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# PARADOXICAL IMPACT OF SPRAWLING INTRA-URBAN HEAT ISLETS: REDUCING MEAN SURFACE TEMPERATURES WHILE ENHANCING LOCAL EXTREMES

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A PREPRINT

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## ABSTRACT

1 Extreme heat is one of the deadliest health hazards that is projected to increase in intensity and  
2 persistence in the near future. Here, we tackle the problem of spatially heterogeneous extreme heat  
3 distribution within urban areas. We develop a novel multi-scale metric of identifying emerging heat  
4 clusters at various percentile-based thermal thresholds and refer to them collectively as *intra-Urban*  
5 *Heat Islets*. Using remotely sensed Land Surface Temperatures, we first quantify the spatial organi-  
6 zation of heat islets in cities at various degrees of sprawl and densification. We then condense the  
7 size, spacing, and intensity information about heterogeneous clusters into probability distributions  
8 that can be described using single scaling exponents (denoted by  $\beta$ ,  $\Lambda_{score}$ , and  $\lambda$ , respectively).  
9 This allows for a seamless comparison of the heat islet characteristics across cities at varying spatial  
10 scales and improves on the traditional Surface Urban Heat Island (SUHI) Intensity as a bulk metric.  
11 Analysis of Heat Islet Size distributions demonstrates the emergence of two classes where the dense  
12 cities follow a Pareto distribution, and the sprawling cities show an exponential tempering of Pareto  
13 tail. This indicates a significantly reduced probability of encountering large heat islets for sprawl-  
14 ing cities. In contrast, analysis of Heat Islet Intensity distributions indicate that while a sprawling  
15 configuration is favorable for reducing the mean SUHI Intensity of a city, for the same mean, it  
16 also results in higher local thermal extremes. This poses a paradox for urban designers in adopting  
17 expansion or densification as a growth trajectory to mitigate the UHI.

## 18 Introduction

19 More than 50% of the world’s population currently resides in cities, and the number continues to increase rapidly with a  
20 projection that 70% of the global population will be urbanized by 2050 (Prudhomme, 2018). Rapid urbanization trends  
21 are manifested in expansion and densification of existing cities and the merging of agglomerations to form megacities,  
22 particularly in South Asia and sub-Saharan Africa (Seto and Shepherd, 2009). Among the numerous challenges that  
23 cities face, a particularly urgent problem due to climate change is that of extreme heat. Urban areas often raise the  
24 local temperatures relative to natural and rural surroundings leading to the phenomenon of Urban Heat Island (UHI)  
25 effect. A synergistic interaction between UHIs and increasingly persistent heat waves further exacerbates the extreme  
26 temperatures within cities (Li and Bou-Zeid, 2013; Zhao et al., 2018). Repercussions of extreme heat include thermal  
27 discomfort (Nikolopoulou and Lykoudis, 2006), increased energy consumption (Santamouris, 2014), and a greater  
28 number of heat-related casualties during heat waves (Semenza et al., 1996; Uejio et al., 2011).

29 The UHI is typically quantified as UHI Intensity, i.e. the difference between the air temperatures of a represen-  
30 tative urban area (point measurement or spatial average) and rural area. However, such an estimate is inadequate to  
31 address the *intra-urban* spatial heterogeneity. Commendable efforts to collect spatially resolved thermal data such as  
32 the Basel Urban Boundary Layer Experiment (BUBBLE) campaign (Rotach et al., 2005) are rare and often limited to  
33 a single city. On the other hand, earth-monitoring satellites such as Landsat and Moderate Resolution Imaging Spec-  
34 troradiometer (MODIS) enable consistent high spatial resolution characterization across multiple cities. As a result,  
35 the Surface UHI (SUHI) estimated using Land Surface Temperatures (LST) has emerged as an alternative approach  
36 which we adopt in this study (Roth et al., 1989; Voogt and Oke, 2003). For satellite sensors, urban features such as  
37 building roof and wall exteriors, surface materials, albedo, impervious and vegetated fractions, and surface moisture  
38 within each pixel determine the resultant LST. Note that while SUHI bears similarity in spatial and temporal patterns  
39 to UHI, spatial patterns of SUHI are more coupled with urban form and function, whereas air temperatures are subject  
40 to the boundary layer wind profiles as well, therefore, a point-to-point correspondence can not be expected (Panwar  
41 et al., 2019).

42 LST, which is strongly determined by the land use land cover properties, emerges from the underlying self-  
43 organization of the urban forms. As a result, the role of spatial organization of urban form in reducing SUHI Intensity  
44 has been a topic of substantial research spanning multitude of spatial scales. At the micro-scale, i.e. within the urban  
45 canyon, the surface temperatures are extremely sensitive to the geometrical details of immediate surroundings, such  
46 as street canyon geometry, sky-view factor, vegetative fraction, solar access and shading (Jamei et al., 2016; Taleghani  
47 et al., 2015; Andreou, 2013). However, at the local scale, i.e. of the order of few kilometers, consistent thermal pat-  
48 terns emerge due to locally homogeneous patches of urban form and function (Stewart and Oke, 2012; Ching et al.,  
49 2018). We do not have a clear understanding of the optimal urban form and function that minimize the urban heat  
50 locally as well as at a city-scale. For instance, studies investigating local scale impacts (Sobstyl et al., 2018; Schwarz  
51 and Manceur, 2014) report that high-density urban development leads to higher local temperatures, whereas, several  
52 others note that sprawling urban development may result in worse thermal conditions (Stone Jr and Rodgers, 2001;  
53 Stone et al., 2010). Debbage and Shepherd (2015) show that regardless of the urban density type *within a patch*, the  
54 relative spatial contiguity of the urban land use patches is a critical variable as well. Despite these recent advances on

55 the intra-urban thermal landscape over the last few decades, a comprehensive framework for the characterization of  
56 intra-urban thermal heterogeneity for diverse city morphologies is still lacking.

