

CONSIDERING THE ROLE OF ADAPTIVE EVOLUTION IN MODELS OF THE OCEAN AND CLIMATE SYSTEM

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1 ABSTRACT

2 Numerical models have been highly successful in simulating global carbon and nutrient cycles in today's
3 ocean, together with observed spatial and temporal patterns of chlorophyll and plankton biomass at the
4 surface. With this success has come some confidence in projecting the century-scale response to continuing
5 anthropogenic warming. There is also increasing interest in using such models to understand the role of
6 plankton ecosystems in past oceans. However, today's marine environment is the product of billions of years
7 of continual evolution – a process that continues today. In this paper, we address the questions of whether an
8 assumption of species invariance is sufficient, and if not, under what circumstances current model projections
9 might break down. To do this, we first identify the key time-scales and questions asked of models. We then
10 review how current marine ecosystem models work and what alternative approaches are available to account
11 for evolution. We argue that for timescales of climate change overlapping with evolutionary timescales,
12 accounting for evolution may lead to very different projected outcomes regarding the timescales of
13 ecosystem response and associated global biogeochemical cycling. This is particularly the case for past
14 extinction events, but may also be true in the future, depending on the eventual degree of anthropogenic
15 disruption. The discipline of building new numerical models that incorporate evolution is also hugely
16 beneficial in itself, as it forces us to question what we know about adaptive evolution, irrespective of its
17 quantitative role in any specific event or environmental changes.

“It is a recognized principle of ecology that the interactions of organisms and environment are reciprocal. The environment not only determines the conditions under which life exists, but the organisms influence the conditions prevailing in the environment.”

Redfield (1958)

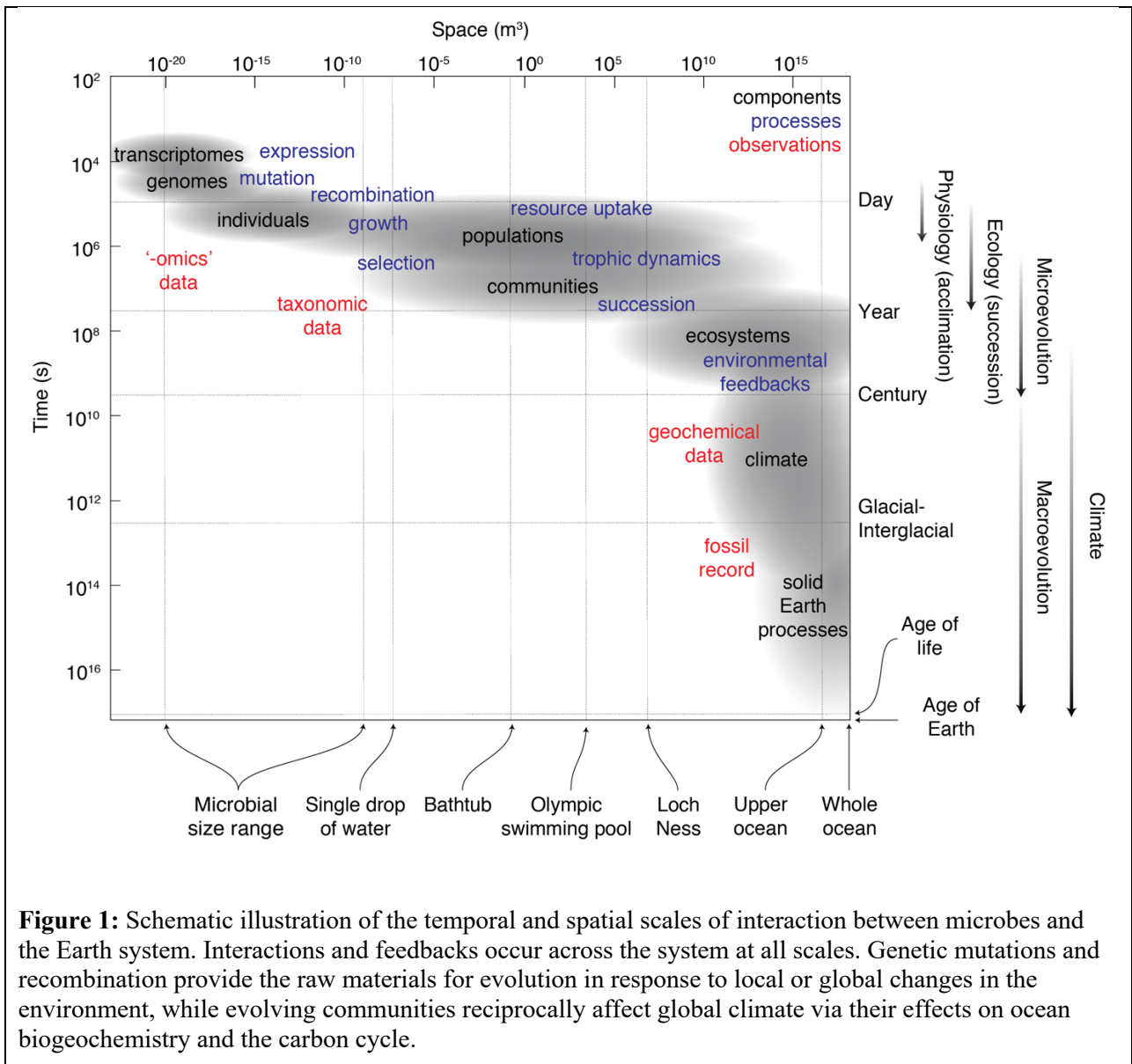
“At every moment natural selection is operating to change the genetic composition of populations in response to the momentary environment, but as that composition changes it forces a concomitant change in the environment itself. Thus organisms and environments are both causes and effects in a coevolutionary process.”

Lewontin (2000)

18 1 INTRODUCTION

19 A particular challenge in Earth system science is to understand and successfully simulate a system in which
20 higher level processes emerge from the interactions of many lower level processes (Levin 1998). Climate is
21 modulated by the global carbon cycle, which is itself driven by the metabolic activity of innumerable
22 interacting organisms. Within the marine realm, more than 10^{27} phototrophic organisms (Flombaum et al.
23 2013) contribute approximately half of total global photosynthesis (Field et al. 1998), while driving the transfer
24 of an estimated $5\text{-}11 \text{ Pg C yr}^{-1}$ of organic carbon into the ocean interior as sinking particles and dissolved
25 molecules (Henson et al. 2011). Interactions and feedbacks within the Earth system therefore occur across an
26 extremely broad range of temporal and spatial scales, from subcellular processes occurring on timescales of
27 minutes or less, to global changes spanning millions of years or more (Figure 1).

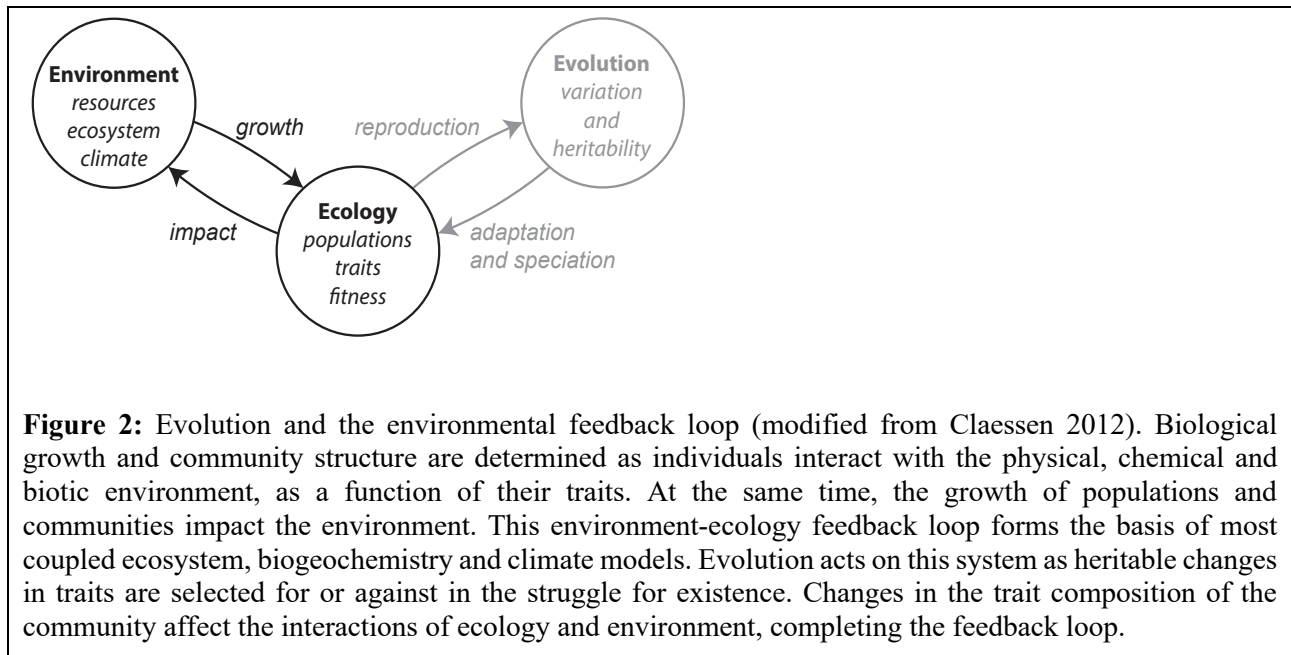
29 Within this coupled system, environmental factors including temperature, predator abundance and the
30 availability of light, nutrients and prey determine the growth of individual organisms. Conversely, the
31 combined growth of those organisms directly impacts the physical, chemical and biological environment, as
32 individuals take up nutrients, absorb light, consume prey and feed predators. As such, there exists a complex
33 feedback loop (Figure 2) with environmental conditions and ecology changing together over a broad range of
34 timescales.



