

1 Microbial indicators and detection of Cu-sulfide ore  
2 mineralization

3 **Bianca P. Iulianella Phillips**<sup>1,2</sup>, **Rachel L. Simister**<sup>3</sup>, **Shane D. Rich**<sup>1,2</sup>, **Craig J.R. Hart**<sup>1,2</sup>,  
4 **Peter A. Winterburn**<sup>1,2</sup>†, and **Sean A. Crowe**<sup>2,3</sup>\*

5 <sup>1</sup>*MDRU-Mineral Deposit Research Unit, University of British Columbia, Vancouver, BC,*  
6 *Canada, V6T 1Z4*

7 <sup>2</sup>*Department of Earth, Ocean, and Atmospheric Sciences, University of British Columbia,*  
8 *Vancouver, BC, Canada, V6T 1Z4*

9 <sup>3</sup>*Department of Microbiology and Immunology, University of British Columbia, Vancouver, BC,*  
10 *Canada, V6T 1Z3*

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12 † Deceased

13 \* Corresponding author: [sean.crowe@ubc.ca](mailto:sean.crowe@ubc.ca)

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29 **ABSTRACT**

30 Rapid electrification of society is placing unprecedented demand on critical mineral and metal  
31 resources. New strategies and technologies are thus needed to promote mineral discovery in  
32 regions where deposits are likely buried deep under soil and glacial till. We show that microbial  
33 communities from different soil types change in composition in response to amendment with  
34 copper. We also show that soil microbial community fingerprinting can detect buried sulfide  
35 mineralization through thick successions of Quaternary surface cover. Indicator species  
36 abundances, indeed, better resolve the known surface projection of mineralization than  
37 geochemical analyses. Therefore, our results indicate that porphyry-style mineralization is  
38 discernible in covered terrains through soil microbial community fingerprints.

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40

## 41 INTRODUCTION

42 Global demand for mineral resources – in particular, growth in green energy alternatives,  
43 including electric vehicles and batteries – is placing increasing pressure on the mining industry to  
44 supply metals such as copper, nickel, and lithium. At the same time, existing mineral deposits are  
45 being exhausted, while the frequency of new deposit discovery is declining (Schodde, 2017).  
46 This is forcing mineral exploration efforts to increasingly focus on concealed targets (Winterburn  
47 et al., 2020), requiring development of tools that detect unexposed mineralization. Whereas  
48 biological approaches to exploration have received minimal attention and uptake in the past,  
49 emerging DNA sequencing technologies offer new opportunities to harness the power of  
50 genomics across the resource sector with strong potential to improve mineral resource discovery  
51 (Simister et al., 2023).

52  
53 Through their growth and metabolism, microorganisms catalyze reactions that drive  
54 fluxes of matter and energy from nano- to global-scales (Newman and Banfield, 2002; Falkowski  
55 et al., 2008; Fierer, 2017). These reactions can promote precipitation and dissolution of minerals  
56 and thus microbial metabolism influences, and is influenced by, the distribution of minerals in  
57 the surface and sub-surface. Soil microorganisms are highly sensitive, detecting and responding  
58 to variability in the physicochemical properties of their habitats. In a single gram of soil there  
59 can be up to tens of thousands of microbial species, each with hundreds to thousands of genes,  
60 that collectively sense and interact with their surroundings (Torsvik and Øvreås, 2002; Fierer,  
61 2017). Mineral deposits in the subsurface likely represent an acute example of this, whereby ore  
62 minerals and constituent elements can be elevated orders of magnitude above background  
63 concentrations (Kelley et al., 2006). As such, microbial community compositions likely reflect

64 variability in the mineralogical composition of soils based on their proximity to, and interaction  
65 with, buried ore minerals (Simister et al., 2023). Furthermore, given the sensitivity of  
66 microorganisms to their surroundings, anomalies in microbial community compositions may  
67 develop, even if differences in physicochemical properties of the host soils are not resolvable  
68 through geochemical analyses.

69

70       Microbial community fingerprinting effectively delineates lithologically controlled  
71 mineralization, but capacity for resolving disseminated and structural ores remains untested. For  
72 example, diamondiferous kimberlites concealed by glacial till can be detected through soil  
73 microbial community anomalies that develop over kimberlite pipes, even where geochemical  
74 signals have been glacially dispersed (Simister et al., 2023). It remains unknown, however,  
75 whether this phenomenon is more broadly extensible, and if other ore types can be detected  
76 through similar applications of DNA sequencing. In this study, we thus tested the application of  
77 soil microbial community fingerprinting to the discovery of porphyry copper deposits (PCDs)  
78 which, when buried, are difficult to locate by traditional methods (Heberlein and Samson, 2010;  
79 Winterburn et al., 2020). We show that PCDs concealed by 10s of meters of soil and glacial  
80 overburden can be detected through soil microbial community analyses.

81

## 82 **METHODS**

### 83 **Incubation Experiments**

84       We conducted two incubation experiments with background soils from the Deerhorn PCD  
85 (British Columbia (BC)), and the sub-arctic tundra (Northwest Territories (NWT)). Amendments  
86 of copper sulfate ( $\text{CuSO}_4$ ) or chalcopyrite ( $\text{CuFeS}_2$ ) were made to aseptically dispensed soils at

87 200 ppm (ambient) or 600 ppm (high) (final concentration in soil), depending on the amendment.  
88 Such concentrations of copper and other ore-related elements reflect values typically detected  
89 (ambient) or those considered anomalous (high) in soil geochemical surveys above PCDs. The  
90 soils were incubated for 35 days (BC) and 85 days (NWT).