57 We use a multi-scale framework wherein we treat the SUHI not as a single entity, but as a collection of heterogeneous  
58 clusters of heat within the city. We refer to these clusters as **intra-urban heat islets**. The objective of this study is  
59 to evaluate the impact of spatial organization of these heat islets on their properties such as size and intensity and  
60 determine if there is a favourable spatial structure for reducing surface temperature extremes at intra-urban spatial  
61 scales. Urban morphology, and as a result LST, emerges via the processes of densification and expansion, albeit  
62 constrained by cultural, geographic and economic factors (Batty, 2013; Mustafa et al., 2018). Different degrees and  
63 combinations of these two processes result in diversity of form and function. Dense urban growth occurs when there  
64 is increased in-fill construction within existing high-density built-up area. Such a process is often driven by economic  
65 and socio-political factors that lead to settlement of new urban regions close to the city center (Andersson et al., 2006).  
66 This is akin to the preferential attachment phenomenon observed in complex networks where a new node is more likely  
67 to agglomerate at the "hub nodes" with the highest density of edges (Barabási and Albert, 1999). We hypothesize that  
68 the densification within urban area results in hot regions getting hotter and larger thereby resulting in a heavy-tailed  
69 size distribution of heat islets. Urban expansion in the form of sprawl, on the other hand, occurs at the periphery of  
70 urban areas in the form of growing sub-urban regions. We hypothesize that this would lead to the emergence of heat  
71 islets that are spread more evenly throughout the city, often interspersed with local heat sinks. This can be detected  
72 in the size distribution as a fast decaying tail, often in the form of an exponential tempering (Amaral et al., 2000).  
73 Similar effects of urban expansion and densification on the power-law node degree observed is observed in several  
74 urban infrastructure systems such as roads and sewage networks (Mohajeri et al., 2015; Yang et al., 2017; Klinkhamer  
75 et al., 2017). Note that we don't refer to the spatial organization of urban assets such as buildings or impervious  
76 areas. Rather, we directly analyze the LST. We implement the framework for a set of 78 cities sampled globally.  
77 Using probabilistic models, we condense the size, spacing, and intensity information about heterogeneous clusters  
78 into distributions that can be described using single scaling exponents. This allows for a seamless comparison of the  
79 heat islet characteristics across cities that represent varying degrees of sprawl or densification. We then assess how the  
80 thermal spatial structure relates to the traditional lumped metric, SUHI Intensity. Lastly, we discuss implications for  
81 desirable thermal configurations for cities to minimize the area and intensity of the heat islets.

## 82 **Data acquisition and Clustering technique**

83 A set of 78 cities were sampled encompassing diverse climatological, geographical, and cultural backgrounds as well  
84 as different realizations of urban form and function (Figure 1). The cities selected range from megalopolises such as  
85 Guangzhou, London, and New York City with a population of over 10 Million and metropolitan areas up to 3000 km<sup>2</sup>,  
86 to smaller cities such as Tbilisi, Bern, and Oslo that span less than 100 km<sup>2</sup>. As a globally standardized dataset of  
87 urban extent, the Urban Land Use layer of Land Cover product from MODIS was used. The exact definition of urban  
88 boundaries and city area plays a significant role in urban scaling laws (Cottineau et al., 2017), therefore, a buffer of  
89 5 km in the rural regions was taken to account for the peri-urban settlements. However, as the heat islets occur well  
90 within the city boundaries, the results were found to be independent of the buffer width.

91 For each city, we selected a Landsat image of a summertime cloud-free day, and derived the LST in the geospatial  
 92 computing environment of Google Earth Engine (Gorelick et al., 2017) using the methodology described in Walawen-  
 93 der et al. (2012) (Walawender et al., 2012). A novel aspect of our methodology is the clustering technique used to  
 94 characterize the LST. The LSTs are treated analogously to topography in Digital Elevation Models (DEM), where the  
 95 temperatures substitute for elevation (Shreevastava et al., 2018). As the cities belong to diverse climatic backgrounds  
 96 (and hence, background temperatures) (Zhao et al., 2014), percentile-based thermal thresholds were chosen for identi-  
 97 fying the relative hottest regions within the urban areas. The areas above a given thermal threshold (e.g., corresponding  
 98 to percentiles of temperatures) were identified, and the connected pixels were grouped into a cluster that we refer to as  
 99 a **heat islet**. Supplementary Information provides code and text describing the methods in more detail.

## 100 Size distribution of heat islets

101 In an exploration of *shapes* of heat islets, we found consistent self-similar, fractal topography across all cities (Shreev-  
 102 astava et al., 2019) (See Supplementary Figure 1). Here, we focus on their *size* distribution by building on the scaling  
 103 laws known for fractal surfaces. According to Korcak’s law, the size distribution of clusters in a fractal topography is  
 104 expected to follow a power law at the percolation threshold (Imre and Novotný, 2016; Mandelbrot, 1975; Isichenko  
 105 and Kalda, 1991). This is mathematically represented as:  $N(a) \propto a^{-\beta}$  where  $N$  is the number of clusters of area,  $a$ ,  
 106 and the scaling exponent is  $\beta$ . Expressed as an exceedance probability we can write it as a Pareto distribution:

$$P(A \geq a) \propto a^{1-\beta}, \quad \forall a \geq a_{min} \quad (1)$$

107 where, for a given area  $a$ , the probability of an islet having an area  $A$  larger than  $a$  is represented by  $P$ , the scaling  
 108 exponent is represented by  $\beta$ , and the minimum area at or above which the power-law is valid is represented as  $a_{min}$ .  
 109 We use Maximum Likelihood Estimation (MLE) to test for and fit the exceedance probability distributions (Clau-  
 110 set et al., 2009) (See Supplementary Text 3). This process is carried out for multiple thermal thresholds (50<sup>th</sup>, 60<sup>th</sup>,  $\dots$ ,  
 111 90<sup>th</sup> percentiles). We find that the estimated exponents ranged between 1.6 to 2.2 with a mean  $\beta = 1.88$ . However,  
 112 for the smaller cities ( $A_{city} \leq 650 \text{ km}^2$ ), the variability in exponents was much larger (Supplementary Figure 2).  
 113 One explanation for this is statistical, wherein for small cities there are not enough islets obtained at 90 m resolution  
 114 which results higher statistical fluctuations about the mean are observed. As the number of islets increases with city  
 115 size, steady averaging is achieved that results in convergence towards the mean. However, from an urban growth  
 116 perspective, this behavior is consistent with several other complex systems that operate within cities (Klinkhamer  
 117 et al., 2017; Barthelemy, 2016). For smaller cities, the variability due to factors unrelated to city size result in more  
 118 detectable fluctuations and it simply indicates that they have not grown in size enough to display self-organization yet  
 119 (Batty, 2008). We, therefore, excluded those from any further analysis and proceed with 49 cities where the internal  
 120 thermal structure could be reliably quantified. For the larger cities, the distributions were well described by Equation  
 121 1 with the same mean exponent and a narrow variability (std. dev. = 0.026).