35

36 This environmental-ecological feedback loop is at the core of all dynamic ocean and Earth system models,
 37 which aim to understand not only how changing environmental conditions might affect marine communities,
 38 but also how such changes might feedback into the broader Earth and climate system. The desire to understand
 39 potentially important feedbacks between marine ecosystems and their environment has been a key motivation
 40 in the development of more complex global ocean ecosystem models (Le Quéré et al. 2005), typically including
 41 a number of highly idealised plankton populations that are assumed to be representative of marine plankton
 42 communities at large. Increasing the degree of ecological complexity included in ocean models has allowed
 43 the exploration of how changes in ecosystem structure can affect both the community response to, and its effect
 44 on, the broader environment (Bopp et al. 2005; Dutkiewicz et al. 2013, 2015). Global models including
 45 representations of key functional groups and traits have led to new insights into the effects of ecological

46 diversity on ecosystem function (Barton et al. 2010; Prowe et al. 2012; Dutkiewicz et al. 2014; Monteiro et al.
47 2016).



48

49 Despite these advances, a notable knowledge gap in current global climate and carbon cycle models is that
50 most do not currently include an explicit representation of evolution, with microbial communities typically
51 represented in terms of a few idealised populations with immutable traits. This is in contrast to diverse marine
52 microbial communities that have evolved over the last 4 billion years of Earth history, and that continue to
53 evolve today on relatively short (seasonal-to-decadal) timescales (Irwin et al. 2015). A key driver of this rapid
54 evolution is the amount of genetic and phenotypic variation that natural selection can potentially act upon. In
55 addition to the generally very large standing variation already present in a community (e.g. Mock et al. 2017),
56 large microbial populations can rapidly produce novel variation through mutation, sexual recombination
57 (Blanc-Mathieu et al. 2017) and horizontal gene transfer (Llorens–Marès et al. 2016; Hall et al. 2017). While
58 beneficial mutations themselves are rare relative to deleterious or neutral ones (Desai and Fisher 2007), the
59 per generation supply of beneficial mutations for marine microbes will be large, simply by virtue of their
60 immense population sizes. For example, the mutation rate for the eukaryotic picoplankton *Ostreococcus* and
61 *Micromonas* is estimated to be on the order of 10^{-10} mutations per genome per generation (Krasovec et al.
62 2017). Even though only a small proportion of these are beneficial, the surface ocean alone is inhabited by
63 more than 10^{27} individual microbes (Flombaum et al. 2013), with typical *in situ* specific net population growth
64 rates on the order of 1 d^{-1} (Laws 2013). Most of this immense community will, therefore, completely turn over

65 every few days or so. On seasonal-to-decadal timescales, microbial populations will go through hundreds to
66 thousands of generations, experiencing novel environments and changing environmental fluctuations (Irwin et
67 al. 2015). A useful supply of novel phenotypes therefore seems assured on these timescales, and laboratory
68 experiments have demonstrated an evolutionary response on timescales as short as weeks (Bell and Collins
69 2008; Collins et al. 2014). Plankton communities are correspondingly high in genetic, phenotypic and
70 functional diversity (Armbrust 2009; Kashtan et al. 2014; de Vargas et al. 2015). The extrapolation of
71 taxonomic data estimates the total number of marine eukaryote species at ~2.2 million (Mora et al. 2011; de
72 Vargas et al. 2015). This high diversity is perhaps to be expected, given the overall complexity of the ecosystem
73 itself, but it is worth noting that such complexity takes time to develop (see for example the recovery after
74 mass extinctions described in Section 3.2).

75 Perhaps as a consequence of this high evolutionary capacity, dominant patterns of changing plankton diversity
76 in the fossil record show that rates of speciation and extinction are most closely tied to long-term environmental
77 factors (Bown et al. 2004; Falkowski et al. 2004), especially the “greenhouse-icehouse” oscillations that have
78 occurred over multi-millions of years (Bown et al. 2004). In contrast, at the (relatively) shorter end of the
79 geological scale (the glacial-interglacial cycles operating on timescales of millennia to hundreds-of-millennia,
80 the sub-millennial oscillations associated with ice sheet dynamics, and the millennial carbon-cycle perturbation
81 of the greenhouse Eocene world), changes in the climate appear to have little or no effect on plankton diversity
82 beyond normal, background levels of turnover (Gibbs et al. 2006). On these timescales, diversity and
83 ecosystem functionality appear to be set by the broader state of the biogeochemical/climate system, particularly
84 latitudinal gradients in temperature and nutrient supply.

85 Nonetheless, it still remains unclear how the changing system properties might have been different in the
86 absence of an evolutionary response. Despite what seems like a virtually limitless ability to ‘evolve’, modern
87 planktonic foraminifera and coccolithophores have surprisingly low taxonomic diversities. It appears that with
88 only ~50 and 190 uniquely-defined morphospecies, respectively (Young et al. 2005; Morard et al. 2018).
89 evolution can become an important limiting factor in response to large and rapid environmental change, such
90 as in the wake of mass extinctions. Following such events, timescales of recovery can be mapped out from
91 different milestones in the fossil record. Taking the Cretaceous–Paleogene extinction event (K/Pg) as an
92 example, the picture that emerges in the recovering ocean is that of a succession of pioneer-type communities,

93 with the first arriving within the first twenty-thousand years (Bown 2005; Schueth et al. 2015). These pioneer
94 communities, dominated by new incoming taxa that are often the first representatives of new lineages, have
95 an atypically low diversity of traits, and are characterised by very small species with opportunistic ecology
96 (Bown 2005; Gallala et al. 2009; Birch et al. 2016). Coccolithophores and planktic foraminifera see new
97 species emerge from a handful of survivor species over the first few hundred thousand years (i.e. ~50 million
98 generations of coccolithophores and perhaps 3-4 million generations of foraminifera), leading to the
99 establishment of globally distributed communities of low diversity and a slow (re-)establishment of stable
100 biogeochemical functions (Payne et al. 2004; Bown 2005; Coxall et al. 2006; Birch et al. 2016). These
101 communities have already displayed a basic expansion of morphologies and genetic diversity, but there is little
102 evidence in the first 300 thousand years of any vertical niche partitioning in the water column, endemism, or
103 a full seasonal spread of productivity (Bown 2005; Birch et al. 2016). Rather, while these opportunist taxa are
104 found throughout the oceans, there is little specialism, with no evidence of niches associated with a fully
105 developed environmental feedback loop, no differentiation along oligotrophic to eutrophic gradients, and no
106 pressure from a complex trophic web. Only much later in the recovery, perhaps after 2 million years (D'Hondt
107 2005; Coxall et al. 2006; Birch et al. 2016), do we see a sufficient co-development of the ecology and
108 environment that niches are rapidly opening up and diversity starts to increase on the trajectory of diversity
109 gain consistent with environmentally-limited models of recovery. Similarly long (multi million years) time-
110 scales appear required for the full re-establishment of ecosystem function and global carbon cycling in the
111 aftermath of the end Permian extinction (Payne et al. 2004).

