

1 Title: Natural and human drivers of specific conductance and major ion composition in United
2 States lakes

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25 **Author contribution statement:** Xinyu Sun and Kendra Spence Cheruvellil conceived the
26 project. Xinyu Sun, Kendra S. Cheruvellil, Patrick J. Hanly, Katherine E. Webster, and Patricia A.
27 Soranno designed the analyses and interpreted the results. Xinyu Sun gathered and processed
28 data. Xinyu Sun and Patrick J. Hanly performed analyses and made figures, tables, and
29 supplements. Xinyu Sun wrote the first draft of the manuscript, and all authors contributed
30 substantially to revisions.

31

32 **Abstract**

33 Specific conductance (SC) and major ion composition are important for understanding and
34 predicting lake water quality and ecosystem responses to global changes and human disturbances.
35 However, little is known about SC and ionic composition for populations of lakes at the
36 continental scale, nor their relationships with natural and human factors operating at multiple
37 spatial scales. We examined the spatial patterns in SC (N=9,784 lakes) and major anion and

38 cation concentrations (N=1,218 lakes) across the conterminous United States, and quantified
39 their relationships with a wide range of multi-scaled natural and human factors. We found
40 substantial spatial variation in ion composition and that lakes with similar SC values can have
41 very different ion composition. Most lakes had relatively low SC (median=206 μ S/cm), with
42 high-SC lakes mainly located in the Plains, Desert Southwest, and Florida. Calcium and
43 bicarbonate were the most common ions in 61% of the study lakes, with the remaining lakes
44 dominated by the cations magnesium or sodium and the anions sulfate or chloride. Lake SC was
45 associated with natural factors including elevation, watershed soils, and hydrology and was
46 influenced by watershed land uses. Ion composition was associated with similar natural factors
47 along with surface connectivity and precipitation, but also strongly affected by road density and
48 urban development. Our results suggest that while geological, hydrological, and climate
49 processes control the ion inputs from natural sources, human disturbances can cause SC and
50 major ions to deviate from their background levels.

51 **Keywords:** Salinity; Conductivity; Salts; Major ions; Geology; Hydrology; Road density;
52 Macroscale

53 **Introduction**

54 Specific conductance (SC) is a critical indicator of lake water quality due to its profound
55 effects on aquatic organisms, drinking water quality, and industrial and recreational water use
56 (Kaushal et al. 2018; Hintz & Relyea 2019; Dugan 2024). SC, which is often used as a proxy for
57 salinity (Calver et al. 2009), has been fluctuating greatly in freshwater ecosystems in some
58 regions of the world with resultant negative impacts; these fluctuations have been attributed to
59 natural and human disturbances such as climate change, extreme climatic events, road salt
60 application, and agricultural activities (Kaushal et al. 2019; Olson 2019; Schacht et al. 2023). A
61 few studies have documented changes in concentrations of individual ions such as chloride in
62 lakes due to human activities (e.g., Dugan et al. 2017; Kaushal et al. 2018). However, we lack a
63 macroscale understanding of the variation of SC across broad ranges of lakes, how ionic
64 composition broadly differs among lakes, and the potential influences of multi-scaled natural
65 environmental and human activities.

66 Studies focusing solely on SC are necessary but not sufficient to build this understanding
67 because the impacts of salts on freshwater communities are related to major ion concentrations
68 and compositions as well as SC (Cañedo-Argüelles et al. 2016). Lab and microcosm experiments
69 have shown that water with the same SC but different ion compositions (e.g., solutions or media
70 dominated by chloride (Cl^-) vs. sulfate (SO_4^{2-}) ions) can lead to divergent effects on organisms
71 and communities (Nostro et al. 2005; Clements & Kotalik 2016; Van Gray & Ayayee 2024).
72 Multiple ions can also interact synergistically to affect aquatic organisms, leading to unexpected
73 outcomes (Elphick et al. 2011). For example, the toxicity of Cl^- on cladocerans was greater in
74 softer water than in hard water (Elphick et al. 2011; Rogalski et al. 2024), and the toxicities of
75 potassium ion (K^+) could be alleviated by high sodium ion (Na^+) concentrations (Mount et al.
76 2016). These findings underscore the necessity of understanding major ion concentrations and

77 composition in lakes for a better assessment and prediction of ions and SC effects on freshwater
78 ecosystems.

79 Variation in the SC and major ions of surface waters has been related to both natural and
80 human factors. For example, SC in streams across the conterminous United States (CONUS)
81 ranged from extremely low ($<2 \mu\text{S}/\text{cm}$) to saline ($>10,000 \mu\text{S}/\text{cm}$), depending on the dominant
82 hydrologic and geologic sources and/or natural ecological context setting (e.g., evaporation;
83 magnesium (Mg^{2+}), calcium (Ca^{2+}), and SO_4^{2-} ions from weathering of rocks; and Na^+ and Cl^-
84 ions from saline groundwater) (Gibbs 1970; Griffith 2014; Olson & Cormier 2019). Studies of
85 lakes in south-central North Dakota revealed that within-region SC and major ion concentrations
86 were related to lake elevation, soils, and groundwater (Swanson et al. 1988). La Baugh et al.
87 (2000) found that SC in lakes and wetlands in central North America varied with evaporation,
88 precipitation, and groundwater fluxes. Landscape position, the location of a lake within
89 hydrologic flowpaths (Kratz et al. 1997), has also been found to affect SC and ions in lakes.
90 Regional studies of lake chains, one measure of landscape position, found that lakes connected to
91 other water bodies (having higher surface connectivity) had higher SC than less connected lakes
92 (Martin & Soranno 2006; Soranno et al. 1999). In addition, lakes without an outflow can be sinks
93 of ions for catchments in some evaporative regions resulting in elevated SC (Saleem et al. 2015;
94 Ding et al. 2024). Moreover, lakes in or near coastal areas can receive substantial amounts of salt
95 inputs from sea-salt influenced precipitation and saline groundwater intrusion (Kiflai et al. 2022;
96 Haque 2023). Human activities and urban development can sometimes cause long-term or
97 permanent changes to SC and ion composition. For instance, salt inputs from irrigation runoff,
98 residential discharge, winter road salt application, wastewater effluents, and mining can increase
99 SC (Oswald et al. 2019; Stets et al. 2020; Dumelle et al. 2024). In particular, agricultural
100 effluents often contain high concentrations of K^+ , Mg^{2+} , Cl^- , and SO_4^{2-} ; road deicing salts can

101 contribute significant amounts of Na^+ , Mg^{2+} , Ca^{2+} , and Cl^- ; and industrial runoff can contribute
102 SO_4^{2-} (Meybeck 2003; Dugan et al. 2017; Dugan 2024).

103 Previous studies of SC and major ions in freshwater systems generally have only
104 included a limited number of predictors, have not considered a range of human influences, and
105 have predominantly been focused on individual waterbodies or watersheds, with broad-scaled
106 studies on streams and rivers, but not lakes (e.g., Meybeck 2003; Griffith 2014; Stets et al. 2020;
107 but see Dugan 2024). Unfortunately, findings from stream studies are not always directly
108 applicable to lakes because streams and lakes differ in morphometry and hydrological pathways
109 that determine water residence time, ion retention time, and evaporation, all of which strongly
110 influence water chemistry (Lottig et al. 2011; Kahlert & Gottschalk 2014). Macroscale studies
111 that incorporate a wide range of multi-scaled (local to regional) factors are therefore required to
112 understand the spatial variation of lake SC and major ions and the influences of natural and
113 human factors on them. This knowledge is critical to establish reasonable standards and goals for
114 management and habitat restoration, as well as to predict lake responses to future changes.

115 In this study, we investigated two questions about lakes across the broad ranges of
116 climate, hydrology, and land use of the CONUS: 1) What are the macroscale spatial patterns and
117 associations among lake SC and major ion concentrations and composition? and 2) What multi-
118 scaled natural and human factors are most strongly related to them? Using water chemistry and
119 ecological context data from multiple data sources for 9,784 (SC) and 1,218 (major ions) lakes of
120 the CONUS, we answered these questions by documenting spatial variation, and conducting
121 random forest and GLMNET analyses to examine the relationships between natural and human
122 factors and both SC and major ions.

123

124 **Methodology**

125 *Data collection*

126 We used data from the LAGOS-US research platform (Cheruvilil et al. 2021) that
127 includes lake, natural environmental, and human activity data for 479,950 lakes ≥ 1 ha surface
128 area across CONUS. We obtained measured epilimnion SC, individual ion concentrations for the
129 cations Ca^{2+} , K^+ , Mg^{2+} , and Na^+ and the anions Cl^- and SO_4^{2-} , and alkalinity (as CaCO_3) from the
130 LAGOS-US LIMNO module (Shuvo et al. 2023). The LIMNO module includes lake surface
131 water quality data from the US Water Quality Portal (WQP; 2021); the 2007, 2012, and 2017 US
132 National Lakes Assessment Surveys (NLA; US Environmental Protection Agency (EPA) 2010,
133 2016, 2022); and the US National Ecological Observatory Network (NEON; Keller et al. 2008).
134 In this study, we used epilimnetic SC (i.e., the electrical conductivity of one cubic centimeter of
135 solution at 25°C), a commonly measured constituent of inland waters, as the proxy for salinity
136 (Calver et al. 2009; Cañedo-Argüelles et al. 2016; Dugan 2024). For natural and human factors,
137 we used lake locational, morphometric, and surface water connectivity data from the LAGOS-
138 US LOCUS module (Cheruvilil et al. 2021; Smith et al. 2021); data about the natural context
139 (e.g., soil texture and climate features) and human factors (e.g., road density and land use) from
140 the LAGOS-US GEO module (Smith et al. 2022); and lake reservoir designation (natural lake or
141 reservoir) from the LAGOS-US RESERVOIR module (Polus et al. 2022; Rodriguez et al. 2023).
142 Additionally, we acquired evapotranspiration and snowpack water equivalent storage data from
143 Blodgett (2023) and livestock manure application data from US EPA EnviroAtlas (2015). We
144 used NEON regions, which are ecoregions classified based primarily on climate (213,800-
145 770,995 km^2 ; Hargrove & Hoffman 1999), as a spatial delineation to account for large-scale
146 geographical variation in the analyses. We removed highly correlated variables, resulting in 46
147 multi-scaled natural and human factors (see Table S1 for a list of factors).

148 *Data processing*

149 We used 2000-2021 epilimnetic measurements of lake SC, major ion concentrations, and
150 alkalinity. Because we had much more SC than ion data, we created two datasets for further
151 analyses: a full dataset with only SC data and a sub-dataset with SC and complete major ion data.
152 We applied water quality QA/QC procedures from LAGOS-NE LIMNO (v. 1.087.3; Soranno et
153 al. 2019) to these datasets. Specifically, we removed lakes with long-term average SC lower than
154 $2 \mu\text{S}/\text{cm}$ (i.e., values so low they could indicate potential measurement errors) and higher than
155 the upper threshold for outliers (i.e., 75th percentile + $15 \times$ interquartile range = $4,969 \mu\text{S}/\text{cm}$),
156 resulting in 11,072 lakes with SC data. We further extracted the SC data collected between April
157 and October (90% of data) during the most recent sampling year for each lake (median year =
158 2015; Figure S1a), computed the mean SC of each lake, and merged those data with multi-scaled
159 natural and human factors, yielding 9,784 lakes with complete data (i.e., the full dataset with no
160 missing values for any factor) that are representative of CONUS lakes (Figure S2). For the 4,581
161 lakes with more than one SC sample in the latest sampling year from April to October, we
162 calculated the coefficient of variation (CV) to represent intra-year temporal variation in SC.
163 These lakes were often sampled from June to September (72% of the samples) with a sampling
164 frequency range of 2-7 times.

165 For the sub-dataset combining SC, ions, and alkalinity data, we extracted April to
166 October data ($\geq 85\%$ of data) and used the most recent concurrent (i.e., taken the same day)
167 samples available (one water sample per lake for 1,498 lakes; median year = 2016). Next, to
168 investigate ion composition in lakes, we converted ion concentrations and alkalinity reported in
169 mg/L units in LIMNO-US to microequivalents per liter ($\mu\text{eq}/\text{L}$; Table S2). In this study, we used
170 bicarbonate (HCO_3^-) (in $\mu\text{eq}/\text{L}$) to represent all carbonate forms of alkalinity including CO_3^{2-} ,
171 which dominates in high pH waters. Major ion data were combined with natural and human

172 factors to generate a 1,218 lakes sub-dataset, of which about 85% were sampled by the National
173 Lakes Assessments (Figure S1).

174 *Data analyses*

175 Data analyses were conducted in R (v4.3.3; R Core Team, 2024). To identify spatial
176 patterns of SC using the full dataset, we plotted the mean and CV of SC over the 17 NEON
177 regions for CONUS. To identify the spatial distribution of lakes with relatively low and high SC,
178 we selected lakes with SC lower than 10% and higher than 90% of all lakes and mapped them
179 with the NEON regions. Next, using the sub-dataset, we calculated the SC and total ion
180 concentration (in mg/L) for lakes with complete major ion data. We computed SC by multiplying
181 the equivalent concentration (in $\mu\text{eq/L}$) of each ion by its corresponding equivalent conductivity
182 (Table S2) and summing up the results. The total ion concentration was calculated by summing
183 up the concentration (in mg/L) of each ion. We then applied linear models to examine the
184 correlations between measured SC and calculated SC, total ion concentration, and major ion
185 concentrations.

186 We used Boruta feature selection ('Boruta' package, v8.0.0, Kurasa & Rudnicki 2022) and
187 random forest (RF, 'randomForest' package, v4.7-1.1, Cutler & Wiener 2022) to examine which
188 and how multi-scaled natural and human factors affect SC in lakes. For these factors, a natural
189 log transformation was applied to non-percent data, and a generalized logit transformation was
190 applied to percent data (Table S1). Two Boruta feature selections with a maximum of 1,000 runs
191 were performed using SC values as the response variable and either natural or human factors as
192 predictors. We ranked natural and human factors separately based on Boruta importance scores,
193 then took the first half (i.e., the top half of factors based on importance) from each and input
194 them into an RF model with 5-fold repeated cross-validation to examine the important natural
195 and human factors that affect lake SC.

