

1 **Taphonomic Controls on a Multi-Element Skeletal Fossil Record**

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31 **Abstract**

32 Animals with multi-element skeletons, including the vertebrates, echinoderms, and
33 arthropods, are some of the most biodiverse and ecologically important animal groups.
34 Understanding the relative impact of the myriad geological and biological factors which impact
35 on the quality of multi-element skeletal fossils is thus crucial for disentangling perceived
36 changes in biodiversity through time and shedding light on gaps in the fossil record. We have
37 characterised the specimen-level taphonomic history of virtually the entire Palaeozoic fossil
38 record of echinoids, the class of echinoderms which includes the living sea urchins. We find
39 that the majority of this fossil record consists of disarticulated skeletal elements and as
40 preservational quality increases, so does the proportion of specimens which can be identified
41 with higher taxonomic precision. We then assessed the relative impacts of multiple geological
42 and biological factors on our specimen quality, identifying that fine-grained host matrix, as well
43 as siliciclastic lithology are the biggest factors in determining quality of fossil echinoid
44 preservation, while aspects intrinsic to specific taxonomic groups also play a role. Differential
45 sampling of the fossil record seems to play little role in influencing the distribution of
46 taphonomic grades, and fluctuations in the North American record of Siliciclastic rocks are
47 positively correlated with fluctuations in taphonomic grades. Our results highlight that the
48 factors controlling the animal macrofossil record are varied, and that the interplay between
49 these variables, taphonomic grade, and taxonomic precision impacts on our ability to use the
50 fossil record to understand macroevolution.

51 **Key words:** Taphonomy, Rock Record, Echinoidea, Echinodermata, Palaeozoic

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53 Not all fossils are preserved equally. This variable preservation impacts our ability to
54 accurately comprehend the ancient past: for example, a single disarticulated tooth and a
55 complete, articulated skeleton provide different amounts of data about the ecology, diversity,
56 and evolutionary history of *Tyrannosaurus rex*. Animals with multi-element skeletons such as

57 vertebrates, echinoderms, and arthropods, are abundant in the fossil record and are a
58 treasure trove of information about how macroevolutionary processes in deep time underlie
59 the origins of modern biodiversity (Legg *et al.* 2013; Mongiardino Koch and Thompson 2021;
60 Kuo *et al.* 2023; Hopkins and Smith 2015). Multi-element skeletons consist of hundreds to
61 thousands of variably-sized parts, which disarticulated rapidly upon death, except in
62 exceptional circumstances (Donovan 1991; Nebelsick 1996). Understanding the controls on
63 differential preservation of multi-element skeletons is thus crucial for our understanding of
64 deep-time biodiversity change.

65 Differential quality in fossil preservation has important influences on our perception of
66 ancient biodiversity patterns, as well as our ability to accurately reconstruct evolutionary
67 histories using these fossils (Smith, 2007; Mannion & Upchurch, 2010; Dean *et al.* 2016).
68 The varied environmental, biological, and geological processes which control the quality of
69 fossil preservation are thus critical for an accurate perception of ancient biodiversity
70 (Nebelsick 1996; Hendy 2011; Allison and Bottjer 2011; Smith *et al.* 2001; Butts and Briggs
71 2011; Nanglu and Cullen 2023; Dean *et al.* 2019). The majority of taphonomic studies which
72 seek to assess the role of these myriad geological and biological controls on fossil
73 preservation focus in exceptional detail on particular localities or basins, interrogating the
74 respective roles of local processes such as depositional, geochemical and diagenetic
75 conditions (Anderson *et al.* 2021; Briggs *et al.* 1996; von Bitter *et al.* 2007). However, in
76 addition to these local signals, broader “megabiases” are also known to influence the fossil
77 record; these include factors such as the age of fossil-bearing deposits (Signor *et al.* 1982;
78 Dean *et al.* 2020), anatomical differences within and between taxonomic groups that
79 influence preservation (Butts and Briggs 2011; Cherns and Wright 2000; Wright *et al.* 2003;
80 Dean, *et al.* 2019; Kowalewski and Flessa 1996), preferential preservation within particular
81 lithologies (Hendy 2011; Cleary *et al.* 2015), spatial variation in the quality of fossil
82 preservation (Dean *et al.* 2019) or spatial sampling heterogeneity more generally (Antell *et*
83 *al.* 2024; Benson *et al.* 2016; Close *et al.* 2020; Benson *et al.* 2021; Jones *et al.* 2021). While

84 local environment and large-scale factors undoubtedly both influence fossil preservation,
85 their impact on both the preservation and evolutionary history of specific taxonomic groups is
86 more difficult to ascertain, due to the difficulty of collecting relevant specimen-level
87 information across the entire fossil record of a group of organisms. Subsequently, the
88 relative importance of these different factors on influencing fossil preservation, and
89 subsequently our understanding of observed fluctuations in palaeodiversity, remains poorly
90 constrained.

91 To characterise global and long-term trends in fossil preservation, and better
92 understand the biological and geological factors which influence fossil preservation
93 throughout deep-time, we herein assess the taphonomic quality of the entire known
94 Palaeozoic fossil record of an invertebrate fossil group with multi-element skeletons.
95 Echinoids are a class of echinoderms which includes the living sea urchins, sand dollars,
96 and heart urchins, and have a fossil record dating back to the Ordovician period (Thompson
97 *et al.* 2022). In contrast to the post-Palaeozoic echinoids, which are renowned for their
98 exceptionally well-sampled fossil record (Foote and Sepkoski 1999), most Palaeozoic
99 echinoid taxa lacked the microstructural interlocking of skeletal plates found in post-
100 Palaeozoic echinoids (Smith 1980; Grun and Nebelsick 2018). This made them prone to
101 disarticulation following death and thus the fossil record of Palaeozoic echinoids exhibits a
102 wide range in preservational quality, from single disarticulated plates (Thompson and
103 Denayer 2017), to articulated individuals preserving most or all of the original skeleton
104 (Schneider *et al.* 2005). Furthermore, while post-Palaeozoic echinoids were abundant
105 members of Mesozoic and Cenozoic ecosystems, their Palaeozoic precursors were
106 comparatively rarer (Schneider 2008), meaning that a near-exhaustive sampling of their
107 occurrence in museum collections is feasible. This variation in preservational state and
108 tractability of building a large near-exhaustive database makes Palaeozoic echinoids an
109 excellent model system to identify the differential importance of factors controlling
110 preservational quality in the fossil record.

111 Building on previous work (Thompson & Bottjer, 2019, Thompson *et al.* 2022) we
112 have produced the most complete specimen-level taphonomic dataset of a temporally and
113 geographically widespread Palaeozoic fossil invertebrate group to date, containing virtually
114 every record of Palaeozoic echinoids (3447 specimens total) collected from the general
115 literature and thirty-three museums around the globe. Each specimen is designated with a
116 taphonomic grade, and is associated with lithological, environmental and geographic
117 information. We then use a variety of statistical and modelling procedures to establish the
118 relative influence of various geological, geographical and biological controls on
119 preservational state through space and time. Lastly, we use this information to refine our
120 understanding of the critical early evolutionary history of echinoids, and the relative impacts
121 of environmental variables on fossil preservation more broadly.

122 **Material and Method**

123 **Fossil dataset**

124 To quantify the preservation of the Palaeozoic echinoid fossil record, we devised a semi-
125 quantitative index to describe the taphonomic grade (TG) of fossil specimens from museum
126 collections and the literature. We visited thirty-three museum collections and, where
127 possible, photographed each specimen and assigned it to one of five taphonomic grades
128 (TG1-5) based upon its completeness (Fig. 1A). These taphonomic grades mirror series of
129 disarticulation and disintegration identified from experiments with living echinoids (Kidwell
130 and Baumiller 1990; Greenstein 1991), and modified for the distinct body plans of
131 Palaeozoic echinoids which lack stereomic interlocking. The lowest taphonomic grade (TG1)
132 represents single, disarticulated skeletal ossicles (including lots in collections of multiple
133 ossicles which we treat conservatively to represent a single disarticulated specimen) while
134 the second lowest grade (TG2) represents disarticulated but associated ossicles. These
135 lowest two taphonomic grades usually consist of disarticulated coronal plates and spines,
136 though more rarely include elements of the Aristotle's lantern. TG3 is the middle taphonomic

137 grade, and represents fragments of articulated or semi-articulated test. TG4 represents
138 articulated tests which have become denuded of spines and represents minimal degradation
139 post-mortem (Kidwell and Baumiller 1990). Finally our highest taphonomic grade, TG5,
140 consists of articulated tests with attached spines, representing highest quality preservation.

141 By basing our analyses on specific specimens from museums and the literature, we
142 are able to link geological variables of the rock matrix, such as grain size and lithology
143 directly with the fossils they entomb. As such, following Thompson and Bottjer (2019) the
144 associated lithology and grain size were recorded for each specimen, alongside data on
145 taxonomy (including most-specific taxonomic rank possible), stratigraphic provenance and
146 locality of collection. These data were further bolstered with data collected from the general
147 literature, as well as all available stratigraphic, taxonomic, lithological and taphonomic data
148 available for each published specimen. All data were collected by one individual (J.R.
149 Thompson), who also vetted all the taxonomic designation of all specimens; as such, there is
150 consistency in approach across the dataset, reducing a potential source of bias (Wagner *et*
151 *al.* 2007). Specimens without precise geographic co-ordinates were georeferenced by
152 locating text-based locality descriptions on Google Maps (available from
153 <http://google.com/maps>); we additionally included the likely precision of this estimate in the
154 degrees-minutes-seconds format. Palaeo-coordinates for these locations were obtained
155 using the 'PALAEOMAP' rotation model through the GPlates API, using the 'palaeorotate'
156 function of the 'palaeoverse' package (Jones *et al.*, 2023). Data were additionally assigned
157 to stage, series and period level time bins using the 'majority' method of the 'bin_time'
158 function from the 'palaeoverse' package (Jones *et al.*, 2023). Our final dataset utilised for
159 analysis contains 3447 unique specimens across 37 countries, and spans approximately the
160 first 300 million years (Ordovician to Permian) of the evolutionary history of echinoids. It is
161 available as supplementary information at
162 https://github.com/ChristopherDavidDean/Echino_taph.