112 The grand challenge outlined in this paper is to extend current global ocean and Earth system models to include
113 an explicit representation of adaptive evolution by natural selection. Just as "the interactions of organisms and
114 environment are reciprocal" (Redfield 1958), we also know that "organisms and environments are both causes
115 and effects in a coevolutionary process" (Lewontin 2000) (Figure 2). In this paper, we discuss whether (or
116 when) evolution should be included in such models of the ocean and climate system, and what we might gain
117 from doing so. In the following sections, we first review current approaches to modelling the ecology of marine
118 plankton communities within Earth system models. We then introduce several techniques that might be used
119 to extend these models to include an evolutionary perspective. From this perspective, we go on to discuss
120 potential applications of evolutionary Earth system models, concluding with a discussion of how the

121 development and application of these models might lead to a better understanding of the Earth system as a
122 whole.

Glossary

Acclimation: Physiological changes to an individual in response to environmental cues. These changes occur within the lifetime of a single organism and are reversible within bounds set by the individuals inherited traits.

Adaptation: Changes in the traits of a population across successive generations that lead to increases in organism fitness. These changes occur as a consequence of heritable changes to the organisms genotype, driven by processes including mutation, recombination and horizontal gene transfer.

Genetic drift: Changes in the relative abundance of different genotypes in a population as a consequence of stochastic processes unrelated to fitness.

Neutral: Changes in genotype or phenotype that have a negligible effect on organism fitness.

Microevolution: Small evolutionary changes within a species or functional group, driven by adaptation and genetic drift.

Macroevolution: Large evolutionary changes including the emergence of new species and functional groups.

Fitness: A measure of the relative reproductive success of genotypes, individuals or populations within a given environment. Organisms with higher fitness are more likely to be selected in subsequent generations.

Trait: Any quantifiable and measurable characteristic of an organism.

Trait space: A hypothetical multi-dimensional space, with each dimension corresponding to a measurable trait.

Fitness Landscape: A metaphorical extension of the trait-space concept, with organism fitness added as an additional dimension. For a two-dimensional trait space, the fitness landscape can be thought of as a mountain range, with fitness corresponding to elevation. Evolution by natural selection changes the values of a population's traits, such that the population ascends the slope of the mountain it sits on.

126 2 MODELLING ECOLOGY & EVOLUTION IN THE EARTH SYSTEM

127 From an Earth system modelling perspective, we might start by asking, why do ESMs even include ecology?
128 For many scientific questions regarding distributions or fluxes of biogeochemical tracers there is little evidence
129 that they need to, as the biological organic carbon pump (the production and transfer of organic matter from
130 the ocean surface to depth) and the carbonate counter pump (the production and downward transfer of
131 alkalinity associated with calcium carbonate) can, to first order, be modelled as a direct function of the
132 prevailing physical and chemical environment at any one point in time and space (Maier-Reimer 1993; Orr et
133 al. 2001; Matsumoto et al. 2004; Ridgwell et al. 2007). Nonetheless, the ecology and evolution of plankton
134 communities are dynamically and reciprocally coupled to their environment on a broad range of timescales,
135 from the diel to the geological (Figure 1). In the absence of any prognostic representation of plankton
136 populations and their interactions, such models are incapable of representing key feedbacks and are thus
137 somewhat inflexible and unavoidably coupled to the (contemporary) systems for which they were
138 parameterised. However, as the base of the marine food web, phytoplankton are also essential to fisheries.

139 Ecosystem models, particularly for future climate change scenarios, are starting to help in understanding the
140 links to fisheries (Stock et al. 2017). This, together with a desire to understand the effect of climate and ocean
141 variations on ecosystem structure and function, has motivated the development of more complex ecological
142 models. In the following section we first give a brief overview of how and why ecological complexity has been
143 accounted for in global ocean models. We then go on to make a similar case for how and why evolutionary
144 processes might be similarly accounted for.

145 2.1 MICROBIAL ECOLOGY IN OCEAN MODELS

146 Models of ocean ecology typically condense the enormous complexity of the marine ecosystem into a limited
147 number of state variables. At the broadest level of organisation, planktonic diversity in ocean models, and
148 especially those used in Earth system models, has frequently been condensed into a very limited number of
149 model populations representing generic phytoplankton producers and zooplankton consumers (Yool et al.
150 2013; Aumont et al. 2015; Kriest and Oschlies 2015). While such models benefit from lower computational
151 costs and require fewer empirical parameters, those parameters are often so highly abstracted that they cannot
152 be directly related to measurable quantities. They must instead be assigned through objective (or subjective)
153 calibration against the state of the modern ocean (Kriest et al. 2017). This empirical tuning of coarsely-resolved
154 model structures means that these models are very closely tied to the ecosystems for which they were
155 parameterised, and as such they will often perform poorly when transferred to novel environments (Friedrichs
156 et al. 2007).

157 This issue has motivated the development of more complex models that describe a broader range of diversity,
158 often focussing on ‘plankton functional types’ (PFTs), such as cyanobacteria, coccolithophores, diatoms,
159 dinoflagellates and various categories of zooplankton (Le Quéré et al. 2005; Bopp et al. 2013). Many of these
160 broad groups have been associated with particular roles in the context of ecosystem function, biogeochemical
161 cycling and climate, and have been incorporated into models with the goal of resolving key ecosystem
162 feedbacks that are necessarily absent from simpler models (Le Quéré et al. 2005; Hood et al. 2006). Such
163 models assume (consistent with the data we have for unicellular species, at least) that there is more variation
164 between functional groups in terms of represented traits such as metabolism, size, edibility and sinking
165 behaviour, than there is within functional groups. Nonetheless, the structure and parameterisation of PFT

166 models is still shaped by our observations and understanding of contemporary ocean ecosystems, and thus
167 remains closely tied to our empirical understanding of those ocean ecosystems.

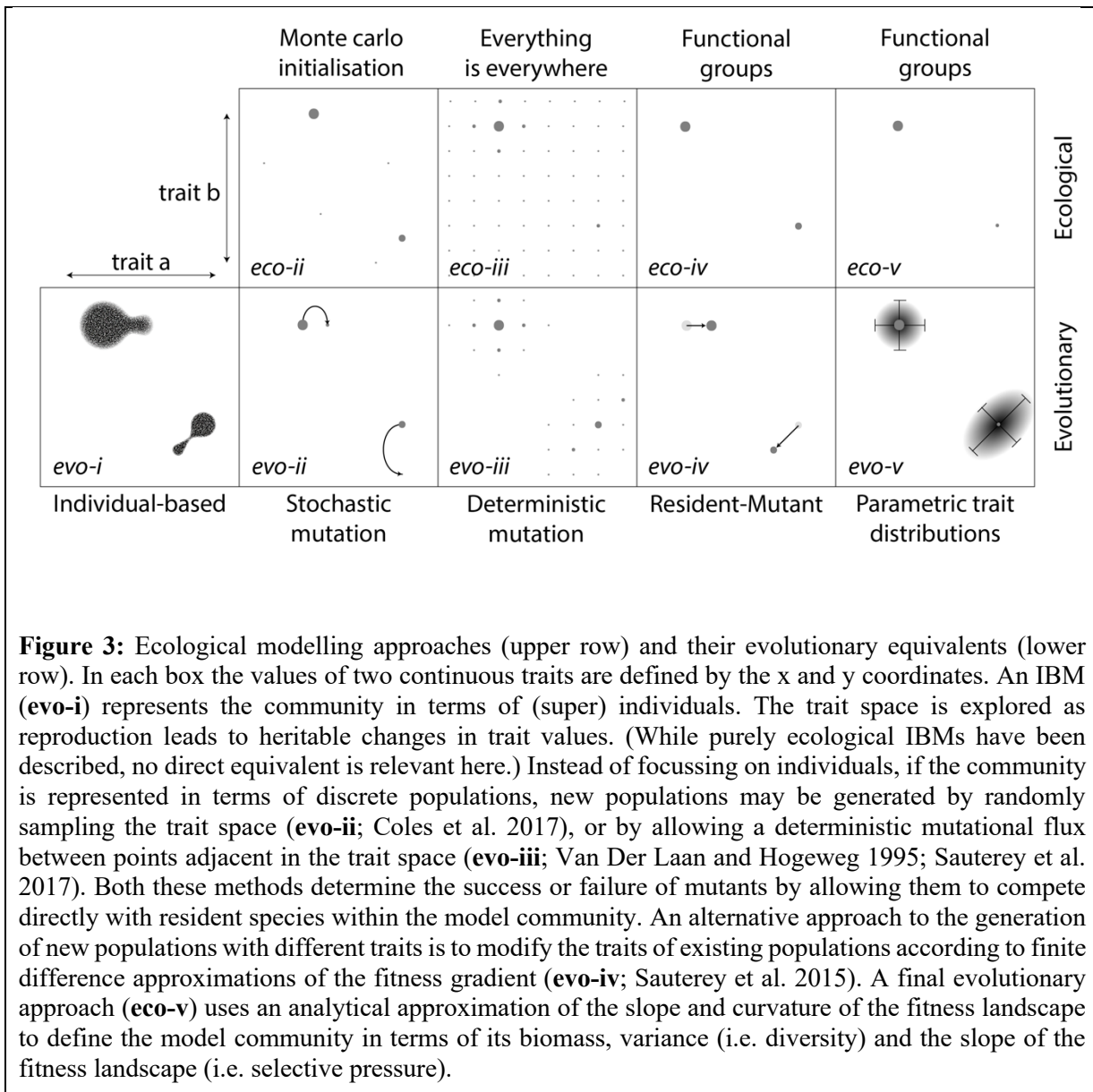
168 More recently, models have been developed to allow a more flexible ecosystem structure, initialising the ocean
169 with many tens or hundreds of potential populations, with ecophysiological and biogeochemical parameters
170 either assigned randomly (Follows et al. 2007; Coles et al. 2017) or discretely sampling across a credible range
171 of the trait-space (Bruggeman and Kooijman 2007; Ward et al. 2012). Instead of focussing on individual
172 species or functional groups, these “trait-based” models aim to identify and quantify how key organismal traits,
173 varying across species and functional groups, might underpin community structure and function. Model
174 communities are allowed to self-organise, with particular traits becoming more or less prevalent as a function
175 of the environmental conditions. In particular, trait-based models have been used to better understand the
176 ecological and biogeochemical impacts of traits related to organism size and shape, photosynthetic efficiency
177 and nutrient acquisition (see Litchman and Klausmeier 2008; Follows and Dutkiewicz 2010).