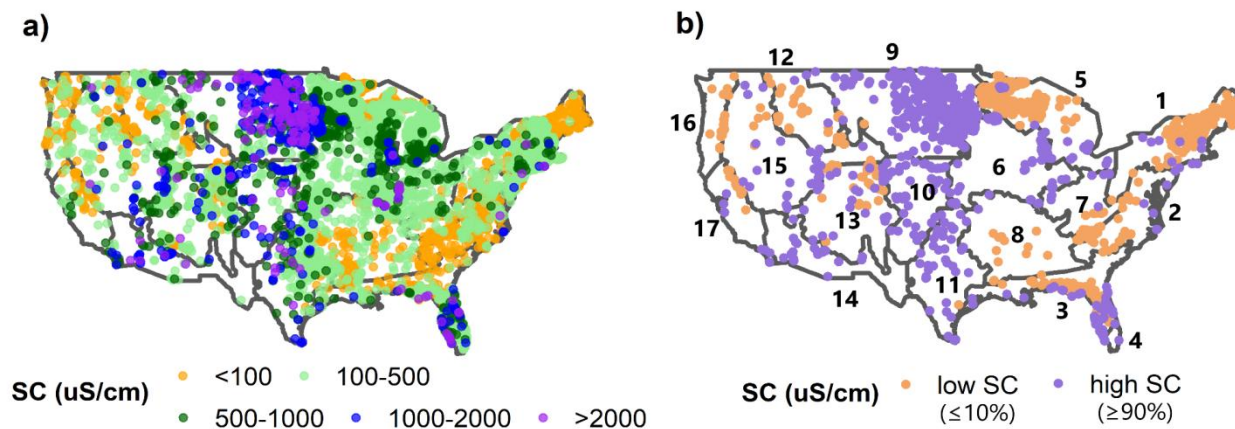
196 The effects of the important factors were assessed through partial dependence plots
197 (PDPs; 'pdp' package, v0.8.1, Greenwell 2022). We identified important factors for PDPs using
198 both the percentage increase in mean squared error (% increase MSE) and the increase in node
199 purity. The % increase MSE represents the increase in MSE when a factor is excluded, and the
200 node purity indicates the before-after change in the residual sum of squares at a splitting node
201 (Cutler & Wiener 2022).

202 To study the spatial patterns of how multi-scaled natural and human factors affect ion
203 composition using the sub-dataset, we first applied hierarchical clustering on \log_{10} -transformed
204 ion equivalent concentrations to identify common patterns among ion concentrations and
205 composition in lakes (Ward's method; Härdle & Simar 2019). We identified 15 clusters (Figure
206 S3) and used those clusters as the response variable in later analyses. Next, we ran a Boruta
207 feature selection using all natural and human factors as predictors and removed the unimportant
208 factors identified by Boruta. We then centered and scaled (into z-scores) all continuous factors
209 before running a multinomial GLMNET model with LASSO regularization ('glmnet' package,
210 v4.1-8, Friedman et al. 2023). Each natural and human factor was assigned to one of 10
211 categories: climate, hydrology, lake and watershed (morphometry), lithology, location, soil,
212 surface connectivity, terrain, human activities, and land use/land cover (LU/LC). We calculated
213 the relative importance of each factor for predicting ion cluster membership by summing the
214 absolute value of the factor's multinomial coefficient across all clusters. The top six factors were
215 then selected for visualization. Finally, to further characterize the clusters, we grouped clusters
216 based on dominant cations and anions (i.e., dominated by Ca^{2+} and HCO_3^- ; Mg^{2+} , Ca^{2+} , and
217 HCO_3^- ; Ca^{2+} , Na^+ , Cl^- , and HCO_3^- ; Mg^{2+} and SO_4^{2-} ; or Na^+ and Cl^-) and applied Wilcoxon tests to
218 examine the differences in the top six factors between each cluster group and the group
219 dominated by Ca^{2+} and HCO_3^- , which was the most common ion composition.

220

221 **Results**222 *SC in lakes across the CONUS*

223 We found that SC varied within and among regions for the 9,784 lakes in the full dataset,
224 much more so than within years. SC ranged from 2.0 to 6,125 $\mu\text{S}/\text{cm}$ SC (mean \pm standard
225 deviation (SD)=343 \pm 511 $\mu\text{S}/\text{cm}$, median=206 $\mu\text{S}/\text{cm}$). Most lakes with relatively high SC (SC
226 higher than 90% of all lakes; SC \geq 696 $\mu\text{S}/\text{cm}$) were located in the Southeast, Prairie Peninsula,
227 Northern Plains, Central Plains, Southern Plains, and Desert Southwest NEON regions (Figure 1).
228 Of these, 202 lakes, mostly located in the Northern Plains, had SC higher than 2,000 $\mu\text{S}/\text{cm}$.
229 Lakes with relatively low SC (SC lower than 10% of all lakes; SC \leq 34 $\mu\text{S}/\text{cm}$) were mostly
230 found in the Northeast, Mid-Atlantic, Southeast, and Great Lakes regions. Among the 17 NEON
231 regions, the Desert Southwest lakes had the highest average SC (1,154 $\mu\text{S}/\text{cm}$), followed by
232 those in the Northern Plains (897 $\mu\text{S}/\text{cm}$), and the Central Plains (894 $\mu\text{S}/\text{cm}$). In contrast, the
233 Pacific Northwest had the lowest mean SC (72 $\mu\text{S}/\text{cm}$), followed by the Northeast (139 $\mu\text{S}/\text{cm}$)
234 and Mid-Atlantic (146 $\mu\text{S}/\text{cm}$) regions. In contrast, most of the lakes with multiple sampling
235 dates within a year had low intra-year temporal variation in SC (CV: mean \pm SD=10% \pm 13%,
236 median=6%), and those with high CV (greater than 100%) were predominantly found in the
237 Northeast, Southeast, and Great Lakes regions (Figure S4).



238

239 **Figure 1.** The mean April-October specific conductance (SC) of each study lake (a; N=9,784)

240 and the location of lakes with low and high SC values (b; 978 lakes with SC lower than 10% of

241 lakes and 979 lakes with SC higher than 90% of lakes, respectively). NEON regions outlined in

242 black are as follows: 1 = Northeast, 2 = Mid-Atlantic, 3 = Southeast, 4 = Atlantic Neotropical, 5

243 = Great Lakes, 6 = Prairie Peninsula, 7 = Appalachians & Cumberland Plateau, 8 = Ozarks

244 Complex, 9 = Northern Plains, 10 = Central Plains, 11 = Southern Plains, 12 = Northern Rockies,

245 13 = Southern Rockies & Colorado Plateau, 14 = Desert Southwest, 15 = Great Basin, 16 =

246 Pacific Northwest, 17 = Pacific Southwest (numbers in plot b).

247

248 Measured and calculated SC were strongly positively correlated for the 1,218 lakes in the

249 sub-dataset with complete SC and major ion data (Figure 2a; $p < 0.001$). The average deviation of

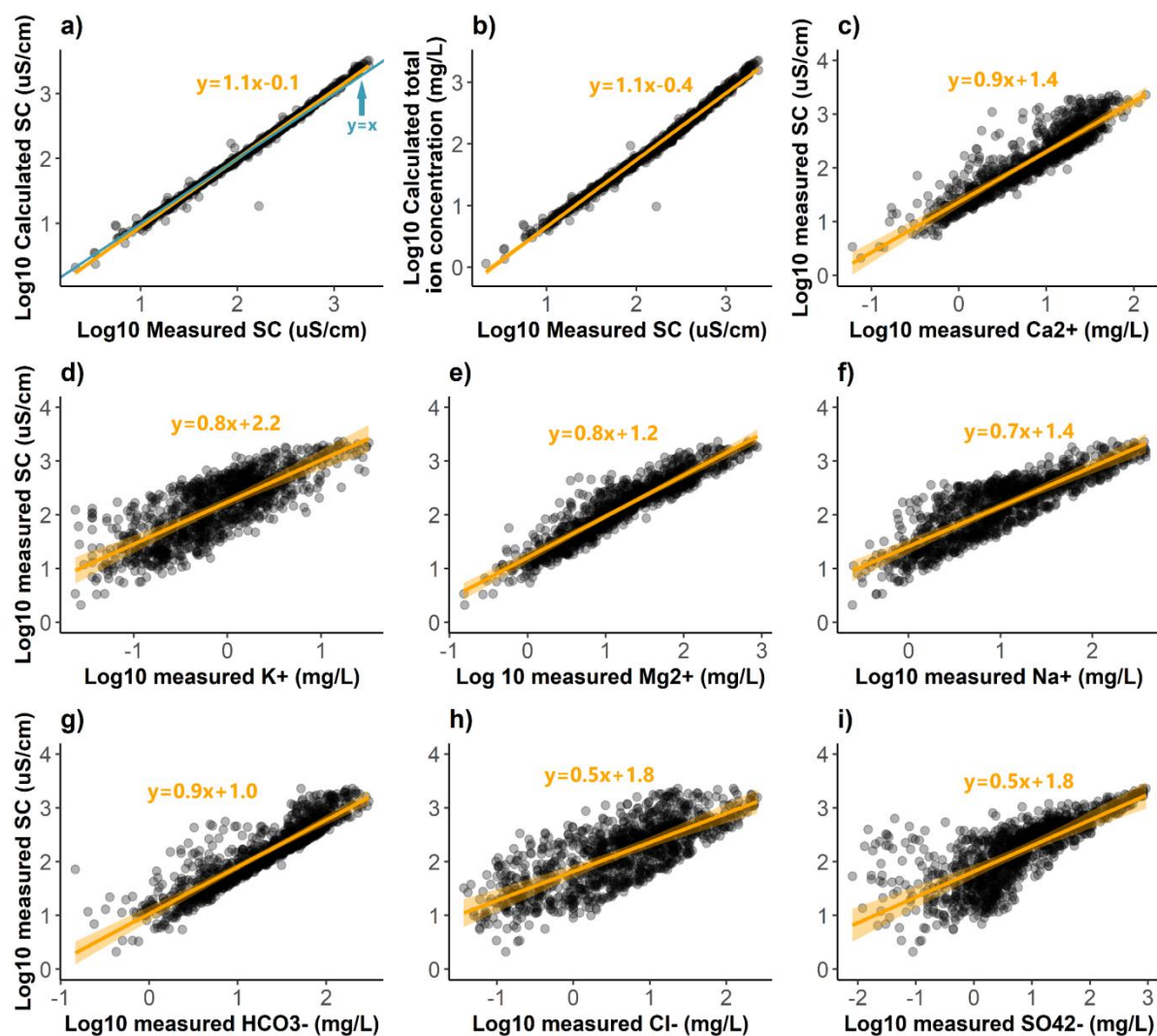
250 calculated to measured SC was about 8%, suggesting that our sub-dataset represented a majority

251 of dominant major ions. We found positive correlations between measured SC and the total ion

252 concentration and between measured SC and each major ion, with the strength of the relationship

253 (i.e., the slope) varying among ions (Figure 2b-i; $p < 0.001$). Specifically, Ca^{2+} , K^+ , Mg^{2+} , Na^+ ,

254 and HCO_3^- had stronger associations with SC compared to those shown for Cl^- and SO_4^{2-} .

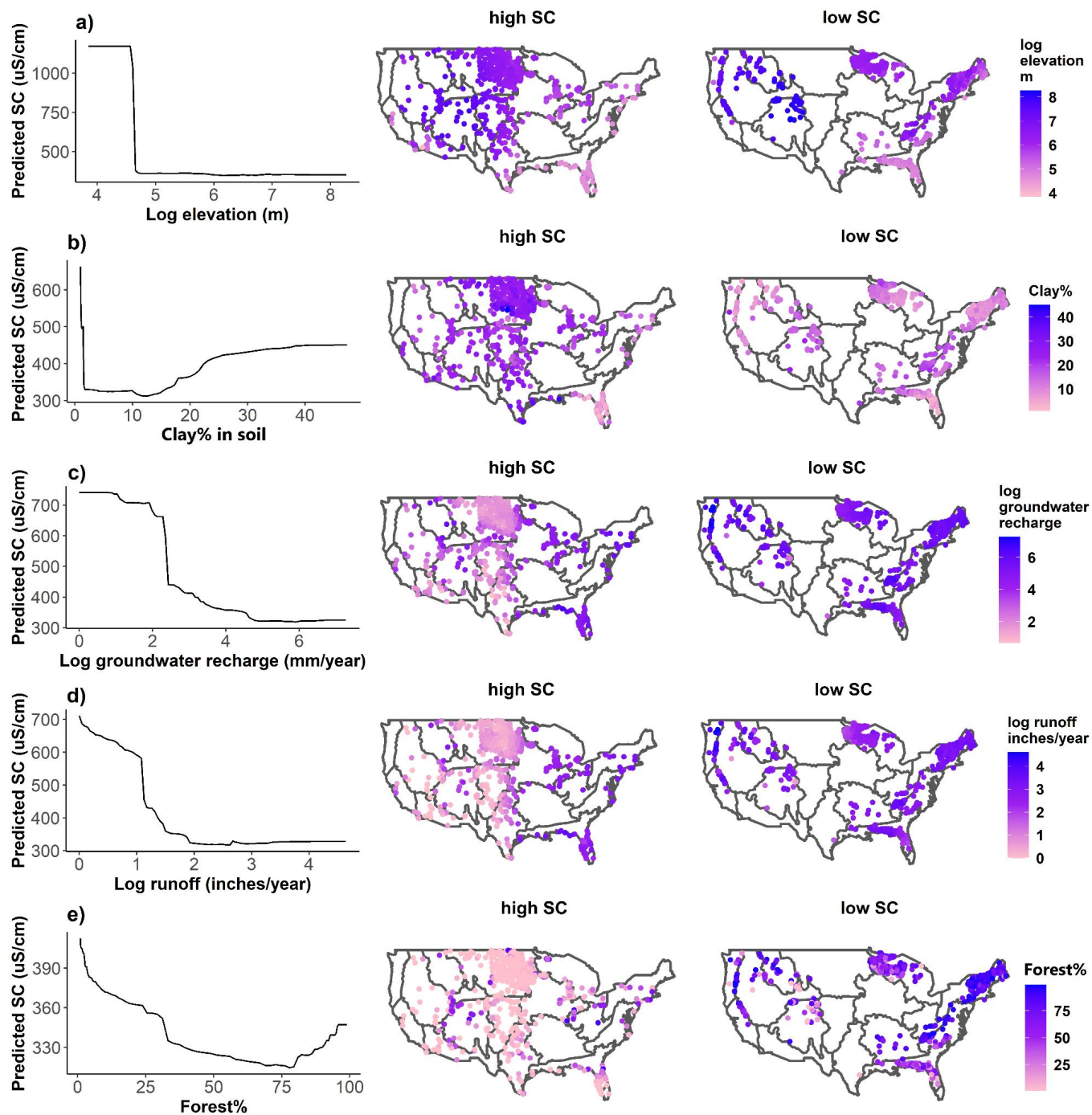


255
 256 **Figure 2.** Relationships between log₁₀-transformed measured SC and calculated SC (a),
 257 calculated total ion concentration (b), and the concentration of each ion (c-i) in 1,218 lakes
 258 across the CONUS. Each dot represents a lake and the orange prediction lines were fitted with a
 259 linear model with the equation above each line. In plot (a), the blue line is $y=x$. The orange
 260 shadow area indicates one standard deviation.
 261

262 *Multi-scaled natural and human factors related to SC*

263 Lake SC was related to both natural and human factors (RF out-of-bag variance
264 explained=59%; N=9,784). Lake elevation was the most important factor based on the RF
265 model % increase MSE, followed by the percentage of clay in the soil and the percentage of
266 watershed forest land cover (Figure S5a). Groundwater recharge was the most important factor
267 identified using the increase in node purity (Gini coefficient), followed by mean annual runoff
268 and percent clay in the soil (Figure S5b).

