
The critical temperature of CO_2 (K).
The critical temperature of CH_4 (K).
Volume of continuum cell (m^3) .
Pore volume of pore body i (m ³).

Name	Definition
W_i	Number of pore bodies connected to continuum cell m (–).
ξ	A constant for dispersion in a cylindrical tube.
Ζ	The compressibility factor (–).

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