

1 **Realistic and simplified models of plant and leaf area indices for a seasonally**
2 **dry tropical forest**

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34 **Abstract**

35 Leaf Area Index (LAI) models that consider all phenological stages have not been developed for the
36 Caatinga, the largest seasonally dry tropical forest in South America. LAI models that are currently used
37 show moderate to high covariance when compared to *in situ* data, but they often lack accuracy in the
38 whole spectra of possible values and do not consider the impact that the stems and branches have over
39 LAI estimates, which is of great influence in the Caatinga. In this study, we develop and assess PAI
40 (Plant Area Index) and LAI models by using ground-based measurements and satellite (Landsat) data.
41 The objective of this study was to create and test new empirical models using a multi-year and multi-
42 source of reflectance data. The study was based on measurements of photosynthetic photon flux density
43 (PPFD) from above and below the canopy during the periods of 2011–2012 and 2016–2018. Through
44 iterative processing, we obtained more than a million candidate models for estimating PAI and LAI. To
45 clean up the small discrepancies in the extremes of each interpolated series, we smoothed out the dataset
46 by fitting a logarithmic equation with the PAI data and the inverse contribution of WAI (Wood Area
47 Index) to PAI, that is the portion of PAI that is actually LAI (LAI_C). LAI_C can be calculated as follows:
48 $LAI_C = 1 - (WAI/PAI)$. We subtracted the WAI values from the PAI to develop our *in situ* LAI dataset
49 that was used for further analysis. Our *in situ* dataset was also used as a reference to compare our models
50 with four other models used for the Caatinga, as well as the MODIS-derived LAI products
51 (MCD15A3H/A2H). Our main findings were as follows: (i) Six models use NDVI (Normalized
52 Difference Vegetation Index), SAVI (Soil-Adjusted Vegetation Index) and EVI (Enhanced Vegetation
53 Index) as input, and performed well, with r^2 ranging from 0.77 to 0.79 (PAI) and 0.76 to 0.81 (LAI),
54 and RMSE with a minimum of 0.41 $m^2 m^{-2}$ (PAI) and 0.40 $m^2 m^{-2}$ (LAI). The SAVI models showed
55 values 20% and 32% (PAI), and 21% and 15% (LAI), smaller than those found for the models that use
56 EVI and NDVI, respectively; (ii) the other models (ten) use only two bands, and in contrast to the first
57 six models, these new models may abstract other physical processes and components, such as leaves
58 etiolation and increasing protochlorophyll. The developed models used the near-infrared band, and they
59 varied only in relation to the inclusion of the red, green, and blue bands. (iii) All previously published
60 models and MODIS-LAI underperformed against our calibrated models. Our study was able to provide
61 several PAI and LAI models that realistically represent the phenology of the Caatinga.

62 Keywords: Caatinga, Landsat, phenology, semi-arid, Woody Area Index, Leaf Area Index.

63 **1. Introduction**

64 The Leaf Area Index (LAI) is a widely adopted parameter in environmental sciences studies.
65 It represents the one-sided area of leaves that covers a specific surface area (Fotis et al., 2018;

66 Knote et al., 2009; Mu et al., 2007; Rodriguez et al., 2009) and is one of the main parameters
67 of both global and regional biosphere models (Arnold et al., 1998; Bieger et al., 2017). LAI is
68 used to scale up from leaf to vegetation photosynthesis and transpiration, energy balance of
69 terrestrial surfaces, and many climatological and hydrological attributes such as atmospheric
70 aerosols, water infiltration, and biogeochemical processes (Bonan, 1995).

71 There are two main approaches used to estimate LAI: (i) direct methods, in which the
72 total leaf canopy is obtained by the summation of direct measurement of all individual leaf
73 areas – this is usually a destructive method because it requires the removal of all leaves and,
74 therefore, is not viable at large scales; and (ii) indirect methods, which may require active or
75 passive sensors to measure parameters that are highly correlated with LAI, such as light
76 extinction coefficient (Jonckheere et al., 2004); or use the litterfall trap method, which is
77 suitable for estimating LAI of deciduous plants (Almeida et al., 2019). Active sensors do not
78 depend on solar radiation as they emit their own electromagnetic signals and capture those
79 reflected, whereas passive sensors depend on solar radiation and are based on estimating the
80 extent to which a given amount of leaf area will reduce radiation transmitted through a stratified
81 arrangement of leaf elements within a canopy (Zheng and Moskal, 2009). This estimation can
82 be determined using a radiative transfer model such as the PROSPECT and SAIL models
83 (Jacquemoud et al., 2009; Jacquemoud and Baret, 1990; Knyazikhin et al., 1998; Verhoef,
84 1985, 1984) or abstracted by coefficients of an empirical model (Bastiaanssen, 1998; Galvncio
85 et al., 2013; Machado, 2014).

86 Radiative transfer models are highly accurate, but require specific inputs, such as
87 pigment concentration, cell diameter, and water content (Jacquemoud et al., 2009; Jacquemoud
88 and Baret, 1990). These parameters can only be obtained with extensive fieldwork, while
89 empirical models are purely statistical fast retrieval algorithms (Zheng and Moskal, 2009). To
90 estimate LAI, the empirical models are mainly composed of regressions that relate LAI values

91 to simple spectral responses and greenness indices (Almeida et al., 2019; Galvncio et al., 2013;
92 Machado, 2014), such as the Soil Adjusted Vegetation Index (*SAVI*) and Normalized
93 Difference Vegetation Index (*NDVI*) (Bastiaanssen, 1998; Galvncio et al., 2013). For LAI
94 estimations at a regional scale, empirical models are generally reliable (Knote et al., 2009).
95 However, in Brazil, more specifically in the seasonally dry tropical forest (SDTF) in the semi-
96 arid region, known as the Caatinga, models that are currently used have not been developed
97 using both intra- and inter-annual field measurements.

98 The Caatinga is the largest continuous SDTF in the Americas, with an open and mostly
99 semi-arid landscape, as seen in many inter-plateau depressions (Ab’Saber, 1974; Silva et al.,
100 2017). The Caatinga covers an area of approximately 900,000 km² (Silva et al., 2017), and
101 exhibits at least 13 different physiognomies ranging from woodlands to sparsely distributed
102 thorny shrubs (Silva et al., 2017). Its climate is characterized by high temperatures and low
103 rainfall rates with high intra- and inter-annual variability both in space and time. The rainfall
104 is normally concentrated over 2–4 months of the year, with the possibility of over 25% of the
105 annual precipitation occurring in a single rainfall event (Miranda et al., 2018). The main
106 landscape units that can be found in the Caatinga are canyons, ravines, mountains, sandy, and
107 clayey plateaus (Leal et al., 2007). Their complex soil mosaics are commonly formed by four
108 dominant soil orders (Latosols, Lithosols, Argisols, and Luvisols) (Menezes et al., 2012). The
109 Caatinga holds over 3,150 species of 930 genera and 152 families of flowering plants (Silva et
110 al., 2017). These plants have unique adaptations to endure conditions of spatiotemporally
111 irregular water availability and extended droughts: approximately 85% of the Caatinga species
112 lose all their leaves during the dry season (Leal et al., 2003; Silva et al., 2017). Thus, methods
113 that attempt to measure the LAI by directly relating it to the intercepted radiation do not reflect
114 only the area of the leaves, but also the surface of the woody area, which is mainly comprised
115 of stems and branches (Cunha et al., 2019). The influence that stems and branches have over

116 the LAI estimates can be addressed by computing the LAI as the difference between Plant Area
117 Index (PAI) and the Woody Area Index (WAI) (Kalacska et al., 2005).

