

1 **Physicochemical and urban land-use characteristics associated**
2 **with resistance to precipitation in estuaries vary across scales**

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14

15 **Abstract.** Estuaries are subject to frequent stressors, including elevated nutrient loading and
16 extreme hydrologic events, which impact water quality and disrupt ecosystem stability and health.
17 The capacity of an estuary to resist changes in function in response to precipitation events is a key
18 component of maintaining estuarine health in our changing climate. However, generalizable
19 patterns in factors related to estuarine responses to extreme precipitation remain unknown. We
20 investigate physicochemical factors and land-use characteristics that are associated with ecological
21 resistance to precipitation – broadly defined as the magnitude of ecosystem change induced by an
22 event – in five disparate estuaries distributed across the continental United States. Using long-term
23 meteorological and water quality data from the National Estuarine Research Reserve System along
24 with land use/land cover and population data, we examine relationships between the resistance
25 index – a proxy for ecosystem stability calculated using dissolved oxygen – and physicochemical
26 and urban land use characteristics on local-to-continental scales. Contrary to our initial hypothesis,
27 we found that more urbanized estuaries were more resistant to precipitation events, and that water
28 temperature, water column depth, turbidity, nitrogen, and chlorophyll-*a* were related to resistance
29 on a continental scale. However, these trends interacted with estuarine salinity and varied across
30 individual estuaries; where we found additional relationships of resistance with salinity, turbidity,
31 phosphate concentrations, N:P, and tree cover. Considering emerging stressors from new climatic
32 scenarios and from urbanization, these results are important for representing the impacts of
33 disturbances in large-scale models and for informing management decisions regarding estuarine
34 water quality.

35 1. Introduction

36 Estuaries are highly dynamic environments that often connect freshwater and saltwater
37 systems, cycle organic matter and nutrients from land to oceans, and provide essential ecosystem
38 services (Bianchi, 2007; He and Silliman, 2019). The function of estuarine ecosystems as unique
39 sites for carbon and nutrient cycling, and habitats for macro- and micro-flora and fauna, relies on
40 stability of a predictable range of dynamic processes like temperature fluctuations, hydrology and
41 nutrient mixing. However, anthropogenic activities and extreme climatic events threaten estuarine
42 ecosystem stability and function (Kemp et al., 2009; Zhang et al., 2010). Predicted increases in
43 urban population size and in the frequency and intensity of precipitation events highlight the
44 urgency to understand the response of estuaries to new urban and climatic scenarios (Kyzar et al.,
45 2021; Li et al., 2019; Martínez et al., 2007; Pickett et al., 2011). Yet, the factors that impact the
46 response of urban estuaries to precipitation are not fully understood.

47 When combined with intense precipitation, watershed urbanization and associated changes
48 in land use/land cover (LULC) often result in increases in stream hydrological flashiness (Gannon
49 et al., 2022; Grimm et al., 2008; Reisinger et al., 2017). Hydrological flashiness induces increased
50 flow rates that cause changes in channel morphology (Booth and Jackson, 1997; Gregory, 2011;
51 Leopold, 1968; Vietz et al., 2016), habitat destruction (Walsh et al., 2005), and disruption of
52 microbial metabolic processes (Reisinger et al., 2017; Uehlinger, 2000). Flashiness can also
53 drastically affect primary production – a regulatory component of dissolved oxygen (DO)
54 dynamics in aquatic environments – through increases in flow velocity, transport of phytoplankton,
55 sediment migration, and light limitation through increased turbidity (Bernot et al., 2010; Fisher et
56 al., 1982; McSweeney et al., 2017; Reisinger et al., 2017; Uehlinger, 2000).

57 While DO is dynamic and depends on a myriad of biological, chemical, and physical
58 processes, it is essential for many estuarine functions and has been widely used as an indicator for
59 overall ecosystem health (Abdul-Aziz et al., 2007; Abdul-Aziz and Gebreslase, 2023; Chapra,
60 2008; Cox, 2003; Kannel et al., 2007; Zhi et al., 2021). Dissolved oxygen is critical to maintaining
61 the life cycles of macro- and micro-fauna, for example, and supports the biogeochemical cycling
62 of carbon and nutrients by serving as a terminal electron acceptor (Bernhardt et al., 2018; Chapra,
63 2008; Zarnetske et al., 2012). Built areas in particular have been previously connected to altered
64 aquatic DO concentrations and other water quality parameters (Bernhardt et al., 2008; Chang,
65 2005; Freeman et al., 2019; Vietz et al., 2016). Given the tight link between aquatic DO, land use,

66 and diverse ecosystem functions, DO is an important indicator and holistic measure of estuarine
67 health.

68 In addition to impacts on DO, extreme precipitation events are often associated with excess
69 nitrogen (N) delivery and changes in salinity, particularly for waterways adjacent to urban-type
70 LULC (Walsh et al., 2005). Nitrogen is central to mediating metabolic processes across a wide
71 range of ecosystems and land covers (Mulholland et al., 2008; Schindler, 1977; Smith, 1984;
72 Vitousek and Howarth, 1991). It is particularly important for primary production in aquatic
73 ecosystems because many phytoplankton species are N-limited (Evans and Seemann, 1989;
74 Howarth, 1988; Vitousek and Howarth, 1991). Further, extreme precipitation events are often
75 associated with influx of freshwater (i.e., runoff) and/or saltwater (i.e., storm surges) into estuaries,
76 which changes salinity, can induce stratification and hypoxia, and impact ecosystem metabolic
77 functions (Rabalais et al., 2010; Wetz and Yoskowitz, 2013; Zhang et al., 2010). However,
78 depending on the temporal and spatial scales of evaluation, the reported trends of estuarine
79 responses to precipitation and salinity changes can be conflicting.

80 Commonalities in responses to major precipitation events across disparate estuaries can
81 help to project long-term estuary health under progressive urbanization and more extreme climate
82 events; but they have been difficult to decipher. This is, in part, due to variation in how the complex
83 and interconnected dynamics within estuaries respond to precipitation, even at local scales. While
84 Ombadi & Varadharajan (2022) report contrasting effects of urbanization on salinity during flood
85 events when regional climate is considered, a continental-scale study by Kaushal et al. (2018)
86 suggests that anthropogenic activity is associated with increasing salinity in waterways. However,
87 the later study recognizes that regional, climatic, LULC, and geologic variabilities also influence
88 salinization patterns. Similarly, continental-scale evaluations showed that waterways within small
89 watersheds appear consistently less flashy than those in large watersheds; and that there is a
90 substantial amount of variability in these relationships at regional scale (Baker et al., 2004; Gannon
91 et al., 2022; Hopkins et al., 2015; Poff et al., 2006). Such variation in relationships across scales
92 may be particularly prevalent in ecosystems influenced by anthropogenic activities (Hopkins et
93 al., 2015; Poff et al., 2006), which demonstrates the importance of considering multiple spatial
94 scales in understanding estuarine responses to changes in precipitation patterns and watershed land
95 use.

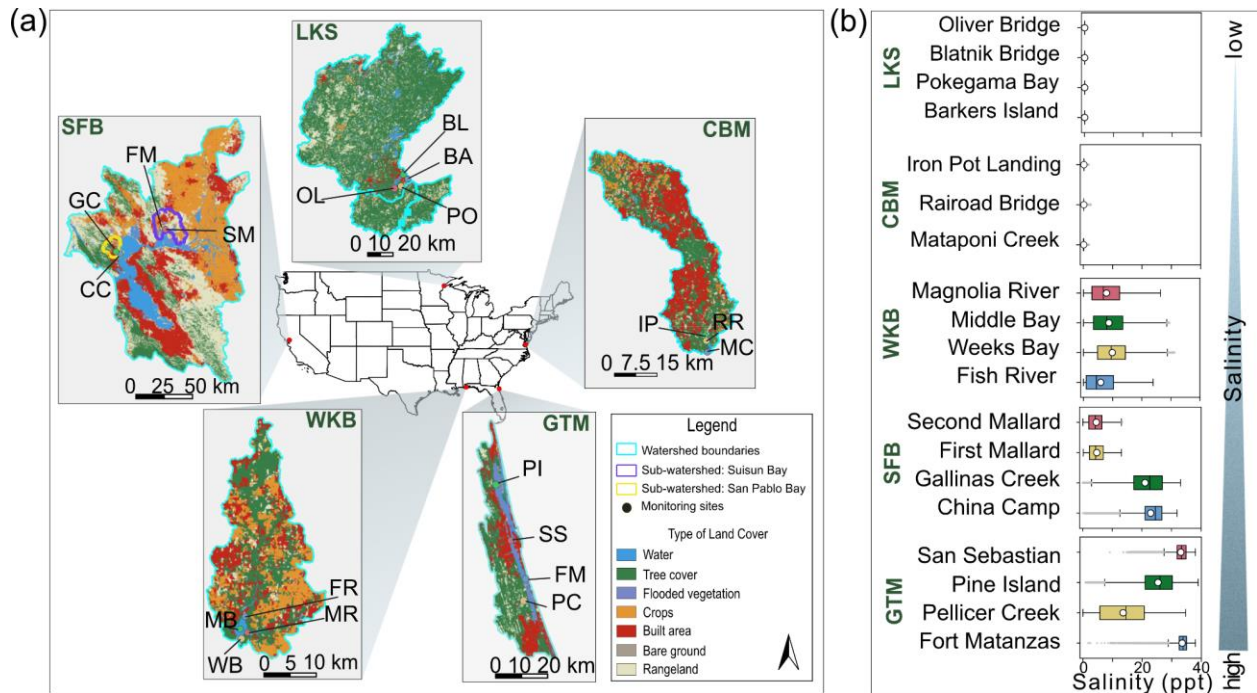
96 We aim to uncover generalizable patterns of responses to large precipitation events across
97 five estuaries spanning a gradient of urbanization and physicochemical properties. Using DO as
98 an indicator of ecosystem health (Abdul-Aziz and Gebreslase, 2023), we evaluate estuary
99 resistance to precipitation – defined here as the ability of estuaries to maintain stability in DO
100 (Isbell et al., 2015; Lake, 2013; McCluney et al., 2014; Pimm, 1984; Utz et al., 2016; Van
101 Meerbeek et al., 2021) – in the context of physicochemical factors and land-use characteristics at:
102 1) the continental-scale (i.e., across all estuaries); 2) across estuaries grouped by salinity; and 3)
103 within each estuary. We hypothesize that urbanization decreases estuarine resistance to
104 precipitation, and further that relationships of resistance with physicochemical and land-use factors
105 will diverge with spatial scale and differences in ambient salinity. This study helps understanding
106 how ongoing changes in climate and urbanization conditions influence estuarine ecosystem health.

107

108 **2. Methods**

109 **2.1 Dataset description.**

110 We used long-term water quality monitoring data from five estuaries in the National
111 Estuarine Research Reserve System (NERRS, 2023) to understand factors associated with their
112 resistance to precipitation events. Lake Superior (LKS), WI; Chesapeake Bay Maryland (CBM),
113 MD (Jug Bay only); Guana Tolomato Matanzas (GTM), FL; Weeks Bay (WKB), AL; and San
114 Francisco Bay (SFB), CA span climatic zones, land uses, and salinity (0.1–35 ppt) (Fig. 1, Table
115 1). Across all estuaries, there were a total of 19 monitoring locations (3 at CBM, and 4 at LKS,
116 GTM, WKB, and SFB).



117
 118 **Figure 1.** Selected National Estuarine Research Reserve (NERR) stations. a) Map of monitoring
 119 locations and land use/land cover within associated watersheds at Lake Superior (LKS) NERR,
 120 Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR, Weeks Bay
 121 (WKB) NERR, and San Francisco Bay (SFB) NERR. b) Salinity from 2012 to 2022 for each
 122 monitoring location at each NERR (all $n > 150,000$). Boxes indicate interquartile range. Means
 123 are shown in white circles. Black lines inside the boxes indicate medians.

124
 125 **Table 1.** Land use/land cover (LULC) and population density in each estuary within 10-km to
 126 water quality and nutrient monitoring locations.

LULC class	Estuary and % LULC class within 10-km to monitoring locations					
	LKS	CBM (Jug Bay)	GTM	WKB	SFB	
					(San Pablo Bay) GC	(Suisun Bay) CC, FM, SM
Water	10.43	2.30	8.58	4.85	5.04	3.59

Tree cover	58.78	53.72	45.11	38.52	26.91	1.16
Flooded vegetation	0.88	1.37	13.04	0.05	1.4	21.33
Crops	0.19	4.63	0.19	30.36	18.05	19.67
Built area	18.84	27.19	23.99	19.33	31.65	21.87
Bare ground	0.06	0.25	0.26	0.1	0.11	1.17
Rangeland	9.82	10.53	8.82	6.76	16.47	31.2
Surface area and population estimates						
Area (km ²)	488.7	218.2	645.6	213.0	105.01	205.86
Estimated population within the area (ppl)	71,111	49,291	148,557	17,404	48,161	100,582
Population density (ppl km ⁻²)	145	225	230	81	458	488
LULC definitions per ESRI (Karra et al., 2021)						
<p><i>Water</i> – areas where water was present throughout the year. Excludes man-made structures like docks</p> <p><i>Tree cover</i> – vegetation with closed/dense canopy ≥ 15 meters.</p> <p><i>Flooded vegetation</i> – areas with intermixing of water and vegetation flooded seasonally or predominantly throughout the year.</p>						

Crops – human planted vegetation (cereals, grasses, and corps) that are not at tree height.

Built area – human made structures like roads, rail road networks, parking spaces, industrial and residential buildings.

Bare ground – areas dominated by rocks, soil, sand (i.e. desert) with sparse to no vegetation throughout the year

Rangeland – homogeneous grasses, mixes of vegetation below tree-height with rocks and soil, clearings in the forests.

127

128 To assess the relationships between resistance and urbanization, we used LULC data at 10-
129 meter resolution from ESRI (2017–2020) (Karra et al., 2021) and population density data at 100-
130 meter resolution from World Population Hub (Bondarenko et al., 2020) (Table 1). We analyzed
131 LULC and population density using QGIS 3.30.3 (QGIS Geographic Information System, 2024)
132 equipped with a semi-automatic classification plug-in.

133 Further, because the resistance index was calculated for precipitation events across short
134 time-scales (i.e., days), which underrepresents the draining time of some watersheds, we
135 considered LULC and population density from a 10-km proximity zone surrounding monitoring
136 locations (Table 1). For SFB NERR, we used 10-km proximity zones from San Pablo and Suisun
137 embayments to quantify LULC and population density as is conventionally done at this watershed
138 due to different hydrologic dynamics within each ebayment.

