

2 **Weak phases production and heat generation controls**  
3 **fault friction during seismic slip**

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6 **The triggering and magnitude of earthquakes is determined by the friction evolution along**  
7 **faults. Experimental results have revealed a drastic decrease of the friction coefficient for**  
8 **velocities close to the maximum seismic one, independently of the material studied<sup>1,2</sup>. Due to**  
9 **the extreme loading conditions during seismic slip, many competing physical phenomena are**  
10 **occurring (like mineral decomposition <sup>3</sup>, nanoparticle lubrication <sup>1</sup>, melting <sup>4</sup> among others)**  
11 **that are typically thermal in origin<sup>5</sup> and are changing the nature of the material.**

12 **Here we show that a large set of experimental data for different rocks can be described by**  
13 **such thermally-activated mechanisms<sup>6</sup>, combined with the production of weak phases. By**  
14 **taking under account the energy balance of all processes during fault movement, we present**  
15 **a framework that reconciles the data, and is capable of explaining the frictional behavior of**  
16 **faults, across the full range of slip velocities ( $10^{-9} - 10$  m/s).**

17 **The similarity of microstructures observed in nature and in experiments suggests that en-**  
18 **ergetic frameworks like the one presented could quantitatively link observations across the**  
19 **scales and provide deep, physics-based insight on the physical mechanisms driving seismic**

20 **slip.**

21 The knowledge of the friction (shear strength) evolution along a pre-existing fault is of major  
22 importance, as it allows extracting many characteristics and features of seismic slip. In particular,  
23 the decrease of the friction with increasing velocity or displacement (a process called frictional  
24 weakening) determines the possible nucleation of earthquakes. If the weakening rate is larger than  
25 a critical value, this leads to the triggering of a dynamic slip at the origin of earthquakes<sup>7</sup>. In  
26 addition to nucleation, the evolution of the friction coefficient- and thus of the fault's shearing  
27 resistance- determines the arrest of the fault slip and governs the form and budget of energy dissi-  
28 pation during seismic slip<sup>8</sup>. The latter is essential as it determines the amount of energy produced  
29 at the fault, which is radiated on the surface through seismic waves and tremors.

30 During the last 20 years, a large set of experimental works has been devoted to reproducing  
31 the extreme conditions of a seismic slip. The development of high velocity shear apparatus al-  
32 lowed the research community to perform experiments at the maximum velocity reached during an  
33 earthquake event (1 – 10 m/s) and, thus, characterize the behaviour of a fault over the full range of  
34 possible slip rates<sup>9</sup>. A drastic decrease of the friction has been observed in most cases for veloc-  
35 ities closed to the maximum slip velocity independently of the material considered<sup>1</sup>, however the  
36 physical mechanisms accompanying this rapid weakening being different for each rock type. Fol-  
37 lowing microstructural observations and measurements in the sheared samples, several thermally  
38 and mechanically activated weakening mechanisms were proposed to understand the experimental  
39 results at seismic slip rates<sup>6</sup>. The common feature of all these weakening mechanisms is phase

40 transformation -like mineral decomposition<sup>3</sup>, nanoparticle lubrication<sup>1</sup>, melting<sup>4</sup>- during which a  
41 change in the nature of the material takes place.

42 To describe the effect of such a weak phase on the frictional behaviour of a mixture of a  
43 strong/weak phase and constrain the influence of phase change on the mechanical behaviour, we  
44 consider first experiments looking at the effect of a weak phase on the frictional response of fault  
45 zones. The weak phases used for the tests are talc or saturated clay materials sheared at low  
46 velocities (lower than  $10^{-5}$  m/s), so that the mechanisms described above are not triggered. The  
47 results are shown in Figure 1, where we may observe that the friction coefficient  $\mu$  decreases as  
48 the weak phase fraction increases. This effect of the weak phase fraction can be captured using  
49 an exponential law  $\mu = \mu_0 + \Delta\mu e^{-\alpha w}$ , where  $\mu_0$  is the friction coefficient of the weak phase,  
50  $\Delta\mu = \mu_s - \mu_0$  is the difference of the friction coefficient of the strong and weak phases,  $w$  is  
51 the weak phase fraction and  $\alpha$  a weakening coefficient ranging from 0.1 to 15 (see supplementary  
52 information). Note that such nonlinear weakening laws are also used in geomechanical engineering  
53 to describe the weathering of calcarenite<sup>10,11</sup>.