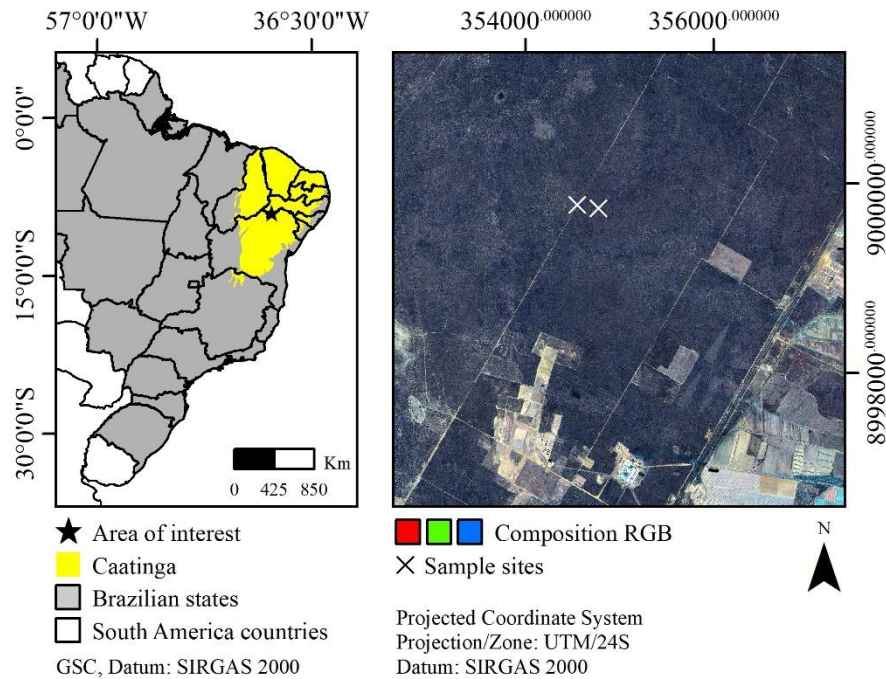
118 Current model estimates of LAI in the Caatinga show moderate to high covariance when
119 compared to *in situ* data ($r^2 = 0.60\text{--}0.93$), but they might lack accuracy in the entire spectra of
120 possible values because they are often developed using observations that do not cover complete
121 intra-annual LAI variations due to phenology (e.g., Almeida et al., 2019; Galvncio et al., 2013;
122 Machado, 2014). In addition, most of these models applied to the Caatinga have not entirely
123 considered the influence of the continuous variation of WAI, which is highly significant in the
124 Caatinga as over 85% of plant's above-ground biomass is composed of stems and branches
125 (Silva and Sampaio, 2008). The oversight of this uniqueness of SDTFs, such as the Caatinga,
126 is likely to cause methodological drawbacks in estimating LAI. By not considering every all
127 phenological stage, the intra-annual LAI changes may be reduced, and PAI can be wrongly
128 addressed as LAI. As a consequence, models may provide unrealistic values for some periods
129 of the year, especially in the dry season.

130 In this study, we aimed to create and test new empirical models using a multi-year and
131 multi-source set of reflectance data. We rely on the premise that by providing multiple
132 reflectance data combinations as input and that by accounting for the WAI component of the
133 PAI we will be able to provide models that are more accurate and better adjusted to the
134 Caatinga. Our objectives were to evaluate the efficiency of new LAI models derived from
135 Landsat reflectance using fitted regressions and field measurements from a typical Caatinga
136 formation area in Brazil, and to test new empirical approaches using previously published
137 models currently used for the Caatinga.

138 ***Study area***

139 Data were collected in an area of shrubby hyperxerophytic Caatinga forest area (Fig. 1) (Kiill,
140 2017), located at the Embrapa Tropical Semiarid Research Station in the state of Pernambuco,
141 Brazil (9°2'33"S, 40°19'16"W; at 350 m a.s.l.). The vegetation in this area consists of shrubs,
142 trees, herbaceous plants, and Cactaceae. The canopy average height is 4.5 m. The plant
143 phenological stages in the Caatinga are usually four: foliar development, maturity, senescence,
144 and dormancy (Rankine et al., 2017), and this cycle follows the rainfall patterns closely (Silva
145 et al., 2017). Most species in the Caatinga are deciduous, and respond quickly to slight changes
146 in soil water availability, breaking the dormancy of wood growth; and that allows the plants to
147 sprout most of their leaves in only a few days in the beginning of the rainy season (Machado
148 et al., 1997). The dominant plant species (approximately 90% of the total relative dominance)
149 in our study area were *Commiphora leptophloeos*, *Schinopsis brasiliensis*, *Mimosa tenuiflora*,
150 *Cenostigma microphyllum*, *Sapium glandulosum*, *Cnidoscullus quercifolius*, *Handroanthus*
151 *spongiosus*, *Manihot pseudoglaziovii*, *Croton conduplicatus*, and *Jatropha mollissima* (Kiill,
152 2017). Although the Cactaceae (*Pilosocereus gounellei* and *Pilosocereus pachycladus*) have a
153 fairly constant vegetative phenology throughout the year, these plants have a relative
154 dominance of less than 5% and an insignificant production of leaves; therefore they were not
155 considered in our LAI estimates. The climate is dry semi-arid (Alvares et al., 2013), with the
156 rainy season between January and April and an average annual temperature of 26°C. Although
157 the average historical annual rainfall is approximately 500 mm, the average rainfall was less
158 than 300 mm during our study period, which is the most severe drought in this region's
159 recorded history. These conditions were particularly interesting for our study, allowing a
160 precise assessment of the WAI influence on the total PAI.

161



162

163 Figure 1. Location of the seasonally dry tropical forest experimental area at the Embrapa
164 Semiarid Research Station in the state of Pernambuco (Brazil).