163 **Multiple logistic regression**

164 To assess the primary controls on the differential fossil preservation seen in our dataset, we
165 used multiple logistic regression, coding each TG as a binary response variable and
166 including lithology (carbonate vs. siliciclastic), grain size (fine-grained vs. coarse grained),
167 age (midpoint of the stage to which the specimen belonged), latitude of the locality of
168 collection, and palaeolatitude as predictor variables. Due to low numbers of occurrences for
169 various families across taphonomic grades, we did not include taxonomic information as a
170 response variable. We are primarily interested in reporting effect sizes, which are gauged by
171 the magnitude of regression coefficients. When interpreting results, note that the chosen
172 taphonomic grade is compared against all other grades (e.g. specimens with a taphonomic
173 grade of 5 [1], vs. all other taphonomic grades [0]) so that positive regression coefficients
174 indicate greater odds of sampling the specific taphonomic grade of choice, and log-odds
175 below 0 indicate reduced odds of sampling. For binary covariates, we report the log odds of
176 one level relative to the other (for instance, fine grained vs. coarse grained), so positive
177 regression coefficients represent greater odds of the named factor influencing the presence
178 of a particular taphonomic grade. An interaction term consisting of a combination of lithology
179 and grain size was additionally added to test for interactions between the covariates. We
180 used the 'dredge' function of the package 'MuMin' (Bartoń 2023) to obtain information about
181 the full model set and the small sample size corrected Akaike Information Criterion (AICc) for
182 final model selection, with lower scores indicating a better model fit (Burnham and Anderson
183 2001). Here we present models with the lowest AICc scores, and only report factors with
184 statistically significant results ($p < 0.05$); full results can be found within Supplementary
185 Information 2 at https://github.com/ChristopherDavidDean/Echino_taph.

186 Statistical Tests

187 We used Pearson's chi-squared tests to establish whether there were statistically
188 significant differences between taphonomic grades with respect to lithology, grain size and
189 time intervals. Mosaic plots of these results were produced using the 'mosaic' function of the
190 'vcd' package (Meyer *et al.* 2006). To test whether preservational state showed a correlation

191 with specific variables through time, we used Spearman's rank correlation tests corrected for
192 large sample sizes with the Benjamini Hochberg procedure (Benjamini and Hochberg 1995)
193 to characterise the strength of correlations between the abundance of taphonomic grades,
194 the available rock record, sampling proxies and genus-level diversity at the stage level. The
195 abundance of carbonate and siliciclastic rocks in North America was obtained from
196 Macrostrat (Peters *et al.* 2018), a database of geological data which includes compilations of
197 the spatial and temporal distribution of North American sedimentary rocks, and was
198 compared to the abundance of North American occurrences (downgraded from the
199 specimen level dataset) of each taphonomic grade from our dataset. This provided a picture
200 of the role of the rock record on the abundance of our preservational states. We used the
201 number of Palaeozoic echinoid, Echinoderm, and invertebrate collections from the
202 Paleobiology Database (PBDB; <https://paleobiodb.org/>) as potential general sampling
203 proxies; we used the latter two as proxies approaching a 'global' sampling signal to reduce
204 the possibility of redundancy (Benton *et al.* 2011). These collections were obtained from
205 PBDB downloads. Species and period level correlations can be found in Supplementary
206 Information 2 at https://github.com/ChristopherDavidDean/Echino_taph.

207 All statistical tests were implemented in R, and lithological data from Macrostrat was
208 obtained in R via the Macrostrat API. Counts for global invertebrate collections were
209 obtained by downloading all Palaeozoic invertebrate records from the PBDB (on
210 08/11/2022). Genus level diversity for echinoids was established using the 'binstat' function
211 of the divDyn package (Kocsis *et al.* 2019). Associated R scripts can be found in
212 Supplementary Information at https://github.com/ChristopherDavidDean/Echino_taph.

213

214 **RESULTS**

215 Differences in quality of fossil preservation

216 Our analyses highlight striking differences in the quality of fossil preservation across
217 the Palaeozoic history of echinoids. The majority of known echinoid specimens are
218 disarticulated (Fig. 1C), with the taphonomic grades in our dataset representing
219 disarticulated remains equating to 61% of the total dataset (TG1: 38%; TG2: 23%). These
220 are twice as abundant than TG4 and TG5, which represent those specimens which are best-
221 preserved (TG4: 14%; TG5: 15%).

222 Across all taphonomic grades Palaeozoic echinoid fossils are significantly more
223 abundant in fine-grained than coarse-grained rocks (Figure 2B,D; X-squared = 159.08, df =
224 4, p-value < 2.2e-16). Mosaic plots reveals that in particular taphonomic grades TG4 and
225 TG5 are noticeably less common than expected in coarse-grained rocks, whereas
226 disarticulated remains (TG1 and TG2) occur more likely than expected in coarse-grained
227 rocks (Fig. 2D). Echinoid fossils are also significantly more abundant in carbonate rocks than
228 siliciclastics (Figure 2A,C; X-squared = 182.6, df = 4, p-value < 2.2e-16). Disarticulated
229 plates (TG1 and TG2) occur less frequently than expected in siliciclastic rocks, whereas
230 individuals with the highest taphonomic grade (TG5) occur more frequently than expected in
231 siliciclastic rocks and less than expected in carbonates (Fig. 2C).

232 Taphonomic grades are not distributed evenly with respect to the lowest taxonomic
233 rank to which a specimen can be assigned (Fig. 3A; X-squared = 1540.9, df = 12, p-value <
234 2.2e-16). At TG1, a higher proportion of specimens are only assignable at the class level
235 (Echinoidea) than at any other taphonomic grades. At this grade, specimens assignable to
236 the species level occur less frequently than expected, while specimens which can be
237 assigned to higher taxonomic levels occur more frequently than expected (Fig. 3C). As
238 preservational score increases, so too does the proportion of specimens assignable to a
239 lower taxonomic rank, with specimens only assignable to higher taxonomic ranks occurring
240 less-frequently than expected. Likewise specimens which can be attributed to the species-
241 level occur more frequently than expected at TG3-TG5 (Fig. 3C).

242 The proportion of different preservational grades also varies temporally (Fig. 3B,D; X-
243 squared = 875.85, df = 16, p-value < 2.2e-16). The Ordovician and Mississippian exhibit
244 surprisingly well preserved records, with higher than anticipated counts of TG4 grade
245 specimens and proportionally less TG1 grades (Fig. 3D). In contrast, the Permian record is
246 shown to be poorly preserved, with a higher proportion of TG1 specimens than anticipated
247 and significantly reduced TG5 specimens. Both the Silurian and Pennsylvanian have a
248 significantly higher proportion of specimens at TG5.

249 Influences on the quality of fossil preservation

250 Comparisons of multiple logistic regression models using AICc consistently identify
251 grain size and palaeo-latitude as predictor variables in the models which best explain
252 preservational state for all five taphonomic grades (Table 1). Lithology is also a covariate in
253 the best model in all grades except for TG3 and latitude, and age (Ma) also appear as
254 covariates in the best model for different preservational states. In taphonomic grades
255 corresponding to disarticulated and semi-articulated fossil remains (TG1 to TG3), log odds of
256 fine grain size are negative (Fig. 4). Inversely, log odds of fine grain size are positive for
257 articulated skeletal remains (TG4 and TG5). Of particular note, fine grain size has the highest
258 positive log odds for any covariate in the best model for TG5, suggesting that the single biggest
259 predictor for the high-quality preservation is an association with fine-grained sediments.

260 Siliciclastic lithology also shows positive log odds for well-preserved individuals (TG4
261 and TG5), and is higher than grain size for TG4. For disarticulated but associated plates
262 (TG2), log odds for siliciclastic lithology are negative, implying higher odds of carbonate
263 lithology accompanying this preservational state. Whilst age, latitude and palaeo-latitude all
264 appear in a number of the top models, their log-odds are often near 0 or they are statistically
265 insignificant, implying little influence on quality of preservation.

266 Comparing preservational state, the rock record and palaeodiversity

267 Because the results of our logistic regression and statistical tests suggest that preservational
268 state varies systematically with respect to geological variables, we used Spearman's rank
269 correlation to compare variation in the counts of Palaeozoic North American siliciclastic and
270 carbonate units with the stage-level abundance of each taphonomic grade of North
271 American Palaeozoic echinoids from our dataset (Table 2). Siliciclastic units report strong
272 and statistically significant positive correlations with the abundance of all taphonomic grades,
273 whereas carbonate units show overall weaker positive correlations, with only TG4 reporting
274 a statistically significant result (Table 2). Disarticulated single plates (TG1) show the lowest
275 strength of correlation across both siliciclastic and carbonate units.

276 To further explore the relationship between taphonomic grade and patterns of
277 taxonomic diversity, we compared the abundance of different preservational states through
278 time to standing diversity of Palaeozoic at the genus level (Table 3). At the stage level, no
279 taphonomic grades show statistically significant positive correlations with genus-level
280 diversity; however, TG5 shows a very strong and statistically significant correlation with
281 diversity at the Period level.

282 We additionally tested the relationship between taphonomic grade and sampling
283 proxies (Table 4). Preservation scores TG1-4 all show strong, positive and statistically
284 significant correlations with counts of echinoid collections through time (Table 4); TG1&2 in
285 particular show very strong correlations. Counts of TG1, 3 and 4 show positive and
286 statistically significant correlations with total Echinodermata collections throughout the
287 Palaeozoic, with TG1 showing the strongest overall. Conversely, correlations between
288 preservation scores and counts of global collections are weak, and none are reported as
289 statistically significant.

290

291 **DISCUSSION**

292 The taphonomic nature of the Palaeozoic echinoid fossil record

293 Our analyses highlight the factors which control the preservation of echinoids, an
294 invertebrate class with multi-element, calcium carbonate skeletons, throughout the
295 Palaeozoic. Simple quantification of different taphonomic grades indicates that the quality of
296 preservation in the echinoid record is non-random, with the majority of recorded specimens
297 belonging to taphonomic grades representing disarticulated plates (Fig. 3A). The prevalence
298 of disarticulated material is in line with species-level analyses of echinoid taphonomic grades
299 (Smith 2007; Greenstein 1992), which typically highlight the abundance of disarticulated
300 spines. These results also align with studies of the vertebrate fossil record; where isolated or
301 fragmentary remains are frequently more abundant than more complete fossils (Gardner *et*
302 *al.* 2016; Schnetz *et al.* 2022; Brown *et al.* 2019). This is seemingly also the case with
303 echinoids and suggests that a relatively low level of phylogenetically or taxonomically useful
304 information is likely to be present in the majority of the sampled fossil record. While the
305 disarticulated plates of Palaeozoic echinoids can often be recognised at the family level
306 (Thompson and Denayer 2017), when we plot the proportion of specimens identifiable to
307 different taxonomic ranks in each taphonomic grade using our taxonomically standardised
308 database (Fig. 3A,C), it is clear that as taphonomic grade increases, so too does the
309 proportion of total specimens identifiable to lower-level taxonomic ranks, such as species or
310 genus.