54 The derived exponential decrease of the mechanical strength from the experimental data of  
55 Figure 1 is then included into a thermo-chemo mechanical model that accounts for the coupled  
56 mechanisms activated at higher velocity conditions (see Methods for the mathematical description  
57 of the model). In this model, the degradation or creation of a weak phase is induced by the energy  
58 input to the system and it is not present before shearing. This unifying approach aims at reconcil-  
59 ing observations across a wide spectrum of materials and velocities. The extensive experimental

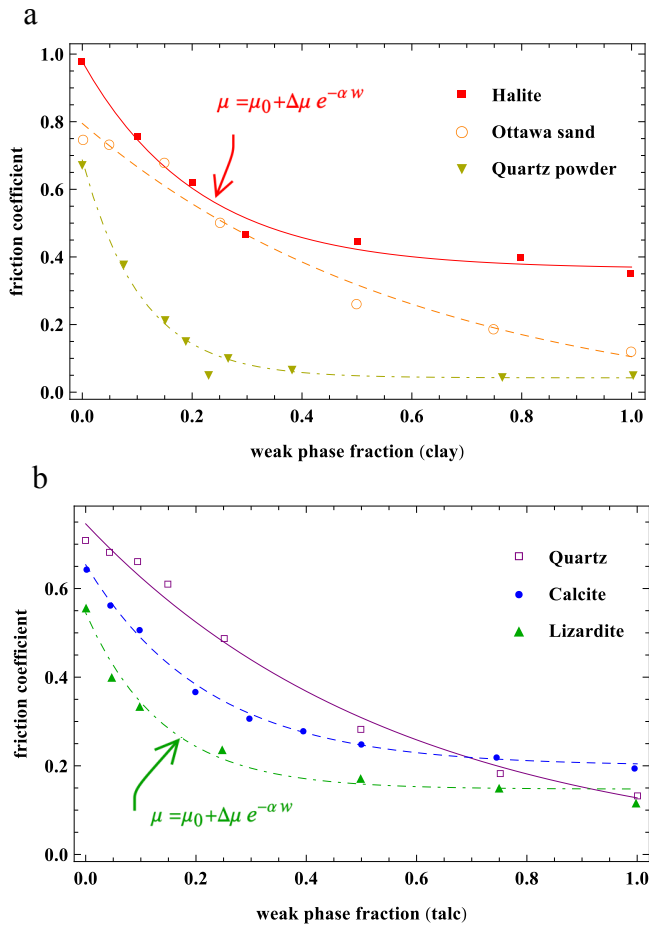


Figure 1: **Effect of the weak phase content on the steady state friction coefficient.** The friction of weak/strong phase binary mixtures is represented as function of the weak phase content in experiments carried out at subseismic sliding velocities and at constant normal stress using triaxial saw cut, double and simple direct shear and rotary shear configurations. **a**, results for clay as weak phase: red corresponds to a muscovite/halite mixture<sup>12</sup>, orange to a crushed Ottawa sand/montmorillonite mixture<sup>13</sup>, dark yellow to a quartz powder/bentonite mixture<sup>14</sup>. **b**, results for talc as weak phase: purple corresponds to quartz as strong phase<sup>15</sup>, blue to calcite<sup>16</sup>, green to Lizardite<sup>15</sup>. In **a-b**, the points represent the experimental data and the solid or dashed lines represent the interpolation using this exponential function.

60 data set used for the comparison corresponds to shear tests performed with rotary shear apparatus  
61 that allows to reach high displacements and therefore the steady state (see Figure 2). These exper-  
62 iments are realized on either gouge granular samples (usually 1mm thick<sup>17</sup>) sandwiched between  
63 two blocks or on bare rock samples<sup>18</sup>. In the latter, a gouge material is formed after only a few  
64 millimetres of displacement<sup>2</sup> with a thickness of 100 to 300  $\mu\text{m}$ . The data are gathered based on  
65 the nature of the material sheared and the physical mechanisms that are inferred to operate during  
66 the experiments<sup>1,19</sup>.

67 The constitutive law for the mechanical behaviour is applied to the gouge material, which ac-  
68 commodates all the deformation and is affected by the temperature, the weak phase fraction and the  
69 state of stress. The weak phase creation is modelled as an endothermic first order chemical trans-  
70 formation affecting the energy balance equation and respecting the mass balance. The geometry  
71 of the model chosen is larger of one or two orders of magnitude than the gouge in order to impose  
72 far field boundary conditions for the temperature and the extend of the phase transformation (see  
73 Figure 2). The steady states of this model can be determined using a continuation algorithm (see  
74 Methods), to test the hypothesis that the combination of thermally activated weakening and the  
75 creation of a weak phase may account for the observed steady state frictional response over many  
76 orders of magnitude of shear velocity.