139 All data were collected using NERRS standard operating procedures. Briefly, water
140 column DO (mg L^{-1} , calculated from % air saturation, temperature, and salinity at the time of
141 measurement; thereby accounting for the impacts of temperature and salinity fluctuations on DO
142 concentration), temperature, turbidity, salinity, and depth were measured at 15-min intervals using
143 synchronized YSI-EXO2 multiparameter sondes. Meteorological conditions, including
144 precipitation, were also measured at 15-min intervals using NERRS standard weather station
145 instrumentation. PO_4^{3-} , NO_3^- , NH_4^+ , NO_2^- and chlorophyll-*a* (Chl-*a*) were measured monthly from
146 grab samples and analyzed in the lab (NERRS, 2023). Samples for nutrients were filtered in the
147 field through $0.7 \mu\text{m}$ glass-fiber filters and analyzed following U.S. Environmental Protection
148 Agency (EPA) methods (O'Dell, 1996b, a; U.S. EPA., 1993a). For Chl-*a* analysis, the samples
149 were collected as whole water, then filtered onto $0.45 \mu\text{m}$ glass-filters and processed following
150 APHA, (2001) and U.S. EPA., (1993b) methods. While Chl-*a* is also measured as Chlorophyll

151 fluorescence using optical sensor with YSI sondes, NERRS recommends using laboratory-based
152 measurements due to the confounding effects of various fluorescent species, organism type, light,
153 temperature and more on sensor-based Chl-*a* measurements (sensor and lab-based measurements
154 often do not correspond to each other). We omitted all data flagged as ‘suspect’ or ‘out of range’.

155 Additionally, we calculated water column depth as the sum of measured water depth plus
156 the distance between the sonde and sediment bed (Table S1). We also calculated the sum of NO_3^-
157 , NO_2^- , and NH_4^+ to assess dissolved inorganic nitrogen (DIN) concentrations and the ratio of DIN
158 to PO_4^{3-} (hereafter, N:P).

159

160 **2.2 Determination of major precipitation events.**

161 Because hydrologic dynamics and related estuarine functions can vary dramatically with
162 annual weather conditions, we first selected one ‘wet’ and one ‘dry’ year for each estuary using
163 precipitation records from nearby airports. The purpose of selecting years with disparate rainfall
164 patterns was to encompass the maximum range of variability in expected estuarine resistance to
165 precipitation. Following Murrell et al. (2018), we calculated the long-term interquartile range
166 (IQR, 1990-2020) of total monthly precipitation for each estuary and then selected relatively
167 wet/dry years based on the number of months plotting above/below IQR and total annual
168 precipitation. Possible wet and dry years were further filtered based on the completeness of
169 NERRS data available for each estuary (Fig. S1). Long-term precipitation records included: Duluth
170 International, Washington Reagan International, Jacksonville International, Birmingham Airport,
171 and San Francisco International airports available from the National Center for Environmental
172 Information.

173 Further, we selected major precipitation events within each wet and dry year by plotting
174 daily precipitation using data from NERR meteorological stations (Fig. S2). Specifically, because
175 the definition of ‘major’ precipitation events is hard to quantify and it varies across estuaries, we
176 first considered hurricanes, tropical storms, Nor’easters, and other major storm events noted within
177 NERR metadata sheets when selecting precipitation events. For example, at GTM we focused on
178 tropical storms Colin, Julia, and Hermine, hurricanes Matthew and Irma, and two Nor’easters
179 (Table S2). Data availability was a second consideration – we removed possible events for which
180 there was a substantial amount of missing data. Lastly, because metabolic and hydrologic
181 processes vary across seasons, we chose to focus on warm season events, with the exception of

182 San Francisco Bay where most precipitation occurs in the cool season but seasonal temperature
183 fluctuations are generally lower than in other systems (Fig. S2 and S3, Table S2).

184

185 **2.3 Calculation of resistance.**

186 To investigate physicochemical and urban land-use characteristics associated with
187 estuarine responses to precipitation events, we calculated the resistance index described in Orwin
188 & Wardle (2004). The resistance index is a normalized parameter (−1 to +1) describing the
189 magnitude of shift in a response variable from an initial condition. It has been used across a wide
190 variety of ecosystems and response variables, including aquatic ecosystems (Thayne et al., 2022,
191 2023; Tsai et al., 2011). The resistance index is calculated as:

$$192 \textit{Resistance} = 1 - \frac{2|D_0|}{(C_0 + |D_0|)} \text{ (eq. 1)}$$

193

194 where, C_0 = concentration of the response variable pre-disturbance, and D_0 = difference between
195 the concentration of the response variable pre- and post-disturbance (P_0) (i.e., $D_0 = C_0 - P_0$). An
196 index value of +1 indicates the highest possible resistance. Index values between 0 and 1 show
197 that the magnitude of response variable shift is less than the magnitude of the baseline (i.e., $|D_0| \leq$
198 C_0). A resistance index of 0 indicates that the shift in the response variable is equivalent to the
199 magnitude of the baseline (i.e., $|D_0| = C_0$), whereas index values between < 0 and -1 reflect that
200 change in the response variable is greater than the magnitude of the baseline (i.e., $|D_0| > C_0$).
201 Overall, index values closer to 1 indicate more resistant systems (Orwin and Wardle, 2004).

202 While the resistance index is indicative of the ability of a system to maintain its pre-
203 disturbance state, we emphasize that it is a normalized value that does not in itself convey
204 information about overall estuarine health. Similarly, the resistance index in itself cannot infer if
205 a shift has a positive or a negative effect on the ecosystem health. The resistance index can infer
206 the ability of a system to maintain its pre-disturbance functions and/or ecological state (i.e.,
207 ecosystem stability). It enables the comparison of the amount of change induced by disturbance
208 across vastly different estuaries. In parallel, it is also useful to consider the absolute value of the
209 response variable (in this case DO) pre- and post-disturbance, which conveys information on the
210 ambient state of an estuary and the directionality of its response to the disturbance. We therefore
211 present C_0 and P_0 (Fig. 2a–e, Table S3) to define differences in DO within and across estuaries, as

212 well as the resistance index, which pairs C_0 and P_0 values for the same event (Fig. 2f–j), to
213 understand the ecosystem stability within and across estuaries after precipitation.

214 Because DO is critical to myriad functions that regulate the health of aquatic ecosystems
215 (Abdul-Aziz et al., 2007; Abdul-Aziz & Gebreslase, 2023; Caffrey, 2004; Mulholland et al., 2001;
216 Murrell et al., 2018; Odum, 1956) and because estuarine metabolism is virtually impossible to
217 model due to bi-directional flow (Loken et al., 2021); we used temperature and salinity adjusted
218 DO (mg L^{-1}) to calculate the resistance index. Additionally, the resistance index is highly sensitive
219 to the researcher-defined baseline and post-disturbance time periods that are used to calculate C_0
220 and D_0 . Therefore, we calculated C_0 as average DO during a manually-curated timespan preceding
221 each precipitation event (~24 hours to 6 days, without precipitation) where DO dynamics looked
222 unremarkable (Fig. S4, Table S2). The timespan for estimating P_0 was also manually selected in
223 the context of each event, defined here as the maximum displacement from C_0 during and after the
224 event (Fig. S4, Table S2). We also verified that the resistance index calculated using DO (mg L^{-1})
225 versus DO (% sat.) showed no notable differences (Fig. S5).

226

227 **2.4 Statistical analysis.**

228 To compare resistance, nutrient concentrations, and concentrations of DO pre- and post-
229 precipitation within and between estuaries, we used ANOVA with post-hoc Tukey HSD or
230 Kruskal-Wallis test, as appropriate based on the Shapiro-Wilk normality test.

231 To test associations of specific physicochemical and land-use factors with estuarine
232 resistance to precipitation, we used linear regressions at continental and local scales independently
233 (i.e., all estuaries combined vs. within each individual estuary) and within salinity-based groups
234 (i.e., using a threshold of average annual salinity less than or above 10 ppt calculated for wet and
235 dry years separately). The low salinity group included all LKS and CBM locations, First Mallard
236 and Second Mallard locations at SFB during wet and dry years, as well as Fish River, Middle Bay,
237 Magnolia Bay, and Weeks Bay (WB: during dry year only) at WKB estuary. The high salinity
238 group included all monitoring locations at GTM, China Camp and Gallinas Creek at SFB estuary
239 for wet and dry years, and Weeks Bay at WKB during the wet year. We used annual mean values
240 for continental-scale, local-scale, and salinity-based groups regressions involving nutrients and
241 Chl-*a* because monthly sampling intervals of lab-based measurements did not always correspond
242 with selected precipitation events. This resulted in two data points per monitoring location for

243 nutrients and Chl-*a* regressions. While not directly associated with any particular precipitation
244 event, relationships of nutrients and Chl-*a* with resistance values analyzed on an annual basis carry
245 valuable information about how the ambient conditions of the system can impact an estuary's
246 response to precipitation. This knowledge is essential for deriving and testing hypotheses that
247 describe why a certain estuary may respond to a storm event in a particular way. For regressions
248 involving sensor-based measurements, we attempted to provide as much resolution as possible into
249 specific events. We therefore used values averaged over the event-specific baseline periods (Table
250 S2) for sensor measurements across scales and salinity-based groupings. Values averaged over
251 event-specific baseline periods were matched with corresponding event-specific resistance index
252 resulting in one data point per precipitation event per each monitoring station. We did not include
253 LULC and population density as predictors of resistance at individual estuaries because of the
254 close proximity of some monitoring locations to one another (< 800 m).

255 Statistical analyses were performed in Python 3.10.11 using `scipy.stats`,
256 `statsmodels.stats.multicomp.pairwise_tukeyhsd`, and `seaborn.regplot` packages.

257

258 **3. Results**

259 **3.1 Land use/land cover and nutrient concentrations across estuaries.**

260 Data describing LULC and population density within 10-km of the monitoring locations
261 for all estuaries are shown in Table 1. Both embayments of the SFB had higher percentages of
262 urban-type land characteristics (e.g., built area) and population density than any other estuary,
263 followed by CBM (Table 1). Agricultural land was more prevalent at WKB (30.36%) compared
264 to other estuaries. GTM and LKS had more mixed LULC, with high proportions of tree cover.
265 With regard to nutrient concentrations, SFB and CBM had high DIN concentrations compared to
266 LKS and GTM ($p < 0.01$). Mean DIN values at all SFB and CBM monitoring locations were >
267 $0.504 \text{ mg-N L}^{-1}$ versus $< 0.16 \text{ mg-N L}^{-1}$ at LKS and GTM (Fig. S6). Phosphate concentrations were
268 the highest at SFB (mean across all monitoring locations = $0.135 \text{ mg-P L}^{-1}$, $SD = 0.08$; means at
269 all other estuaries $< 0.03 \text{ mg-P L}^{-1}$, $p < 0.001$, Fig. S5). Overall, mean N:P across LKS, CBM,
270 GTM, WKB, and SFB estuaries were 28.04 ($SD = 28.08$), 26.56 ($SD = 10.75$), 2.84 ($SD = 0.91$),
271 83.82 ($SD = 55.27$), and 5.10 ($SD = 1.4$), respectively (Fig. S7), with GTM and SFB indicating N
272 limiting conditions based on a 16N:1P Redfield ratio (Redfield, 1934).

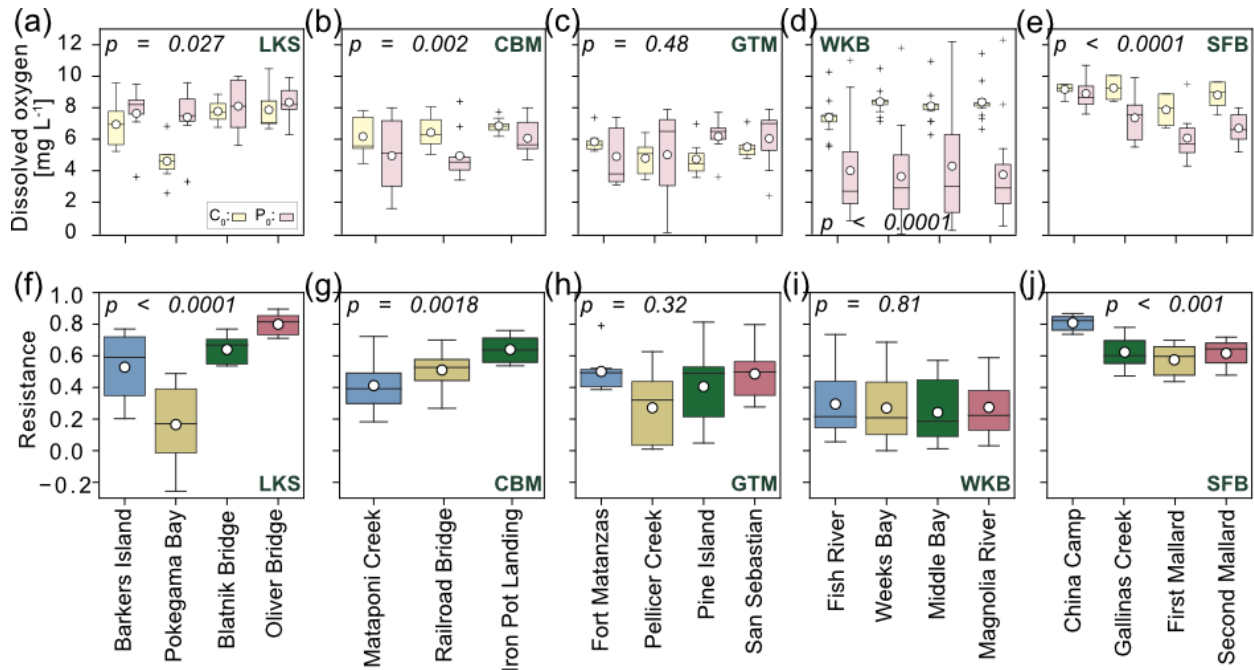
273

274 **3.2 Changes in dissolved oxygen and resistance to precipitation across estuaries and**
275 **monitoring locations.**

276 DO concentrations pre- and post-precipitation differed across all estuaries, when
277 evaluating the overall trends (pre- (C_0): $F = 45.6$, $p < 0.0001$; and post- (P_0): $F = 17.1$, $p < 0.0001$,
278 Fig. 2a–e). Across all precipitation events (i.e., non-specific to an event), SFB had the highest pre-
279 precipitation DO concentration of all estuaries. Generally, at SFB and CBM (more urban) and at
280 WKB (more agricultural), DO declined following precipitation ($p < 0.01$). At LKS, precipitation
281 events significantly increased DO concentration ($F = 5.1$, $p = 0.027$). There was no significant
282 difference between the overall pre- and post-precipitation DO concentrations at GTM ($F = 0.5$, p
283 $= 0.48$).

