Title: Local water year values for the conterminous United States

Running title: CONUS local water years

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- 11
- 12 Author Contribution Statement: XS and KSC conceived the project and designed the overall
- 13 approach. XS collected and processed the data, wrote the code, and conducted the spatial
- 14 interpolation. XS wrote the first draft of the manuscript and both authors reviewed and revised

15 the manuscript.

# 36 Abstract:

Quantifying and predicting precipitation and water flow and their influences is challenged by the 37 dynamic relationships between and timing of precipitation and water fluxes. To help with these 38 challenges, scientists use "water year" to examine and predict the impacts of precipitation and 39 relevant extreme climatic and hydrological events on ecosystems. However, traditional water 40 year definitions used in the U.S. lack a consideration of areal variation in climate and hydrology, 41 which is needed when studying ecosystems at regional or national scales. We developed local 42 water year (LWY) values that consider spatial variation using existing definitions whereby the 43 44 water year begins in the month with the lowest or highest average monthly streamflow. We employed spatial interpolation to assign LWY start and end months to 202 subregions across the 45 conterminous U.S. that range from 4,384 to 134,755 km<sup>2</sup>. This dataset can be linked with diverse 46 47 climate, terrestrial, and aquatic data for broad-scale studies.

48

49 Keywords: Water year; Streamflow; Spatial variation; Spatial interpolation; Macroscale;

50 Precipitation

51

#### 52 Background & Motivation

Precipitation plays a crucial role in shaping hydrology and ecosystems. Precipitation can 53 be highly spatially variable, especially when considering macroscale spatial extents of regions to 54 continents (Mock 1996; Augustine 2010). As a result of climate change, there has been greater 55 inter-annual variability in precipitation in many regions worldwide (IPCC 2021), causing 56 57 increased frequency and intensity of extreme climatic and hydrological events such as drought and flooding (Easterling et al. 2000; Grimm & Natori 2006; Prein et al. 2016; Kundzewicz et al. 58 2020). In addition, water fluxes are sometimes asynchronous with precipitation and extreme 59 events can occur during the transition between years, complicating hydrological estimations. For 60 example, rainfall in late fall can be retained in the soil and influence water fluxes the following 61 spring (Pike 1964; Kamps & Heilman 2018). Despite such spatial and temporal variation and 62 asynchronicity in precipitation, a calendar year timeframe (from January 1st to December 31st) 63 has often been used to examine and predict the impacts of precipitation and relevant extreme 64 65 climatic and hydrological events on aquatic systems.

To more accurately predict water flow, researchers adopted a "water year" that usually 66 spans two standard calendar years. For example, the U.S. Geological Survey(USGS) water year, 67 68 which was adopted a century ago, starts on October 1st and ends on September 30th of the next year (Henshaw et al. 1915). This USGS water year is applied to the whole U.S. and intends to 69 account for the influence of snowfall from October to December on the next year's streamflow 70 71 (Henshaw et al. 1915). However, different regions of the U.S. have different timing of precipitation (including snowfall) and hydrology, as well as varying topographic patterns, all of 72 73 which affect relationships between (and timing of) precipitation and water fluxes (Nicótina et al. 2008; Condon & Maxwell 2015; Torre Zaffaroni et al. 2023). These facts mean that a more
localized timeframe is needed rather than applying a single water year to the macroscale.

Researchers have started to use various definitions for water year, and subsequent 76 start/end times of that water year depend on the locations, ecosystem types, and research 77 purposes or questions. For example, Olson et al. (2013) started their water year in April when 78 79 analyzing the methane and carbon dioxide fluxes of a temperate peatland, Kamps and Heilman (2018) started their water year in September to match annual precipitation with water and carbon 80 budgets in Central Texas, and Caruso (2000) started water years from July or October so that the 81 82 low-streamflow periods in the Otago region in New Zealand could be fully captured. These (and other) studies use water years for a relatively local spatial extent (e.g., watershed or single region 83 ). However, organisms and ecological processes are influenced by multi-scale factors, from local 84 (e.g., lake morphometry) to regional (e.g., land use) and macroscale (e.g., climate), and these 85 factors can sometimes interact to affect ecosystems (Heffernan et al. 2014; Rose et al. 2017; 86 87 LaRue et al. 2021). Thus, it is crucial to investigate and predict how ecosystems respond to environmental changes, such as precipitation variability and relevant extreme events, across 88 multiple spatial and temporal scales. Therefore, localized water year timeframes are needed for a 89 90 range of research purposes at regional to continental scales.