269 We observed a range of relationships between explanatory factors and lake SC using
270 partial dependence plots (PDPs). Lakes with the highest SC at the lowest elevations were
271 located mostly in the Southeast region. There was a subsequent steep decline in SC for lakes
272 located above 120 m; thereafter SC gradually declined and eventually became stable as lake
273 elevation increased (Figure 3a). The percentage of clay in the soil was negatively associated with
274 SC until 2% clay content, then the SC remained low and started to increase when clay content
275 exceeded 12% (Figure 3b). Most high-SC lakes had low groundwater recharge, indicating
276 dominance by surface water pathways, and a negative association was found between SC and
277 groundwater recharge until about 150 mm/year, at which point SC stabilized at a low level
278 (Figure 3c). Similarly, lake SC was negatively associated with annual runoff and became stable
279 at a low SC level when runoff exceeded about 20 inches/year (Figure 3d). Finally, we observed
280 that most high-SC lakes had low watershed forest land cover (which can represent areas with
281 high human disturbance), whereas SC was lowest in lakes with watersheds with moderate forest
282 cover. Lake SC, however, showed a small increase as forest cover reached about 80% (Figure
283 3e).



284

285 **Figure 3.** PDPs showing the relationships between predicted SC and the most important natural
 286 and human factors (left) and maps showing the values of the important factors of low SC (SC
 287 lower than 10% of all lakes; right) and high (SC higher than 90% of all lakes; middle) SC lakes.

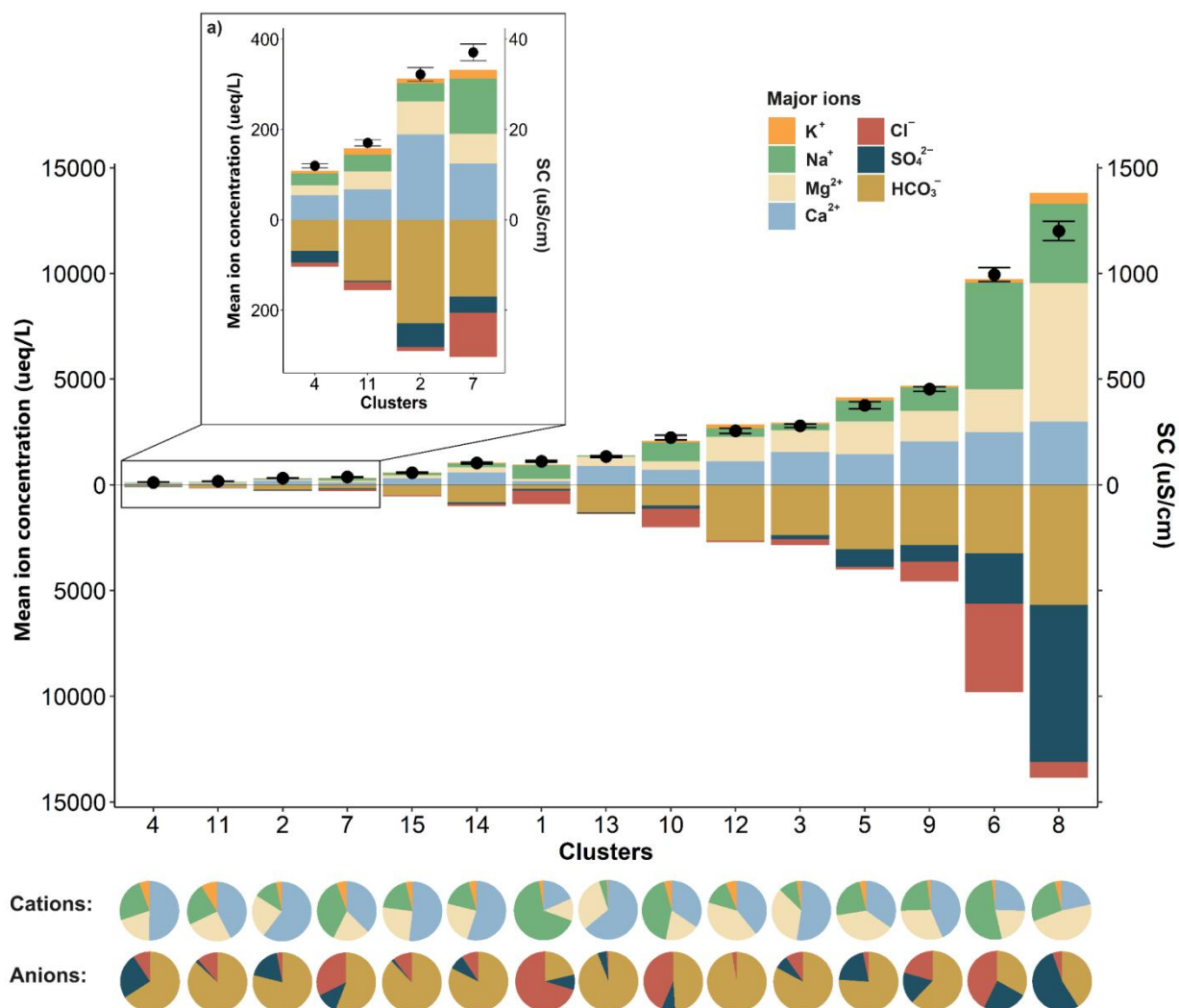
288 'Log' = natural log transformation.

289

290 *Ion composition in US lakes and associations with multi-scaled natural and human factors*

291 We hierarchically classified the 1,218 lakes in the sub-dataset with both SC and major
292 ion data into 15 well-defined clusters with divergent ion concentrations and ion compositions
293 (Figures 4, Table 1). In most clusters (61% of lakes; clusters 4, 11, 2, 15, 14, 13, 3, and 9,
294 ordered based on low to high mean SC within a cluster), and despite differing SC, Ca^{2+} and
295 HCO_3^- were the most abundant cation and anion, respectively. However, about 8% of lakes were
296 dominated by Mg^{2+} and SO_4^{2-} (cluster 8); Ca^{2+} and Mg^{2+} co-dominated in 8% of lakes (clusters
297 12 and 5); Na^+ and Cl^- were the most abundant cation and anion, respectively, in 5% of lakes
298 (clusters 6 and 1); and Ca^{2+} and Na^+ had similar proportions among cations in 17% of lakes
299 (clusters 7 and 10). Among all the clusters, cluster 8 had the highest SC, followed by clusters 6
300 and 9; and cluster 4 had the lowest SC, followed by clusters 11 and 2.

301



302

303 **Figure 4.** For each cluster, the mean equivalent concentrations ($\mu\text{eq/L}$) of major ions, mean SC

304 ($\mu\text{S/cm}$), and proportions of cations and anions. Stacked bars represent ion concentrations with

305 cations above and anions below the zero line, respectively, with corresponding values shown on

306 the left y-axis; black dots represent mean cluster SC and correspond to the right y-axis; error bars

307 represent one standard deviation in each direction; and pie charts at the bottom represent the

308 proportions of each cation and anion within a cluster. To aid in visualization, subplot a) shows

309 ion concentrations and SC in clusters 4, 11, 2, and 7.

310

311 **Table 1.** Summary of each cluster's dominant cation(s) and anion(s), percentage of study lakes
 312 within each cluster, and average SC of each cluster. Clusters are ordered according to dominant
 313 ions.

Cluster	Dominant cation(s)	Dominant anion(s)	Percentage of study lakes	Average SC \pm SD (μ S/cm)
4	Ca ²⁺	HCO ₃ ⁻	5	12 \pm 4
11	Ca ²⁺	HCO ₃ ⁻	2	17 \pm 7
2	Ca ²⁺	HCO ₃ ⁻	7	32 \pm 15
15	Ca ²⁺	HCO ₃ ⁻	5	57 \pm 26
14	Ca ²⁺	HCO ₃ ⁻	12	103 \pm 40
13	Ca ²⁺	HCO ₃ ⁻	2	134 \pm 44
3	Ca ²⁺	HCO ₃ ⁻	15	279 \pm 75
9	Ca ²⁺	HCO ₃ ⁻	13	453 \pm 111
12	Ca ²⁺ , Mg ²⁺	HCO ₃ ⁻	2	255 \pm 119
5	Ca ²⁺ , Mg ²⁺	HCO ₃ ⁻	6	377 \pm 172
7	Ca ²⁺ , Na ⁺	HCO ₃ ⁻	8	37 \pm 18
10	Ca ²⁺ , Na ⁺	Cl ⁻ , HCO ₃ ⁻	9	224 \pm 113
8	Mg ²⁺	SO ₄ ²⁻	8	1,201 \pm 448
1	Na ⁺	Cl ⁻	2	111 \pm 40
6	Na ⁺	Cl ⁻	3	995 \pm 330

314

315 The diverse spatial patterns in lake ion composition and cluster assignments of the study
316 lakes were determined by both natural and human factors (GLMNET; Figure 5; Figure S6; Table
317 S3). Ion clusters showed stronger associations with factors in the categories of soil, terrain,
318 climate, human activities, and LU/LC than those in other categories. Hydrology, particularly
319 groundwater-related features, and lake location were only moderately associated with ion
320 composition. Additionally, surface water connectivity was found to be associated with lake
321 membership in only a subset of ion clusters. Below we describe the spatial distribution of each
322 ion cluster as well as their characteristics using the top six natural and human factors (i.e.,
323 surface water connectivity, percent clay in the soil, precipitation, runoff, road density, and
324 percent forest land cover) from the GLMNET analysis.

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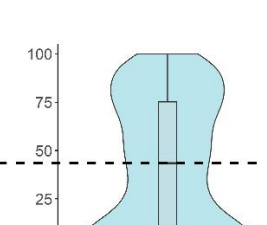
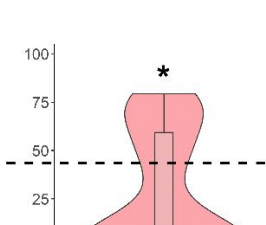
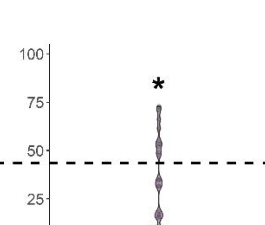
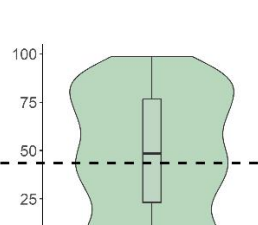
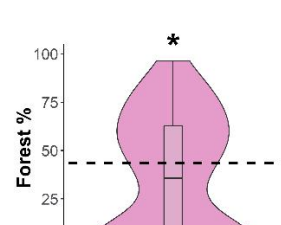
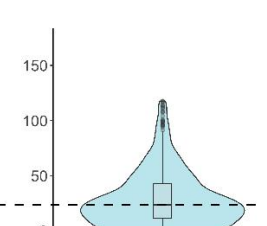
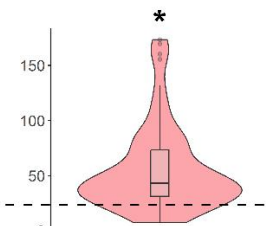
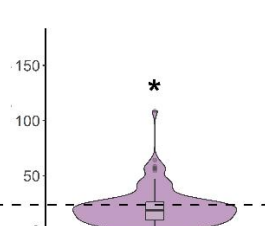
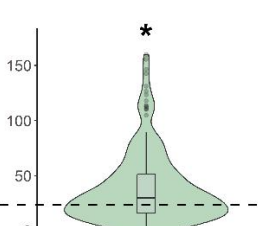
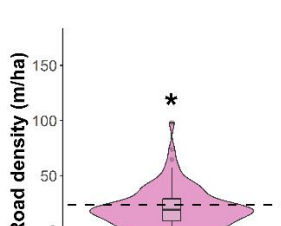
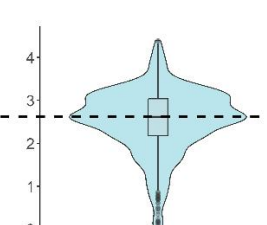
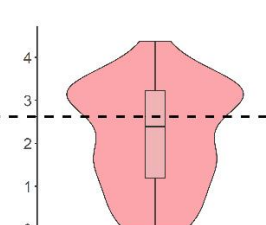
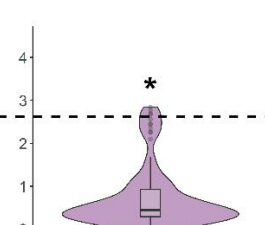
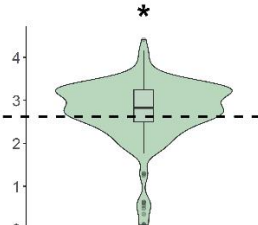
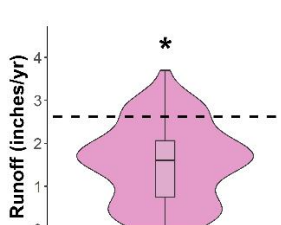
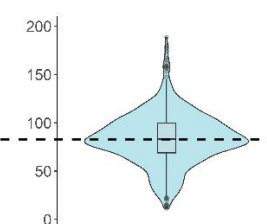
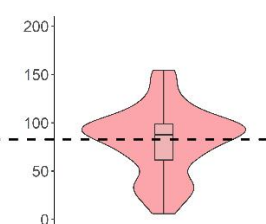
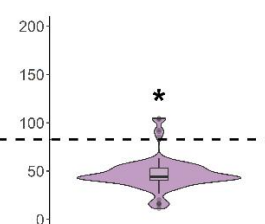
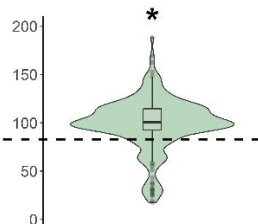
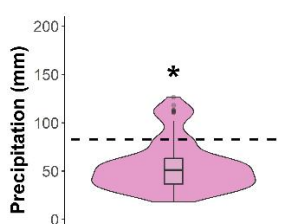
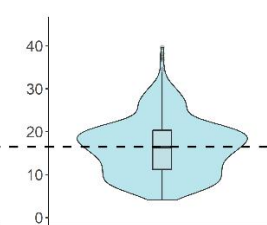
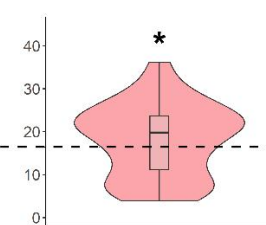
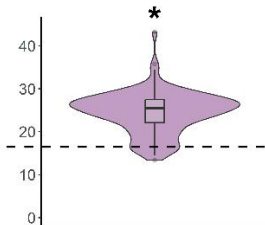
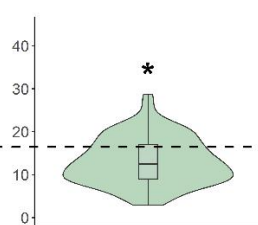
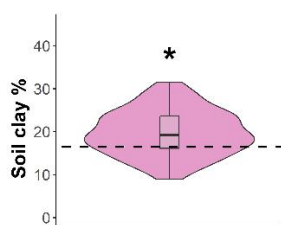
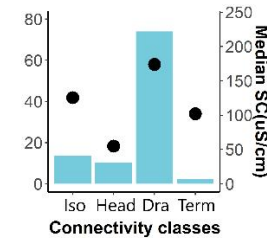
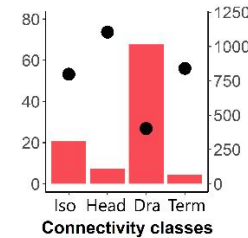
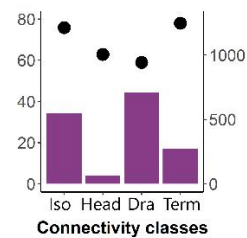
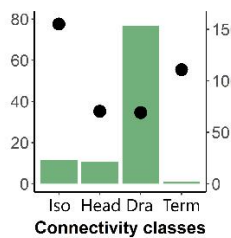
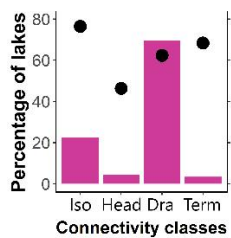
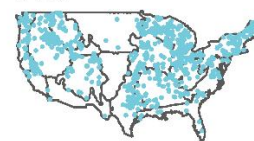
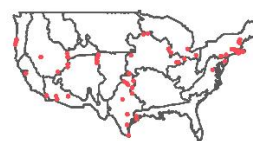
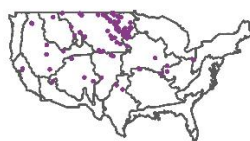
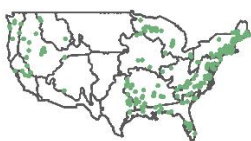
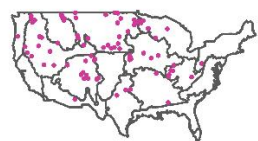
clusters 12 and 5
Ca, Mg, HCO₃
N=89

