Slow true polar wander around varying equatorial axes since 320 Ma

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Key points:

- New estimates of the magnitude and rate of true polar wander during the last 320 million years
- True polar wander mostly occurred about two nearly orthogonal equatorial axes
- No evidence for fast (>1°/Ma) true polar wander rotations of >5 Ma since 320 Ma

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Abstract

True polar wander (TPW), the rotation of the solid Earth relative to the spin axis, is driven by changes in the Earth's moment of inertia induced by mantle convection and may have influenced past climate and life. Long-term TPW is typically inferred from large polar shifts in paleomagnetic apparent polar wander paths or computed directly by rotating them in a mantle reference frame. However, most apparent polar wander paths do not incorporate uncertainties in paleomagnetic data, which may bias estimates of TPW. Here, we provide new quantitative estimates of TPW since 320 Ma by placing a recent global apparent polar wander path corrected for age bias and with improved uncertainty quantification in existing mantle reference frames. We find large amplitude (>10°) but slow TPW rotations that predominantly occurred about two equatorial axes that are approximately orthogonal. During the Triassic and Jurassic, a ~24° TPW oscillation occurred about an axis at ~15°W, close to the present-day TPW axis at ~10°E. In contrast, the TPW axis was located at ~85°E during a smaller oscillation (~6–10°) over the past ~80 Ma, as well as between 260 and 320 Ma. We propose that these varying TPW axes reflect changes in the distribution and flux of subduction in the Tethyan and Pacific realms. We find no evidence for previously postulated fast (>1°/Ma) TPW oscillations in the Cretaceous and Jurassic. Finally, we propose that calibrating mantle convection models against reconstructed TPW will improve our understanding of mantle dynamics and the drivers of TPW itself.

Plain Language Summary

True polar wander is the rotation of the Earth's crust and mantle relative to the spin axis. On geological timescales, true polar wander is caused by movements of heavier or lighter material in the mantle, such as sinking tectonic plates. To compensate for these movements, the Earth rebalances itself by a rotation around an axis located on the equator. These rotations change the position of all continents simultaneously, influencing their latitude and potentially climate and life. Additionally, the location of the axis and the speed of true polar wander can provide important insights on the structure and movement of material in the Earth's mantle. In this study, we calculated the magnitude, speed, and axis of true polar wander for the last 320 million years. Our results suggest that while the rotations were large (>10°), they were relatively slow. True polar wander mainly occurred around two different equatorial axes: one close to the present-day axis and the other about 90° away. The changing axis may be caused by changes in the location and amount of sinking tectonic plates over time. Lastly, we propose that models of mantle convection could help us better understand the drivers of true polar wander.

1 Introduction

Determining the rate of convection of the Earth's mantle is key in deciphering the drivers of plate tectonics and volcanism but is notoriously difficult to quantify from kinematic observations alone. True polar wander (TPW) provides an avenue to kinematically constrain mantle convective processes. TPW is the rotation of the Earth's mantle and crust (i.e., the solid Earth) relative to the spin axis, such that the axis of maximum nonhydrostatic moment of inertia remains closely aligned with the spin axis. On geological timescales, TPW is driven by the redistribution of density anomalies within the Earth's mantle (Goldreich & Toomre, 1969; Evans, 2003), which include sinking subducted slabs or rising mantle plumes (Steinberger and Torsvik, 2010). TPW rotations occur, by definition, about an axis located in the equatorial plane, bringing excess masses towards the equator and mass deficits towards the geographic poles (Gold, 1955; Evans, 2003). The magnitude and rate of TPW are primarily controlled by the spatial distribution and magnitude of mass heterogeneities and by the viscosity of the mantle (e.g., Spada et al., 1992; Tsai & Stevenson, 2007; Rose & Buffett, 2017). Quantitative estimates of TPW therefore provide kinematic constraints on the structure, rheology, and dynamics of the Earth's mantle. The primary tool to quantify TPW is paleomagnetism (e.g., Besse et al., 2021). However, estimates of TPW obtained through paleomagnetism vary significantly, and the magnitude and rate at which TPW occurred in the geological past are uncertain.

Long-term (≥5 Ma) TPW is often quantified using apparent polar wander paths, which track the motion of the time-averaged paleomagnetic pole, assumed to coincide with the Earth's spin axis, relative to a tectonic plate (e.g., Besse and Courtillot, 2002; Torsvik et al., 2008). Phanerozoic TPW rotations have been determined by identifying common rotations of major continents observed from paleomagnetic data (e.g., Jurdy & Van der Voo, 1974; Steinberger & Torsvik 2008; Torsvik et al., 2012, 2014). This approach has yielded relatively slow (<1°/Ma) TPW rates, but large in amplitude (up to >20° in the Paleozoic and Mesozoic; Torsvik et al., 2014). On the other hand, rapid shifts in the position of the paleomagnetic pole have also been interpreted as evidence for fast TPW (>1°/Ma). Proposed episodes of fast TPW include a short-lived episode at ~84 Ma (Gordon, 1983; Sager & Koppers, 2000; Mitchell et al., 2021), a ~30-40° polar shift during the Late Jurassic referred to as the 'Jurassic monster polar shift' (e.g., Kent & Irving, 2010; Kent et al., 2015; Muttoni & Kent, 2019), a ~50° polar shift in the Ordovician-Silurian (Jing et al., 2022) as well as a ~90° degrees oscillation in the Ediacaran (e.g., Mitchell et al., 2011; Robert et al., 2017). The occurrence of such fast and large amplitude TPW events not only has large implications for the structure and dynamics of the Earth's interior but may also have had profound consequences for e.g., the biosphere, sea level, climate and the geodynamo (e.g., Evans, 2003; Raub et al., 2007; Biggin et al., 2012; Muttoni et al., 2013; Jing et al., 2022; Domeier et al., 2023; Wang & Mitchell, 2023). However, the existence of these fast TPW episodes requires high-resolution paleomagnetic data with small age uncertainties and therefore remains controversial: several recent studies have questioned whether these polar shifts truly represent rapid TPW or rather represent paleomagnetic artifacts, noise, or non-dipole behavior of the Earth's magnetic field (e.g., Kulakov et al., 2021; Cottrell et al., 2023; Domeier et al., 2023).

The most direct way to quantify the rate and magnitude of TPW is through the comparison of paleomagnetic reference frames derived from an APWP and mantle reference frames that estimate plate tectonic motions relative to the ambient mantle, such as a hotspot reference frame (e.g., Livermore et al., 1984; Andrews, 1985; Besse & Courtillot 2002; Doubrovine et al., 2012). Placing an APWP in a mantle reference frame enables the construction of a TPW path that provides a direct estimate of the motion between the spin axis and 'mean' solid Earth. However, the record of well-defined hotspot tracks so far limited the reliable application of this approach to computing TPW back to the Early Cretaceous (~120 Ma; Torsvik et al., 2008; Doubrovine et al., 2012). Studies using the most recent hotspot reference frames obtained TPW paths that show slow (<1.0°/Ma) but significant (up to 10-20° in magnitude) TPW back to Mesozoic times (e.g., Besse & Courtillot, 2002; Doubrovine et al., 2012).

A clear limitation of most published estimates of TPW is that they were obtained using conventional, pole-based apparent polar wander paths, in which spatial and temporal uncertainties in the underlying data were not incorporated (e.g., Besse & Courtillot, 2002; Torsvik et al., 2012). The recent global APWP of Vaes et al. (2023) is the first in which these sources of uncertainty were propagated. One of the key steps forward of that global APWP is that estimates of APW rates may now be corrected for temporal bias, by computing the 'effective' age of the reference pole positions of the APWP from the age distribution of the paleomagnetic data used as input for those poles. Vaes et al. (2023) showed that observed peaks in APW rate that have previously been interpreted as phases of relatively rapid TPW, such a spike between 110 and 100 Ma (Steinberger & Torsvik, 2008; Torsvik et al., 2012; Doubrovine et al., 2012), disappeared after correcting for the 'effective' age difference between successive reference poles of the APWP.

Here, we derive new quantitative estimates of the magnitude, direction, and rate of TPW since 320 Ma by comparing the recent global APWP of Vaes et al. (2023) with previously published mantle reference frames. We first compute TPW paths that track the motion of the time-averaged paleomagnetic pole relative to the deep mantle for the last 120 Ma using mantle frames based on hotspots with different underlying assumptions (Müller et al., 1993; Torsvik et al., 2008; Doubrovine et al., 2012). Next, we compute TPW back to 320 Ma using a recent mantle reference frame of Müller et al. (2022) that is based on a set of tectonic 'rules' and was computed for the last 1 Ga. We test previous observations that show that TPW was limited since the mid-Cretaceous and analyze whether the axis of TPW has remained approximately stable, as previously proposed, or whether TPW rotations occurred around changing equatorial axes since 320 Ma. Next, we re-assess whether fast polar shifts that were previously interpreted as phases of TPW may truly represent rapid TPW or rather may have resulted from noise induced by age uncertainty or the use of paleopole averages in determining polar wander. Finally, we discuss the implications of our results for analyzing mantle dynamics and the role of deep-mantle structure in determining the axis of TPW.

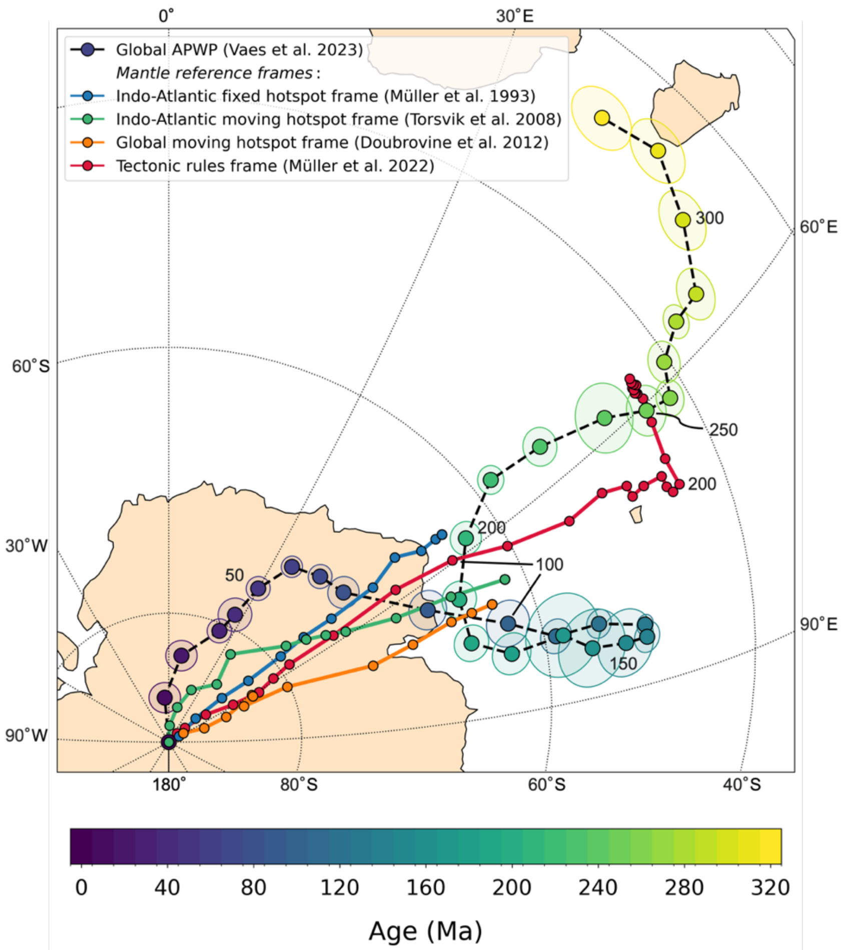

Fig. 1. Comparison between the global APWP for the last 320 Ma of Vaes et al. (2023) and the four mantle reference frames used in this study to derive TPW. The global APWP is plotted in South African coordinates, thus showing the position of the Earth's spin axis relative to a fixed South Africa. The pole positions were computed at 10 Ma steps using a 20 Ma sliding window (Fig. S1, Table S2). Reference poles and 95% confidence regions are colored by their age. The four polar wander paths constructed using the mantle reference frames indicate the motion the 'mean' mantle relative to a fixed (South) African plate (Table S3). In the absence of TPW, this would correspond to the motion of the spin axis relative to the Africa plate. Note that these paths based on the mantle reference frames only go back to 120 Ma, except for the tectonic rules frame of Müller et al. (2022), which is plotted here back to 320 Ma.