178 An essential component of the trait-based approach is to identify ecophysiological constraints or ‘trade-offs’.
179 Under the reasonable assumption that fundamental physical, chemical and biological constraints prevent the
180 simultaneous optimisation of all potentially beneficial traits, it is expected that certain beneficial characteristics
181 will come at the expense of others. Although such trade-offs are often hypothetical, a number have empirical
182 support (e.g. Litchman et al. 2007; Edwards et al. 2013), and have been shown to explain observed ecological
183 successions along environmental gradients (Tilman 1990; Ward et al. 2013). As such, trait-based models have
184 achieved notable success in terms of predicting large-scale patterns of species biogeography, and to a slightly
185 lesser extent, ecosystem function. Nonetheless, it is worth noting that our current level of understanding is not
186 yet at a point where we can use purely continuous trait-based approaches to account for highly discontinuous
187 differences between disparate lineages (e.g. diatoms vs. coccolithophores), and there remains a very important
188 role for taxonomic approaches based on discrete classifications.

189 Regardless of their degree of complexity, the parameters of ocean ecosystem models are most typically
190 constants. It is however worth noting that some models include a degree of phenotypic flexibility, with the
191 allocation of cellular resources to different metabolic functions (e.g. nutrient uptake, light harvesting and
192 photosynthetic carbon fixation) optimised in order to maximise some imposed measure of fitness, which is
193 most typically the gross population growth rate (Geider et al. 1996; Smith and Yamanaka 2007; Pahlow and

194 Oschlies 2009). These optimal acclimation models go some way to capturing the dynamic behaviour of the
195 marine microbial community and its potential to change (Smith et al. 2009). This response is, however, more
196 or less immediate and changes will not be inherited to allow for the generation of new lineages.

197 While the ecological models described above may differ in their underlying construction, they all share the
198 same mechanism of operation, in that community structure and function are determined by the outcome of
199 competition among a predefined set of model populations. Despite the potential for the emergence of what
200 look like new species within the community, such models should not be thought of as reflecting Darwinian
201 evolution. All populations are defined *a priori*, with essentially immutable traits. (A notable exception to this
202 limitation is the model of Coles et al. (2017), in which new species with randomly generated traits replace
203 those that have become extinct). There is no heritable variation of traits, and the potential for growth of
204 different populations is limited strictly to those populations that were included in the model initially. Within
205 this characteristic is the implicit assumption that the rate of evolution is not limiting to the emergence of new
206 phenotypes. If a niche is available, it will be filled, and it is not necessary to account for the time taken for a
207 species or function to evolve, or for it to disperse globally. In other words, “Everything is everywhere, but the
208 environment selects” (Baas-Becking 1934). Is this consistent with the observation that it has taken perhaps
209 twenty thousand years for the first new species to appear after the K/Pg event? A fundamental quality of
210 evolution is that it can only act on what is already present (Jacob 1977), and if we are interested in examining
211 whether the state of the marine ecosystem is sensitive to its evolutionary history, it is essential that our models
212 are able to reliably capture both standing variability and the limited generation of novel phenotypes.



213

214 2.2 MICROBIAL EVOLUTION IN OCEAN MODELS

215 At its core, Darwinian evolution requires two key features: a struggle for existence and the heritable variation
 216 of traits (Lewontin 1970). In isolation, the struggle for existence will lead to the regulation of population
 217 growth, and the reorganisation of communities, reflected in the dominance of some species at the expense of
 218 others (Claessen 2012). While the struggle for existence encompasses all facets of life, it is typically
 219 represented in ocean ecosystem models in terms of resource competition and, to a lesser extent, predation (e.g.
 220 Prowe et al. 2012; Ward et al. 2012; Vallina et al. 2014b).

221 The ecological models described in the previous section include the struggle for existence, but any variation
 222 was not heritable. In practice, the inheritance of variation can be modelled either through the generation of

223 new individuals or populations with variant traits, or the same effects can be represented by allowing the traits
224 of extant populations or communities to vary on intergenerational timescales. In the following, we will review
225 a number of different evolutionary models, that move from the individual to the community as the modelled
226 unit of selection (Lewontin 1970).

227 *2.2.1 INDIVIDUAL-BASED MODELS (IBMs)*

228 By modelling the ecology and behaviour of individuals it is, in principal, relatively straightforward to
229 incorporate evolution by allowing the traits of offspring to deviate from those of their parents. Better adapted
230 offspring will more likely prevail, allowing both adaptation and speciation. In practice, the evolution of marine
231 microbial communities is determined by the interaction of more than 10^{27} unique individuals in a highly
232 interconnected system. It is obviously impossible to resolve anything even close to this number of agents in a
233 model system, so microbial IBMs have typically focussed on homogenised ‘super-individuals’, which are
234 assumed to represent the behaviour of a large number of identical individuals that exhibit the same responses
235 to environmental conditions. The dynamics of super-individual growth may be represented in a number of
236 ways (as reviewed by Hellweger et al. 2016), but a particular challenge for modelling their evolution in marine
237 communities is to maintain a computationally tractable number of super-individuals in the face of rapid
238 dispersal. As a consequence of the highly interconnected nature of marine ecosystems, local populations in a
239 3D circulation model are likely to become very rapidly diversified by a large number of immigrant populations.
240 One way to handle this is to merge the least abundant, or alternatively the most similar, super-individuals at
241 each location (Woods 2005; Clark et al. 2011). This approach may, however, make it difficult to maintain
242 anything but the most abundant species across a global model, and to date, global plankton IBMs have avoided
243 horizontal mixing of plankton communities (Clark et al. 2013; Daines et al. 2014).

244 Nonetheless, a particular advantage of IBMs is that each (super-) individual can be directly linked to its
245 evolutionary antecedents, such that it should be possible to retrace the complete spatial and temporal
246 evolutionary history of the community (given sufficient computational resources). It is also relatively
247 straightforward to incorporate sexual reproduction, which is more problematic in population- or community-
248 based models.