77 The resulting steady state response of the model, in terms of friction and velocity, is depicted  
78 in Fig. 2. We can identify five distinct regimes of the system response to loading velocity: (I)  
79 **Static**. At low velocities, the material remains at static friction. Negligible temperature increase

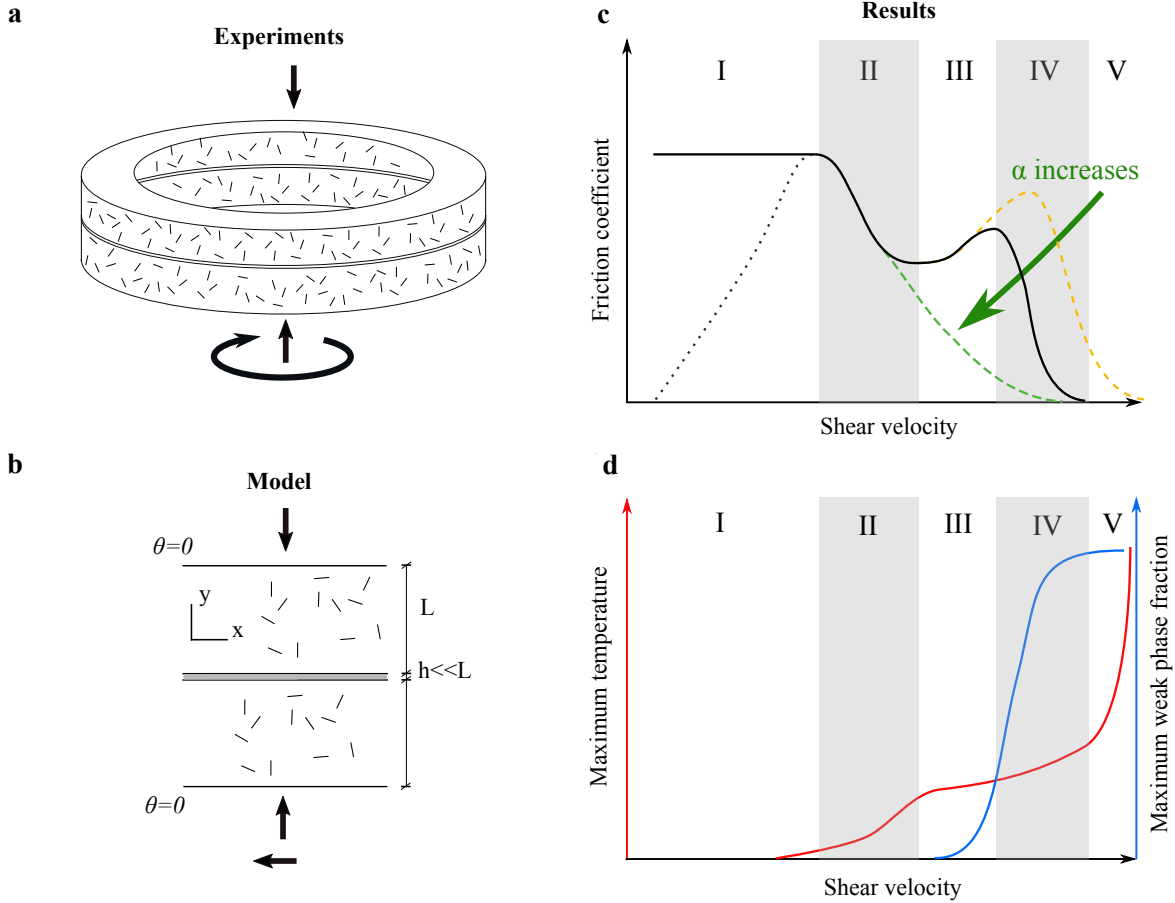
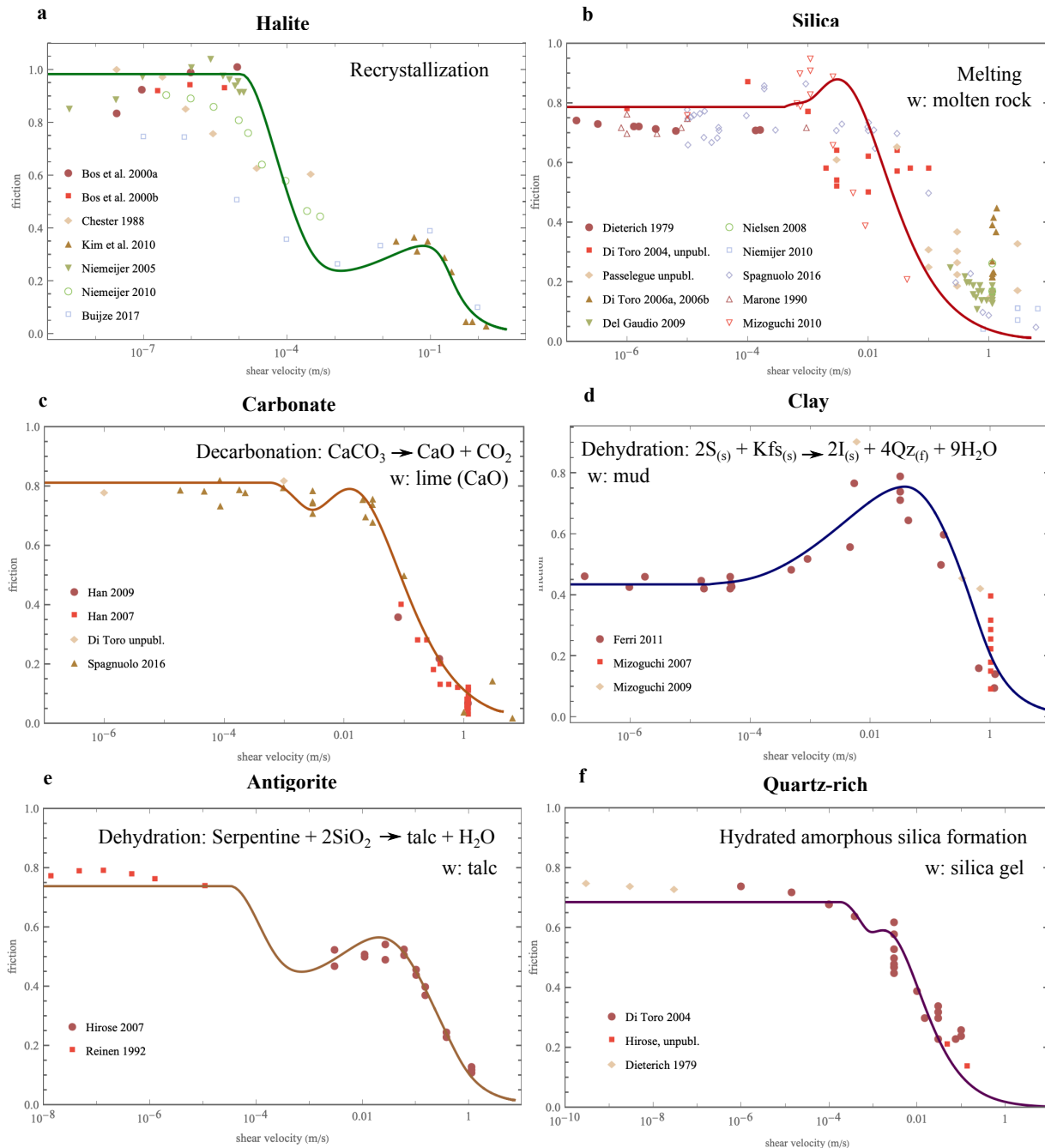