One challenge to creating localized water year timeframes has been limited data for variables such as snow melting time, ice-off dates, and annual gross primary productivity (e.g., Olson et al. 2013, Kamps & Heilman 2018). However, Wasko et al. (2020) proposed a climateand hydrology-relevant local water year (LWY) timeframe that solely used streamflow data that are available for most areas globally. This LWY provides a site-specific timeframe beginning in the month with the lowest average monthly streamflow to capture the concurrent and lagged associations between precipitation and hydrology (Wasko et al. 2020). Using this localized
timeframe, they predicted the timing and trends of flooding and streamflow at the global scale
and demonstrated an improved accuracy of estimation compared with using a calendar year
timeframe (Wasko et al. 2020).

101 Although a big step forward for global studies relying on water year data, these data do 102 not completely cover the conterminous U.S. (CONUS), which may limit regional to CONUSscale research. Thus, we build on their work by extending this local water year timeframe to 103 cover the CONUS. We used recent (1990 to 2018) streamflow data, the same method used by 104 105 Wasko et al. (2020), and a spatial interpolation method to construct a CONUS-scale LWY timeframe. In addition, given that the most appropriate definition of water year varies depending 106 107 on research purposes, we applied the same process and generated a second LWY timeframe 108 starting from the month with the highest average monthly streamflow, an approach often used in studies of low streamflow and hydrological drought (e.g., Caruso 2000; Chagas et al. 2024). To 109 create these LWYs, we used subregions that were created based on the drainage features by the 110 USGS (Seaber et al. 2007). This hierarchical regionalization framework divides and subdivides 111 the U.S. into successively smaller hydrologic units (HUs); we used the HU4 subregion, which is 112 113 the second-level classification that delineates large river basins (USGS 2024). We included 202 HU4s in the CONUS that range in area from 4,384 to 134,755 km<sup>2</sup>. These LWY data will help 114 advance the understanding of the impacts of variability in precipitation and streamflow on inland 115 116 waters at the regional and national U.S. scales.

117

118 **Data Description** 

This data product consists of two datasets housed on the Environmental Data Initiative 119 (EDI) repository (https://doi.org/10.6073/pasta/c27e57749f856bd24dc7c7559b9b316b), as well 120 as our R code (file name: local water year code.R). The first dataset (file name: water years 875 121 sites.csv) was used to develop and evaluate the LWY end month across the CONUS. This file 122 includes the identifier from the Global Runoff Data Centre, which is the archive where we 123 124 obtained streamflow data ("grdc no" column), end month of the LWY that begins from the month with the lowest average monthly streamflow ("end.month.lowest" column), end month of 125 the LWY that begins from the month with the highest average monthly streamflow 126 ("end.month.highest" column), and locational information (longitude ("lon" column), latitude 127 ("lat" column), altitude ("altitude" column), and the name of the river ("river" column) and 128 station ("station" column) of each gauging site where the daily streamflow data were measured) 129 (the process of site selection will be described in the next section). There are a total of 875 sites. 130 When the LWY starts from the lowest-flow month, the most common LWY end month among 131 132 these sites is July (228 sites), followed by August (219 sites), December (160 sites), and September (78 sites) (Figure 1a & 2a). When the LWY starts from the highest-flow month, the 133 most common end month is April (214 sites), followed by February (175 sites), March (145 134 135 sites), and May (134 sites) (Figure 3a & 4a).