249 *2.2.2 STOCHASTIC GENERATION OF NEW PHENOTYPES*

250 Plankton communities are often represented in ocean ecosystem models as a set of competing populations.
251 Typically, these populations are fixed, but some models allow the generation of new populations with new
252 phenotypes. For example, Coles et al. (2017) describe a model with the random generation of new phenotypes
253 from within a predefined range of credible traits. While it should be noted that this process is not strictly
254 analogous to Darwinian evolution by natural selection (the generation of new phenotypes is independent of
255 the existing members of the community), it could be adapted to include the generation of new phenotypes from
256 existing members of the community by mutation, sexual recombination or lateral gene transfer. For example,
257 mutant traits of a new population could be drawn from a normal distribution, centred on the appropriate trait
258 value of the ancestral population.

259 In the interests of maintaining a computationally tractable model, the generation of new model populations
260 must be balanced by the compensatory removal (i.e. extinction) of populations. In the Coles et al. (2017)
261 model, this was achieved by replacing any populations accounting for less than 1% of the local biomass at any
262 location in the model in a given period. While this approach successfully allowed the exploration of over 2000
263 different phenotypes in just 20 years, the defined "extinction threshold" is somewhat arbitrary, and it will be
264 important to assess the sensitivity of results to the assumed definition of extinction, and the spatial and temporal
265 scales across which it is defined.

266 *2.2.3 DETERMINISTIC GENERATION OF NEW PHENOTYPES*

267 The stochastic generation of new phenotypes allows the evolutionary exploration of the trait space, with
268 individual traits treated as continuous variables. An alternative approach is to divide the trait space into a
269 discrete grid, with mutations treated as a small diffusional flux between adjacent points in the grid (Polechová
270 and Barton 2005; Leimar et al. 2008). This approach is similar to discretised trait-based models (Bruggeman
271 and Kooijman 2007), with a key difference being that a small fraction of reproduction in each population is
272 diverted to neighbouring populations in the trait space (Sauterey et al. 2017). Additionally, the assumption that
273 no species can go extinct is removed, with new populations only emerging via mutation (or through an explicit
274 representation of spatial dispersal).

275 It is however worth noting that the need to resolve the entire trait space as a discrete grid may be
276 computationally intractable for high resolution ocean models, especially when multiple traits are considered.
277 In addition, the diffusive nature of traits may make it difficult to provide a robust definition of species –
278 although this may in fact be a desirable (or at least realistic) characteristic in the microbial realm (Rosenberg et
279 al. 2013).

280 2.2.4 CHANGING THE TRAITS OF A POPULATION

281 As an alternative to modelling adaptation as the succession of individuals or populations with different traits,
282 it is possible to treat a population as a fixed entity, allowing its traits to vary. This is achieved by allowing a
283 population to undergo incremental changes to its traits, in order to maximise some defined fitness metric. This
284 ‘Adaptive Dynamics’ approach (Geritz et al. 1998; Litchman et al. 2009; Kremer and Klausmeier 2017) works
285 by comparing an ecologically-established ‘resident’ population to a ‘mutant’ population with slightly modified
286 traits. This requires the definition of some metric of ‘invasion fitness’, which is most typically defined as the
287 net population growth rate integrated over some temporal (and possibly spatial) scale. Under the assumption
288 that the resident population is at (or at least close to) an equilibrium or limit-cycle, its invasion fitness will be
289 zero. A positive invasion fitness for a mutant can therefore be taken as evidence that it is better adapted to the
290 local conditions, and the mutant is assumed to replace the resident (Geritz et al. 2002). In practice, this can be
291 achieved by assigning the adapted traits of the mutant to the resident population. This allows for adaptation and
292 speciation, with the latter occurring if mutations in opposite directions of trait space are both associated with
293 a positive invasion fitness (Fig. 3, panel *evo-iv*).

294 In practical terms, the calculation of mutant invasion fitness is greatly simplified by the assumption that the
295 initial abundance of mutants is sufficiently small that they have no effect on the ecosystem. The absence of
296 any feedbacks means that the invasion fitness of mutants can be calculated as a prognostic function of the
297 broader ecosystem, without any requirement to assess their reciprocal impact on the system.

298 This approach requires a number of basic assumptions that can be questioned, especially in a large-scale
299 spatially-resolved model. First, the fact that the resident species is assumed to be at or close to equilibrium
300 requires the artificial separation of ecological and evolutionary timescales in the model, with mutations and
301 replacements only occurring after re-equilibration of the ecological model from the previous event. Second, it
302 is assumed that a positive mutant invasion fitness always leads to replacement of the resident (Geritz et al.

303 2002). This may not occur in reality if, for example, a locally maladapted species is maintained in a region by
304 immigration. More generally, the selection of a suitable fitness metric and the temporal and spatial scales over
305 which it is defined is an arbitrary process, and one that is likely to impact the evolutionary trajectory and
306 endpoint in the model system.

307 *2.2.5 MODELLING THE DISTRIBUTION OF TRAITS IN A COMMUNITY*

308 In the models described above, microbial diversity is represented by a finite number of individuals or
309 populations, each corresponding to a single point in the phenotypic trait space. As an alternative, we can
310 represent the diversity within a community as a continuous distribution of biomass across the trait space,
311 resolving state variables for the total community biomass alongside the mean and variance of the trait
312 distribution (e.g. Bruggeman 2009). Using this continuous approach, the fitness landscape is defined by the
313 biomass-specific (net or gross) growth rate, itself a continuous function of organism traits and the current
314 environment. This function may be quite complicated, and it is typically approximated by a Taylor polynomial,
315 under certain simplifying assumptions. In general,

- 316 1. The biomass-specific community growth rate is approximated as the area under the fitness landscape,
317 which is often assumed to follow a Gaussian distribution that is centred symmetrically around the mean
318 trait value.
- 319 2. The mean trait value is assumed to move up the fitness gradient, at a rate proportional to the slope defined
320 at the mean trait value and the variance of the biomass in trait space. This reflects the rate of adaptation
321 increasing with both selection pressure and the standing diversity of the community.
- 322 3. Diversity (i.e. trait variance) increases if the curvature of the fitness landscape is positive. Positive
323 curvature implies that selection pressure increases towards the optimum, with the traits of better adapted
324 members of the community diverging from those of less well adapted members. The converse is true if
325 the curvature of the fitness landscape is negative. As this is always the case at a fitness maximum,
326 equilibrium diversity tends to collapse to zero, unless it is maintained by external variability or some
327 additional mutation or immigration term (this is not necessarily unrealistic given the assumptions of the
328 model, and certainly not an exclusive feature of this kind of model).

329 This ‘moment-based’ representation of community traits is computationally efficient, and has been generalised
330 to be applicable in a spatial context (Bruggeman 2009; Chen et al. 2019). It is worth noting that the approach

331 does not distinguish between ecological and evolutionary changes. Furthermore, while the approach has
332 generally been applied to describe the ecological and evolutionary dynamics of plankton ‘communities’, these
333 model communities are really more like ecological guilds, because the continuous representation of traits is
334 best suited to the description of groups of organisms exploiting similar resources. While the approach could in
335 principal be extended to represent the interaction of multiple guilds (or functional groups), it has so far only
336 been used to represent generic phytoplankton (Merico et al. 2009; Chen et al. 2019). It also remains uncertain
337 as to whether the approach can be extended to explicitly incorporate diverse predator-prey dynamics, for which
338 disruptive selection (i.e. speciation) is likely to play a key role (Vallina et al. 2014b). In such cases where the
339 slope of the fitness landscape is zero while its curvature is positive, the modelled variance would rapidly
340 increase towards infinity.

341 3 WHAT CAN WE GAIN?

342 We know that adaptive evolution plays a key role in shaping the whole Earth system, but to what extent, if at
343 all, is it advantageous to include it explicitly in models of the ocean and Earth system? While the ecological
344 models described in Section 2.1 have all been applied within global ocean ecosystem and Earth system models,
345 this is not the case for the evolutionary models summarised in Section 2.2. An outstanding challenge is,
346 therefore, to take these idealised evolutionary models and integrate them into Earth system models, so that
347 potential feedbacks can be explored. If we consider that the evolution of marine microbial communities takes
348 place within a much broader and highly interconnected system, there is clearly potential for new insights. In
349 particular, regarding how microbial evolution affects the Earth system, and how the Earth system affects
350 microbial evolution. However, the question remains: can we reliably develop a predictive ecosystem model
351 that is structured and parameterised according to our (limited) observations of phenotypes living in
352 contemporary ocean communities? The marine ecosystem we see today is the current state of a co-evolving
353 system encompassing life and the planet. In this regard, it might be desirable to develop models based on the
354 fundamental constraints that have shaped past and present ecosystems, rather than to develop models encoding
355 the current state itself. Evolution is a central part of this process, although it remains to be seen whether
356 modelling evolutionary processes can deliver improved predictions or novel insights that would not be

357 available from current Earth system and ecosystem models. In the following we will address key areas where
358 an evolutionary perspective might help to improve our understanding of both evolution and the Earth system.

