

GARPOS: analysis software for the GNSS-A seafloor positioning with simultaneous estimation of sound speed structure

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11 **Abstract**

12 Global Navigation Satellite System – Acoustic ranging combined seafloor geodetic technique
13 (GNSS-A) has extended the geodetic observation network into the ocean. The key issue for analyzing
14 the GNSS-A data is how to correct the effect of sound speed variation in the seawater. We
15 constructed a generalized observation equation and developed a method to directly extract the
16 gradient sound speed structure by introducing appropriate statistical properties in the observation
17 equation, especially the data correlation term. In the proposed scheme, we calculate the posterior
18 probability based on the empirical Bayes approach using the Akaike’s Bayesian Information
19 Criterion (ABIC) for model selection. This approach enabled us to suppress the overfitting of sound
20 speed variables and thus to extract simpler sound speed field and stable seafloor positions from the
21 GNSS-A dataset. The proposed procedure is implemented in the Python-based software “GARPOS”
22 (GNSS-Acoustic Ranging combined POSitioning Solver).

23 **1 Introduction**

24 **1.1 Basic configurations of the GNSS-A observation**

25 Precise measurements of seafloor position in the global reference frame opens the door to the
26 “global” geodesy in the true sense of the word. It extended the observation network for crustal
27 deformation into the ocean and has revealed the tectonic processes in the subduction zone including
28 megathrust earthquakes (e.g., Bürgmann and Chadwell, 2014; Fujimoto, 2014, for review). Many
29 findings have been reported especially in the northwestern Pacific along the Nankai Trough (e.g.,
30 Yokota et al., 2016; Yasuda et al., 2017; Yokota and Ishikawa, 2020), and the Japan Trench (e.g.,
31 Sato et al., 2011; Kido et al., 2011; Watanabe et al., 2014; Tomita et al., 2015). These achievements
32 owe to the development of GNSS-A (Global Navigation Satellite System – Acoustic ranging
33 combined) seafloor positioning technique, proposed by Spiess (1980).

34 Observers can take various ways to design the GNSS-A observation for the positioning of the
35 seafloor benchmark. They have to solve the difficulties not only in the technical realizations of

36 GNSS-A subcomponents such as the acoustic ranging and the kinematic GNSS positioning, but also
37 in designing the observation configurations and analytical models to resolve the strongly correlated
38 parameters. For example, because the acoustic ranging observations are performed only on the sea
39 surface, the errors in sound speed perturbations are strongly correlated with the relative distance,
40 typically the depths of the benchmark.

41 In the very first attempt for the realization, Spiess et al. (1998) derived horizontal displacement using
42 a stationary sea-surface unit which was approximately placed on the horizontal center of the array of
43 multiple seafloor mirror transponders. They determined the relative positions and depths of the
44 transponders in advance. The relative horizontal positions of the sea-surface unit to the transponder
45 array can be determined by acoustic ranging data, to be compared with the global positions
46 determined by space geodetic technique. In this “stationary” GNSS-A configuration, the temporal
47 variation of sound speed is less likely to affect the apparent horizontal position under the assumption
48 that the sound speed structure is horizontally stratified. Inversely, comparing the residuals of acoustic
49 travel time from multiple transponders, Osada et al. (2003) succeeded in estimating the temporal
50 variation of sound speed from the acoustic data. Kido et al. (2008) modified the expression to
51 validate the stationary configuration for a loosely tied buoy even in the case where the sound speed
52 has spatial variations. The stationary GNSS-A configuration is applied mainly by the groups in the
53 Scripps Institution of Oceanography (e.g., Gagnon et al., 2005; Chadwell and Spiess, 2008) and in
54 the Tohoku University (e.g., Fujimoto et al., 2014; Tomita et al., 2015).

55 On the other hand, Obana et al. (2000) and Asada and Yabuki (2001) took a “move-around”
56 approach where the 3-dimensional position of single transponder can be estimated by collecting the
57 acoustic data from various relay points on the sea surface. Figure 1 shows the schematic image of
58 move-around configuration. The move-around GNSS-A configuration is developed and practicalized
59 mainly by the collaborative group of the Japan Coast Guard and the University of Tokyo, and the
60 Nagoya University. Unlike the stationary configuration, the horizontal positions of transponders are
61 vulnerable to bias errors of sound speed field. Fujita et al. (2006) and Ikuta et al. (2008) then
62 developed the methods estimating both the positions and the temporal variations of sound speed.

63 Similar to the effects of distribution of the GNSS satellites on the positioning, well-distributed
64 acoustic data is expected to decrease the bias errors of the estimated transponders’ positions in the
65 move-around configuration. By implementing the sailing observations where the sea-surface unit
66 sails over the transponder array to collect geometrically symmetric data, positioning accuracy and
67 observation efficiency have improved (Sato et al., 2013; Ishikawa et al., 2020).

68 In order to enhance the stability of positioning, an assumption that the geometry of transponder array
69 is constant over whole observation period is usually adopted (e.g., Matsumoto et al., 2008; Watanabe
70 et al., 2014; Yokota et al., 2018). Misestimates of sound speed cause the positional biases parallel to
71 the averaged acoustic-ray direction, which results in the distortion of the estimated array geometry.
72 Constraining the array geometry contributes to reducing the bias error in the sound speed estimates
73 and the transponders’ centroid position.

74 It should be noted that these two configurations are compatible under the adequate assumptions and
75 constraints. Recently, the group in the Tohoku University uses not only the stationary but also the
76 move-around observation data collected for determining the array geometry (Honsho and Kido,
77 2017).

78 1.2 Recent improvements on GNSS-A analytical procedures

79 In the late 2010s, analytical procedures with the estimation of the spatial sound speed gradient for the
80 move-around configuration have been developed. In the earlier stage of the move-around GNSS-A
81 development, the spatial variations of sound speed were approximated as the temporal variations,
82 because most of the sound speed change are confined in the shallowest portion along the acoustic ray
83 paths (e.g., Watanabe and Uchida, 2016). Actually, Yokota et al. (2019) extracted the sound speed
84 gradient in the shallow layer from the temporally expanded sound speed corrections. However, the
85 smoothly expanded temporal variations cannot represent the transponder-dependent variation which
86 is caused by the sound speed gradient in the relatively deeper portion. Therefore, Yokota et al. (2019)
87 extracted the transponder-dependent correction term from the residuals of the results derived by the
88 conventional method of Fujita et al. (2006).

89 Yasuda et al. (2017) took a different approach where the sound speed structure shallower than 1000
90 m is assumed to be one-dimensionally inclined due to the Kuroshio current flowing near their sites in
91 the offshore region south of Kii Peninsula, Japan. Because their model reflects the specific
92 oceanographic feature, the estimated parameters are easier to be interpreted than that of Yokota et al.
93 (2019) which has higher degree of freedom to extract the oceanographic features as shown in Yokota
94 and Ishikawa (2019).

95 Meanwhile, Honsho et al. (2019) showed a more general expression for one-dimensional sound
96 speed gradient. As they mentioned, the gradient terms in their formulation correspond to the
97 extracted features in Yokota et al. (2019). The work by Honsho et al. (2019) showed the possibility to
98 connect all the GNSS-A configurations into a unified GNSS-A solver. However, due to the limitation
99 in resolving the general gradient structure, an additional constraint was taken for the practical
100 application, which concludes to essentially the same formulation as Yasuda et al. (2017).

101 In this study, to overcome the limitation above, we propose a method to directly extract the gradient
102 sound speed structure by introducing appropriate statistical properties in the observation equation.
103 This paper first shows the reconstructed general observation equation for GNSS-A, in which only the
104 continuity of the sound speed field in time and space is assumed. The generalized formulation
105 approximately includes the practical solutions in the previous studies by Yokota et al. (2019), Yasuda
106 et al. (2017), and Honsho et al. (2019) as special cases. We then describe the analytical procedure to
107 derive the posterior probability based on the empirical Bayes approach using the Akaike's Bayesian
108 Information Criterion (ABIC; Akaike, 1980) for model selection. We obtain the solution which
109 maximizes the posterior probability under the empirically selected prior distribution. This is
110 implemented in the Python-based software "GARPOS" (GNSS-Acoustic Ranging combined
111 POsitioning Solver; Watanabe et al., 2020a, available at <https://doi.org/10.5281/zenodo.3992688>).

112 **2 Methodology**

113 **2.1 Positioning of sea-surface transducer**

114 The key subcomponent of the GNSS-A is the global positioning of the transducer, generally realized
115 by GNSS observation. Whereas acoustic measurement determines the relative positions of the
116 seafloor transponders and the sea-surface transducer, GNSS plays a role to align them to the earth-
117 centered, earth-fixed (ECEF) coordinates such as the International Terrestrial Reference Frame
118 (ITRF). In terms of GNSS positioning, the transducer's position, $\mathbf{P}(t)$, is assumed as the orbit of the
119 GNSS satellites. When $\mathbf{P}(t)$ is determined in the GNSS's reference frame, a realization of the ITRF,
120 the global positions of transponders can be estimated.

121 It should be noted that the transponders' positions are generally a function of time, including the
122 solid earth tide as well as global and local crustal deformation (e.g., IERS Conventions, 2010). For
123 the purpose of detecting crustal deformation, it is better to determine the seafloor positions in the
124 solid-earth-tide-free coordinates. Because the observation area is limited to several-kilometers-width,
125 solid-earth-tide-free solutions can be obtained when the trajectory of the transducer is determined in
126 the solid-earth-tide-free coordinates. Hereafter, the positions are expressed in solid-earth-tide-free
127 coordinates in this paper.

128 In order to determine $\mathbf{P}(t)$ in the ECEF coordinates, a set of GNSS antenna/receiver and a gyro
129 sensor should be mounted on the sea-surface unit. The positions of GNSS antenna, $\mathbf{Q}(t)$, can be
130 determined using any of appropriate kinematic GNSS solvers. The gyro sensor provides the attitude
131 of the sea-surface platform, $\boldsymbol{\theta}(t) = [\theta_r \ \theta_p \ \theta_h]^T$, i.e., roll, pitch, and heading (Figure 2). Because
132 the attitude values are aligned to the local ENU coordinates, it is convenient to transform $\mathbf{Q}(t)$ from
133 ECEF to local ENU coordinates, i.e., $\mathbf{Q}(t) = [Q_e \ Q_n \ Q_u]^T$. Using the relative position of the
134 transducer to the GNSS antenna in the gyro's rectangular coordinate (called "ATD offset" hereafter;
135 Figure 2), $\mathbf{M} = [M_r \ M_p \ M_h]^T$, we obtain the transducer's position in the local ENU coordinates
136 as,

$$137 \quad \mathbf{P}(t) = \mathbf{Q}(t) + R(\boldsymbol{\theta}(t))\mathbf{M} \quad (1.1)$$

138 with,

$$139 \quad R(\boldsymbol{\theta}) = \begin{bmatrix} 0 & 1 & 0 \\ 1 & 0 & 0 \\ 0 & 0 & -1 \end{bmatrix} \begin{bmatrix} \cos \theta_h & -\sin \theta_h & 0 \\ \sin \theta_h & \cos \theta_h & 0 \\ 0 & 0 & 1 \end{bmatrix} \begin{bmatrix} \cos \theta_p & 0 & \sin \theta_p \\ 0 & 1 & 0 \\ -\sin \theta_p & 0 & \cos \theta_p \end{bmatrix} \begin{bmatrix} 1 & 0 & 0 \\ 0 & \cos \theta_r & -\sin \theta_r \\ 0 & \sin \theta_r & \cos \theta_r \end{bmatrix} \quad (1.2)$$

140 The ATD offset values should be measured before the GNSS-A observation.

141 2.2 Underwater acoustic ranging

142 Another key subcomponent is the technique to measure the acoustic travel time between the sea-
143 surface transducer and the seafloor transponders. The techniques for the precise ranging using
144 acoustic mirror-type transponders had been developed and practicalized in early studies (e.g., Spiess,
145 1980; Nagaya, 1995). Measuring round-trip travel time reduces the effect of advection of the media
146 between the instruments.