The second dataset (file name: hu4 water years with notes.csv) contains the local water year data for each subregion. This file includes the start ("start.month.lowest" column) and end month ("end.month.lowest" column) of the LWY that begins from the month with the lowest average monthly streamflow as well as the start ("start.month.highest" column) and end month ("end.month.highest" column) of the LWY that begins from the month with the highest average monthly streamflow for each of the 202 subregions (i.e., 4-digit Hydrologic Unit; HU4) across

the CONUS ("hu4.code" column). This dataset also includes two "notes" columns 142 ("notes.lowest" and "notes.highest") that provide details about whether there were streamflow 143 data in the subregion and the method we used to determine the LWY end month. There are three 144 categories in this column: 1) dominant interpolation, which indicates that there were streamflow 145 data and we based the end month on the single, dominant interpolated LWY end month value in 146 147 the subregion; 2) local sites based, which indicates that there were streamflow data and multiple LWY end month values in the subregion; therefore, end month was based on the site-specific 148 LWY data; and 3) ND interpolation, which indicates that there were no streamflow data in the 149 150 area and the end month was determined based on the dominant LWY end month value from interpolation of nearest sites. More details about the methodology can be found in sections 3 and 151 4. 152

The results of this work are 404 LWYs, two for each subregion across the CONUS. 153 There are spatial differences in LWY end months (Figure 2b & 4b). For example, along the 154 eastern and western edges of the CONUS, the LWY that begins from the month with the lowest 155 streamflow (hereafter referred to as LWY-lowest) usually ends in July or August, except for the 156 very southeast where it ends in April. In contrast, there is much more heterogeneity in LWY-157 158 lowest in the central U.S., with November and December being the most common end months. The most common end month for LWY-lowest is August (64 subregions), followed by July (49 159 subregions), December (34 subregions), and November (16 subregions) among all subregions 160 161 (Figure 1b). When LWY starts from the highest-flow month (hereafter referred to as LWYhighest), the end months are often February and March in the eastern U.S. and April and May in 162 163 the central U.S., with more heterogeneity in the western U.S. Among all subregions, the most

164 common LWY-highest end month is February (48 subregions), followed by April (42

subregions), March (39 subregions), and May (33 subregions) (Figure 3b).

166

167 Methods

We first generated subregion-specific LWYs based on the definition and method proposed by Wasko et al. (2020) (i.e., LWY-lowest), in combination with daily streamflow data and spatial interpolation. Then, using the same process, we generated a second LWY timeframe based on a different definition (i.e., LWY-highest). Data processing was performed in R (R Core Team 2024).

We used daily streamflow data from the Global Runoff Data Centre (GRDC; GRDC 173 2023) to calculate the monthly streamflow at each site (i.e., river gauge station). The GRDC is an 174 open-access archive of international data that has been widely used in regional, multinational, 175 and global hydrological studies (e.g., Hong et al. 2007; Wasko et al. 2021; Brunner & Slater 176 2022). We first downloaded data from 1990 to 2018 for the CONUS. Then, we filtered the 177 streamflow data for sites that met four criteria to avoid big missing gaps in data, to take into 178 account potential inter-annual variation in streamflow features, and to ensure that the data are 179 180 relatively 'recent': 1) with at least 10 years of data, 2) with at least eight months of data from at least half of the years, 3) average daily streamflow data missing rate was  $\leq 80\%$  across all the 181 years, and 4) the last year of data is post-2000. This process resulted in 875 sites spread across 182 183 the CONUS.

184 *LWY beginning with the lowest streamflow month (LWY-lowest)* 

For each site, we calculated the average monthly streamflow data and compared these monthly averages to determine the month with the lowest streamflow. This month became the

start month of a site's LWY (i.e., each site has its own lowest-streamflow-month, which is the 187 start month of an LWY; Wasko et al. 2020). Interestingly, although a previous study suggested 188 that low-river-flow timing in some European and U.S. regions exhibit slight inter-annual 189 variation (Floriancic et al. 2021), the end month of the LWY-lowest was consistent from 1990 to 190 191 2018 across all the sites in our dataset. 192 We then applied ordinary kriging to interpolate site (river gaging station) LWY-lowest end month data (months as integers, 1 through 12) to the whole CONUS using gstat (v2.1-1, 193 Pebesma & Graeler 2023) and raster (v3.6-23, Hijmans et al. 2023) R packages. Ordinary kriging 194 195 (OK) is a geostatistical technique commonly used to interpolate and map data for unsampled locations and areas (e.g., Sanabria et al. 2013; Boudibi et al. 2019; Li et al. 2023). OK generally 196 involves three steps: computing the semivariogram, defining a semivariogram model, and 197 interpolating based on the semivariogram model (Gimond 2023). 198