2 Methods

We construct true polar wander paths (TPWPs) using four different mantle reference frames (Fig. 1, Table S1): the Indo-Atlantic fixed hotspot reference frame of Müller et al. (1993), the Indo-Atlantic moving hotspot reference frame of Torsvik et al. (2008), the global moving hotspot reference frame of Doubrovine et al. (2012), and the recently published mantle reference frame of Müller et al. (2022)

that was constructed using an optimization approach based on a set of 'tectonic rules' (Tetley et al., 2019). We computed each TPWP as an apparent polar wander path in a coordinate frame in which the mantle is kept fixed, following the approach of many previous workers (e.g., Livermore et al., 1984; Andrews, 1985; Besse & Courtillot, 2002; Doubrovine et al., 2012). To this end, we placed the recent global APWP of Vaes et al. (2023) in each of the mantle reference frames. We computed the TPWPs using a time step of 10 Ma, which is the temporal resolution at which both paleomagnetic and mantle reference frames are typically computed, as well as estimates of TPW that are derived from those (e.g., Steinberger & Torsvik, 2008; Torsvik et al., 2014). To account for the uneven age distribution of the paleomagnetic input data, Vaes et al. (2023) computed the 'effective' age of the reference poles and showed that this age is often significantly different from the center age of the 20 Ma time window used to compute each reference pole. To construct a global APWP at exact 10 Ma time steps, we re-computed the global APWP here using same paleomagnetic database, global plate circuit and iterative approach used by Vaes et al. (2023). As an additional step, we interpolated between the reference poles generated for each iteration to obtain a cloud of simulated reference poles with the same age (e.g., 10, 20, 30 Ma; Fig. S1). We emphasize that in this approach, both the spatial and temporal uncertainties in the underlying paleomagnetic data are propagated into the confidence regions of the global APWP. We quantify the 95% confidence regions of this interpolated APWP by the P_{95} cone of confidence, which includes 95% of the simulated reference poles computed for each time step, following the approach developed by Vaes et al. (2022, 2023) (Fig. S1). Finally, to compute the TPWPs, we placed the interpolated global APWP in each mantle reference frame by rotating it using the Euler rotation poles (defined at 10 Ma steps) that describe the motion of the South African plate relative to the mantle. The resulting polar wander path then describes the motion of the time-averaged paleomagnetic pole relative to the ambient mantle, in the same way as a conventional APWP describes the motion of the pole relative to a fixed tectonic plate.

It is important to note that in our computation of the TPWP, the uncertainties associated with the mantle reference frames are not (yet) incorporated. The uncertainties of the rotation poles for Africa relative to the mantle were not quantified for the fixed hotspot reference frame of Müller et al. (1993) and the tectonic rules frame of Müller et al. (2022). 95% confidence regions for the two moving hotspot reference frames were provided by Doubrovine et al. (2012) and Torsvik et al. (2008). In addition, we note that the plate circuit and time scale used to determine the mantle reference frames differs between all four models and is also different from the updated plate circuit as well as the time scale (of Gradstein et al., 2020) used for the construction of the global APWP of Vaes et al. (2023). Because the four reference frames are based on different input datasets, underlying assumptions, and methodologies, they each provide a different perspective on the motion of South Africa relative to the deep mantle.

The three hotspot-based mantle frames used here are assumed to determine the 'absolute' motion of tectonic plates relative to the deep mantle using hotspot tracks: linear chains of intraplate volcanoes that show a clear progression of eruption ages (e.g., Wilson, 1963; Morgan, 1971; 1981).

The geometry and age progression of these hotspot tracks have been widely used to reconstruct lithospheric motions relative to mantle upwellings (or 'plumes', Morgan, 1971), which are either assumed to be stationary in a 'mean mantle' frame (i.e., a fixed hotspot reference frame, e.g., Morgan, 1981; Müller et al., 1993) or corrected for slow relative motions using a numerical model of mantle convection (i.e., a moving hotspot reference frame; O'Neill et al., 2005; Torsvik et al., 2008; Doubrovine et al., 2012). Moving hotspot reference frames, however, do not provide independent kinematic constraints on mantle dynamics because a mantle convection model, constrained by a global model of plate tectonic motions and mantle density heterogeneities, is used as input, after which the reference frame is iteratively constructed to fit hotspot tracks (e.g., Torsvik et al., 2008; Doubrovine et al., 2012). In addition, the disadvantage of hotspot reference frames is that the availability of linear hotspot tracks only allows the determination of plate motions relative to hotspots back to the Early Cretaceous (~120 Ma).

The recent mantle reference frame of Müller et al. (2022) was constructed in an entirely different way. Building on the work of Tetley et al. (2019), they used an iterative approach in which the absolute plate motion of the reference plate Africa is obtained by optimizing a global plate model using four different criteria. These criteria include the restriction of the net lithospheric rotation rate to values deemed reasonable from numerical experiments, the minimization of global trench migration velocities, a global continental median plate velocity below 6.0 cm/a, and the minimization of the spatio-temporal misfit between observed and model-predicted hotspot tracks (only used for the last 80 Ma). We note that this type of mantle frame may be biased by subjective choices related to the relative weighting of the different rules and the imposed limits. On the other hand, the joint incorporation of multiple kinematic constraints enables the construction of a mantle frame that is less likely to suffer from errors in or the overfitting of a single kinematic observation such as a hotspot track (Tetley et al., 2019). Using the global plate model for the last billion years of Merdith et al. (2021) as input for their model, Müller et al. (2022) constructed a tectonic rules-based mantle reference frame since 1000 Ma, using 5 Ma time steps. No attempts were made to quantify the uncertainties for this frame, however. Nonetheless, taking this frame at face value enables the computation of a TPW path from 320 Ma to present, using the global APWP for the last 320 Ma of Vaes et al. (2023). Finally, to assess whether fast TPW phases may have occurred during the last 320 Ma, we also computed a TPWP at a 5 Ma resolution, following the same procedure as described above, but using the 5-Ma-resolution-APWP of Vaes et al. (2023) instead.

3 Results

3.1 True polar wander paths for the last 120 Ma

We first computed four true polar wander paths for the last 120 Ma using the four different mantle reference frames (Fig. 2, Table S3). The paths show the motion of the time-averaged paleomagnetic pole position, which is assumed to coincide with the spin axis, relative to the 'mean' mantle. The

TPWPs thus visualize the motion of the spin axis in a fixed mantle frame, providing direct estimates of the magnitude, rate, and direction of TPW. The four TPWPs show a similar first-order geometry for the last 120 Ma (Fig. 2a-d). All four paths show a back-and-forth motion of the spin axis during the last 80 Ma, roughly along the 180° meridian. From 80 to 60 Ma, all paths show a shift in the pole position away from the geographic pole (in a fixed mantle frame), with the net TPW angle peaking at 60 Ma in all four TPWPs. Except for the path based on the global moving hotspot reference frame of Doubrovine et al. (2012), a near stillstand of the pole position is observed between 60 and 30 Ma, as indicated by the TPW rate of ~0.1°/Ma (Fig. 2e). The net angle of TPW of >10° at 60 Ma is notably larger for the Doubrovine et al. (2012) model than the ~6° for Müller et al. (1993) and Torsvik et al. (2008) models. The three TPW paths based on a hotspot reference frame show a sharp cusp at 80 Ma. Prior to 80 Ma, the trajectory of the TPWPs based on the Müller et al. (1993) and Doubrovine et al. (2012) frames is sub-parallel to the great circle around an equatorial TPW axis at 11°E, which is the proposed long-term TPW axis of Torsvik et al. (2014) that closely corresponds to the present-day axis of TPW (Pavoni, 2008). The TPWP computed using the Torsvik et al. (2008) frame shows a similar direction between 80 and 100 Ma but is perpendicular to this direction between 100 and 120 Ma. We observe very similar TPW rates for all four paths (Fig. 2e). For most of the last 120 Ma, including the entire Cenozoic, the TPW rate is between ~0.1°/Ma and 0.4°/Ma. The highest TPW rates are obtained for the Late Cretaceous, peaking at a rate of 0.5-0.6°/Ma. Notable, we find that the TPW rate since 120 Ma – on a timescale of 10 Ma - stays well below 1.0°/Ma, which we defined as the threshold for 'fast' TPW, following Cottrell et al. (2023).

3.2 A true polar wander path back to 320 Ma

We computed a TPWP back to 320 Ma by placing the interpolated global APWP of Vaes et al. (2023) in the tectonic rules frame of Müller et al. (2022) (Fig. 3, Table S4). The resulting TPWP can be roughly divided into five segments (Fig. 4a). For the last 130 Ma, the TPW path shows an oscillatory motion with a cusp at 60 Ma. During this time interval, the path runs approximately parallel to the 5°W meridian, depicted in Fig. 4a by the red line. The Jurassic-Triassic portion of the TPWP shows two smooth segments, from 130 to 200 Ma and from 200 to 260 Ma, that are nearly parallel. The trajectory corresponds to successive large TPW rotations around an equatorial axis located at ~15°W (blue line in Fig. 4a). During both these time spans, the spin axis is estimated to move ~24° relative to the stable mantle. Note that the position of the spin axis relative to a fixed mantle is close to the present-day position at 260-250 Ma, just before the large back-and-forth TPW rotations. The oldest part of the TPWP, from 260 to 320 Ma, runs roughly parallel to the 5° W/175°E meridian, similar to the 0-130 Ma segments. The largest net TPW angle of approximately 21° is computed at 200 Ma and 320 Ma. The TPW rates computed for the last 320 Ma stay below 0.7°/Ma for this entire time span, with an average of 0.34°±0.14°/Ma (1σ).