clusters 7 and 10
Ca, Na, Cl, HCO₃
N=214

cluster 8
Mg, SO₄
N=99

clusters 1 and 6
Na, Cl
N=68

clusters 4, 11, 2, 15, 14,
13, 3, and 9
Ca, HCO₃
N=748



327 **Figure 5.** Maps (top), barplots (middle), and violin plots (bottom) of lakes in clusters aggregated
328 into five major ion cluster groups based on the dominant cation(s) and anion(s). Maps show the
329 location of lakes within each ion group, with the text listing the ion clusters we included in each
330 major ion cluster group, the dominant cations and anions, and the number of lakes. Barplots
331 show the percentage of lakes (bars, left axis) and median SC (black dots, right axis) in the four
332 surface connectivity classes (Iso=isolated; Head=headwater; Dra=drainage; Term=terminal).
333 Violin plots show data distribution of top natural and human factors by major ion cluster group.
334 In the embedded boxplots, the bold line indicates the median and the upper and lower whiskers
335 indicate the 75th and 25th percentiles. The horizontal dashed lines are the median values of the
336 cluster group dominated by Ca^{2+} and HCO_3^- (i.e., the most common ion composition; clusters 4,
337 11, 2, 15, 14, 13, 3, and 9). The star above the violin plot indicates a significant difference
338 ($p \leq 0.05$) in that factor between the corresponding cluster group and the major ion cluster group
339 dominated by Ca^{2+} and HCO_3^- (the rightmost panel).

340
341 Lakes dominated by Ca^{2+} and HCO_3^- (hereafter referred to as the common ion group;
342 clusters 4, 11, 2, 15, 14, 13, 3, and 9) were widely distributed across multiple NEON regions,
343 except for the Northern Plains, Central Plains, Desert Southwest, and Pacific Southwest. Lakes
344 dominated by Mg^{2+} , Ca^{2+} , and HCO_3^- (clusters 12 and 5) were mostly found in the Northern
345 Plains, Prairie Peninsula, Southern Rockies & Colorado Plateau, and Great Basin. Compared to
346 the common ion group, lakes in these clusters had higher percentages of clay in the soil and
347 lower precipitation, runoff, road density, and forest cover (Wilcoxon, $p \leq 0.01$). Lakes dominated
348 by Ca^{2+} , Na^+ , Cl^- , and HCO_3^- (clusters 7 and 10) were commonly found in the Northeast, Mid-
349 Atlantic, Southeast, Atlantic Neotropical, Great Lakes, Ozarks Complex, and Great Basin. These
350 lakes had lower percentages of clay and higher precipitation, runoff, and road density than the

351 common ion group ($p < 0.001$). Lakes dominated by Mg^{2+} and SO_4^{2-} (cluster 8) were mostly found
352 in the Northern Plains. These lakes were distinct in their surface connectivity classes as, this
353 cluster had the highest proportions of lakes with no inflow and no outflow (i.e., isolated lakes;
354 34% of all cluster 8 lakes) or with only inflow (i.e., terminal lakes; 17% of all cluster 8 lakes)
355 (Table S4). Additionally, lakes in this cluster had higher percentages of clay and lower
356 precipitation, runoff, road density, and forest cover compared with the common ion group
357 ($p \leq 0.003$). Finally, lakes dominated by Na^+ and Cl^- (clusters 1 and 6) were commonly found in
358 the Northeast, Mid-Atlantic, Prairie Peninsula, and Southern Plains. These lakes had higher
359 percentages of clay and road density and lower forest cover than the common ion group
360 ($p \leq 0.04$).

361

362 **Discussion**

363 We describe general spatial patterns in SC and major ion composition across the broad
364 ecoclimatic zones of the conterminous US lakes that help provide context for interpreting the
365 potential for future changes. We also examine the relationships between multi-scaled natural and
366 human factors and SC and major ions that can be used to predict and support the management of
367 lake water quality. Although most lakes had low SC (median=206 $\mu S/cm$), 2% of the lakes had
368 SC higher than 2,000 $\mu S/cm$. Sixty-one percent of lakes were dominated by Ca^{2+} and HCO_3^- and
369 the remaining were dominated by the cation Mg^{2+} or Na^+ and the anion SO_4^{2-} or Cl^- . SC and ionic
370 composition and concentration were strongly associated with factors including lake elevation,
371 surface connectivity, soil composition, hydrology, climate, and watershed land use that explained
372 the spatial patterns.

373 *Spatial variation in SC and major ions in US lakes*

374 Low SC lakes (SC lower than 10% of all lakes) were mostly found in the Northeast, Mid-
375 Atlantic, Great Lakes, and Northwest regions, which aligns with findings from previous studies
376 on US streams (Griffith 2014; Olson & Cormier 2019). High SC lakes (SC higher than 90% of
377 all lakes) were mainly in the Plains, Desert Southwest, and Southeast regions, particularly
378 Florida. Previous macroscale studies of stream conductivity found relatively high SC in the
379 Plains and Desert Southwest but did not find high SC in the Southeast (Griffith 2014; Olson &
380 Cormier 2019). This difference could demonstrate the water chemistry differences between lakes
381 and streams in the Southeast and/or be attributed to the large difference in the number of systems
382 studied between previous and our research (e.g., 315 Coastal Plain (a National Rivers and Stream
383 Assessment region) stream sites in Olson & Cormier 2019 compared to 1,521 lakes in the
384 Southeast in our dataset). Previous regional studies have suggested that the high SC in the
385 Southeast, specifically in Florida, could be attributed to ion inputs from precipitation and
386 seawater intrusion as they were in or near coastal areas (Kiflai et al. 2022; Haque 2023) as well
387 as from discharges from stormwater pipes and detention ponds (Beckingham et al. 2019).
388 Understanding that high-SC lakes occur in the southeastern US has the potential to change SC
389 prediction and management in this region as further increases could lead to biological responses
390 depending on the dominant cation and anions. Furthermore, differences in regional patterns
391 between our study and the stream study could also be due to temporal variations in SC, as some
392 lakes in the Southeast had relatively high CVs (Figure S4). However, most lakes were only
393 sampled once each year in the summer, which highlights the need for sampling during other
394 seasons in order to obtain a comprehensive understanding of long-term inter- and intra- annual
395 variation in lake SC.

396 Although SC is an important measure of water quality, major ion concentration and
397 composition are also critical for understanding lake SC. We observed a linear relationship

398 between SC and the total ion concentration (calculated by summing up the concentration of each
399 ion), which could be applied to predict total ion concentration using SC or vice versa. However,
400 these relationships differed among different major ion compositions, particularly in high-SC
401 lakes (Figure S7). Moreover, the same SC levels can be made up of very different major ion
402 compositions and concentrations. For example, clusters 14 and 1 had similar average SC
403 ($103\mu\text{S}/\text{cm}$ and $111\mu\text{S}/\text{cm}$, respectively), but cluster 14 was dominated by Ca^{2+} and HCO_3^- and
404 cluster 1 was dominated by Na^+ and Cl^- . These differences are important because ionic
405 composition can determine the aquatic species present in water, the impacts of SC and ion level
406 changes on organisms (Elphick et al. 2011; Mount et al. 2016; Huber et al. 2024), the usage of
407 water (Tiri et al. 2018; Zaman et al. 2018), and management strategies to mitigate the negative
408 effects of elevated ions in water.

409 Management agencies in the US responsible for monitoring lake water quality (which is
410 our primary data source) are more likely to measure SC than major ions, perhaps due to the high
411 costs associated with sampling and measuring ions compared to SC. Our results highlight the
412 importance of these agencies considering approaches to supplement their SC data, particularly
413 for lakes with high SC and those in the Southeast (e.g., Florida) and Central Plains. Researchers
414 can also help fill this data gap by developing predictive models for ion composition in lakes that
415 can enhance our understanding of major ion levels and their potential impacts on lake
416 ecosystems. Our estimation of the relationships between major ions and SC and our models,
417 which will be discussed later, that documented relationships between major ion composition and
418 key natural and human features provide an important first step in this direction.

419 We found that ion concentration and composition varied greatly within and among
420 regions. As expected (Dugan 2024), Ca^{2+} , and sometimes Na^+ , was the dominant cation in many
421 lakes. However, in about 16% of lakes located mostly in the Northern Plains, Mg^{2+} was the

422 dominant cation or equally abundant to Ca^{2+} (a few others were located in the Prairie Peninsula
423 and West regions). We also found substantial within-region variation in ion levels, which is
424 intriguing because earlier stream studies suggested that the composition of ions was directly
425 controlled by natural mechanisms (e.g., precipitation, rock dominance, and evaporation), which
426 often vary regionally (Griffith 2014; Olson & Cormier 2019). However, previous regional lake
427 studies also found heterogeneity in ion composition that they attributed to both natural processes
428 and human activities (e.g., groundwater, soil texture, and acid deposition) (Baker et al. 1991;
429 DeSellas et al. 2023). Our research findings that each NEON region included multiple ion
430 clusters further support the idea that major ions are influenced by multiple multi-scaled natural
431 and human factors that vary both locally and regionally.

432 *Multi-scaled natural and human factors related to SC and major ions*

433 SC, major ion concentrations, and ion composition were related to both natural and
434 human factors. This result suggests that factors such as geology, hydrology, and climate are key
435 determinants of ion inputs from natural sources and that, in combination with the water balance
436 in the lake, control background SC and ions. Human disturbances that influence export ions to
437 lakes cause ions to deviate from those background levels and vary across lakes within the region.
438 For example, Cl^- was strongly and mostly influenced by human activities (Figure S8). Therefore,
439 effective management strategies designed to manage water quality and lake ecosystems should
440 target specific measures of SC and specific ions.

441 We found that a range of multi-scaled natural factors were important for understanding
442 SC and major ions. For example, at the lake scale, elevation was negatively related to SC.
443 However, because these low-elevation and high SC lakes also had higher percentages of
444 developed land use and road density in their catchments, this relationship could be signaling
445 greater human ion inputs to low-elevation lakes compared to those at higher elevations (Table

446 S5). A similar pattern of higher SC in lower-elevation lakes was observed by Müller et al. (1998),
447 which was attributed to higher rock weathering and dissolution rates and lower relative water
448 inputs from rainfall with decreasing elevations. Additionally, our lakes with low elevation and
449 high SC were mostly in Florida and the Atlantic Neotropical region and had relatively high
450 proportions of Na^+ and Cl^- , which could also be related to current or past seawater intrusion
451 (Kiflai et al. 2022).

452 We anticipated surface connectivity, such as measures of landscape position, to be
453 important for explaining differences in SC and major ions because it influences the relative
454 importance of water and ion inputs from precipitation, surface water, and groundwater sources
455 (Riera et al. 2000; Bennett et al. 2007; Dumelle et al. 2024). However, it was only strongly
456 associated with lake membership in some ion clusters. For instance, cluster 8 (high SC,
457 dominated by Mg^{2+} and SO_4^{2-}) had the highest proportions of terminal (only surface inflows) and
458 isolated (no surface inflow or outflow) lakes among all clusters; and 41% of all terminal lakes
459 and 20% of all isolated lakes were in cluster 8 (Table S4). Previous research found that isolated
460 and terminal lakes can be ion sinks for the catchment, particularly in evaporative regions
461 (Saleem et al. 2015; Cotner et al. 2022; Ding et al. 2024). We also found that the SC of drainage
462 lakes (those with both inflows and outflows) was higher than that of other connectivity classes in
463 the most common ion group (i.e., clusters dominated by Ca^{2+} and HCO_3^-), but SC for drainage
464 lakes was lower than that of other connectivity classes in other cluster groups (Figure 5). This
465 result implies that although the water inputs from tributaries can import ions into lake (e.g., from
466 rock weathering) and increase ionic concentrations, other stronger natural and human processes
467 (e.g., groundwater discharge and human land use) may mask this effect by causing greater
468 increases, in which case tributary water inputs could have a dilution effect on SC and ion
469 concentrations.