199 We computed the semivariogram, which depicts the spatial correlation between the 200 neighboring values, using equation (1),

201 
$$\gamma(h) = \frac{1}{2n} \sum_{i=1}^{n} [Z(x_i) - Z(x_i + h)]^2$$
 (1)

where  $\gamma(h)$  is the semivariogram; Z(x<sub>i</sub>) and Z(x<sub>i</sub> + h) are the data at locations x<sub>i</sub> and x<sub>i</sub> + h, respectively; and n is the number of pairs of data separated by distance h (Li & Heap 2011; Sanabria et al. 2013). Second, we fit a mathematical model to the semivariogram. The spherical function was used in our model, and we adjusted parameter values (e.g., partial sill, range, and nugget) to improve the model fit. Third, we applied this semivariogram model to interpolate the LWY-lowest end-month data, by using equations (2) and (3) to estimate the local data (at the unsampled location) using neighboring data,

209 
$$Z^*(x_0) = \sum_{i=1}^n \lambda_i Z(x_i)$$
 (2)

210  $var\{Z^*(x_0) - Z(x_0)\} = minimum$  (3)

where  $Z^*(x_0)$  is the estimated value at location  $x_0$ ;  $Z(x_i)$  is the data value at location  $x_i$ ; and  $\lambda_i$  is the weighting factor that is determined by minimizing the variance (equation 3). Finally, we overlaid the interpolated end month LWY-lowest values with subregion polygons for the CONUS to assign the LWY-lowest end month for each subregion.

The resulting 202 subregion LWY-lowest values include 156 subregions that were 215 labeled "dominant interpolation" in the dataset (hu4 water year with notes.csv, "notes.lowest" 216 column). These subregions had streamflow data and a single dominant interpolated LWY-lowest 217 end month in the subregion, so the end month was chosen based on the dominant value. There 218 were 29 subregions labeled as "local sites based", which indicates that there were streamflow 219 data but multiple different LWY-lowest end month values in the subregion. It was difficult to 220 determine the dominant month of these subregions based on interpolation results, thus the 221 decision of the LWY-lowest end month of the subregion was made by checking the site-specific 222 223 LWY data in each of the subregions and determining the dominant month of each subregion. For the 17 subregions labeled "ND interpolation", there was no streamflow data and the month was 224 225 determined solely based on interpolation results and the dominant interpolated LWY-lowest end 226 month value.

227 *LWY beginning with the highest streamflow month (LWY-highest)* 

For each site, we used the calculated average monthly streamflow data and compared these monthly averages to determine the month with the highest streamflow, which became the start month of a site's LWY. When the highest-streamflow-month varied among years for a site, the start month of a site's LWY-highest was the highest-streamflow-month with the highest frequency of occurrence. Then, following the processes described above, we applied spatial interpolation and obtained 202 subregion LWY-highest values, including 178 subregions that
were labeled "dominant\_interpolation" in the dataset (hu4 water year with notes.csv,
"notes.highest" column), seven subregions labeled as "local\_sites\_based", and 17 subregions
labeled "ND\_interpolation" (without streamflow data).

237

### 238 **Technical Validation**

We assessed the performance of the spatial interpolation method using a leave-one-out 239 cross-validation approach (Sanabria et al. 2013). Firstly, we randomly chose a site (i.e., river 240 241 gaging station) and removed its LWY data from the dataset. Then, we applied the ordinary kriging method described above to the new dataset, re-estimated the LWY end month of the 242 removed site, and compared the new estimated LWY end month value with the actual end month. 243 We repeated this process 10 times on 10 different, spatially-separated sites. For both LWY 244 definitions, we found that the estimated and the actual end month of these 10 sites were either the 245 same or differed by one month (mean absolute difference = 0.4 months for LWY-lowest and 0.5 246 months for LWY-highest), depending on the streamflow data density of the subregion. 247 Subregions with a higher data density had higher accuracies than areas with a lower density of 248 249 data.