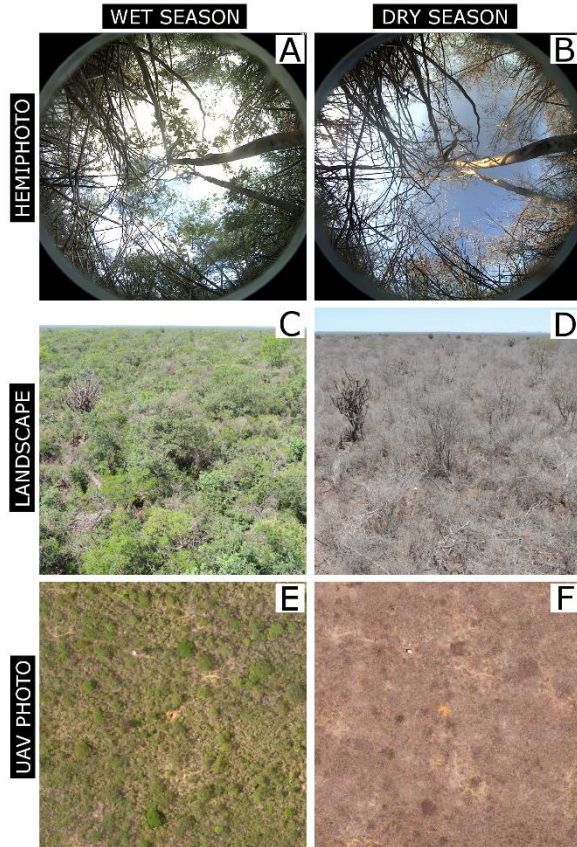
165 2. Methodology

166 *Field measurements*

167 LAI was derived from field measurements of photosynthetic photon flux density (PPFD) taken
168 from above and below the canopy using two different non-destructive methods. The
169 measurements were conducted throughout the year in order to cover all plant phenological
170 stages and covered five years: 2011–2012 and 2016–2018. The first method measured PPFD
171 using three quantum sensors (one LI-190SA sensor to measure the above-canopy PPFD, and
172 two LI-191 sensors for the below-canopy data) installed in a 16-m meteorological tower in the
173 study area. All sensors were connected to a data acquisition system (CR1000, Campbell
174 Scientific Inc.), which was programmed to compute averages of 30-s measurements taken at
175 30-min intervals from January 2011 to December 2012. In order to maximize the quality of our
176 measurements, we filtered all data, considering only the average of the measurements between

177 10 a.m. and 2 p.m. each day (GMT -3), when the zenith angle is close to zero. The second
178 measurement approach was applied on a weekly basis (68.97% of the entire dataset) from
179 January 2016 to November 2018, with the following exceptions: 19.54% (≥ 8 days of interval
180 between measurements – DBM), 8.05% (≥ 14 DBM) and 3.44% (≥ 21 DBM). The dataset
181 consisted of LAI estimates based on the transmission of light through the canopy at various
182 angles by using an AccuPAR ceptometer (AccuPAR[®] LP-80, Decagon Devices). The
183 AccuPAR has a linear ceptometer with 80 sensors, capable of measuring PPFD at the
184 photosynthetically active radiation (PAR) range (400–700 nm wavelength) from 0 to 2500
185 $\mu\text{mol m}^{-2}$. The above-canopy PPFD and solar zenith angle measurements were obtained in a
186 nearby (about 10 m away) clear area, and the below-canopy PPFD was acquired by holding the
187 AccuPAR beneath the canopy at approximately 0.4 m above ground. The dataset from this
188 approach was linearly interpolated to produce the daily data required to match the satellite
189 overpass times. We used the data collected to predict scattered and transmitted PPFD, as well
190 as to predict light extinction, as proposed by Norman (1979).

191



192

193 Figure 2. Contrast in the Caatinga between its wet (A, C and E) and dry (B, D and F) conditions.

194 A–B are hemispheric photos taken from below the vegetation in 12/18/2018 and 9/27/2018

195 respectively; C–D are landscape photos taken horizontally in 2/5/2016 and 10/20/2017 at the

196 height of 14 m; E–F are orthophotos taken by drone (unmanned aerial vehicle) at 80 m height

197 in 02/16/2018 and 10/20/2017, respectively.

198 *Plant Area Index (PAI) partitioning*

199 In our study, we defined PAI as the sum of WAI and LAI (Magalhães et al., 2018), and

200 the WAI as the contribution of woody material such as stems, branches, and trunks to the light

201 interception of PAI. In order to carry out this partition of our data, we first took the minimum

202 LAI (LAI_{MIN}) value of each year as the WAI, which was verified by visual evaluation of

203 hemispheric photos from a phenological monitoring database (Fig. 2); then we fixed this value

204 from the day of the LAI_{MIN} to the first subsequent day with rainfall over 2.5 mm. Based on

205 field observations, we assumed that low-precipitation ($\leq 2.5 \text{ mm d}^{-1}$) events did not cause any
206 significant phenological change in the ecosystem. The WAI was assumed to change between
207 sequential dry seasons gradually; we gap-filled the WAI dataset with a linear interpolation
208 between the fixed-value periods of each year. To avoid small discrepancies in the extremes of
209 each interpolated series, we smoothed the dataset by fitting a logarithmic equation (Eq. 1) with
210 the PAI data and the inverse contribution of WAI to PAI, which is the percentage of PAI that
211 is actually LAI (here called LAI_C). LAI_C was calculated as follows: $\text{LAI}_C = 1 - (\text{WAI}/\text{PAI})$.
212 The WAI values were subtracted from the PAI to develop our *in situ* LAI dataset, which is used
213 for further analysis.

214

$$\text{WAI} = \{1 - [\ln(\text{PAI}) \times 0.5]\} \times \text{PAI} \quad (1)$$

215 ***Landsat data processing***

216 We selected the Landsat Surface Reflectance Level-2 products for the entire study
217 period (total of 110 candidate images). These products are designed to provide atmospherically
218 and geometrically corrected reflectance data with 30-m resolution for every 16 days. These
219 data are generated using the auxiliary climate data from MODIS (e.g., water vapor, ozone,
220 geopotential height, and aerosol optical thickness) and two different algorithms: 1) the Second
221 Simulation of a Satellite Signal in the Solar Spectrum (6S) algorithm to the data derived from
222 Landsat 5 Thematic Mapper (TM) and Landsat 7 Enhanced Thematic Mapper Plus (ETM+)
223 images; and 2) a unique radiative transfer model to the Landsat 8 Operational Land Imager
224 (OLI) data. The data were extracted from two sample sites (Fig. 1), and all clear pixels were
225 filtered using the respective Quality Band (QA band) of each product (L5–7 = 66, and L8 =
226 322), resulting in a 70-record dataset. The dataset was then submitted to an iterative model-

227 fitting approach to create new PAI and LAI models. The Landsat Collection 2 Level-2 products
 228 include reflectance values derived from three sensors (TM/Landsat 5, ETM+/Landsat 7, and
 229 OLI/Landsat 8) with 30-m spatial resolution. The different bands were matched to create an
 230 equivalent dataset of reflectance across all sensors (Table 1). These products are freely
 231 available through the LSDS Science Research and Development (LSRD) database of the U.S.
 232 Geological Survey (<https://espa.cr.usgs.gov/>).

