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Preprint

39 **Title:** Regional Source-type Discrimination Using Nonlinear Alignment Algorithms

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Preprint

50 **Abstract**

51 The discrimination problem in seismology aims to accurately classify different underground
52 source types based on local, regional and/or teleseismic observations of ground motion. Typical
53 discriminant approaches are rooted in fundamental, physics-based differences in radiation
54 pattern or wave excitation, which can be frequency dependent and may not make use of the full
55 waveform. In this paper, we explore whether phase and amplitude distances derived from
56 dynamic time warping (DTW) and elastic shape analysis (ESA) can inform event discrimination.
57 We demonstrate the ability to distinguish underground point-sources using synthetic waveforms
58 calculated for a 1-D Earth model and various source mechanisms. We then apply the method to
59 recorded data from events in the Korean Peninsula, which includes declared nuclear explosions,
60 a collapse event, and naturally occurring earthquakes. Phase and amplitude distances derived
61 from DTW and ESA are then used to classify the event types via dendrogram and k-nearest
62 neighbor clustering analyses. Using information from the full waveform, we show how different
63 underground sources can be distinguished at regional distances. We highlight the potential of
64 these nonlinear alignment algorithms for discrimination and comment on ways we can extend
65 the framework presented here.

67 **Introduction**

68 Source type discrimination is needed to classify events for nuclear treaty or seismic
69 hazard monitoring purposes. Traditional discriminants exploit physics-based intuition that the
70 radiation pattern of double-couple sources (e.g., earthquakes) should be fundamentally
71 different from explosion-like sources (e.g., chemical/nuclear tests, mining blasts). This is due to
72 the difference in the repartition of energy release that exists between shear slip and a pressure
73 pulse acting on the rock (Ben-Menahem and Singh, 1981). Approaches to discrimination
74 between seismic events include moment tensor inversion (e.g., Alvizuri and Tape, 2018;
75 Pasyanos and Chiang, 2022), body to surface-wave magnitude ratios ($M_s:m_b$, e.g., Stevens and
76 Day, 1985), spectral amplitude ratios (e.g., Tibi, 2021) and recently, machine learning methods
77 (e.g., Kong et al., 2022; Maguire et al., 2024; Linville et al., 2019). Many of these established
78 discrimination methods are successful because specific diagnostic parts of the seismic
79 waveform can be analyzed in the time or frequency domain using narrow passbands, or, in the
80 case of moment tensor inversion, the full moment tensor can be estimated, and the proportion
81 of double-couple and non-double couple components determined. Limitations to the above
82 methods may include difficulty to achieve satisfactory phase separation at local distances (e.g.,
83 Tibi et al., 2023), constraints on event size (i.e., total moment release) or in the case of machine
84 learning, analysts may not have enough large, labeled datasets to draw from for a given region
85 of interest.

86 Nonlinear dynamic programming techniques, broadly referred to as dynamic time
87 warping (DTW), have been used to determine what temporal shifts are necessary to optimally

88 align time series (Anderson and Gaby, 1983; Kumar et al., 2022; Müller, 2021). DTW has been
89 applied in seismic exploration (Hale, 2013), ambient-noise interferometry (Mikesell et al., 2015;
90 Yuan et al., 2021) and linear seismic inversion problems (Tan and Langston, 2022). DTW
91 algorithms capture phase variability quite well but may not be as robust to amplitude variability
92 (Müller et al., 2021). This is a concern for seismic event monitoring and especially in the context
93 of signal window selection, which can contain both body and surface waves with markedly
94 different amplitudes. A novel representation of functional data that addresses both phase and
95 amplitude variability within a time series is elastic shape analysis of curves (ESA; Tucker et al.,
96 2013). The ESA method aligns signals to one another after applying a square root slope function
97 (SRSF; Joshi et al., 2007; Srivastava et al., 2011). The SRSF is a distance-preserving
98 transformation between metric spaces (isometry) and yields a proper distance in either phase
99 or amplitude space. Our hypothesis is that two signals sharing the same source mechanism and
100 similar Green's functions will have a lower phase or amplitude distance whereas signals that do
101 not will have correspondingly higher distances. Moreover, nonlinear alignment methods are not
102 constrained by phase separation, narrow time windows, or event size.

103 This manuscript explores whether nonlinear time series alignment algorithms might
104 assist event type discrimination. First, we set up a simple synthetic test relevant to regional
105 distance monitoring to see if DTW and ESA can distinguish double couple from non-double
106 couple (e.g., explosive) signals. Because the synthetic test supports this hypothesis, we then
107 analyze a real seismic dataset of earthquake, explosion, and collapse events in the Democratic
108 People's Republic of Korea (DPRK) and apply hierarchical clustering and k-nearest neighbor (knn)
109 analysis to the DTW and ESA distances to see whether the method works to distinguish event
110 types using a real discrimination scenario. We compare the DTW and ESA methods, contrast our
111 approach to existing discriminants, and discuss recommendations to extend this preliminary
112 analysis.