359 3.1 INTERPRETING THE PAST

360 With the plankton fossil record (planktic foraminifera, coccolithophores, and, to a lesser extent, diatoms,
361 radiolarian and dinoflagellates) we have unrivalled taxonomic and stratigraphic completeness compared with
362 any other organismal groups, extinct or extant (Bown et al. 2004; Falkowski et al. 2004). We are able to track
363 both micro- and macroevolutionary patterns from millennial-scale records of skeletal size and morphological
364 variations (e.g. Schmidt et al. 2004; Finkel et al. 2007; O’Dea et al. 2014), up to the major landmarks of
365 physiological innovation such as evolution of carbon concentrating mechanisms, acquisition of
366 photosymbionts and new biomineralisation strategies (e.g. Bolton and Stoll 2013; Birch et al. 2016; Monteiro
367 et al. 2016). The planktonic fossil record is, therefore, a valuable source of information on both the past history
368 of the marine ecosystem, and the state of the Earth system. However, our interpretation of that record is shaped
369 by assumptions regarding the links between the environment, ecology and evolution – from geological records
370 alone we can only hypothesize causation, while the quantitative (or even qualitative) outcome of the interaction
371 of multiple feedbacks is often impossible to diagnose.

372 An important advance would therefore be the ability to draw self-consistent quantitative links between precise
373 hypotheses and paleontological/geochemical observations. One goal might therefore be the development of
374 quantitative models describing the behaviour of an evolving ecosystem in response to environmental
375 perturbations over the appropriate timescales. For example, it is not known to what extent the pace of recovery
376 after mass extinctions is limited by climatic, biogeochemical or evolutionary processes, or by feedbacks among
377 all three (D’Hondt 2005; Hull 2015). Taking the example of the ecological and Earth system response to the
378 impact at the end of the Cretaceous, the marine carbon cycle transitioned into an apparently radically different
379 ‘mode’, characterised by some combination of decreased export production and shallower recycling of carbon
380 and nutrients, leading to a weaker oceanic carbon sequestration. This state persisted for over several millions
381 of years before recovery was achieved (Coxall et al. 2006). Why? What prevented rapid recovery of the system
382 once the initial perturbation had subsided? Does this simply reflect the time-scale for the re-evolution of
383 important traits or the reestablishment of community structures? Or does it reflect the evolution of a new quasi-

384 steady-state of tightly coupled ecology and carbon and nutrient cycling that was only slowly ‘eroded’, with
385 full ecological function only much later re-attained?

386 In the most general terms, does the emergence of new species occur rapidly with the emergence of new climates
387 and habitats or is the functional response of the ecosystem limited by the evolutionary response of the plankton.
388 Temporarily leaving aside questions of feasibility (Section 2), such questions and hypotheses about the past
389 can only be explicitly and quantitatively tested, explored, and falsified in a model system that includes adaptive
390 evolution alongside a representation of global climate and carbon cycling. Placing an evolutionary model
391 within a coupled Earth system model might also provide a means to generate testable predictions that can be
392 compared to the palaeoceanographic record. Alongside the evaluation of hypotheses within a constrained
393 environmental context, such an approach might additionally help to constrain unknown parameters concerning
394 the rate and capacity for evolution in marine microbial communities.

395 3.2 UNDERSTANDING THE PRESENT

396 Evolution has shaped marine ecosystems and the Earth’s biosphere over the past 4 billion years (Falkowski et
397 al. 2004) and continues to do so in the present, on timescales of decades or even less (Collins et al. 2014; Irwin
398 et al. 2015). Changes in the fitness of different phenotypes can be brought about by changes in any part of the
399 overall system, across a broad range of timescales. As such, it is always likely that some kind of adaptive
400 change is under way. In practical terms, this means that the set of phenotypes present in a community is not
401 fixed, nor is it necessarily stable. So instead of thinking of wild populations as unchangeable archetypes upon
402 which we should base our models, we should perhaps instead consider the current state of the system as a
403 waypoint along a constantly developing trajectory. The key challenge in this regard is to identify the main
404 constraints that shape the evolutionary trajectory.

405 The mechanisms and potential impacts of evolutionary adaptation would likely also benefit from consideration
406 within a realistic environmental context. One possibility would be to examine the structural and functional
407 sensitivity of microbial food-webs (Loeuille and Loreau 2005) under different assumptions of environmental
408 dispersal, ecological productivity and evolutionary rate. Such a model would give new insights into how
409 ecosystems might respond to, and recover from, perturbations brought about by ocean warming and changes
410 in circulation. It may also be revealing to examine the evolutionary stability of mechanisms related to the

411 stoichiometry of inorganic and organic nutrient elements (Redfield 1958; Tyrrell 1999; Klausmeier et al.
412 2004). Evolutionary models may also offer new insights regarding community assembly and coexistence of
413 marine plankton (Sauterey et al. 2015, 2017; and Figure 3), their resilience to environmental fluctuations
414 (Kremer and Klausmeier 2017), and their potential for catastrophic regime shifts or functional collapse (Lenton
415 et al. 2008).

416 In analogy, the development of complex ecological models within ocean and Earth system models has allowed
417 environmental scientists to explore the large-scale impact of ecological processes that might previously have
418 only been understood in highly idealised contexts. For example, concepts from the metabolic theory of
419 ecology, resource competition theory and ecological stoichiometry have recently been assessed within models
420 accounting for complex community ecology within a heterogenous fluid environment (Göthlich and Oschlies
421 2012; Ayata et al. 2014; Vallina et al. 2014a). While these concepts have a rich theoretical background, their
422 application within more realistic environmental frameworks has allowed closer comparison to observations
423 and better assessment of global impacts.

424 Adding an explicit representation of evolution to contemporary models of the ocean and Earth system will
425 provide new constraints on the current structure and function of the marine ecosystem, in particular providing
426 new context for a wealth of bioinformatic data. For example, while some genomic data points to the stable
427 coexistence of many hundreds of phenotypically distinct subpopulations of *Prochlorococcus* (Kashtan et al.
428 2014), considerable uncertainty remains as to how much observed diversity is actually driven by natural
429 selection, rather than genetic drift and neutral evolution (Hellweger et al. 2014).