Figure 2: **a-b, Schematic view of the geometry of the high velocity rotary shear experiments and the model. c-d, Steady state of the thermo-chemo mechanical model.** Response of the mathematical system of equations (Eqs. 5-6 in Methods) at steady state. **c**, the steady state friction coefficient as a function of velocity, for varying weak phase sensitivity coefficient  $\alpha$ . **d**, the steady-state temperature and weak phase content dependency on the velocity, plotted for the black line of **c**. The five regimes (I - V) that can be observed in the response of the friction coefficient are correlated with the temperature and weak phase production processes, as explained in the main text.

80 or weak phase production is observed. (II) **Thermo-mechanical weakening.** The temperature  
81 increase leads the friction coefficient to drop, in the absence of any weak phase production. (III)  
82 **Thermo-chemical stabilization.** With increasing velocity and temperature, small fractions of the  
83 weak phase are produced, absorbing the excess temperature and reducing the thermal softening  
84 effect on the friction coefficient. Depending on the value of the weakening coefficient, the friction  
85 coefficient can experience an intermediate increase. (IV) **Chemo-mechanical weakening.** Even-  
86 tually, weak phase content will be produced until it reaches a critical value that will dominate the  
87 friction coefficient and lead the material to unconstrained weakening. (V) **Runaway instability.**  
88 Once the reactants are depleted ( $w = 1$ ), the temperature is increasing uncontrollably and the  
89 friction coefficient drops towards zero.

90 After identifying the regimes of the steady frictional response of faults, the model is applied  
91 to experimental data. Figure 3 summarizes the results of the model for the steady state friction  
92 coefficient as a function of the velocity for six sets of materials<sup>1,9,18-21</sup>. The experimental data  
93 are a collection of several independent studies at different experimental conditions. As shown in  
94 the supplementary information, the normal stress shows no clear effect on the value of the friction  
95 coefficient reported here and for this reason it is not further investigated in this study. One of  
96 the interesting features of the model is the reproduction of strengthening observed experimentally  
97 at intermediate velocities (regime III in Figure 2) without supposing any additional hardening  
98 mechanism. For each material, the inferred phase transformation and the resulting weak product  
99 are highlighted. The model enables to capture accurately the observations and uses as input the  
100 material parameters listed in the supplementary information. Moreover, it enables us to retrieve



**Figure 3: Results of the model for different types of materials.** Application of the mathematical model (solid lines) to literature data (dots) of the friction coefficient as a function of the velocity. See the Methods for the mathematical description, Fig. 2 for the qualitative steady-state response and the supplementary information for the references of the experimental data. For each material, the phase transformation and the associated weak phase are indicated. **a**, for halite rock. The model reproduces the experimental behaviour with  $\alpha = 0$ . **b**, for silicate rocks ( $\alpha = 7.5$ ). **c**, for carbonate rocks ( $\alpha = 5.3$ ). **d**, for clay-rich rocks ( $\alpha = 2$ ).



101 information for the parameters of the different processes such as the activation energies, together  
102 with an assessment of the temperature evolution. This can be used as a basis to compare with  
103 the microstructural observations of the samples after the experiments. Indeed, in Figure 4 we  
104 are summarizing the evolution of the Temperature and weak phase ratio  $w$  required to obtain the  
105 friction coefficient results of Figure 3. Based on these two figures, we can compile the processes  
106 underpinning the macroscopic response of the frictional resistance of the different materials.

107 The temperature predicted by the model when the weak phase begins to appear can be  
108 compared to the theoretical activation temperature of the phase transformation (decarbonation<sup>22</sup>,  
109 melting<sup>23</sup>, dehydration<sup>24</sup> or clay type transition<sup>9</sup>), when available. In all cases, the temperature of  
110 the model is lower than the theoretical one (e.g. 530° C against 720° C for carbonates) imply-  
111 ing that the phase transition is triggered locally at the contact of the grains where the temperature  
112 can be higher than the bulk temperature. Moreover, these local phase changes are hard to detect  
113 even though essential for the mechanical behaviour. This also explains why the evidences of the  
114 phase transformations from specific sensors or microstructural observations (recrystallized halite  
115 grains<sup>17</sup>, increase of CO<sub>2</sub><sup>22</sup> or humidity next to the tested sample<sup>18</sup>, melted asperities<sup>25</sup> and white  
116 flakes due to silica gel<sup>20</sup>) are retrieved for higher velocities in experiments than predicted by the  
117 model. A notable case is halite, for which the weakening factor is  $\alpha = 0$  implying that any weak  
118 phase generated during shear does not affect the friction coefficient. As the material undergoes  
119 recrystallisation<sup>17</sup> during shearing which is a phase transformation that produces the same mineral  
120 with different grain sizes. Despite not producing a weak phase directly though, recrystallisation af-  
121 fects the energy budget and, thus the temperature produced (Figure 4) and therefore the mechanical

122 behaviour of the gouge.

123       These results suggest that a thermally activated creep and the transformation of the material  
124 inside the fault zone may be the dominant mechanisms during seismic slip of dry rocks. In the case  
125 where pressurized fluids are present -a scenario difficult to study experimentally- the static friction  
126 response<sup>9</sup> and the transient behaviour through thermal pressurization<sup>26,27</sup> are likely to be affected.  
127 In that case a detailed comparison between phase transition and thermal pressurization needs to be  
128 carried, to assess the dominant mechanism of slip <sup>28</sup>.