233

OLI/Landsat 8 (nm)	ETM+/Landsat 7 and TM/Landsat 5 (nm)	Equivalent bands for this study (nm)
-	$\rho_1^{ETM+/TM} = 450-520$	$\rho_1 = [\rho_2^{OLI}, \rho_1^{ETM+/TM}]$
$\rho_2^{OLI} = 452-512$	$\rho_2^{ETM+/TM} = 520-600$	$\rho_2 = [\rho_3^{OLI}, \rho_2^{ETM+/TM}]$
$\rho_3^{OLI} = 533-590$	$\rho_3^{ETM+/TM} = 630-690$	$\rho_3 = [\rho_4^{OLI}, \rho_3^{ETM+/TM}]$
$\rho_4^{OLI} = 636-673$	$\rho_4^{ETM+/TM} = 770-900$	$\rho_4 = [\rho_5^{OLI}, \rho_4^{ETM+/TM}]$
$\rho_5^{OLI} = 851-879$	$\rho_5^{ETM+/TM} = 1,550-1,750$	$\rho_5 = [\rho_6^{OLI}, \rho_5^{ETM+/TM}]$
$\rho_6^{OLI} = 1,566-1,651$	-	-
$\rho_7^{OLI} = 2,107-2,294$	$\rho_7^{ETM+/TM} = 2,090-2,350$	$\rho_7 = [\rho_7^{OLI}, \rho_7^{ETM+/TM}]$

234 Table 1 - Equivalence table of the bands of the sensors TM/Landsat 5, ETM+/Landsat 7 and
 235 OLI/Landsat 8.

236 **Model calibrations**

237 We developed PAI and LAI models based on the combinations of bands (ρ_1 to ρ_7); vegetation
 238 indices (*NDVI*, *SAVI* and *EVI*; Eqs. 2 to 4); transformation functions, i.e., x , $1/x$, $\ln(x)$,
 239 $\log_{10}(x)$, \sqrt{x} , x^2 , e^x ; and basic mathematical operations. These models were obtained by using
 240 an exhaustive training iteration process ($> 10^6$ iterations) that selected the best results based on
 241 the highest coefficient of determination (r^2) with the lowest Root Mean Square Error (RMSE).
 242 We used the Percent Bias (PBIAS) and the concordance correlation coefficient (ρ_c) as auxiliary
 243 performance indices. In our regression analysis, we used linear, logarithmic, exponential, and
 244 power functions to fit the observed data. We obtained *NDVI*, *SAVI*, and Enhanced Vegetation
 245 Index (*EVI*) using Eqs. 2 to 4, where $C1$ (6) and $C2$ (7.5) are the coefficients of the aerosol

246 resistance, G (2.5) is a gain factor, and L is the soil effect constant, according to Rouse et al.
 247 (1974) and Huete (1988). Our L for the EVI and $SAVI$ were set using a sensitivity analysis,
 248 varying the factor L from -1 to 1 with intervals of 0.01. The best L -value occurred when
 249 simulated data achieved the highest r^2 with the lowest RMSE. The L -values found were 0.07
 250 (for PAI models) and 0.37 (for the LAI models) for equations using $SAVI$; and 1 for both PAI
 251 and LAI models for equations using EVI . The number of models evaluated can be calculated
 252 using Eq. 5, where nc is the number of parameters entered into the model. All independent data
 253 were previously tested with the Variance Inflation Factor ($VIF = 1/(1 - r^2)$) to avoid any
 254 significant multicollinearity. We considered data to be independent when $VIF < 10$. All
 255 processing was performed using an interpreter Python 2.7.15 with only basic modules installed
 256 (freely available at <https://github.com/razeayres/correlator>).
 257

$$NDVI = \frac{(\rho_4 - \rho_3)}{(\rho_4 + \rho_3)} \quad (2)$$

$$SAVI = \frac{(1 + L) \times (\rho_4 - \rho_3)}{L + \rho_4 + \rho_3} \quad (3)$$

$$EVI = G \times \frac{\rho_4 - \rho_3}{\rho_4 + C1 \times \rho_3 - C2 \times \rho_1 + L} \quad (4)$$

$$f(nc) = C_{\begin{bmatrix} \rho_1 & \rho_4 & NDVI \\ \rho_2 & \rho_5 & SAVI \\ \rho_3 & \rho_7 & EVI \end{bmatrix} + (nc-1)}^{nc} \times C_{\begin{bmatrix} x & 1/x & \ln(x) \\ \log_{10}(x) & \sqrt{x} & x^2 \\ e^x & & \end{bmatrix} + (nc-1)}^{nc} \quad (5)$$

$$\times \begin{cases} 1, & nc = 1 \\ C_{\begin{bmatrix} + & - \\ \times & \div \end{bmatrix} + (nc-2)}^{(nc-1)}, & nc \geq 2 \end{cases}$$

258 **Models verification**

259 To verify the accuracy of all models in this study, we first assessed the applicability of
 260 parametric statistics to all data with the Shapiro–Wilk (for normality) and Brown–Forsythe (for

261 homoscedasticity) tests (Zar, 1996), and then we conducted a comparison between the remotely
262 sensed data and the estimates from the field observations using the Monte Carlo cross-
263 validation technique (Xu and Liang, 2001), considering 91 different sampling sizes varying
264 from 5 to 95% of the total data at 1% intervals. Each sample was evaluated by its r^2 and
265 computed as the mean of 50 random repetitions. The methods of cross-validation are widely
266 adopted, and they were used to check whether models tend to over-adjust to the *in situ* dataset
267 distribution (Hawkins, 2004). This over-adjustment would mean that excellent results would
268 be obtained only in calibration (Shao, 1993), while during verification, the accuracy of the
269 model would drastically drop. This approach allows for a good calibration (Shao, 1993). In
270 addition, we used the models proposed by Bastiaanssen (1998) (Eq. 6), Galvncio et al. (2013)
271 (Eq. 7), Machado (2014) (Eq. 8), and Almeida et al. (2019) (Eq. 9), and derived from MODIS
272 data (MCD15A3H/A2H) to produce independent data required for comparing with our field
273 observations and for testing our models. Except for the model developed by Bastiaanssen
274 (2018), these other models were specifically developed for the Caatinga. However, the model
275 of Bastiaanssen (1998) has been widely used to estimate LAI in this region (e.g., Bezerra et al.,
276 2014; Oliveira et al., 2015; Santos et al., 2017).