122 The impact of a dense or sprawling spatial organization becomes apparent in how the exceedance probability dis-  
 123 tributions change as the threshold increases. The large metropolitan regions of Lagos and Jakarta are selected as

124 representatives of dense cities, where Chicago and Guangzhou are chosen to represent sprawling cities (Figure 2a,b).  
 125 The Pareto size distribution is consistent at lower thresholds for all cities. At 90<sup>th</sup> percentile threshold, however, Lagos  
 126 and Jakarta show a pronounced aggregation of heat islets indicative of dominance of a dense urban center, whereas  
 127 Chicago and Guangzhou are more dispersed (Figure 2c,d). In agreement with our initial hypothesis, Lagos and Jakarta,  
 128 display a Pareto heat islet size distribution across all the thresholds (Figure 2e). However, for Chicago and Guangzhou,  
 129 the heat islet size distributions deviate significantly from the Pareto in the form of an exponential tempering (Figure  
 130 2f), such that their distributions more closely follow:

$$P(A \geq a) \propto a^{1-\beta} \cdot e^{-c \cdot a}, \quad \forall a \geq a_{min} \quad (2)$$

131 where  $c$  represents the exponential tempering coefficient (Supplementary Table 3).

132 Such behavior is explained by invoking percolation theory (Isichenko, 1992; Sahimi and Sahimi, 2014). Percolation  
 133 theory is the study of random clusters and their spatial connectivity at a given threshold. The coagulation of dispersed  
 134 clusters into a contiguous component is referred to as percolation, and the largest cluster is identified as the percolating  
 135 cluster. In fractal landscapes, the Pareto size distribution of clusters holds within a finite range (Percolation Transition  
 136 Range) of thresholds, i.e., until the percolating cluster retains its identity. We computed the percolation transition range  
 137 by identifying the inflection points in the size of the largest cluster as a function of temperature threshold (Figure 2g,h).  
 138 The range was then normalized using the minimum and maximum temperatures for each city such that the range is  
 139 restricted to 0 and 1. We refer to this as the Normalized Percolation Range (NPR) (Supplementary Figure 3). In case  
 140 of the aggregated cities (e.g. Jakarta and Lagos) as the temperature threshold is increased, the largest connected islet  
 141 decreases in size gradually, and the resulting NPR is large (Figure 2g). Conversely, in the case of sprawling cities  
 142 (e.g., Chicago and Guangzhou) there is a much sharper decrease in the size of the percolating cluster (Figure 2h)  
 143 resulting in a narrow NPR (Figure 2i,j). As the 90<sup>th</sup> percentile thresholds in these cases fall outside the NPR (Figure  
 144 2j), exponential tempering is observed.

145 From the perspective of the size distribution of heat islets alone, as the thermal threshold is increased, fewer and  
 146 smaller heat islets are captured. Therefore, an exponential tempering presents a reduced probability of encountering  
 147 large heat islets of higher temperatures. This suggests that a sprawling spatial structure is favourable for reducing  
 148 the size of extreme heat islets. Thus far, we have characterized the size distribution of these islets, not their spatial  
 149 organization. We now introduce a metric to quantify and analyze the relationship between the spacing of the urban  
 150 heat islets and the characteristics we observed in their size distributions.

## 151 **Quantifying Aggregated vs Dispersed heat islets: Lacunarity**

152 A built-up patch in a city acts as a source of increased sensible heat flux, as well as anthropogenic heat flux due to  
 153 human activities such as air-conditioning. Likewise, the gaps between the patches (also referred to as spacing in this  
 154 work), often water or vegetation, act as heat sinks that absorb the excess heat generated. Therefore, characterizing this  
 155 spacing between the urban patches is an essential step towards ameliorating heat stress (Debbage and Shepherd, 2015).  
 156 Particularly, the impact of the relative sizes and strengths of such sources and sinks on the overall thermal landscape

157 has been relatively understudied and requires further investigation. Since the present study focuses on the thermal  
 158 landscape characterized by LST, we can directly quantify the spacing between the identified heat islets. Popular  
 159 metrics such as root mean square distances work well for Gaussian systems, but for fractal landscapes, lacunarity is  
 160 a better-suited metric of spatial structure (Plotnick et al., 1996). Lacunarity ( $\Lambda$ ) is a scale-dependent measure of the  
 161 aggregation of spaces between the heat islets (Mandelbrot, 1982; Plotnick et al., 1996). A ‘gliding box’ algorithm for  
 162 the calculation of  $\Lambda$  as a function of box size ( $r$ ), as described in Plotnick et al. (1993), was adopted here (Methods  
 163 section). While the absolute values of  $\Lambda$  offer little insight, the appropriate way to interpret lacunarity is in the context  
 164 of the rate of change of  $\Lambda$  as a function of  $r$ . If the value of  $\log(\Lambda(r))$  decreases at any scale (quantified with  $\log(r)$ ),  
 165 the presence of spacing corresponding to that length scale is indicated. The two extremes of lacunarity curvature can  
 166 be best conceptualized as a chessboard-type homogeneous distribution of small-scale spacing, and a single contiguous  
 167 cluster. Essentially, the length scales corresponding the steepest slopes should be interpreted as the dominant scale of  
 168 spacing.