311 A number of studies have highlighted the importance of so called “hidden diversity”
312 that can be recognised from disarticulated skeletal elements and which have been proposed
313 to will greatly increase our understanding of palaeobiodiversity of fossil groups (Twitchett and
314 Oji 2005; Pietsch *et al.* 2019; Oji and Twitchett 2015; Martins *et al.* 2023). Because of the
315 limited taxonomic utility of disarticulated specimens, for echinoids at least, it is unlikely that
316 disarticulated specimens will significantly alter our understanding of deep time biodiversity
317 change, and suggests that a minority of the record, those specimens from higher
318 taphonomic grades, will be useful for analyses requiring higher taxonomic specificity.
319 Notably, this contrasts with recent work on vertebrates, which has highlighted the utility of

320 disarticulated skeletal elements for taxonomic identification and phylogenetic analysis
321 (Cashmore *et al.* 2021; Woolley *et al.* 2022). The echinoid body plan is comprised of
322 hundreds of skeletal elements, many of which are added indeterminately throughout growth
323 (Smith 2005; Thompson *et al.* 2021), and almost all of which exist in multiples of at least five.
324 Because of this, identifying homologous skeletal ossicles across different species from
325 disarticulated remains is difficult. This is further complicated because of intraspecific and
326 ontogenetic variation (Raup 1956; David and Laurin 1996). This situation contrasts to
327 bilaterally symmetrical vertebrates where (with the exception of axial elements) in many
328 species a single bone can be readily identified and provide precise taxonomic information
329 due to bone-specific autapomorphies (Cashmore, *et al.* 2021). The taxonomic utility of
330 disarticulated fossils, and thus the impact of taphonomic variables on fossil preservation, are
331 likely to be felt differently across different taxonomic groups.

332 Our results furthermore show that different higher Palaeozoic echinoid taxa vary
333 systematically in their preservation, and that taxon-specific differences in preservational
334 quality may be related to factors intrinsic to the animals, such as differential aspects of
335 skeletal anatomy. Different groups of Post-Palaeozoic echinoids, most classically the regular
336 and irregular echinoids, are well known for their differential preservation in the fossil record
337 (Greenstein 1993; Kier 1977), associated with differences in microstructural interlocking of
338 test plates (Smith 1984; Grun and Nebelsick 2018). The relative over- and under-
339 representation of particular families in different taphonomic grades suggests similar
340 differences may exist for some Palaeozoic echinoid groups. One example is the
341 Palaechinidae, which have characteristic tessellate test plating and are over-represented as
342 articulated tests with denuded spines (TG4), or as fragments of tests (TG3). This is in stark
343 contrast to other Palaeozoic families, such as the archaeocidarids, that have imbricate
344 plating and are frequently preserved at the lowest taphonomic grades (Fig. 1C). This
345 suggests that the tessellate plating in the Palaechinids may be less-prone to disarticulation,
346 instead favouring fragmentation as is the case in many extant euechinoids (Kidwell and

347 Baumiller 1990). It further suggests the presence of micro-structural interlocking, as present
348 in many crown group echinoids (Smith, 1980, Grun and Nebelsick, 2018), may have been
349 present in Palaechinids.

350 Factors influencing multi-element skeletal preservation

351 Because fossils at higher taphonomic grades impact more on our perceptions of
352 palaeodiversity, we sought to identify the factors that influence the taphonomic grades of
353 individual specimens, as well as their relative impact on preservational state. Our
354 interrogation of the various factors affecting fossil preservation reveals that, while taxonomy
355 position, grain size, and lithology all effect fossil preservation, of the factors included in our
356 multiple logistic regression, the most important is grain size. Fossils at lower taphonomic
357 grades are less likely to be associated with fine-grained sediments, while fossils at higher
358 taphonomic grades are positively associated with fine-grained sediments (Fig. 4). After
359 death, echinoid skeletons disarticulate rapidly (Kidwell and Baumiller 1990), and due to their
360 porous test microstructure have low residence times on the seafloor relative to other more
361 robust, taxa (Kowalewski *et al.* 2018). Rapid burial is thus generally evoked as the single
362 most relevant factor in generating high-quality preservation of fossil echinoderms, as lower
363 residence time on the seafloor is more likely to prevent disarticulation from disruption by
364 scavengers, bioturbators, or erosion (Brett *et al.* 1997). Many of specimens in our dataset of
365 the highest taphonomic grades are likely the result of obrution deposits where a layer of
366 sediment has rapidly covered the animal during or soon after life (Brett and Seilacher 1991).
367 These deposits are often associated with turbidites and tempestites (Ausich and
368 Sevastopulo 1994), and often found in offshore slope or deltaic environments (Brett, *et al.*
369 1997; Brett and Seilacher 1991). While we did not attempt to characterise
370 palaeoenvironment across our dataset due to the nature of historical museum collections
371 data, many of the fine-grained occurrences in our dataset were likely deposited in deep-
372 water settings. In contrast to fine-grained preservation associated with event beds, coarse-
373 grained environments have been previously associated with, typically more near-shore, and

374 higher-energy preservation of disarticulated echinoderms (Ausich and Sevastopulo 1994;
375 Brett, *et al.* 1997) Recent experimental work has shown that during rapid burial, coarse-
376 grained sediments are more likely to be associated with disarticulated or fragmentary
377 echinoderm remains while finer-grained, more thixotropic sediments are more likely to
378 impede movement during escape, and thus result in higher-quality preservation (Fraga and
379 Vega 2024). While we have no examples of soft-tissue preservation in our dataset, rapid
380 burial with fine-grained sediments has also been shown to be conducive to the high-quality
381 soft-tissue preservation which characterises many Lagerstätte (Gaines *et al.* 2012; Allison
382 1988). Our global and temporally expansive analyses thus align with the numerous studies
383 focusing on local sedimentological and geochemical conditions of numerous Lagerstätte and
384 obrution deposits which have demonstrated the association between grain size taphonomic
385 grade of fossil preservation.

386 At higher taphonomic grades, lithology also positively influences preservation, with
387 well-preserved fossils associated with siliciclastic rocks. This mirrors a study of the
388 ichthyosaur fossil record, which found the highest average completeness of specimens in
389 fine grained siliciclastic (but not carbonate) sediments (Cleary, *et al.* 2015). Additionally, it
390 has been noted that for similar depositional environments, siliciclastic settings tend to result
391 in slightly better preserved fossil echinoderms than carbonate settings (Brett, *et al.* 1997)
392 due to higher overall rates of sedimentation (Enos 1991; Schindelz 1980). The high rates of
393 sedimentation in some siliciclastic environments, such as deltas, slopes, and channels, may
394 explain this association in our dataset. Indeed many of our specimens at the highest
395 taphonomic grades were recorded from well-known siliciclastic palaeoenvironments such as
396 the deltaic Edwardsville Formation, the channel-shaped deposits of the Winchell Formation
397 and the offshore downslope Ordovician Ladyburn Starfish beds (Lane 1973; Schneider, *et al.*
398 2005). In our dataset, however, the relationship between lithology and taphonomic grade
399 breaks down at the lowest taphonomic grade (Fig. 4), suggesting that differences in lithology
400 have little impact on dictating the distribution of disarticulated remains. This is likely because

401 single disarticulated elements are more robust to biostratinomic processes such as transport
402 and reworking (Brett, *et al.* 1997), and are thus more likely to be found across a range of
403 environments. We suggest these are general patterns found across the marine skeletal fossil
404 record, though further studies in other fossil groups will be necessary to verify this.

405 We have also identified factors which have little bearing on the quality of fossil
406 preservation. While these covariates are present in many of the best models identified for
407 each preservational score, latitude, palaeolatitude, and fossil age all have little impact on
408 fossil preservation based upon our analyses. Given that modern latitude has little to do with
409 the original environmental conditions of a specific locality in the Palaeozoic, the lack of
410 influence of latitude on specimen is not surprising. Furthermore, while variables linked to
411 palaeolatitude, such as temperature, have been determined to influence specimen
412 preservation in live-dead comparisons (Kidwell *et al.* 2005), and experiments (Kidwell and
413 Baumiller 1990), given how global temperatures have fluctuated across different latitudes
414 and times throughout the Palaeozoic (Goldberg *et al.* 2021; Korte *et al.* 2008), it is perhaps
415 not surprising that specific latitudes are not associated with a given taphonomic grade or
416 grades. In contrast, the age of fossil bearing-deposits has been identified as a megabias
417 which influences the quality of the fossil record, and has been shown to influence
418 taphonomic state of fossils in a non-random way (Hendy 2011). This impact of age on
419 preservation is largely linked to specific intervals of time, such as the Neogene, with near-
420 recent levels of fossil quality, and the Permian which is associated with high degrees of
421 silicification (Butts and Briggs 2011; Hendy 2011). Age could play less of a role in our
422 analysis because all of our fossils are all Palaeozoic, and the Permian in particular lacks
423 many well-preserved Palaeozoic echinoids (see discussion below). All in all, our analyses
424 suggests that larger-scale determinants, such as latitude, palaeolatitude, and age have less
425 of an impact on fossil preservation than more localised sedimentological and
426 palaeoenvironmental factors, such as the above-discussed lithology and grain size.

427 **Taphonomy, the rock record and perceived palaeodiversity**

428 We found no overall correlation between the abundance of different taphonomic grades and
429 taxonomic diversity at the stage and period levels (Table 3). This is in agreement with
430 studies comparing patterns of skeletal completeness with raw taxonomic diversity, which
431 have typically identified an absence of, or statistically weak, correlations between specimen
432 or species quality and richness (Cashmore and Butler 2019; Cleary, *et al.* 2015; Schnetz, *et*
433 *al.* 2022; Brown, *et al.* 2019), though see Dean *et al.* (2016) for a contrasting pattern
434 observed in pterosaurs. The only exception to this in our dataset was a strong correlation
435 between period-level genus richness and the abundance of fossils at the highest taphonomic
436 grade, TG5. This implies that, at least at coarse temporal resolution, the abundance of well-
437 preserved fossils could influence patterns of observed palaeodiversity. This is likely due to
438 the aforementioned ability to more readily assign the best-preserved fossils to lower
439 taxonomic levels (Fig. 3A,C). However, overall, it appears that quality of preservation does
440 not significantly influence our understanding of Palaeozoic echinoid diversity.