277

$$LAI = -\frac{\ln\left[\frac{(0.69 - SAVI)}{0.59}\right]}{0.91} \quad (6)$$

$$LAI = e^{1.426 + \frac{-0.542}{NDVI}} \quad (7)$$

$$LAI = 0.102 \times e^{5.341 \times NDVI} \quad (8)$$

$$LAI = 9.555 \times EVI - 1.324 \quad (9)$$

278 For Eqs. 6 to 9, we used the same Landsat dataset produced for the models calibrations; for the
279 MODIS MCD15A2H/A3H products, we used all images for the entire study period (total of
280 830 candidate images). These products are designed to provide data with a spatial resolution
281 of 500 m every four days (MCD15A3H) or every eight days (MCD15A2H). They are based
282 on a complex algorithm that uses both the daily surface reflectance values of the MODIS sensor
283 on one or both of the Terra and Aqua satellites and the data from a radiative transfer model,
284 which are stored in a two-dimensional lookup table (Yang et al., 2006). These reflectance data
285 are already corrected for atmospheric interferences such as atmospheric gases and aerosols,
286 and they are freely available through the Earth Explorer online tool of the U.S. Geological
287 Survey (<https://earthexplorer.usgs.gov/>). For all products, scale corrections were performed
288 using the Geospatial Data Abstraction Library and clear land dry forest pixels were filtered
289 using the Quality Band (QA band, value 0).

290 **3. Results and discussion**

291 Six of our selected models use *NDVI*, *SAVI* and *EVI* as input (Eqs. 15 to 17, and 23 to 25 in
292 Table 2). These models exhibited r^2 values ranging from 0.77 to 0.79 for PAI and 0.78 to 0.81
293 for LAI, and RMSE with a minimum of $0.41 \text{ m}^2 \text{ m}^{-2}$ for PAI and $0.40 \text{ m}^2 \text{ m}^{-2}$ for LAI. The
294 *SAVI* models (Eqs. 15 and 23) showed RMSE values smaller than the ones found for the
295 models that use *EVI* and *NDVI*. We ascribe the better accuracy with the *SAVI* models over the
296 other vegetation indices to the fact that *SAVI* takes into consideration the effects of soil
297 background, while not showing high variability as *EVI* does for sparsely vegetated areas, which
298 in turn produces infrared reflectance at low levels due to dry soil background (Lu et al., 2015).
299 In addition, *SAVI* better reflects the surface roughness, which affects momentum, heat, and
300 water vapor fluxes (Bastiaanssen, 1998), and varies according to the phenological stages of the
301 Caatinga (Teixeira et al., 2008). These models are useful because they allow easy retrieval of

302 the PAI or LAI from remote sensing data. For example, many *NDVI* products, using a large
303 variety of sensor data, are freely available, and they can be used to acquire physical information
304 for large forest areas.

305 Our models presented a better performance when fitted linearly rather than in any other
306 non-linear form (Table 2). This is the opposite of what was shown by some *NDVI*–LAI
307 relationship models (Liu et al., 2012; Tavakoli et al., 2014). Liu et al. (2012) conducted an
308 experiment in the Ningxia Hui Autonomous District, in Northwest China, and they found
309 saturation of *NDVI* at high LAI values. Tavakoli et al. (2014), in 16 plots of winter wheat
310 (*Triticum aestivum* L., cv. Cubus) in an experimental station located in Marquardt in Germany,
311 found the best *NDVI*–LAI relation when fitting data logarithmically. In fact, this saturation of
312 LAI as function of *NDVI* is commonly expressed by a logarithmic relationship. However,
313 *NDVI* values tend to be poorly associated with those from ground observations in SDTFs
314 (Guzmán et al., 2019). In our study, the vegetation indices did not exhibit saturation related to
315 the LAI of the Caatinga vegetation, which resulted in a linear covariance as reflect in Eqs. 23–
316 25. Magalhães et al. (2018) showed that a linear model simulates better LAI in a SDTF by
317 arguing that *NDVI* can only saturate in vegetation types with LAI above $5 \text{ m}^2 \text{ m}^{-2}$. That supports
318 our findings since this threshold is above the values we used to develop LAI models for the
319 Caatinga.

320 In this study, absolute non-saturated simulated LAI values varied from 0 to $4.56 \text{ m}^2 \text{ m}^{-2}$
321 (0.61 to $5.23 \text{ m}^2 \text{ m}^{-2}$ for PAI values), while Bastiaanssen (1998) exhibited values from 0 to
322 ca. $4.45 \text{ m}^2 \text{ m}^{-2}$ (considering only the non-saturated values), Galvêncio et al. (2013) from 0.63
323 to $1.98 \text{ m}^2 \text{ m}^{-2}$, Machado (2014) from 0.25 to $3.7 \text{ m}^2 \text{ m}^{-2}$, and Almeida et al. (2019) from ca. 0
324 up to $4.26 \text{ m}^2 \text{ m}^{-2}$ in average. All of these previously models do not consider the temporal
325 variations due to the phenological stages of the Caatinga on a continuous multi-year basis, thus
326 the range of possible simulated values is smaller when compared to our models. Bastiaanssen

327 (1998) derived LAI using different equations for only seven types of land use cover types
 328 (cotton, maize, soy, wheat, fruit trees, vegetables, and native forests), none of which were
 329 similar to the dry forest in our study area. The study of Galvıncio et al. (2013) was based on a
 330 comparison of data obtained using an AccuPAR analyzer with indices created from
 331 spectroradiometry from a single day of measurements. The model proposed by Machado
 332 (2014) was developed in a Caatinga area of the National Park of Catimbau using only one
 333 Landsat 5 TM image, combined with 54 field-derived LAI measurements acquired three times
 334 over 20 days using simultaneous averages of diffuse light interception at five different zenith
 335 angles using sensors with fisheye lens. Almeida et al. (2019) created LAI models using the
 336 litterfall trap method in the Caatinga, and collected data for three representative species.
 337 Although the LAI models of Almeida et al. (2019) show high correlation to their field
 338 measurements, they were not able to consider the entire growth cycle in their analysis due to
 339 limitations of the litterfall fall method such as the discrete distribution of the measurements
 340 over time, and misestimation of the foliar development because of the appearance of new leaves
 341 between assessments.