430 3.3 PREDICTING THE FUTURE

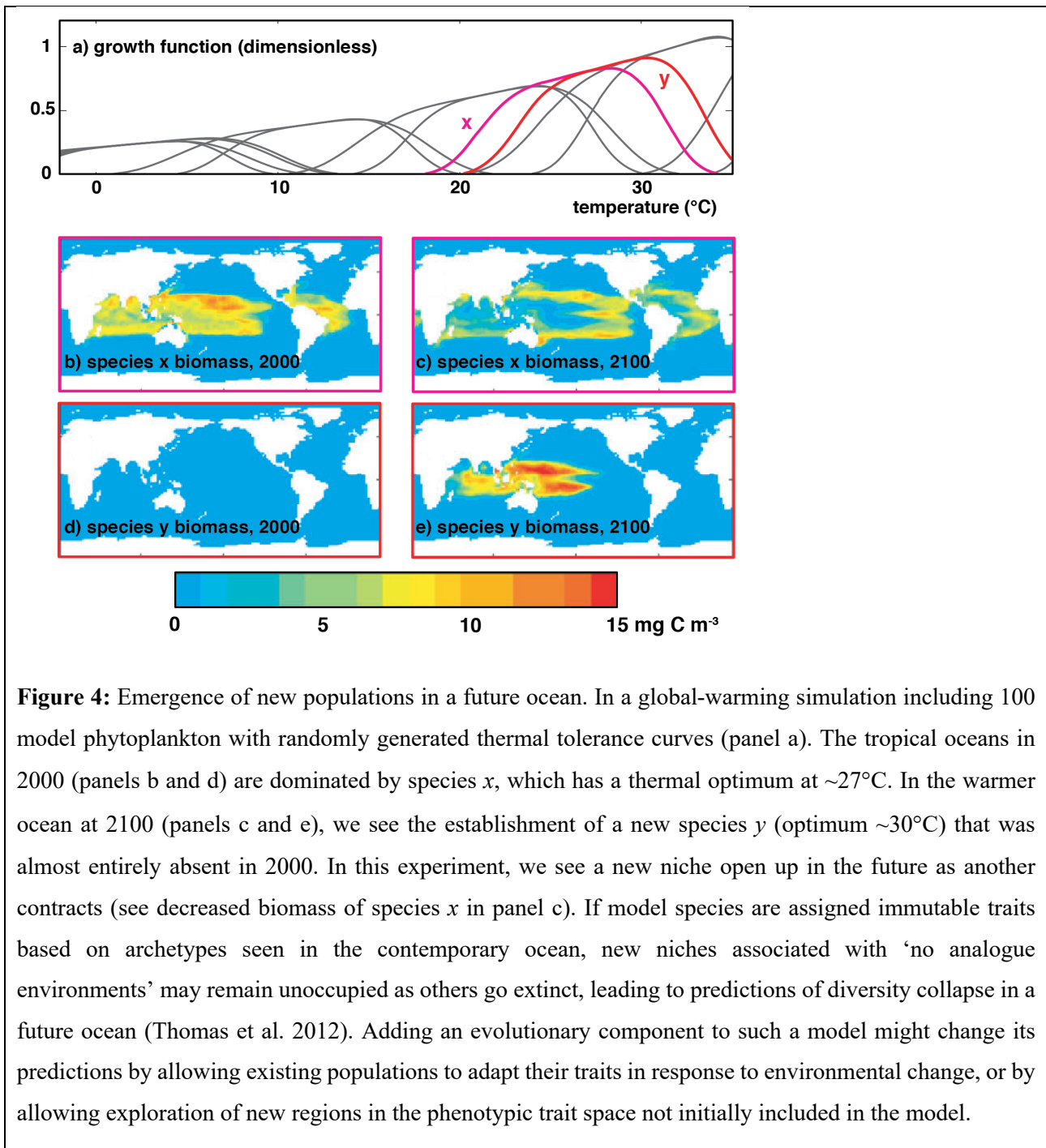
431 Under a future ‘business-as-usual’ scenario, the current generation of climate models predict changes in global
432 sea surface temperature of approximately 3-5°C (Gruber 2011; e.g. Bopp et al. 2013), manifested in a poleward
433 shift of isotherms over the course of the 21st century (Dutkiewicz et al. 2013). On these timescales there will
434 be increased stratification of the surface waters (Bopp et al. 2001; Gruber 2011). Coupled to a predicted
435 slowdown of the meridional overturning circulation (Schmittner 2005), these changes are expected to perturb
436 the light environment and decrease the oceanic supply of nitrate and phosphate to the euphotic zone, with an

437 associated decrease in ocean ventilation (Shepherd et al. 2017). As pCO₂ increases, oceanic pH is predicted to
438 fall (Doney et al. 2009; Gruber 2011), while warming temperatures will drive a further retreat of sea-ice.

439 Projected changes in global temperature are small in comparison to global temperature range, so it is likely
440 that the dominant response to global warming in a well-connected ocean (Jönsson and Watson 2016) will
441 continue to be range shifts (Dutkiewicz et al. 2013, 2015; Barton et al. 2016). That said, higher temperatures
442 will affect global average growth rates, while decreased nutrient supply is generally predicted to shift
443 communities towards smaller organisms. Such changes are expected to drive a poleward expansion in the
444 dominance of small, warm-adapted species (Bopp et al. 2005; Dutkiewicz et al. 2013; Barton et al. 2016).

445 Decreasing ocean pH is also expected to affect the role of calcifying plankton with regard to carbon
446 sequestration (Hofmann and Schellnhuber 2009), as well as affecting phytoplankton metabolic rates (Mackey
447 et al. 2015). Unlike temperature, open ocean pH changes are going to be larger than the current global range
448 (Bopp et al. 2013; Dutkiewicz et al. 2013). Increasing hypoxia will diminish the metabolic capacity of animals
449 (Deutsch et al. 2015). Alongside these largely physiological effects, we will likely see both extinctions and the
450 opening up of new niches, especially at the extreme frontiers of environmental change, in the polar and
451 equatorial oceans (Thomas et al. 2012).

452 Most current marine ecosystem models use a power law like function (e.g. Eppley 1972) for increased
453 phytoplankton growth rate with temperature. Thus there is a play off between reduction in growth rates with
454 lower nutrient supplies, and increased growth with higher temperatures in future world scenarios (e.g. Taucher
455 and Oschlies 2011; Dutkiewicz et al. 2013; Laufkötter et al. 2015). But different species of phytoplankton
456 cannot grow over the full temperature range, and in fact die out quickly beyond an optimal maximum (Thomas
457 et al. 2012; Boyd et al. 2013; also see representation in Fig 4). Thus, current models potentially over-estimate
458 phytoplankton adaptability. Almost no models considered changes in metabolic rates with increased pH, and
459 as such also over-estimate adaptability. On the other hand, though, in one study that does include pH effects
460 on growth rates (Dutkiewicz et al. 2015), extinctions of some species occurred due to differing pH responses.
461 However, including evolution might not allow such extinctions.



462

463 How might adaptive capacity change the future projections made by climate models? It is important to reiterate
 464 that model predictions are based upon the characteristics of plankton communities in today’s ocean. If some
 465 species were able to adapt to lower nutrient conditions, to grow even faster with increased temperatures, or to
 466 tolerate decreased pH or O₂ concentrations, some more pessimistic model predictions might be somewhat
 467 tempered. Predictions of the decline in larger phytoplankton with subsequent impact on fisheries (Stock et al.
 468 2017) might also be altered if those phytoplankton could adapt to some of the changing conditions. The
 469 estimated vulnerability of the global plankton community might well be very different if their ability to adapt

470 changing communities is fully and realistically accounted for. The current predictions of climate models must
471 be remembered to be derived from a view that is restricted by the lack of evolution. Studies to explore what
472 and how much evolution could change these results are therefore timely and important. The case for adaptive
473 evolution in ocean models

474 While it is clear that evolution shapes marine microbial communities on all timescales from the ecological to
475 the geological, the question remains as to whether adaptive evolution should be included in marine ecosystem
476 and Earth system models. Arguably the strongest signal present in the fossil record is that background species
477 turnover and diversification following mass extinctions is paced by the coupling of ecology and the
478 environment (Bown et al. 2004; Falkowski et al. 2004). Adaptation and the emergence of new species have
479 occurred in lockstep with the vast majority of environmental changes over the last several hundred million
480 years, and fossil evidence suggests that evolution will eventually fill any feasible niches that are opened up by
481 environmental change, as long as there is enough time to do so. Among microbial species in particular, the
482 ability to inhabit extreme environments, such as hot springs and hypersaline lakes suggests that the
483 evolutionary capacity to occupy diverse niches extends beyond the range of conditions we are likely to consider
484 in ocean models. For simulations across the longest geological timescales (millions of years and beyond), as
485 long as we are reasonably certain that it can do so within those timescales, we are perhaps less interested in
486 how evolution fills a new niche, and more in what the endpoints of the evolutionary process turn out to be. In
487 particular, while it is generally assumed that these evolutionary endpoints are (at least functionally) singular,
488 the presence of nonlinear eco-evolutionary feedbacks (Figure 2) suggest we should not rule out the possibility
489 of bifurcations and multiple stable states. However, any attempt to explicitly model the evolutionary
490 emergence of new functional groups on such timescales is at odds with the assumed unpredictability of such
491 large macroevolutionary changes. Instead, it might only be necessary to ensure that models include sufficient
492 ecological flexibility to allow the restructuring of communities in response to environmental change. This sort
493 of approach is already possible with some trait-based models (Bown et al. 2004; Falkowski et al. 2004), but it
494 is of course essential that constraints on the ecological community are correctly described.