147 The round-trip travel time for the i th acoustic signal to the j th transponder, T_i , is calculated as a
148 function of the relative position of the transponder to the transducer and the 4-dimensional sound
149 speed field, $V(e, n, u, t)$, i.e.,

$$150 \quad T_i = T_i^c \left(\mathbf{P}(t_{i+}), \mathbf{P}(t_{i-}), \mathbf{X}_j, V(e, n, u, t) \right) \quad (2)$$

151 where t_{i+} , t_{i-} , and \mathbf{X}_j are the transmitted and received time for the i th acoustic signal, and the
152 position of seafloor transponder numbered j , respectively. Note that j is a function of i .

153 Although the concrete expression is provided as the eikonal equation (e.g., Jensen et al., 2011; Sakic
154 et al., 2018), it requires much computational resources to numerically solve. When the sound speed
155 structure is assumed to be horizontally stratified, we can apply a heuristic approach based on the

156 Snell's law (e.g., Hovem, 2013), which has an advantage in computation time (e.g., Chadwell and
157 Sweeney, 2010; Sakic et al., 2018).

158 Therefore, we decomposed the 4-dimensional sound speed field into a horizontally stratified stationary
159 sound speed profile and a perturbation to obtain the following travel time expression:

$$160 \quad T_i^c(\mathbf{P}(t_{i+}), \mathbf{P}(t_{i-}), \mathbf{X}_j, V(e, n, u, t)) = \exp(-\gamma_i) \cdot \tau_i(\mathbf{P}(t_{i+}), \mathbf{P}(t_{i-}), \mathbf{X}_j, V_0(u)) \quad (3)$$

161 where τ_i and $V_0(u)$ denote the reference travel time and the reference sound speed profile,
162 respectively. $V_0(u)$ is given as a piecewise linear function of height, so that the propagation length
163 along the radial component and the propagation time can be calculated for the given incidence angle
164 according to the Snell's law (e.g., Hovem, 2013; Sakic et al. 2018). The expression of the correction
165 coefficient, $\exp(-\gamma_i)$, is selected for the simplification in the following expansion. It represents the
166 discrepancy ratio of the actual travel time to the reference, which caused by the spatial and temporal
167 perturbations of the sound speed field.

168 In the right-hand side of equation 3, γ_i and \mathbf{X}_j are assigned as the estimator. Equation 1 gives the
169 transducer's position $\mathbf{P}(t)$ as a function of the GNSS antenna's position $\mathbf{Q}(t)$, the attitude vector
170 $\boldsymbol{\theta}(t)$, and the ATD offset \mathbf{M} . The time-independent parameter \mathbf{M} can be also assigned as the
171 estimator when the variation of the attitude value is large enough to resolve the parameter. Hence, the
172 reference travel time can be rewritten as $\tau_i = \tau_i(\mathbf{X}_j, \mathbf{M} | \mathbf{Q}(t), \boldsymbol{\theta}(t), V_0(u))$, where the variables on
173 the left and right sides of the vertical bar indicate the estimators and the observables, respectively.

174 2.3 Sound speed perturbation model

175 In seawater, sound speed is empirically determined as a function of temperature, salinity, and
176 pressure (e.g., Del Grosso, 1974). Because these variables strongly depend on the water depth, the
177 vertical variation of the sound speed is much larger than the horizontal variation in the observation
178 scale. Thus, $|\gamma_i| \ll 1$ will be satisfied in most cases where the reference sound speed appropriately
179 represents the sound speed field. In such cases, the average sound speed along the actual ray path is
180 expressed as $\bar{V}_0 + \delta V_i \sim \bar{V}_0 + \gamma_i \bar{V}_0$, where \bar{V}_0 denotes the average sound speed of the reference
181 profile.

182 Recalling that the sound speed field is continuous and usually smooth in time and space compared to
183 the sampling rates of acoustic data, the acoustic ray path also has continuity in time and positions of
184 both ends, within the observation scale. It means that the acoustic rays from/to the neighboring ends
185 transmitted at almost the same time will take almost the same paths. Thus, γ_i can be modeled with a
186 smooth function of time and acoustic instruments' positions for the transmission and return paths,
187 i.e., $\gamma_i \equiv \frac{1}{2} \sum_{l=i+, i-} \Gamma(t_l, \mathbf{P}(t_l), \mathbf{X}_j)$. The function $\Gamma(t, \mathbf{P}, \mathbf{X})$ can be called the sound speed
188 perturbation model.

189 For simplification, we put the sound speed perturbation model as a linear function in space as
190 follows:

$$191 \quad \Gamma(t, \mathbf{P}, \mathbf{X}) \equiv \alpha_0(t) + \boldsymbol{\alpha}_1(t) \cdot \frac{\mathbf{P}}{L^*} + \boldsymbol{\alpha}_2(t) \cdot \frac{\mathbf{X}}{L^*} \quad (4)$$

192 where L^* indicates the characteristic length of the observation site (typically in several kilometers).
193 $\alpha_0(t)$, $\alpha_1(t)$ and $\alpha_2(t)$ are the time-dependent coefficients for each term. Because the vertical
194 variation of \mathbf{P} and \mathbf{X} are much smaller than the horizontal variation, we can practically ignore the
195 vertical component of $\alpha_1(t)$ and $\alpha_2(t)$. Thus, $\alpha_1(t)$ and $\alpha_2(t)$ are reduced to a 2-dimensional
196 vector to denote the horizontal gradient.

197 Each coefficient can be represented by a linear combination of basis functions $\Phi_k(t)$:

$$198 \quad \begin{cases} \alpha_0(t) = \sum_{k=0}^{K_a} a_k^{(0)} \Phi_k^{(0)}(t) \\ \alpha_1(t) = \sum_{k=0}^{K_b} (a_k^{(1E)} \Phi_k^{(1E)}(t), a_k^{(1N)} \Phi_k^{(1N)}(t), 0) \\ \alpha_2(t) = \sum_{k=0}^{K_c} (a_k^{(2E)} \Phi_k^{(2E)}(t), a_k^{(2N)} \Phi_k^{(2N)}(t), 0) \end{cases} \quad (5)$$

199 where $a_k^{(\cdot)}$ are the coefficients of the k th basis function, $\Phi_k^{(\cdot)}(t)$, for each term named $\langle \cdot \rangle$. E and N in
200 $\langle \cdot \rangle$ denote the eastward and northward components of the vector, respectively. For simplification, we
201 compile these coefficients into vector \mathbf{a} , hereafter.

202 Because the values for \mathbf{M} and \mathbf{X}_j are usually obtained in the precision of less than meters prior to the
203 GNSS-A analysis, \mathbf{P} and \mathbf{X}_j in Γ can be approximated with the prior, i.e., \mathbf{M}^0 and \mathbf{X}_j^0 . This reduces
204 the number of estimation parameters in the correction term, i.e., $\gamma_i = \gamma_i(\mathbf{a} | \mathbf{X}_j^0, \mathbf{M}^0, \mathbf{Q}(t), \boldsymbol{\theta}(t))$.

205 2.4 Rigid array constraints

206 Usually, the local deformation within the transponders' array is assumed to be sufficiently small, so
207 that the same array geometry parameters can be used throughout all visits. Because the relative
208 positions of the transponders are strongly coupled with the sound speed variable and positional
209 offsets, constraining the array geometry is expected to stabilize the GNSS-A solutions. Matsumoto et
210 al. (2008) developed the rigid-array constraint, which has been adopted in the subsequent studies
211 (e.g., Watanabe et al., 2014; Yokota et al., 2016) except in the cases where the rigid-array assumption
212 is inadequate (e.g., Sato et al., 2011).

213 To implement the rigid-array constraint, slight change in the observation equation is needed. We
214 divide the transponders' positions as $\mathbf{X}_j = \bar{\mathbf{X}}_j + \Delta \mathbf{X}_c$, where $\bar{\mathbf{X}}_j$ and $\Delta \mathbf{X}_c$ denote the relative positions
215 of each transponder for the arbitrary origin, and the parallel translation of the transponder array,
216 respectively. The array geometry, $\bar{\mathbf{X}}_j$, should be determined prior to the analytical procedure, using
217 the data of multiple observation visits.

218 Meanwhile, $\bar{\mathbf{X}}_j$ can also be determined simultaneously with the positioning procedure by combining
219 the data vectors, model parameter vectors, and observation equation for all series of the observation
220 visits, as the original formulation of Matsumoto et al. (2008). However, it requires huge
221 computational resources to solve all the parameters, as the number of observations increases.
222 Therefore, we are not concerned in this paper and code with the simultaneous determination of the
223 array geometry.

224 3 Analytical procedures

225 3.1 Observation equation

226 In the GNSS-A analysis, observed travel time, T_i^o , are compared with the model, T_i^c . In order to
227 expand the range of travel time from $(0, \infty)$ to $(-\infty, \infty)$, we took the logarithms of travel time.
228 Summarizing the above expansion, we put the following observation equation for i th acoustic round-
229 trip travel time:

$$230 \quad \log(T_i^o/T^*) = \log(\tau_i(\mathbf{X}_j, \mathbf{M}|\mathbf{Q}, \boldsymbol{\theta}, V_0)/T^*) - \gamma_i(\mathbf{a}|\mathbf{X}_j^0, \mathbf{M}^0, \mathbf{Q}, \boldsymbol{\theta}) + e_i \quad (6.1)$$

231 or in the form with the rigid-array constraint,

$$232 \quad \log(T_i^o/T^*) = \log(\tau_i(\Delta\mathbf{X}_c, \mathbf{M}|\bar{\mathbf{X}}_j, \mathbf{Q}, \boldsymbol{\theta}, V_0)/T^*) - \gamma_i(\mathbf{a}|\mathbf{X}_j^0, \mathbf{M}^0, \mathbf{Q}, \boldsymbol{\theta}) + e_i \quad (6.2)$$

233 where T^* is the characteristic travel time and e_i is the observation error vector. Figure 3 indicates the
234 summary for constructing the observation equation. It should be noted that, in this formulation, only
235 the continuity of sound speed field is assumed.

236 This section shows the algorithm to estimate the model parameters from the nonlinear observation
237 equation 6. We took a Bayesian approach because of its simple expression when incorporating prior
238 information. Furthermore, it provides a well-defined index for the model selection, i.e., the Akaike's
239 Bayesian Information Criterion (ABIC; Akaike, 1980). The expansion shown in this section is based
240 on Tarantola and Valette (1982) and Matsu'ura et al. (2007).