342

Model	$r^{2\ 1}$	RMSE ²	ρ_c	PBIAS ²
Eq. 10 $y = 10.1 \times (\rho_4 - \sqrt{\rho_3}) + 3.1$	0.79	0.41	0.88	0.33
Eq. 11 $y = -13.2 \times (\sqrt{\rho_2} - \rho_4) + 3.1$	0.77	0.44	0.87	1.84
Eq. 12 $y = -13.5 \times \left(\frac{\log_{10}(\rho_4)}{\ln(\rho_3)} \right) + 6.1$	0.77	0.43	0.87	-1.84
Eq. 13 $y = -20.3 \times (\rho_3 - \rho_4^2) + 3$	0.77	0.43	0.87	-0.83
Eq. 14 $y = -3.2 \times (\ln(\rho_3) \times \sqrt{\rho_4}) - 1.4$	0.79	0.41	0.88	-0.22
Eq. 15 ³ $y = 3.5 \times (e^{SAVI}) - 2.7$	0.79	0.41	0.88	1.10
Eq. 16 $y = 4.8 \times (e^{EVI}) - 3.7$	0.77	0.45	0.86	3.72
PAI Eq. 17 $y = 5 \times (NDVI^2) + 1.3$	0.79	0.43	0.89	1.04
LAI Eq. 18 $y = \left(\frac{\rho_4^2}{\rho_1} \right) - 0.1$	0.79	0.41	0.88	-0.01

Eq. 19	$y = -9.7 \times \left(\frac{\log_{10}(\rho_3)}{\left(\frac{1}{\rho_4}\right)} \right) - 1.2$	0.78	0.42	0.88	-4.84
Eq. 20	$y = 11.2 \times (\sqrt{\rho_4} - e^{\rho_3}) + 8.3$	0.76	0.44	0.86	7.15
Eq. 21	$y = 12.2 \times (\sqrt{\rho_4} - \sqrt{\rho_2}) - 1.2$	0.76	0.44	0.86	-0.73
Eq. 22	$y = 19.6 \times (\rho_4^2 - e^{\rho_3}) + 21.4$	0.78	0.42	0.87	-3.01
Eq. 23 ³	$y = 11 \times (SAVI^2) + 0.2$	0.81	0.40	0.89	0.04
Eq. 24	$y = 6.5 \times (EVI) - 0.4$	0.78	0.42	0.88	-5.71
Eq. 25	$y = 4.9 \times (NDVI^2) + 0.1$	0.80	0.41	0.89	4.39

¹ Significant at $p = 0.05$

² RMSE is in $m^2 m^{-2}$, and PBIAS is showed as percentage.

³ L-values in the SAVI calculations were 0.07 (for the PAI) and 0.37 (for the LAI).

343 Table 2 - Calibration of PAI and LAI models created through an iterative process using Landsat
344 reflectance data.

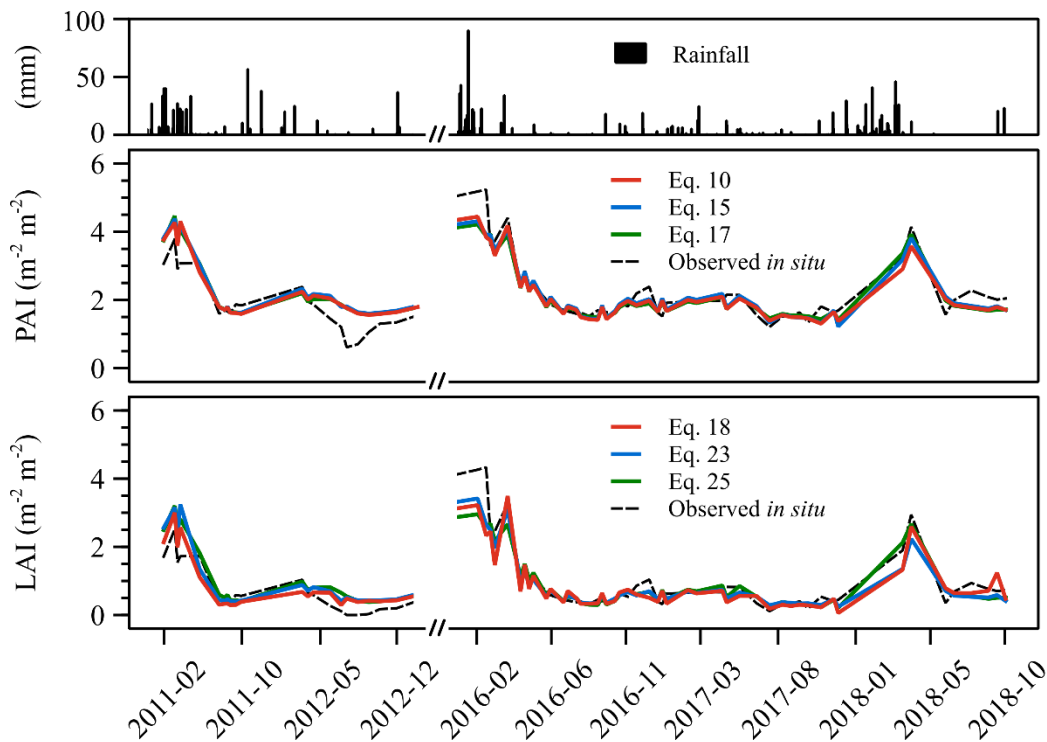
345 The best-performing new models that use different band combinations were Eqs. 10 to
346 14, and 18 to 22 (Table 2). These equations may represent other physical processes and
347 components, such as leaf etiolation and increasing protochlorophyll, which is reported to
348 influence the blue band of the visible spectrum (ρ_1 in Eq. 17) (Gates et al., 1965). Medeiros et
349 al. (2019) suggested the near-infrared (NIR) band may be a good indicator of leaf radiation
350 reflectance patterns among different species, which reflect variations in leaf size, form, and
351 type, and even plant habit. Our models used the NIR band, and they varied only to the inclusion
352 of the red, green, and blue bands. The amount of energy reflected or absorbed in these bands
353 varies according to the physicochemical and biophysical properties of the target (Edwards et
354 al., 2013). All bodies reflect or emit electromagnetic radiation at different wavelengths and in
355 different ways, and the result is a reflectance curve or spectral signature (Schmugge et al.,
356 2002). This set of unique interactions restricts the bands that distinguish certain characteristics
357 of a target and allows various parameters quantification (e.g., pigment concentration and plant
358 structure complexity) (Blackburn, 2007; Dawson et al., 1998; Schmugge et al., 2002). Usually,

359 vegetation reflects about half of the incident radiant flux in the NIR band (Zhao et al., 2007);
360 therefore, this is a band very sensitive to biomass and LAI. Leaves predominantly absorb
361 energy at the blue–red spectrum and reflect the energy in the green and NIR bands because of
362 the interaction with chlorophyll, carotenoids, and the mesophyll itself (Gates et al., 1965).
363 Thus, the green and NIR bands are considered bands of high reflectance (Fan et al., 2018). In
364 comparison to the green band, the NIR has a relatively higher multiple reflectance through
365 within-canopy layers, which reduces the canopy light extinction coefficient (Zheng and
366 Moskal, 2009).