441 Comparisons between the number of collections and abundance of different
442 taphonomic grades through time can reveal whether preservational quality is predominantly
443 controlled by sampling variation. Whilst TG1-TG4 reported strong and statistically significant
444 correlations with the number of echinoid collections through time, this was strongest for the
445 lowest taphonomic grades; conversely TG5 showed a near miss for statistical significance,
446 and the lowest strength of correlation (Table 4). This pattern indicates that the abundance of
447 poorly preserved specimens closely mirrors echinoid sampling intensity, and that additional
448 sampling will produce an equivalent amount of incomplete fossil material. At higher
449 taphonomic grades, the strength of this correlation reduces, indicating that high quality
450 preservation is less-influenced by sampling effort. This is likely because higher taphonomic
451 grades likely are likely the result of event-based taphonomic conditions preserving ecological
452 snapshots (Ausich 2016), as opposed to background conditions where specimens are
453 exposed to a greater degree of biostratigraphic processes (Brett, *et al.* 1997; Brett and Baird
454 1986). A more pronounced variation of this pattern is observed with comparisons to total

455 Echinodermata collections (Table 4). Conversely, the lack of significant correlations between
456 total invertebrate collections and any taphonomic grades indicates that preservation is not
457 dependent on 'global' measure of sampling intensity during the Palaeozoic, but more
458 focused sampling on localities, formations and facies likely to record echinoderm fossils.

459 Because lithology was shown to positively influence the preservation of higher
460 taphonomic grades, we compared the abundance of North American occurrences of different
461 taphonomic grades to the abundance and outcrop area of carbonate and siliciclastic
462 lithological units in Macrostrat (Table 2). Our results appear somewhat counterintuitive.
463 Despite echinoid fossils being more abundant in carbonate rocks (Figure 1d, 2a) and overall
464 showing an affinity for carbonate substrates (Thompson and Bottjer, 2019), variations in the
465 number of siliciclastic units through time report stronger and more significant positive
466 correlations with the abundance of different taphonomic grades than do carbonates (Table
467 2). This correlation is weakest at the lowest taphonomic grade, in agreement with our results
468 obtained from multiple logistic regression, which found that lithology plays little role in
469 preservation at lower grades, but that siliciclastic lithology influences higher taphonomic
470 grades (Fig. 4). At higher taphonomic grades, the stronger positive correlations between
471 siliciclastic sediment area and North American echinoid taphonomic grades may, as noted
472 above, reflect higher rates of sedimentation in siliciclastic environments (Schindelz 1980;
473 Enos 1991). Furthermore analyses of tropical mollusc death assemblages have shown
474 higher rates of skeletal degradation and reduced skeletal condition in carbonate
475 environments than in siliciclastic environments as a result of pore water chemistry and
476 disturbance from bioeroders or bioturbators (Kidwell, *et al.* 2005; Best *et al.* 2007). These
477 may explain why the distribution of North American taphonomic grades more-closely mirrors
478 siliciclastics. Though we did not analyse secular variations in the abundance of fine-grained
479 sediments through time we hypothesize, based on our analyses, that variations in the
480 proportion of fine to coarse-grained rocks, coupled with rates of sedimentation as evident in

481 siliciclastic settings, may be an overall regulator of fossil quality through time, at least for
482 taxa with multi-element skeletons.

483 Taphonomic impact on the evolutionary history of echinoids

484 The Palaeozoic is a crucial interval for echinoids, encompassing both their initial
485 diversification in their Ordovician (Thompson, *et al.* 2022), as well as the origin of the
486 echinoid crown group in the late Palaeozoic (Mongiardino Koch and Thompson 2021;
487 Thompson *et al.* 2020). Furthermore, the Permian, which yields some of the earliest crown
488 group echinoids, includes species which have been used to calibrate divergence time
489 estimates for the origin of the crown group echinoids (Koch *et al.* 2022; Thompson *et al.*
490 2015). Recent phylogenomic analyses have placed the origins of multiple extant echinoid
491 groups in the Permian (Mongiardino Koch *et al.* 2022), despite the absence of these
492 lineages from the Palaeozoic fossil record (Kier 1965). Identifying the number of lineages
493 that survived the end-Permian mass extinction and subsequently diversified in the Mesozoic
494 is a question with profound impacts on understanding the origins of modern echinoid
495 biodiversity and the role of the end-Permian extinction in dictating their evolutionary history
496 (Thompson *et al.* 2018; Thuy *et al.* 2017; Hagdorn 2017). While disarticulated remains are
497 relatively abundant in the Permian, the lack of specimens at the highest taphonomic grades
498 highlights that that majority of the sampled echinoid fossil record during the Permian are not
499 identifiable to lower taxonomic ranks. There is also relative paucity of fine-grained or
500 siliciclastic echinoid occurrences in our dataset which our analyses suggest are most
501 strongly associated with high-quality preservation. Taken together, this suggests that
502 differences in fossil preservation may explain the numerous ghost-lineages that plague our
503 understanding of the evolutionary history of echinoids in the Late Palaeozoic and Early
504 Mesozoic. More generally, this exemplifies how variations in the preservation of different
505 environments in the fossil record can influence not only our perception of ancient biodiversity
506 (Smith, *et al.* 2001), but also the factors which are conducive to the high-quality fossil
507 preservation which is necessary to understand the origins of living clades.

508

509 **CONCLUSION**

510 Multi-element skeletons are abundant across the tree of life and their preservation in the
511 fossil record sheds insight into fundamental mechanisms of macroevolution and the origin of
512 diverse modern clades (Koch, *et al.* 2022; Simoes *et al.* 2018; Legg, *et al.* 2013; Hopkins
513 and Smith 2015). Our results provide clear evidence of the dominant controls impacting the
514 preservational quality of the known Palaeozoic fossil record for an entire major taxonomic
515 group with a multi-element skeleton. We find that the majority of the Paleozoic echinoid fossil
516 record is disarticulated, and that the dominant controls on fossil preservation are grain size
517 and lithology. Counterintuitively, the abundance of North American siliciclastic rocks is more
518 closely linked to fluctuations in fossil preservation than are those of carbonates. The
519 influence of fossil preservation on our understanding of ancient biodiversity is likely to be
520 mediated through the ability to more precisely identify well-preserved fossils. The geological
521 and biological factors we herein implicate in controlling their fossil record are likely to be yet
522 another mechanism by which perceived deep time fluctuations in biodiversity influenced by
523 non-random fossil preservation.

524

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562

563 DATA ARCHIVING STATEMENT

564 Data for this study, including the underlying database for all analyses, data downloaded from
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567

568 REFERENCES

- 569 ALLISON, P. A. 1988. Konservat-Lagerstätten: cause and classification. *Paleobiology*, **14**,
570 331-344.
- 571 ALLISON, P. A. and BOTTJER, D. J. 2011. Taphonomy: bias and process through time.
572 *Taphonomy: Process and bias through time*, 1-17.
- 573 ANDERSON, R. P., TOSCA, N. J., SAUPE, E. E., WADE, J. and BRIGGS, D. E. 2021. Early
574 formation and taphonomic significance of kaolinite associated with Burgess Shale
575 fossils. *Geology*, **49**, 355-359.
- 576 ANTELL, G. T., BENSON, R. B. and SAUPE, E. E. 2024. Spatial standardization of taxon
577 occurrence data—a call to action. *Paleobiology*, **50**, 177-193.
- 578 AUSICH, W. 2016. Fossil Species as Data: A Perspective from Echinoderms. 301-311. In
579 ALLMON, W. D. and YACOBUCCI, M. M. (eds). *Species and Speciation in the*
580 *Fossil Record*. University of Chicago Press, Chicago, pp. Custom 7.
- 581 AUSICH, W. I. and SEVASTOPULO, G. D. 1994. Taphonomy of Lower Carboniferous
582 crinoids from the Hook Head Formation, Ireland. *Lethaia*, **27**, 245-256.
- 583 BARTON, K. 2023. MuMIn: Multi-Model Inference_. R package version 1.47.5.
- 584 BENJAMINI, Y. and HOCHBERG, Y. 1995. Controlling the false discovery rate: a practical
585 and powerful approach to multiple testing. *Journal of the Royal statistical society:*
586 *series B (Methodological)*, **57**, 289-300.
- 587 BENSON, R. B., BUTLER, R., CLOSE, R. A., SAUPE, E. and RABOSKY, D. L. 2021.
588 Biodiversity across space and time in the fossil record. *Current Biology*, **31**, R1225-
589 R1236.
- 590 BENSON, R. B., BUTLER, R. J., ALROY, J., MANNION, P. D., CARRANO, M. T. and
591 LLOYD, G. T. 2016. Near-stasis in the long-term diversification of Mesozoic
592 tetrapods. *PLoS biology*, **14**, e1002359.
- 593 BENTON, M. J., DUNHILL, A. M., LLOYD, G. T. and MARX, F. G. 2011. Assessing the
594 quality of the fossil record: insights from vertebrates. *Geological Society, London,*
595 *Special Publications*, **358**, 63-94.
- 596 BEST, M. M., KU, T. C., KIDWELL, S. M. and WALTER, L. M. 2007. Carbonate preservation
597 in shallow marine environments: unexpected role of tropical siliciclastics. *The Journal*
598 *of Geology*, **115**, 437-456.
- 599 BRETT, C. E. and BAIRD, G. C. 1986. Comparative taphonomy: a key to
600 paleoenvironmental interpretation based on fossil preservation. *Palaios*, 207-227.

601 BRETT, C. E., MOFFAT, H. A. and TAYLOR, W. L. 1997. Echinoderm taphonomy,
602 taphofacies, and Lagerstätten. *The Paleontological Society Papers*, **3**, 147-190.

603 BRETT, C. E. and SEILACHER, A. 1991. Fossil Lagerstätten: a taphonomic consequence of
604 event sedimentation. 283-297. In EINSELE, G., RICKEN, W. and SEILACHER, A.
605 (eds). *Cycles and events in stratigraphy*. pp. Custom 7.

606 BRIGGS, D. E., RAISWELL, R., BOTTRELL, S., HATFIELD, D. T. and BARTELS, C. 1996.
607 Controls on the pyritization of exceptionally preserved fossils; an analysis of the
608 Lower Devonian Hunsrueck Slate of Germany. *American Journal of Science*, **296**,
609 633-663.