241 3.2 Prior information

242 The observation equation can be rewritten as,

$$243 \quad \mathbf{y} = \mathbf{f}(\mathbf{x}) + \mathbf{e} \quad (7)$$

244 where $\mathbf{x} = [\mathbf{X}_j^T \quad \mathbf{M}^T \quad \mathbf{a}^T]^T$, $y_i = \log(T_i^o/T^*)$, and $f_i = \log(\tau_i/T^*) - \gamma_i$. Let us consider the direct
245 prior information for the model parameters \mathbf{X}_j and \mathbf{M} written as,

$$246 \quad \begin{bmatrix} \mathbf{X}_j^0 \\ \mathbf{M}^0 \end{bmatrix} = \begin{bmatrix} \mathbf{X}_j \\ \mathbf{M} \end{bmatrix} + \begin{bmatrix} \mathbf{d}_X \\ \mathbf{d}_M \end{bmatrix} \quad (8)$$

247 where \mathbf{X}_j^0 , \mathbf{M}^0 and $\mathbf{d} = [\mathbf{d}_X^T \quad \mathbf{d}_M^T]^T$ denote the predicted model parameter vector and the error
248 vector, respectively. Let us assume that \mathbf{d}_X and \mathbf{d}_M follow a normal distribution with a variance-
249 covariance of $D_X(\rho^2)$ and $D_M(\rho^2)$, whose scale can be adjusted by a hyperparameter ρ^2 , i.e., $D_X =$
250 $\rho^2 \widetilde{D}_X$ and $D_M = \rho^2 \widetilde{D}_M$, respectively. The prior probability density function (pdf) for the constraints
251 can be written as,

$$252 \quad p(\mathbf{X}_j, \mathbf{M}; \rho^2) = c \cdot \exp \left[-\frac{1}{2} \left(\begin{bmatrix} \mathbf{X}_j^0 \\ \mathbf{M}^0 \end{bmatrix} - \begin{bmatrix} \mathbf{X}_j \\ \mathbf{M} \end{bmatrix} \right)^T \begin{bmatrix} D_X(\rho^2) & 0 \\ 0 & D_M(\rho^2) \end{bmatrix}^{-1} \left(\begin{bmatrix} \mathbf{X}_j^0 \\ \mathbf{M}^0 \end{bmatrix} - \begin{bmatrix} \mathbf{X}_j \\ \mathbf{M} \end{bmatrix} \right) \right] \quad (9)$$

253 where c denotes the normalization constant.

254 For the model parameter \mathbf{a} , an indirect prior information can be applied that the temporal change of
255 sound speed perturbation model Γ is small. Specifically, the roughness which can be defined by the
256 derivatives of each term in equation 4 should be small. In this study, we use the square of second

281 temporal correlation because the sound speed variation is modeled by a smooth function of space and
282 time. Thus, we assumed the following covariance terms:

$$283 \quad E_{ij} = \begin{cases} \sqrt{E_{ii}E_{jj}} \exp\left(-\frac{|t_i - t_j|}{\mu_t}\right) & \text{if the transponders for } i \text{ and } j \text{ are the same} \\ \mu_{MT} \sqrt{E_{ii}E_{jj}} \exp\left(-\frac{|t_i - t_j|}{\mu_t}\right) & \text{for others} \end{cases} \quad (14)$$

284 whose formulation refers to Fukahata and Wright (2008). Equation 14 means that the densely
285 sampled data would have smaller weights in the model than the isolated data. A factor $\mu_{MT} \in [0, 1]$
286 was introduced to suppress the error correlation between the different transponders because the
287 acoustic rays for different transponders take separate paths as the depths increases. Consideration of
288 the non-diagonal components of the data variance-covariance contributes to reduce the complexity of
289 the model against the excessively high-rate data sampling.

290 On the other hand, the diagonal component of E controls the weight of individual data. Because the
291 measurement errors of acoustic travel time are caused by mis-reading of the return signal, it is
292 independent on the travel time value. Therefore, we apply $E_{ii} = (T^*/T_i^o)^2$, so that all measured data,
293 T_i^o , has the same weight in the real scale.

294 3.4 Posterior probability

295 The posterior pdf after the data acquisition, which can be defined to be equal to the likelihood of the
296 model parameter given the data, can be written as,

$$297 \quad p(\mathbf{x}; \sigma^2, \mu_t, \mu_{MT}, \rho^2, \lambda^2 | \mathbf{y}) = c \cdot (2\pi\sigma^2)^{-\frac{(n+g)}{2}} |E|^{-\frac{1}{2}} \|\widetilde{\Lambda}_G\|^{\frac{1}{2}} \exp\left[-\frac{1}{2\sigma^2} s(\mathbf{x})\right] \quad (15.1)$$

298 with,

$$299 \quad s(\mathbf{x}) = (\mathbf{y} - \mathbf{f}(\mathbf{x}))^T E^{-1} (\mathbf{y} - \mathbf{f}(\mathbf{x})) + (\mathbf{x}^0 - \mathbf{x})^T \widetilde{G} (\mathbf{x}^0 - \mathbf{x}) \quad (15.2)$$

300 where $\widetilde{G} = \sigma^2 G(\rho^2, \lambda^2)$ and $\|\widetilde{\Lambda}_G\|$ represents the absolute value of the product of non-zero
301 eigenvalues of \widetilde{G} .

302 Defining $\widehat{\mathbf{x}}(\sigma^2, \mu_t, \mu_{MT}, \rho^2, \lambda^2)$ as \mathbf{x} that maximizes the posterior probability (equation 15) under the
303 given hyperparameters, the partial derivative of $p(\mathbf{x} | \mathbf{y})$ with respect to \mathbf{x} should be zero for $\mathbf{x} = \widehat{\mathbf{x}}$.
304 Hence, $\widehat{\mathbf{x}}$ should satisfy the following equation:

$$305 \quad A(\widehat{\mathbf{x}})^T E^{-1} (\mathbf{y} - \mathbf{f}(\widehat{\mathbf{x}})) + \widetilde{G} (\mathbf{x}^0 - \widehat{\mathbf{x}}) = \mathbf{0} \quad (16.1)$$

306 where $A(\mathbf{x})$ is the Jacobian matrix at point \mathbf{x} defined as,

$$307 \quad A(\mathbf{x}) = \begin{bmatrix} \frac{\partial f_1}{\partial x_{k1}}(\mathbf{x}) & \cdots & \frac{\partial f_1}{\partial x_{km}}(\mathbf{x}) \\ \vdots & \ddots & \vdots \\ \frac{\partial f_n}{\partial x_{k1}}(\mathbf{x}) & \cdots & \frac{\partial f_n}{\partial x_{km}}(\mathbf{x}) \end{bmatrix} \quad (16.2)$$

308 We can solve the nonlinear equation 16 numerically by performing an iterative method, where \mathbf{x}_k is
 309 corrected in each step with the following algorithm:

$$310 \quad \mathbf{x}_{k+1} = \mathbf{x}_k + (A(\mathbf{x}_k)^T E^{-1} A(\mathbf{x}_k) + \tilde{G})^{-1} (A(\mathbf{x}_k)^T E^{-1} (\mathbf{Y} - \mathbf{f}(\mathbf{x}_k)) + \tilde{G}(\mathbf{x}^0 - \mathbf{x}_k)) \quad (17)$$

311 to satisfy the following convergence criteria:

$$312 \quad A(\mathbf{x}_k)^T E^{-1} (\mathbf{Y} - \mathbf{f}(\mathbf{x}_k)) + \tilde{G}(\mathbf{x}^0 - \mathbf{x}_k) \ll 1 \quad (18)$$

313 Ignoring the term $O((\mathbf{x} - \hat{\mathbf{x}})^2)$ in $f(\mathbf{x})$ around $\hat{\mathbf{x}}$, $s(\mathbf{x})$ can be rewritten as,

$$314 \quad s(\mathbf{x}) \sim s(\hat{\mathbf{x}}) + (\mathbf{x} - \hat{\mathbf{x}})^T (A(\hat{\mathbf{x}})^T E^{-1} A(\hat{\mathbf{x}}) + \tilde{G})(\mathbf{x} - \hat{\mathbf{x}}) \quad (19)$$

315 Therefore, the linearized variance-covariance matrix around $\hat{\mathbf{x}}$ can be obtained as,

$$316 \quad \hat{C} = \sigma^2 (A(\hat{\mathbf{x}})^T E^{-1} A(\hat{\mathbf{x}}) + \tilde{G})^{-1} \quad (20)$$

317 3.5 Hyperparameter tuning

318 The appropriate values of the hyperparameters can be determined by minimizing Akaike's Bayesian
 319 Information Criteria (ABIC; Akaike, 1980),

$$320 \quad \text{ABIC} = -2 \log \int p(\mathbf{y}|\mathbf{x}; \sigma^2, \mu_t, \mu_{MT}) p(\mathbf{x}; \rho^2, \lambda^2) d\mathbf{x} + 2N_{\text{HP}} \quad (21)$$

321 where N_{HP} denotes the number of hyperparameters. Although it is difficult to analytically calculate
 322 the integral for the marginal likelihood because of the nonlinearity in $\mathbf{f}(\mathbf{x})$, the Laplace's method can
 323 be applied in this case where the degree of freedom is sufficiently large and $s(\mathbf{x})$ can be almost
 324 unimodal. Thus, an approximated form for ABIC is obtained as follows:

$$325 \quad \text{ABIC} \cong (n + g - m) \log s(\hat{\mathbf{x}}) - \log|E^{-1}| - \log\|\Lambda_G\| + \log|A(\hat{\mathbf{x}})^T E^{-1} A(\hat{\mathbf{x}}) + \tilde{G}| + \text{const.} \quad (22)$$

326 where m is the number of model parameters. For the derivation, we used the following relationship:

$$327 \quad \sigma^2 = \frac{s(\hat{\mathbf{x}})}{n + g - m} \quad (23)$$

328 which is derived from the condition that the partial derivative of ABIC with respect to σ^2 should be
 329 zero. We can tune the hyperparameters to minimize the approximated ABIC value defined in
 330 equation 22, to obtain the solution $\mathbf{x}^* = \hat{\mathbf{x}}(\sigma^{2*}, \mu_t^*, \mu_{MT}^*, \rho^{2*}, \lambda^{2*})$, where \cdot^* denotes the selected
 331 hyperparameters.

332 4 Features of "GARPOS"

333 GARPOS (Watanabe et al., 2020a; available at <https://doi.org/10.5281/zenodo.3992688>) has been
 334 developed to implement the GNSS-A analysis procedure. GARPOS is compatible with Python 3,
 335 with other packages NumPy, SciPy, pandas, and matplotlib. These packages are pre-installed in most
 336 of the Python distributions such as Anaconda. Sample scripts and data for testing GARPOS are also
 337 stored in the repository.

338 GARPOS is distributed as a series of files, which requires a driver script to run. The toolset consists
 339 of multiple Python files and a Fortran90 library for ray tracing. GARPOS requires the following
 340 input files:

341 (I-1) Initial site parameter file (in Python's configuration format),

342 (I-2) Acoustic observation data file (in csv format),

343 (I-3) Reference sound speed data file (in csv format),

344 (I-4) Setting file (in Python's configuration format).