169 As the differences in the spatial organization of heat islets are most apparent at higher temperature thresholds, here,  
 170 we characterized the spatial structure obtained at the 90<sup>th</sup> percentile of LST for all cities. By extension, the total  
 171 *islet area* under consideration corresponds to the hottest 10% of the total city area. Lacunarity curves for the four  
 172 representative cities investigated in the previous section are highlighted in Figure 3. The cities that have a dominance  
 173 of larger spacing between the islets lay above the diagonal. Conversely, a dispersed spatial structure of the heat islets  
 174 manifests as smaller spacing, fall under the diagonal. We assign a single score ( $\Lambda_{score}$ ) to the convexity of the curves  
 175 in Figure 3a such that positive scores indicate larger spacing and vice-versa. This is achieved using the following  
 176 empirical equation:

$$\log_{10}(\Lambda(r)) = \left(1 - \frac{\log_{10}(r)}{2}\right)^{2^{\Lambda_{score}}} \quad (3)$$

177 where constants 1 and 2 are used to fix the end points of the curve at  $\log(\Lambda(r)) = 1$  and  $\log(r) = 2$ , and the exponent,  
 178  $\Lambda_{score}$  is scale-independent measure of the shape of the lacunarity curve (See Methods section). The 49 cities have  
 179  $\Lambda_{score}$  ranging between -0.9 to 0.6, and distributed normally (Figure 3b; See Supplementary Table 4).

180 Using  $\Lambda_{score}$ , we compare the relationship between the islet spacing and their NPR (and by extension, likely expo-  
 181 nential tempering at higher thresholds). We find that the dense cities associated with an aggregated heat islet structure  
 182 (positive  $\Lambda_{score}$ ) display a larger NPR ( $\geq 0.25$ ; Figure 3c). Whereas, sprawling and disaggregated cities (negative  
 183  $\Lambda_{score}$ ) have a smaller NPR ( $< 0.25$ ; Figure 3c) and consequently an exponential tempering of the power law tail  
 184 (Figure 2f). An exception to this pattern are cities with a negative  $\Lambda_{score}$  despite having an NPR  $\geq 0.25$  (shown in  
 185 yellow in Figure 3c). Upon examination, we found these to have a significant river flowing right through them. Under  
 186 such a scenario, the percolating heat cluster is divided structurally into two halves by a heat sink, irrespective of the  
 187 threshold (Supplementary Figure 4). This results in a negative  $\Lambda_{score}$  due to the spacing introduced by the river despite  
 188 an aggregation of heat islets on either side of the river. Thus, Figure 3c serves to quantitatively affirm the correlation  
 189 between the spatial configuration of cities (dense and/or sprawling) and the 2 classes of size distributions of the heat  
 190 islets.

191 Note that for any given size distribution, the islets can be spatially arranged in several ways. In order to examine  
 192 the variability in islet size and spacing of the various cities, we define two scale-independent metrics to characterize  
 193 size: Mean ( $A_M$ ) and Largest ( $A_L$ ) Relative Heat Islet Sizes, calculated as a percentage of the total city area. First,  
 194 we observe that there is a weak positive correlation ( $R^2 = 0.4$ ) between  $A_M$  and spacing of the heat-islets (Figure  
 195 3d). This is expected because a positive  $\Lambda_{score}$  as well as a high  $A_M$  corresponds to dense cities, and a negative  
 196  $\Lambda_{score}$  and low  $A_M$  corresponds to sprawling cities. More noteworthy is the horizontal spread about the diagonal in  
 197 Figure 3d which reflects the different spatial configurations (characterized by  $\Lambda_{score}$ ) that are possible for any given  
 198 size distribution. This spread may be explained by  $A_L$ , which increases with  $\Lambda_{score}$  (illustrated using marker size in  
 199 Figure 3d; Supplementary Figure 5). In the bottom-left, both  $A_M$  and  $A_L$  are small. This is because negative  $\Lambda_{score}$   
 200 corresponds to sprawling cities where large clusters were absent in the islet-size distribution (as inferred from the  
 201 exponential tempering of Pareto). In the bottom-right, however, the dominance of the largest aggregated islet results  
 202 in a positive  $\Lambda_{score}$  despite a low  $A_M$  value. A schematic diagram drawn to represent each of the vertices of this  
 203 plot is given as Supplementary Figure 6. The phase plot of  $A_M$  and  $\Lambda_{score}$  may be useful for city planners to gauge  
 204 the current spatial structure of the thermal landscape of their cities and to determine mitigation strategies to achieve a  
 205 more desirable state.

## 206 Islet Intensity distribution

207 Apart from the size and spacing of heat islets, we now focus on the temperatures obtained within the heat islets. To  
 208 address this, we first use the well-known indicator of excess heat in urban areas, the SUHI Intensity in the traditional  
 209 sense i.e., the difference between the mean urban and rural temperatures (Schwarz et al., 2011) to evaluate the average  
 210 excess heat within cities. We find that larger  $\Lambda_{score}$  values (representative of aggregated heat islets) tend to be associ-  
 211 ated with higher SUHI Intensity (Figure 4b). This suggests that sprawling cities, with a larger number of heat sinks to  
 212 match the heat sources, are a better configuration for reducing the *overall* SUHI Intensity. This is in agreement with our  
 213 findings based on the size distribution of extreme heat islets as well as Debbage and Shephard (2015) findings based  
 214 on discontinuity of urban patches calculated using National Land Cover Dataset (NLCD) (Debbage and Shepherd,  
 215 2015). Traditional estimates of the UHI Intensity that simply use the difference between the **mean** temperatures over  
 216 an urban area and the surrounding non-urban environment fail to address the intra-urban heterogeneity adequately.

217 For a more comprehensive assessment of the thermal variability within cities, we introduce a novel *Heat Islet*  
 218 *Intensity* distribution metric. First, we compute the excess heat ( $\Delta T$ ) for each islet as the difference between the mean  
 219 *islet* temperature and the *threshold* temperature. We refer to this term as the *Islet Intensity*. We find that the mean and  
 220 standard deviation of  $\Delta T$  were equal (Supplementary Figure 7) which, along with the shape of its distribution (Figure  
 221 4a), were indicative that  $\Delta T$  is exponentially distributed, i.e:

$$P(\Delta T \geq x) \propto 1 - e^{-\lambda x} \quad (4)$$

222 where, for a given islet intensity  $x$ , the probability of an islet having a temperature  $\Delta T$  larger than  $x$  is represented  
 223 by  $P$ , an exponential distribution characterized by  $\lambda$ . By extension,  $1/\lambda$  is the mean islet intensity. Lower values of

224  $\lambda$  correspond to an increased probability of higher temperatures within the islets. Therefore, a single metric,  $\lambda$  can be  
 225 used as an indicator to capture the intra-urban thermal variability across islets. This is represented as the color bar in  
 226 Figure 4b.