495 As timescales constrict, the likelihood of evolution finding those endpoints decreases, and the pathways of
496 evolution become much more important. Where we are concerned with more extreme and rapid climate change
497 (on the order of a few tens to hundreds of thousands of years), there are hints in the fossil record that evolution

498 acts as a general pacer for the emergence of new species (Gibbs et al. 2006), perhaps pointing to a millennial-
499 scale pacing of the accumulation of species-specific trait differentiation. On the even shorter timescales of
500 anthropogenic climate change, it appears that species are able to adapt certain traits to decadal changes in the
501 environment (Irwin et al. 2015). This is in contrast to most empirical and mechanistic model projections for
502 the next century, which typically assume that species traits and niches will remain fixed (Thomas et al. 2012;
503 Barton et al. 2016). A key unknown at this stage is, therefore, how an evolutionary response might modify
504 future projections made by climate models, in terms of the both the community response and future climate
505 projections. For example, in a changing environment, do constraints associated with the adaptive modification
506 of traits affect the functional sensitivity and overall resilience of the ecosystem? If they do, is there a critical
507 rate of environmental change above which the ecosystem loses its ability to keep pace with environmental
508 change?

509 More generally, an important question to answer is whether or not integrating microevolutionary processes
510 into ocean and Earth system models is useful. From a purely practical standpoint, allowing microevolution of
511 traits can increase efficiency by removing the need to include a very large number of potential phenotypes, the
512 majority of which are likely to be inviable at any given time (Coles et al. 2017). More importantly perhaps,
513 will adding a representation of microevolution change anything relative to current models? Accounting for
514 this microevolutionary response on relatively short time-scales is undoubtedly challenging, but it is probably
515 a more realistic short-term goal than modelling the (somewhat unpredictable) macroevolutionary emergence
516 of new functional groups. Even though individual genetic changes may be stochastic and unpredictable, the
517 aforementioned large population sizes coupled to strong selective pressures mean that functionally similar
518 adaptations are likely to occur in similar environments, even if the genetic or metabolic basis for those changes
519 may be less convergent. The more fundamental a trait is, the more likely it is to be a case of convergent trait
520 evolution. Thus, while evolutionary trajectories are strongly history dependent, evolution in fundamental traits
521 that are strongly correlated with fitness can evolve convergently when selection is intense enough.
522 Nonetheless, natural selection can only act on what is present in the community, so the evolution of that
523 community will always be somewhat constrained by its past history. This is something that the ‘everything is
524 everywhere’ approach cannot account for, because it imposes no limits to the emergence of new populations
525 (beyond those dictated by prior constraints on the feasible trait-space). If everything is not everywhere, then
526 the whole-system response to (for example) extinction events will depend on the rate and degree to which

527 functional traits can reach new environments, either by evolution *in situ*, or through immigration from other
528 locations. Given that the adaptive response is inherently stochastic, it will also be important to assess the
529 associated uncertainty in the evolutionary pathway.

530 Since plankton are represented in models by their functional traits, some of the key challenges to integrating
531 microevolution into trait-based models revolve around understanding how natural selection acts on the traits
532 used in marine ecosystem models. Different patterns of environmental change, including the amplitude,
533 frequency and predictability of fluctuations influence both the strength of natural selection and the phenotypes
534 that it favours (Kremer and Klausmeier 2017). Natural selection is driven by differences in fitness, but we tend
535 instead to measure functional traits, often with only a partial understanding of how well they correlate with
536 fitness in complex environments and over entire organismal life cycles (Rengefors et al. 2017). Because of
537 this, representing competition between phytoplankton types in models through resource competition and
538 temperature tolerance curves may not predict the correct succession or dominance on the timescales of interest.
539 For example, across several types of bacteria, performance in variable thermal environments is not predicted
540 by performance across a range of constant temperature environments (Saarinen et al. 2018), raising concerns
541 as to the utility of these parameterisations in certain models.

542 Whether we are dealing with ecological or evolutionary trait-based models, one of the key things we need to
543 know are the fundamental limits of the feasible trait space (i.e. the full set of physically attainable and
544 biologically viable trait combinations). While plankton traits and trade-offs are still poorly constrained,
545 significant advances have been made in recent years thanks to multi-species compilations of organismal traits
546 (Edwards et al. 2012, 2013; Marañón et al. 2013). These have been worth considerably more than the sum of
547 their experimental parts, because they have helped to identify the limits of the trait space across contemporary
548 communities. Targeted lab experiments also provide useful constraints, particularly those that identify
549 fundamental limits, such as temperatures of protein degradation, the minimum elemental composition of a
550 viable cell (Finkel et al. 2016), or the fundamental size-dependence of metabolism (Savage et al. 2007).

551 One of the limits of trait-based approaches is that ultimately, the modeler must decide what traits to include in
552 a model, and what are the limits and trade-offs associated with those traits. These decisions are heavily
553 weighted towards our understanding of contemporary marine communities, for which natural selection will
554 have likely culled the realised trait variability to a significant extent. As a result, trait distributions in extant

555 populations may not represent the full range of trait values accessible to taxa, and in a changing ocean it is
556 possible that previously unseen trait values or combinations will be (or have been) favoured by natural
557 selection. While it seems likely that past or future environments with “no contemporary analogue” may be of
558 limited importance over most of the open ocean, trait-based models parametrised using extant communities
559 may be unable to predict how some key taxa will respond to novel environments. Here laboratory-based
560 evolutionary experiments may help to bridge the gap between experimental data based on extant species, and
561 the fundamental limits of plankton physiology. While such experiments are almost certainly not going to reveal
562 what the “future” version of a particular organism will look like, they will help to identify genetic and statistical
563 correlations between traits along the evolutionary trajectory. Additionally, they can show how far the
564 evolutionary process can take certain traits along that trajectory before the fundamental limits of evolution are
565 reached.

566 4 CONCLUDING REMARKS

567 Over the longest timescales spanning millions of years, we would expect that the high evolutionary capacity
568 of marine microbes should lead to convergent evolution, at least amongst the most fundamental traits, such as
569 organism size, maximum growth rate or optimal temperature, even if the genetic or even physiological basis
570 for those changes is more contingent to chance. We therefore expect that over geological timescales, ocean
571 and climate models should give reasonably robust results regardless of whether evolution is explicitly included,
572 as long as the model ecosystem is flexible enough to reflect all potential changes in ecosystem structure
573 (Dutkiewicz et al. 2013).

574 On the other hand, including an explicit representation of evolution by natural selection is likely to become
575 more important when model timescales overlap with the likely trajectory of evolutionary change. This would
576 occur for questions pertaining to the recovery from mass extinctions, and the contemporary ecosystem’s
577 response to anthropogenic climate change. In the former case, an evolutionary model would allow us to
578 disentangle the interrelated effects of environmental change and the evolutionary emergence of new
579 phenotypes and communities. In the latter case, adding evolution to climate models would allow us to assess
580 the effects of evolution on ecosystem stability, its resistance to change and its resilience following change,
581 exploring how the magnitude and rate of global or local change might affect these responses.

582 Finally, the geological record reveals the most profound changes occurring in marine species and ecosystems,
583 oxygenation of the atmosphere and ocean, and climate over the past 3-4 billion years of Earth history. These
584 changes have been relatively slow (overs 10-100s of millions of years) at times, and at others, extremely rapid.
585 Some have been directional and monotonic, others partially or even completely reversed. Many, if not all, are
586 characterized by correlations between life and the environment. One of the most profound questions in
587 understanding our planet is the causality of these relationships – to what degree does one drive the other, or
588 more likely, there is a tight coupling in the colloquial ‘co-evolution of life and the planet’. Understanding the
589 environmental feedback loop (Figure 2) *necessitates* that the interaction elements (ecology, evolution and
590 environment) are dynamically represented and linked in models. The time-scales concerned and the sparse
591 constraints on past ecosystem and environmental conditions present immense challenges to models, both in
592 terms of raw computational cost as well as developing appropriate parametrisations for only partially known
593 elements. Hence in our opinion, to fully address such fundamental questions about our planet’s history, to
594 better contextualise our ecological understanding of contemporary communities, and to develop more reliable
595 projections of future marine ecosystem impacts and recovery, we need to start to incorporate the dynamics of
596 adaptive plankton evolution into Earth system models.

597

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615 ACKNOWLEDGEMENTS

616 This work is dedicated to the memory of David Claessen. We thank Mark Moore and Andrew Yool for
617 helpful comments on earlier drafts of the manuscript. BAW is supported by A Royal Society University
618 Research Fellowship. This work was also partly supported by the European Research Council
619 “PALEOGENiE” project (ERC-2013-CoG-617313) and NASA grant NNX16AR47G

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