113 **Methodology**

114 *Dynamic Time Warping*

115 DTW estimates time shifts between signals to estimate geophysical parameters and can
116 overcome strong cycle-skipping even in the presence of low signal to noise-ratio (SNR), which is
117 an advantage over windowed cross-correlation and linear trace stretching methods that may
118 estimate incorrect lag times (Mikesell et al., 2015). However, to optimally align signals,
119 unrealistically large dilation of the original time series can occur and strategies to constrain the
120 dynamic programming algorithm (i.e., global or local constraints on the permissible warping
121 function) to reasonable dilation values are not always easy to set a-priori. Despite these choices,
122 DTW allows one to calculate a non-Euclidian distance metric that gives a measure of how much
123 warping was needed for optimal alignment, the DTW distance (DTW_{dist}), defined below as,

124
$$DTW_{dist}(f_1, f_2) = \min\{w_p(f_1, f_2)\} \quad (1)$$

125 where w_p is the warping path that aligns $f_1(t)$ and $f_2(t)$ after an accumulated distance matrix is
 126 computed (Müller, 2021).

127 *Elastic Shape Analysis*

128 ESA separates amplitude and phase information uniquely by first computing the SRSF,
 129 $q(t)$,

130
$$q(t) = \text{sign}(f'(t))\sqrt{|f'(t)|} \quad (2)$$

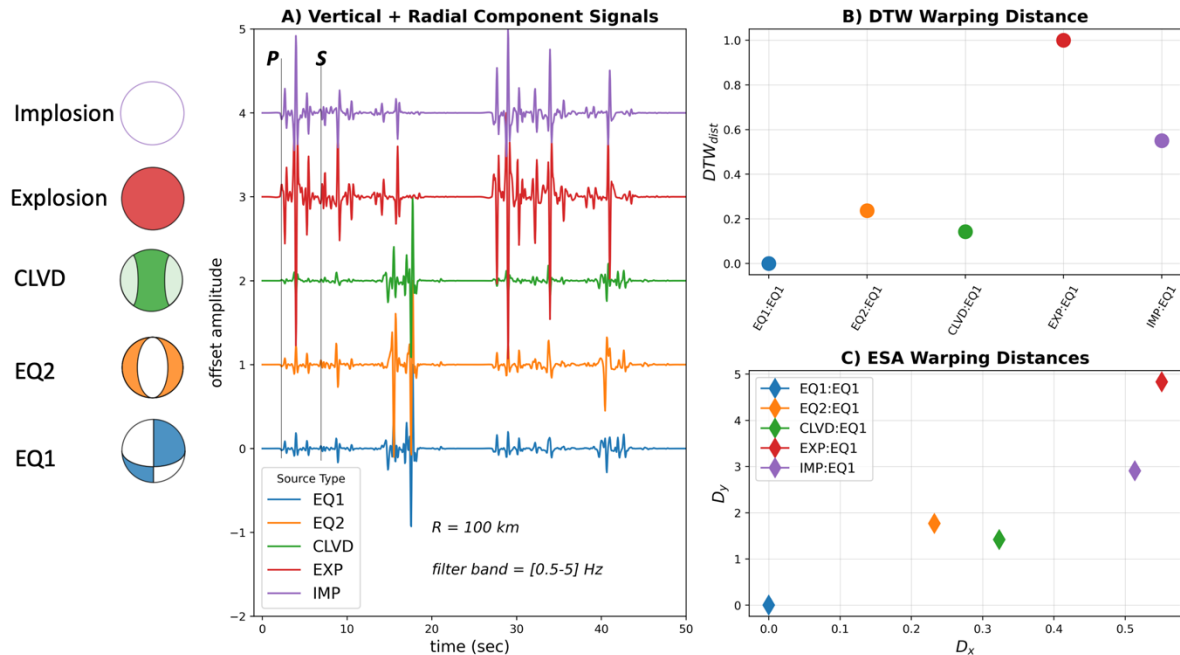
131 where $q(t)$ is the transformed signal and $f'(t)$ is the first derivative of the original signal with
 132 respect to time (Srivastava et al., 2011). The amplitude distance (D_y) between two functions $f_1(t)$
 133 and $f_2(t)$ is defined as,

134
$$D_y(f_1, f_2) = \inf_{\gamma \in \Gamma} \| q_1 - (q_2 \circ \gamma)\sqrt{\dot{\gamma}} \| \quad (3)$$

135
 136 where q_1 and q_2 are the SRSF of $f_1(t)$ and $f_2(t)$, respectively, and γ is the warping function that
 137 best aligns them. The double-bars " $\| \cdot \|$ " denote the L2-norm and Γ represents the complete
 138 set of invertible functions that map a smooth surface to each another such that both the
 139 function and its inverse are well defined over $[0,1]$. The phase distance (D_x) is then defined as
 140 the distance between warping functions according to,

141
 142
$$D_x(\gamma_1, \gamma_2) = d_\psi(\psi_1, \psi_2) \equiv \cos^{-1}(\int_0^1 \psi_1(t)\psi_2(t) dt) \quad (4)$$

143
 144 where $\psi(t)$ represents a mapping of the warping function to Hilbert space (\mathcal{H}) and D_x is thus the
 145 arc-length between the corresponding SRSF on a \mathcal{H} unit sphere. The theory behind elastic
 146 distances is rich, and we refer the interested reader to Wu and Srivastava (2011), Srivastava et
 147 al., (2011) or Tucker et al., (2013) for the in-depth, mathematical underpinnings of D_x and D_y .
 148 The important properties of ESA that distinguish this method from typical DTW algorithms are:
 149 1) D_y and D_x are independent of one another, 2) the SRSF transformation guarantees a
 150 mathematically proper distance, and 3) the distances are invariant to warping order.