250

## 251 Data Use and Recommendations for Reuse

This local water year dataset is intended to provide localized, continental-scale water year timeframes that can be used for studying the features and impacts of precipitation and hydrology across the CONUS. It is important to note that precipitation and hydrological dynamics and patterns can vary by the water year definition (e.g., Figure 5). Using the Little Fork River in

Minnesota (USA) as an example, if a researcher was studying the peak streamflow in April 2001, 256 it would be in water year 2002 when using either LWY-lowest or LWY-highest, but in water 257 year 2001 when using the water year created by USGS (Oct 1st - Sep 30th). Thus, it is crucial to 258 choose a water year definition that matches the context and research question being asked. The 259 LWY-lowest can be useful for studying the relationship between precipitation and runoff and 260 261 local long-term hydrological cycles (e.g., water replenishment and depletion cycle). The LWYhighest can provide more relevant insights for research focused on dry or low-flow periods 262 because it covers the entire low-streamflow period. Finally, in some cases, an alternative 263 definition could be more useful. For example, the USGS water year definition that starts from 264 October 1st might be appropriate for hydrological studies in snow-dominated regions. 265

Here, we provide an example of using these three different timeframes to identify the 266 water years with the lowest and highest annual average streamflow from water years 1991-2018. 267 We assigned each of the 875 sites two LWY end months (LWY-lowest and LWY-highest) 268 according to the subregion they are located in (i.e., all the sites in the same subregion share the 269 same end month; Sun & Cheruvelil 2024), calculated the annual average streamflows of each site 270 based on these two LWY timeframes, and then determined the water years with the highest and 271 272 lowest streamflow for each site and timeframe. Then, we calculated the site-specific annual average streamflow based on the water year that ends on September 30th (USGS) and 273 determined the water years with the highest and lowest streamflow for each site using that 274 275 definition of water year. Finally, we compared the highest and lowest streamflow water years between LWY-lowest and the Oct-Sep water year (by USGS) and between LWY-highest and the 276 277 Oct-Sep water year (by USGS). For all the three definitions, the water year was named by the 278 calendar year in which it ended (e.g., the 12-month period from August 1st, 2010 to July 31st,

279	2011 = LWY 2011). We found that, for some sites across the CONUS, the water years with the
280	lowest and highest annual average streamflow were consistent across water year definitions,
281	while for others, these years varied (Figure 6). Additionally, the comparison between LWY-
282	highest and USGS Oct-Sep water year definition resulted in more sites with different highest and
283	lowest streamflow years (Figure 6 b&d 75% and 68% were different for the highest and lowest
284	streamflow years, respectively) than the comparison between LWY-lowest and Oct-Sep water
285	year created by USGS (Figure 6 a&c 25% and 24% for the highest and lowest streamflow years,
286	respectively). These results suggest that the water year definition can influence the identification
287	of extreme streamflow events and highlight the importance of selecting appropriate definitions.
288	This LWY dataset considers areal variations and can be used in various meteorological,
289	hydrological, and ecological studies to identify and predict trends in precipitation, extreme
290	events (drought and flooding), and water fluxes as well as investigate their effects on ecosystems
291	(e.g., Kamps & Heilman 2018) and human communities (e.g., calculating hydropower generation
292	capacity; Bongio et al. 2016). Future users of the subregion-specific LWYs can combine these
293	data with a wide range of climatic, as well as terrestrial and aquatic abiotic and biotic data, by
294	linking our dataset with other data products, such as LAGOS-US modules (e.g., Cheruvelil et al.
295	2021) or USGS datasets (e.g., Blodgett 2023) using subregion identifiers (i.e., HU4 codes).
296	Moreover, our R code is available for download at the EDI repository so that users can apply a
297	similar method to other regions around the world to generate site or region-specific LWY
298	timeframes. As such, these data will be a valuable addition to the literature that can contribute to
299	building macroscale understanding of precipitation and streamflow variability and their
300	influences on a variety of systems.