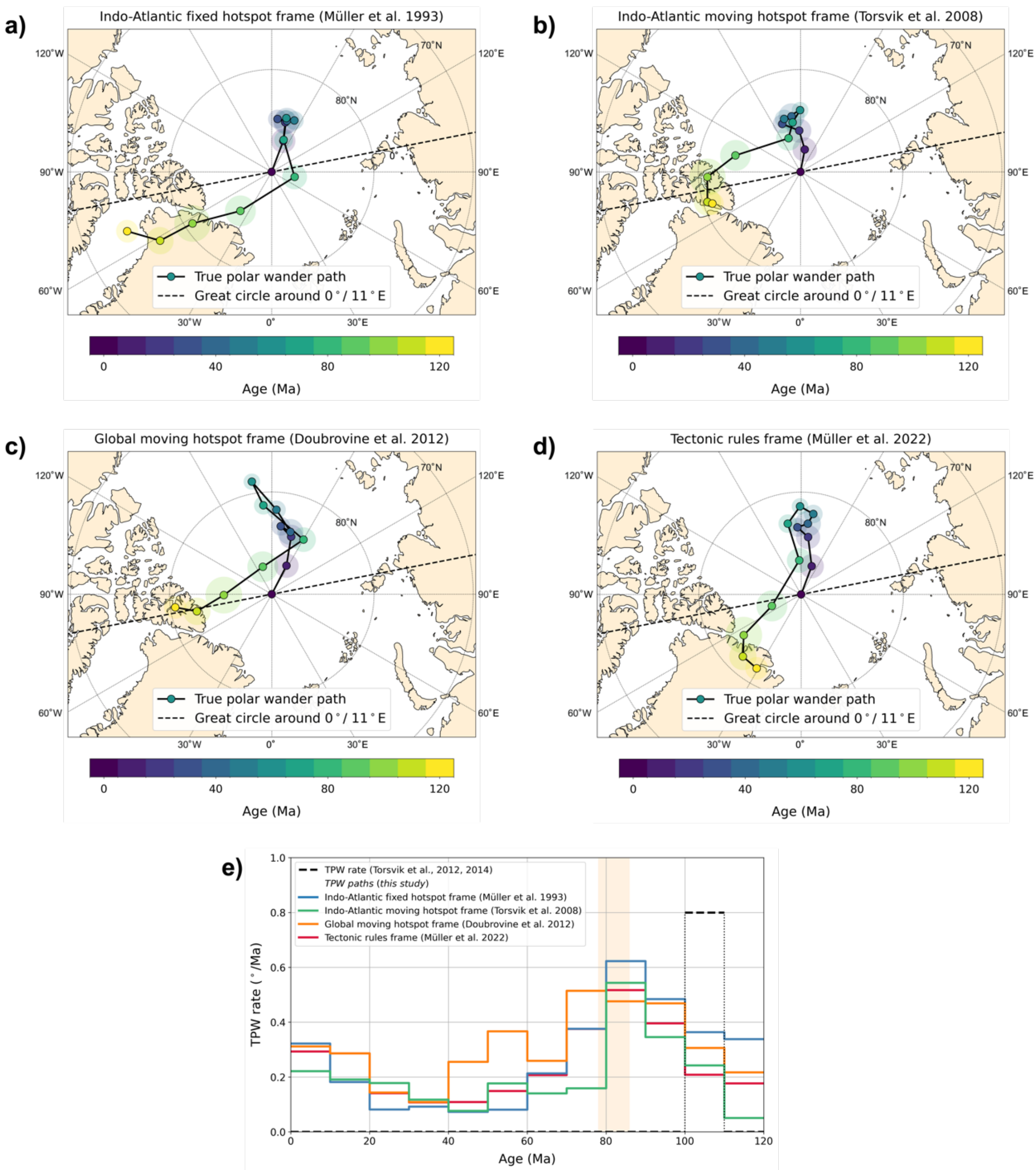

Fig. 2. True polar wander paths for the last 120 Ma, computed using four different mantle reference frames **(a-d).** The poles represent the location of the Earth's spin axis relative to a fixed 'mean' mantle at 10 Ma steps and are colored by their age. The confidence regions correspond to the paleomagnetic uncertainty only (that is, the P_{95} of the reference poles of the global APWP, see Vaes et al., 2023). The dashed black line shows the great circle around an equatorial rotation axis at 11°E, corresponding to the preferred TPW axis of Torsvik et al. (2012, 2014). The TPWPs are listed in Table S4. **e)** TPW rate for the last 120 Ma, estimated from each of the four TPW paths. The proposed phases of rapid TPW between 86 and 78 Ma by Mitchell et al. (2021) and between 110 and 100 Ma by Torsvik et al. (2012, 2014) are indicated by the colored band and dashed black line, respectively.

The TPWP computed at a 5 Ma resolution shows a similar overall trend but is more irregular and has much larger confidence regions (Fig. 5, Table S5). Again, we emphasize that these confidence regions are solely defined by the uncertainty in the global APWP and do not incorporate those of the mantle reference frame – the true uncertainty must thus be larger. This is a direct consequence of the lower amount of paleomagnetic data that underlies each pole position, given that a smaller time window of 10 Ma is used instead of 20 Ma. The main difference with the 10 Ma-resolution TPWP is that there are segments of the path showing a zig-zag pattern, for instance between 10 and 60 Ma (Fig. 5a). However, whether these motions are truly representative of TPW is questionable considering the relatively large and often overlapping (minimum) confidence regions. The larger uncertainty is also clearly reflected in the confidence regions of the APW rate of the 5-Ma-resolution-APWP (Fig. 5b). The TPW rates obtained from this higher resolution TPWP are ~50% higher than for the 10-Ma-resolution path, with a mean rate of 0.50°±0.27°/Ma (Fig. 5c). In addition, five peaks are observed with a TPW rate of >0.8°/Ma, with three peaks above >1.0°/Ma. It should be noted, however, that the latter peaks are observed for the pre-260 Ma portion of the TPWPs, where the 5 Ma-resolution-APWP is the least robust and where the tectonic rules mantle frame likely has higher uncertainty than for younger times.

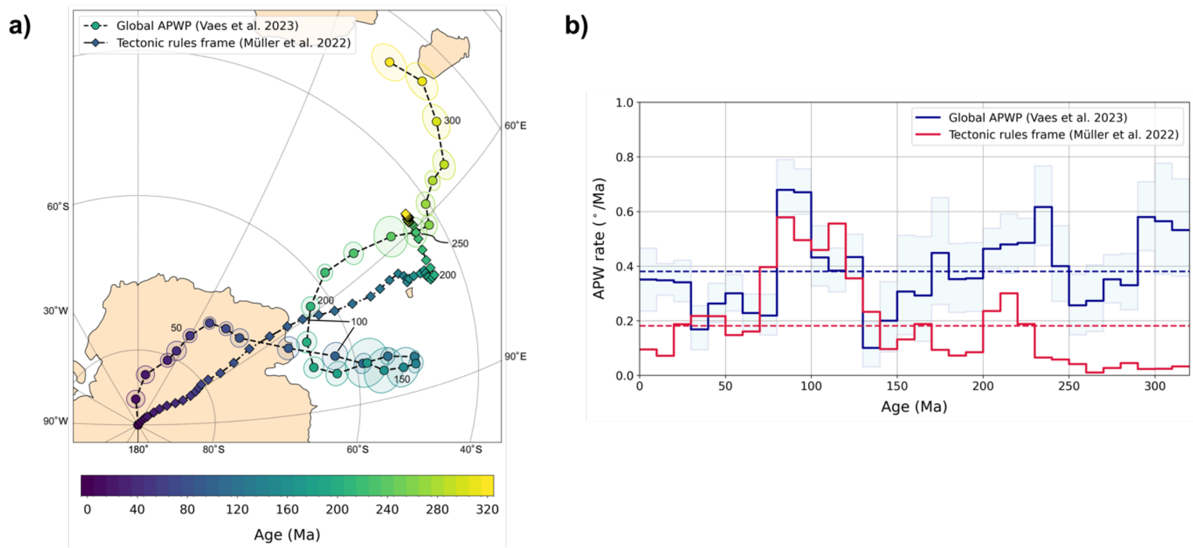

Fig. 3. a) Orthographic plot of the global APWP of Vaes et al. (2023) and the tectonic rules mantle reference frame of Müller et al. (2022). The poles indicate the position of the Earth's spin axis and 'mean' mantle relative to a fixed African plate, respectively. **b)** Comparison of the apparent polar wander rate derived from the pole paths shows in a). The horizontal dashed lines show the mean rate for the last 320 Ma. The light blue band indicates the 95% confidence regions on the rate that is computed from the uncertainty of the global APWP.

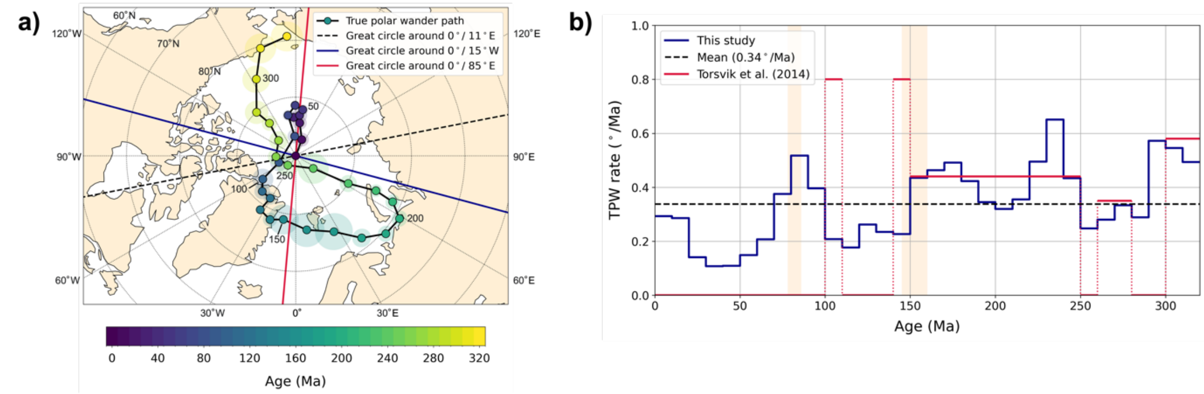

Fig. 4. a) True polar wander path for the last 320 million years, computed using the tectonic rules reference frame of Müller et al. (2022). Colors and confidence regions are the same as in Fig. 2. Great circles around three different TPW axes are shown in blue, red and black (dashed) lines, respectively. **b)** The estimated rate of TPW since 320 Ma, computed from the TPW path shown in a). The mean TPW rate (0.34°/Ma) is indicated by the black dashed line. Previously inferred TPW rates by Torsvik et al. (2024) are shown in red. Note that this study identified multiple 'phases' of TPW, interspersed by time intervals without significant TPW. Time intervals for which rapid TPW (>1°/Ma) has been proposed are highlighted by the colored bands.

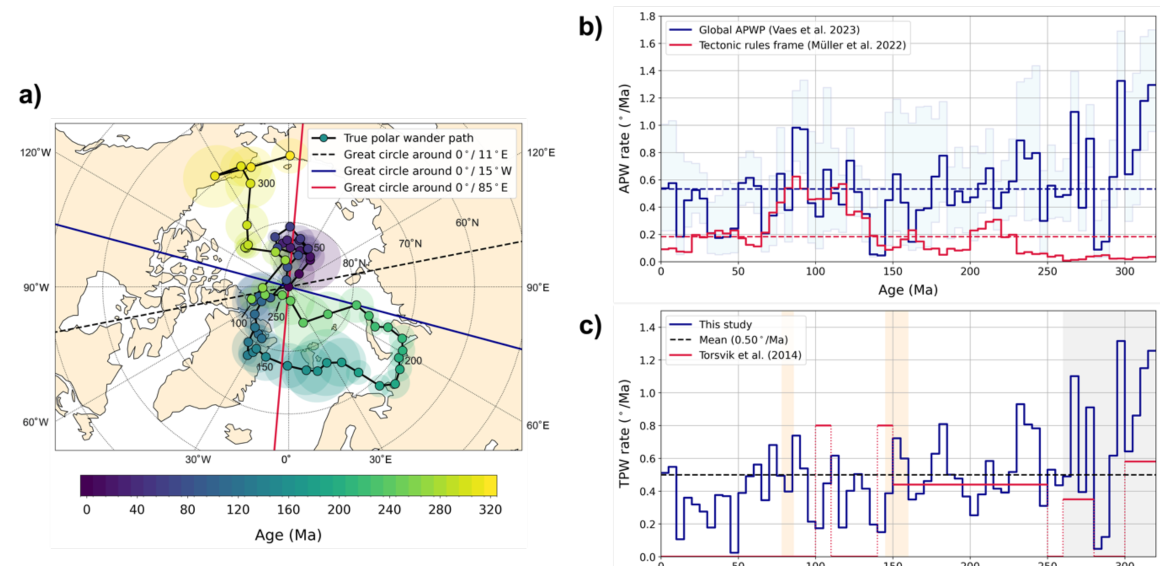

Fig. 5. a) TPW path computed at a 5 Ma resolution using the global APWP from Vaes et al. (2023) calculated at a 5 Ma time step using a 10 Ma time window (Table S5). **b)** Comparison of the polar wander rates derived from the 5-Ma resolution global APWP and tectonic rules mantle frame of Müller et al. (2022). The horizontal dashed lines show the mean rates. The light blue band indicates the 95% confidence regions on the rate that is computed from the uncertainty in the 5 Ma-resolution global APWP. **c)** Same as Fig. 4c, but now showing the TPW rate derived from the 5-Ma-resolution TPW path shown in **a)**. The grey area highlights the time interval for which the TPW path is less robust, leading to large variations in TPW rate that are likely the result of noise.