91

## 92 **Field Area and Survey Design**

93 The Deerhorn Cu-Au calc-alkaline PCD, BC, Canada (Figure 1b), hosts chalcopyrite  
94 ( $\text{CuFeS}_2$ ), enargite ( $\text{Cu}_3\text{AsS}_4$ ), and minimal bornite ( $\text{Cu}_5\text{FeS}_4$ ) mineralization within monzonite  
95 intrusions, and country rock consists of andesite breccias and thick beds of volcanic derived  
96 sandstones (del Real et al., 2017). The mineral resource is delineated by a 0.2 g/t Au equivalent  
97 grade shell, including most rocks that are  $>0.22\%$  Cu, with a surface area expression diameter of  
98  $\sim 500$  m at the bedrock-till interface (Figure 1a) (Sherlock, R., Blackwell, J., and Skinner, T., 2013).  
99 The Deerhorn area was glaciated; the PCD is variably covered by soils, organic deposits, and  
100 glacial materials (mapped based on classification schemes: (Howes et al., 1988; Taylor and  
101 Eggleton, 2001)), which cover mineralization (10 – 60 m) and host rocks (up to 169 m). 150 soil  
102 samples (B-horizon) were collected in three 2.5 km transects (50 m spacing; 200 m fence-lines)  
103 across mineralization and background, aligned with recent glaciation (Figure 1a).

104

## 105 **Geochemical and Geobiological Analyses**

106 Element concentrations were determined through inductively coupled plasma mass  
107 spectrometry (ICP-MS) following aqua-regia digests (Table S1 in the Supplemental Material<sup>1</sup>),  
108 with physicochemical variables (e.g., pH) measured *in situ*. Microbial community fingerprints of  
109 the same Deerhorn soils and soils from incubation experiments were determined via amplicon

110 (515f and 806r) (Caporaso et al., 2011; Apprill et al., 2015) sequencing of the 16S rRNA gene  
111 (V4) using the Illumina MiSeq. Sequences were analyzed – including richness (Table S2 in the  
112 Supplemental Material<sup>1</sup>) – using QIIME2 (Bolyen et al., 2019) and the SILVA taxonomic  
113 reference database (release 138). Differential abundance algorithms, ANCOM (analysis of  
114 composition of microbes) (Mandal et al., 2015) and LEfSe (linear discriminant analysis effect size)  
115 (Segata et al., 2011), were implemented in QIIME2 and Mothur (Schloss et al., 2009), respectively  
116 (Table S3 in the Supplemental Material<sup>1</sup>).

117

## 118 **MICROBIAL COMMUNITY RESPONSES TO COPPER AMENDMENTS**

119 Soil amendments with copper sulfate and chalcopyrite elicit changes in both microbial  
120 community structure and composition. Differences in species richness are more apparent in  
121 copper amended tundra soils (Figure 3) than Deerhorn soils (Figure 2). Richness metrics, such as  
122 the number of observed species (defined by amplicon sequence variants (ASVs)), are similar in  
123 Deerhorn baseline soils ( $1090 \pm 403$ ) and those amended with copper sulfate or chalcopyrite  
124 ( $918 \pm 137$ ) (Figure 2b). Conversely, there is a noticeable reduction in species richness in tundra  
125 soils amended with chalcopyrite ( $513 \pm 40$  and  $290 \pm 113$ , respectively; Figure 3b). A richness  
126 reduction in tundra soils may reflect growth and/or decay of taxa depending on sensitivity to the  
127 amendments. The dominant phyla in each soil differed in abundance, with Acidobacteriota,  
128 Proteobacteria, and Verrucomicrobiota dominating in Deerhorn soils (68% avg.), and  
129 Proteobacteria and Actinobacteriota in tundra soils (74% avg.). Both soil compositions are  
130 typical of soils globally (Fierer, 2017), with tundra and Deerhorn soils exhibiting similar  
131 composition to other northern (Johnston et al., 2019; Hale et al., 2019; Frank-Fahle et al., 2014)  
132 and temperate soils (Yang et al., 2022; Kaiser et al., 2016), respectively. The tundra soils had

133 more pronounced shifts in community composition with copper amendment that could be  
134 resolved at the phylum-level, with an increase in relative abundance of Proteobacteria (46% to  
135 69%) and a decrease in Acidobacteriota (5% to 3%), Bacteroidota (6% to 2%), and  
136 Verrucimicrobiota (6% to 2%) (Figure 3a), relative to baseline controls. Deerhorn soils remained  
137 largely unchanged at the phylum level—with the notable exception of minor shifts in  
138 Acidobacteriota (31% baseline; 24% copper amendment) and Actinobacteriota (5% baseline; 8%  
139 copper amendment) (Figure 2a). Tundra soils are more sensitive to copper amendment than the  
140 Deerhorn soils possibly due to lower overall species richness in the baseline soils, and thus  
141 differences in structure and composition induced by copper amendments are more readily  
142 resolved. Tundra soils also contain lower baseline concentrations of copper than Deerhorn soils  
143 (Table S4 in the Supplemental Material<sup>1</sup>), and the tundra soil microbial community may thus be  
144 more sensitive to copper amendment. Microbial community responses to copper amendment  
145 were further assessed through hierarchical clustering analyses. Baseline control soils and copper  
146 amended soils largely clustered independently (Figure 2d and Figure 3d). The coherent  
147 clustering of copper amended soils reveals a consistent response to copper amendment with  
148 changes in composition that are statistically well-resolved.

149

150 Multiple species-level differences in relative abundance develop in response to copper  
151 amendments. For example, both the Deerhorn and tundra soil incubations experienced  
152 enrichments in taxa belonging to the candidate phylum AD3 and the orders Xanthomonadales,  
153 Ktedonobacterales, and Micrococcales, and depletion in members of the candidate phylum WPS-  
154 2, and orders Acidobacteriales and Diplorickettsiales (Figure 2c and Figure 3c). In each  
155 experiment, members of the genus *Rhodanobacter* are particularly sensitive to elevated copper

156 concentrations (Figure 2c and Figure 3c). Members of this genus indeed have conspicuously high  
157 abundances in environments with elevated copper and other metals (Green et al., 2012; Cho et  
158 al., 2017; Carlson et al., 2018). Despite greater changes to the sub-arctic tundra soils in response  
159 to copper amendment than Deerhorn soils, species-level differences were detectable in both.  
160 These species thus represent naturally occurring indicators that respond to copper exposure.