367 We consider Eqs. 10, 15, and 17 to be optimal solutions for the estimation of PAI, and
368 Eqs. 18, 23, and 25 for the estimation of LAI in the Caatinga (Table 2). Although studies have
369 highlighted the dubious quality of data acquired by remote sensing in the blue band because of
370 wavelength-dependent atmospheric interference (e.g., Carter et al., 2009; Motohka et al.,
371 2009), Eq. 18 has performed very well with $r^2 = 0.79$ and $RMSE = 0.41 \text{ m}^2 \text{ m}^{-2}$, with values
372 comparable to Eq. 23 (which does not use the blue band) with $r^2 = 0.81$ and $RMSE = 0.40 \text{ m}^2$
373 m^{-2} . The greatest contribution of Eq. 18 is its natural proximity to a 1:1 relation to *in situ*
374 measurements ($\rho_c = 0.88$, $PBIAS = -0.01$), which provides greater ability to simulate values
375 near to zero. Although our models require observations in the NIR band, many images have
376 NIR sensors and, if well calibrated, they allow for LAI to be estimated based on reflectance
377 from spectral mixture or coarse resolution compositions. These images include those captured
378 by phenological cameras, unmanned aerial vehicles, and high-resolution monitoring satellites
379 (e.g., QuickBird and IKONOS).

380 The accuracy of our best models can be visualized when plotting their estimates
381 alongside observed PAI and LAI data (Fig. 3). Our best performance models were able to
382 emulate the variance of LAI in our study period (Table 2). Eqs. 10, 15 and 23 are biased towards
383 overestimation ($PBIAS = 0.33$, 1.10 and 0.04 respectively) and Eq. 18 presented a small

384 underestimation bias (PBIAS = -0.01). In general, based on our findings, models developed
385 with independent bands of a sensor produced low-magnitude values near-optimal zero,
386 indicating accurate model simulation, while the models created using vegetation indices
387 exhibited moderate bias.
388

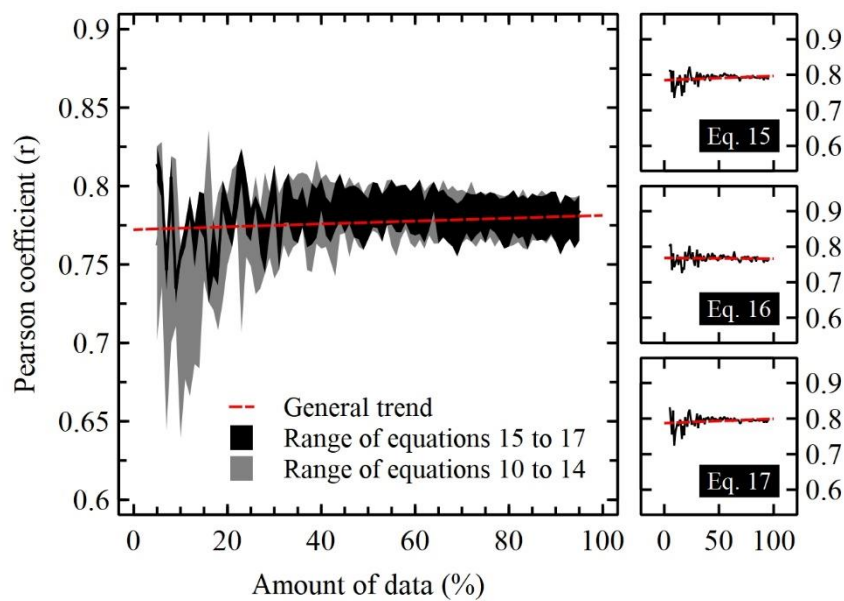


389
390 Figure 3. Comparison of temporal variation of PAI (in the middle) and LAI (bottom) observed
391 *in situ* to the best simulated models created through an iterative process using Landsat
392 reflectance data.

393 In the cross-validation analyses, Eq. 10 produced maximum values ($r_{max} = 0.81$) and
394 minimum values ($r_{min} = 0.68$) similar to Eq. 18 ($r_{max} = 0.83$ and $r_{min} = 0.59$). In comparison,
395 the *NDVI*, *SAVI* and *EVI* models yielded higher maximum and minimum values for both PAI
396 ($r_{max} = 0.83$, 0.82 , and 0.80 , respectively; $r_{min} = 0.73$ for all models) and LAI ($r_{max} = 0.84$,
397 0.85 , and 0.83 ; $r_{min} = 0.71$, 0.71 , and 0.69 , respectively). The other models presented r values
398 ranging from 0.64 to 0.84 (for PAI), and 0.6 to 0.83 (for LAI). The 0.01 standard deviation of

399 r was the same for all models, indicating that they are reliable. When varying the amount of
400 data taken for cross-validation from 5 to 95%, the r values tend to show minimal variations
401 (Figs. 4 and 5). However, in our verification we did not observe any statistically significant
402 pattern in correlation owing to the removal of data from the calibration of the models, which
403 confirms that these models are highly robust to estimate LAI.

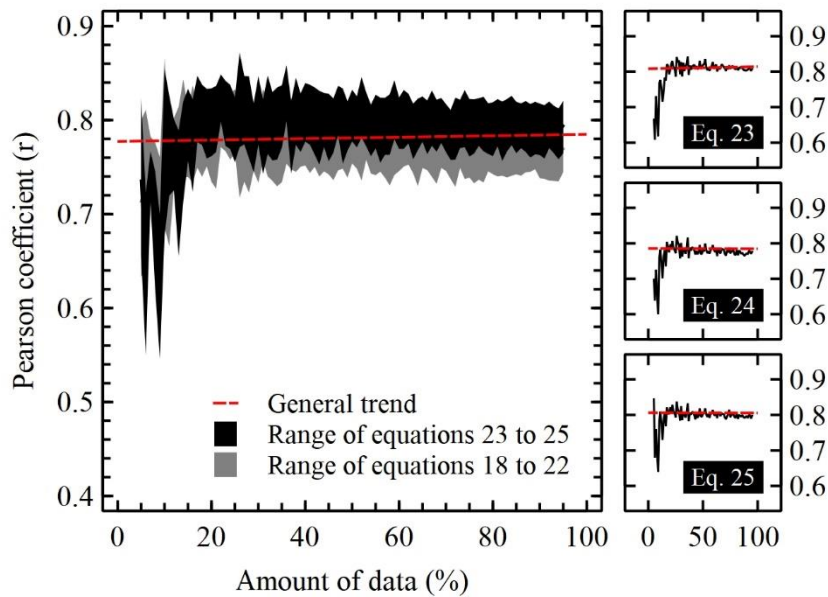
404



405

406 Figure 4. Cross-validation of the PAI models created through an iterative process using Landsat
407 reflectance data. Detailed validations of Eq. 15 to 17 are on the right side.

408



409

410 Figure 5. Cross-validation of the LAI models created through an iterative process using Landsat
411 reflectance data. Detailed cross-validations of Eqs. 23 to 25 are on the right side.

