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Journal:	<i>Journal of Glaciology</i>
Manuscript ID	Draft
Manuscript Type:	Communication
Date Submitted by the Author:	n/a
Complete List of Authors:	Mallett, Robbie; UiT The Arctic University of Norway, Physics and Technology
Keywords:	Snow, Sea ice, Snow physics
Abstract:	No abstract for communications

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Manuscripts

# A methodologically robust seasonal snow densification function from Soviet North Pole drifting station data

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Snow on sea ice plays a critical role in the polar ocean's energy balance, but also in satellite retrievals of sea ice thickness among other variables. The density of snow on sea ice evolves over the winter season, generally increasing as grains become rounder and the snowpack settles due to the effect of overburden. It is therefore desirable to form a simple equation for the snow density as a function of the time-of-year. In order to investigate the role of snow in radar-derived estimates of sea ice thickness, such an equation was put forward by Mallett and others (2020, henceforth M20):

$$\rho_s = 6.5t_m + 274.51 \quad (1)$$

Where  $\rho_s$  is the snow density in  $\text{kgm}^{-3}$ , and  $t_m$  is the number of months since October. The equation has now been used in several publications (e.g. Dong and others, 2022; Shi and others, 2023; Jiang and others, 2023; Dong and others, 2023; Sievers and others, 2023; Fredensborg Hansen and others, 2024; Chen and others, 2024).

Equation (1) was computed as follows: a large dataset of snow depth and snow water equivalent (SWE) was compiled from in-situ measurements at Soviet North Pole (NP) drifting stations by Warren and others (1999), and monthly quadratic fits were published for both variables. Following common practice in radar altimetry processing chains, M20 divided the quadratic fits for SWE by those for depth to produce spatial distributions for snow density. The spatial average of these density distributions in a subdomain of the Arctic Ocean was then computed, producing one mean snow density value for each winter month. These values were then regressed against the month number to generate Equation (1) of this manuscript. The above method has several drawbacks; their impact and remediation are the subject of this communication.

The first limitation of the method described above concerns the original quadratic fits for SWE and depth themselves, the parameters of which were published by Warren and others (1999). Values are

Mallett:

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26 sometimes negative in the marginal seas of the Arctic, and are not inherently “snow conserving” (i.e. the  
27 mean value in the Arctic Ocean is not inherently the mean value of the underlying values, particularly since  
28 the spatial definition of the Arctic Ocean is not well defined). Furthermore, it is sub-optimal to compute  
29 monthly spatial distributions for density by dividing those for SWE by those for depth: it would be better  
30 to compute the density distributions directly from the density measurements and their positions in the  
31 month concerned. Unfortunately the data to do this have not historically been easily accessible online.

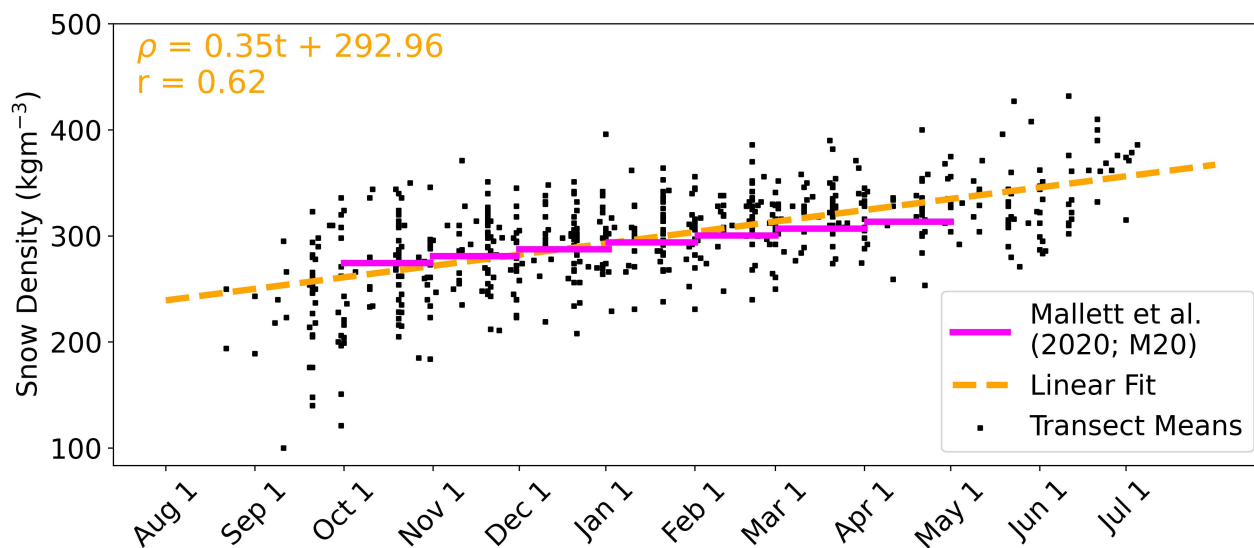
32 Further drawbacks exist in the averaging and regression process underpinning Equation 1: the area over  
33 which M20 averaged the density distributions in each month goes beyond that covered by the NP station  
34 data: for example, it includes the Laptev Sea which stations rarely visited. It was also only performed in  
35 the months of October - April, when the source data from NP stations potentially would allow a function  
36 to apply beyond those months. Finally,  $t$  in equation 1 represents the integer number of months since  
37 October, indicating that the formula is not weighted for the variable lengths of the winter months. In a  
38 sense, it is linear in month number, and thus not strictly linear in time.