161

## 162 **DELINEATION OF A CONCEALED PCD**

163       Microbial community composition in soils that overlie mineralization at the Deerhorn  
164 PCD are indistinguishable, at the phylum-level, from those in background soils. These phylum-  
165 level compositions are also similar overall, to those used in the incubation experiments (Figure  
166 2a and Figure 4a). Across all Deerhorn soils, Acidobacteriota (28%), Proteobacteria (25%), and  
167 Verrucomicrobiota (16%) were the most abundant (Figure 3a)—typical of temperate soils around  
168 the world (Janssen, 2006; Fierer, 2017; Delgado-Baquerizo et al., 2018). Observed species  
169 richness ( $1034 \pm 454$ ; Figure 4b) is also typical of similar such soils (Carini et al., 2016;  
170 Thompson et al., 2017), with no appreciable differences between soils that overlie the surface  
171 expression of mineralization and background (Figure 4b). It thus appears that species richness  
172 and phylum-level community compositions are largely insensitive to the buried Deerhorn PCD.

173

174       Differential abundance – or indicator species – analyses reveal a number of species-level  
175 enrichments in soils above mineralization. Given the gradational and disseminated nature of  
176 porphyry copper mineralization and grade thresholds, we conducted indicator species analyses of  
177 soils at Deerhorn in multiple tiers, with soils that are considered in the mineralized zone varying  
178 at 5-10 m increments from the known surface projection of mineralization (Figure 4g). In this

179 range – with minimum and maximum distances from mineralization in rock set at 0 m and 60 m,  
180 respectively – indicator species can be defined in the 0 m to 15 m distances from the surface  
181 expression, with the number of indicator species declining between 20 m and 60 m (Figure 4d).  
182 Indeed, there are no indicator species found through ANCOM beyond 15 m and the number of  
183 LEfSe indicator species drops by an average of 42% (Figure 4d). Only the indicator species in 0  
184 m – 15 m distances from the surface projection of mineralization were further considered in this  
185 work.

186  
187       Microbial community fingerprinting delineates the surface expression of mineralization  
188 through the spatial distribution of anomalies that emerge at the species level. These indicator  
189 species were derived from both differential abundance analyses at Deerhorn (0 m – 15 m) and  
190 copper amendment incubation experiments. These indicator species exhibit significantly (LDA  
191 score > 2 (LEfSe); W score rejecting null hypothesis (ANCOM)) higher relative abundances in  
192 soils above mineralization, and together (normalized sum), reveal strong anomalies at the surface  
193 (Figure 4d). Of the indicator species that delineate mineralization, 29 originate from the  
194 incubation experiments and 66 originate from the field-based observations supporting the idea  
195 that indicators identified through laboratory experiments are effective vectors in the field.  
196 Compared to indicator/pathfinder elements associated with sulfide mineralization, the microbial  
197 anomalies more accurately resolve the buried PCD, even where the glacial till cover is thick (>  
198 25 m) (Figures 4c and 4e).

199

200 **CONCLUSIONS AND IMPLICATIONS FOR MINERAL EXPLORATION**

201           We've shown that anomalies in soil microbial community composition develop in  
202 proximity to the surface expression of a PCD. Indicator species derived from incubation  
203 experiments and field-based analyses of soils above a PCD resolve the surface expression of the  
204 buried mineral deposit even where inorganic geochemical signals are weak. Given that 60% of  
205 global copper is extracted from PCDs (Tabelin et al., 2021), microbial community fingerprinting  
206 may help in the discovery of new copper deposits improving copper and associated critical metal  
207 (e.g., Bi, Te) supply chains. Furthermore, our results extend the potential power of DNA  
208 sequencing and microbial community fingerprinting of surface soils to a wider range of mineral  
209 resources, suggesting the approach may be extensible to other commodities (e.g., Ni, Zn, Pb) that  
210 are critical to the energy transition.

211

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219

220 <sup>1</sup>Supplemental Material. Supplemental methods, Tables S1-S4. Please visit  
221 <https://doi.org/10.1130/XXXX> to access the supplemental material, and contact  
222 [editing@geosociety.org](mailto:editing@geosociety.org) with any questions.

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308

309 **Figure 1. Deerhorn location and survey design. (A)** Soil sample locations (pink bubbles) and  
310 surficial materials map. Mineralization defined by >0.2 g/t Au equivalent ore (black line). **(B)**  
311 Deerhorn Cu-Au PCD location and BC geology (BCGS Open File 2017-8, 9p, access 2019-12-  
312 19 (Cui et al., 2017)).

313

314 **Figure 2. Microbial community composition and richness in Deerhorn soil incubations. (A)**  
315 Phylum-level microbial community composition (percentage of total 16S rRNA gene sequence  
316 reads) for each soil sample. **(B)** Number of observed species (ASVs), grouped as control (beige),  
317 or copper amended soils (blue). **(C)** Species-level (ASV) shifts in relative abundance across  
318 treatments over time. **(D)** Hierarchical relationships among soils based on Euclidean distance of  
319 ASV abundances, obtained with the unweighted pair group method with arithmetic mean  
320 (UPGMA) clustering algorithm. Node labels indicate the timepoint and treatment.

321

322 **Figure 3. Microbial community composition and richness in tundra soil incubations. (A)**  
323 Phylum-level microbial community composition (percentage of total 16S rRNA gene sequence  
324 reads) for each soil sample. **(B)** Number of observed species (ASVs), grouped as control (beige),  
325 or copper amended soils (blue). **(C)** Species-level (ASV) shifts in relative abundance across  
326 treatments over time (error bars = standard deviation). **(D)** Hierarchical relationships among soils  
327 based on Euclidean distance of ASV abundances, obtained with the unweighted pair group  
328 method with arithmetic mean (UPGMA) clustering algorithm. Node labels indicate the timepoint  
329 and treatment.