4. Discussion

4.1 True polar wander since 320 Ma

We have quantified the motion of the time-averaged paleomagnetic pole relative to the mantle using the global APWP of Vaes et al. (2023) and four different mantle reference frames. If these motions correspond to the movement of the Earth's spin axis relative to the mantle, they quantify the magnitude, rate, and direction of TPW. The similarities between the first-order geometry of the four TPWPs for the last 120 Ma indicates that the observed trends in polar motion are reproducible, even though these paths were constructed using mantle reference frames based on different datasets, plate circuits and approaches. Our results clearly show that TPW since 120 Ma is non-negligible: each TPWP yields an angular deviation of the spin axis relative to the mantle of >5° for both the early Cenozoic (50-60 Ma) and Early Cretaceous. This would argue for a rejection of the null hypothesis that no significant TPW occurred since the Late Cretaceous, in contrast to the conclusions reach in a recent paper by Cottrell et al. (2023), as well as to previous TPW estimates of Torsvik et al. (2014) (Fig. 4b). If the TPW rotations computed here are artifacts, this would imply a large and systematic error in either the global APWP or mantle reference frames. Given the well-constrained and small uncertainty of the global APWP of Vaes et al. (2023), this would suggest that there are large inadequacies in all mantle reference frames used here, for instance due to a large drift of all the hotspots or large, systematic errors in the relative plate motion circuit, which we consider unlikely.

The four TPWPs of the last 120 Ma all indicate a back-and-forth motion of the mantle relative to the spin axis during the last 80 Ma (Fig. 2a-d). This oscillatory motion was previously computed by Doubrovine et al. (2012) in a comparison of their moving hotspot frame and the global APWP of Torsvik et al. (2012). Intriguingly, the direction of TPW for the last 80 Ma is nearly orthogonal to the great circle about an equatorial axis at 11°E (Fig. 6), which is the axis of TPW rotation that would be expected if the Earth's moment of inertia is mainly controlled by the basal mantle structures referred to as large low shear-wave velocity provinces (LLSVPs; Steinberger & Torsvik, 2008; Torsvik et al., 2012, 2014) and that is close to the present-day TPW axis at ~10°E (e.g., Pavoni, 2008). The small component of TPW about this axis for the last 80 Ma is consistent with the findings of Torsvik et al. (2014), who modelled no significant TPW rotations around an equatorial axis located at 0°/11°E for the last 100 Ma (Fig. 7a). Instead, our TPWP for the last 320 Ma shows a ~9° oscillatory TPW rotation about an axis located at ~85°E (Figs. 6, 7a).

However, the estimated TPW axis between 80 and 120 Ma is closer to the 0°/11°E axis, with an estimated longitude of the rotation axis based on all four TPWPs of ~34° (Fig. 6a). The TPW axis determined from the three hotspot-based mantle frames is particularly in agreement with this TPW axis, which is also observed from the (sub-)parallel trajectories of the TPWPs to the great-circle about that axis (Fig. 2a-c). The TPWP based on the fixed hotspot frame of Müller et al. (1993) shows the largest rotation of ~15° about this axis, between 80 and 120 Ma, although we note that the fixed hotspot model prior to 80 Ma is considered much less reliable (e.g., Torsvik et al., 2008). The previously identified phase of relatively fast TPW (~0.8°/Ma) at 110-100 Ma (Steinberger & Torsvik, 2008; Torsvik et al., 2012, 2014) is, however, absent in our TPWPs (Fig. 2e). Vaes et al. (2023) showed that a spike in APW between 110 and 100 Ma in the global APWP of Torsvik et al. (2012) was likely the result of temporal bias, and we therefore interpret this inferred 110-100 Ma TPW event as a result of this artifact in the underlying APWP.

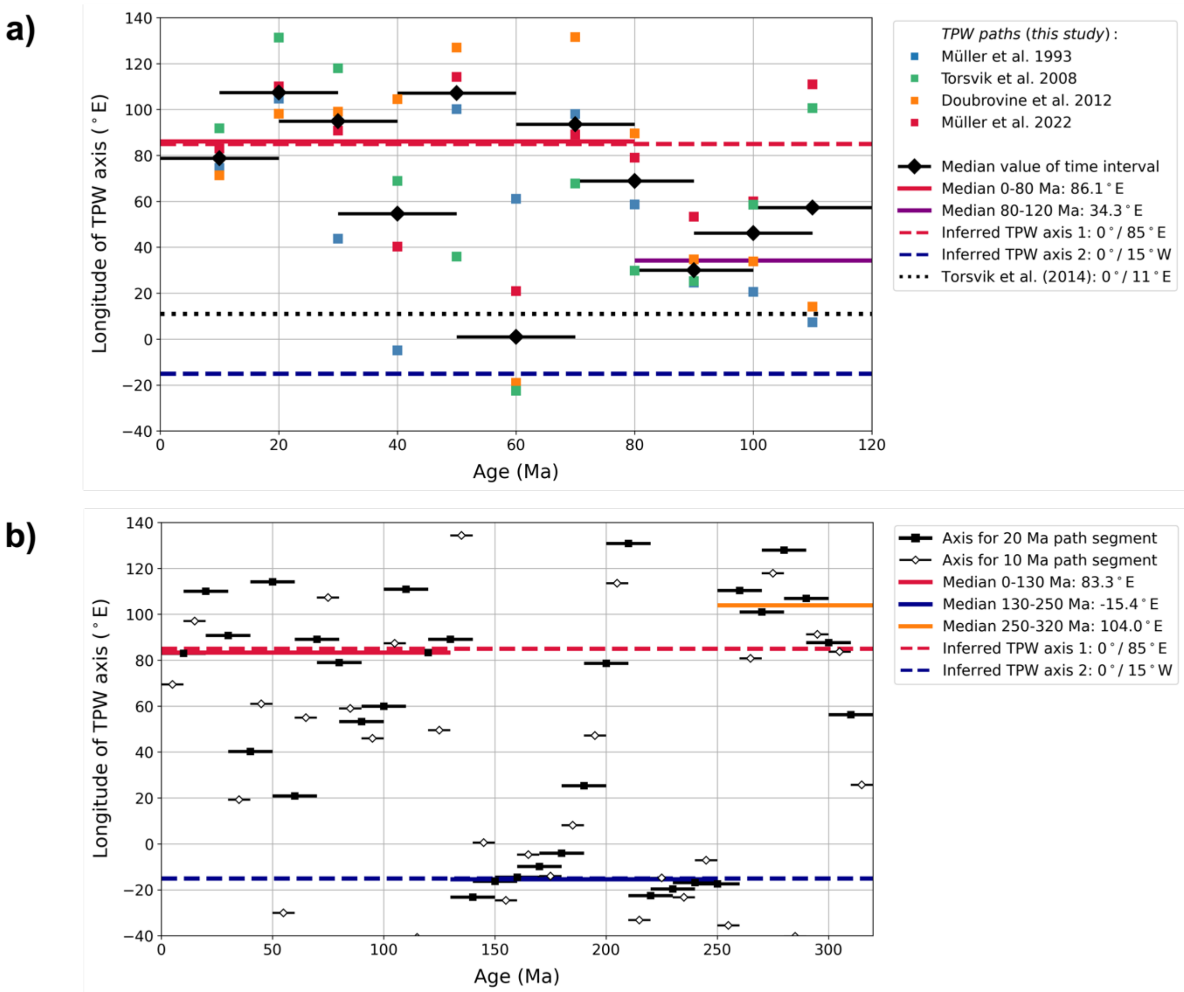

Fig. 6. Estimated position of the equatorial axis of TPW. **a)** Longitude of the TPW axis computed for 20 Ma segments (e.g., 0-20 Ma, 10-30 Ma) of each TPW path shown in Figs. 2a-d. The median value for each time interval is depicted with the thick black lines and diamonds. The median longitude for the 0-80 Ma and 80-120 Ma time intervals are plotted as red and purple horizontal lines. Inferred main TPW axes are shown as dashed lines. The dotted black line is the preferred axis of TPW from Torsvik et al. (2014). **b)** Longitude of the TPW axis determined for both 10 and 20 Ma segments of the TPW path for the last 320 Ma, shown in Fig. 4a. Median longitude values are shown for the 0-130, 130-250 and 250-320 Ma intervals.