284 Resistance also differed across estuaries ($F = 21.6$, $df1 = 4$, $df2 = 172$, $p < 0.0001$, Fig. 2).
285 SFB monitoring locations had the highest mean resistance to precipitation (mean = 0.68), while
286 monitoring locations at LKS were the least resistant to precipitation (mean = 0.47). Within
287 individual estuaries, resistance varied across monitoring locations at LKS ($F = 13.9$, $df1 = 3$, $df2$
288 $= 24$, $p < 0.0001$), CBM ($F = 7.9$, $df1 = 2$, $df2 = 27$, $p < 0.01$), and SFB ($F = 10.2$, $df1 = 3$, $df2 =$
289 27 , $p < 0.001$) but was not significantly different between monitoring locations within GTM ($F =$
290 3.5 , $df = 3$, $p = 0.32$), WKB ($F = 0.9$, $df = 3$, $p = 0.81$) (Fig. 2f–j). Resistance was most variable
291 across monitoring locations at LKS (–0.26 to 0.89) and least variable across locations at WKB (0
292 to 0.73).

293



294

295 **Figure 2:** Variation in pre- and post-disturbance distribution of dissolved oxygen and resistance
 296 within individual estuaries. Boxes show the quartiles of the dataset and the whiskers show the rest
 297 of the distribution. Means are shown in white circles, and medians are shown in black solid lines.
 298 *P*-values are at the top of each panel. a-e) Distribution of dissolved oxygen concentrations prior to
 299 (C₀, solid boxes) and post (P₀, dashed boxes) precipitation. f-j) Resistance index across monitoring
 300 locations at: Lake Superior (LKS) NERR (all n = 7), Chesapeake Bay (CBM) NERR (all n = 10),
 301 Guana Tolomato-Matanzas (GTM) NERR (all n = 7), Weeks Bay (WKB) NERR (all n = 15), San
 302 Francisco Bay (SFB) NERR (n = 8 except Second Mallard (SM) n = 7).

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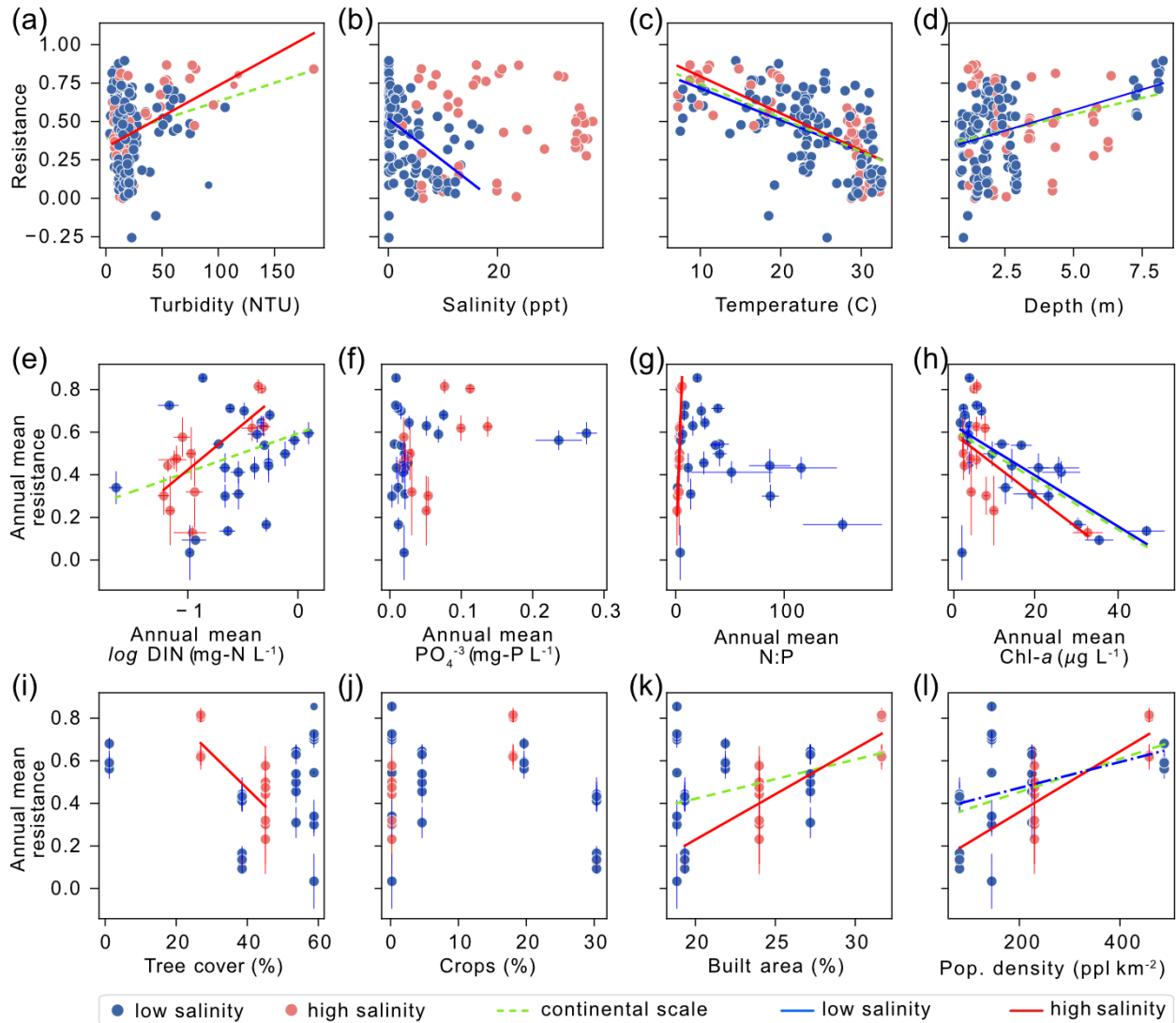
304 **3.3 Continental, salinity-based, and local relationships of resistance with physicochemical**
 305 **factors, land use/land cover, and population density.**

306 When data from the five estuaries were considered together (i.e., continental-scale), we
 307 found significant positive relationships of resistance with water column depth ($p < 0.0001$; $R^2 =$
 308 0.10), turbidity ($p = 0.0015$; $R^2 = 0.06$), log(DIN) ($p = 0.031$; $R^2 = 0.12$), percent built area ($p =$
 309 0.02; $R^2 = 0.14$), and population density ($p = 0.001$; $R^2 = 0.27$); and significant negative
 310 relationships to water temperature ($p < 0.0001$; $R^2 = 0.39$) and Chl-*a* ($p < 0.0001$; $R^2 = 0.39$) (Fig.
 311 3, Table S4).

312 When grouped by salinity, estuarine resistance to precipitation was more tightly correlated
 313 to physicochemical and land-use factors (Fig. 3, Table S4). Within ‘low-salinity’ estuaries,

314 resistance was positively related with depth of the water column ($p < 0.0001$; $R^2 = 0.1$), which is
315 consistent with the trend observed on continental-scale; and negatively related to salinity ($p <$
316 0.0001 ; $R^2 = 0.214$), a relationship not found on continental-scale. Within ‘high-salinity’ estuaries,
317 mean resistance showed positive relationships to annual mean log(DIN) ($p = 0.004$; $R^2 = 0.55$) and
318 to percent built area ($p = 0.0002$, $R^2 = 0.73$), which also was consistent with continental-scale
319 results. Also, resistance in high-salinity estuaries was positively associated with turbidity ($p <$
320 0.0001 , $R^2 = 0.29$). Observations present in ‘high-salinity’ estuaries but absent from continental-
321 scale evaluations included negative relationships of mean resistance with tree cover ($p = 0.011$; R^2
322 $= 0.46$), and negative relationships to N:P ($p < 0.0001$; $R^2 = 0.81$). Additionally, mean resistance
323 in both low- and high-salinity groups was positively related to population density ($p = 0.051$; $R^2 =$
324 0.16 , and $p = 0.0002$; $R^2 = 0.73$, respectively), and negatively related to water temperature ($p =$
325 0.001 ; $R^2 = 0.38$, and $p = 0.002$; $R^2 = 0.6$, respectively) and Chl-*a* ($p = 0.0002$, $R^2 = 0.46$, and p
326 $= 0.044$; $R^2 = 0.32$, respectively). The temperature and Chl-*a* trends were consistent with
327 continental-scale observations. Generally, ‘low-salinity’ estuaries showed fewer relationships of
328 resistance with physicochemical factors, LULC, and population density compared to ‘high-
329 salinity’ estuaries.

330

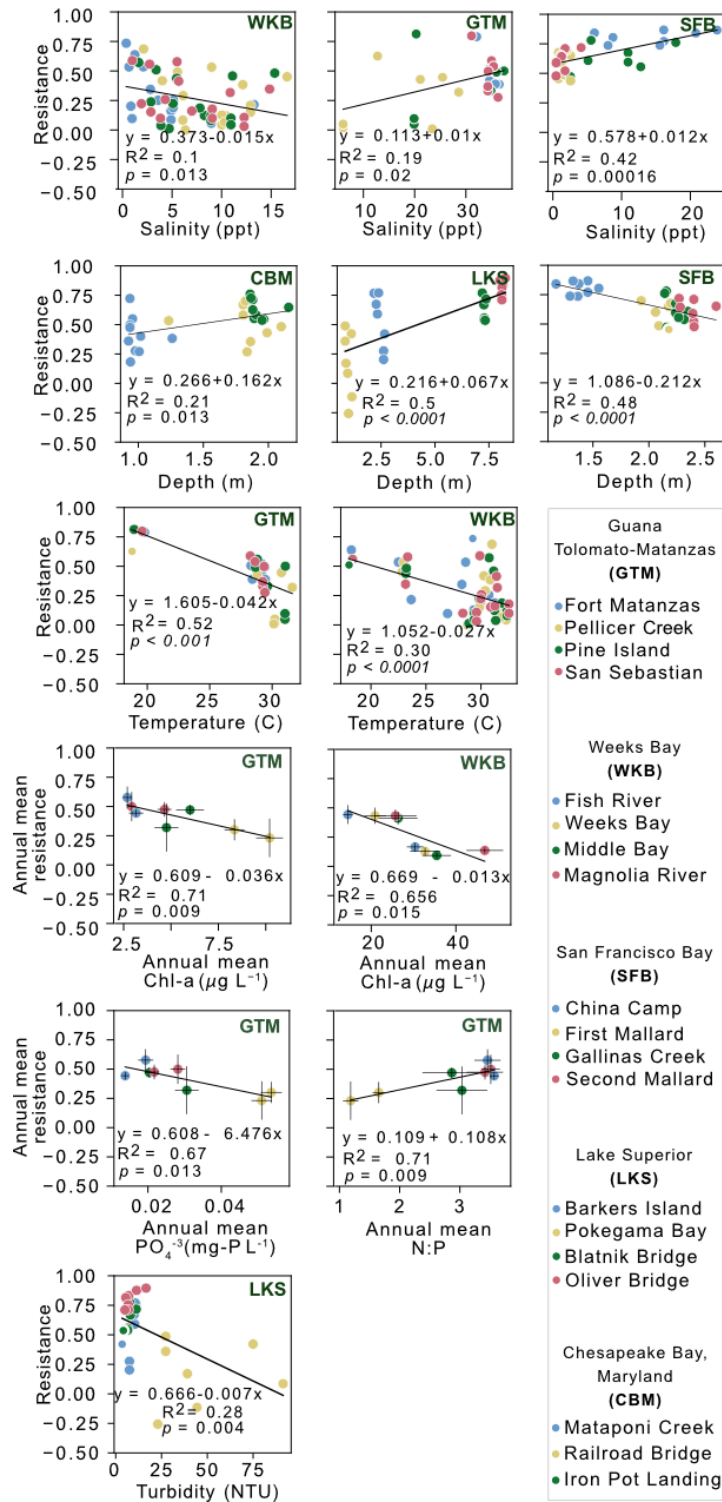


331
 332 **Figure 3.** Relationships of continental-scale and salinity-based resistance with physicochemical
 333 factors, land use/land cover, and population density. Continental-scale regressions considered all
 334 monitoring locations across all estuaries. Significant relationships ($p < 0.05$) are shown in green,
 335 red, and blue for continental-scale, high-salinity and low-salinity estuaries, respectively. Standard
 336 errors of the mean are shown in vertical and horizontal lines. The LULC parameters and
 337 population density were used from within the 10-km radius adjoined to the monitoring locations.
 338 Please refer to Table S4 for resulting statistics and to Fig. S8 for information on covariance among
 339 predictor variables.

340
 341 At local scales (i.e., within each estuary), resistance was related to some physicochemical
 342 factors that were not observed in continental-scale or salinity-based relationships. Moreover, the

343 number, strength, and trends of relationships between local-scale resistance and physicochemical
344 factors varied substantially across estuaries (Fig. 4, Figs. S9, S10). Resistance at GTM had the
345 most relationships to physicochemical factors. It was negatively related to water temperature ($p <$
346 0.001 ; $R^2 = 0.52$), PO_4^{-3} ($p = 0.013$; $R^2 = 0.67$) and Chl-*a* concentrations ($p = 0.009$; $R^2 = 0.71$);
347 and positively related to salinity and N:P ($p = 0.02$, $R^2 = 0.19$ and $p = 0.009$, $R^2 = 0.71$,
348 respectively). In contrast, at CBM, resistance was related only to water column depth (positive, p
349 $= 0.013$; $R^2 = 0.21$). At LKS, resistance was positively related to water column depth ($p < 0.0001$;
350 $R^2 = 0.5$), and negatively related to turbidity ($p = 0.004$; $R^2 = 0.28$). At WKB, relationships of
351 resistance with water temperature, salinity, and Chl-*a* concentrations were all negative ($p < 0.0001$,
352 $R^2 = 0.30$; $p = 0.013$, $R^2 = 0.1$; and $p = 0.015$, $R^2 = 0.66$, respectively). At SFB, resistance was
353 negatively related to water column depth ($p < 0.0001$; $R^2 = 0.48$), which opposed the general trend,
354 and positively related to salinity ($p = 0.00016$; $R^2 = 0.42$). There was no overarching relationship
355 between resistance and total precipitation amount of each event, except for a significant but weak
356 negative relationship at GTM estuary ($p = 0.026$, $R^2 = 0.18$, Fig. S11).