330

331 **Figure 4. Soil microbial community composition, richness, and microbial and geochemical**  
332 **anomalies above the Deerhorn PCD. (A)** Phylum-level microbial community composition  
333 (expressed as a percentage of total reads) for each soil sample. **(B)** Number of observed species  
334 (ASVs), grouped by origin: “above background” (beige) or “above mineralization” (blue). **(C)**  
335 Response ratio of indicator/pathfinder elements vs. indicator species from copper amendment  
336 incubation experiments and Deerhorn soil analyses (average “above mineralization” over the  
337 average “above background” relative to an equivalent ratio of 1, expressed in %). **(D)** The  
338 number of indicator species from multi-tiered differential abundance analyses (ANCOM and  
339 LEfSe) – soils “above mineralization” or “above background” for each tier are determined by  
340 soil sample proximity to projected mineralization. Anomaly maps at the Deerhorn PCD, showing  
341 the normalized sum of all indicator species **(E)** and Cu, Mo, K, and As, **(F)** with each element  
342 normalized to  $C_{\text{organic}}$  (abundant in some areas of the field site). Individual indicator species **(E)**  
343 or elements **(F)** are normalized to the mean prior to summation, with anomaly intervals based on

344 probability plots. **(G)** Gradational distances from mineralization used for differential abundance  
345 analyses.

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347

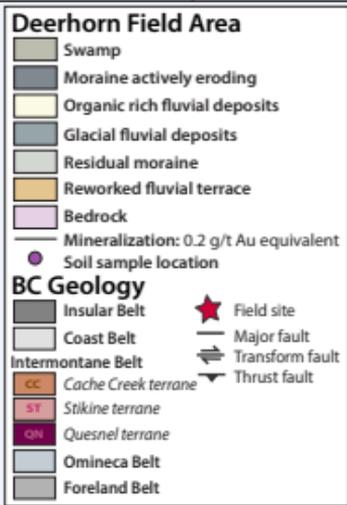
348

# Iulianella Phillips et al. **Figure 1** - MS#

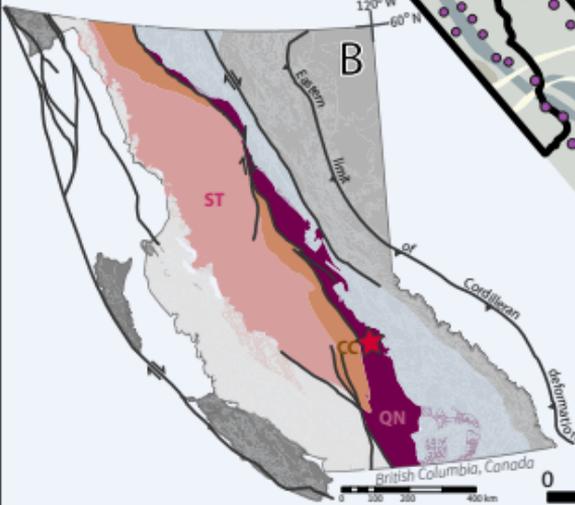
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121°21'04.88" W