302

303 **Conflict of interest**: The authors declare no conflict of interest.

304

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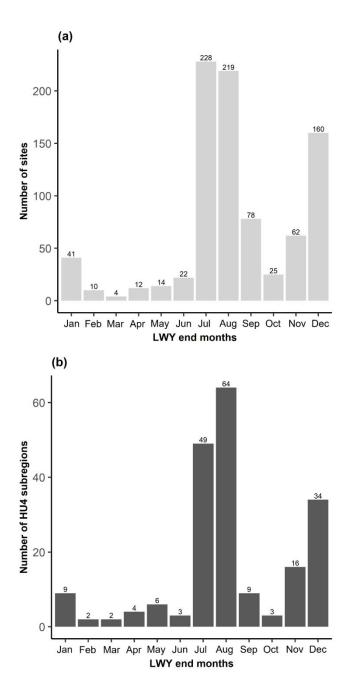


Figure 1. The number of stream gauging sites (a) and subregions (b) by the end month of the local water year (LWY) that starts from the month with the lowest average monthly streamflow (LWY-lowest). The numbers above each bar indicate the number of sites (top) or subregions (bottom). Subregion = HU4 (Seaber et al. 2007). More information about HU4s can be found on the USGS website: <u>https://water.usgs.gov/GIS/huc.html</u>, last accessed September 2023.

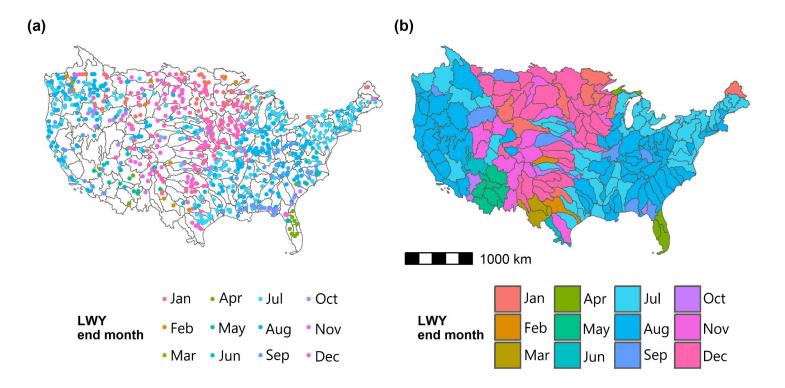


Figure 2. Maps showing the LWY-lowest end month of each site overlaid with subregion (HU4) polygons (a) and the end month for each subregion (b). The LWY starts from the month with the lowest average monthly streamflow. In plot (a), there are 17 subregions without streamflow data. Colors represent the end months.

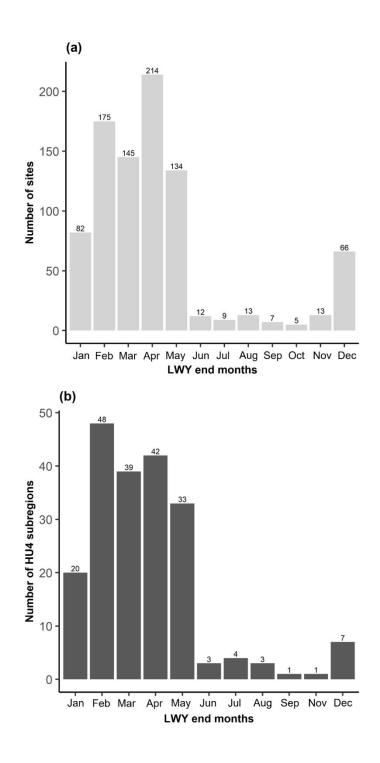


Figure 3. The number of stream gauging sites (a) and subregions (b) by the end month of the local water year (LWY) that starts from the month with the highest average monthly streamflow (LWY-highest). The numbers above each bar indicate the number of sites (top) or subregions (HU4; bottom).