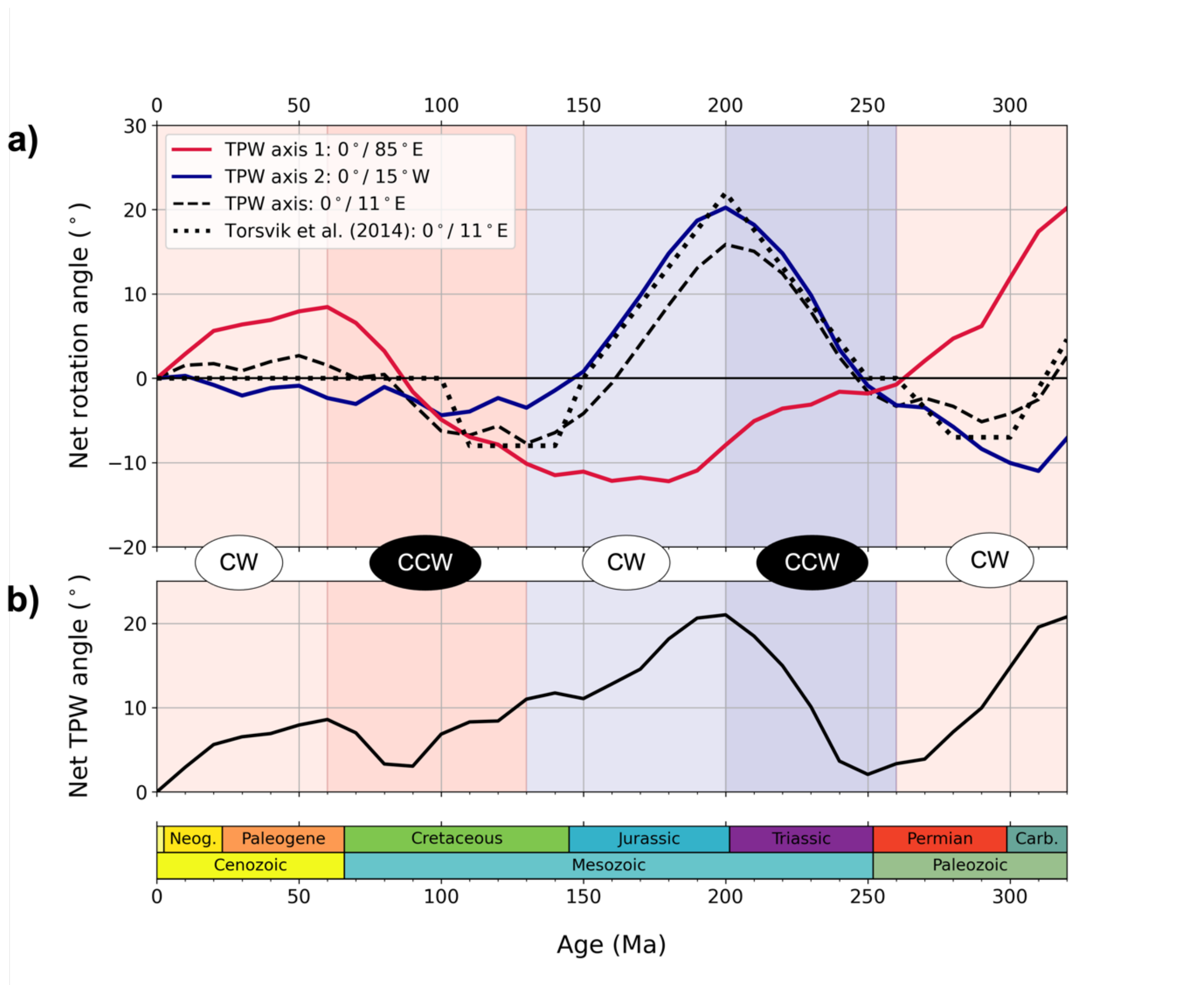

Fig. 7. Magnitude of TPW since 320 Ma. **a)** Net rotation angle around three different equatorial TPW axes. The angles are computed from the TPW path for the last 320 Ma (Fig. 4a) as the total rotation around each axis. Positive values indicate a net clockwise (CW) rotation since that time. Five major, long-term TPW rotations are identified for the last 320 Ma. Clockwise/counterclockwise (CCW) rotations around the inferred TPW axis 1 (0°/85°E) or TPW axis 2 (0°/15°W) are shown highlight by light/dark red or blue shading, respectively. The dotted black line shows the estimated net TPW rotations around the preferred TPW axis of 0°/11°E from Torsvik et al. (2014). **b)** Net TPW angle since 320 Ma, computed as the difference between the pole position (that is, the spin axis) relative to the mantle (that is, the geographic pole in Fig. 4a).

The TPWP computed for the last 320 Ma shows that TPW predominantly occurred around two equatorial rotation axes that are approximately orthogonal: one axis at ~85°E and a second axis located at ~15°W (Fig. 6). Remarkably, the timing and magnitude of true polar wander along the first axis are very similar to the TPW rotations determined by Torsvik et al. (2012, 2014) based on paleomagnetic data alone (Fig. 7), which built upon earlier work by Steinberger & Torsvik (2008). Our new TPWP shows a motion of the spin axis of ~24° away from the geographic pole between 260 and 200 Ma, and back parallel to the same great circle and with similar magnitude between 200 and 130

Ma (Fig. 7). The inferred rotation axes at a longitude of 15°W is very close to the estimated axes by Steinberger & Torsvik (2008), Torsvik et al. (2012) and Mitchell et al. (2012). These motions would correspond to a coherent counterclockwise rotation of the solid Earth relative to the spin axis between 260 and 200 Ma, followed by a clockwise rotation between 200 and 130 Ma, in agreement with previous findings of e.g., Torsvik et al. (2014). Our results confirm previous inferences by Torsvik et al. (2012, 2014) that cumulative TPW for the last 250 Ma is close to zero (Fig. 7b). Finally, we observe a TPW rotation of ~20° between 320 and 260 Ma, yielding an angular deviation of the spin axis relative to the mantle of ~20° for the Late Carboniferous (~320-310 Ma). This northward motion of Pangea towards the equator has previously been interpreted as the result of TPW (e.g., Le Pichon et al., 2021). However, in the approach of Torsvik et al. (2012, 2014), Africa is assumed to have remained longitudinally more or less stable and the axis of TPW is fixed at 0°/11°E, i.e., within Africa. With those underlying assumptions, those studies cannot determine paleolatitudinal motion of Africa (and Pangea) as TPW. The very low absolute plate motion of Africa (and thus Pangea) in the mantle reference frame of Müller et al. (2022) (Fig. 3b), which mostly results from minimizing continent motions relative to the mantle, suggests that most of Pangea's paleolatitudinal motion between 320 and 260 Ma is a result of TPW, which has some important implications for the structure of the deep mantle that will be discussed in section 4.3.

4.2 True polar wander: not so fast?

Phases of rapid TPW are frequently proposed based on rapid shifts in paleomagnetic pole positions. The existence of episodic, and possibly oscillatory, phases of rapid TPW are actively investigated by the paleomagnetic community. They are particularly interesting because it may significantly impact regional climate, paleoenvironment, sea level and life (e.g., Raub et al., 2007; Muttoni et al., 2013; Jing et al., 2022; Wang & Mitchell, 2023; Muttoni et al., 2024). Whether such fast polar shifts truly represent rapid TPW is not always straightforwardly determined, and it may often be difficult to distinguish signal from noise (e.g., Evans, 2003; Kulakov et al., 2021; Cottrell et al., 2023; Domeier et al., 2023). Importantly, the many inherent uncertainties in paleomagnetic data may result in a significant deviation of individual paleomagnetic pole positions from the estimated time-averaged pole given by an APWP, as reflected in the dispersion of poles obtained from similar-aged rocks (Rowley, 2019; Vaes et al., 2022). TPW shifts derived from the data of a single tectonic plate and/or from a small number of paleomagnetic poles, each based on different numbers of paleomagnetic samples/sites and thus averaging the magnetic field to different extents (Vaes et al., 2022), may be prone to unrecognized biases and any conclusions drawn from such analyses should be approached with caution.

For the last 120 Ma, all our TPWPs show a peak in TPW rate during the Late Cretaceous (~80-90 Ma), with a rate of 0.5-0.6°/Ma (Fig. 2e). This peak coincides with a phase of fast TPW around 84 Ma that was already proposed decades ago (e.g., Gordon, 1983; Sager & Koppers, 2000), although these results are not without controversy (e.g., Cottrell & Tarduno, 2000). A recent study by Mitchell et al. (2021) argued for a phase of rapid, oscillatory TPW (~3°/Ma) between ~86 and 78 Ma, based on high-resolution paleomagnetic records from two partly overlapping stratigraphic sections in northern Italy. Their TPW rates are much higher than obtained in this study (Fig. 4b, 5b). However, Cottrell et al. (2023) recently showed argued these records contain a secondary overprint overlooked by Mitchell et al. (2021) and interpreted that the section was likely affected by local block rotations, causing a bias in the paleomagnetic directions and casting doubt on this interpreted fast TPW event.

The most notable potential episode of rapid TPW is the Jurassic 'monster' polar shift: a ~30-40° shift in pole position in the APWP of North America between ~160 and 145 Ma (Kent et al., 2015). High-quality paleomagnetic data from Adria, under the assumption that Adria was rigidly attached to Africa since the opening of the Ionian Sea between 170 and 154 Ma (Channell et al., 2022), confirm a ~30° shift in pole position in North American coordinates (Kent & Muttoni, 2019; Muttoni et al., 2024). However, the reliability of key paleomagnetic poles from North American kimberlites as well as the assumption that Adria was rigidly attached to northern Africa have been challenged (e.g., van Hinsbergen et al., 2020; Kulakov et al., 2021; Mirzaei et al., 2021). Since TPW affects all tectonic plates and continents simultaneously, the existence of a rapid TPW phase may be tested at different (paleo-)locations. Intriguingly, different paleomagnetic datasets from a range of tectonic plates provide both evidence in support of fast polar motion during the Late Jurassic, such as from North China (Yi et al., 2019), the Lhasa terrane (Ma et al., 2022) and the Pacific plate (Fu & Kent, 2018), or against it, such as from North China (Gao et al., 2021), Greenland (Kulakov et al., 2021) and South America (Ruiz González et al., 2022).

Our results, based on a global APWP with high spatiotemporal data coverage and with propagated age uncertainty and spatial uncertainty in the underlying paleomagnetic data, do not provide any evidence for rapid (>1.0°/Ma) TPW during the last 320 Ma (Figs. 4b, 5c). We acknowledge that the absence of relatively fast TPW may, in theory, be a consequence of using a 20 Ma time window for the global APWP that underlies the TPWPs presented here. This may smoothen or obscure potential short-term (<10 Ma) and/or small-amplitude phases of polar wander (e.g., Muttoni et al., 2005, 2024; Mitchell et al., 2021). It is interesting to note, however, that the highest TPW rates are observed during the middle of a long, large-amplitude TPW rotation (Fig. 4b). This is consistent with theoretical inferences that the fastest TPW is expected to occur during longer term episodes of TPW (Goldreich & Toomre, 1969; Tsai & Stevenson, 2007; Cottrell et al., 2023). We indeed observe that TPW rate accelerates until a peak velocity is reached before decelerating again, such that the fastest TPW rate during each long-term TPW rotation (see Fig. 7) is greater than the average TPW rate for that rotation (Fig. 4b). Then again, even when computing a TPWP at a 5 Ma resolution and with a 10 Ma time window, we do not find clear evidence for TPW occurring at velocities beyond 1°/Ma. The peaks in TPW rate of more than 1.0°/Ma observed for the 5-Ma-resolution TPWP (Fig. 5c) are not statistically significant and do not follow the expected velocity pattern clearly observed for the 10-Ma-resolution TPWP, and we therefore conservatively interpret these as noise rather than a true 'signal'.

To establish whether short (<5 Ma) phases of fast TPW have occurred since the Late Paleozoic requires an increase in the resolution of the APWPs used to quantify TPW, as well as the quantification of the uncertainty in the mantle reference frame. Before these become available, such rapid changes in pole position may instead be identified by collecting well-dated, paleomagnetic datasets from sedimentary sequences with a high sampling resolution, which may provide high-resolution records of shifts in the paleomagnetic declination and inclination (e.g., Mitchell et al., 2021; Vaes et al., 2021). However, observed shifts in the paleomagnetic direction may result from paleomagnetic and/or tectonic artifacts, and should be thoroughly tested for potential biases. The collection of multiple high-resolution paleomagnetic records from stratigraphic sections of overlapping age and from different tectonic plates provides a way to test whether an observed rapid polar shift may indeed represent TPW. In addition, higher-resolution APWPs computed from site-level paleomagnetic data using a bottom-up uncertainty propagation scheme (Gallo et al., 2023) may allow the identification of fast and short-lived phases of TPW.

4.3 Linking true polar wander and mantle dynamics

The rate of TPW provides first-order kinematic constraints on the rate of mantle convection, particularly on the velocity at which density anomalies, such as sinking lithospheric slabs, move through the Earth's mantle. Fu et al. (2022) presented a compilation of paleomagnetically estimated TPW rates since the late Mesoproterozoic (~1100 Ma), proposing a direct link between the secular change of the TPW rate and the thermal structure and nature of convection in the mantle. Fast TPW (>1°/Ma) has often been inferred for pre-Pangean TPW, suggesting a different geodynamic regime that may correlate with the supercontinent cycle (e.g., Evans, 2003; Mitchell, 2014; Fu et al., 2022). Testing pre-Mesozoic TPW hypotheses is not straightforward, as plate motion-induced polar wander may be difficult to distinguish from TPW and thus obscure past TPW signals (Evans, 2003; Torsvik et al., 2014). Moreover, rapid polar shifts of up to ~90° during the Early Cambrian and Ediacaran have also been attributed to TPW (e.g., Kirschvink et al., 1997; Mitchell et al., 2011; Robert et al., 2017), but may alternatively be explained by unusual geomagnetic field behavior (Domeier et al., 2023; Robert et al., 2023).