227 We find that while cities with a higher degree of sprawl have a lower mean temperature, for the same SUHI (Y-axis in  
 228 Figure 4b), cities with lower  $\Lambda_{score}$  also experience higher likelihood of encountering thermal extremes. For example,  
 229 dense cities such as Lagos and Jakarta have a steeper exponential decaying rate than Chicago and Guangzhou, which  
 230 drastically reduces the probability of local thermal extremes within their heat islets. While the probability of a heat  
 231 islet being hotter than the mean by  $1^\circ\text{C}$  is almost zero for the first two, the likelihood increases to roughly 20% for the  
 232 latter two (Figure 4a). As the larger heat islets are often associated with the highest islet intensity as well, this can result  
 233 in a significantly large areas of extreme heat especially for megacities like Guangzhou and Chicago. Such a finding  
 234 reveals that while mean SUHI Intensity decreases with sprawling cities, for the same mean, they also experience higher  
 235 local thermal extremes. As a result, in addition to the mean SUHI Intensity, it is essential to characterize the thermal  
 236 heterogeneity within the cities, and the islet intensity distribution can be adopted as a complementary metric.

## 237 **Summary and Conclusions**

238 Cities grow through a combination of parallel and sequential episodes of expansion and densification at different  
 239 rates. Depending on local preferences and constraints, neighborhoods are built with different spatial patterns, for  
 240 example, from dense downtowns to sprawling suburbs. Factors like geographical topography, coastline, and intra-  
 241 urban commuting time constrain expansion, whereas other factors such as local building laws limit densification.  
 242 While there are several objective functions such as commuting travel time distribution, net carbon emissions, and  
 243 socio-economical factors which urban form and functions are optimized for, here, we focus on the aspect of urban  
 244 heat. More specifically, the spatial heterogeneity of extreme heat islets within urban areas. Towards that, we present  
 245 a novel multi-scale framework that allows us to identify intra-urban heat islets for several thermal thresholds. Using  
 246 this framework, we evaluate the impact of spatial organization, characterized by a Lacunarity-based metric,  $\Lambda_{score}$ .  
 247 We do not find a bi-modal distribution of  $\Lambda_{score}$  corresponding to the two classes of sprawl or dense cities. Rather,  
 248 the  $\Lambda_{score}$  was normally distributed around a mean value close to zero indicating that most cities display a balance  
 249 between sprawl and dense heat islet structure. Different realizations and degrees of expansion and densification yield  
 250 a diverse array of spatial structures.

251 We condense the size, spacing, and intensity information about heterogeneous clusters into probabilistic distribu-  
 252 tions that can be described using single scaling exponents. This allows for a seamless comparison of the intra-urban  
 253 heat islet characteristics across cities at several spatial scales ranging from 90 meters (resolution of Landsat 8 and  
 254 corresponding to several urban blocks) up to few thousand sq. km (total area of large cities). We implement this  
 255 framework for 78 globally representative cities to answer the following key questions. First, how many and how big  
 256 are the emergent heat islets at multiple thermal thresholds? Second, how much hotter than the threshold are these  
 257 heat islets? From the size distribution analysis, we demonstrate that islet sizes in dense cities follow and maintain  
 258 a power law tail across all temperature thresholds, whereas the sprawling cities show an exponential tempering of  
 259 tails at higher thresholds. Such a tempering is favourable as it indicates a reduced emergence of large heat islets in



260 sprawling and dispersed spatial configurations. Additionally, a dispersed configurations results in lower *mean SUHI*  
261 *Intensity* over the city. However, from the islet intensity distribution analysis, we find that heat islet intensities ( $\Delta T$ )  
262 can be modelled as exponential distributions, where dispersed configurations result in higher rate parameters,  $\lambda$ . This  
263 implies a significantly higher probability of encountering extreme temperatures within the islets. As a result, while a  
264 sprawling configuration is favorable for reducing the mean temperature of a city, for the same mean SUHI intensity, it  
265 results in higher local thermal extremes. The implications of this from a design perspective are that: (i) while design-  
266 ing sprawling cities, higher intra-urban variability should be expected. (ii) With changes in urban morphology, we can  
267 attempt to control the precise spatial location of the extreme temperatures. It is of utmost importance that the spatial  
268 locations of these extremities do not coincide with the high risk zones where most vulnerable populations reside, such  
269 as densely populated downtowns, or areas without access to air-conditioning such as urban slums.

270 At this juncture, we reiterate that while the spatial characterization of temperature is informative for urban heat  
271 assessments, it does not inform the overall risk to the concerned population. Risk is a combination of hazard (i.e.  
272 extreme heat-stress in this case, a combination of air temperature and humidity (Oleson et al., 2015)), the time period  
273 of exposure to heat stress, and vulnerability. Prior research shows that in the absence of hydrological processes, the  
274 changes in radiative flux, determined by LST, contribute more to the near-surface temperature changes than the turbu-  
275 lent heat flux (Brownlee et al., 2017). As a result, similar spatial patterns of heat islets derived from air temperatures  
276 can be expected. Vulnerability assessments will require further consideration of demographic factors such as old age,  
277 low educational attainment, high poverty levels, poor health, and lack of air conditioning (Cutter et al., 2009; Uejio  
278 et al., 2011; Bradford et al., 2015). An investigation of the spatio-temporal dynamics of risk to extreme heat is beyond  
279 the scope of the present study but it certainly warrants further research.