345 Initial site parameter file (I-1) contains the initial values of the transponders' positions, the ATD
 346 offset and the relevant prior covariance information, as well as the metadata for the observation site
 347 and conditions. Acoustic observation data file (I-2) contains the list of the observation data associated
 348 with each acoustic ranging, such as travel time, positions, attitude and other metadata. Reference
 349 sound speed data file (I-3) contains the reference sound speed profile approximated into a polygonal
 350 curve. Setting file (I-4) contains the parameters to control the analysis procedures including the hyper
 351 parameters. Users can put the lists of candidates of hyperparameters in which the best combination
 352 may be within. The parameters $nmp0$, $nmp1$, and $nmp2$ in the setting file control the number of
 353 basis functions, K_a , K_b , and K_c in equation 5.

354 The results are written in the following output files:

355 (O-1) Estimated site parameter files (in Python's configuration format),

356 (O-2) Modified acoustic observation data file (in csv format),

357 (O-3) Model parameter list file (in csv format),

358 (O-4) Posterior variance-covariance matrix file (in csv format).

359 Estimated site parameter files (O-1) is written in the same format as the file (I-1). Modified acoustic
 360 observation data file (O-2) contains the calculated travel time data and the coefficients of sound
 361 speed perturbation model, as well as the original data/metadata set in (I-2). Model parameter list file
 362 (O-3) and posterior variance-covariance matrix file (O-4) contain the whole estimated model
 363 parameter vector and its variance-covariance, respectively.

364 Major input/output parameters and hyperparameters for GARPOS are listed in Tables 1 and 2,
 365 respectively.

366 We developed GARPOS to be compatible with both observation configurations. When handling the
 367 GNSS-A data collected in the stationary configurations, we should process data with some
 368 constraints on model parameters. Specifically, (1) upward components of transponders' positions
 369 should be fixed to zero, and (2) spatial gradient components of the sound speed perturbation model
 370 should not be solved, i.e., $nmp1 = nmp2 = 0$, because these parameters cannot be well resolved in
 371 the stationary configuration. Although further parameter tuning may be required for optimization,
 372 users can solve the seafloor position by GARPOS with the stationary data in addition to the move-
 373 around data.

374 5 Applications to the actual data

375 5.1 Data and settings

376 In order to verify the proposed analytical procedure, we reanalyzed the GNSS-A data at the sites
 377 named “TOS2” and “MYGI” (Table 3, Figure 4) in 2011-2019. The test sites were selected for
 378 several reasons: (1) whereas TOS2 is expected to move at almost constant rate, MYGI will show the
 379 transient displacement due to the postseismic crustal deformation of the 2011 Tohoku-oki
 380 earthquake; (2) the oceanographic environments are different, i.e., the effect of the Kuroshio current
 381 is dominant at TOS2; but (3) the depths of both sites are almost the same. The observation epochs
 382 used in this study is listed in Supplementary Tables 1 and 2. The datasets used in this study are
 383 available at <https://doi.org/10.5281/zenodo.3993912> (Watanabe et al., 2020b).

384 Acoustic round-trip travel times were measured on the survey vessel using the hull-mounted acoustic
 385 transducer (e.g., Ishikawa et al., 2020). Processing delays in the acoustic devices were subtracted
 386 from the acoustic data beforehand.

387 Solid-earth-tide-free positions of GNSS antenna $\mathbf{Q}(t)$ were determined at 2 Hz by the open source
 388 software RTKLIB version 2.4.2 (Takasu, 2013) in post-processing kinematic Precise Point
 389 Positioning (PPP) mode, using the precise satellite orbit and the 30-sec satellite clock solutions (final
 390 products) provided by the International GNSS Service (International GNSS Service, a; b), in the
 391 same procedures as Watanabe et al. (in press). The ATD offset values for each vessel, \mathbf{M} , were
 392 measured by leveling, distance, and angle surveys before the first GNSS-A observation cruise, to be
 393 used as \mathbf{M}^0 .

394 Along with the acoustic observations, the profiles of temperature and/or conductivity were measured
 395 by CTD, XCTD or XBT probes several times. The reference sound speed profile, $V_0(u)$, was
 396 calculated from the observed temperature and salinity profiles using the empirical relationship
 397 proposed by Del Grosso (1974). To save the computational cost for ray tracing, the profile was
 398 approximated into a polygonal curve with several tens of nodes (Figure 5).

399 During a GNSS-A survey, the vessel sails on a pre-determined track over the seafloor transponder
 400 array to collect geometrically balanced acoustic data (e.g., Figure 1). The along-track observation is
 401 repeated several times by reversing the sailing direction in order to reduce the bias due to the errors
 402 in the ATD offset. The along-track observation (called “subset”, hereafter) is repeated several times,
 403 with reversed sailing direction in order to reduce the bias due to the errors in the ATD offset.

404 During an observation cruise, it occasionally took more than a few weeks to collect sufficient
 405 acoustic data at a single site due to weather conditions or other operational restrictions. Even so, we
 406 compiled a single dataset per site per cruise for the static seafloor positioning in practice, because the
 407 positional changes should be too small to detect. We call the collection of a single GNSS-A dataset
 408 “observation epoch” or “epoch”, hereafter.

409 We set the parameters for the numbers of basis functions, K_a , K_b , and K_c , in equation 5, as $nmp0 =$
 410 $nmp1 = nmp2 = 15$ for both preprocess and main process.

411 5.2 Array geometry determination

412 In order to calculate the proper array geometry $\overline{\mathbf{X}}_j$ for the rigid-array constraint, we first determined
 413 the positions of each transponder for all observations. Note that not all transponders are used in each

414 observation, for example, because of additional installation of transponders for replacing
415 transponders which were decommissioned due to battery outage. $\bar{\mathbf{X}}_j$ and the positional difference of
416 the array center for n th observation, $\mathbf{c}^{(n)}$ were calculated by solving the following simultaneous
417 equations:

$$418 \quad \begin{cases} \mathbf{X}_j^{(n)} = \delta_j^{(n)} \bar{\mathbf{X}}_j + \delta_j^{(n)} \mathbf{c}^{(n)} & (\text{for } j = 1 \dots J \text{ and } n = 1 \dots N) \\ 0 = \sum_{n=1}^N \mathbf{c}^{(n)} \end{cases} \quad (24.1)$$

419 with,

$$420 \quad \delta_j^{(n)} = \begin{cases} 1 & \text{if the transponder } j \text{ is used in } n \text{ th observation} \\ 0 & \text{others} \end{cases} \quad (24.2)$$

421 where J and N are the number of transponders and observations, respectively, and $\mathbf{X}_j^{(n)}$ denotes the
422 predetermined transponders' positions for the n th observation.

423 The preliminary array-free positioning was also used for the verification of the collected data. We
424 eliminated the outliers whose discrepancies from the preliminary solution were larger than the
425 arbitrary threshold. We set the threshold to be 5 times as large as the root mean square value (RMS)
426 of the travel time residuals.

427 5.3 Hyperparameter search

428 In order to get the solution \mathbf{x}^* , we should determine the appropriate values for the various
429 hyperparameters, i.e., σ^2 , μ_t , μ_{MT} , ρ^2 , λ_0^2 , λ_{1E}^2 , λ_{1N}^2 , λ_{2E}^2 , and λ_{2N}^2 . In the scheme of the ABIC
430 minimization, σ^2 can be determined analytically by equation 23. It is reasonable to assume $\lambda_g^2 \equiv$
431 $\lambda_{1E}^2 = \lambda_{1N}^2 = \lambda_{2E}^2 = \lambda_{2N}^2$ because these hyperparameters control the smoothness of the spatial
432 sound speed structure. For the purpose of single positioning, ρ should be a large number, for example
433 in meter-order. The large ρ hardly changes the ABIC value and thus the solution.

434 In order to save the computational resources, we should further reduce the number of
435 hyperparameters. We tentatively put $\mu_m = 0.5$. For the sound speed variations, we had to assume the
436 strong constancy of spatial sound speed structure to resolve them with the single transducer GNSS-A.
437 For this reason, we selected the ratio of λ_0^2 and λ_g^2 , as $\lambda_g^2 = 0.1 \lambda_0^2$. The last two hyperparameters,
438 μ_t and λ_0^2 , were determined with the grid search method. The tested values for μ_t and λ_0^2 are $\mu_t =$
439 (0 min., 0.5 min., 1 min., 2 min., 3 min.) and $\lambda_0^2 = (10^{-3}, 10^{-2}, 10^{-1}, 10^0, 10^1, 10^2)$,
440 respectively.

441 5.4 Results

442 Figure 6 shows the time series of the estimated positions at sites TOS2 and MYGI. The positions are
443 aligned to the ITRF 2014 (Altamimi et al., 2016) and transformed into local ENU coordinates.
444 Comparing the time series derived by the existing scheme (SGOBS version 4.0.2; used in Yokota et
445 al., 2019), GARPOS reproduced almost the same trends for both sites.

446 TOS2 is located offshore in the south of Shikoku Island, southwestern Japan, above the source region
447 of the 1946 Nankaido earthquake (e.g., Sagiya and Thatcher, 1999) along the Nankai Trough.

448 According to Yokota and Ishikawa (2020), who investigated the transient deformations at the GNSS-
 449 A sites along the Nankai Trough, no significant signal was detected at TOS2. The results by the
 450 proposed method show the same trends as the conventional results. Although the trend of horizontal
 451 displacement seems to be changed in 2018 or 2019, careful inspection is needed because the
 452 transponders had been replaced during this period.

453 MYGI is located in the offshore east of Miyagi Prefecture, northeastern Japan, which experienced the
 454 2011 Tohoku-oki earthquake (Sato et al., 2011). After the earthquake, significant westward
 455 postseismic movement and subsidence due to the viscoelastic relaxation has been observed at MYGI
 456 (Watanabe et al., 2014). The postseismic movements continue but appear to decay. It is true that the
 457 changes in the displacement rate at these sites are crucial in seismic and geodetic researches, but
 458 discussing these matters is beyond the scope of the present paper. The point is that the seafloor
 459 positioning results were well reproduced by the proposed method.

460 **6 Discussions**

461 **6.1 Interpretations for the correction coefficient**

462 As mentioned in Section 2.3, it is convenient to relate the correction coefficient to the sound speed
 463 perturbation by assuming the case for $|\gamma_i| \ll 1$ for better understanding, though observation equation
 464 6 is valid for arbitrary value of γ_i . For the relationship $\delta V_i \sim \gamma_i \bar{V}_0$, we can convert each term of Γ into
 465 the dimension of speed and speed gradient as, $\delta V_0(t) \equiv \bar{V}_0 \alpha_0(t)$, $\mathbf{g}_1(t) \equiv \bar{V}_0 \alpha_1(t)$, and $\mathbf{g}_2(t) \equiv$
 466 $\bar{V}_0 \alpha_2(t)$.