Reproducing observed TPW by numerical modeling has also proven difficult, as this requires detailed knowledge on the structure of the mantle, such as the volumes and distribution of subducted lithospheric slabs (see Steinberger & Torsvik, 2010). Paleomagnetism-based estimates of TPW rates are frequently compared to TPW speed limits inferred from geodynamic modelling. However, these predictions vary substantially. For instance, Tsai and Stevenson (2007) provide a theoretical speed limit of 2.4°/Ma but find a maximum TPW shift of 8° for a period of 10 Ma (corresponding to 0.8°/Ma). On the other hand, Greff-Lefftz and Besse (2014) argue that TPW may occur at rates of up to 10°/Ma for larger mass reorganizations. These estimates heavily rely on the choice of rheological parameters, which are poorly constrained, particularly for deep geological time (Tsai & Stevenson, 2007; Steinberger & Torsvik, 2010). Moreover, whether such geodynamical models used to determine the magnitude and rate of TPW have slab sinking rates that are consistent with independent kinematic constraints on the slab sinking velocity versus depth, such as those obtained from seismic tomographic images of subducted slabs (e.g., van der Meer et al., 2018), it not always clear (Steinberger & Torsvik, 2010; van der Wiel et al., 2024).

Recent modelling studies show that TPW rotations are likely stabilized by the contribution to the moment of inertia by the two antipodal LLSVPs, with important contributions from subducted slabs, whose relative contribution to the moment of inertia strongly change with depth (e.g., Steinberger & Torsvik, 2010; Steinberger et al., 2017). Observed TPW rotations about an axis of minimum moment of inertia that is close to the combined center of mass of the LLSVPs suggests that TPW may be controlled by these structures (Torsvik et al., 2012; 2014). TPW rotations along an axis nearly orthogonal to this 0°/11°E axis (see Figs. 6, 7) instead indicate that the reorientation of the solid Earth may (also) be controlled by other contributions to the moment of inertia tensor, e.g., resulting from large density anomalies caused by subducting slabs at relatively high latitudes ~90° from this alternative axis (Steinberger & Torsvik, 2010). If driven by changes in subduction configuration and slab sinking, the two approximately orthogonal axes of TPW identified here (Figs. 6, 7) roughly coincide with low-latitude subduction in the Tethyan realm for the ~0°/15°E axis, whereas the ~0/85°E axis would correspond to changes in subduction flux in the Pacific realm. Our calculations show that the magnitudes of TPW are roughly equal for the two, or even slightly larger for the effects of Tethyan subduction, even though Tethyan subduction occurs at much lower latitudes than Pacific subduction (e.g., Seton et al., 2012; Vaes et al., 2019; Boschman et al., 2021). This may illustrate the stabilizing effects of LLSVPs and could allow the detailed computation of their absolute density from their contribution to the moment of inertia.

Intriguingly, our results show a ~20° TPW rotation for the period between 320 and 260 Ma around an axis orthogonal to 0°/11°E axis (Fig. 7a). This would identify the well-constrained northward motion of Pangea during this time interval as largely a result of TPW (Marcano et al., 1999; Le Pichon et al., 2021). This is surprising, given that the LLSVPs are thought to dominate the present-day moment of inertia of the Earth (Steinberger & Torsvik, 2010) and may have remained stable and antipodal throughout the entire Phanerozoic (Torsvik et al., 2014). If correct, this TPW rotation would imply that the contribution of other density anomalies than the LLSVPs, likely high-latitude slabs, may contribute sufficiently to the total moment of inertia to allow them to rotate ~20° relative to the spin axis towards the equatorial plane. This could be explained by either a different mantle viscosity and/or density structure during the (Late) Paleozoic, e.g., driven by a large volume of subducting slabs in the upper mantle at high latitude, driving the supercontinent and its surrounding subduction zones towards the equator. Alternatively, it may be explained by LLSVPs that were smaller, less dense, or mobile prior to ~260 Ma. We note, however, that the estimates of pre-260 Ma TPW are strongly influenced by the choices in either reference frame computation (the 'optimized' mantle reference frame of Müller et al. (2022) that keeps Pangea nearly fixed to the ambient mantle to reduce continental plate motion) or TPW computation (keeping the TPW axis fixed at 0°/11°E, Torsvik et al., 2014).

Finally, the computation of TPW requires knowledge of absolute plate motion in the paleomagnetic and mantle reference frames. The former requires that the time-averaged geomagnetic field aligns with the Earth's spin axis and that paleomagnetic data provide an accurate approximation of the mean pole position. Recent analysis has shown that even higher-resolution and more precise APWPs may be computed when uncertainty is propagated from site-level paleomagnetic data using a bottom-up approach (Gallo et al., 2023). Moreover, we emphasize that to establish whether observed TPW is statistically significant, and to assess the robustness of TPW rates and its implications for mantle dynamics, it is key to improve the quantification and incorporation of uncertainties in mantle reference frames. Defining a robust mantle reference frame with uncertainty quantification remains a key challenge for solid Earth science. All our attempts to define a mantle reference frame will inevitably remain approximations, since we are treating a convecting mantle as a fixed frame of reference. However, in a mantle frame based on simple 'tectonic rules', TPW becomes an intrinsic property of the mantle reference frame. This way, numerical models of global mantle convection driven by plate tectonics in an assumed mantle reference frame may then be calibrated against TPW, and iterated, along with hotspot tracks in an iterative modeling approach as in e.g., Doubrovine et al. (2012). This may not only provide novel constraints on mantle dynamics research but also allow the identification of the dynamic drivers of TPW. Lastly, we foresee that improved observational constraints on the timing, rate, and magnitude of TPW will contribute to exploring the exciting links between TPW and Earth’s climate, hydrosphere, geodynamo, and biosphere in the geological past.

5. Conclusions

Here, we present new quantitative estimates of true polar wander since 320 Ma. We find that TPW since 320 Ma occurred as large (>10°) but slow rotations at rates typically below 0.5°/Ma, with a mean TPW rate of 0.34°±0.14°/Ma (1σ). The TPW paths computed using four different mantle reference frames all show significant (>5°) TPW since 100 Ma, in contrast to some recent studies. The TPW path back to 320 Ma supports the existence of multiple large TPW oscillations of up to ~20° since the Permian. We find no evidence for phases of fast (>1°/Ma) TPW on timescales of >5 Ma, suggesting that previously observed, and heavily debated, rapid TPW may be an artifact. Our results show that TPW since 320 Ma predominantly occurred around two nearly orthogonal equatorial axes. We confirm a previously constrained oscillatory TPW rotation of ~24° in the Triassic and Jurassic around an axis that is close to the present-day center of mass of the antipodal LLSVPs in the lowermost mantle. In contrast, TPW is shown to have occurred about an axis at ~85°E during the last ~80 Ma and between 260 and 320 Ma, implying that other contributions to the moment of inertia, such as those exerted by subducting lithospheric slabs, controlled the nature of TPW. We tentatively explain the changes in the dominant axis of TPW since 320 Ma to changes in the contribution to the moment of inertia of low-latitude Tethyan subduction versus higher-latitude subduction in the Pacific realm. The >20° TPW oscillation around an axis close to the center of LLSVPs may confirm the stabilizing effects of LLSVPs on Earth's moment of inertia. Finally, we highlight that the coupling of kinematic constraints on TPW with absolute plate motion models provides an opportunity for calibration of numerical experiments of mantle dynamics.

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Conflict of Interest

The authors have no conflicts of interest to declare.

Data Availability Statement

The Jupyter Notebook used for the analyses in this study are publicly available on GitHub (https://github.com/bramvaes/TPW) and will be archived on Zenodo upon acceptance of this manuscript. We acknowledge the use of the freely available paleomagnetic software package PmagPy (Tauxe et al., 2016) in the Python codes used to perform the calculations. The complete paleomagnetic database that underpins the global APWP of Vaes et al. (2023) is available as Supplementary Datafile to the online version of that publication, as well as on the Reference database portal on APWP-online.org (Vaes et al., 2024).

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Supporting Information for

Slow true polar wander around varying equatorial axes since 320 Ma

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Introduction

The supplementary material provided in this file consists of two additional figures and five additional tables. Figure S1 shows the construction of the global APWP of Vaes et al. (2023) interpolated to exact 10 Ma and 5 Ma steps. Figure S2 shows the apparent polar rate of these interpolated APWPs and how to compare to the rates of the published global APWP of Vaes et al. (2023). Table S1 provides the Euler rotation parameters for the four mantle reference frames. The interpolated global APWP for the last 320 Ma is tabulated in South African coordinates in Table S2. Table S3 lists the polar wander paths constructed from each of the four mantle reference frames, which are shown in Figure 1. Table S4 provides the true polar wander paths computed at a 10-Ma-resolution that are plotted in Figure 2. Finally, Table S5 gives lists the interpolated global APWP and true polar wander path computed at a 5-Ma-resolution.

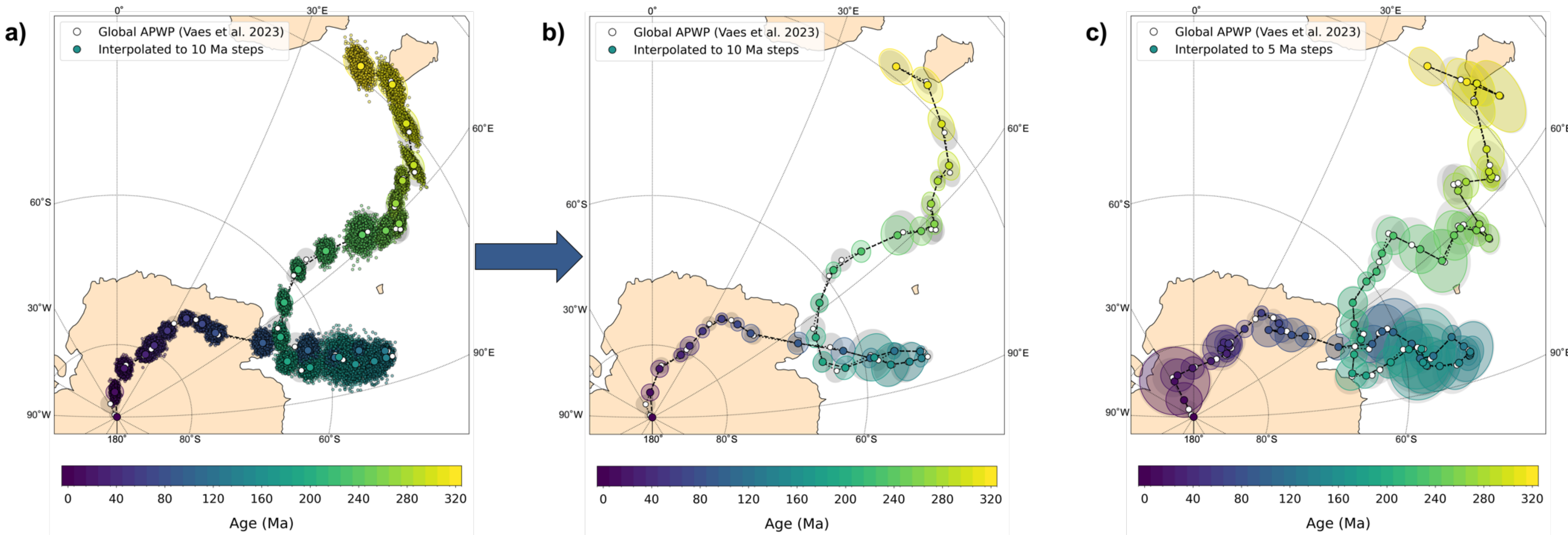

Figure S1. a) Orthographic plot showing the interpolated reference poles obtained by 1000 iterations. For each iteration, interpolated reference poles are determined at exact 10 Ma steps by interpolation along the great circle between successive reference poles. The initially computed reference poles have an 'effective' age computed by taking the mean age of the re-sampled virtual geomagnetic poles (VGPs) that fall within the 20 Ma window around the interpolation age. This 'effective' age therefore differs from the center age of the time window. See Vaes et al. (2023) for more details. **b)** The global APWP of Vaes et al. (2023) in South African coordinates (in white) compared to the interpolated path at 10 Ma steps used in this study. **c)** Same as b) but using a 5 Ma temporal resolution and a 10 Ma time window.