280 Lastly, our analysis here is limited to the structural heterogeneity of heat sources and sinks, and not the functional  
281 heterogeneity. If we assume a uniform heat capacity of land use, the sizes of heat islets are then indicative of the  
282 strength of the sources, and the length-scale of spacing is indicative of the sink strengths. Consideration of the *func-*  
283 *tional* heterogeneity will require a treatment of the variability in heat capacities and thermal conductivities of the land  
284 use which jointly determine heat dissipation from sources, which is possible using models such as Weather Research  
285 Forecast (WRF) (Chen et al., 2011; Salamanca et al., 2011). In such a scenario, instead of LST, heat fluxes can be  
286 treated as DEM for such an analysis. It may also then be beneficial to study the spatial correlation between source  
287 strength and sink strength to evaluate thermal dissipation. To conclude, our work presents the first steps toward a  
288 multi-scale characterization of the complex intra-urban thermal landscape and we hope that it opens new vistas for  
289 future investigations.

## 290 **Methods**

### 291 **Study area and data sources**

292 Land surface temperature (LST) data was derived using a Single Channel Algorithm as detailed in (Walawender et al.,  
293 2012) from Landsat 8 at a resolution of 90 m. The geo-spatial analysis environment of Google Earth Engine (GEE)  
294 was used to filter out cloud free summer time days with an incident solar angle of at least 60 degrees (Gorelick et al.,

295 2017). See Supplementary Text S1 and S2 for algorithms, and Table S1 for list of cities and information on Landsat  
 296 scenes used. For coastal cities the Large Scale International Boundary (LSIB) dataset provided by United States Office  
 297 of the Geographer was used to crop out the oceans and delineate urban boundaries within GEE environment. Urban  
 298 area was estimated using MODIS's Land Cover Type dataset - MCD12Q1.

### 299 **Statistical modelling of size and intensity distributions**

300 For fitting probability distribution functions (pdfs) to cluster size and intensity distributions, a combination of  
 301 maximum-likelihood estimation (mle) with goodness-of-fit tests based on the Kolmogorov-Smirnov (KS) statistic  
 302 and likelihood ratios were used (Clauset et al., 2009). See Supplementary Text S3 for details and Table S2 for results.

### 303 **Lacunarity**

304 First, the landscape was sliced at a thermal threshold and an islets map was obtained. For each box size ( $1 < r <$   
 305  $A_{city}$ ), the number of occupied pixels (islets) was measured. The number of occupied sites was referred to as the  
 306 box mass. The box was then moved one column to the right and the box mass was again counted. This process was  
 307 repeated over all rows and columns producing a frequency distribution of the box masses. The number of boxes of size  
 308  $r$  containing  $S$  occupied sites was designated by  $n(S,r)$  and the total number of boxes of size  $r$  by  $N(r)$ . This frequency  
 309 distribution was converted into a probability distribution:  $Q(S,r) = \frac{n(S,r)}{N(r)}$ . Lacunarity was a measure of variability  
 310 in the calculated occupancy for each box size.

$$\Lambda(r) = \frac{\text{Variance}Q(S,r)}{\text{Mean}Q(S,r)^2} + 1 \quad (5)$$

311 For all cities, Lacunarity score was calculated only for the 90<sup>th</sup> percentile thermal threshold. As a result, 90% of the  
 312 total area in all cases comprised of spaces and the  $\Lambda(r)$  value for box size = 1 was the same for all cities. The largest  
 313 box size taken under consideration is normalized from 0 to 100 in order to account for the variable sizes of cities. Note  
 314 that the curvature of Lacunarity curve was unaffected by these transformations.

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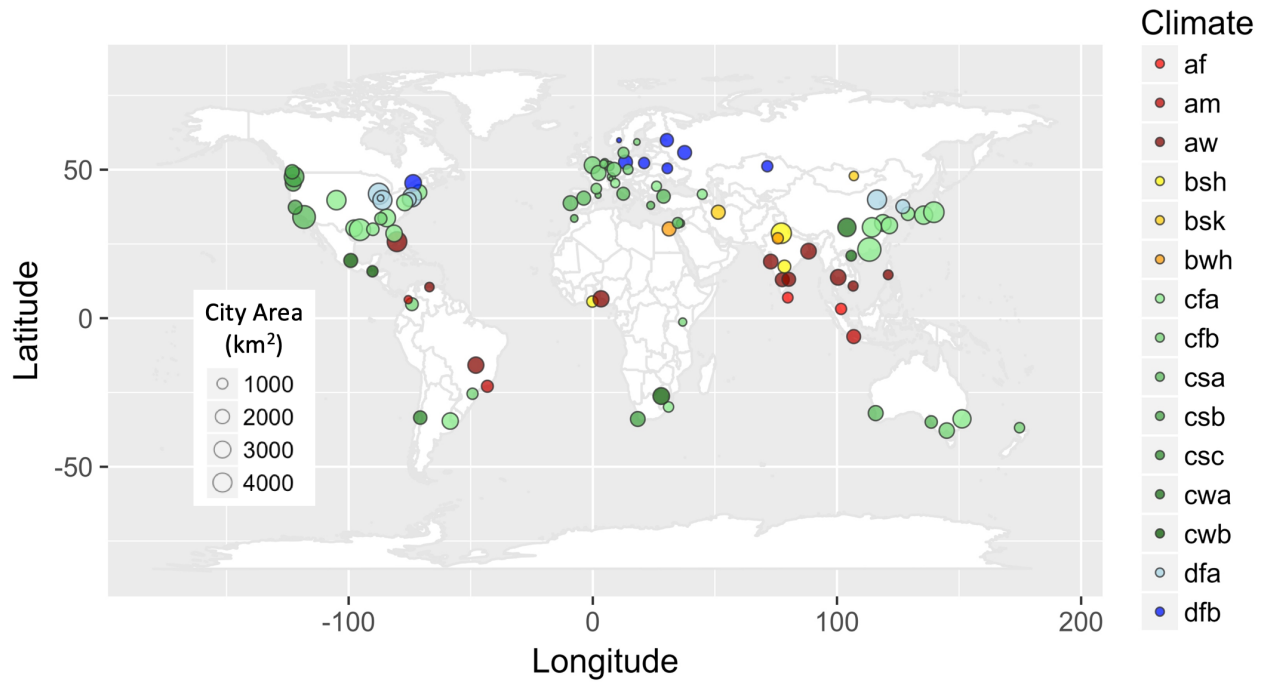
431 **Figures**

Figure 1: World map showing the locations of 78 cities considered in this study. The marker size is representative of the city size, and the colour represents their Koppen-Geiger climate classification Peel et al. (2007). Description of Koppen-Geiger climate types are given in Supplementary Table 1.