467 The early models by Fujita et al. (2006) and Ikuta et al. (2008) took only the term $\delta V_0(t)$ into
 468 account. Whereas Ikuta et al. (2008) used the cubic B-spline functions as basis functions, Fujita et al.
 469 (2006) applied the multiple 2nd degree polynomial functions with 10-20-minute time windows.
 470 Although these models do not include any transponder dependent term $\mathbf{g}_2(t)$, the transponder
 471 independent spatial gradient $\mathbf{g}_1(t)$ can be indirectly extracted as shown by Yokota et al. (2019).

472 In addition to estimating the term identical to δV_0 , Yokota et al. (2019) implemented the additional
 473 process to estimate \mathbf{g}_2 from the residuals of the solution by the method of Fujita et al. (2006).
 474 Strictly, the derived parameters in their scheme, i.e., ΔV_1 and ΔV_2 in Yokota et al. (2019), are the
 475 same as $\mathbf{g}_1 + \mathbf{g}_2$ and \mathbf{g}_2 in this study, respectively. For these parameters, our team have already
 476 made a qualitative interpretation in Yokota and Ishikawa (2019).

477 In order to show the relationship with other conventional models, we expand the proposed
 478 formulation to those by Honsho et al. (2019), Yasuda et al. (2017) and Kinugasa et al. (2020).
 479 Because Honsho et al. (2019) practically assumed 1-dimensional sound speed gradient, they
 480 constructed the model basically in the 2-dimensional plane spanned by the gradient direction and
 481 vertical direction.

482 For simplification, we assume that the ray path is a straight line connecting both ends. Putting L^*
 483 equal to the depth of the observation site, the emission angle θ defined in Figure 3 of Honsho et al.
 484 (2019) can be expressed as,

$$485 \quad \frac{X_j}{L^*} - \frac{P(t)}{L^*} = \tan \theta \quad (25)$$

486 Furthermore, assuming that the transmit/reception positions are the same and that the difference
487 between transmit/reception time is so small that $\alpha_0(t)$, $\alpha_1(t)$ and $\alpha_2(t)$ hardly change, γ_i can be
488 written as,

$$489 \quad \gamma_i = \alpha_0(t) + (\alpha_1(t) + \alpha_2(t)) \frac{P(t)}{L^*} + \alpha_2(t) \tan \theta \quad (26)$$

490 Because δT defined in equations 2 and 5 of Honsho et al. (2019) is equivalent to $T_i^c - \tau_i$ in our
491 formulation, we have,

$$492 \quad (\exp(-\gamma_i) - 1)\tau_i = \frac{1}{\cos \theta} (c_0(t) + g(t)x_0 + w(t) \tan \theta) \quad (27)$$

493 where $c_0(t)$, $g(t)$, $w(t)$ and $x_0 = P$ are defined in equations 6, 7, 8 of Honsho et al. (2019) and the
494 transducer's position in their formulation, respectively. Recalling that the slant range of acoustic ray
495 path is $2L^*/\cos \theta$, the reference round trip travel time can be written as,

$$496 \quad \tau_i = \frac{2L^*}{V_0(u) \cos \theta} \quad (28)$$

497 Considering the case where $|\gamma_i| \ll 1$, equation 27 is approximated to,

$$498 \quad -\frac{2L^*}{V_0} \gamma_i = c_0(t) + g(t)x_0 + w(t) \tan \theta \quad (29)$$

499 From equations 26 and 29, the following relationships are derived:

$$500 \quad c_0(t) = -\frac{2L^*}{V_0} \alpha_0(t) \quad (30.1)$$

$$501 \quad g(t) = -\frac{2}{V_0} (\alpha_1(t) + \alpha_2(t)) \quad (30.2)$$

$$502 \quad w(t) = -\frac{2L^*}{V_0} \alpha_2(t) \quad (30.3)$$

503 In Honsho et al. (2019), $w(t)$ is extended to a 2-dimensional vector, i.e.,

$$504 \quad \mathbf{w}(t) = -\frac{2L^*}{V_0} \boldsymbol{\alpha}_2(t) \quad (31.1)$$

505 Similarly, when extending $g(t)$ to the 2-dimensional vector, we can use the following vector form:

$$506 \quad \mathbf{g}(t) = -\frac{2}{V_0} (\boldsymbol{\alpha}_1(t) + \boldsymbol{\alpha}_2(t)) \quad (31.2)$$

507 though they consequently use the assumption that $\mathbf{g}(t)$ is parallel to $\mathbf{w}(t)$. It is equivalent to the case
508 that $\boldsymbol{\alpha}_1$ is parallel to $\boldsymbol{\alpha}_2$ in the proposed formulation.

509 Honscho et al. (2019) supposed the physical model where a spatially homogeneous 1-dimensional
 510 gradient of slowness lies in the certain layer, from sea-surface to the depth D , in the water. In such
 511 cases, $\mathbf{w}(t)$ is proportional to $\mathbf{g}(t)$, as $\mathbf{w} = (D/2)\mathbf{g}$. This is exactly the same assumption as the
 512 model by Yasuda et al. (2017). The model of Kinugasa et al. (2020) is the special case of those
 513 models where D equals to the water depth.

514 In the proposed method, the sound speed field is approximately interpreted by their models when the
 515 unit vector of $\boldsymbol{\alpha}_1$ is supposed to be same as that of $\boldsymbol{\alpha}_2$ and $|\boldsymbol{\alpha}_1| \geq |\boldsymbol{\alpha}_2|$. The depth of the gradient
 516 layer is calculated as,

$$517 \quad D = \frac{2L^*}{1 + \boldsymbol{\alpha}_1/\boldsymbol{\alpha}_2} \quad (32)$$

518 When $\boldsymbol{\alpha}_1 = \boldsymbol{\alpha}_2$, it concludes to the model of Kinugasa et al. (2020). Conversely, when $|\boldsymbol{\alpha}_2| \ll |\boldsymbol{\alpha}_1|$,
 519 sound speed gradient lies in the thin layer near the surface.

520 In addition to the simple model above, the proposed method can extract more complicated sound
 521 speed field, which partly described by Yokota and Ishikawa (2019). Extracted parameters for the
 522 sound speed perturbation indicate the complicity of oceanographic structure, as shown in the next
 523 section.

524 **6.2 Validity of extracted sound speed perturbation model**

525 Typical examples for the estimation results for each observation, i.e., the time series of travel time
 526 residuals, and sound speed perturbation interpreted from the correction coefficient, are shown in
 527 Figure 7. Results for all the datasets are available in Supplementary Figure 1.

528 In the most cases for site TOS2, both terms of the estimated sound speed gradient vector stably direct
 529 south to southeast. Because the sound speed increase with the water temperature, it means that the
 530 water temperature is higher in the southern region. The results that \mathbf{g}_2 is comparable with \mathbf{g}_1 in many
 531 cases indicate that the gradient of water temperature continues to the deeper portion, as discussed in
 532 the previous section. This is consistent with the fact that the Kuroshio current continuously flows on
 533 the south of TOS2.

534 In contrast, the directions of gradient terms at MYGI have less constancy than TOS2. Unlike the area
 535 around TOS2 where the Kuroshio current dominantly affects the seawater structure, MYGI is located
 536 in an area with a complicated ocean current system (e.g., Yasuda, 2003; Miyazawa et al., 2009).
 537 Watanabe and Uchida (2016) have also shown that the temperature profiles at MYGI vary widely
 538 with observation epochs. These features cannot be resolved by the simpler model with single sound
 539 speed gradient parameter.

540 The complexity in sound speed variation at MYGI tends to cause overfitting due to large variations in
 541 the residual travel time. Nevertheless, the proposed method successfully extracted the smooth sound
 542 speed structure for many observation epochs, except a few epochs such as June 2013
 543 (MYGI.1306.kaiyo_k4) and June 2019 (MYGI.1906.meiyo_m5). In these epochs, relatively larger
 544 values for the hyperparameter λ_0^2 were adopted. Possible causes of this include the systematic errors
 545 in other observation subcomponents such as the random walk noise in GNSS positioning, the drifts of
 546 gyro sensor, or the time synchronization error between the devices.

547 Preferred models for all the tested epochs had positive values for data correlation length, μ_t . It
 548 contributed to avoiding overfitting of the correction coefficient γ_i . It is considered that the plausible
 549 estimation of sound speed is realized by introducing the statistic information criteria and the
 550 information of data covariances.

551 Figure 8 shows the examples of the cases for the models without assuming the data correlation, i.e.,
 552 $\mu_t = 0$. The preferred models were selected from $\lambda_0^2 = (10^{-3}, 10^{-2}, 10^{-1}, 10^0, 10^1, 10^2, 10^3, 10^4)$.
 553 It is clear that the preferred models without assuming the data correlation have larger λ_0^2 . Although
 554 the residuals of travel time were reduced in these models, overfittings occurred for each term of Γ .
 555 Comparing the preferred and less-preferred results, the existence of data covariance components
 556 contributes to the selection of a model with less perturbation by decreasing the impact of individual
 557 data on model parameters.

558 Finally, we confirm the stability of the seafloor positioning results. The differences of seafloor
 559 position for the tested models from the most preferred models are summarized in Figure 9. The
 560 differences in estimated positions for most of the tested models converged into several centimeters.
 561 For both sites, variations in vertical component tend to be larger for larger values of λ_0^2 . It indicates
 562 that finer hyperparameter tuning is not required when considering the application to seafloor
 563 positioning.

564 7 Conclusions

565 We reconstructed the GNSS-A observation equation and developed the Python-based software
 566 GARPOS to solve the seafloor position as well as the sound speed perturbations using the empirical
 567 Bayes approach. It provides a stable solution for a generally ill-posed problem caused by the
 568 correlation among the model parameters, by introducing the hyperparameter tuning based on the
 569 ABIC minimization and data covariance to rationalize the normalization constant of the posterior pdf.

570 The most important point is that the proposed method succeeded in directly extracting the time-
 571 dependent sound speed field with two end members of spatial gradient terms, which are roughly
 572 characterized by depths, even when the observers used only one sea-surface unit. Statistical approach
 573 allowed us to suppress the overfitting and thus to obtain simpler sound speed field from densely
 574 collected dataset. It successfully reproduced the stationary southward sound speed gradient at TOS2,
 575 which is consistent with the Kuroshio current.

576 On the other hand, model overfits were shown in several epochs. These overfits can be caused not
 577 only by the actually complicated sound speed field but also by other error sources which were not
 578 well included in the model. It means that the hyperparameter tuning also plays a role in the
 579 verification of dataset and model. Error analyses in such cases might rather help improving the
 580 GNSS-A accuracy and methodology.

581 We suggested a simplified formatting for the GARPOS input files. Researchers can enter into the
 582 field of seafloor geodesy by collecting the listed data with adequate precision. Since each
 583 subcomponent of GNSS-A technique, i.e., GNSS positioning, acoustic ranging, and so on, has been
 584 well established, observers can combine them on their platform. Especially, GNSS-A is expected to
 585 be practicalized in the near future with an unmanned surface vehicle (Chadwell, 2016) or a buoy
 586 (e.g., Tadokoro et al., 2020; Kinugasa et al., 2020). Even in the case of the stationary observation due
 587 to small cruising speed, GARPOS may provide the solutions by making a slight modification in the
 588 prior variance-covariance matrix.