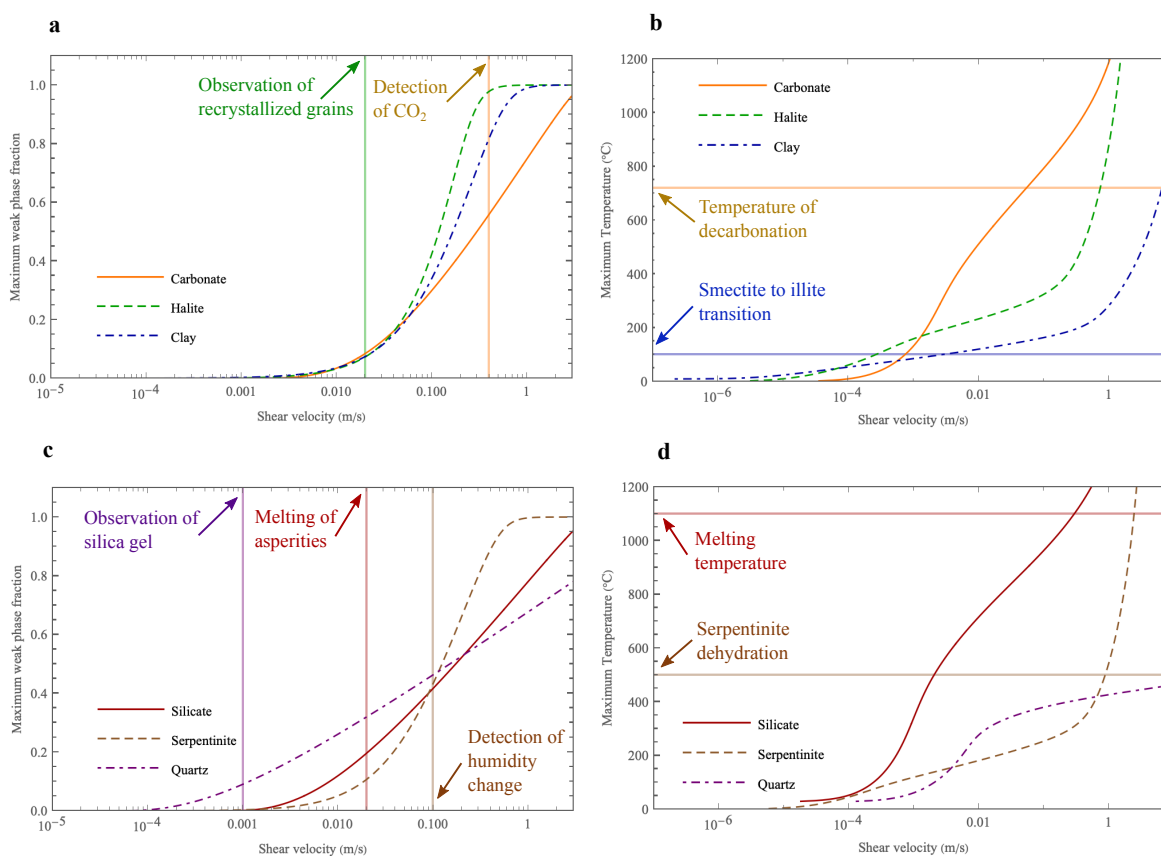


Figure 4: **Temperature and weak phase fraction evolution with the velocity for the experiments of Fig.**

**3. a and c,** weak phase fraction in the middle of the sample. Note that weak phase production has been approached as a first order chemical reaction, thus reaching its maximum value of one (100% weak phase present) when the reaction is depleted. **b and d,** maximum temperature in the middle of the sample.

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**Methods**

**Description of the Mathematical Model.** When the shear velocity applied to a rock or a granular sample is increased, thermal effects tend to dominate the frictional response<sup>5,20,29,30</sup>. A critical velocity is required to activate this mechanism that is related to the processes at asperities or grain contacts<sup>5</sup>. In this paper, we focus on the response of the material for intermediate and high velocities where the thermal and chemical effect are important (other models have been developed for lower velocities and the nucleation of earthquakes<sup>29,31</sup>). Therefore, we consider here that for low velocities experiments the strength of the material is determined by the static friction of the materials in contact or the internal static friction of the granular assembly. For velocities larger than the critical one, the shear stress of the system is calculated by solving a thermo-chemo-mechanical model inside the deforming zone. The critical velocity is retrieved as a result from this model and can be approximated by an analytical solution (see supplementary information). Physically, it corresponds to the critical velocity for which thermal weakening becomes significant.