357



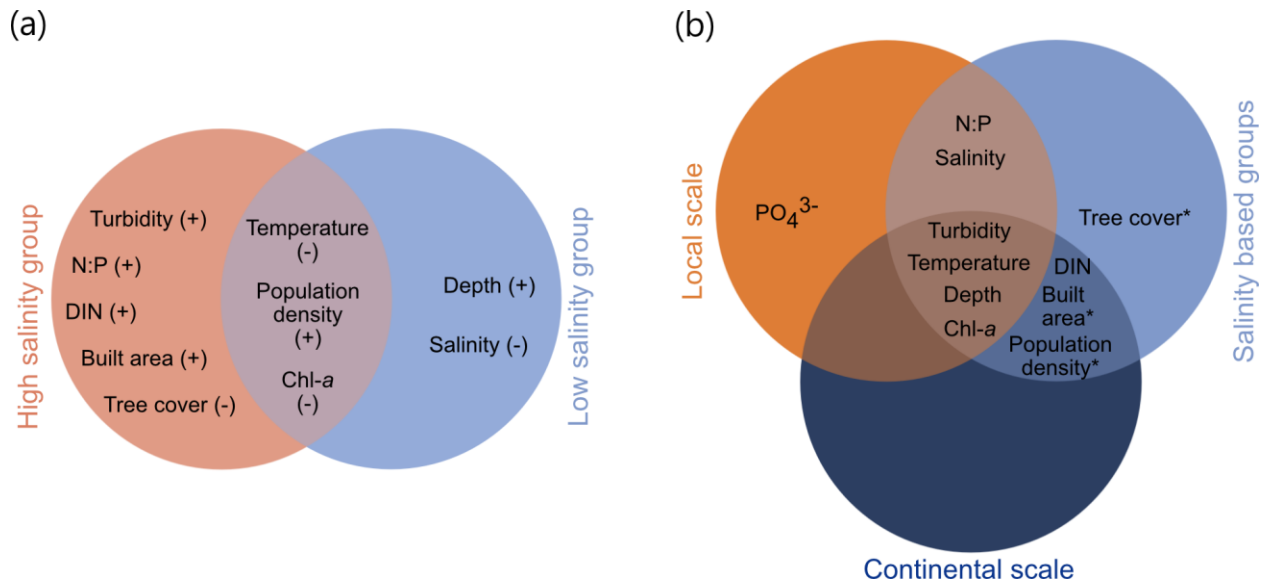
358

359 **Figure 4.** Relationships between resistance and physicochemical factors for each estuary. National
 360 Estuarine Reserve System (NERR) estuary abbreviations: Lake Superior (LKS) NERR,
 361 Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR, Weeks Bay
 362 (WKB) NERR, and San Francisco Bay (SFB) NERR. Significant correlations ($p < 0.05$) are shown

363 with black lines. Standard errors of the mean are shown in vertical and horizontal black lines for
 364 relationships using annual means for chlorophyll-*a* (Chl-*a*), PO₄³⁻, and N:P. For additional results
 365 see Figs. S9, S10, and S12.

366
 367 In summary, some relationships of estuarine resistance with physicochemical factors and
 368 urban land use appeared to be more universal while others varied across scales (Fig. 5). For
 369 instance, resistance was related to water temperature, water column depth, turbidity, and Chl-*a*
 370 across continental- and local-scales, and within salinity-based estuary groups. Dissolved inorganic
 371 N, percent built area, and population density were all related to resistance at the continental scale
 372 and in salinity-based groups. Both local and salinity-based evaluations revealed correlations of
 373 N:P, salinity, and turbidity with resistance to precipitation. Unique relationships included PO₄³⁻ at
 374 the local scale (GTM only) and tree cover in ‘high-salinity’ estuaries.

375



376
 377 **Figure 5.** Cross-scale relationships of estuarine resistance with physicochemical factors and land-
 378 use characteristics. a) Venn diagram of resistance relationships to physicochemical factors and
 379 land-use characteristics in high- vs. low-salinity estuaries. Positive or negative relationships are
 380 indicated with ‘+’ and ‘-’ respectively. b) Venn diagram of resistance relationships with
 381 physicochemical and land-use factors in continental, local, and salinity-based groups. Estuarine
 382 resistance with land use/land cover and population density marked with asterisks (*) were not
 383 evaluated at local-scale due to overlap in the spatial domains of some monitoring locations.

384

385 4. Discussion

386 Understanding patterns in estuarine responses to precipitation is important for predicting
387 the impacts of urbanization and climate change on estuaries as a whole. Previous studies have
388 shown that patterns identified at large scales may not be applicable across different climatic
389 conditions, regional geology, and other ecosystem factors (Baker et al., 2004; Gannon et al., 2022;
390 Hopkins et al., 2015; Kaushal et al., 2018; Ombadi and Varadharajan, 2022; Poff et al., 2006). Our
391 results underscore the importance of cross-scale evaluations that can elucidate commonalities in
392 estuarine response to precipitation, as well as variability in the factors associated with resistance
393 to precipitation across individual estuaries.

394 We show that while relationships between estuarine resistance, urbanization, and DIN may
395 prevail at the continental-scale, they may not correspond to resistance in individual estuaries. This
396 is because local resistance is associated with myriad specific factors in addition to many factors
397 identified at larger scales. Also, in contrast to our overarching hypothesis, we find that urbanized
398 estuaries tend to have higher resistance than more pristine estuaries. These results could suggest
399 that the effects of watershed urbanization may impact estuarine stability by either providing some
400 mechanism that allows estuaries responding to major precipitation events to dampen large shifts
401 in DO; or by disturbing the baseline DO to an extent where even a major precipitation event would
402 not produce a significant shift in DO, making the system appear highly resistant.

403 404 **4.1 Dissolved oxygen dynamics differ in estuaries with urban or agricultural land use/land** 405 **cover**

406 There were vast differences in geometry, circulation, and hydrologic conditions among
407 urbanized estuaries versus estuaries surrounded by agricultural land in this study. Yet, they
408 exhibited similar patterns in DO response to precipitation. Precipitation generally reduced DO
409 concentrations at WKB, SFB, and CBM (Fig. 2a–e). Simultaneously, the estuary with the least
410 amount of urbanized LULC – LKS – overall, experienced an increase in DO concentration
411 following precipitation.

412 While a wide range of physical factors can impact DO in estuaries including channel
413 geometry and river discharge (Kemp and Boynton, 1980; Raymond et al., 2012; Raymond and
414 Cole, 2001), wind (Scully, 2010; Zheng et al., 2024), and circulation (Raimonet & Cloern, 2017),
415 urban estuaries in particular are often impacted by a combination of these processes. Urban

416 estuaries often serve as basins for wastewater treatment outflows, which can supply continued
417 freshwater discharge and nutrients when river discharge is low. Both urban and agriculturally
418 influenced estuaries are prone to increased nutrient loading during and shortly after precipitation
419 events (Chapin et al., 2004; Costanzo et al., 2003; Mallin et al., 2009), which impacts primary
420 production and microbial metabolism and may lead to declines in DO concentrations (e.g., algal
421 blooms).

422 Therefore, closely examining the directionality of change in DO response to precipitation
423 is important for informing the environmental significance of storm events for a specific system.
424 An overall increase in DO following precipitation could reduce hypoxia through processes like
425 reaeration of the water column (Bianucci et al., 2018; Bohórquez-Bedoya et al., 2024) and have
426 positive effects on nutrient cycling (Harris et al., 2015) and growth conditions for macro- and
427 micro- fauna. A decrease in DO following precipitation may reflect the opposite effects (e.g.,
428 hypoxia), which are associated with changes in microbial community structure, disruptions in
429 nutrient cycles, and fish die-off (Diaz et al., 1992; Llansó, 1992). Further, the degree to which
430 ecosystem health is associated with DO changes is intertwined with its original ecological state
431 and with the magnitude of change in DO. An increase in DO following a storm may be less
432 consequential for estuaries with DO-replete ambient conditions as compared to estuaries with
433 chronically depleted DO. This implies that to more fully assess the effects of storms on estuarine
434 ecosystem stability and health, estuarine responses should be evaluated both through the lens of
435 the resistance index, historical DO conditions and through the magnitude and directionality of DO
436 change.

437 For example, resistance index values for the more pristine and tree-cover dominated
438 Pokegama Bay at LKS was comparable to most monitoring stations at WKB, which were
439 associated with agriculture; however, Pokegama Bay and WKB experienced different patterns in
440 DO change following precipitation (Figs. 2, S4). While the full nature of the causes and impacts
441 of DO change in each estuary is beyond the scope of this study, the differences in absolute DO
442 change suggests that different underlying processes influence stability in estuarine dynamics in
443 response to precipitation across different estuaries. Our goal here is to unravel common factors
444 that are able to predict estuarine resistance at the continental-scale, despite this variation in
445 processes, as well as to reveal specific predictors in estuaries with contrasting salinity and in

446 individual estuaries to help elucidate certain dynamics that may be more impactful in one estuary
447 relative to another.

448

449 **4.2 Urbanization and inorganic nitrogen correspond with elevated resistance to precipitation** 450 **at the continental-scale.**

451 Higher resistance to precipitation in the most urbanized estuaries and overarching
452 relationships of urban LULC with resistance across all estuaries reflect that estuaries within urban
453 watersheds typically experience relatively small shifts in DO following precipitation (Figs. 2-3).
454 These results contradict our hypothesis that urban estuaries should show low resistance to
455 precipitation because of greater physical and chemical disturbances (flashiness, streambed
456 scouring, N loading, turbidity, etc.) (Bernhardt et al., 2008; Groffman et al., 2004; Hession et al.,
457 2003; Hopkinson & Vallino, 1995; Walsh et al., 2005).

458 It is possible that watershed urbanization could equip estuaries with adaptations that help
459 dampen precipitation impacts on shifts in DO. For instance, an increase in flashiness in waterways
460 would increase flow velocity and reaeration of the water column in estuaries (Raymond et al.,
461 2012; Raymond and Cole, 2001) and contribute to phytoplankton removal via transport or
462 turbidity-driven light attenuation (Caffrey, 2004; Pennock and Sharp, 1986). This impact may be
463 particularly important for estuaries whose baseline conditions are influenced by algal blooms,
464 which induce large DO fluctuations by overproducing DO during the day and severely depleting
465 DO at night (Chapin et al., 2004; Ni et al., 2020). Such control for overgrowth of phytoplankton
466 would help maintain DO near baseline. Supporting this explanation, turbidity-driven limitations
467 on phytoplankton were previously reported for the SFB estuary (Cloern, 1987). We also found
468 positive and negative relationships of turbidity and Chl-*a*, respectively, with built area and
469 population density (Fig. S13).

470 Relatively high levels of ambient N and low ambient DO in urban estuaries may also
471 facilitate comparatively low DO fluctuations in response to precipitation. Watershed urbanization
472 is often related to high N export to estuarine environments including in the San Francisco Bay area
473 (Bettez et al., 2015; Dugdale et al., 2007; Hopkinson and Vallino, 1995; Parker et al., 2012;
474 Reisinger et al., 2016). Here, estuarine inorganic N concentrations were related to urbanization
475 and also were a major predictor of resistance at the continental-scale (Figs. 3 and S13). While
476 moderate levels of N support biological metabolisms (Howarth, 1988; Howarth and Marino, 2006;

477 Vitousek and Howarth, 1991; Zhang et al., 2021), nutrient loading is a significant problem for
478 urban aquatic environments (Beman et al., 2005; Bernot et al., 2010; Black et al., 2011; Mulholland
479 et al., 2008). For urban estuaries, where N loading is associated with chronically low DO
480 availability, this could mean that precipitation-driven increases in DO under simultaneous high N
481 input would be small, which would result in higher resistance. Nitrogen limitation, in contrast, is
482 prevalent in coastal marine systems and can negatively impact overall estuarine health by
483 constraining biological activity (Elser et al., 2007; Guildford and Hecky, 2000; Paerl, 2018; Paerl
484 and Piehler, 2008). Relationships of resistance with N and urbanization were particularly evident
485 in more coastal (high-salinity) estuaries where we found positive associations of resistance with
486 N:P and built area (Fig. 3). Thus, altered N dynamics in urban estuaries may be key to
487 understanding their comparatively low change in response to precipitation.

488 Alternatively, we also consider if baseline DO concentrations in urban estuaries are already
489 disturbed to such an extent that precipitation events cause minimal further disruption. Urbanization
490 itself can promote large diel DO fluctuations, as suggested by (Gold et al., 2020), and if a
491 precipitation event induces fluctuations that are similar in magnitude, post-precipitation DO
492 concentrations will not deviate significantly from baseline conditions. Under such a scenario,
493 urban estuaries will appear more resistant to precipitation events compared to more natural
494 estuaries.

495 **4.3 Water column depth, temperature, turbidity, and Chl-*a* are related to resistance across** 496 **all scales.**

497 We found generalizable patterns (across all scales), in which resistance was mostly
498 positively related to water column depth and turbidity, and negatively related to water temperature
499 and Chl-*a* (Figs. 3–5). As discussed above, phytoplankton dynamics may be an important factor
500 in regulating DO in response to disturbance due to their tight linkage with N loading, which is
501 consistent with results previously reported by others (Thayne et al., 2023). This is further
502 underscored by the existence of a relationship between Chl-*a* and resistance across continental,
503 salinity-based, and local analyses. In parallel, the relationship between resistance and turbidity was
504 positive at the continental scale and in the high salinity group (Fig. 3), while at LKS resistance and
505 turbidity correlated negatively (Figs. 3, 4). The positive associations of resistance with turbidity
506 and simultaneous negative relationship with Chl-*a* can be attributed to turbidity-driven light

507 attenuating conditions that restrict phytoplankton growth, which is further confirmed through
508 negative correlation between turbidity and Chl-*a* (Fig. S8).

509 Water column depth can also influence the calculated resistance of estuaries to
510 precipitation through physical and hydrologic impacts on DO. Through dilution, deeper estuaries
511 could have a greater capacity to resist hydrologic changes in response to precipitation, for instance
512 by attenuating the influx of oxygen-saturated rain water; nutrient loading; turbulence; and water-
513 atmosphere gas exchange. Similarly, deep estuaries have longer equilibration time with
514 environmental conditions and less diel variability in parameters like temperature (Caissie, 2006;
515 Macan, 1958), which also affects diel and seasonal DO dynamics, possibly resulting in more stable
516 DO baselines and more moderate responses to precipitation. Deeper estuaries can also be
517 associated with urbanization, which also shows high resistance in this study, because shifts towards
518 more urban and agricultural LULC generally deepen estuarine channels through increased runoff,
519 sediment transport, hydrological flashiness, channel dredging, and/or other anthropogenic
520 activities (O'Driscoll et al., 2010; Simon and Rinaldi, 2006; Walsh et al., 2005). Therefore, water
521 column depth appears to be a critical factor in regulating the resistance to precipitation in estuaries.