589 There is a room for improvement in setting the prior information for transponders' positions, X_j^0 . For
 590 instance, the displacement of transponder array from the previous epoch is predicted as small as
 591 several centimeters when the interval of observation visits is short. Such assumption leads to the
 592 application of the inter-epoch filtering. Furthermore, it has a possibility to progress to the kinematic
 593 seafloor positioning, as shown by Tomita et al. (2019). We expect that the publication of GARPOS
 594 on the open-access repository will enhance the researchers' engagement and the future development
 595 on the GNSS-A technique.

596 **8 Conflict of Interest**

597 The authors declare no conflict of interest.

598 **9 Data Availability**

599 The datasets analyzed in this study can be found in an open access repository at
 600 <https://doi.org/10.5281/zenodo.3993912> (Watanabe et al., 2020b). The code developed in this study
 601 is available at <http://doi.org/10.5281/zenodo.3992688> (Watanabe et al., 2020a)

602 **10 Author Contributions**

603 SW designed the study and wrote the manuscript. SW developed “GARPOS” and processed the data.
 604 SW, TI, YY, and YN discussed about the methodology and commented to improving the manuscript.

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607 **12 Abbreviations**

608 ABIC, Akaike Bayesian Information Criterion; ATD offset, Antenna – Transducer offset; GNSS,
 609 Global Navigation Satellite System; GNSS-A, Global Navigation Satellite System – Acoustic
 610 Ranging combined technique; pdf, probability density function.

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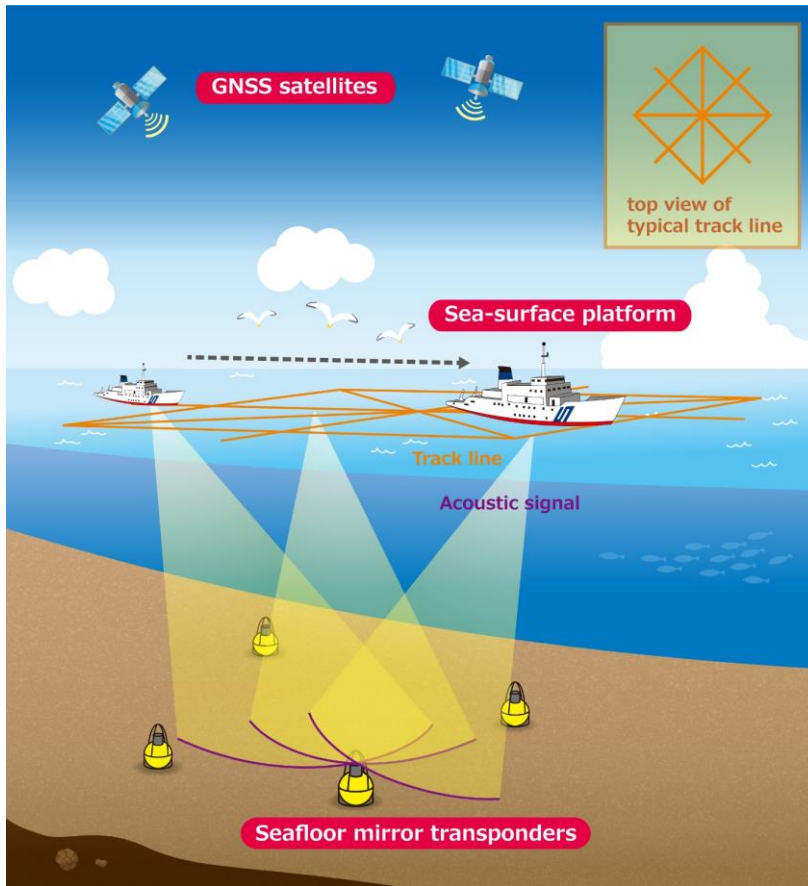
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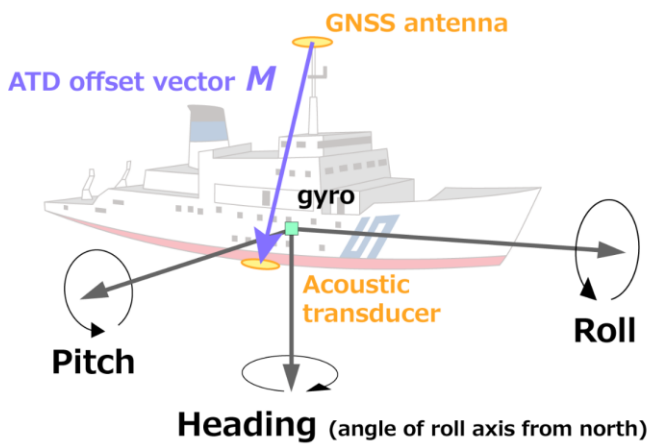
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771 15 Figures



772
773 Figure 1. Schematic image of the GNSS-A system in the move-around configuration.

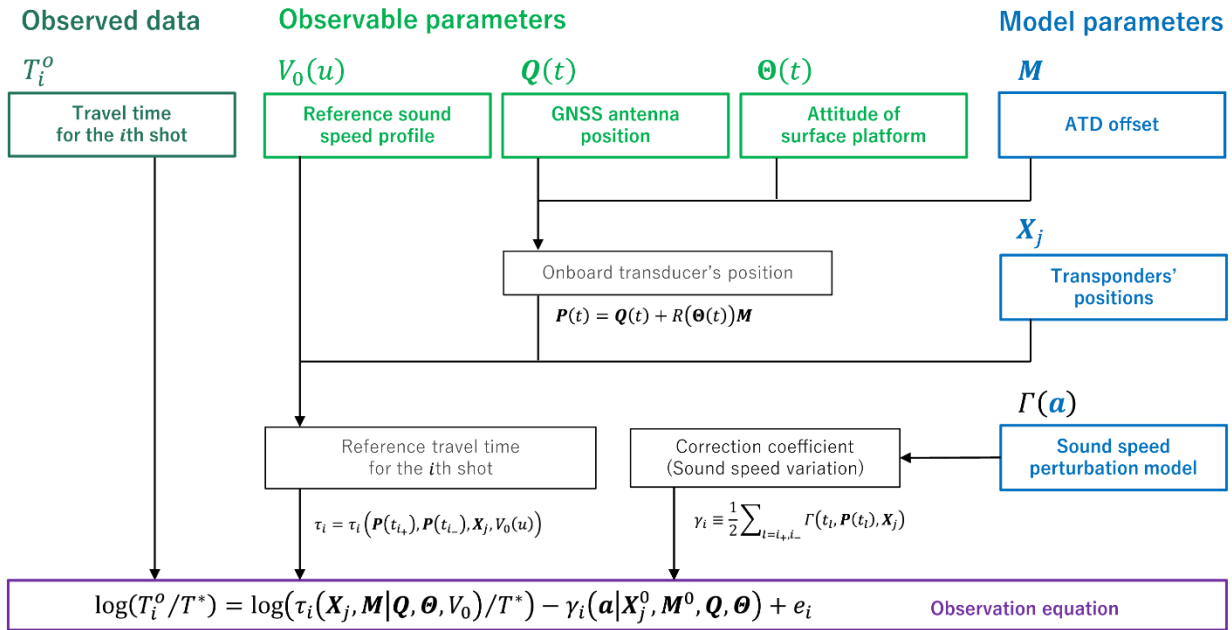
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776 Figure 2. Definitions of the attitude parameters and the ATD offset vector for the sea-surface
777 platform. Heading is zero when the roll axis directs to the north. The roll and pitch axes direct
778 forward and rightward (portside) of the vessel, respectively.

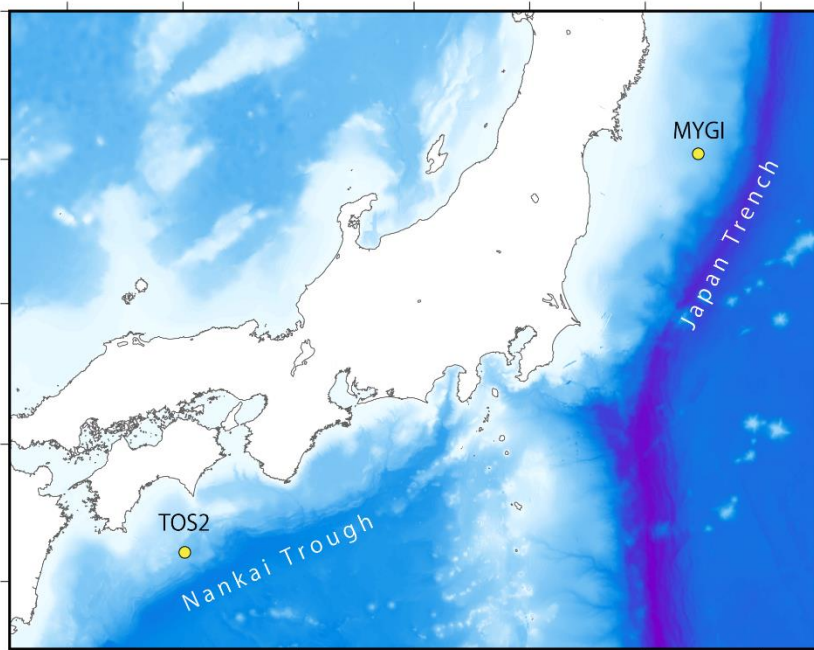
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782 Figure 3. Flow chart to construct the observation equation.