The mathematical model consists of solving the momentum, mass and energy balance equation at steady state, for an infinite sheared layer. The equations are briefly summarized here for easiness in reproducibility of the results<sup>3</sup>.

The momentum balance equations are considered and we neglect the inertia terms<sup>32</sup>:

$$\sigma_{ij,j} = 0 \quad (1)$$

where  $\sigma_{ij}$  is the stress tensor. These equations lead in the case of one dimensional shear zone to a constant normal and shear stress in space inside the layer.

148 The constitutive law for the mechanical behaviour is a rigid elastic-viscoplastic law with  
 149 the most generic form: an Arrhenius-power law dependency<sup>3,33</sup>. This law is only considered in  
 150 a layer of thickness  $h$ , much smaller than the total thickness of the layer  $L$  (see Figure 2). This  
 151 enables to describe the fact that after only a few millimetres of slip during the shear experiments  
 152 on bare rocks, a thin layer of gouge materials forms. This layer composed of crushed grains from  
 153 the initially rough surfaces has generally a thickness of 100-300  $\mu\text{m}^2$  and accommodates all the  
 154 deformation. However, as the boundary conditions for the weak phase and the temperature are  
 155 not well defined for this gouge layer, a domain of 1 cm is considered in order to apply Dirichlet  
 156 boundary conditions for these fields.

$$\dot{\epsilon}^{vp} = \dot{\epsilon}^0 \left( \frac{\tau}{\tau_y} \right)^m e^{-Q/RT} \quad (2)$$

157 where  $\dot{\epsilon}^0$  is a reference strain rate,  $m$  is the exponent of the power law,  $\tau$  is the shear stress,  $\tau_y$  is  
 158 the yield stress,  $Q$  is an activation enthalpy for the microscopic mechanism inducing a nonlinear  
 159 behaviour,  $R$  is the perfect gas constant and  $T$  is the temperature. This law allows to include more  
 160 physics into the hardening evolution as in the theory of plasticity for metals<sup>33</sup>. The Arrhenius  
 161 dependency of the flow law enables to introduce multi-physical couplings such as the effect of  
 162 heat generation on the frictional strength or more generally interface phenomena between the solid  
 163 skeleton and the pores<sup>34,35</sup>.

164 The effect of the non-mechanical state variables on the mechanical behaviour of the system  
 165 can be expressed as a single scalar function called the weathering index<sup>11</sup>,  $X_d$ . The strength of the  
 166 material depends on both the plastic strain and this weathering index. It is assumed that the two

167 effects are uncoupled<sup>10,11</sup> and a multiplicative structure of the yield stress is postulated:

$$\tau_y = T_y(\epsilon^p)T_y(X_d) \quad (3)$$

168 In our case, we do not consider any purely mechanical hardening law, so that the function  $T_y(\epsilon^p)$  is  
 169 constant. Moreover,  $X_d$  is considered to be a weak phase volume fraction. As shown in Figure 1,  
 170 the presence of a weak phase induces an exponential decrease of the frictional strength along with  
 171 the weak phase fraction. Assuming negligible shear strength for the weak phase, the final form of  
 172 the constitutive law is therefore:

$$\dot{\epsilon}^{vp} = \dot{\epsilon}^0 \left( \frac{\tau}{\tau_0} \right)^m e^{-Q/RT} e^{\alpha m w} \quad (4)$$

173 where  $\tau_0$  is the yield strength of the strong phase and  $\alpha$  is the weak phase sensitivity coefficient.

174 For a material consisting of two species: a weak and a strong phase, occupying volumes  $V_w$   
 175 and  $V_s$  respectively, we may define the volume ratio  $w = \frac{V_w}{V_w + V_s}$ . Inside a one dimensional shear  
 176 zone yield a system of two equations<sup>3,36</sup> obtained from the mass balance of the weak phase fraction  
 177 and the energy balance equations:

$$\frac{\partial T}{\partial t} = c_{th} \frac{\partial^2 T}{\partial y^2} + F(y) \frac{\tau \dot{\epsilon}^{vp}}{\rho C} - \frac{\Delta H r_F}{\rho C} \quad (5)$$

$$\frac{\partial \rho_1}{\partial t} + \frac{\partial J_w}{\partial y} = r_F \quad (6)$$