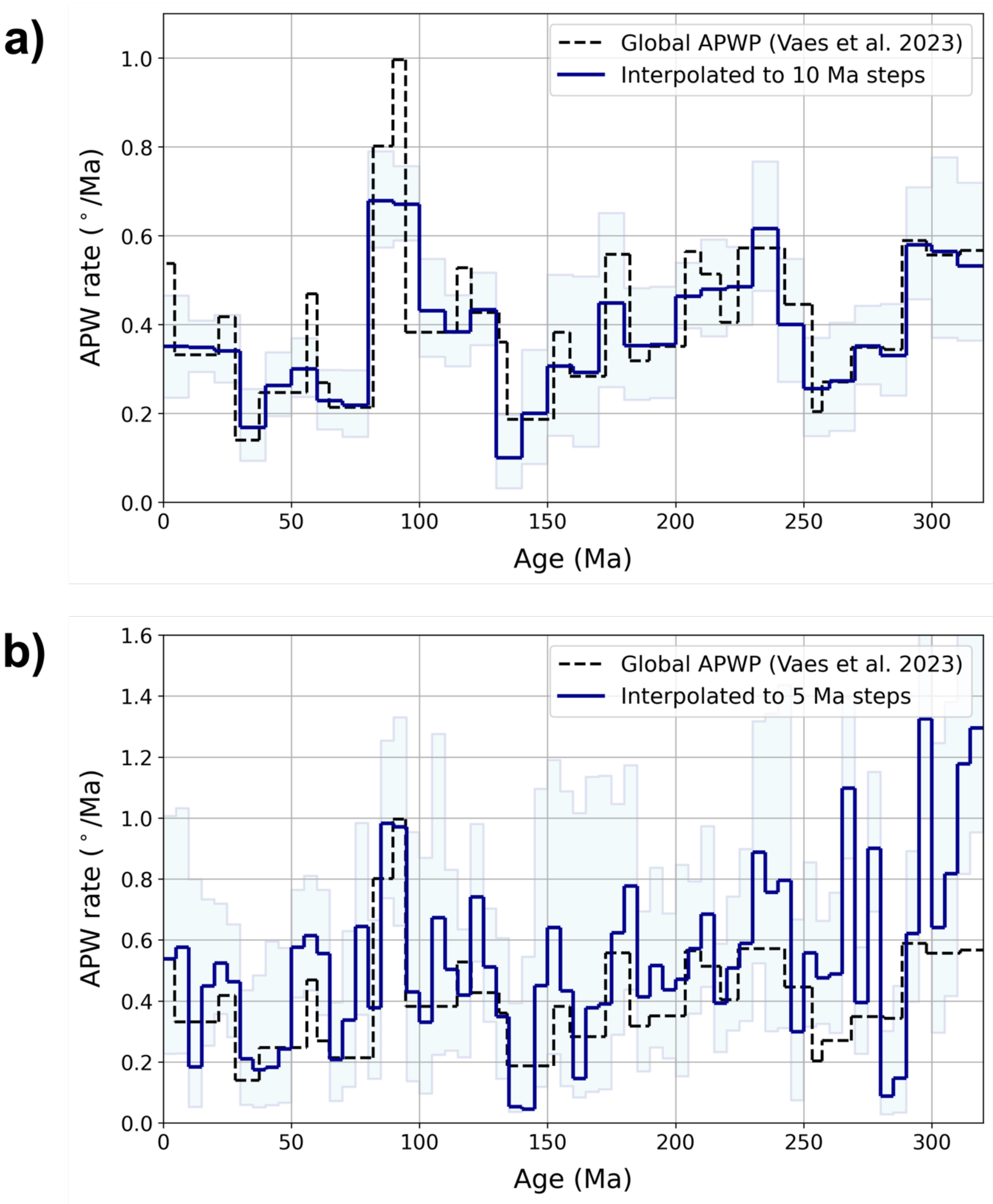

Figure S2. a) APW rates derived from the interpolated global APWP (in South African coordinates) per 10 Ma age interval (blue line) The light blue band represents the 95% confidence regions. The original APW rates obtained by Vaes et al. (2023) – corrected for the effective age – are shown by the dashed black line. **b)** Same as a) but for the 5-Ma-resolution APWP.

Müller et al. (1993)

Age*	Lat	Lon	Angle
0	0	0	0
10.9	59.3	-31.6	-1.89
20.1	50.9	-44.5	-4.36
33.1	40.3	-43.0	-7.91
40.1	37.7	-41.2	-9.65
47.9	32.8	-40.8	-12.09
55.9	30.1	-41.7	-13.89
67.7	26.4	-40.9	-16.23
73.0	22.3	-39.6	-17.80
79.1	18.0	-38.9	-19.98
83.5	19.0	-40.9	-21.53
89.9	19.4	-41.9	-23.31
100.5	18.9	-39.5	-25.35
111.1	17.7	-39.5	-26.71
120.4	18.7	-39.7	-27.37
130.7	16.7	-37.5	-28.52

Torsvik et al. (2008)

Age	Lat	Lon	Angle
0	0	0	0
5	30.3	326.0	-1.0
10	46.2	272.1	-1.9
15	45.8	276.8	-3.0
20	45.2	281.5	-4.0.
25	44.4	2860	-5.1
30	43.5	290.3	-6.1
35	44.3	297.9	-7.1
40	44.6	305.7	-8.1
45	40.8	303.3	-9.2
50	37.0	301.1	-10.3
55	30.6	310.1	-11.4
60	23.7	317.9	-12.5
65	22.2	319.4	-13.2
70	20.7	320.9	-13.8
75	19.2	322.4	-14.4
80	17.7	323.9	-15.0
85	16.2	325.3	-15.6
90	14.6	326.7	-16.2
95	14.5	328.6	-18.2
100	14.4	330.4	-20.1
105	15.1	331.1	-22.5
110	15.7	331.7	-24.9
115	16.4	332.4	-27.3
120	17.0	333.0	-29.7
125	17.5	333.5	-32.1
130	18.0	334.0	-34.6

Doubrovine et al. (2012)

Age	Lat	Lon	Angle
0	0	0	0
10	-36.66	146.72	1.61
20	-35.58	156.74	3.52
30	-36.19	153.99	5.85
40	-37.63	147.30	9.09
50	-42.80	147.61	9.83
60	-56.84	149.31	11.47
70	-51.40	149.04	15.75
80	-30.59	156.20	19.06
90	-22.76	156.10	21.48
100	-16.58	156.03	24.22
110	-17.11	155.90	26.11
120	-17.57	155.79	28.01

Müller et al. (2022)

Age	Lat	Lon	Angle
0	0	0	0
5	90.00	0.00	0.00
10	69.69	-51.67	-1.31
15	68.18	-50.15	-2.54
20	68.94	-48.57	-3.65
25	69.64	-45.38	-4.74
30	63.92	-42.45	-5.97
35	59.78	-40.73	-6.97
40	55.01	-37.80	-7.89
45	51.56	-34.44	-9.04
50	49.64	-33.01	-10.25
55	48.67	-33.70	-11.78
60	48.17	-35.10	-12.94
65	46.56	-36.43	-13.41
70	44.77	-37.64	-13.72
75	42.55	-37.84	-14.67
80	40.07	-37.18	-16.47
85	37.78	-38.14	-18.70
90	35.16	-39.60	-21.37
95	32.25	-39.04	-24.34
100	29.83	-38.33	-26.80
105	28.90	-37.59	-29.11
110	27.33	-35.89	-31.12
115	25.86	-34.34	-33.12
120	24.25	-33.33	-35.60
125	22.21	-32.16	-38.07
130	20.85	-32.21	-39.55
135	19.28	-32.26	-41.09
140	18.35	-32.19	-42.55
145	18.64	-31.31	-43.27
150	19.46	-30.87	-43.75
155	20.55	-30.68	-43.91
160	20.05	-30.90	-44.28
165	19.22	-30.60	-44.91
170	18.50	-30.13	-45.90
175	17.65	-29.93	-46.46
180	17.43	-29.19	-46.46
185	17.09	-28.35	-46.36

190	17.04	-28.07	-46.51
195	17.02	-27.49	-46.73
200	16.67	-27.55	-47.06
205	16.33	-27.63	-47.41
210	16.00	-29.05	-47.17
215	15.65	-30.57	-46.96
220	15.24	-32.45	-47.05
225	14.77	-34.41	-47.23
230	14.47	-35.53	-47.35
235	14.13	-36.74	-47.51
240	13.15	-36.69	-47.22
245	12.06	-36.66	-46.90
250	11.00	-36.57	-46.68
255	9.86	-36.47	-46.42
260	10.33	-36.43	-46.53
265	10.83	-36.33	-46.61
270	10.83	-36.31	-46.58
275	10.83	-36.28	-46.51
280	10.65	-36.32	-46.49
285	10.32	-36.43	-46.51
290	9.83	-36.33	-46.63
295	9.29	-36.23	-46.79
300	9.08	-36.24	-46.68
305	8.82	-36.25	-46.57
310	8.59	-36.26	-46.50
315	8.30	-36.33	-46.40
320	8.00	-36.39	-46.42

Table S1. Euler rotation parameters of the mantle reference frames used in this study, representing the total reconstruction poles of South Africa relative to the ambient mantle. Ages in million years ago (Ma). Lat/lon = latitude/longitude (in degrees). Angle in degrees; a positive angle indicates a counterclockwise rotation. * The ages of the Indo-Atlantic fixed hotspot frame of Müller et al. (1993) were updated by Torsvik et al. (2008).

Age	P_{95}	Longitude	Latitude	APW rate	APW rate (low)	APW rate (high)
0	0.00	0.0	-90.0			
10	1.18	355.15	-86.49	0.35	0.24	0.47
20	1.17	8.04	-83.18	0.35	0.29	0.41
30	0.92	23.84	-80.58	0.34	0.27	0.42
40	1.27	26.93	-78.99	0.17	0.09	0.26
50	0.91	29.89	-76.44	0.26	0.19	0.34
60	0.80	35.08	-73.75	0.30	0.24	0.37
70	0.91	42.69	-73.05	0.23	0.16	0.30
80	1.21	50.00	-72.67	0.22	0.15	0.30
90	1.50	65.09	-68.10	0.68	0.57	0.79
100	1.79	74.39	-62.61	0.67	0.59	0.76
110	1.33	79.35	-59.04	0.43	0.33	0.55
120	1.11	79.97	-55.22	0.38	0.31	0.47
130	0.77	82.41	-51.14	0.43	0.35	0.52
140	1.21	84.01	-51.09	0.10	0.03	0.20
150	2.61	83.81	-53.08	0.20	0.09	0.34
160	3.00	82.93	-56.11	0.31	0.15	0.51
170	3.23	79.62	-58.42	0.29	0.12	0.51
180	1.62	79.54	-62.91	0.45	0.26	0.65
190	1.40	75.06	-65.85	0.35	0.23	0.48
200	1.40	66.45	-65.56	0.36	0.23	0.49
210	1.27	57.95	-62.76	0.46	0.39	0.54
220	1.08	53.54	-58.48	0.48	0.37	0.59
230	1.45	54.97	-53.70	0.48	0.40	0.58
240	2.63	57.95	-47.83	0.62	0.48	0.77
250	1.90	60.64	-44.30	0.40	0.27	0.55
260	1.38	61.41	-41.80	0.26	0.15	0.37
270	1.58	58.24	-40.47	0.27	0.16	0.40
280	1.29	56.36	-37.27	0.35	0.26	0.44
290	1.99	56.13	-33.97	0.33	0.24	0.45
300	2.38	50.95	-30.18	0.58	0.46	0.71
310	2.67	45.75	-26.88	0.56	0.37	0.78
320	2.71	39.94	-28.21	0.53	0.36	0.72

Table S2. Interpolated global APWP for the last 320 Ma in South African coordinates, calculated using a 20 Ma window. The P_{95} represents the 95% confidence region of the reference pole (in degrees). APW rates (with their 95% confidence limits) are given in degrees/Ma. Note that the APW rate provided for e.g., 30 Ma is the rate determined for the interval between 30 and 20 Ma.