522 Lastly, because temperature controls various chemical and biological processes that impact
523 DO availability (e.g., microbial growth, oxygen solubility), many studies have focused on the
524 impact of rising temperature on DO dynamics in estuaries (Apple et al., 2006; Caffrey, 2003;
525 Caffrey et al., 2014). Highlighted results link climate change to thermal pollution of aquatic
526 systems following rain events (Zahn et al., 2021) and show connections of elevated global
527 temperatures with decreased primary production (Song et al., 2018). As such, estuaries with
528 elevated ambient temperatures may have a decreased capacity to resist disturbances to DO
529 dynamics relative to estuaries with more moderate temperatures.

530

531 **4.4 High variability in factors associated with resistance at local scales.**

532 The contrasting relationships between physicochemical factors and resistance across
533 monitoring locations in individual estuaries highlight substantial fine-scale variation in
534 precipitation responses (Fig. 4). For example, resistance at CBM is related to one factor (water
535 column depth), while at GTM, resistance is related to five factors. The variability in factors
536 associated with resistance suggests that individual estuaries may need to consider factors beyond
537 water temperature, water column depth, turbidity, and Chl-*a* in water-quality management and

538 conservation strategies. For instance, Chl-*a* and dissolved inorganic phosphorus have been shown
539 to respond to storms in some estuaries more than in others due to differences in light limitation,
540 grazing, and nutrient concentrations (Chen et al., 2015; Cloern, 2001; Cloern and Jassby, 2010;
541 Dix et al., 2013, 2008; Liao et al., 2021; Zhang et al., 2022). Likewise, water temperature and
542 salinity also have variable responses to precipitation (Buelo et al., 2023; Chen et al., 2015; Dix et
543 al., 2008), leading to differences in biological processes and phytoplankton activity in different
544 estuaries (Apple et al., 2008). While there are myriad potential estuary-specific interactions that
545 may result in different responses to stressors, our results highlight that system-variability is
546 important when identifying parameters involved in estuarine resistance to precipitation. We also
547 acknowledge that other scales of investigation (e.g., regional) could introduce additional insight
548 into the predictors of estuarine resistance. Moreover, system dependencies on groundwater
549 discharge as a driver of nutrient dynamics and control on DO dynamics (Brookfield et al., 2021;
550 Kornelsen and Coulibaly, 2014) as well as system dependencies on microbial community structure
551 (Cheng et al., 2021) should also be appraised when evaluating resistance of urban estuaries to
552 precipitation.

553

554 **5. Conclusions**

555 In light of increasing urbanization and emerging climatic scenarios, cross-scale evaluations
556 of the responses of estuaries to precipitation events are imperative for developing effective
557 management strategies. We find that urban estuaries appear to be more resistant to changes in DO
558 in response to precipitation than more pristine estuaries; and that the depth of the water column,
559 water temperature, turbidity, and Chl-*a* are generalizable cross-scale predictors for estuarine
560 resistance. However across different scales, we find that system variability results in additional
561 factors that are important to consider when managing the responses to major precipitations events
562 of individual estuaries. Based on our results, we suggest that investigations targeting N, nutrient
563 delivery mechanisms, and microbial activity could help understand what makes urban estuaries
564 more resistant to large shifts in DO following precipitation. We also highlight that high-resolution
565 water quality and nutrient data surrounding precipitation events, along with careful considerations
566 of local variabilities and models for system responses to precipitation, are needed to help elucidate
567 the underlying mechanisms for high resistance of urban estuaries. This study serves as a platform

568 for improvement of guidelines and predictive capabilities addressing system response to future
569 climatic and urbanization scenarios.

570

571 **Data availability:** This study used publicly available datasets, which included: 1) Long-term
572 estuarine water quality, nutrients and meteorological conditions and watershed boundaries for
573 Lake Superior (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato
574 Matanzas (GTM) NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB) NERR stations
575 from <https://cdmo.baruch.sc.edu>; 2) Long-term precipitation data from U.S. airports from
576 <https://www.ncei.noaa.gov/cdo-web/datasets>; 3) Land use/land cover maps from
577 <https://livingatlas.arcgis.com/landcover/>; 4) U.S. population data from
578 <https://www.worldpop.org/>. Data and data processing code are also available at
579 <https://figshare.com/s/49d2f3dca084d885638a>.

580

581 **Author Contribution:** E.B.G. and A.B.T. developed the study and interpreted the results. A.B.T.
582 performed data analysis and drafted the manuscript. All authors contributed to manuscript editing.

583

584 **Competing interest:** The authors declare that they have no competing interest.

585

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597

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966

967 **Supplementary Tables.**

968

969 **Table S1:** Multiparameter sonde position within the water column at five National Estuarine
 970 Research Reserve Systems (NERRS).

Estuary	Monitoring location	Sonde position above sediment bed (m)
Lake Superior (LKS)	Barker’s Island (BA)	0.5
	Pokegama Bay (PO)	0.25
	Blatnik Bridge (BL)	5.5
	Oliver Bridge (OL)	6.5
Chesapeake Bay, Maryland (CBM)	Iron Pot Landing (IP)	0.25
	Railroad Bridge (RR)	0.25
	Mataponi Creek (MC)	0.25
Weeks Bay (WKB)	Magnolia River (MR)	0.5
	Middle Bay (MB)	0.5
	Weeks Bay (WB)	0.5
	Fish River (FR)	0.5
Guana Tolomato-Matanzas (GTM)	San Sebastian (SS)	1.0
	Pine Island (PI)	1.0
	Pellicer Creek (PC)	1.0

	Fort Matanzas (FM)	1.0
San Francisco Bay (SFB)	First Mallard (FM)	0.25 - 0.5
	Second Mallard (SM)	0.25 - 0.5
	Gallinas Creek (GC)	0.25 - 0.5
	China Camp (CC)	0.25 - 0.5

971

972 **Table S2.** Datetime and threshold values for major precipitation events, and breakdown of the
973 number of events during wet and dry years at each National Estuarine Research Reserve (NERR)
974 station.

NERR Station	Datetime used to select dissolved oxygen and precipitation measurements to calculate the resistance index		Precip. event (mm)	Wet/Dry year	precip. threshold (mm/day)
	Prior to disturbance (C ₀)	During and post disturbance (P ₀)			
Lake Superior, WI (LKS)	2017/06/24 00:00:00 -2017/06/27 23:45:00	2017/06/28 06:00:00 -2017/06/29 23:45:00	34.4	Wet (4)	> 25
	2017/07/30 00:00:00 -2017/08/02 23:45:00	2017/08/03 00:00:00 -2017/08/04 23:45:00	38.1		
	2017/08/23 00:00:00 -2017/08/25 23:45:00	2017/08/26 00:00:00 -2017/08/28 23:45:00	61.0		
	2017/09/29 00:00:00 -2017/10/01 23:45:00	2017/10/02 00:00:00 -2017/10/05 23:45:00	44.2		
	2020/07/14 00:00:00 -2020/07/17 23:45:00	2020/07/18 00:00:00 -2020/07/19 23:45:00	42.9	Dry (3)	
	2020/07/14 00:00:00 -2020/07/17 23:45:00	2020/07/21 00:00:00 -2020/07/23 23:45:00	29.6		
	2020/08/04 00:00:00 -2020/08/06 23:45:00	2020/08/07 12:00:00 -2020/08/11 23:45:00	74.9		
	Chesapeake Bay, MD (CBM)	2018/05/10 00:00:00 -2018/05/11 12:00:00	2018/05/16 00:00:00 -2018/05/21 23:45:00	117.1	Wet (7)
2018/05/24 00:00:00 -2018/05/26 23:45:00		2018/05/27 15:00:00 -2018/05/28 23:45:00	40.2		
2018/05/24 00:00:00 -2018/05/26 23:45:00		2018/06/03 06:00:00 -2018/06/07 23:45:00	56.2		
2018/06/16 00:00:00 -2018/06/18 23:45:00		2018/06/19 12:00:00 -2018/06/25 23:45:00	41.8		
2018/07/13 00:00:00 -2018/07/16 23:45:00		2018/07/21 12:00:00 -2018/07/27 23:45:00	218.1		
2018/09/04 00:00:00 -2018/09/06 23:45:00		2018/09/09 00:00:00 -2018/09/10 23:45:00	49.1		
2018/09/13 00:00:00 -2018/09/15 23:45:00		2018/09/23 00:00:00 -2018/09/25 23:45:00	58.6	Dry (3)	
2016/06/24 00:00:00 -2016/06/27 23:45:00		2016/07/01 12:00:00 -2016/07/02 23:45:00	41.0		
2016/09/15 00:00:00 -2016/09/18 23:45:00		2016/09/19 00:00:00 -2016/09/19 23:45:00	25.2		
2016/09/23 00:00:00 -2016/09/25 23:45:00		2016/09/28 00:00:00 -2016/09/30 23:45:00	76.4		
Guana Tolomato	2017/08/19 00:00:00 -2017/08/22 23:45:00	2017/09/10 00:00:00 -2017/09/21 23:45:00	222.9	Wet (3)	*
	2017/08/19 00:00:00 -2017/08/22 23:45:00	2017/09/30 00:00:00 -2017/10/13 23:45:00	270.7		

Matanzas, FL (GTM)	2017/11/15 00:00:00 -2017/11/22 23:45:00	2017/11/23 00:00:00 -2017/11/25 23:45:00	123.1		
	2016/06/01 00:00:00 -2016/06/04 23:45:00	2016/06/05 12:00:00 -2016/06/07 23:45:00	127.9	Dry (4)	
	2016/08/21 00:00:00 -2016/08/27 23:45:00	2016/08/28 00:00:00 -2016/09/06 23:45:00	67.2		
	2016/09/05 00:00:00 -2016/09/08 23:45:00	2016/09/14 00:00:00 -2016/09/19 23:45:00	27.3		
	2016/09/21 00:00:00 -2016/09/25 23:45:00	2016/09/28 00:00:00 -2016/10/17 23:45:00	193.3		
Weeks Bay, AL (WKB)	2018/05/19 00:00:00 -2018/05/22 23:45:00	2018/05/23 00:00:00 -2018/05/27 23:45:00	90.9	Wet (8)	> 30
	2018/06/07 00:00:00 -2018/06/09 23:45:00	2018/06/11 00:00:00 -2018/06/13 23:45:00	47.0		
	2018/06/28 00:00:00 -2018/06/30 23:45:00	2018/07/01 12:00:00 -2018/07/09 23:45:00	86.3		
	2018/07/12 00:00:00 -2018/07/14 23:45:00	2018/07/16 09:00:00 -2018/07/18 23:45:00	53.7		
	2018/08/25 00:00:00 -2018/08/26 23:45:00	2018/09/01 00:00:00 -2018/09/03 23:45:00	56.6		
	2018/08/25 00:00:00 -2018/08/26 23:45:00	2018/09/04 00:00:00 -2018/09/08 23:45:00	128.6		
	2018/09/19 00:00:00 -2018/09/20 23:45:00	2018/09/21 12:00:00 -2018/09/23 23:45:00	46.7		
	2018/09/19 00:00:00 -2018/09/20 23:45:00	2018/09/24 00:00:00 -2018/09/29 23:45:00	66.5		
	2019/04/01 00:00:00 -2019/04/03 23:45:00	2019/04/04 09:00:00 -2019/04/04 23:45:00	40.0	Dry (7)	
	2019/04/15 00:00:00 -2019/04/17 23:45:00	**2019/04/26 15:00:00 -2019/04/28 23:45:00	32.9		
	2019/06/01 00:00:00 -2019/06/04 23:45:00	2019/06/06 00:00:00 -2019/06/13 23:45:00	116.1		
	2019/06/20 00:00:00 -2019/06/25 23:45:00	2019/07/13 00:00:00 -2019/07/15 23:45:00	89.1		
	2019/08/07 00:00:00 -2019/08/10 23:45:00	2019/08/15 12:00:00 -2019/08/16 23:45:00	37.0		
	2019/08/22 00:00:00 -2019/08/24 23:45:00	2019/08/26 06:00:00 -2019/08/27 06:00:00	48.2		
2019/10/21 00:00:00 -2019/10/24 23:45:00	2019/10/30 00:00:00 -2019/11/01 23:45:00	61.8			
San Francisco Bay, CA (SFB)	2017/01/01 00:00:00 -2017/01/02 06:00:00	2017/01/03 00:00:00 -2017/01/06 00:00:00	38.7	Wet (5)	> 20
	2017/01/01 00:00:00 -2017/01/02 06:00:00	2017/01/07 00:00:00 -2017/01/12 23:45:00	126.4		
	2017/01/01 00:00:00 -2017/01/02 06:00:00	2017/01/18 00:00:00 -2017/01/25 00:00:00	110.8		
	2017/03/11 00:00:00 -2017/03/17 23:45:00	2017/03/20 00:00:00 -2017/03/23 23:45:00	43.0		
	2017/04/01 00:00:00 -2017/04/05 06:00:00	2017/04/06 20:00:00 -2017/04/07 23:45:00	37.0		
	2018/01/06 00:00:00 -2018/01/07 23:45:00	2018/01/08 00:00:00 -2018/01/09 23:45:00	73.6	Dry (3)	
	2018/02/20 00:00:00 -2018/02/23 23:45:00	***2018/03/01 00:00:00 -2018/03/01 23:45:00	30.7		
	2018/04/04 12:00:00 -2018/04/05 12:00:00	2018/04/06 00:00:00 -2018/04/07 23:45:00	54.6		
<p>* For GTM 2016 the selected events were: Colin, Julia, Hermine, and Matthew. For 2017 the events were Irma and two Nor'easters.</p> <p>** This resistance calculation does not include dissolved oxygen measurements during the actual rain event from 2019/04/25 only dissolved oxygen after the event, because data during the event is missing from MB monitoring location.</p> <p>*** Dissolved oxygen data for the SM site is missing. No resistance was calculated for SM during that time.</p>					

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976

977 **Table S3:** Average pre-disturbance (C_0) and post-disturbance (P_0) dissolved oxygen (mg L^{-1}) and
978 the resistance index values for each major precipitation event across all monitoring locations at
979 five estuaries. (.csv file)

980

981

982 **Table S4:** Resulting statistical parameters for linear regression analysis conducted for low- and
 983 high-salinity groups and on continental scale for five estuaries (this table accompanies Fig. 3).