121°21'05.96" W



Most recent direction of glacial movement  
305°



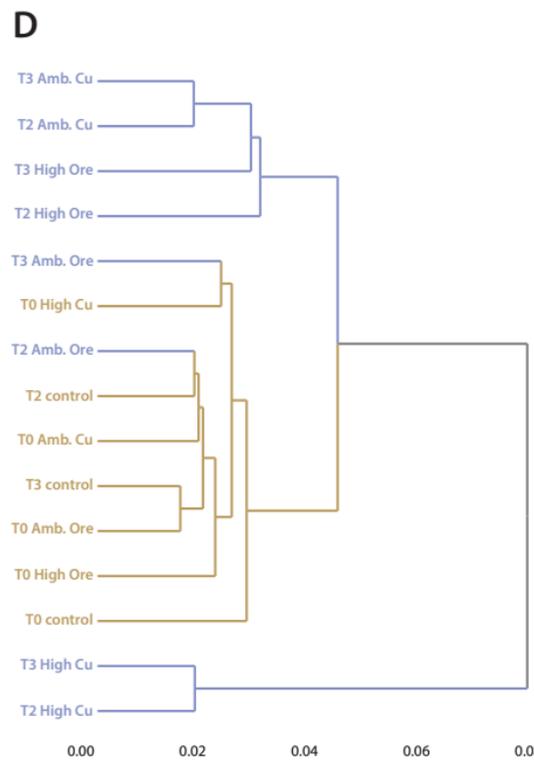
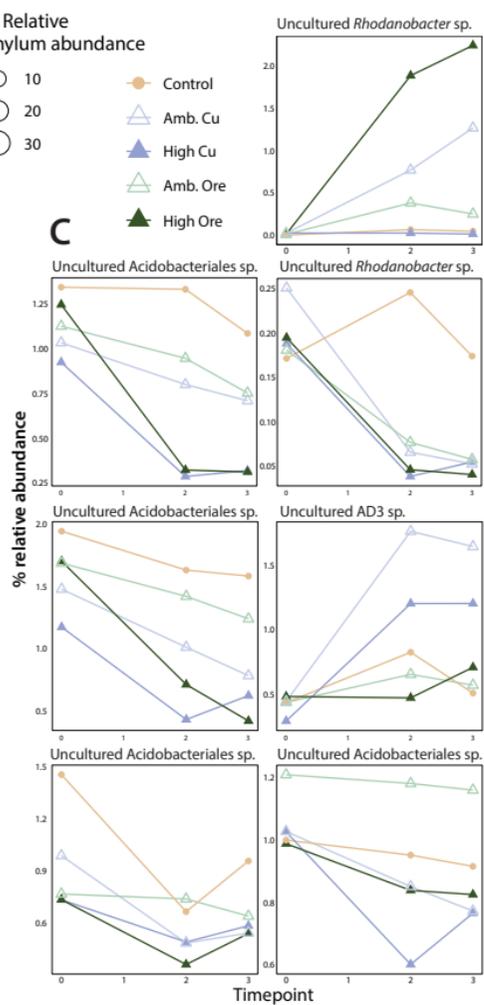
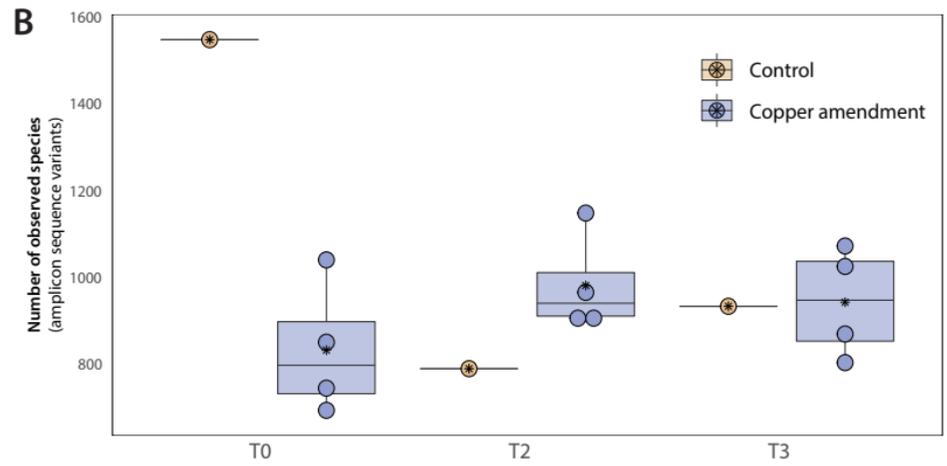
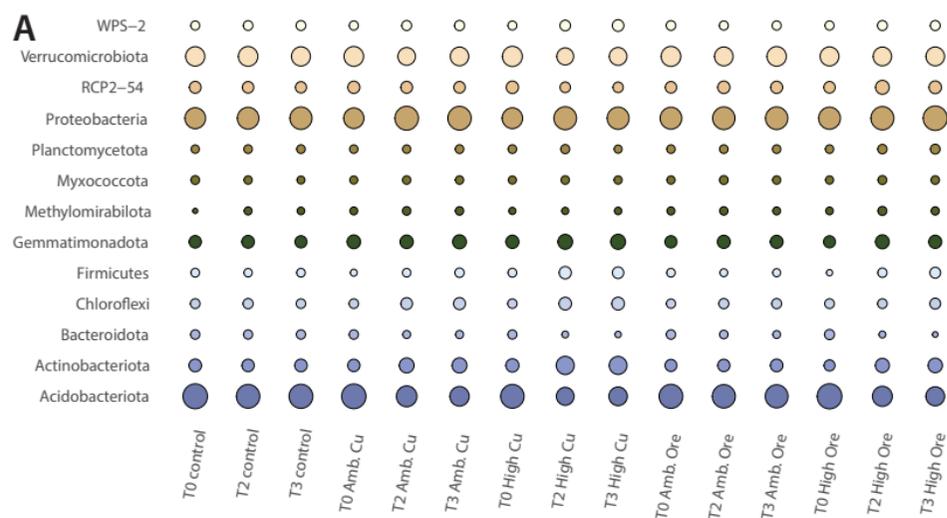
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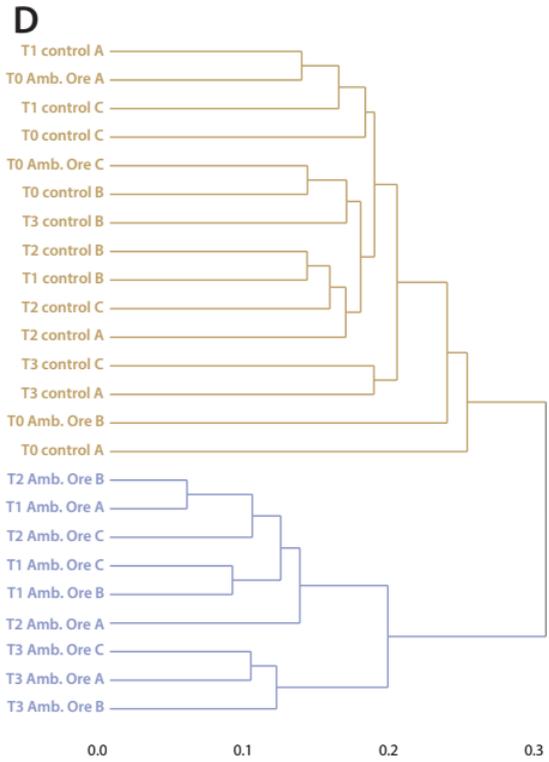
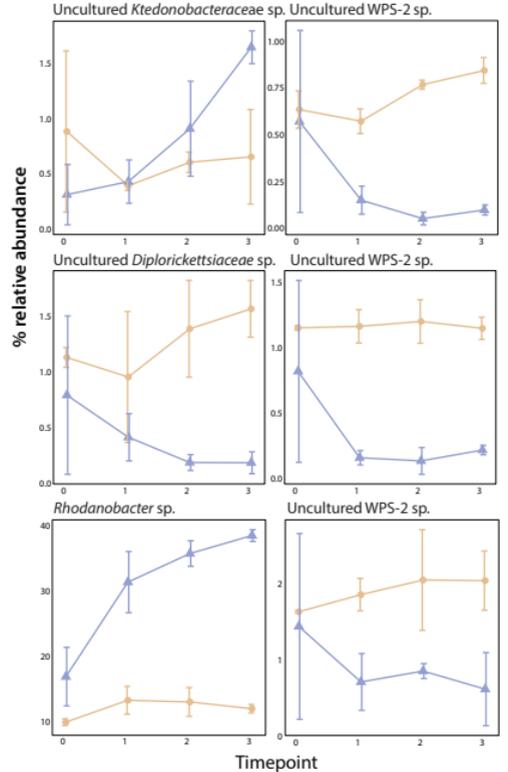