610 BROWN, E. E., CASHMORE, D. D., SIMMONS, N. B. and BUTLER, R. J. 2019. Quantifying
611 the completeness of the bat fossil record. *Palaeontology*, **62**, 757-776.

612 BURNHAM, K. P. and ANDERSON, D. R. 2001. Kullback-Leibler information as a basis for
613 strong inference in ecological studies. *Wildlife research*, **28**, 111-119.

614 BUTTS, S. H. and BRIGGS, D. E. G. 2011. Silicification Through Time. 411-434. In
615 ALLISON, P. A. and BOTTJER, D. J. (eds). *Topics in Geobiology, Taphonomy:
616 Process and Bias Through Time*. 11 pp. Custom 7.

617 CASHMORE, D. D. and BUTLER, R. J. 2019. Skeletal completeness of the non-avian
618 theropod dinosaur fossil record. *Palaeontology*, **62**, 951-981.

619 CASHMORE, D. D., BUTLER, R. J. and MAIDMENT, S. C. 2021. Taxonomic identification
620 bias does not drive patterns of abundance and diversity in theropod dinosaurs.
621 *Biology Letters*, **17**, 20210168.

622 CHERNS, L. and WRIGHT, V. P. 2000. Missing molluscs as evidence of large-scale, early
623 skeletal aragonite dissolution in a Silurian sea. *Geology*, **28**, 791-794.

624 CLEARY, T. J., MOON, B. C., DUNHILL, A. M. and BENTON, M. J. 2015. The fossil record
625 of ichthyosaurs, completeness metrics and sampling biases. *Palaeontology*, **58**, 521-
626 536.

627 CLOSE, R., BENSON, R. B., SAUPE, E., CLAPHAM, M. and BUTLER, R. 2020. The spatial
628 structure of Phanerozoic marine animal diversity. *Science*, **368**, 420-424.

629 DAVID, B. and LAURIN, B. 1996. Morphometrics and cladistics: measuring phylogeny in the
630 sea urchin *Echinocardium*. *Evolution*, **50**, 348-359.

631 DEAN, C. D., ALLISON, P. A., HAMPSON, G. J. and HILL, J. 2019. Aragonite bias exhibits
632 systematic spatial variation in the late Cretaceous Western Interior Seaway, North
633 America. *Paleobiology*, **45**, 571-597.

634 DEAN, C. D., CHIARENZA, A. A. and MAIDMENT, S. C. 2020. Formation binning: a new
635 method for increased temporal resolution in regional studies, applied to the Late
636 Cretaceous dinosaur fossil record of North America. *Palaeontology*, **63**, 881-901.

637 DEAN, C. D., MANNION, P. D. and BUTLER, R. J. 2016. Preservational bias controls the
638 fossil record of pterosaurs. *Palaeontology*, **59**, 225-247.

639 DONOVAN, S. K. 1991. The taphonomy of echinoderms: calcareous multi-element
640 skeletons in the marine environment. 241-269. In DONOVAN, S. K. (ed.) *The
641 processes of fossilization*. Belhaven Press, London, pp. Custom 7.

642 ENOS, P. 1991. Sedimentary parameters for computer modeling. *Bulletin (Kansas
643 Geological Survey)*, 63-99.

644 FOOTE, M. and SEPKOSKI, J. J. 1999. Absolute measures of the completeness of the fossil
645 record. *Nature*, **398**, 415-417.

646 FRAGA, M. C. and VEGA, C. S. 2024. How does rapid burial work? New insights from
647 experiments with echinoderms. *Palaeontology*, **67**, e12698.

648 GAINES, R. R., HAMMARLUND, E. U., HOU, X., QI, C., GABBOTT, S. E., ZHAO, Y.,
649 PENG, J. and CANFIELD, D. E. 2012. Mechanism for Burgess Shale-type
650 preservation. *Proceedings of the National Academy of Sciences*, **109**, 5180-5184.

651 GARDNER, E. E., WALKER, S. E. and GARDNER, L. I. 2016. Palaeoclimate, environmental
652 factors, and bird body size: a multivariable analysis of avian fossil preservation.
653 *Earth-Science Reviews*, **162**, 177-197.

- 654 GOLDBERG, S. L., PRESENT, T. M., FINNEGAN, S. and BERGMANN, K. D. 2021. A high-
655 resolution record of early Paleozoic climate. *Proceedings of the National Academy of*
656 *Sciences*, **118**, e2013083118.
- 657 GREENSTEIN, B. J. 1991. An integrated study of echinoid taphonomy: Predictions for the
658 fossil record of four echinoid families. *Palaios*, **6**, 519-540.
- 659 GREENSTEIN, B. J. 1992. Taphonomic bias and the evolutionary history of the family
660 Cidaridae (Echinodermata: Echinoidea). *Paleobiology*, **18**, 50-79.
- 661 --- 1993. Is the fossil record of regular echinoids really so poor? A comparison of living and
662 subfossil assemblages. *Palaios*, 587-601.
- 663 GRUN, T. B. and NEBELSICK, J. H. 2018. Structural design of the minute clypeasteroid
664 echinoid *Echinocyamus pusillus*. *Royal Society open science*, **5**, 171323.
- 665 HAGDORN, H. 2017. Slipped through the bottleneck: *Lazarechinus mirabeti* gen. et sp. nov.,
666 a Paleozoic-like echinoid from the Triassic Muschelkalk (late Anisian) of East France.
667 *PalZ*, 1-16.
- 668 HENDY, A. J. 2011. Taphonomic overprints on Phanerozoic trends in biodiversity:
669 lithification and other secular megabiases. *Taphonomy: process and bias through*
670 *time*, 19-77.
- 671 HOPKINS, M. J. and SMITH, A. B. 2015. Dynamic evolutionary change in post-Paleozoic
672 echinoids and the importance of scale when interpreting changes in rates of
673 evolution. *Proceedings of the National Academy of Sciences*, **112**, 3758-3763.
- 674 JONES, L. A., DEAN, C. D., MANNION, P. D., FARNSWORTH, A. and ALLISON, P. A.
675 2021. Spatial sampling heterogeneity limits the detectability of deep time latitudinal
676 biodiversity gradients. *Proceedings of the Royal Society B*, **288**, 20202762.
- 677 KIDWELL, S. M. and BAUMILLER, T. 1990. Experimental disintegration of regular echinoids:
678 roles of temperature, oxygen, and decay thresholds. *Paleobiology*, **16**, 247-271.
- 679 KIDWELL, S. M., BEST, M. M. and KAUFMAN, D. S. 2005. Taphonomic trade-offs in tropical
680 marine death assemblages: differential time averaging, shell loss, and probable bias
681 in siliciclastic vs. carbonate facies. *Geology*, **33**, 729-732.
- 682 KIER, P. M. 1965. Evolutionary trends in Paleozoic echinoids. *Journal of Paleontology*, 436-
683 465.
- 684 --- 1977. The poor fossil record of the regular echinoid. *Paleobiology*, **3**, 168-174.
- 685 KOCH, N. M., THOMPSON, J. R., HILEY, A. S., MCCOWIN, M. F., ARMSTRONG, A. F.,
686 COPPARD, S. E., AGUILERA, F., BRONSTEIN, O., KROH, A. and MOOI, R. 2022.
687 Phylogenomic analyses of echinoid diversification prompt a re-evaluation of their
688 fossil record. *Elife*, **11**, e72460.
- 689 KOCSIS, Á. T., REDDIN, C. J., ALROY, J. and KIESSLING, W. 2019. The R package
690 divDyn for quantifying diversity dynamics using fossil sampling data. *Methods in*
691 *Ecology and Evolution*, **10**, 735-743.
- 692 KORTE, C., JONES, P. J., BRAND, U., MERTMANN, D. and VEIZER, J. 2008. Oxygen
693 isotope values from high-latitudes: Clues for Permian sea-surface temperature
694 gradients and Late Palaeozoic deglaciation. *Palaeogeography, Palaeoclimatology,*
695 *Palaeoecology*, **269**, 1-16.
- 696 KOWALEWSKI, M., CASEBOLT, S., HUA, Q., WHITACRE, K. E., KAUFMAN, D. S. and
697 KOSNIK, M. A. 2018. One fossil record, multiple time resolutions: disparate time-
698 averaging of echinoids and mollusks on a Holocene carbonate platform. *Geology*, **46**,
699 51-54.
- 700 KOWALEWSKI, M. and FLESSA, K. W. 1996. Improving with age: the fossil record of
701 lingulide brachiopods and the nature of taphonomic megabiases. *Geology*, **24**, 977-
702 980.
- 703 KUO, P. C., BENSON, R. B. and FIELD, D. J. 2023. The influence of fossils in
704 macroevolutionary analyses of 3D geometric morphometric data: A case study of
705 galloanseran quadrates. *Journal of Morphology*, **284**, e21594.
- 706 LANE, N. G. 1973. Paleontology and paleoecology of the Crawfordsville fossil site (upper
707 Osagian: Indiana) with sections by J. L. Matthews, E.G. Driscoll, and E. L.
708 Yochelson. *University of California Publications in Geological Sciences*, **32**, 1-36.

709 LEGG, D. A., SUTTON, M. D. and EDGEcombe, G. D. 2013. Arthropod fossil data
710 increase congruence of morphological and molecular phylogenies. *Nature*
711 *communications*, **4**, 1-7.

712 MARTINS, L., COSTA, K. B. and TOLEDO, F. 2023. Disarticulated ossicles of sea
713 cucumbers from the Campos Basin, Brazil: A new perspective into the discovery of
714 diversity of Holothuroidea (Echinodermata). *The Holocene*, 09596836231197764.

715 MEYER, D., ZEILEIS, A. and HORNIK, K. 2006. The strucplot framework: visualizing multi-
716 way contingency tables with vcd. *Journal of Statistical Software*, **17**, 1-48.

717 MONGIARDINO KOCH, N. and THOMPSON, J. R. 2021. A total-evidence dated phylogeny
718 of Echinoidea combining phylogenomic and paleontological data. *Systematic Biology*,
719 **70**, 421-439.

720 NANGLU, K. and CULLEN, T. M. 2023. Across space and time: A review of sampling,
721 preservational, analytical, and anthropogenic biases in fossil data across
722 macroecological scales. *Earth-Science Reviews*, 104537.

723 NEBELSICK, J. H. 1996. Biodiversity of shallow-water Red Sea echinoids: implications for
724 the fossil record. *Journal of the Marine Biological Association of the United Kingdom*,
725 **76**, 185-194.