412 The approaches proposed by Bastiaanssen (1998) (Eq. 6), Galvncio et al. (2013) (Eq.
413 7), Machado (2014) (Eq. 8), and Almeida et al. (2019) (Eq. 9), as well as the MCD15A3H/A2H
414 data underperformed compared to our models, when compared to our *in situ* PAI and LAI data
415 (Table 3), even though correlations were significant ($p < 0.05$). Eqs. 6 and 7 performed the best
416 in terms of accuracy, while Eqs. 8 and 9 presented the highest covariances. Although the
417 MODIS products are supposed to reproduce the LAI seasonality well, we observed that they
418 do not respond well for the Caatinga during the dry season; the lowest values were around 0.5
419 $\text{m}^2 \text{m}^{-2}$ when real LAI values were practically zero. We tested the fitness of MODIS products
420 against our *in situ* PAI and LAI datasets. The lower correlation with the MCD15A3H product
421 in comparison to the MCD15A2H product indicates that the proportion of high-quality data for
422 the dry forest area is lower for periods of composition of four days than for the eight-day
423 version (Table 3). These periods of composition are created from the highest value observed *in*
424 *situ*; thus, the greater the number of values available for the determination of LAI of each pixel,
425 the higher the probability of it being an accurate value. This is because data obtained by

426 satellites are influenced by a number of different atmospheric factors such as water vapor, cloud
 427 cover, and aerosols, thus any given bit of satellite data may not yield accurate results (Yang et
 428 al., 2006). Our results indicate that studies that rely on LAI from MODIS products for
 429 vegetation assessment in SDTFs, such as the Caatinga, are likely to incorporate bias due to
 430 unrealistically high LAI values during the dry season, when plants lose most of their leaves
 431 and the vegetation consist predominantly of non-photosynthetic biomass (Leal et al., 2003;
 432 Silva et al., 2017). The reduction of photosynthesis has direct consequences for the
 433 evapotranspiration and gross primary productivity, affecting the carbon storage and CO₂
 434 exchange capacity (Morais et al., 2017; Nagler et al., 2003).

435

Reference	Parameter	r^{21}	RMSE ²	ρ_c	PBIAS ²
Bastiaanssen (1998)	PAI	0.73	1.54	0.35	-69.44
	LAI	0.75	0.51	0.84	-26.53
Galvncio et al. (2013)	PAI	0.73	1.32	0.33	-57.44
	LAI	0.72	0.52	0.75	2.33
Machado (2014)	PAI	0.76	1.28	0.61	-40.48
	LAI	0.78	1.01	0.72	43.12
Almeida et al. (2019)	PAI	0.77	1.66	0.40	-73.74
	LAI	0.77	0.65	0.81	-36.80
MCD15A3H	PAI	0.66	1.39	0.29	-60.08
	LAI	0.65	0.57	0.71	-4.00
MCD15A2H	PAI	0.77	1.26	0.38	-55.07
	LAI	0.78	0.46	0.82	8.04
Summary of selected models					
Eq. 10	PAI	0.79	0.41	0.88	0.33
Eq. 15 ³	PAI	0.79	0.41	0.88	1.10
Eq. 17	PAI	0.79	0.41	0.89	1.04
Eq. 18	LAI	0.79	0.41	0.88	-0.01
Eq. 23 ³	LAI	0.81	0.40	0.89	0.04
Eq. 25	LAI	0.80	0.41	0.89	4.39

¹ Significant at $p = 0.05$

² RMSE is in $m^2 m^{-2}$, and PBIAS is showed as percentage.

³ L-values in the SAVI calculations were 0.07 (for the PAI) and 0.37 (for the LAI).

436 Table 3 - Comparison of the previously published models, and the MCD15A3H/A2H products
 437 with *in situ* data.

438 The LAI in this study, as seen in all models for the Caatinga, can be defined as effective
439 LAI, which is the portion of LAI that effectively intercepts the light, not directly considering
440 grouped foliage. This grouping of leaves can be quantified by a vegetation dispersion parameter
441 Ω (clumping index) (Nilson, 1971), which often can be determined by a random distribution
442 (Chen and Black, 1992). The “true” LAI is not easy to achieve; it requires intensive fieldwork
443 and systematic sampling, using all possible allometric relationships (Frazer et al., 1997; Weiss
444 et al., 2004). Since the approaches of estimating PAI and LAI used in our study are based on
445 the light extinction, the WAI values as a result of the difference between PAI and LAI are
446 likely to be underestimated when LAI is high (Nackaerts et al., 2000; Stenberg, 1996). This is
447 attributed to the fact that when LAI values are very high, the leaves cover the woody area and
448 reduce the role of light interception of the branches and stems (Chen et al., 1997), which in
449 turn leaves the PAI and LAI values very similar (e.g., Feb 2012 in Fig. 3).

450 Our models are easy-to-use PAI and LAI predictors that can be applied to estimate these
451 indices for the Caatinga. The models also can be used to simulate other Caatinga types (such
452 as in transitional areas), but since they rely on calibration coefficients, minor adjustments might
453 be required to approximate minimum and maximum LAI. Regional applicability can be
454 considered as moderate-high, because the phyto-physiognomy dominated by shrubs is the main
455 and most abundant in the Caatinga (Silva et al., 2017). However, at a regional scale, our models
456 may be used as backup models in a physical approach that does not require calibration to
457 achieve maximum generalization. Further improvements may include (i) pooling coefficients
458 adjusted for other areas of Caatinga with different levels of degradation, which could be similar
459 to what was made by Bastiaanssen (1998) when developing LAI models; (ii) the adjustment of
460 these equations using field data from other types of Caatinga vegetation, where some plants,
461 such as Cactaceae and Bromeliaceae, may have a more significant presence, and the soil
462 exposure may be different; (iii) the removal of the influence of non-photosynthetic plant

463 material, such as flowers, fruits and petioles, on LAI measurements; and (iv) approximation of
464 LAI to more realistic values, developing and introducing a new Ω to more efficiently account
465 for leaf dispersion directly in the models, instead of abstracting it in regression coefficients.
466 This could solve systematic problems, such as misestimation of LAI at a given phenological
467 stage.

468 **4. Conclusions**

469 Our study developed and assessed several PAI and LAI models to be realistically representative
470 for the phenology of a typical Caatinga ecosystem. Given the high frequency of our *in situ*
471 measurements (mostly measured on a daily or weekly basis), all Caatinga phenological stages
472 were covered and reproduced in our models. The joint usage of ground and satellite data
473 presented an efficient way to assess both PAI and LAI models. The results included
474 parameterizations with the visible and infrared spectral bands, which allowed the use of many
475 currently available datasets to estimate LAI.

476 The models produced results with high accuracy (up to $r^2 = 0.81$ and $RMSE = 0.41 \text{ m}^2$
477 m^{-2}). The significant improvement of our models over the others used for the Caatinga is due
478 to the consideration of WAI, which previously had not been considered in calibrations for the
479 Caatinga, and the temporal variations of LAI, which allowed us to create more generalist
480 models that can be used during different phenological stages of the Caatinga vegetation.

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