Predictor variables regressed against mean resistance index values	Linear regression parameters: Coefficient of determination (R^2) Significance value (p)		
	Low salinity group	High-salinity group	Continental-scale
Turbidity (NTU)	$R^2 = 0.00$ $p = 0.933$	$R^2 = \mathbf{0.29}$ $p < \mathbf{0.0001}$	$R^2 = \mathbf{0.06}$ $p = \mathbf{0.0015}$
Salinity (ppt)	$R^2 = \mathbf{0.214}$ $p < \mathbf{0.0001}$	$R^2 = 0.01$ $p = 0.495$	$R^2 = 0.001$ $p = 0.513$
Water temperature (C)	$R^2 = \mathbf{0.30}$ $p < \mathbf{0.0001}$	$R^2 = \mathbf{0.59}$ $p < \mathbf{0.0001}$	$R^2 = \mathbf{0.39}$ $p < \mathbf{0.0001}$
Water depth (m)	$R^2 = \mathbf{0.1}$ $p < \mathbf{0.0001}$	$R^2 = 0.00$ $p = 0.997$	$R^2 = \mathbf{0.10}$ $p < \mathbf{0.0001}$
$\log(\text{DIN})$ (mg L^{-1})	$R^2 = 0.05$ $p = 0.268$	$R^2 = \mathbf{0.55}$ $p = \mathbf{0.004}$	$R^2 = \mathbf{0.12}$ $p = \mathbf{0.031}$
PO_4^{-3} (mg L^{-1})	$R^2 = 0.03$ $p = 0.416$	$R^2 = 0.06$ $p = 0.31$	$R^2 = 0.07$ $p = 0.134$
N:P	$R^2 = 0.14$ $p = 0.077$	$R^2 = \mathbf{0.81}$ $p < \mathbf{0.0001}$	$R^2 = 0.08$ $p = 0.101$
Chl- <i>a</i> ($\mu\text{g L}^{-1}$)	$R^2 = \mathbf{0.46}$ $p = \mathbf{0.0002}$	$R^2 = \mathbf{0.32}$ $p = \mathbf{0.044}$	$R^2 = \mathbf{0.39}$ $p < \mathbf{0.0001}$
Trees (%)	$R^2 = 0.01$ $p = 0.72$	$R^2 = \mathbf{0.46}$ $p = \mathbf{0.011}$	$R^2 = 0.03$ $p = 0.30$

Crops (%)	$R^2 = 0.14$ $p = 0.07$	$R^2 = 0.05$ $p = 0.48$	$R^2 = 0.04$ $p = 0.22$
Built area (%)	$R^2 = 0.03$ $p = 0.41$	$R^2 = 0.73$ $p = 0.0002$	$R^2 = 0.14$ $p = 0.02$
Population density (ppl km ⁻²)	$R^2 = 0.16$ $p = 0.051$	$R^2 = 0.73$ $p = 0.0002$	$R^2 = 0.27$ $p = 0.001$

984 Note: Significant correlations (i.e., $p < 0.05$) are shown in bold.

985 **Supplementary figures.**

986

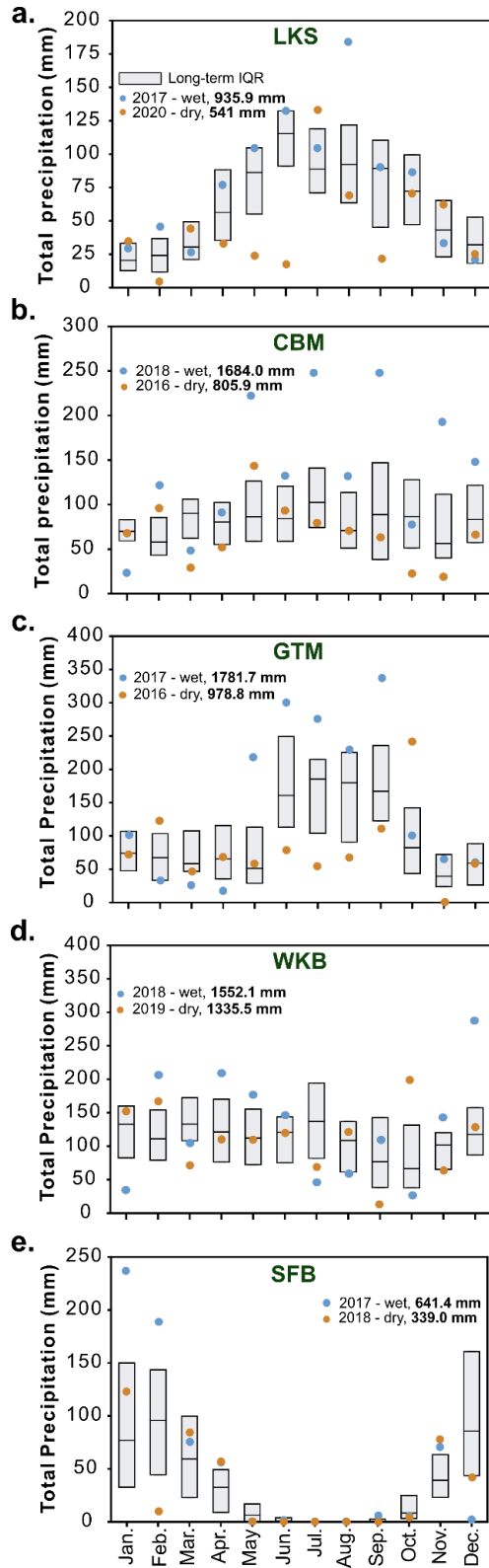
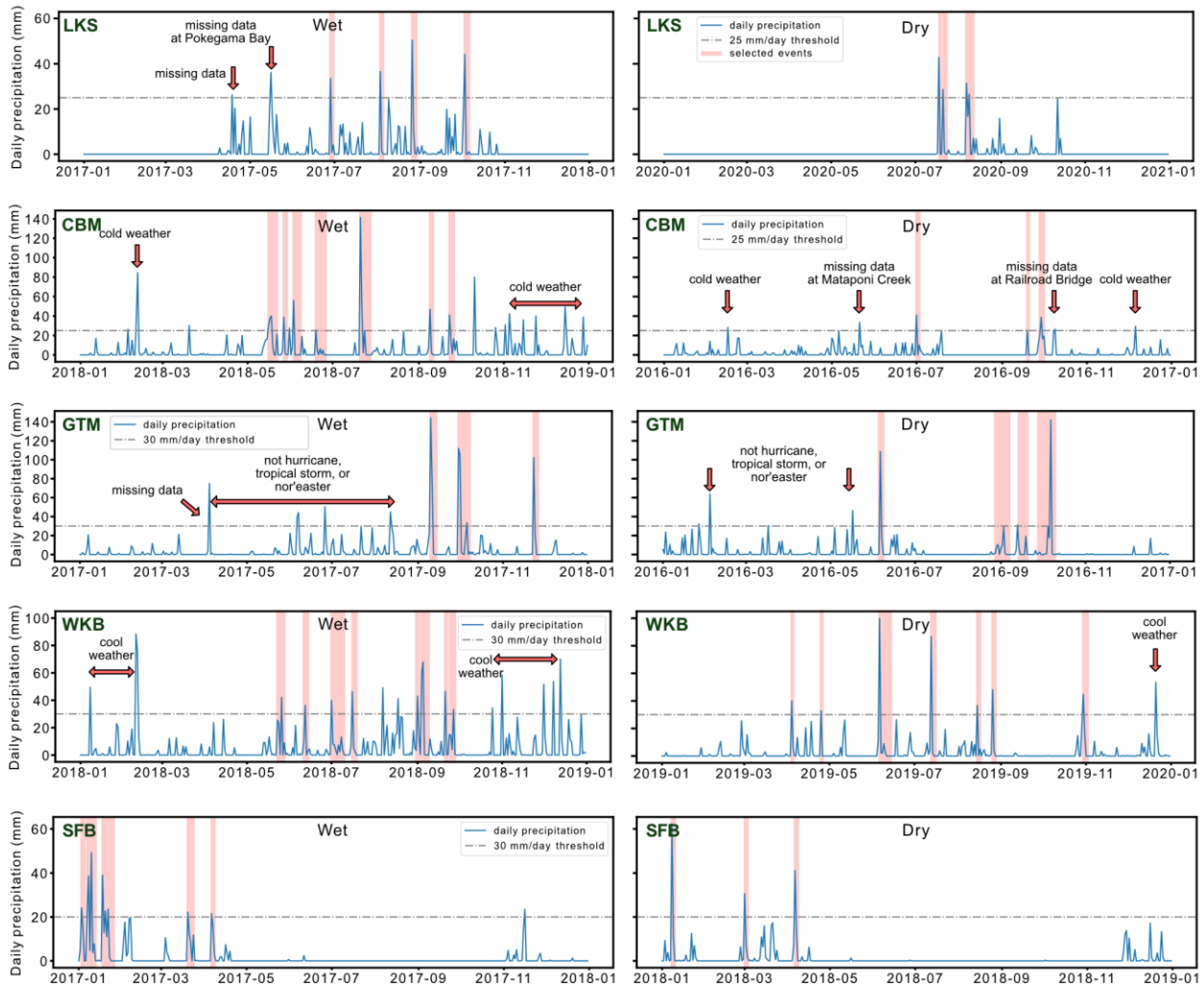


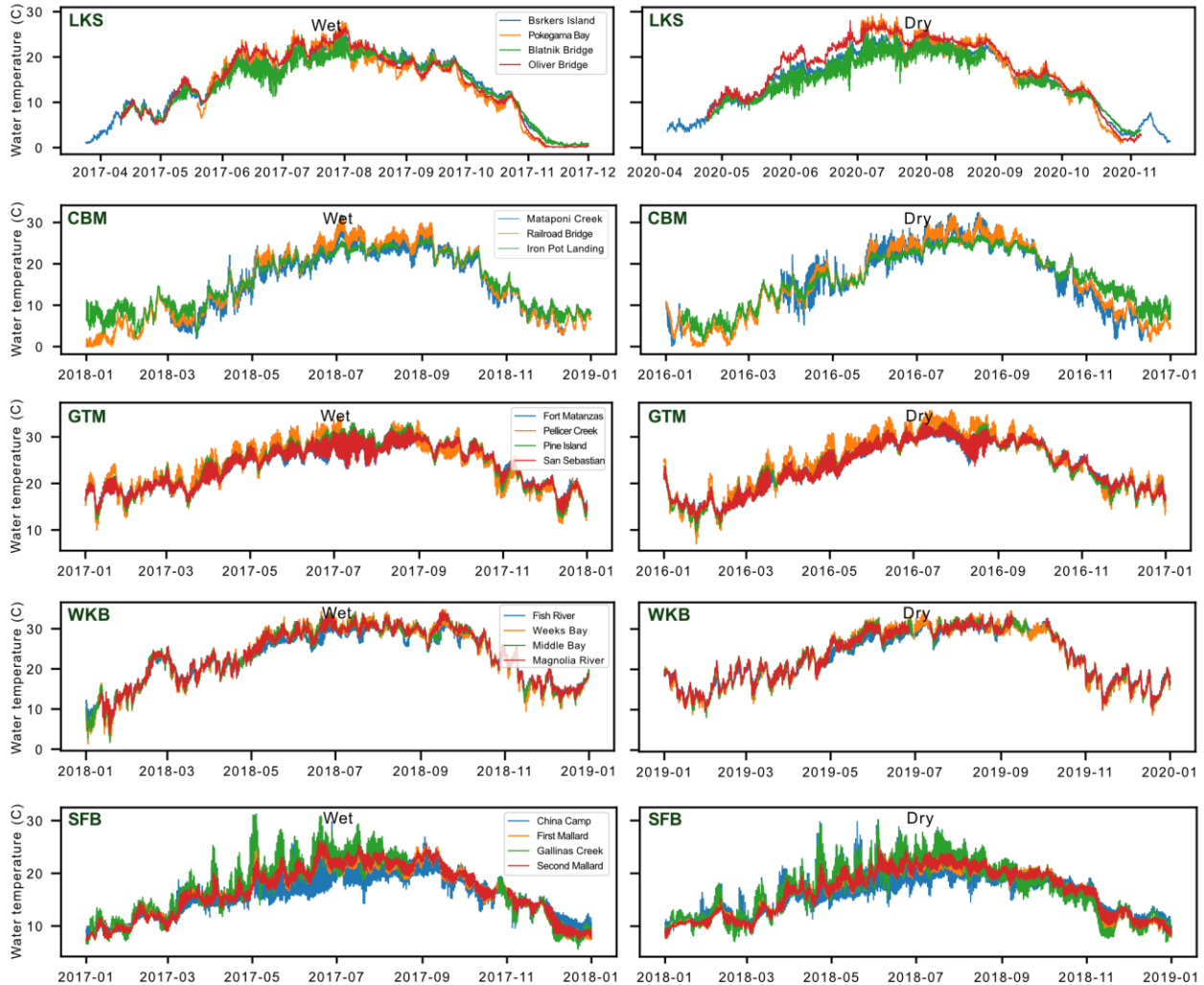
Fig. S1. Long-term (30-year) interquartile range and monthly precipitation during relatively wet and dry years. Long term precipitation data obtained from airports located in the vicinity of each selected estuarine station. a) Precipitation records from Duluth International airport were used to infer wet/dry years at Lake Superior (LKS) station. b) Washington Reagan International Airport precipitation records were used for Chesapeake Bay (CBM) station. c) Jacksonville International Airport precipitation record was used for Guana Tolomato Matanzas (GTM) station. d) Birmingham Airport precipitation records were used for Weeks Bay (WKS) station. e) San Francisco International Airport precipitation records were used for San Francisco Bay (SFB) station.