470 At the watershed and regional scales, soil composition variables, measures of overland
471 surface water flow and groundwater recharge, and precipitation were important in models of lake
472 SC and ion composition. We found that lakes with high percentages of clay tend to have higher
473 SC and/or Mg^{2+} concentrations than other lakes. This result could be explained by the higher
474 levels of agricultural activities in areas with higher clay in the soil along with the finding that soil
475 with a higher proportion of clay often contains higher amounts of cations such as Mg^{2+} (Ross et
476 al. 2008). We found negative relationships between SC and groundwater recharge and runoff,
477 which may indicate greater influence of precipitation and surface water inputs for low SC lakes
478 (Webster et al. 2006) compared to the influences of groundwater as a source of ions to high SC
479 lakes in many regions (Li et al. 2020; Dugan 2024). Additionally, runoff and regional
480 precipitation were both lower in lakes with high proportions of Mg^{2+} (clusters 12, 5, and 8),
481 suggesting that other factors such as evaporation may be more important for these lakes (La
482 Baugh et al. 2000). These findings imply the significant, complex roles that watershed geology
483 and hydrology and regional climate play in affecting lake hydrologic and chemical budgets.

484 Lake SC and major ions were also related to multi-scaled human factors (and those that
485 indicate the lack of disturbance), including watershed land use/cover and road density. For
486 example, forest land cover, which can be used to identify lakes with fewer human disturbances,
487 was negatively related to SC and was lower in clusters 12, 5, 8, 1, and 6, which were not
488 dominated by Ca^{2+} and HCO_3^- . Moreover, one of the most crucial factors that affected ion
489 composition was road density, which was higher in clusters 7, 10, 1, and 6. These same clusters,
490 which are located in regions where winter road salt applications happen (e.g., the Northeast,
491 Mid-Atlantic, Great Lakes, and Prairie Peninsula regions; Dugan et al. 2017), also had higher
492 concentrations and proportions of Na^+ and Cl^- than other clusters. Therefore, this pattern is likely
493 due to road deicing salt applications in winter that are retained in lakes, causing increases in salt

494 concentrations throughout the year (Kaushal et al. 2021; Solomon et al. 2023). Supplementing a
495 long-term study that found an increasing trend of Cl^- in 125 of 371 lakes in North America
496 (Dugan et al. 2017), our results demonstrate the importance of reducing road salt applications in
497 regions with high impervious surface area and road density near lakes.

498 Although our models considered the effects of all the natural and human factors, we did
499 not examine the joint effects of these factors or the influence of temporal changes in those factors
500 on SC and ions. Very few studies have looked at the combined effects of such factors on other
501 water quality parameters (Kernan & Helliwell 2001; Nobre et al. 2020; Lin et al. 2021), and no
502 or limited information exists on SC and major ions, despite the likely complex interactive effects.
503 Additionally, human activities can influence some of the natural factors and cause them to
504 change through time (Meybeck 2003), which may affect these factors' relationships with ions.
505 Therefore, future research efforts should attempt to disentangle the complex underlying
506 mechanisms affecting SC and ion concentrations and composition in lakes.

507

508 **Conclusion**

509 We documented spatial variation in specific conductance (SC) and major ion
510 concentration and composition in 9,784 and 1,218 lakes, respectively, across the continental US.
511 We found substantial spatial variation in ion composition and that lakes with similar SC values
512 can have very different ion composition, which can result in differential effects on lake biota,
513 may change differently in the future, and likely require different strategies to manage. These
514 findings highlight the importance of considering all major ions when studying lake water
515 chemistry. We also found that variation in SC was related to a wide range of local, watershed,
516 and regional factors, such as lake elevation, soil texture, hydrology, precipitation, and forest land
517 cover. Major ions were strongly associated with both human and natural factors. In particular,

518 elevated Cl⁻ levels were predominantly related to road density and urban development, which can
519 be managed to address issues such as freshwater salinization. Our results suggest that while
520 geology, hydrology, surface connectivity, and climate control ion inputs from natural sources,
521 human disturbances directly and indirectly alter ion export to lakes, causing SC and major ions to
522 deviate from their background levels.

523

524

525 **Conflict of interest:** The authors declare no conflict of interest.

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530

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Supplementary Figures

Paper title: Natural and human drivers of specific conductance and major ion composition in United States lakes

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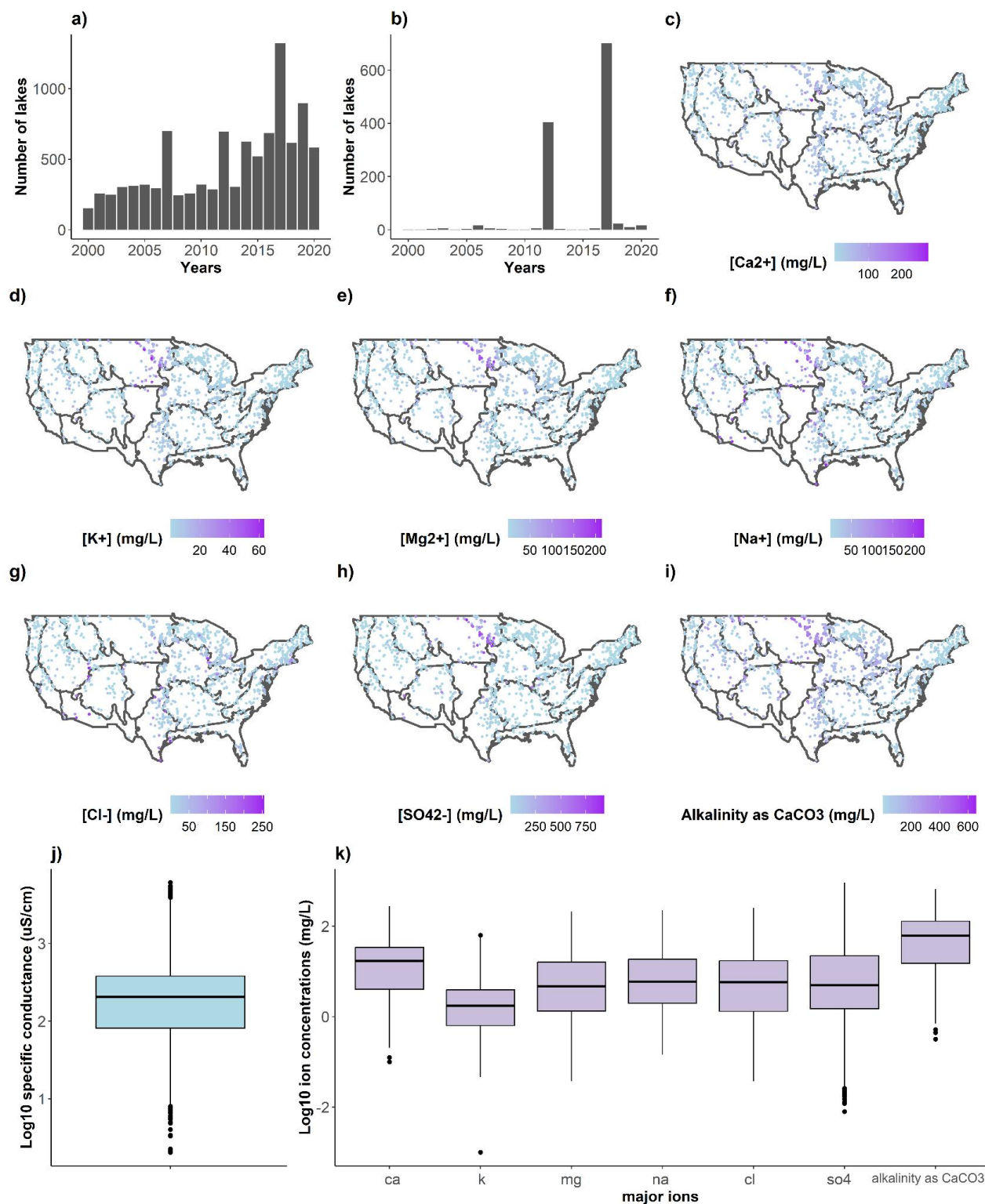


Figure S1. Summary of specific conductance (SC) and major ion data: a) number of lakes sampled in each latest SC sampling year; b) number of lakes sampled in each latest major ion sampling year; c) - i) maps showing major ion concentrations and alkalinity levels as CaCO₃ (mg/L) by NEON region; j) boxplot showing the distribution of log₁₀-transformed SC data; and k) boxplots showing the distribution of log₁₀-transformed major ion concentrations. See Figure 1 caption for NEON region names.

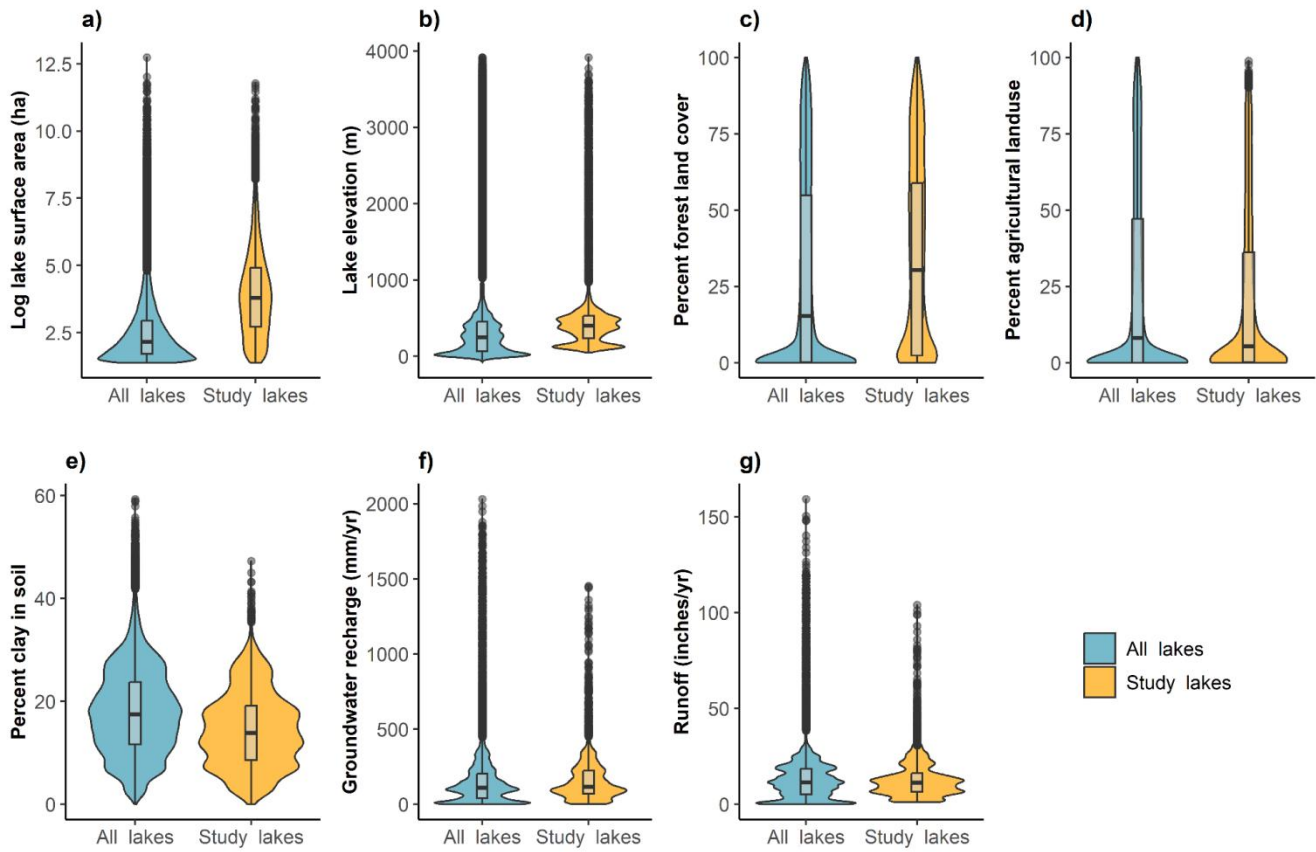


Figure S2. Comparison of all lakes > 4 ha in the U.S. (All lakes, blue) relative to the study lake population (Study lakes, orange). The violin plots show all data for the two groups of lakes, and the boxes within the violin plots are the interquartile ranges with the median. The ‘All lakes’ category was derived from the LAGOS lake population > 4 ha.

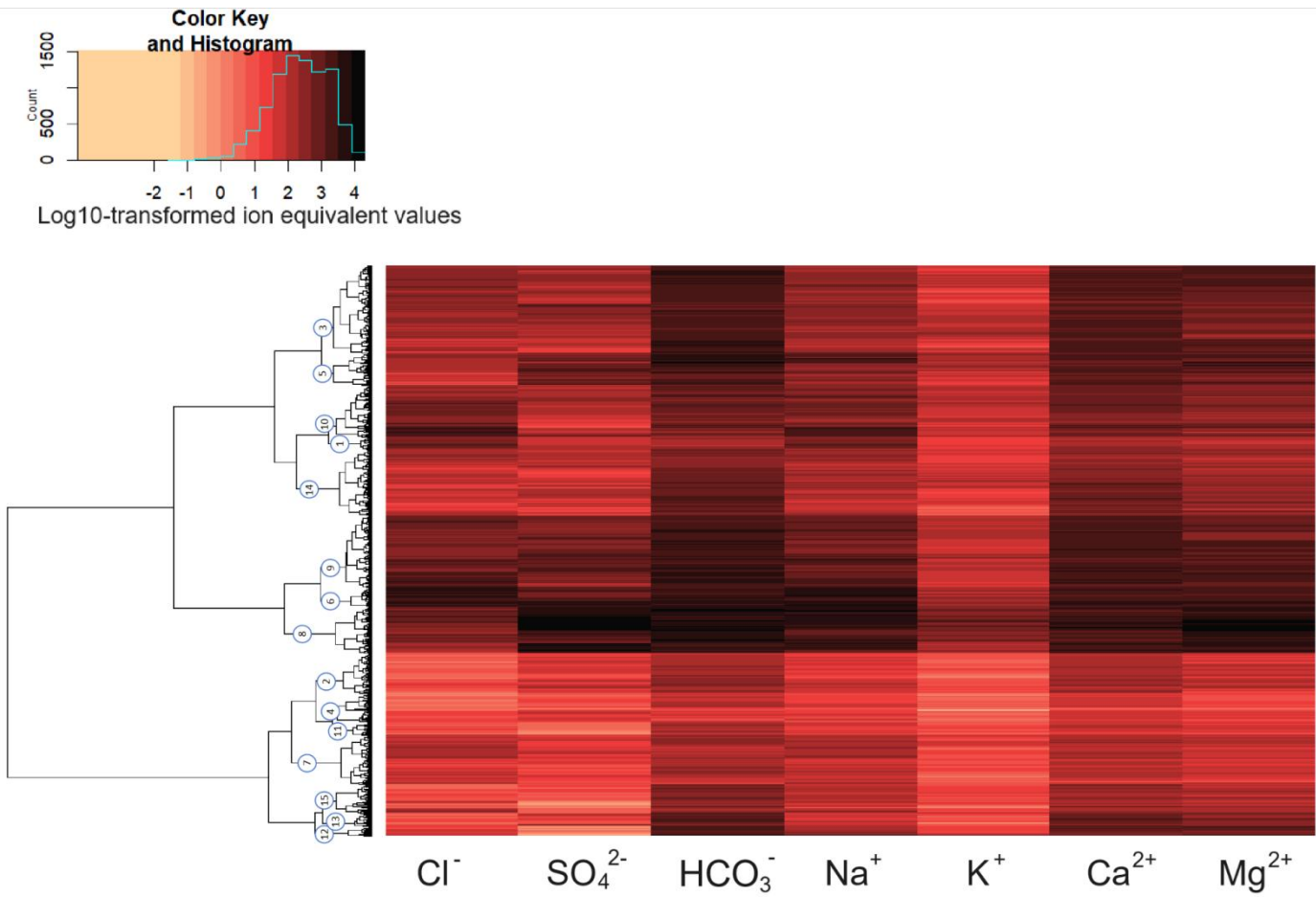


Figure S3. Heatmap showing log₁₀-transformed major ion equivalent values grouped by clusters with a dendrogram (output of the hierarchical clustering using the Ward's method) on the left edge. The histogram at the top shows the color key and distribution of data.