178 where  $T$  is the temperature,  $c_{th}$  the thermal diffusivity,  $\rho C$  the heat capacity of the mixture con-  
 179 sidered constant here,  $J_w$  the diffusion flux of the weak phase,  $\Delta H$  the enthalpy of the phase  
 180 change reaction considered endothermic,  $r_F$  the reaction rate and  $\rho_1 = \rho_w w$ .  $F(y)$  is a function  
 181 which value is 1 for  $y \in [-h/2, h/2]$  and 0 otherwise. The reaction rate is expressed as first order

182 chemical reaction with an Arrhenius law.

$$r_F = (1 - w) \frac{\rho_s}{M_s} k_F e^{-Q_c/RT} \quad (7)$$

183 where  $\rho_s$  and  $M_s$  are the density and molar mass of the strong phase.  $k_F$  and  $Q_c$  are the preexpo-  
 184 nential factor and activation energy of the chemical reaction. Using Equations 4, 7, considering a  
 185 Fick's law for the diffusion flux (defining a diffusivity  $c_w$ ) and the steady state of Equations 5 and  
 186 6, we obtain a system of two differential equations in space. This system is written in a dimension-  
 187 less form for the purpose of reducing the number of parameters to study and to enable a clearer  
 188 understanding of the main features of the system:

$$\frac{\partial^2 \theta}{\partial \bar{y}^2} + F(\bar{y}) Gr e^{\frac{Ar \theta}{1+\theta}} e^{\alpha m w} - Da(1 - w) e^{\frac{Arc \theta}{1+\theta}} = 0 \quad (8)$$

$$\frac{\partial w}{\partial \bar{y}^2} + \mu Da(1 - w) e^{\frac{Arc \theta}{1+\theta}} = 0 \quad (9)$$

189 where,  $\theta$  is the dimensionless temperature.  $Gr$ ,  $Da$ ,  $Ar$  and  $Arc$  are called the Gruntfest, Damköhler,  
 190 Arrhenius and chemical Arrhenius numbers respectively. They are defined by:

$$Ar = \frac{Q}{R T_0}, \quad Arc = \frac{Q_c}{R T_0} \quad (10)$$

$$Gr = \frac{\tau_0 \dot{\epsilon}_0 L^2}{\rho C c_{th} T_0} \left(\frac{\tau}{\tau_0}\right)^{m+1} e^{-Ar} \quad (11)$$

$$Da = \frac{\Delta H k_F \rho_s L^2}{\rho C M_s c_{th} T_0} e^{-Arc} \quad (12)$$

$$\mu = \frac{\rho C c_{th} T_0 M_w}{\Delta H \rho_w c_w} \quad (13)$$

191 **Numerical Bifurcation of the Steady State friction coefficient.** The solutions of this nonlinear  
 192 system of differential equations are approximated numerically using pseudospectral methods. The  
 193 temperature and weak phase fraction fields are interpolated in space using Chebyshev polynomials

194 of the first kind:

$$\theta(\bar{y}) = \sum_{i=1}^N a_i (\phi_{2i}(\bar{y}) - 1) \quad (14)$$

$$w(\bar{y}) = \sum_{i=1}^N b_i (\phi_{2i}(\bar{y}) - 1) \quad (15)$$

195 where  $\phi_{2i}$  are the Chebyshev polynomials of degree  $2i$ . Note that only the even degree Chebyshev  
 196 polynomials are kept here as the solution is symmetric about the origin. Moreover, a basis recom-  
 197 bination is used by considering interpolation functions of the form  $\psi_{2i}(y) = \phi_{2i}(y) - 1$ , allowing  
 198 to enforce a zero Dirichlet boundary conditions implicitly<sup>38</sup>.  $N$  is the number of polynomials used  
 199 to simulate the solutions. A convergence analysis has been conducted in each case to verify that  $N$   
 200 is high enough to have a negligible error on the solution.  $a_i$  and  $b_i$  are the interpolation coefficients  
 201 for the temperature and the weak phase fraction respectively. The interpolation points used for the  
 202 resolution are the Gauss-Lobato points defined by:

$$x_j = \cos\left(\frac{(2j-1)\pi}{4N}\right), \quad j = 1, \dots, N \quad (16)$$

203 The nonlinear system of algebraic equations obtained is solved using the Newton-Raphson method.  
 204 In order to capture all the steady state solutions of the system for the different values of the stress, a  
 205 continuation pseudo-arclength algorithm is used. The continuation parameter chosen is the Grunt-  
 206 fest number<sup>3</sup>.

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