726 OJI, T. and TWITCHETT, R. J. 2015. The oldest post-Palaeozoic crinoid and Permian-
727 Triassic origins of the Articulata (Echinodermata). *Zoological Science*, **32**, 211-215.

728 PETERS, S. E., HUSSON, J. M. and CZAPLEWSKI, J. 2018. Macrostrat: A platform for
729 geological data integration and deep-time earth crust research. *Geochemistry*,
730 *Geophysics, Geosystems*, **19**, 1393-1409.

731 PIETSCH, C., RITTERBUSH, K. A., THOMPSON, J. R., PETSIOS, E. and BOTTJER, D. J.
732 2019. Evolutionary models in the Early Triassic marine realm. *Palaeogeography*,
733 *Palaeoclimatology, Palaeoecology*, **513**, 65-85.

734 RAUP, D. M. 1956. Dendraster: a problem in echinoid taxonomy. *Journal of Paleontology*,
735 685-694.

736 SCHINDELZ, D. E. 1980. Microstratigraphic sampling and the limits of paleontologic
737 resolution. *Paleobiology*, **6**, 408-426.

738 SCHNEIDER, C. L. 2008. The importance of Echinoids in late Paleozoic ecosystems. 71-90.
739 In AUSICH, W. and WEBSTER, G. (eds). *Echinoderm Paleobiology*. Indiana
740 University Press, Bloomington, pp. Custom 7.

741 SCHNEIDER, C. L., SPRINKLE, J. and RYDER, D. 2005. Pennsylvanian (Late
742 Carboniferous) Echinoids from the Winchell Formation, North-Central Texas, USA.
743 *Journal of Paleontology*, **79**, 745-762.

744 SCHNETZ, L., BUTLER, R. J., COATES, M. I. and SANSOM, I. J. 2022. Skeletal and soft
745 tissue completeness of the acanthodian fossil record. *Palaeontology*, **65**, e12616.

746 SIGNOR, P. W., LIPPS, J. H., SILVER, L. T. and SCHULTZ, P. H. 1982. Sampling bias,
747 gradual extinction patterns and catastrophes in the fossil record. *Geological Society*
748 *of America Special Paper*, **190**, e296.

749 SIMOES, T. R., CALDWELL, M. W., TAŁANDA, M., BERNARDI, M., PALCI, A.,
750 VERNYGORA, O., BERNARDINI, F., MANCINI, L. and NYDAM, R. L. 2018. The
751 origin of squamates revealed by a Middle Triassic lizard from the Italian Alps. *Nature*,
752 **557**, 706-709.

753 SMITH, A. 2005. Growth and Form in Echinoids: The Evolutionary Interplay of Plate
754 Accretion and Plate Addition. In BRIGGS, D. E. G. (ed.) *Evolving Form and Function:*
755 *Fossils and Development*. Yale Peabody Museum, New Haven, Connecticut, pp.
756 Custom 7.

757 SMITH, A. B. 1980. Stereom microstructure of the echinoid test. *Special Papers in*
758 *Palaeontology*, **25**, 1-85.

759 --- 1984. *Echinoid palaeobiology*. George Allen and Unwin, London, 190 pp.

760 --- 2007. Intrinsic versus extrinsic biases in the fossil record: contrasting the fossil record of
761 echinoids in the Triassic and early Jurassic using sampling data, phylogenetic
762 analysis, and molecular clocks. *Paleobiology*, **33**, 310-323.

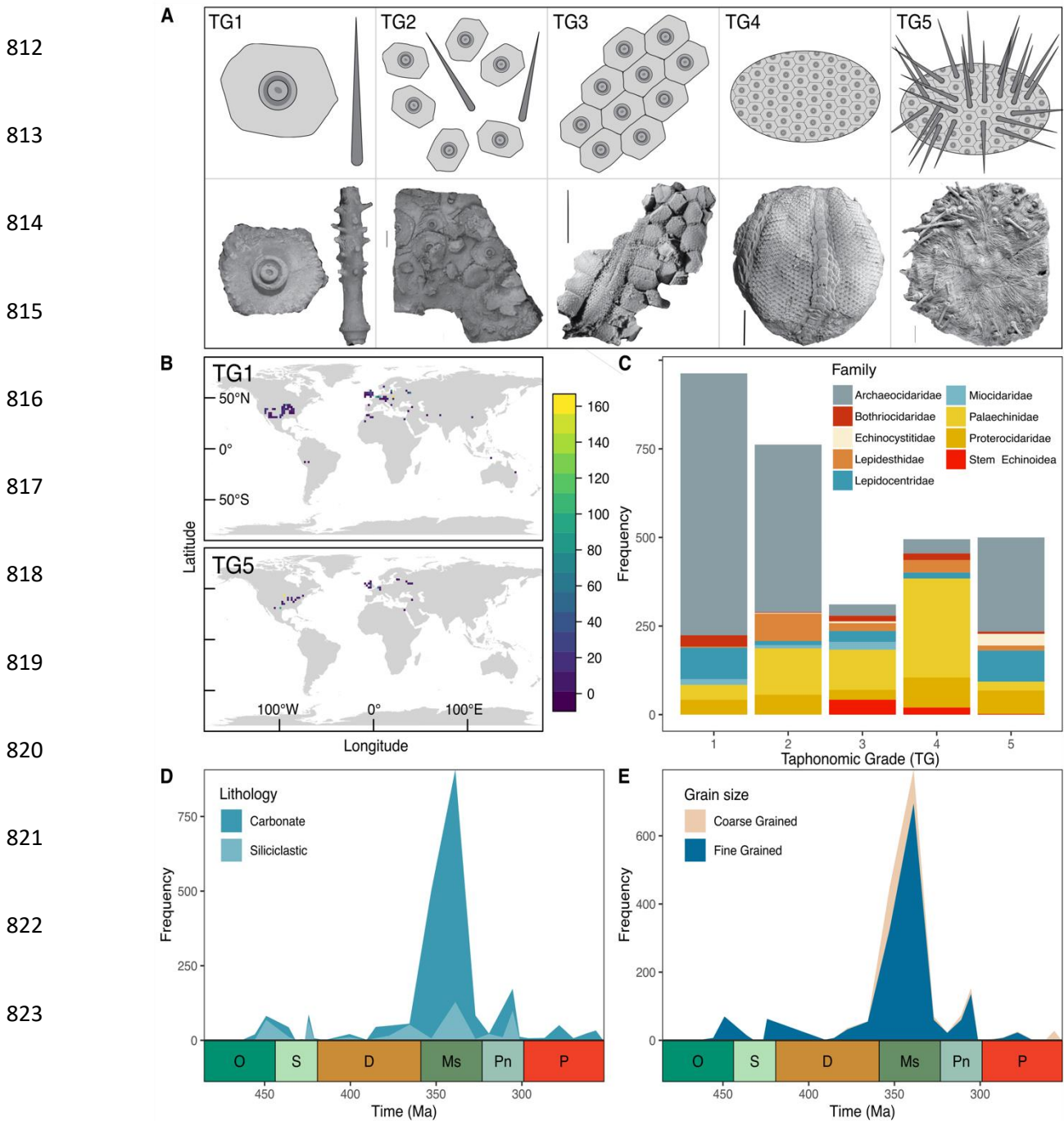
- 763 SMITH, A. B., GALE, A. S. and MONKS, N. E. 2001. Sea-level change and rock-record bias
764 in the Cretaceous: a problem for extinction and biodiversity studies. *Paleobiology*, **27**,
765 241-253.
- 766 THOMPSON, J. R. and BOTTJER, D. J. 2019. Quantitative analysis of substrate preference
767 in Carboniferous stem group echinoids. *Palaeogeography, Palaeoclimatology,*
768 *Palaeoecology*, **513**, 35-51.
- 769 THOMPSON, J. R., COTTON, L. J., CANDELA, Y., KUTSCHER, M., REICH, M. and
770 BOTTJER, D. J. 2022. The Ordovician diversification of sea urchins: Systematics of
771 the Bothriocidaroida (Echinodermata: Echinoidea). *Journal of Systematic*
772 *Palaeontology*, **19**, 1395-1448.
- 773 THOMPSON, J. R. and DENAYER, J. 2017. Revision of echinoids from the Tournaisian
774 (Mississippian) of Belgium and the importance of disarticulated material in assessing
775 palaeodiversity. *Geological Journal*, **52**, 529-538.
- 776 THOMPSON, J. R., HU, S.-X., ZHANG, Q.-Y., PETSIOS, E., COTTON, L. J., HUANG, J.-Y.,
777 ZHOU, C.-Y., WEN, W. and BOTTJER, D. J. 2018. A new stem group echinoid from
778 the Triassic of China leads to a revised macroevolutionary history of echinoids during
779 the end-Permian mass extinction. *Royal Society Open Science*, **5**, 171548.
- 780 THOMPSON, J. R., MIRANTSEV, G. V., PETSIOS, E. and BOTTJER, D. J. 2020.
781 Phylogenetic analysis of the Archaeocidaridae and Palaeozoic Miocidaridae
782 (Echinodermata: Echinoidea) and the origin of crown group echinoids. *Papers in*
783 *Palaeontology*, **6**, 217-249.
- 784 THOMPSON, J. R., PAGANOS, P., BENVENUTO, G., ARNONE, M. I. and OLIVERI, P.
785 2021. Post-metamorphic skeletal growth in the sea urchin *Paracentrotus lividus* and
786 implications for body plan evolution. *EvoDevo*, **12**, 1-14.
- 787 THOMPSON, J. R., PETSIOS, E., DAVIDSON, E. H., ERKENBRACK, E. M., GAO, F. and
788 BOTTJER, D. J. 2015. Reorganization of sea urchin gene regulatory networks at
789 least 268 million years ago as revealed by oldest fossil cidaroid echinoid. *Scientific*
790 *reports*, **5**.
- 791 THUY, B., HAGDORN, H. and GALE, A. S. 2017. Paleozoic echinoderm hangovers: Waking
792 up in the Triassic. *Geology*, G38909. 1.
- 793 TWITCHETT, R. J. and OJI, T. 2005. Early Triassic recovery of echinoderms. *Comptes*
794 *Rendus Palevol*, **4**, 531-542.
- 795 VON BITTER, P. H., PURNELL, M. A., TETREAULT, D. K. and STOTT, C. A. 2007.
796 Eramosa Lagerstätte—Exceptionally preserved soft-bodied biotas with shallow-
797 marine shelly and bioturbating organisms (Silurian, Ontario, Canada). *Geology*, **35**,
798 879-882.
- 799 WAGNER, P. J., ABERHAN, M., HENDY, A. and KIESSLING, W. 2007. The effects of
800 taxonomic standardization on sampling-standardized estimates of historical diversity.
801 *Proceedings of the Royal Society B: Biological Sciences*, **274**, 439-444.
- 802 WOOLLEY, C. H., THOMPSON, J. R., WU, Y.-H., BOTTJER, D. J. and SMITH, N. D. 2022.
803 A biased fossil record can preserve reliable phylogenetic signal. *Paleobiology*, **48**,
804 480-495.
- 805 WRIGHT, V. P., CHERNS, L. and HODGES, P. 2003. Missing molluscs: Field testing
806 taphonomic loss in the Mesozoic through early large-scale aragonite dissolution.
807 *Geology*, **31**, 211-214.