39 All the methodological issues described above can be reduced (and some resolved), by directly regressing  
40 the mean densities calculated from the original transect data against the day-of-year on which they were  
41 generated. After some data cleaning (see below), this regression yields:

$$\rho_s = 0.35t_d + 292.96 \quad (2)$$

42 Where  $t_d$  is the day of year (with a breakpoint on 1st August), and  $\rho_s$  remains the snow density in  
43  $\text{kgm}^{-3}$  as in Equation 1. Five data points have been removed in the months of July and August, four of  
44 which are  $>500 \text{ kg}^{-3}$  and one of which is  $25 \text{ kgm}^{-3}$  (this is likely a measurement error). These extreme  
45 values exist near the break-point of the analysis, and their inclusion makes the slope of the regression  
46 highly sensitive to the choice of this date. Because of their removal, it is inadvisable to use generate snow  
47 densities from Equation 2 in these months. Despite this, it is clear that Equation 2 can sensibly be used to  
48 produce values outside of the “cold season” considered by the M20 calculation, for instance in September,  
49 May and June.

50 Figure 1 also makes clear that the new regression slope is not very different from the M20 function  
51 in a quantitative sense. Density calculations in the publications cited above using M20 can therefore be  
52 trusted. So why make a new one? The first reason is that the new, simpler, more robust methodology can  
53 be better trusted in future to represent the underlying data, and in more months of the year. In addition,



**Fig. 1.** Transect-mean snow densities ( $n=573$ ; black scatter), with the M20 values shown as magenta lines. Linear regression through the scatter points shown in orange. Where possible, NP station transects were performed at ten day intervals on the 10th, 20th and 10th of each month, generating a periodic distribution of scatter along the time-axis. The Pearson  $r$  correlation coefficient (0.62) is listed for completeness.

54 the new function also takes a continuous input of day-of-year rather than the month number, aiding its  
 55 utility as described above.

56 This new densification function retains some key limitations. It still relies on data collected by  
 57 Soviet NP drifting stations that operated on multiyear ice, and overwhelmingly in the Central Arctic,  
 58 East Siberian and Chukchi seas (See Figure 2 of Mallett and others, 2021, for trajectories of stations  
 59 contributing measurements to this analysis). Snow in the multiyear ice environment may well have a  
 60 different densification rate to that in the first-year ice environment due to its relative lack of salinity and  
 61 the rougher underlying ice. Relatedly, the high latitude of the measurements means that the densification  
 62 rate in Equation 2 may not reflect that of lower latitudes where periods of diurnal cycling are more  
 63 protracted and temperatures are often higher.

#### 64 Code and Data Availability

65 All code and data required to reproduce this analysis can be downloaded from:

66 <https://github.com/robbiemallett/densification>.

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