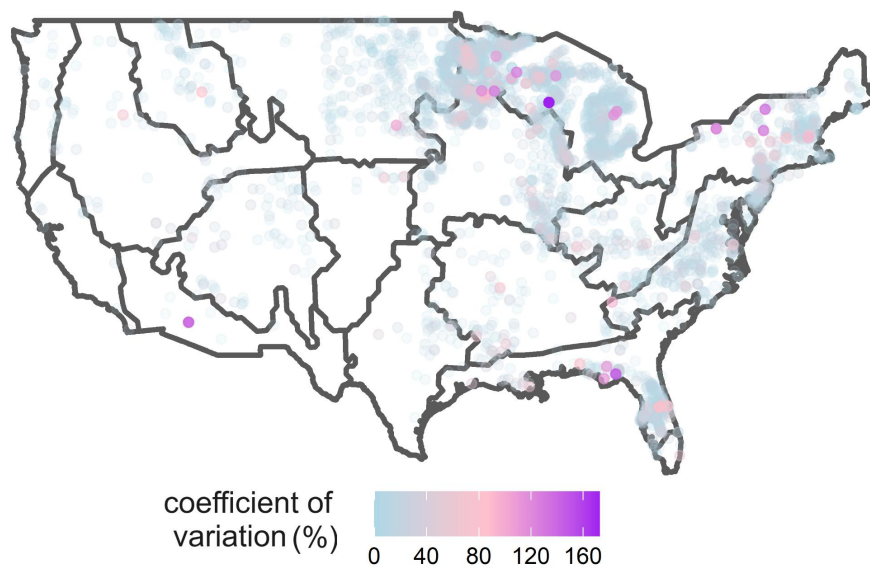


Figure S4. Map showing the intra-year SC coefficient of variation ($N=4,581$) of each lake by NEON region. Each dot represents a lake. See Figure 1 caption for NEON region names.

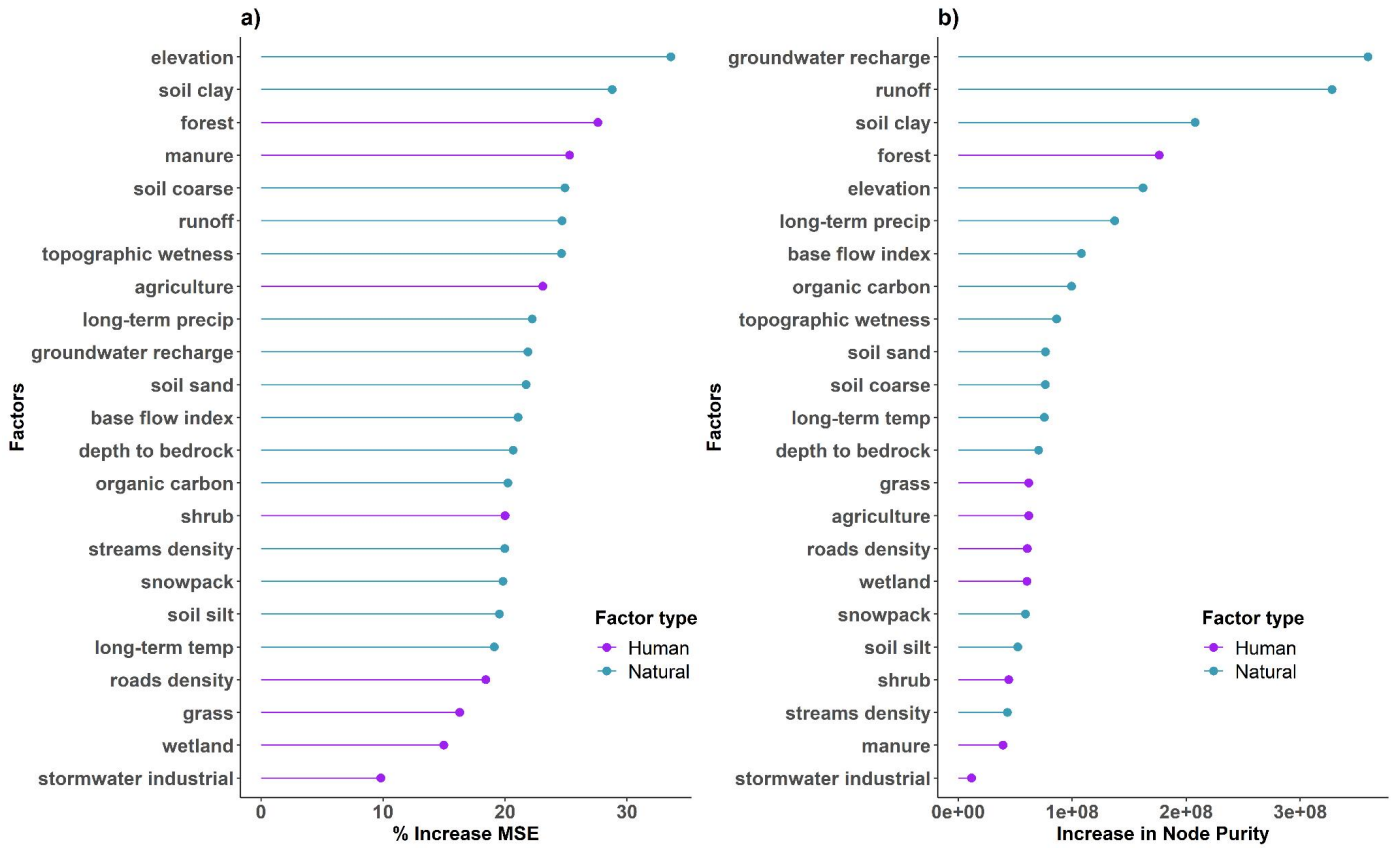


Figure S5. Results from random forest regression model (N=9,784) of natural and human factors and their relationships to lake SC: a) importance scores of important natural and human factors based on the percentage increase in mean squared error (MSE); b) importance scores of important natural and human factors based on the increase in node purity (from Gini Index).

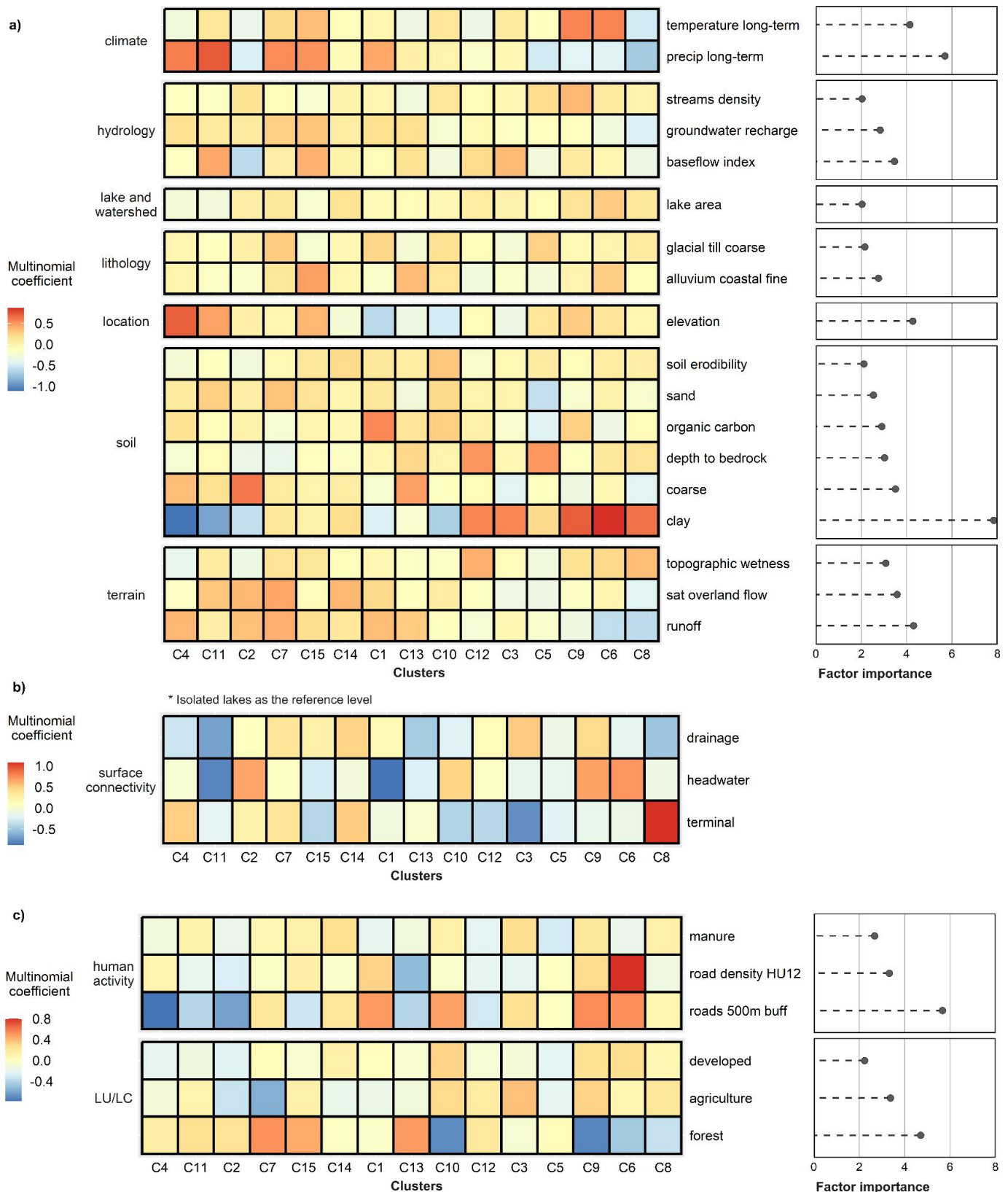


Figure S6. Heatmaps showing multinomial coefficients generated from a single GLMNET model predicting ion composition cluster assignment based on each natural (a,b) and human factor (c), with the factor importance plotted on the right (lollipop plots) that was calculated as the absolute sum of the cluster-specific coefficients. Numeric factors in plots a) and c) were centered and scaled. The text on the left of the heatmaps indicates factor

categories, and the text on the right indicates factor names. Surface connectivity (plot b) was input as a categorical natural variable in the model; we plotted this factor separately to avoid confusion and did not calculate its importance score. Only factors with importance values ≥ 2 were plotted (18 of 30 inputted natural factors, surface connectivity was counted as one factor, and 6 of 12 inputted human factors). A full table of coefficients can be found in Table S3.