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810 **FIGURES**

811 **Figure 1**



824 Figure 1. The distribution of different taphonomic grades spatially, taxonomically, and with
 825 respect to lithology and grain size. A: Illustrations and photographs of Palaeozoic echinoid
 826 fossils demonstrating examples of our different taphonomic grades. TG1 is specimen
 827 MCZ:IP:103620, a disarticulated spine and interambulacral plate of the archaeocidarid
 828 *Archaeocidaris* sp. TG2 is specimen LACMIP 2020, a pile of disarticulated interambulacral

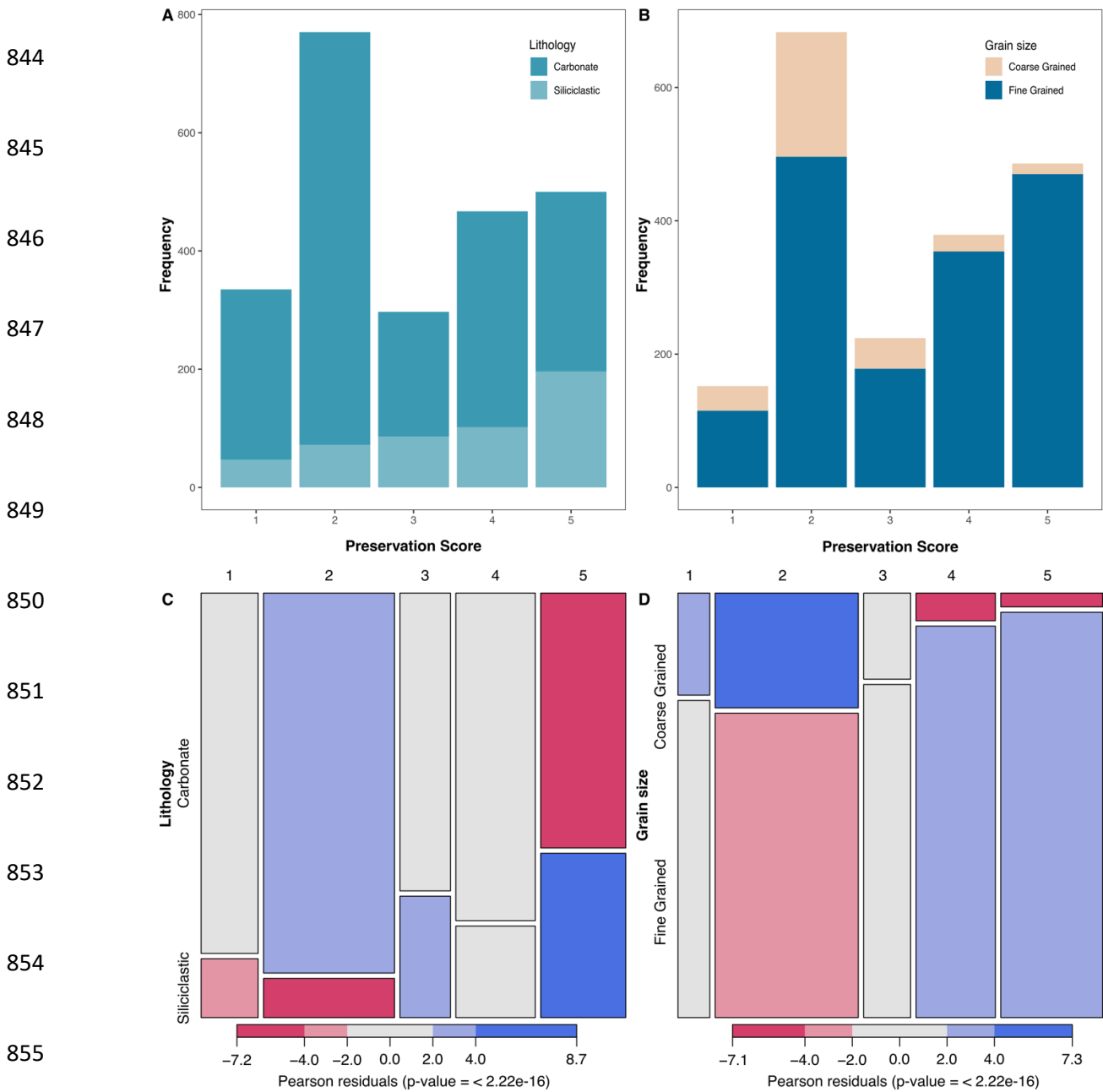
829 plates and spines from *Archaeocidaris* sp. TG3 is specimen MCZ:IP:101936, a fragment of a
830 test of the palaechinid *Maccoya intermedia* showing ambulacral and interambulacral plates.
831 TG4 shows specimen YPM 26552, a denuded test of the lepidesthid echinoid *Lepidesthes*
832 *colletti*. TG5 shows specimen NHMUK Pal E 76888, the holotype of *Archaeocidaris*
833 *whatleyensis*, which is an articulated test preserved with covering spines. B: the differential
834 global distribution of specimens of our lowest taphonomic grade (TG1, top) and our highest
835 taphonomic grade (TG5, bottom). C: bar chart showing the differential frequency of
836 taphonomic grades across our dataset broken down by family (or unit approximating family
837 for stem echinoids). D: frequency of specimens associated with carbonate or siliciclastic
838 sediments plotted through time. E: frequency of specimens associated with fine- or coarse-
839 grained sedimentary rocks plotted through time.

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843 **Figure 2**

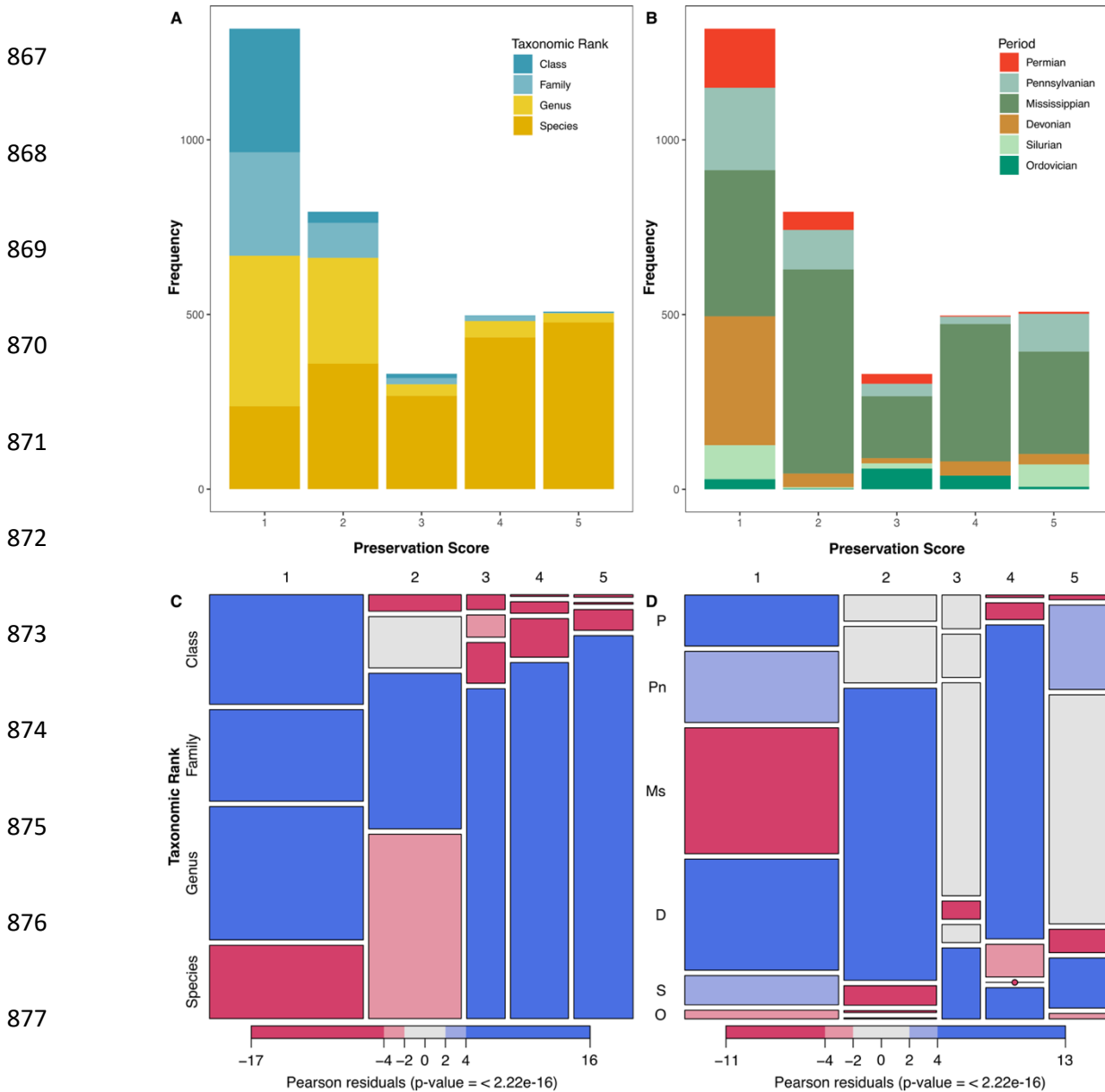


856 Figure 2. The distribution of different taphonomic grades according to lithology and grain-
 857 size. A: the frequency of different taphonomic grades according to carbonate or siliciclastic
 858 lithology. B: the frequency of different taphonomic grades according to grain size. C: mosaic
 859 plot showing the relative proportions of preservation scores within each lithological category,
 860 compared against expected levels, with colours relating to Pearson residuals from a log-
 861 linear model. D: mosaic plot showing the relative proportions of preservation scores within
 862 each grain size category, compared against expected levels, with colours relating to Pearson

863 residuals from a log-linear model. For mosaic plots, red indicates fewer observations than
 864 expected, whereas blue represents more observations than expected.