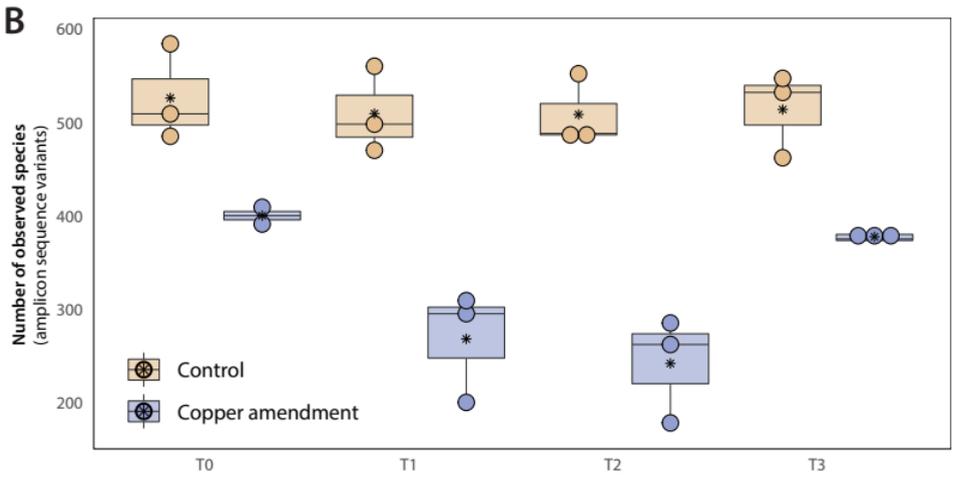
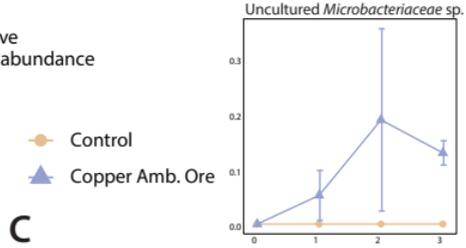
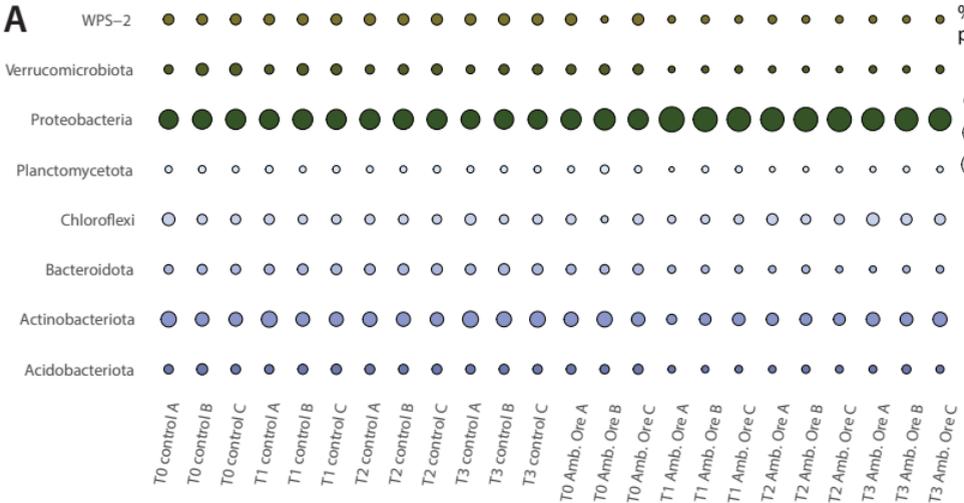
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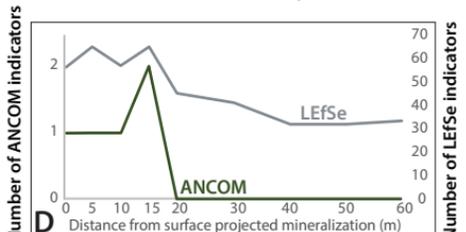
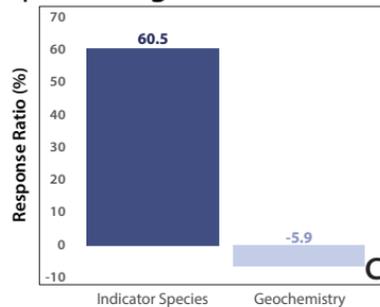
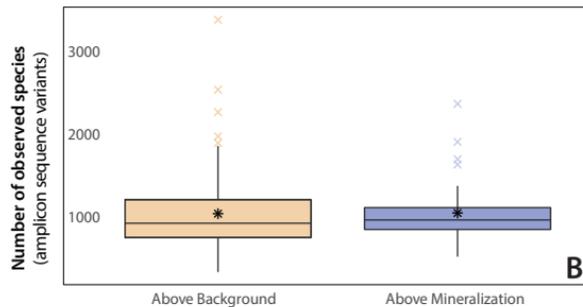
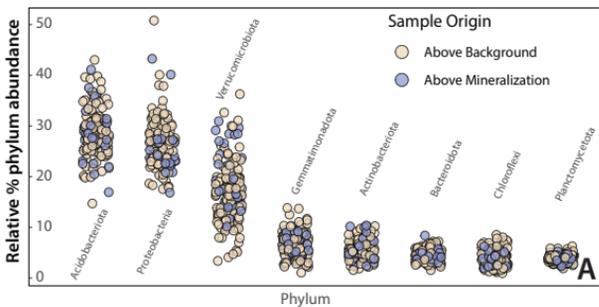
5°24'10.54" N