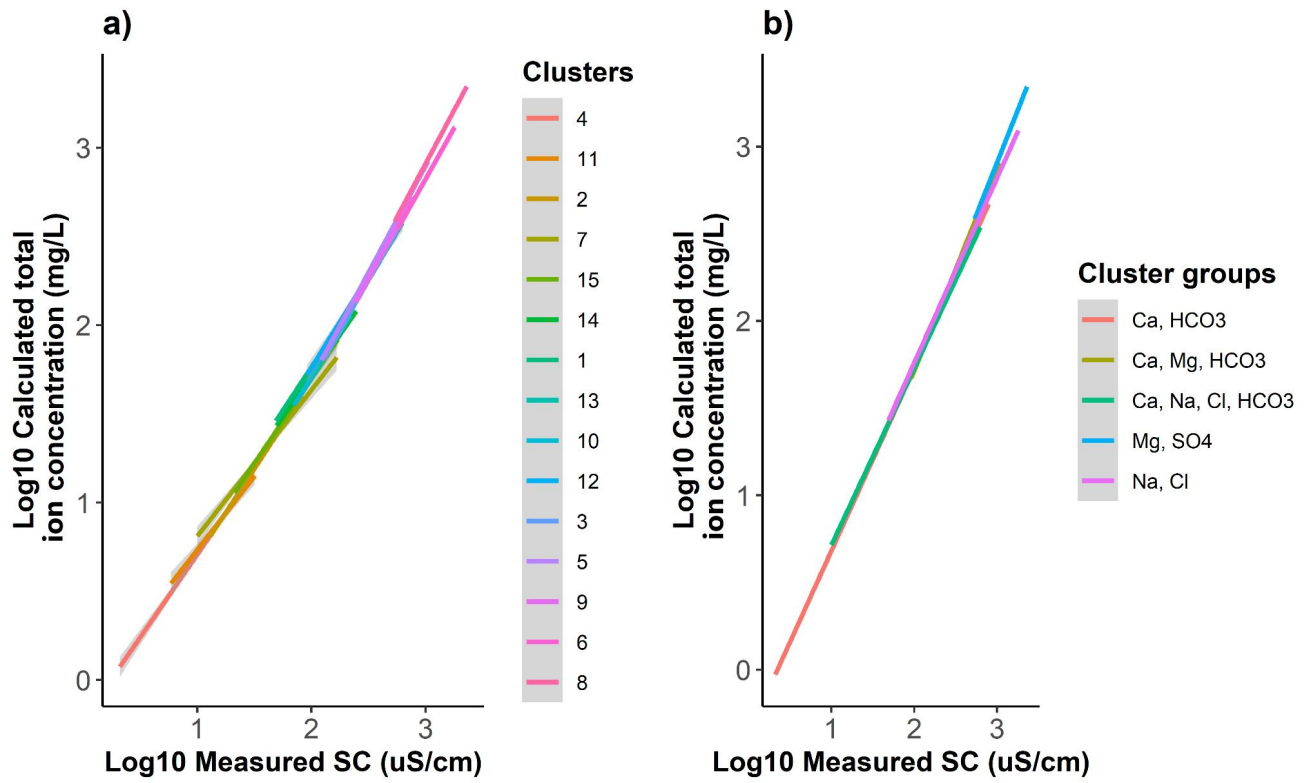
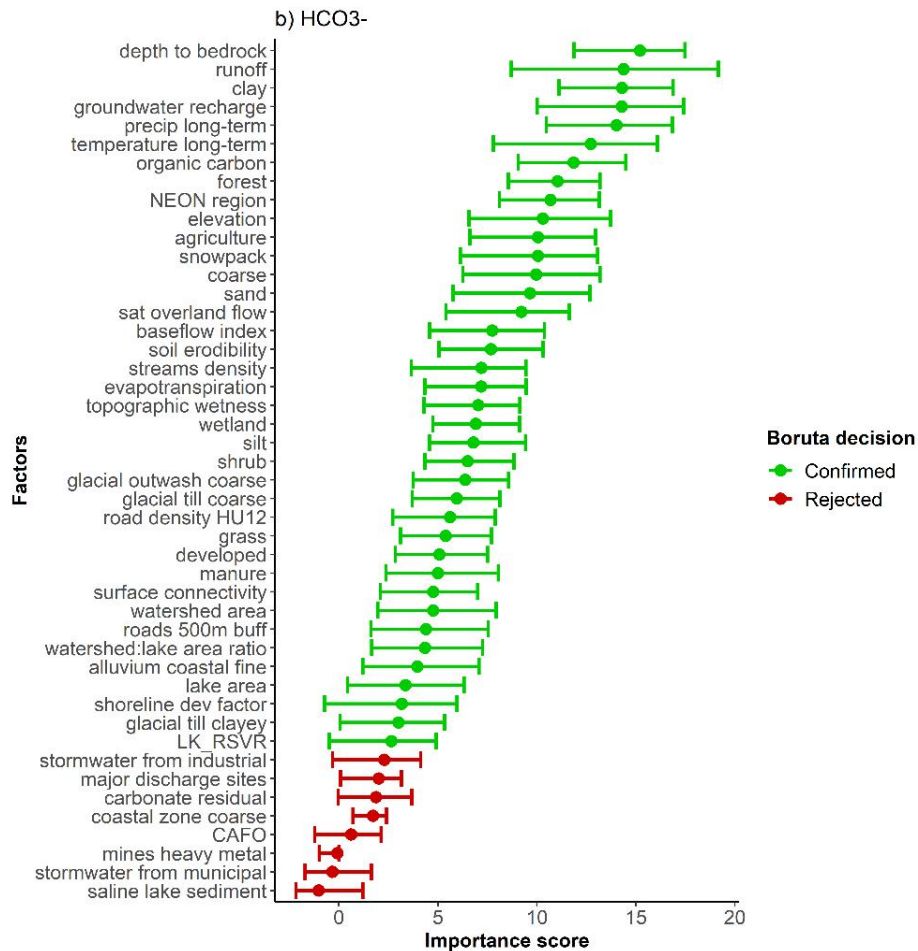
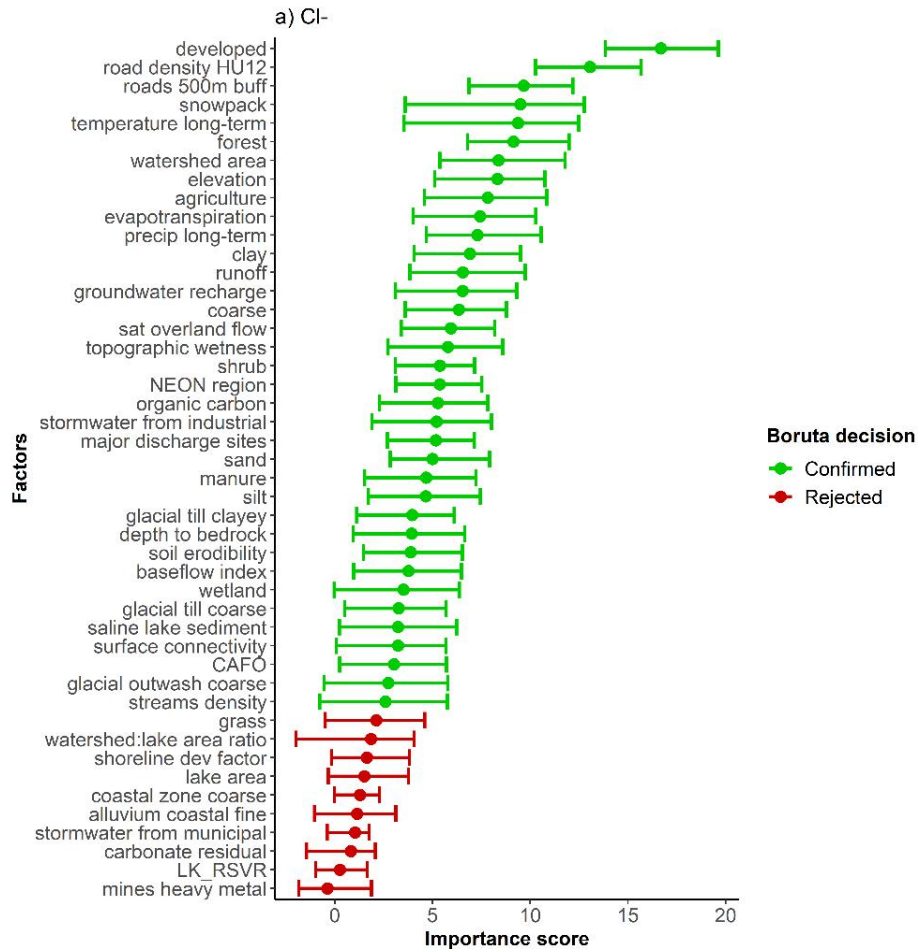


Figure S7. Relationships between log₁₀-transformed measured SC and calculated total ion concentration by clusters (a; 15 clusters; ordered based on SC, from the lowest (cluster 4) to the highest (cluster 8)) and cluster groups (b; 5 groups based on dominant cations and anions). The prediction lines were fitted with linear models.



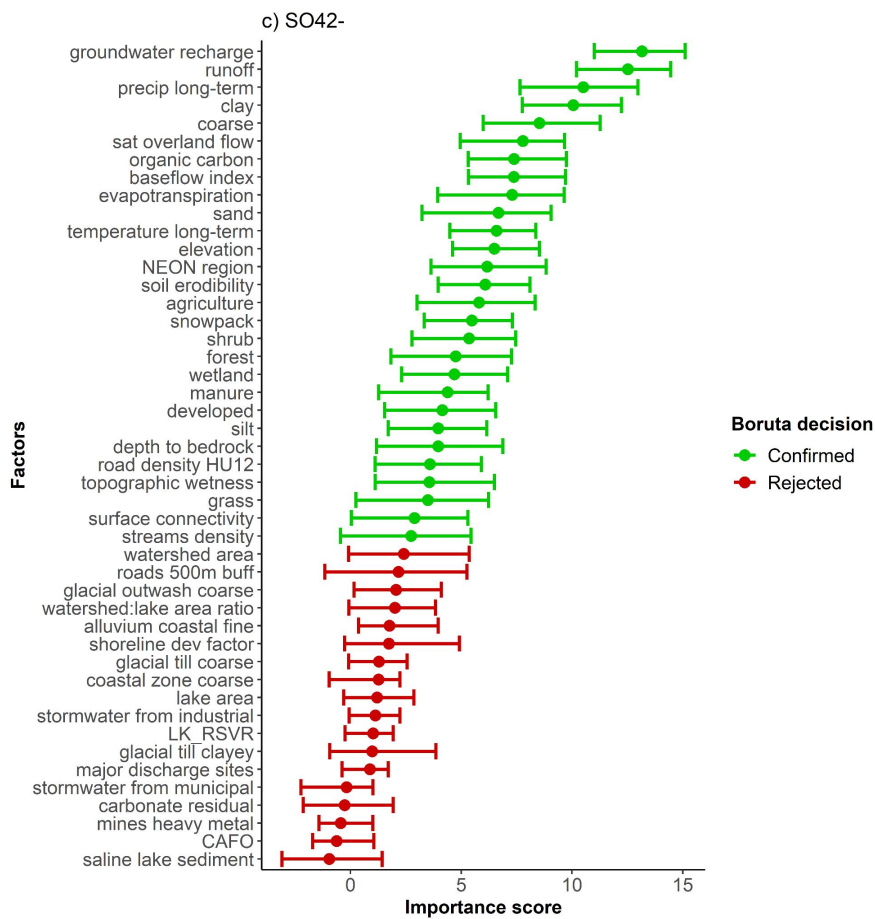


Figure S8. Importance scores from Boruta feature selection for the three anions: Cl⁻ (a), HCO₃⁻ (b), and SO₄²⁻ (c). Green dots and bars indicate that the factors were identified as ‘important’ by Boruta and red dots and bars indicate that Boruta rejected the factors. Dots are the mean Boruta importance values and bars are the minimum and maximum Boruta importance values. Boruta feature selection identifies relevant predictors by comparing the importance of the test feature with the importance of permuted copies of the data known as shadow features (Kursa et al. 2010).

Kursa, M. B., A. Jankowski, and W. R. Rudnicki. 2010. Boruta—A system for feature selection. *Fundam. Inform.* 101: 271–285. doi:10.3233/FI-2010-288

Supplementary Tables

Paper title: Natural and human drivers of specific conductance and major ion composition in United States lakes

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Table S1. Table of natural and human factors with variable sources, spatial scale, group, explanation, and data transformation information.

Factors name in figure	Factors original name	Spatial scale of measurement	Group	Natural Human	Data sources	Explanation	Transformation in RF
connectivity	lake_connectivity_class	lake	surface connectivity	N	LAGOS-US LOCUS	Hydrologic connectivity class of the focal lake	NA
elevation	lake_elevation_m	lake	location	N	LAGOS-US LOCUS	Mean elevation in the zone	natural log
reservoir	lake_rsvr_class	lake	lake and watershed	N	LAGOS-US RESERVOIR	The classification of a lake into either natural lake or reservoir	NA
shoreline dev factor	lake_shorelinedevfactor	lake	lake and watershed	N	LAGOS-US LOCUS	A measure of the deviation of lake shape from 1, which is a perfect circle, calculated from the lake water area and perimeter	natural log
lake area	lake_waterarea_ha	lake	lake and watershed	N	LAGOS-US LOCUS	The water area of the lake contained within the outer shoreline	natural log
manure	ManureMean	HU12	human activity	H	EnviroAtlas	Manure application (kg N/ha/yr)	natural log
evapotranspiration	NA, calculated	HU12	climate	N	Blodgett 2023	Long-term average evapotranspiration data by HU12 region	NA
snowpack	NA, calculated	HU12	climate	N	Blodgett 2024	Long-term average snowpack water equivalent storage data by HU12 region	NA
NEON region	neon_zoneid	NEON region	location	N	LAGOS-US LOCUS	National Ecological Observatory Network (NEON) region id	NA
precip long-term	NA, calculated	HU12	climate	N	LAGOS-US GEO	Long-term average precipitation data (5 years before sampling year)	NA
roads 500m buff	road.500buff	500m buffer	human activity	H	LAGOS-US GEO	Road density in 500 m buffer around the shoreline of a LAGOS-US lake polygon	natural log
temperature long-	NA, calculated	HU12	climate	N	LAGOS-US	Long-term average air temperature	NA

term					GEO	data (5 years before sampling year)
baseflow index	value.baseflowindex_pct	HU12	hydrology	N	LAGOS-US GEO	Mean within the zone of the percentage of streamflow that can be attributed to groundwater discharge into streams, calculated as baseflow divided by total flow using data from 1951-1980 logit, generalized (-0.01 to 1.01)
groundwater recharge	value.groundwaterrecharge_mmper_yr	HU12	hydrology	N	LAGOS-US GEO	Mean within the zone of annual groundwater recharge calculated as baseflow multiplied by mean annual runoff using data from 1951-1980 natural log
alluvium coastal fine	value.lith_alluviumcoastalfine_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as alluvium and fine textured coastal zone sediment logit, generalized (-0.01 to 1.01)
carbonate residual	value.lith_carbonateresid_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as carbonate residual material logit, generalized (-0.01 to 1.01)
coastal zone coarse	value.lith_coastalzonecoarse_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as coastal zone sediment, coarse textured logit, generalized (-0.01 to 1.01)
glacial outwash coarse	value.lith_glacialoutwashcoarse_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as glacial outwash and glacial lake sediment, coarse textured logit, generalized (-0.01 to 1.01)
glacial till clayey	value.lith_glacialtillclayey_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as glacial till, clayey logit, generalized (-0.01 to 1.01)
glacial till coarse	value.lith_glacialtillcoarse_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as glacial till, coarse textured logit, generalized (-0.01 to 1.01)
saline lake sediment	value.lith_salinelake_pct	HU12	lithology	N	LAGOS-US GEO	Percent of zone with lithology classified as saline lake sediment logit, generalized (-0.01 to 1.01)
mines heavy metal	value.mines_heavy_metal_n	HU12	human activity	H	LAGOS-US GEO	Count of heavy metal mines active in 2003 within the zone natural log
CAFO	value.npdes_cafon_n	HU12	human activity	H	LAGOS-US GEO	Count of concentrated animal feed operations within the zone natural log
major discharge sites	value.npdes_major_discharge_n	HU12	human activity	H	LAGOS-US GEO	Count of facilities with major discharges within the zone natural log
stormwater from industrial	value.npdes_stormwaterindustrial_n	HU12	human activity	H	LAGOS-US GEO	Count of industrial storm water sewers within the zone natural log