	M93		T08		D12		M22	
Age	**Plon**	**Plat**	**Plon**	**Plat**	**Plon**	**Plat**	**Plon**	**Plat**
0	0	-90	0	-90	0	-90	0	-90
10	59.15	-89.11	2.79	-88.68	57.20	-88.71	41.04	-89.05
20	47.23	-87.27	12.92	-87.18	67.76	-87.14	46.84	-88.35
30	49.17	-84.75	22.40	-85.58	65.72	-85.28	52.29	-86.49
40	51.73	-82.39	38.55	-84.23	60.08	-82.80	59.10	-84.38
50	52.27	-79.36	34.20	-81.78	60.95	-82.79	60.72	-82.23
60	52.15	-77.12	50.42	-78.56	64.12	-83.73	58.45	-80.79
70	53.23	-74.64	53.35	-77.09	65.21	-80.19	57.13	-79.20
80	53.87	-70.73	56.19	-75.71	71.08	-73.61	57.62	-75.25
90	52.05	-68.01	58.75	-74.33	70.30	-70.21	57.52	-69.46
100	54.55	-66.13	62.92	-70.54	69.53	-66.80	59.57	-64.58
110	54.63	-64.73	65.12	-66.04	69.80	-65.06	63.05	-60.28
120	54.76	-64.12	67.43	-61.63	70.10	-63.32	65.27	-54.85
130							64.79	-51.31
140							65.92	-49.10
150							67.37	-49.02
160							67.14	-47.71
170							67.48	-45.85
180							68.82	-45.80
190							69.72	-45.43
200							69.41	-44.61
210							66.12	-44.88
220							61.95	-44.42
230							59.39	-44.01
240							58.52	-44.19
250							57.73	-44.30
260							58.30	-44.26
270							58.33	-44.36
280							57.98	-44.28
290							57.77	-43.86
300							57.53	-44.01
310							57.22	-44.11
320							56.78	-44.01

Table S3. Polar wander paths constructed from the mantle reference frames in Table S1, plotted in Fig. 1. Plon/plat = pole longitude and latitude of the spin axis relative to a fixed South Africa plate in the absence of TPW. M93= Müller et al. (1993); T08 = Torsvik et al. (2008); D12 = Doubrovine et al. (2012); M22 = Müller et al. (2022).

		M93		T08		D12		M22	
Age	**P_{95}**	**Plon**	**Plat**	**Plon**	**Plat**	**Plon**	**Plat**	**Plon**	**Plat**
0	0.00	0.0	90.0	0.0	90.0	0.0	90.0	0.0	90.0
10	1.18	159.4	86.8	169.2	87.8	152.8	86.9	159.4	87.1
20	1.17	164.6	85.0	-178.2	86.0	161.4	84.1	173.1	84.4
30	0.92	173.6	84.8	-159.1	85.0	172.5	83.3	-176.8	83.5
40	1.27	163.4	84.8	-170.9	84.5	163.4	83.7	174.5	83.1
50	0.91	156.2	84.5	-162.8	84.6	176.8	81.7	171.4	82.1
60	0.80	164.7	84.6	-179.6	84.0	-170.0	78.8	-179.3	81.4
70	0.91	159.4	86.7	-170.9	85.1	-174.6	81.3	-169.3	83.0
80	1.21	77.7	87.7	-160.4	86.5	149.8	83.8	-176.9	86.7
90	1.50	-38.7	85.1	-104.2	83.5	-161.9	87.2	-68.4	86.9
100	1.79	-56.9	80.8	-86.9	80.9	-88.9	85.3	-54.6	83.1
110	1.33	-58.4	77.1	-72.1	80.4	-77.1	82.5	-43.2	81.7
120	1.11	-67.7	74.6	-70.2	80.9	-82.2	80.5	-31.0	81.6
130	0.77							-33.3	79.0
140	1.21							-21.9	78.3
150	2.61							-11.0	78.9
160	3.00							8.5	77.2
170	3.23							26.7	75.4
180	1.62							38.7	71.8
190	1.40							49.1	69.4
200	1.40							58.8	69.0
210	1.27							64.6	71.5
220	1.08							66.4	75.0
230	1.45							62.2	79.9
240	2.63							54.2	86.3
250	1.90							-39.8	87.9
260	1.38							-87.2	86.6
270	1.58							-131.9	86.1
280	1.29							-141.3	82.9
290	1.99							-138.1	80.0
300	2.38							-152.9	75.2
310	2.67							-161.9	70.4
320	2.71							-175.8	69.2

Table S4. True polar wander paths computed by placing the interpolated global APWP into the mantle reference frames in Table S1, as plotted in Fig. 2a-d and Fig. 4a. The P_{95} represents the paleomagnetic uncertainty only and is thus a minimum confidence region. See captions of the tables above for the meaning of the abbreviations.

		Interpolated APWP at 5 Ma steps					TPWP	
Age	P_{95}	Plon	Plat	APW rate	APW rate (low)	APW rate (high)	Plon	Plat
0	0.00	0.0	-90.0	0.54	0.23	1.01	0.0	90.0
5	2.42	331.46	-87.31	0.58	0.23	1.03	140.73	87.45
10	4.79	332.79	-84.43	0.18	0.05	0.80	140.91	84.71
15	2.26	339.52	-83.82	0.45	0.25	0.73	144.29	84.28
20	1.72	358.76	-83.21	0.52	0.36	0.72	162.21	84.18
25	1.16	16.61	-81.93	0.46	0.28	0.65	177.46	83.94
30	1.09	26.85	-80.24	0.21	0.06	0.44	-171.94	83.23
35	1.48	26.00	-79.20	0.17	0.05	0.48	-178.24	82.79
40	2.36	21.49	-79.01	0.18	0.06	0.59	166.99	82.62
45	2.24	21.40	-79.92	0.24	0.07	0.57	153.50	83.50
50	1.89	24.89	-78.89	0.58	0.41	0.76	152.94	83.40
55	1.15	29.69	-76.20	0.61	0.43	0.81	166.22	82.35
60	0.87	33.02	-73.25	0.56	0.35	0.76	178.65	80.71
65	1.14	42.55	-72.99	0.21	0.07	0.39	-164.71	82.02
70	1.50	45.63	-73.53	0.34	0.12	0.63	-164.06	83.74
75	2.19	40.71	-74.55	0.64	0.35	0.98	163.01	84.93
80	2.23	50.61	-72.90	0.38	0.14	0.67	-176.65	86.97
85	1.39	55.05	-71.58	0.98	0.74	1.25	-162.33	88.91
90	1.76	66.25	-68.48	0.97	0.65	1.33	-58.94	86.72
95	2.55	70.87	-63.99	0.43	0.14	0.95	-67.82	84.12
100	3.64	68.87	-62.05	0.33	0.10	0.87	-71.69	85.12
105	4.60	70.61	-60.62	0.67	0.22	1.28	-61.47	85.26
110	2.48	75.86	-58.55	0.50	0.24	0.83	-51.01	83.27
115	1.52	80.42	-57.80	0.42	0.21	0.63	-32.82	81.35
120	1.32	81.03	-55.74	0.74	0.53	0.98	-27.48	81.00
125	1.63	78.49	-52.35	0.51	0.33	0.70	-39.71	81.86
130	0.93	81.62	-50.69	0.35	0.14	0.61	-36.08	79.41
135	1.47	84.05	-49.89	0.05	0.04	0.39	-30.97	77.61
140	2.30	83.69	-49.75	0.05	0.04	0.55	-28.56	78.45
145	3.80	83.34	-49.75	0.45	0.12	1.10	-27.54	79.17
150	3.78	84.82	-51.80	0.64	0.23	1.19	-18.02	78.57
155	3.60	83.99	-54.97	0.43	0.12	1.09	-0.84	77.74
160	3.97	82.75	-57.02	0.15	0.08	1.02	11.96	76.74
165	5.01	81.68	-57.46	0.38	0.10	1.13	18.94	76.13
170	5.27	78.18	-57.32	0.39	0.11	1.14	26.66	76.78
175	3.84	76.94	-59.16	0.62	0.29	1.05	35.06	75.62
180	1.96	78.48	-62.19	0.78	0.44	1.17	39.22	72.70

185	1.82	79.93	-66.02	0.41	0.17	0.69	42.57	68.81
190	2.40	77.58	-67.87	0.52	0.25	0.84	47.66	67.19
195	1.98	71.38	-66.88	0.44	0.21	0.69	53.42	68.13
200	1.51	67.85	-65.23	0.47	0.15	0.85	57.19	69.36
205	2.06	62.25	-65.10	0.57	0.36	0.79	60.77	69.44
210	2.13	56.48	-63.71	0.69	0.45	0.97	65.53	70.40
215	1.20	54.33	-60.43	0.39	0.23	0.59	67.85	73.23
220	1.41	54.21	-58.47	0.51	0.30	0.74	65.20	75.17
225	1.59	51.85	-56.27	0.59	0.33	0.90	70.13	76.72
230	2.75	50.91	-53.37	0.89	0.52	1.32	74.86	79.11
235	3.50	57.46	-51.42	0.76	0.30	1.34	55.09	82.60
240	4.58	62.58	-49.51	0.80	0.31	1.44	22.06	84.01
245	3.73	60.19	-45.86	0.30	0.07	0.67	7.79	87.81
250	2.38	60.00	-44.37	0.56	0.32	0.85	-37.30	88.37
255	2.07	62.46	-42.24	0.48	0.33	0.66	-71.18	86.15
260	1.85	65.17	-40.98	0.49	0.29	0.76	-71.57	83.98
265	1.80	61.96	-41.32	1.10	0.87	1.39	-87.77	86.00
270	2.36	54.66	-41.23	0.39	0.23	0.58	-172.89	85.86
275	2.09	54.71	-39.25	0.90	0.69	1.15	-159.23	84.27
280	1.64	57.68	-35.42	0.09	0.03	0.30	-132.42	81.13
285	1.49	57.56	-34.99	0.15	0.03	0.45	-132.18	80.90
290	2.15	56.67	-34.89	0.62	0.40	0.89	-135.96	80.99
295	3.35	53.93	-32.77	1.33	1.00	1.69	-145.96	78.46
300	3.95	47.92	-28.62	0.64	0.30	1.25	-159.70	72.80
305	3.24	46.78	-25.58	0.82	0.42	1.38	-158.45	69.62
310	5.06	50.84	-23.81	1.18	0.82	1.65	-146.36	69.04
315	3.40	45.23	-26.81	1.30	0.95	1.70	-162.55	70.30
320	2.92	38.49	-29.37	0.54	0.23	1.01	179.38	69.37

Table S5. Interpolated global APWP (in South African coordinates) at a 5-Ma-resolution (shown in Fig. S1c) and true polar wander path computed at a 5-Ma-resolution that was constructed using the mantle reference frame of Müller et al. (2022). See captions of tables above for the meaning of the abbreviations.