A

B







**Swamp:** thick organic horizons; well-sorted fine sand, clast supported organic-rich material; low lying areas.

**Moraine actively eroding:** erosional component of the glacially deposited residual moraine.

**Organic rich fluvial deposits:** slow, but active fluvial deposition with accumulation of organic-rich materials.

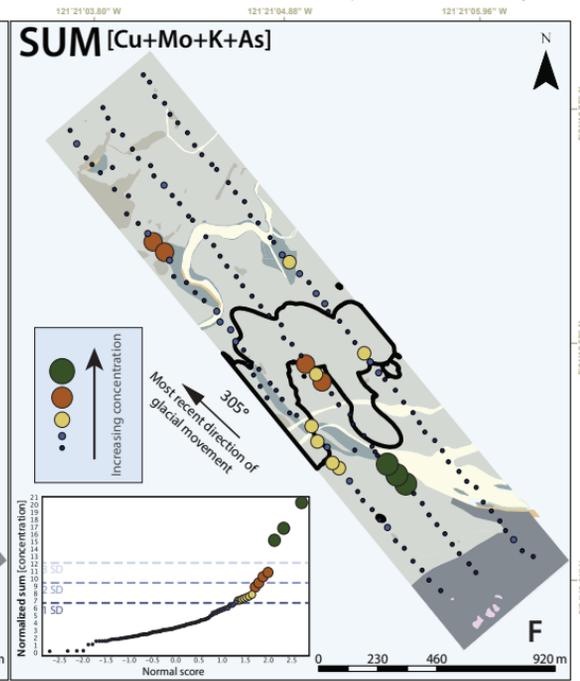
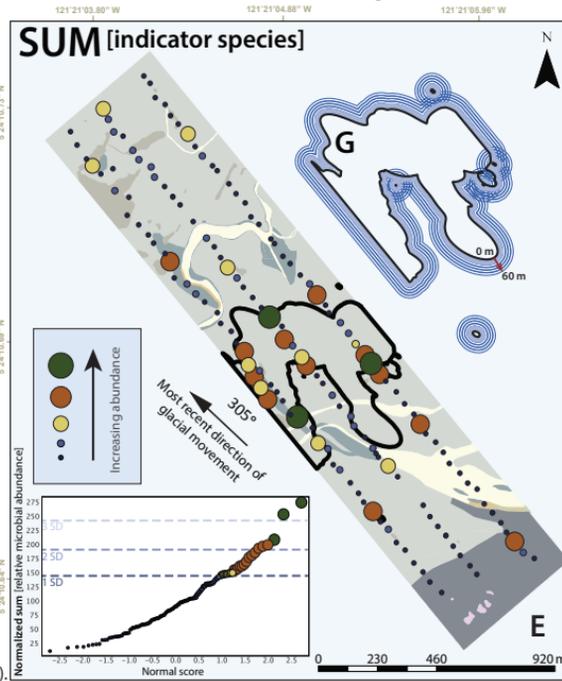
**Glacial fluvial deposits:** fine-medium to coarse sand, moderate to well-sorted, clast supported; depositional and erosional materials; abundant kame structures.

**Residual moraine:** dominant material blanketing bedrock; typically planer, but can be undulating; poorly-sorted till, matrix supported (clay), rounded-angular clasts.

**Reworked fluvial terrace**

**Bedrock:** Nicola Group volcanic sandstone outcrops.

**Mineralization:** 0.2 g/t Au equivalent (incl. >0.22% Cu).



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## **Supplemental Material**

**Supplemental Methods.**

**Figure captions for data Tables S1-S4.**

# Microbial indicators and detection of Cu-sulfide ore mineralization

**Bianca P. Iulianella Phillips<sup>1,2</sup>, Rachel L. Simister<sup>3</sup>, Shane D. Rich<sup>1,2</sup>, Craig J.R. Hart<sup>1,2</sup>, Peter A. Winterburn<sup>1,2†</sup>, and Sean A. Crowe<sup>2,3</sup>**

<sup>1</sup>*MDRU – Mineral Deposit Research Unit, University of British Columbia, Vancouver, BC, Canada, V6T 1Z4*

<sup>2</sup>*Department of Earth, Ocean, and Atmospheric Sciences, University of British Columbia, Vancouver, BC, Canada, V6T 1Z4*

<sup>3</sup>*Department of Microbiology and Immunology, University of British Columbia, Vancouver, BC, Canada, V6T 1Z3*

## SUPPLEMENTAL MATERIAL

### Field sampling and QA/QC

Three soil sampling transects (2.5 km long with 50 m sample spacing and 200 m fence-lines) were established over a porphyry-style Cu-Au mineral deposit (Deerhorn) buried beneath glacial till (Figure 1a). Mineralization is well-defined by drilling with a resource grade of 0.2 g/t Au equivalent ore, which includes most rocks >0.22% Cu. The survey was designed such that it captured up-ice background materials, down-ice background materials and materials that directly overlie the surface expression of mineralization. Soils for DNA sequencing and microbial community fingerprinting at Deerhorn were sampled with sanitized equipment without field screening, to preserve the microbial community as much as possible. Descriptions were documented for *in situ* physicochemical variables at each sample site for every observed soil horizon in the profile (Rich, 2016). Soils at the field sites are derived from the breakdown of till by surface-weathering processes *in situ*, so the soils are considered residual weathering products of the till blanket. B-horizon soils were targeted for each soil sample, although multiple horizons (including O, Ah, Ae, and C) were taken, where possible, for future analyses. Soils for microbial community fingerprinting were frozen at  $-20^{\circ}\text{C}$  upon return to the laboratory at The University of British Columbia (UBC) after 1-2 weeks in field storage ( $-5 - 0^{\circ}\text{C}$ ) and transit, prior to DNA extraction. Sub-samples of the soils used in microbial community profiling were also collected for geochemical analysis. Field measurements (methods described in: (Rich, 2016)) consisted of slurry tests for pH and oxidation-reduction potential (ORP) after field sieving to below  $180\ \mu\text{m}$  (80 mesh). Soil samples for elemental geochemical analysis ( $\sim 1\ \text{kg}$ ) were sent to ALS Global (North Vancouver, BC) for aqua-regia acid digestion and deionized water digestion prior to inductively coupled mass spectrometry (ICP-MS). Field duplicates, CRMs certified reference materials (CRMs), and blanks were inserted into the analytical stream every 20 samples.

### Incubation experiments

Bulk background soil samples from the Deerhorn field area (Sample ID 154414; 52.274, -121.379) and from the Canadian sub-arctic tundra (63.484, -109.08) (Slave Craton, Northwest Territories) were used as baseline soils for the copper incubation experiments. These areas contained background-level metal concentrations and were collected from the upper B-horizon under aseptic conditions. Samples are considered as representative of background soil if its geochemical indicator and pathfinder elements are below the anomalous threshold determined statistically for the surveyed area. The soils were packed into a sealed Poly Ore sample bag and stored at ambient temperature.