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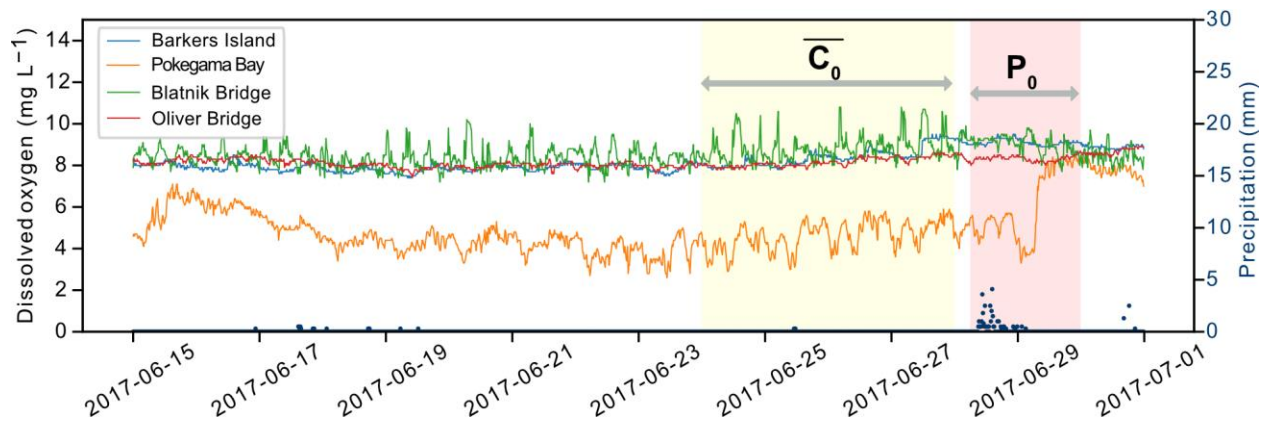
990 **Fig. S2.** Major precipitation events for each estuary during selected wet and dry years. The major
 991 precipitation events (shaded in red) identified by plotting National Estuarine Research Reserve
 992 (NERR) precipitation data collected at each estuary. The events were down-selected based on
 993 water quality data availability for each estuary, within warmer seasons, and if the event was noted
 994 in NERR metadata sheets and/or reported as hurricane, tropical storm, or nor'easter. Estuary
 995 abbreviations: Lake Superior (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana
 996 Tolomato Matanzas (GTM) NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB)
 997 NERR. For details about events used for resistance index calculations, please see Table S2.



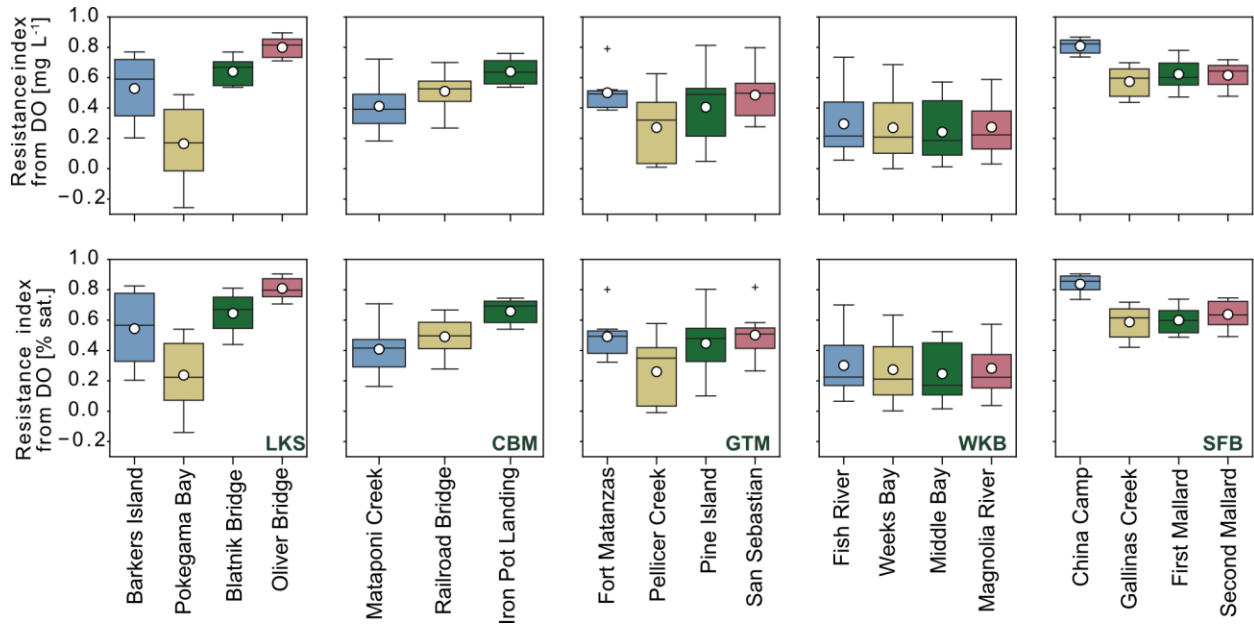
998

999 **Fig. S3.** Water temperature during selected wet and dry years for five National Estuarine Research
 1000 Reserve (NERR) estuaries. Estuary abbreviations: Lake Superior (LKS) NERR, Chesapeake Bay,
 1001 Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR, Weeks Bay (WKB) NERR,
 1002 and San Francisco Bay (SFB) NERR. Note: At LKS, the water temperature record extends from
 1003 April to December for 2017, and from April to November for 2020 because the St. Louis River
 1004 freezes over and no measurements are collected.

1005



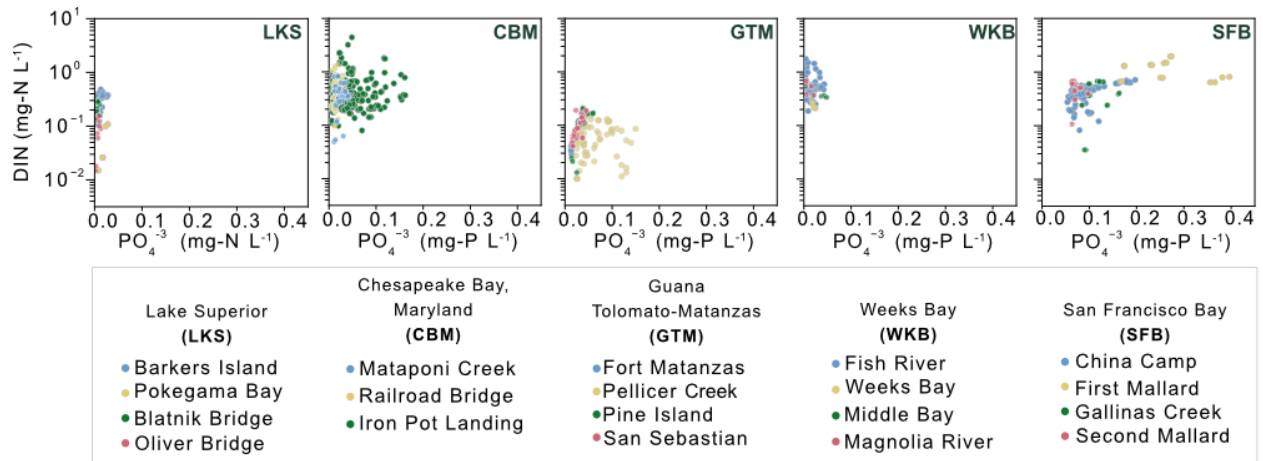
1006
 1007 **Fig. S4.** Example for visualizing the selection of variables for resistance index calculations.
 1008 Dissolved oxygen concentrations and precipitation at Lake Superior (LKS) NERR prior to and
 1009 during a precipitation event on 06-28-2017. Yellow shaded box indicates the time used to calculate
 1010 average pre-disturbance dissolved oxygen concentration (average C_0). Red shaded box indicated
 1011 the time used to identify dissolved oxygen concentration post-disturbance (P_0). P_0 was identified
 1012 as the maximum displacement from average C_0 .



1013

1014 **Fig. S5.** Resistance index at each site calculated using dissolved oxygen (DO) concentration and
 1015 DO percent saturation. (top row) Resistance index calculated using DO mg L⁻¹. (bottom row)
 1016 Resistance index calculated using DO % saturation.

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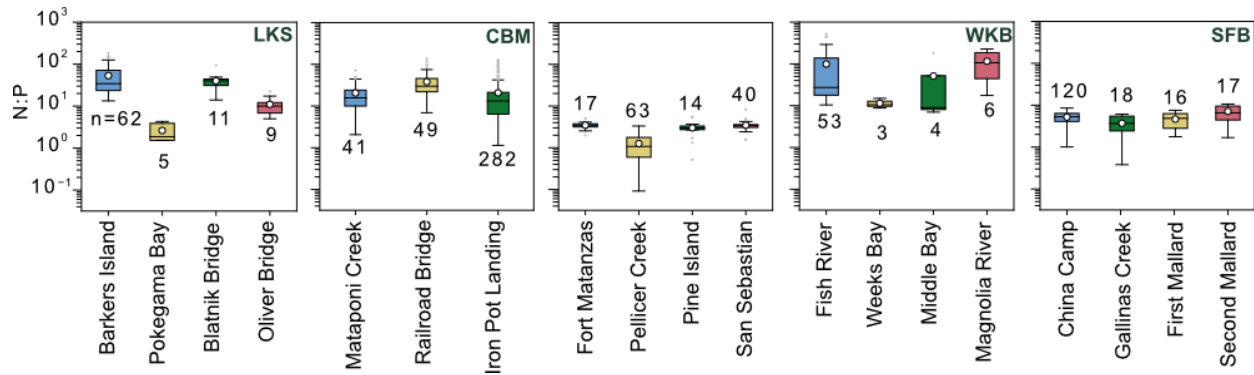
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Fig. S6. Dissolved inorganic nutrient concentrations at each estuary. Dissolved inorganic nitrogen (DIN) (i.e., NO₃⁻ + NO₂⁻ + NH₄⁺). Estuary abbreviations: Lake Superior (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB) NERR. For LKS estuary, the NH₄⁺ measurements for dry year (2020) were missing at all monitoring locations. For WKB, the PO₄³⁻ measurements for wet year (2018) were missing at WB, MB, and MR monitoring locations.

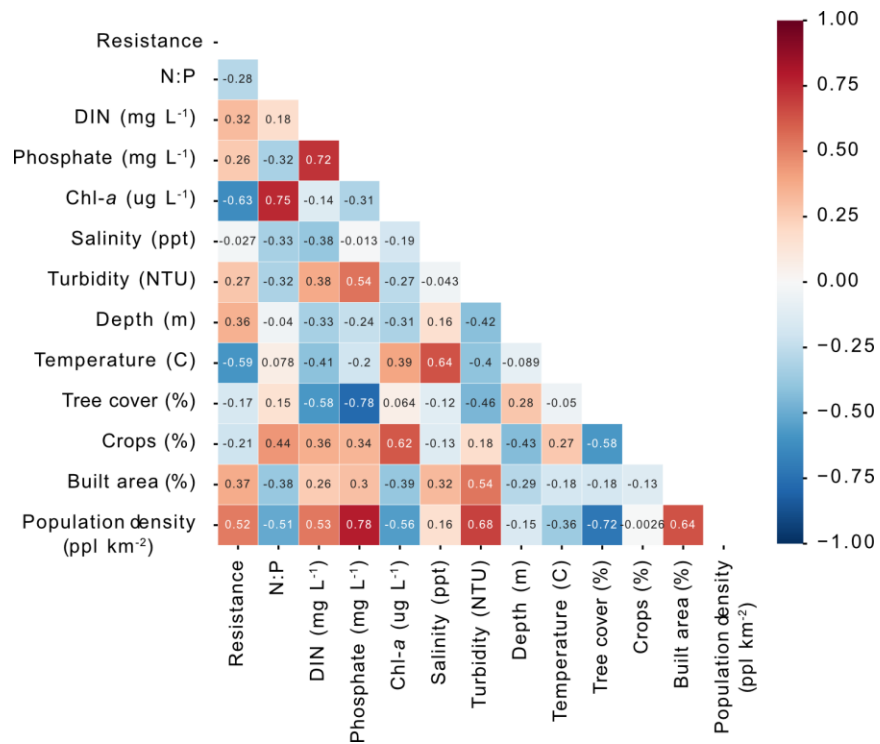
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1027 **Fig. S7.** Distribution of nitrogen to phosphorus ratio (N:P) within individual estuaries. Estuary
1028 abbreviations: Lake Superior (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana
1029 Tolomato Matanzas (GTM) NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB)
1030 NERR. Means are shown in white circles, and medians are shown in black solid lines. Boxes
1031 show the quartiles of the dataset and the whiskers show the rest of the distribution. The
1032 stoichiometric N:P was calculated using dissolved inorganic nitrogen species (i.e., $\text{NO}_3^- + \text{NO}_2^-$
1033 $+ \text{NH}_4^+$) and phosphate. We note that because of missing measurements for PO_4^{3-} during the wet
1034 year at WKB- WB, MB, and MR – the N:P at these locations were calculated only for the dry
1035 year.