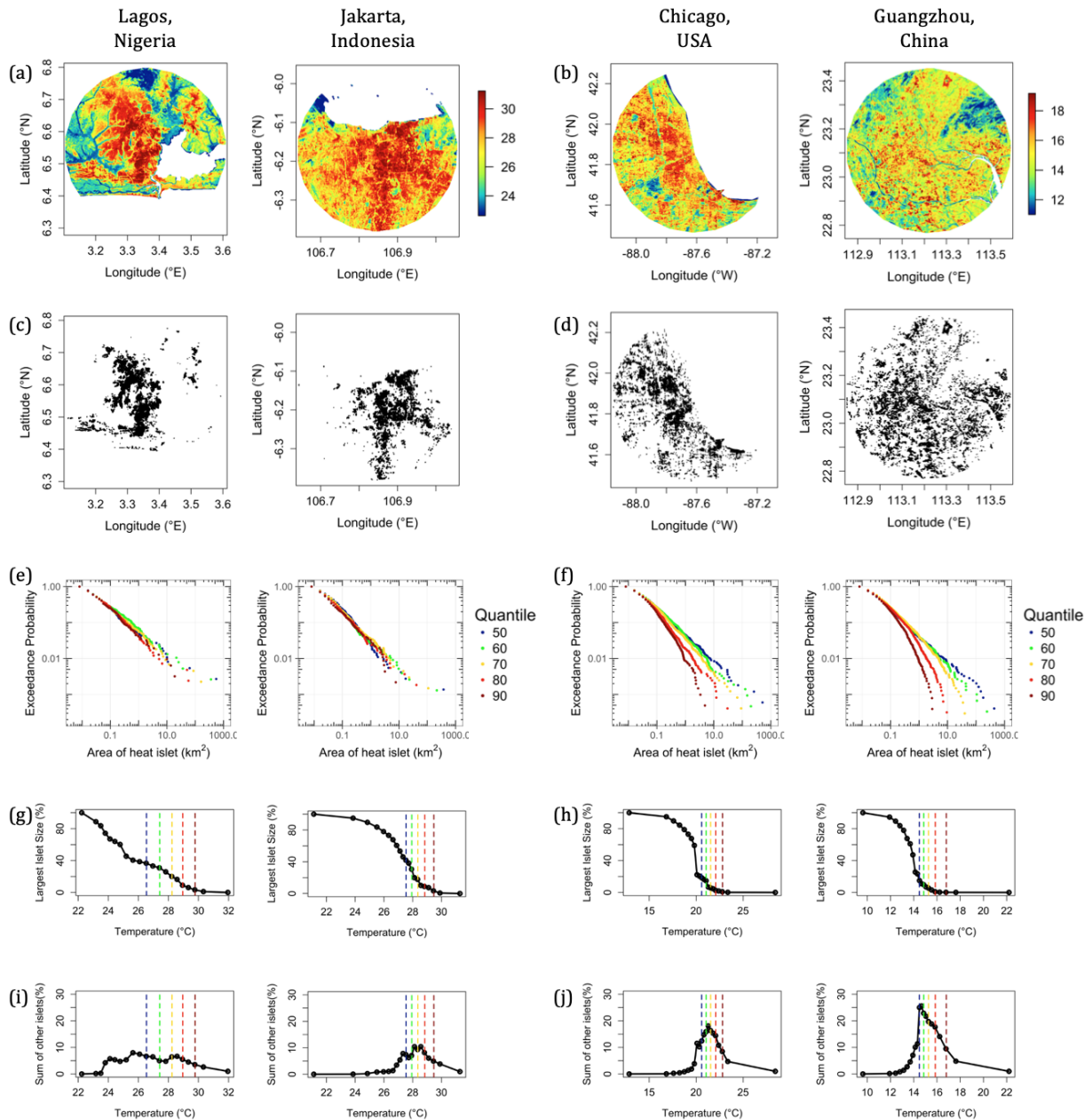


Figure 2: Two groups of cities emerge based on the size distributions of heat islets at incremental thermal thresholds. Two representative cities for each group - Jakarta, Indonesia and Lagos, Nigeria for dense cities, and Chicago, USA, and Guangzhou, China for sprawling cities - are shown. (a,b) Land Surface Temperature map (in °C), (c,d) Heat islets that emerge at the 90<sup>th</sup> percentile thermal threshold, (e,f) Exceedance probability plots for heat islets at several thermal thresholds (50<sup>th</sup>, . . . , 90<sup>th</sup>). Note the leftward shift in size distribution as the thresholds increase, especially the exponential tempering evident in sprawling cities, (g,h) Largest islet size, and (i,j) sum of remaining islets (as a % of total city area), as a function of thermal threshold. The vertical dashed coloured lines mark the temperatures corresponding to the percentiles used in (e,f).