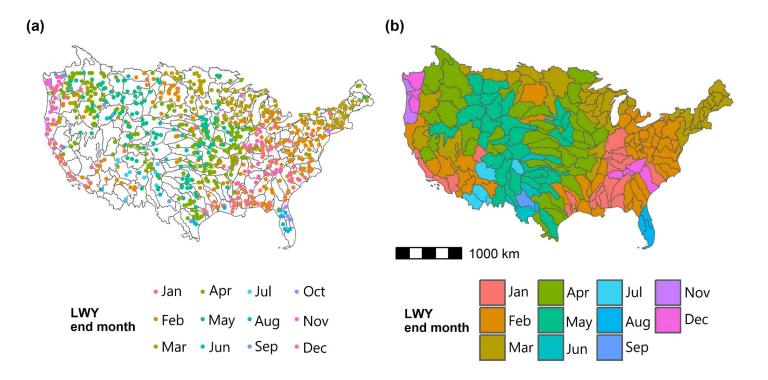


Figure 4. Maps showing the LWY-highest end month of each site overlaid with subregion (HU4) polygons (a) and the end month for each subregion (b). The LWY starts from the month with the lowest average monthly streamflow. Colors represent the end months.

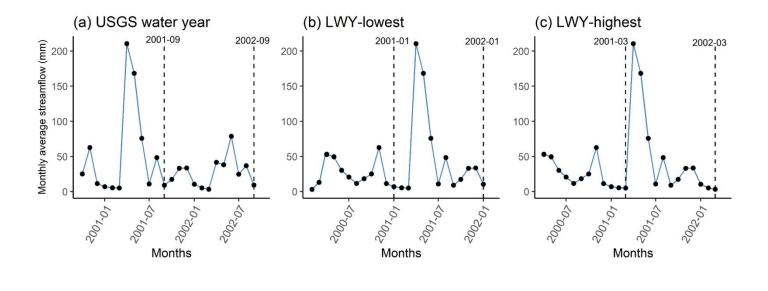


Figure 5. Monthly average streamflow data of Little Fork River (Minnesota, USA,
latitude=48.3958, longitude=-93.5493) in the water years 2001 and 2002 using three LWY
definitions: Oct 1st - Sep 30th water year (sensu USGS) (a), LWY-lowest (b), and LWY-highest
(c). The dashed lines indicate the end month of the water years.

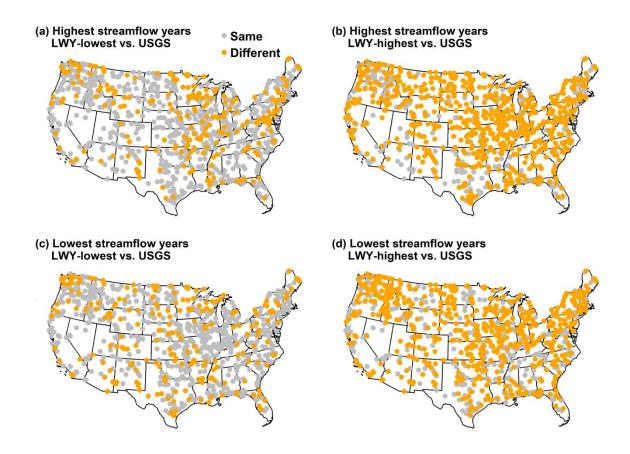


Figure 6. Maps showing the comparison of the highest/lowest streamflow water years across the three water year definitions. Plots (a) and (b) show whether the year with the highest streamflow was the same or different for each site when using different water year definitions. Plot (a) compares LWY-lowest with the Oct - Sep (USGS) water year definition, while plot (b) compares LWY-highest with the Oct - Sep water year. Plots (c) and (d) show whether the year with the lowest streamflow was the same or different for each site when using different water year definitions. Plot (c) compares LWY-lowest with the Oct - Sep water year. Sep water year definition, while plot (d) compares LWY-highest with the Oct - Sep water year. Grey dots indicate that the years were the same between the definitions, and orange dots indicate that the years were different.