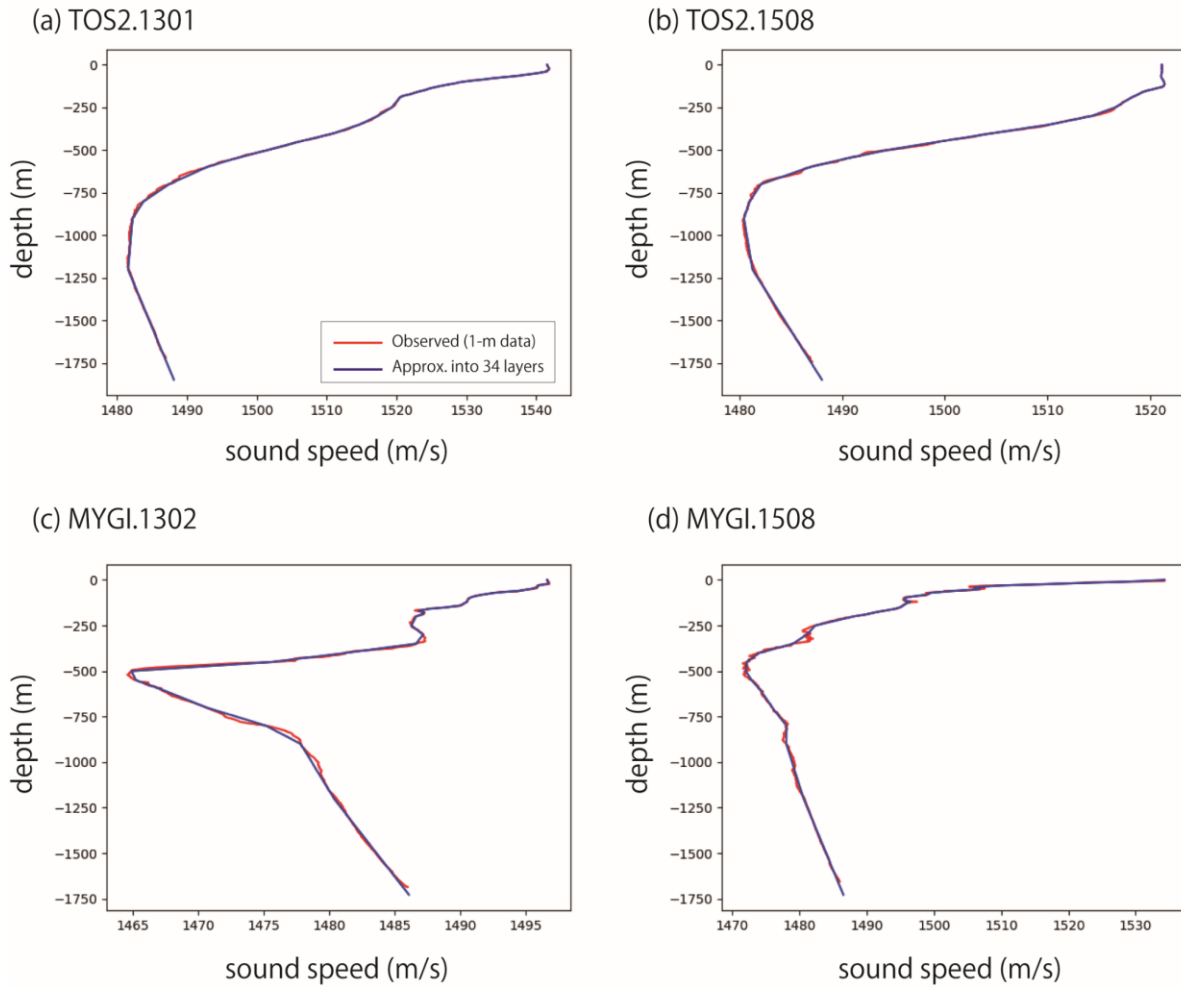


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784 Figure 4. Locations of the tested GNSS-A sites TOS2 and MYGI.

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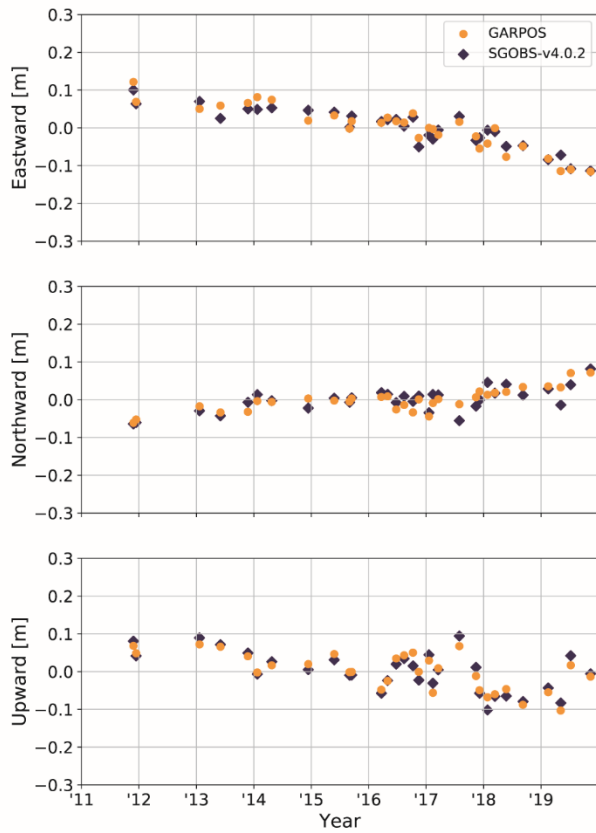
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788 Figure 5. Reference sound speed profiles (blue lines) for epochs (a) TOS2.1301 (Jan. 2013), (b)
 789 TOS2.1508 (Aug. 2015), (c) MYGI.1302 (Feb. 2013), and (d) MYGI.1508 (Aug. 2015). Red
 790 lines indicate 1-m sound speed profiles obtained from the 1-m layered XBT/XCTD data.

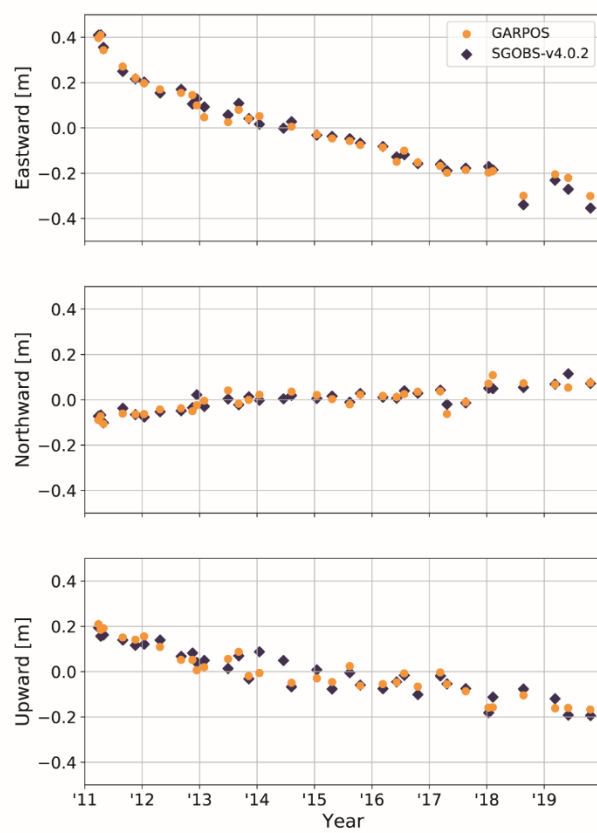
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(a) TOS2



(b) MYGI

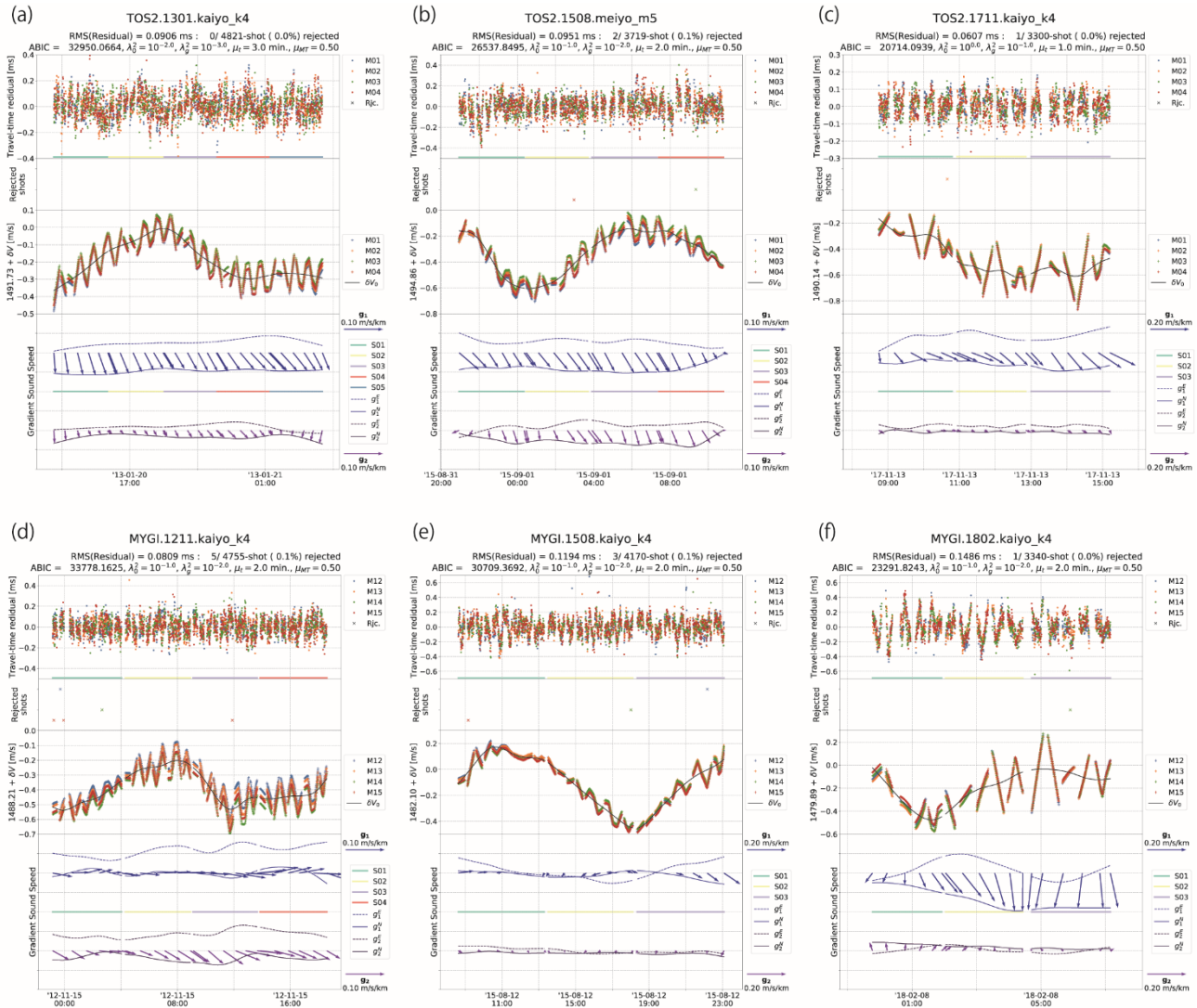


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794 Figure 6. Time series of displacement at (a) TOS2 and (b) MYGI solved by GARPOS (orange
795 circles) and SGOBS version 4.0.2 (blue squares). The positions are aligned to the ITRF 2014.