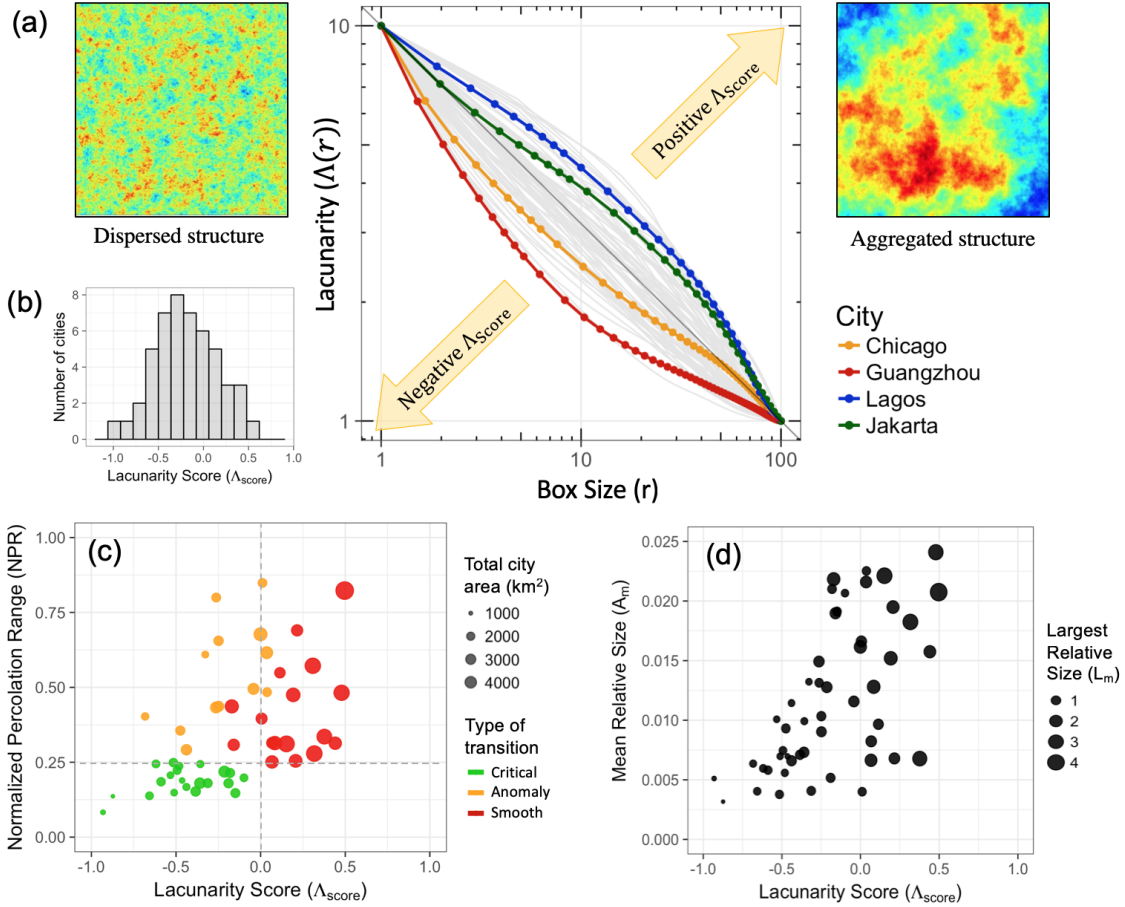


Figure 3: (a) Lacunarity curves of 49 cities (in grey) and the four archetype cities (in colour) shown on a  $\log(\Lambda)$  vs  $\log(r)$  plot. The cities with a concave downwards shape in the upper side of the diagonal indicate larger and more aggregated gaps, whereas cities underneath the curve indicate a more uniform dispersed pattern of islets and smaller gaps. (b) Histogram of  $\Lambda_{score}$  of 49 cities (mean = 0.04, s.d. = 0.38). (c) Scatter plot of percolation transition range and Lacunarity score. This figure illustrates the classification of cities into the 2 classes based on Lacunarity Score and the type of transition. (d) Scatter plot of Mean Relative Heat Islet Size ( $A_M$ ) versus  $\Lambda_{score}$ . Additionally, since the islet-size distribution is heavy tailed, in addition to the  $A_M$ , the largest islet size (as a percentage of the total city area) is indicated using the marker size. The  $A_M$  and the largest-heat islet size ( $A_L$ ) serve to illustrate the size distribution of the hottest islets occupying the ten percent of the city area.



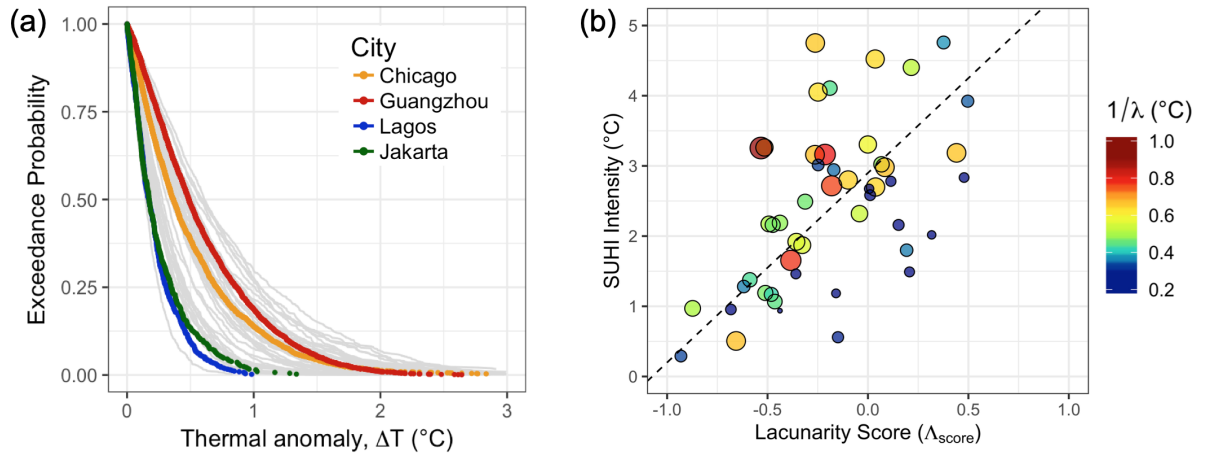


Figure 4: (a) Empirical pdf of  $\Delta T$  for the 4 archetype cities shown on at their 90<sup>th</sup> percentile thermal thresholds respectively. The same for all 49 cities is shown in grey in the background. Each  $\Delta T$  distribution was well described as an exponential distribution characterized by the parameter:  $\lambda$ . (b) Scatter plot of mean SUHI Intensity, defined as the difference between mean urban and rural temperatures, versus Lacunarity Score ( $\Lambda_{score}$ ) is shown. A weak positive correlation ( $R^2 = 0.344$ ) is detected shown as dashed regression line. The color as well as size of the marker indicates the inverse of rate parameter ( $\lambda$ ) from Equation 4 which is equal to the mean Heat Islet Intensity for each distribution. Increasing size indicates higher temperatures within the heat islets.