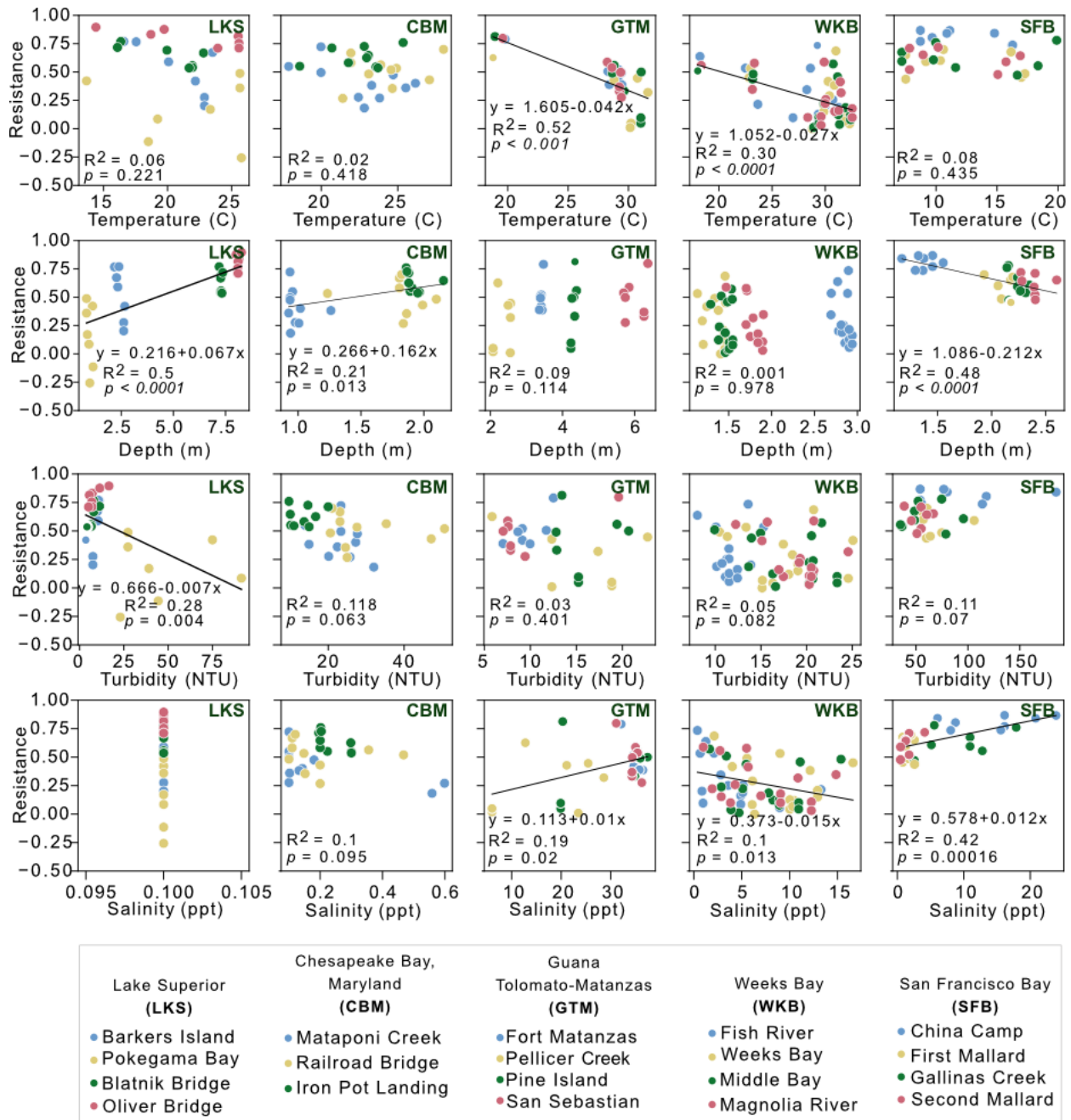
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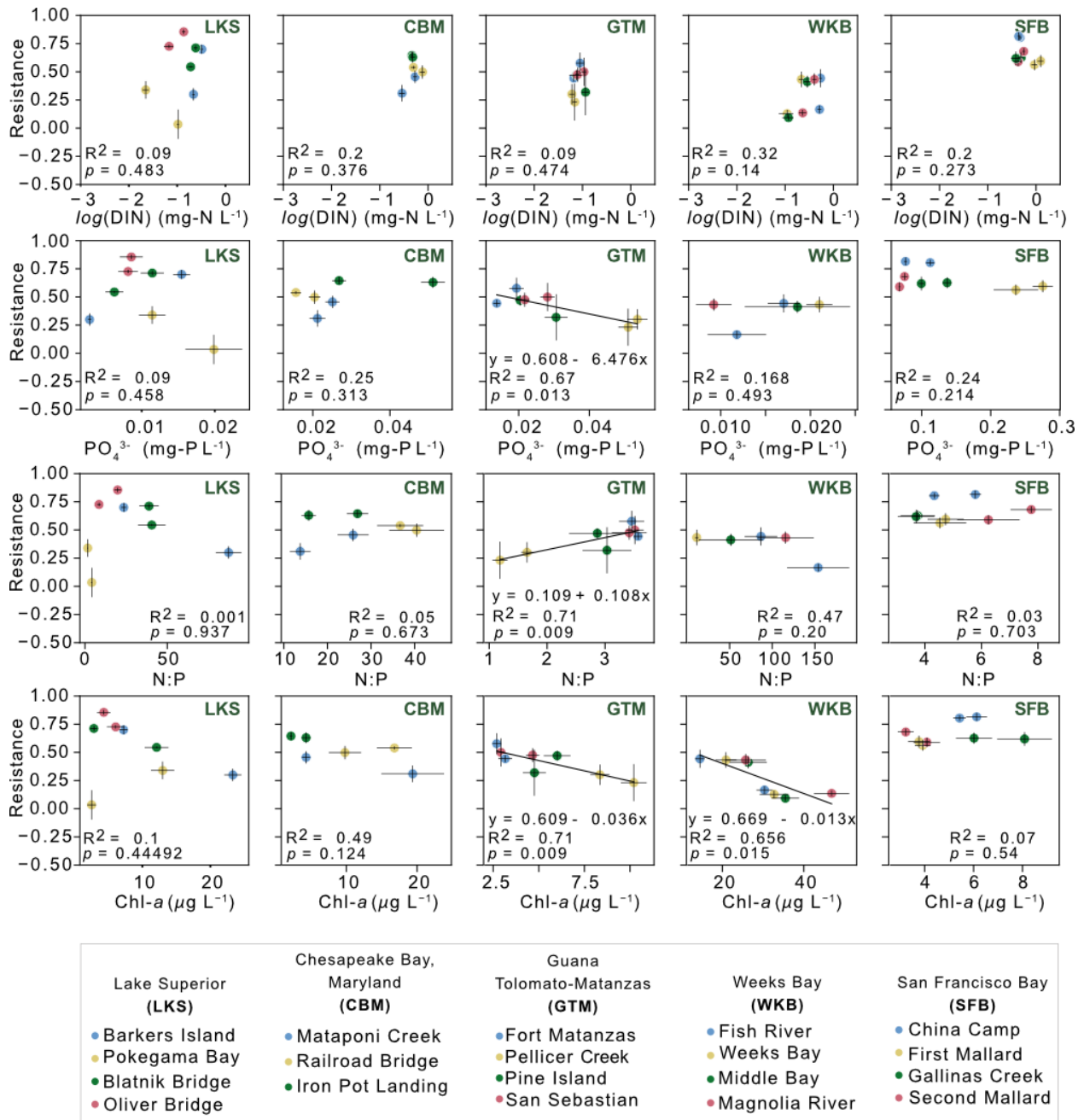
1038 **Fig. S8.** Correlation matrix for resistance, water quality parameters, and land use/land cover across

1039 combined estuaries.



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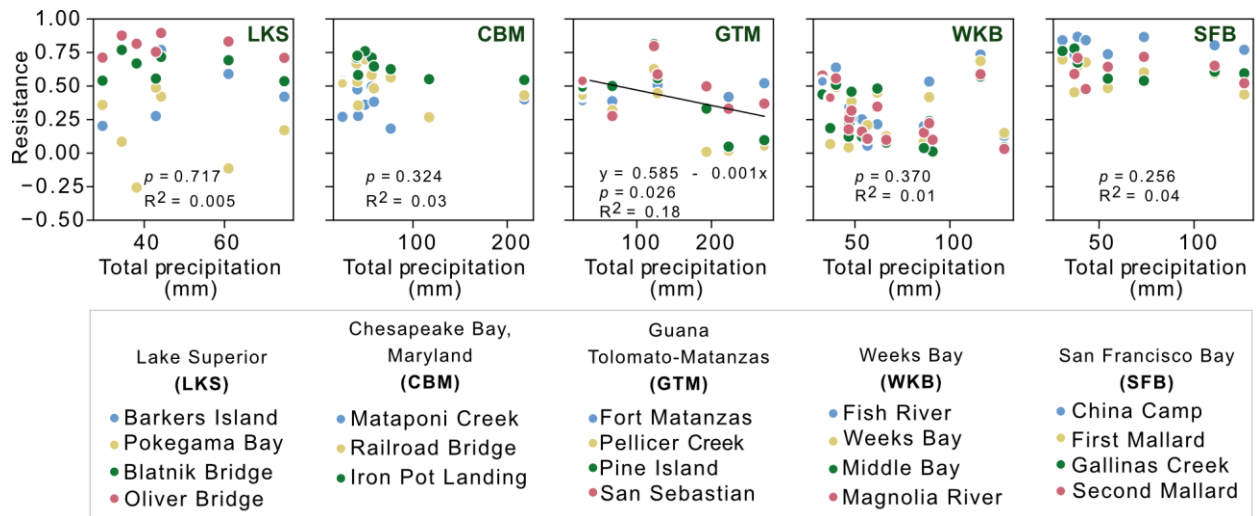
1042 **Fig. S9.** Relationships of resistance to water temperature, water column depth, turbidity, and
 1043 salinity at each estuary. Estuary abbreviations: Lake Superior (LKS) NERR, Chesapeake Bay,
 1044 Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR, Weeks Bay (WKB) NERR,
 1045 and San Francisco Bay (SFB) NERR. Significant correlations ($p < 0.05$) are shown in black lines.
 1046 Relationships consider mean values of physicochemical factors in context of precipitation events
 1047 used in resistance index calculations (see dates in Table S2).



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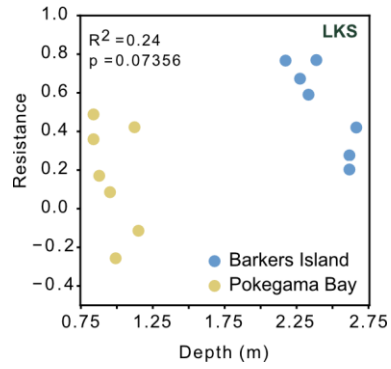
1050 **Fig. S10.** Relationships of resistance to dissolved inorganic nitrogen (DIN), phosphate (PO₄³⁻),
 1051 N:P, and chlorophyll-*a* (Chl-*a*) at each estuary. Estuary abbreviations: Lake Superior (LKS)
 1052 NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM) NERR,
 1053 Weeks Bay (WKB) NERR, and San Francisco Bay (SFB) NERR. Significant correlations ($p <$
 1054 0.05) are shown in black. Standard errors of the mean are shown in horizontal and vertical black

1055 lines. Relationships consider annual mean values of resistance and annual mean values for
1056 nutrients, N:P, and Chl-*a* calculated for wet and dry years separately.



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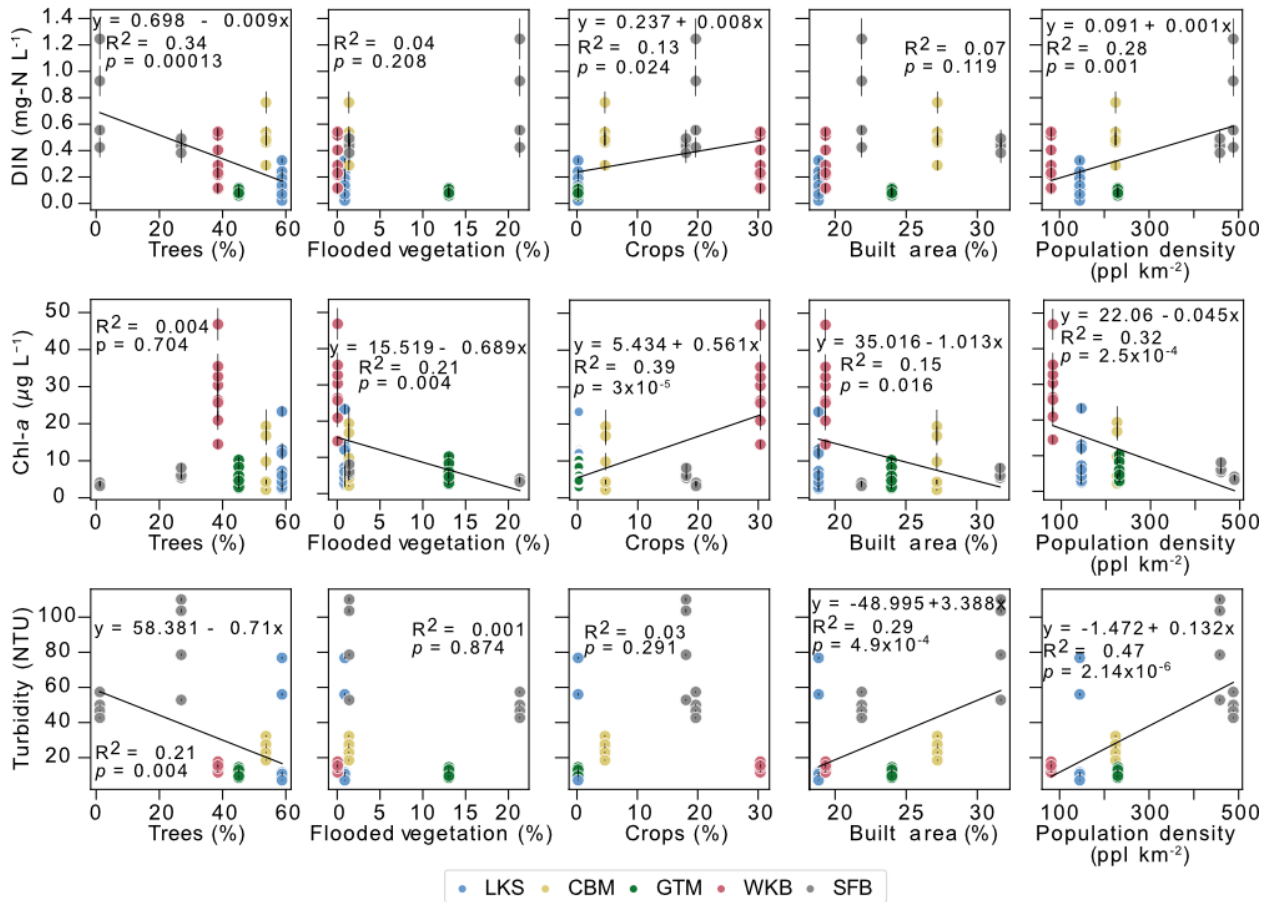
1058 **Fig. S11.** Relationships between the resistance index and total precipitation during each event at
 1059 each estuary. National Estuarine Reserve System (NERR) estuary abbreviations: Lake Superior
 1060 (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas (GTM)
 1061 NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB) NERR. Significant correlations
 1062 ($p < 0.05$) are indicated with a black line.



1063

1064 **Fig. S12.** Resistance versus depth relationship at Lake Superior NERR (LKS) with Oliver Bridge
 1065 and Blatnik Bridge omitted due to positioning of the sondes within the water column (see Table
 1066 S2).

1067



1068

1069 **Fig. S13.** Relationships of dissolved inorganic nitrogen (DIN), chlorophyll-*a* (Chl-*a*), and turbidity
 1070 to land use/land cover and population density across estuaries. Estuary abbreviations: Lake
 1071 Superior (LKS) NERR, Chesapeake Bay, Maryland (CBM) NERR, Guana Tolomato Matanzas
 1072 (GTM) NERR, Weeks Bay (WKB) NERR, and San Francisco Bay (SFB) NERR. Regressions for
 1073 significant relationships (p-value < 0.05) are shown in black lines. All relationships use annual
 1074 means for physicochemical factors calculated for wet and dry years separately and land use/land
 1075 cover characteristics adjoined to monitoring locations at each estuary.