stormwater from municipal	value.npdes_stormwatermunicipal_n	HU12	human activity	H	LAGOS-US GEO	Count of Phase I MS4: Municipal separate storm sewer systems within the zone	natural log
road density HU12	value.roads_mperha	HU12	human activity	H	LAGOS-US GEO	Road density within each HU12 region	natural log
runoff	value.runoff_inperyr	HU12	terrain	N	LAGOS-US GEO	Mean within the zone of annual runoff; data from 1951 to 1980	natural log
sat overland flow	value.satoverland_flow_pct	HU12	terrain	N	LAGOS-US GEO	Mean within the zone of the average percentage of saturation overland flow in total streamflow using data from 1951-1980	logit, generalized (-0.01 to 1.01)
clay	value.soil_clay_pct	HU12	soil	N	LAGOS-US GEO	Average percentage mass fraction of clay, 0 to 2 micrometers, in the 0 to 5 cm depth soil layer within the zone	logit, generalized (-0.01 to 1.01)
coarse	value.soil_coarse_pct	HU12	soil	N	LAGOS-US GEO	Average percentage by volume of coarse fragments in the 0 to 5 cm soil depth layer within the zone	logit, generalized (-0.01 to 1.01)
depth to bedrock	value.soil_depth_tobedrock_cm	HU12	soil	N	LAGOS-US GEO	Average absolute depth to bedrock within the zone	natural log
k factor	value.soil_kffact	HU12	soil	N	LAGOS-US GEO	Average soil erodibility factor, not adjusted for the effect of rock fragments, within the zone	natural log
organic carbon	value.soil_organicarbon_gperkg	HU12	soil	N	LAGOS-US GEO	Average organic carbon content, fine earth fraction, in the 0 to 5 cm soil layer within the zone	natural log
sand	value.soil_sand_pct	HU12	soil	N	LAGOS-US GEO	Average percentage mass fraction of sand, 50 to 200 micrometers, in the 0 to 5 cm depth soil layer within the zone	logit, generalized (-0.01 to 1.01)
silt	value.soil_silt_pct	HU12	soil	N	LAGOS-US GEO	Average percentage mass fraction of silt, 2 to 50 micrometers, in the 0 to 5 cm depth soil layer within the zone	logit, generalized (-0.01 to 1.01)
streams density	value.streams_all_mperha	HU12	hydrology	N	LAGOS-US GEO	Density of all streams within the zone, calculated as the sum of the stream length divided by the zone area	natural log
topographic wetness	value.topographic_wetness	HU12	terrain	N	LAGOS-US GEO	Mean topographic wetness index of cells within the zone	natural log
watershed area	ws_area_ha	watershed	lake and	N	LAGOS-US	Area of zone polygon	natural log

			watershed	LOCUS			
watershed:lake area ratio	ws_lake_arearatio	lake	lake and watershed	N	LAGOS-US LOCUS	Ratio between watershed area and lake water area	natural log
agriculture	NA, combined nlcd_cultcrop82_pct with nlcd_past81_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as cultivated crops or pasture and hay using 2006 data	logit, generalized (-0.01 to 1.01)
grass	nlcd_grass71_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as grassland or herbaceous using 2006 data	logit, generalized (-0.01 to 1.01)
shrub	nlcd_shrub52_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as shrub and scrub using 2006 data	logit, generalized (-0.01 to 1.01)
developed	NA, combined nlcd_devopen21_pct, nlcd_devlow22_pct, nlcd_devmed23_pct, and nlcd_devhi24_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as developed, including open space and low, medium, and high intensity, using 2006 data	logit, generalized (-0.01 to 1.01)
forest	NA, combined nlcd_forcon42_pct, nlcd_fordec41_pct, and nlcd_formix43_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as evergreen forest, deciduous forest, or mixed forest using 2006 data	logit, generalized (-0.01 to 1.01)
wetland	NA, combined nlcd_wetemerg95_pct with nlcd_wetwood90_pct	watershed	LU/LC	H	LAGOS-US GEO	Percent of zone classified as emergent herbaceous or woody wetlands using 2006 data	logit, generalized (-0.01 to 1.01)

Table S2. Converting factors applied to convert major ion concentrations from mg/L to ueq/L and to calculate SC. Converting factors were extracted from Hem (1985).

Ions	Converting factor (from mg/L to ueq/L)	Converting factor for calculating SC using ion equivalent concentrations
Ca	49.9	52.0
Mg	82.29	46.6
Na	43.5	48.9
K	25.58	72.0
Cl	28.21	75.9
SO ₄	20.82	73.9
Alkalinity (as CaCO ₃)	19.98	(as HCO ₃ ⁻) 43

Hem, J.D. 1985. Study and Interpretation of the Chemical Characteristics of Natural Water. U.S. Geological Survey, Water Supply Paper 2254. Retrieved from <https://pubs.usgs.gov/wsp/wsp2254/html/pdf.html>

Table S3. Multinomial coefficients from a GLMNET model predicting ion composition cluster assignment. C1-C15 indicate ion clusters. Groups indicate natural and human factor categories.

Factors	Group	N_H	C1	C10	C11	C12	C13	C14	C15	C2	C3	C4	C5	C6	C7	C8	C9
drainage	surface connectivity	N	0.16	-0.21	-0.67	0.14	-0.47	0.47	0.25	0.10	0.52	-0.31	-0.11	-0.17	0.37	-0.51	0.45
headwater	surface connectivity	N	-0.88	0.47	-0.81	0.07	-0.24	-0.05	-0.27	0.66	-0.16	-0.01	-0.15	0.72	0.08	-0.11	0.68
terminal	surface connectivity	N	-0.07	-0.42	-0.20	-0.40	0.03	0.51	-0.40	0.23	-0.76	0.50	-0.22	-0.12	0.37	1.08	-0.15
elevation	location	N	-0.61	-0.49	0.45	-0.08	-0.32	-0.25	0.37	0.04	-0.33	0.70	0.14	0.15	-0.03	-0.03	0.29
shoreline dev factor	lake and watershed	N	0.33	0.12	-0.10	-0.25	-0.04	-0.07	-0.33	-0.01	0.05	0.02	-0.05	0.07	-0.02	0.15	0.14
lake area	lake and watershed	N	-0.07	-0.04	-0.27	0.07	-0.08	0.17	-0.21	0.06	0.01	-0.26	-0.08	0.30	0.12	0.13	0.17
manure	human activity	H	-0.20	0.13	0.12	-0.23	-0.13	0.27	0.13	-0.18	0.27	-0.12	-0.29	-0.18	0.07	0.13	0.21
evapotranspirat ion	climate	N	0.05	0.07	0.06	-0.05	0.15	-0.05	-0.21	0.11	0.11	0.03	0.00	-0.11	0.19	-0.10	-0.27

snowpack	climate	N	-0.06	-0.06	-0.05	-0.05	0.18	-0.13	0.15	0.05	-0.10	0.06	0.09	-0.01	0.05	-0.05	-0.08
NEON region	location	N	0.00	-0.03	0.10	0.07	0.01	0.02	0.09	-0.01	0.01	-0.07	-0.05	-0.02	-0.04	-0.04	-0.05
precip long-term	climate	N	0.43	0.00	0.73	-0.18	0.04	-0.07	0.51	-0.47	-0.05	0.60	-0.51	-0.44	0.54	-0.70	-0.43
road density 500m buff	human activity	H	0.50	0.48	-0.40	-0.30	-0.40	0.22	-0.32	-0.62	0.26	-0.78	-0.01	0.53	0.21	0.09	0.55
temperature long-term	climate	N	-0.02	0.14	0.13	-0.34	-0.32	-0.10	0.34	-0.36	0.10	-0.29	-0.12	0.57	0.23	-0.51	0.57
baseflow index	hydrology	N	-0.08	-0.27	0.42	0.22	0.17	0.02	0.39	-0.60	0.36	-0.14	-0.28	-0.03	0.04	-0.34	0.10
groundwater recharge	hydrology	N	0.19	-0.23	0.10	-0.06	0.19	0.07	0.32	0.09	-0.14	0.19	-0.11	-0.30	0.27	-0.46	-0.12
alluvium coastal fine	lithology	N	-0.18	0.15	-0.16	-0.29	0.35	-0.03	0.46	-0.18	-0.17	0.00	-0.26	0.28	0.14	-0.10	-0.01
carbonate residual	lithology	N	-0.03	-0.02	-0.04	0.02	0.15	0.10	0.07	0.03	0.08	-0.03	-0.12	-0.12	-0.06	-0.05	0.03
coastal zone coarse	lithology	N	0.24	0.24	-0.04	-0.01	0.00	-0.15	-0.10	-0.02	-0.07	0.00	0.00	0.00	-0.01	0.00	-0.07

glacial outwash coarse	lithology	N	0.15	-0.13	0.01	0.03	0.06	-0.11	-0.04	0.04	0.09	-0.03	-0.11	-0.06	-0.08	-0.11	0.28
glacial till clayey	lithology	N	0.05	0.01	-0.06	-0.10	0.18	-0.12	0.02	-0.06	-0.01	0.00	-0.14	0.36	-0.17	-0.15	0.19
glacial till coarse	lithology	N	0.24	0.18	-0.12	-0.10	-0.24	-0.05	-0.22	0.04	-0.25	-0.03	0.26	-0.01	0.29	0.08	-0.06
saline lake sediment	lithology	N	0.00	0.04	0.12	-0.04	-0.02	-0.06	-0.06	-0.06	0.06	-0.03	-0.09	0.20	-0.01	0.00	-0.05
major discharge sites	human activity	H	0.01	0.00	-0.05	-0.04	-0.05	-0.08	0.01	-0.06	0.04	0.01	-0.15	0.27	-0.03	0.05	0.08
stormwater from industrial	human activity	H	0.14	-0.02	-0.06	-0.04	-0.03	0.20	0.02	-0.06	-0.06	0.12	-0.07	-0.14	-0.11	-0.01	0.11
stormwater from municipal	human activity	H	-0.05	0.01	0.00	0.00	0.00	0.09	-0.03	0.00	0.18	0.00	0.00	-0.10	-0.06	0.08	-0.10
road density HU12	human activity	H	0.32	-0.04	-0.19	-0.21	-0.53	-0.04	0.15	-0.28	-0.21	0.10	0.00	0.81	-0.03	-0.14	0.29
runoff	terrain	N	0.35	-0.14	0.05	-0.24	0.29	0.13	0.20	0.33	-0.05	0.37	-0.19	-0.58	0.40	-0.60	-0.32
sat overland flow	terrain	N	0.23	0.08	0.31	-0.15	-0.13	0.37	-0.09	0.37	-0.32	-0.16	-0.33	-0.02	0.43	-0.43	-0.16

clay	soil	N	-0.47	-0.67	-0.89	0.57	-0.21	0.20	0.01	-0.57	0.56	-1.11	0.23	0.86	0.13	0.64	0.72
coarse	soil	N	-0.22	-0.13	0.19	-0.08	0.46	0.06	-0.01	0.62	-0.42	0.35	-0.12	-0.05	0.08	-0.41	-0.32
depth to bedrock	soil	N	-0.03	0.00	-0.08	0.50	0.23	-0.16	-0.11	-0.35	-0.05	-0.22	0.49	0.12	-0.37	0.17	-0.14
soil erodibility	soil	N	0.12	0.30	-0.13	-0.19	0.07	0.22	0.13	-0.29	-0.06	-0.24	0.09	0.06	-0.06	0.06	-0.09
organic carbon	soil	N	0.55	0.28	-0.10	0.02	0.15	-0.05	-0.02	0.02	-0.20	0.17	-0.44	-0.33	-0.24	-0.08	0.27
sand	soil	N	0.13	0.23	0.27	0.00	-0.28	-0.01	0.16	0.05	0.00	0.09	-0.57	0.01	0.32	-0.20	-0.20
silt	soil	N	-0.11	-0.14	0.00	-0.12	0.18	0.00	-0.08	0.01	0.16	0.01	0.04	-0.01	-0.07	0.12	0.00
streams density	hydrology	N	-0.03	0.14	-0.15	-0.08	-0.30	0.01	-0.20	0.18	-0.05	-0.12	0.22	0.10	-0.09	0.00	0.36
topographic wetness	terrain	N	-0.04	-0.16	0.08	0.40	-0.15	-0.10	0.14	-0.33	-0.08	-0.37	-0.30	0.23	0.17	0.35	0.17
watershed area	lake and watershed	N	-0.03	0.18	-0.13	-0.09	0.04	0.02	-0.11	-0.04	0.01	-0.25	-0.04	0.25	0.00	0.11	0.10
watershed:lake area ratio	lake and watershed	N	0.01	0.17	0.02	-0.10	0.07	-0.06	0.01	-0.06	0.00	-0.09	0.00	0.07	-0.06	0.03	0.00

agriculture	LU/LC	H	-0.17	0.29	0.11	0.17	-0.12	-0.18	0.13	-0.33	0.38	-0.11	-0.21	0.09	-0.57	0.21	0.31
grass	LU/LC	H	0.04	-0.16	-0.06	-0.20	0.00	0.09	-0.07	0.06	-0.07	0.12	0.07	0.11	0.10	0.04	-0.10
shrub	LU/LC	H	-0.03	0.40	0.14	0.02	-0.04	0.13	-0.07	-0.11	-0.23	0.14	-0.01	-0.05	0.00	-0.13	-0.16
developed	LU/LC	H	0.02	0.31	-0.15	-0.08	-0.06	0.13	-0.06	-0.24	-0.10	-0.21	-0.23	0.28	0.03	0.08	0.26
forest	LU/LC	H	-0.03	-0.70	0.26	0.20	0.50	-0.02	0.45	0.24	-0.08	0.15	0.04	-0.46	0.53	-0.34	-0.72
wetland	LU/LC	H	0.03	-0.08	-0.10	0.11	-0.13	0.02	0.07	0.00	0.00	-0.12	0.06	-0.02	0.00	0.02	0.14

Table S4. A table showing the number of lakes in each surface connectivity class within each cluster.

Cluster	Isolated	Headwater	Drainage	Terminal	Total
4	8	17	34	3	62
11	12	2	12	1	27
2	6	17	60	3	86
7	7	12	84	1	104
15	13	6	43	1	63
14	11	8	118	4	141
1	4	0	23	0	27
13	7	4	16	1	28
10	18	11	80	1	110
12	8	2	10	0	20
3	31	10	145	1	187
5	12	2	52	3	69
9	14	12	125	3	154
6	10	5	23	3	41
8	34	4	44	17	99
Total	195	112	869	42	1218

Table S5. Results of two-sample t-tests that examined the differences in developed land use and road density between lakes with high and low elevations. Low-elevation lakes had higher developed land use and road density than high-elevation lakes. The breakpoint of the elevation was 120m which was determined based on the partial dependence plot.

	Percentage developed land use	Road density
t-score	35.81	18.46
p-value	<0.001	<0.001