The Deerhorn (BC) soil was digested using a multi-acid near total digestion and the digestate analyzed by inductively coupled plasma-mass spectrometry (ICP-MS) ALS Global (North Vancouver, BC) to determine that the soil contained 9 ppm Cu, 1 ppm As, and 0.32 ppm Mo. The bulk soil was not dried prior to the start of the experiment. The soil was dispensed aseptically into sterile containers for each treatment and thoroughly mixed with copper-bearing materials. Amendment concentrations were chosen to represent either concentrations of copper that are routinely detected in geochemical surveys over buried mineral deposits (ambient or '(Am)') or very high levels of copper that might be expected in highly anomalous soils (high or '(Hi)'). The amendments were as follows: 1) 'Hi-ore' soil was amended with chalcopyrite ( $\text{CuFeS}_2$ ) ore at 600 ppm Cu; 2) 'Am-ore' soil was amended with chalcopyrite ore at 200 ppm Cu; 3) 'Hi-Cu' soil was amended with copper in the form of  $\text{CuSO}_4$  (dissolved in Milli-Q®

filtered water) at 600 ppm Cu; and 4) ‘Am-Cu’ soil was amended with copper in the form of  $\text{CuSO}_4$  to 200 ppm Cu. Soil was sampled at T = 0, T = 1 (14 days), and T = 2 (35 days).

The tundra soils (NWT) were digested using a near total multi-acid digestion and the digestate was analyzed by ICP-MS at ALS Minerals Laboratories Ltd (North Vancouver, BC) to determine that the soil contains 8 ppm Cu, < 1 ppm As, and 0.73 ppm Mo. The bulk soil was not dried prior to the start of the experiment. Soil was dispensed aseptically into sterile containers for each treatment and amended with the same chalcopyrite ore described above at 200 ppm. Soil was sampled at T = 0, T = 1 (15 days), T = 2 (55 days), and T = 3 (85 days).

### **DNA extraction and QA/QC**

DNA extraction and quantification was performed as described in: (Simister et al., 2023).

### **SSU rRNA gene amplification and DNA amplicon sequencing**

Gene amplification and DNA amplicon sequencing were performed as described in: (Simister et al., 2023).

### **Bioinformatics**

DNA sequences were processed using the QIIME2 amplicon sequence analysis program (Bolyen et al., 2019) with the DADA2 denoising pipeline (Callahan et al., 2016). Sequences were clustered to amplicon sequence variants (ASV) and classified using SILVA reference taxonomy database (release 138). ASVs that had less than 1 read were filtered from analysis. Species richness measures (number of observed features) were calculated in QIIME2. Sequences were deposited into the Sequence read archive (SRA) under accession numbers PRJNA698256 and PRJNA698251.

### **Anomaly identification and mapping**

LEfSe indicator species analyses (Segata et al., 2011) were performed in the Mothur program (subsampling to 11132 sequences) (Schloss et al., 2009) and ANCOM indicator species analyses (Mandal et al., 2015) were performed in the QIIME2 program, where indicator species (ASVs) are considered statistically significant if the LDA score > 2 (LEfSe) or a W score rejects null hypothesis (ANCOM). Sample groups for the copper amendment incubation experiments are based on unamended “control soils” and amended “copper-bearing soils”. Sample groups were set for field analyses based on their origin from “background soil” or “soils above mineralization”. These groups are defined based on underlying geology whereby “background soils” come from above the andesite breccias and thick beds of volcanic derived sandstones (del Real et al., 2017) and “soils above mineralization” come from above the surface projection of mineralization (0.2 g/t Au equivalent) as defined by drilling.

Differential abundance analyses of Deerhorn soils were multi-tiered with the soil samples considered “above mineralization” and “above background” defined based on the distance from the drilling-defined surface projection of mineralization. Soil samples “above mineralization” in each analysis ranged from 0 m – 60 m (0 m, 5 m, 10 m, 15 m, 20 m, 30 m, 40 m, 50 m, and 60 m). The combined collection of indicator species derived from 0 m – 15 m were used downstream to create the microbial anomaly map.

Incubation-derived indicator species showing an enrichment in copper amended soils were curated to plot at Deerhorn. Indicator species with > 1 average reads per sample in the incubation experiment and positive response ratios at the Deerhorn field site were included. Response ratios for indicator species were calculated as the ratio between the average relative abundance in “soils above mineralization” and the average relative abundance of “background soils”. Indicator species predicted from the Deerhorn field site were not curated further, thus each indicator species output was included in the generation of the anomaly maps.

Both map data plots were created using relative abundances of indicator species from 16S rRNA gene sequencing and pathfinder/indicator element concentrations from aqua-regia acid digest ICP-MS results. Individual indicator species and pathfinder/indicator elements were normalized to the mean prior to summation. Response ratio bar plots of the normalized sums of indicator species and pathfinder elements are expressed by the following equation:

$$\left( \left( \text{average} \left( \frac{\text{on deposit}}{\text{off deposit}} \right) \right) - 1 \right) * 100$$

Anomaly identification through probability plots was done in the Reflex/Imdex ioGAS software (version 8.0), and mapping of anomalies and surficial geology was performed in the ESRI ArcGIS software. QGIS was used to determine soil sample distance from the surface projection of mineralization.

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## SUPPLEMENTAL DATA TABLES S1-S4 FIGURE CAPTIONS

**Table S1.** Soil geochemical data from the Deerhorn field survey. Elemental concentrations from ALS Global (North Vancouver, BC), with digest/analysis codes indicated above the concentration value for each analyte.

**Table S2.** Species richness data from soil microbial communities (number of observed species/features/ASVs) from the soil incubation experiments and the Deerhorn field survey.

**Table S3.** List of indicator species and their relative abundances for the Deerhorn porphyry copper deposit.

**Table S4.** Baseline geochemistry for the starting soils for each incubation experiment (Deerhorn (BC) and tundra (NWT)). Digest/analysis codes for ALS Global (North Vancouver, BC) are indicated above the concentration value for each analyte.