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866 **Figure 3**



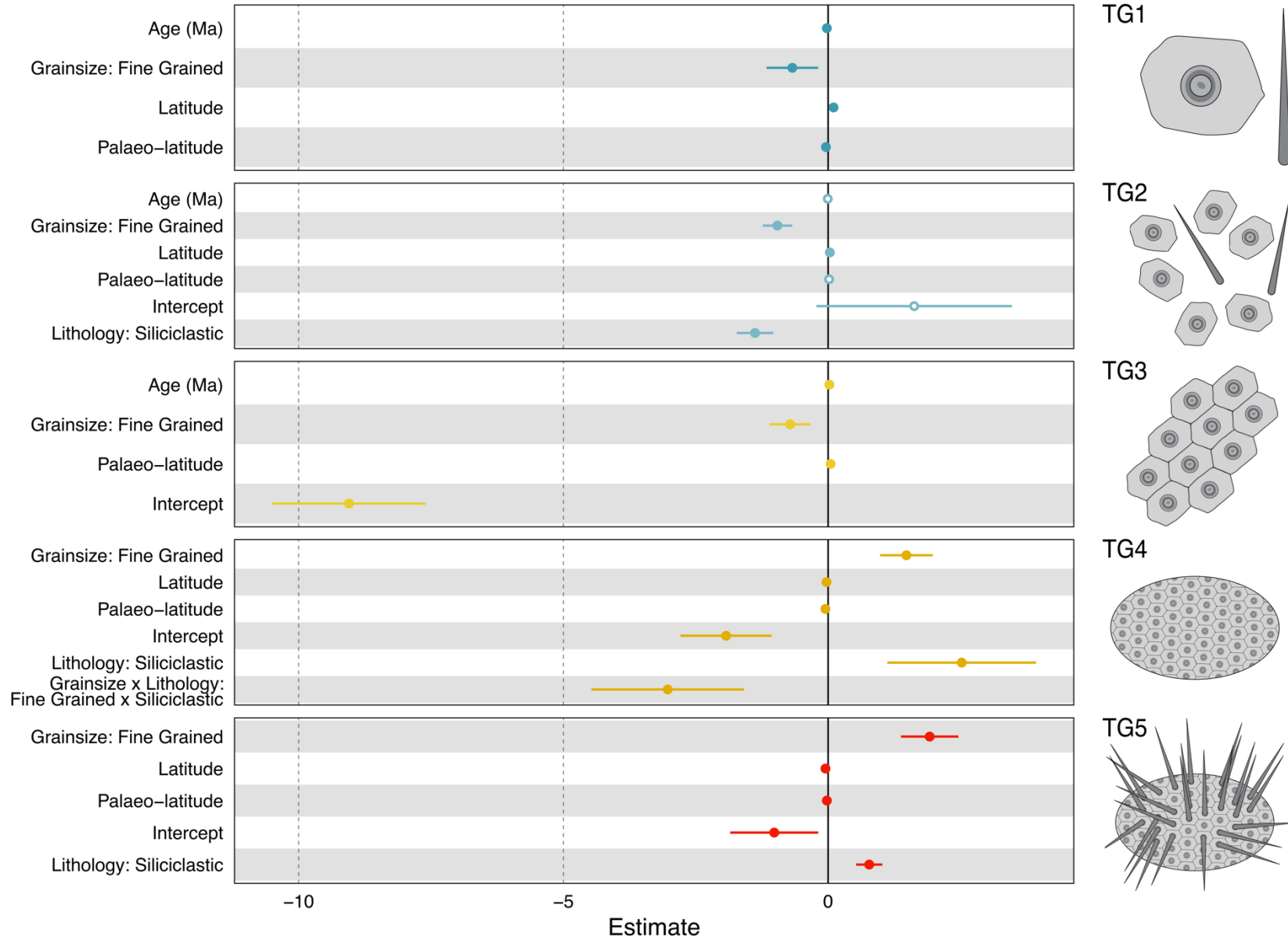
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879 Figure 3. The distribution of specimens of different taphonomic grades according to the
 880 lowest assignable taxonomic rank and geological time. A: the distribution of specimens of
 881 different taphonomic grades according to the lowest taxonomic rank to which each specimen

882 could be assigned. B: the frequency of different taphonomic grades according to the Period
883 of geological time from which they were sampled. C: mosaic plot showing the relative
884 proportions of preservation scores within each taxonomic rank, compared against expected
885 levels, with colours relating to Pearson residuals from a log-linear model. D: mosaic plot
886 showing the relative proportions of preservation scores within geological Period, compared
887 against expected levels, with colours relating to Pearson residuals from a log-linear model.
888 For mosaic plots, red indicates fewer observations than expected, whereas blue represents
889 more observations than expected.

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891 **Figure 4**



892 Figure 4. Forest plot showing best fitting model results for multiple logistic regression models
 893 for each taphonomic grade. Positive values indicate that the chance of observing a specimen
 894 at the associated taphonomic grade increases with the associated covariate, whereas
 895 negative values indicated the opposite. The output estimate is provided in log-odds scale.

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897 **TABLES**

898 **Table 1**

899 Top ranked multiple logistic regression models for each Taphonomic grade. TG: Taphonomic
 900 Grade. df: Degrees of freedom. AICc: small sample corrected Akaike Information Criterion.

<i>TG</i>	<i>Model</i>	<i>df</i>	<i>Log-likelihood</i>	<i>AICc</i>	<i>Weight</i>
1	~ Age (Ma) + Palaeo-latitude + Grain size + Latitude + Lithology + Grain size:Lithology	7	-394.816	803.7	0.675
2	~ Age (Ma) + Palaeo-latitude + Grain size + Latitude + Lithology	6	-974.809	1961.7	0.306
3	~ Age (Ma) + Palaeo-latitude + Grain size	4	-569.936	1147.9	0.395
4	~ Palaeo-latitude + Grain size + Latitude + Lithology + Grain size:Lithology	6	-796.565	1609.2	0.618
5	~ Age (Ma) + Palaeo-latitude + Grain size + Latitude + Lithology	5	-897.518	1805.1	0.302

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906 **Table 2**

907 Spearman's rank correlations between Taphnomic Grade and counts of North North
 908 American siliclastic and carbonate rock units in Macrostrat through time. TG: Taphnomic
 909 grade. BH: Benjamini Hochberg corrected p-value. Signif.: statistical significance, evaluated
 910 using Benjamini Hochberg corrected p-value < 0.05.

<i>Temporal Resolution</i>	<i>TG</i>	<i>vs. (test)</i>	<i>Rho</i>	<i>p</i>	<i>BH</i>	<i>Signif.</i>
Stage	1	Macrostrat carbonate (count)	0.312	0.056	0.137	
Stage	2	Macrostrat carbonate (count)	0.332	0.042	0.111	
Stage	3	Macrostrat carbonate (count)	0.346	0.033	0.092	
Stage	4	Macrostrat carbonate (count)	0.439	0.006	0.024	*
Stage	5	Macrostrat carbonate (count)	0.365	0.024	0.078	
Stage	1	Macrostrat siliclastic (count)	0.49	0.002	0.008	*
Stage	2	Macrostrat siliclastic (count)	0.644	0.00001	0.0003	*
Stage	3	Macrostrat siliclastic (count)	0.524	0.001	0.005	*
Stage	4	Macrostrat siliclastic (count)	0.627	0.00003	0.0003	*
Stage	5	Macrostrat siliclastic (count)	0.548	0.0004	0.003	*

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922 **Table 3**

923 Spearman's rank correlations between Taphonomic Grade and generic diversity at Stage
 924 and Period temporal resolutions. TG: Taphonomic grade. BH: Benjamini Hochberg corrected
 925 p-value. Signif.: statistical significance, evaluated using Benjamini Hochberg corrected p-
 926 value < 0.05.

<i>Temporal Resolution</i>	<i>TG</i>	<i>vs. (test)</i>	<i>Rho</i>	<i>p</i>	<i>BH</i>	<i>Signif.</i>
Stage	1	Diversity (genera)	0.462	0.047	0.119	
Stage	2	Diversity (genera)	0.541	0.017	0.058	
Stage	3	Diversity (genera)	0.343	0.151	0.280	
Stage	4	Diversity (genera)	0.258	0.286	0.391	
Stage	5	Diversity (genera)	0.402	0.088	0.196	
Period	1	Diversity (genera)	0.618	0.191	0.340	
Period	2	Diversity (genera)	0.618	0.191	0.340	
Period	3	Diversity (genera)	0.224	0.670	0.734	
Period	4	Diversity (genera)	0.265	0.612	0.690	
Period	5	Diversity (genera)	0.971	0.001	0.006	*

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938 **Table 4**

939 Spearman's rank correlations between Taphonomic Grade and counts of collections through
 940 time. TG: Taphonomic grade. BH: Benjamini Hochberg corrected p-value. Signif.: statistical
 941 significance, evaluated using Benjamini Hochberg corrected p-value < 0.05. echinoid and
 942 global collections, and counts of North American siliclastic and carbonate rock units.

<i>Temporal resolution</i>	<i>TG</i>	<i>vs. (test)</i>	<i>Rho</i>	<i>p</i>	<i>BH</i>	<i>Signif.</i>
Stage	1	Echinoid collections	0.776	0.00000046	0.00002	*
Stage	2	Echinoid collections	0.823	0.00000002	0.000002	*
Stage	3	Echinoid collections	0.645	0.0001	0.001	*
Stage	4	Echinoid collections	0.685	0.00003	0.0003	*
Stage	5	Echinoid collections	0.436	0.016	0.058	
Stage	1	Echinodermata collections	0.589	0.00011	0.0012	*
Stage	2	Echinodermata collections	0.361	0.031	0.091	
Stage	3	Echinodermata collections	0.562	0.00036	0.0025	*
Stage	4	Echinodermata collections	0.539	0.0007	0.0044	*
Stage	5	Echinodermata collections	0.298	0.077	0.18	
Stage	1	Global collections (pbdb)	0.207	0.212	0.352	
Stage	2	Global collections (pbdb)	0.0481	0.774	0.815	
Stage	3	Global collections (pbdb)	0.324	0.047	0.119	
Stage	4	Global collections (pbdb)	0.119	0.475	0.603	
Stage	5	Global collections (pbdb)	0.0113	0.947	0.958	

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