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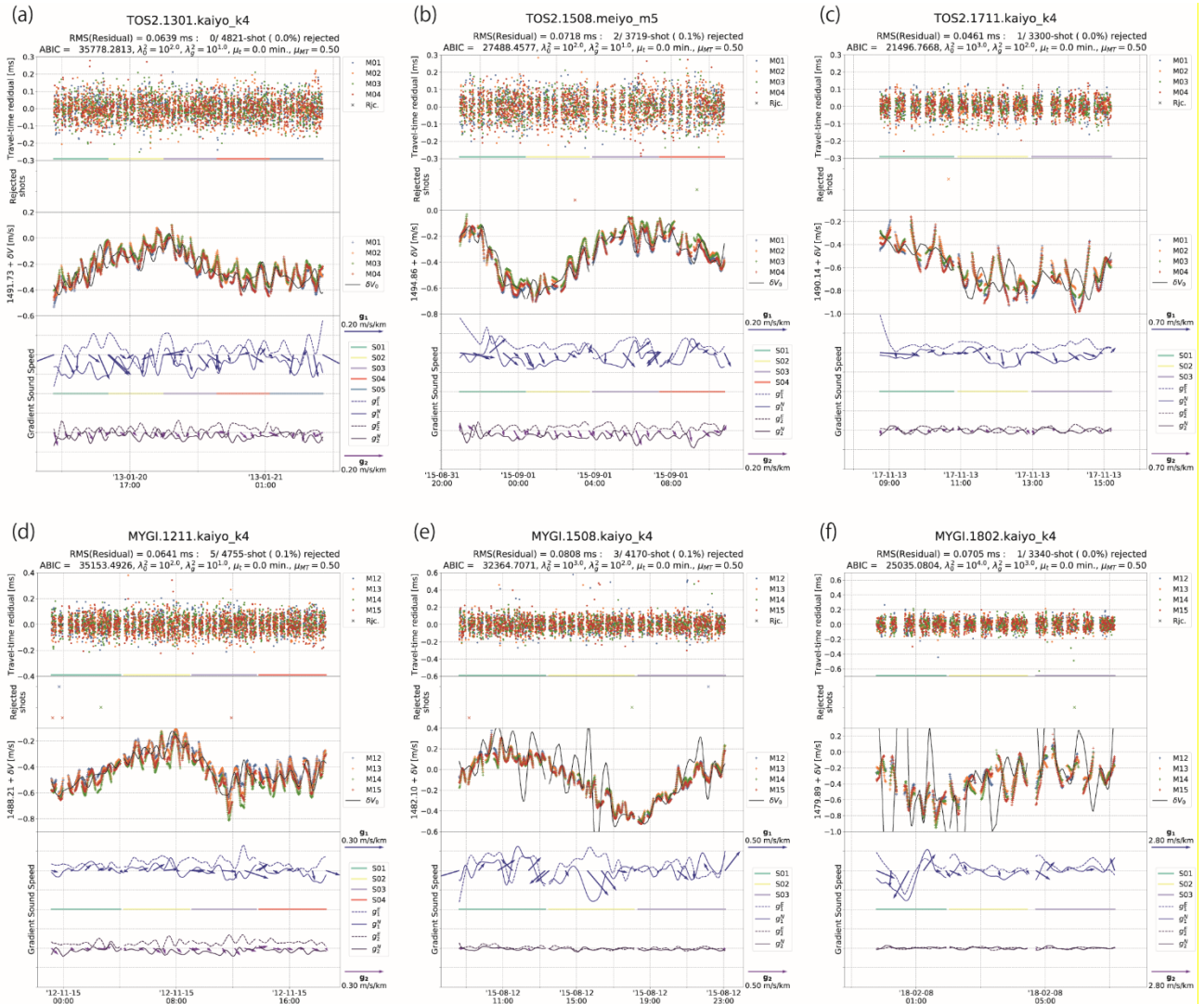


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799 Figure 7. Estimated results of the most preferred model for epochs (a) TOS2.1301.kaiyo_k4, (b)
 800 TOS2.1508.meiyo_m5, and (c) TOS2.1711.kaiyo_k4 (d) MYGI.1211.kaiyo_k4, (e)
 801 MYGI.1508.kaiyo_k4, and (f) MYGI.1802.kaiyo_k4. The top panels show the model residuals
 802 of the round-trip travel time. The second panels show the rejected acoustic data in the
 803 preprocessing step for determining the array geometry. The third panels indicate the sound speed
 804 perturbations, i.e., $\gamma_i \bar{V}_0$ (the crosses), and $\delta V_0(t) \equiv \bar{V}_0 \alpha_0(t)$ (black line). The colours of the
 805 symbols in these panels identify the target transponders. The blue and purple arrows on the
 806 bottom panels indicate the spatial gradient of the sound speed perturbations in north-up
 807 expression, i.e., $\mathbf{g}_1(t) \equiv \bar{V}_0 \alpha_1(t)$, and $\mathbf{g}_2(t) \equiv \bar{V}_0 \alpha_2(t)$, respectively. Dotted lines and solid
 808 lines show the temporal variations of eastward and northward components, respectively. The
 809 colored horizontal lines denote the ranges of the observation subsets.

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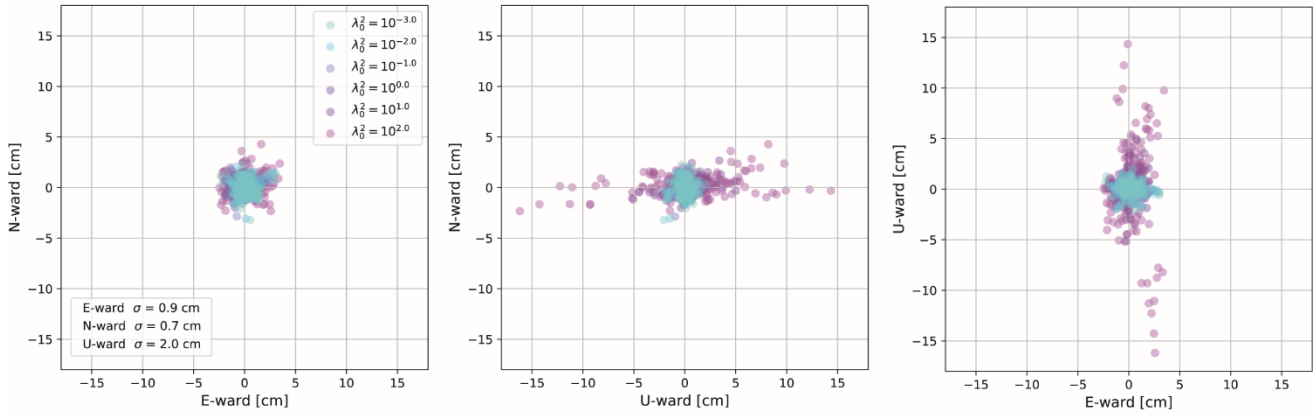
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Figure 8. Same as Figure 7, but for the most preferred model in the models with $\mu_t = 0$.

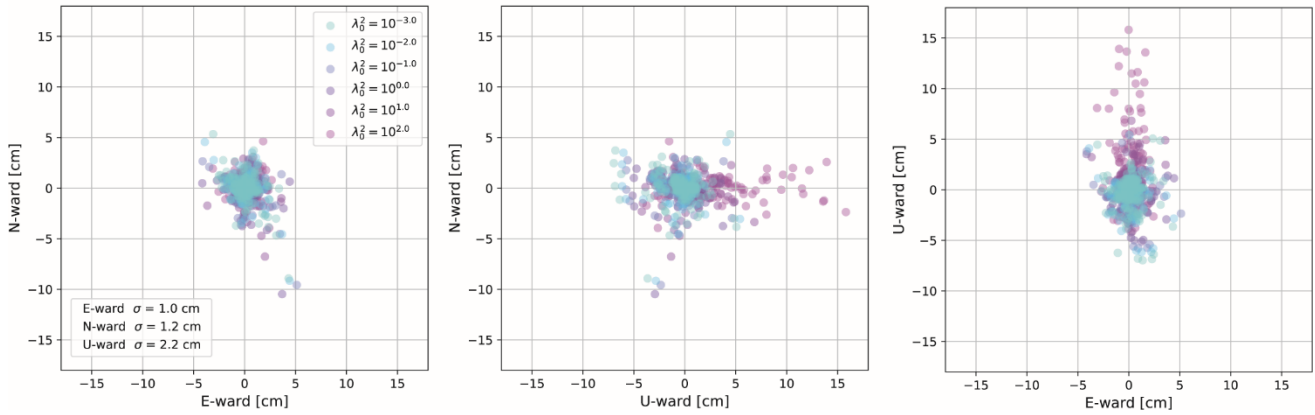
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(a) TOS2



(b) MYGI



816

817 Figure 9. Distributions of differences of positions of the tested models from the preferred ones at (a)
 818 TOS2 and (b) MYGI for northward-eastward (left), northward-upward (center), and upward-
 819 eastward (right) components. The colours of circles indicate the value of λ_0^2 .

820

821 16 Tables

822 Table 1. List of observable and estimation parameters used in GARPOS.

Parameter	Description	Name in I/O file	I/O file	type	unit
t_{i+}	transmit time	<i>ST</i>	I-2	obs	s
t_{i-}	reception time	<i>RT</i>	I-2	obs	s
$Q(t_{i+})$	Position of GNSS antenna at t_{i+} in ENU coordinates	<i>ant_e0</i> <i>ant_n0</i> <i>ant_u0</i>	I-2	obs	m
$Q(t_{i-})$	Position of GNSS antenna at t_{i-} in ENU coordinates	<i>ant_e1</i> <i>ant_n1</i> <i>ant_u1</i>	I-2	obs	m
$\Theta(t_{i+})$	Attitude of platform at t_{i+}	<i>roll0</i> <i>pitch0</i> <i>head0</i>	I-2	obs	deg.
$\Theta(t_{i-})$	Attitude of platform at t_{i-}	<i>roll1</i> <i>pitch1</i> <i>head1</i>	I-2	obs	deg.
γ_i	Correction coefficient	<i>gamma</i>	O-2	est	-
M^0	Prior ATD offset	<i>ATDoffset</i>	I-1	obs	m
X_j^0	Prior position of transponder	<i>M{j}_dPos</i>	I-1	obs	m
ΔX_c^0	Prior offset of transponder array	<i>dCentPos</i>	I-1	obs	m
\hat{M}	Posterior ATD offset	<i>ATDoffset</i>	O-1	est	m
\hat{X}_j	Posterior position of transponder	<i>M{j}_dPos</i>	O-1	est	m
$\hat{\Delta X}_c$	Posterior offset of transponder array	<i>dCentPos</i>	O-1	est	m
$V_0(u)$	Reference sound speed profile	CSV table	I-3	obs	m/s
K_a	Number of internal knots for α_0	<i>nmp0</i>	I-4	setting	-
K_b	Number of internal knots for α_1	<i>nmp1</i>	I-4	setting	-
K_c	Number of internal knots for α_2	<i>nmp2</i>	I-4	setting	-

823 * Note that $K_{\begin{Bmatrix} a \\ b \\ c \end{Bmatrix}} = nmp \begin{Bmatrix} 0 \\ 1 \\ 2 \end{Bmatrix} \times (\text{number of subset})$ in GARPOS.

824

825 Table 2. List of hyperparameter in GARPOS.

Hyper-parameter	Description	Formulation set in (I-4)	Name in Setting file	unit
μ_t	Correlation length of data	μ_t	<i>mu_t</i>	min.
μ_{MT}	Data correlation coefficient b/w the different transponders	μ_{MT}	<i>mu_mt</i>	-
λ_0^2	Smoothness parameter for α_0	$\log_{10} \lambda_0^2$	<i>Log_Lambda0</i>	-
λ_{1E}^2	Smoothness parameter for α_{1E}	$\log_{10} \left(\frac{\lambda_{(\cdot)}^2}{\lambda_0^2} \right)$	<i>Log_gradLambda</i>	-
λ_{1N}^2	Smoothness parameter for α_{1N}			-
λ_{2E}^2	Smoothness parameter for α_{2E}			-
λ_{2N}^2	Smoothness parameter for α_{2N}			-
σ^2	Scale of measurement error	N/A	N/A	-
ρ^2	Scale of a priori positioning error	N/A	N/A	m ²

826 * Note that σ^2 is calculated analytically, and that ρ^2 is set in (I-2).

827

828 Table 3. Locations and observation periods of the GNSS-A observation sites used in this study.

Site	Latitude	Longitude	Height	Number of epochs	Observation period
TOS2	32.43 °N	134.03 °E	-1740 m	31	2011.904 – 2019.863
MYGI	38.03 °N	142.92 °E	-1640 m	33	2011.238 – 2019.803

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