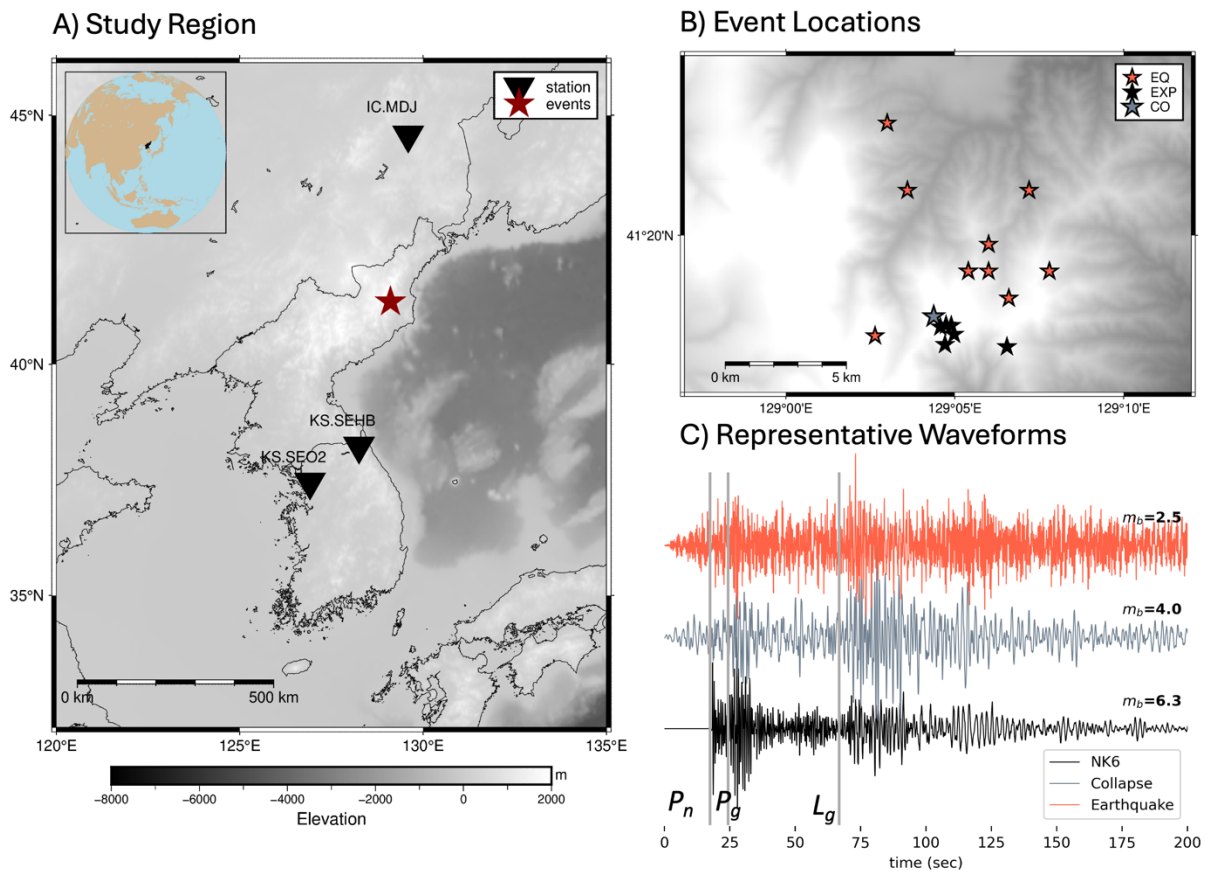


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 152 *Figure 1. A) Concatenated vertical (Z) and radial (R) component synthetics calculated for the Ford et al. (2009) velocity*
 153 *model recorded at a source-receiver distance of 100 km assuming an azimuth of 30 degrees and hypocenter depth of*
 154 *1 km. Source mechanisms range from purely double-couple (bottom) to implosive (top) sources. The waveforms are*
 155 *normalized to their respective maximum amplitude and distinct phase arrival times are denoted by thin gray lines on*
 156 *Z components. Direct P and S are labeled. B) Dynamic time warping distance (DTW_{dist}) between each waveform pair*
 157 *relative to EQ1, normalized by the largest DTW_{dist} . C) Elastic phase (D_x) and amplitude (D_y) distances for each warping*
 158 *pair (unnormalized). Legend in C applies to subfigures B and C.*
 159
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161 Synthetic Data Experiment

162 We generate near-regional synthetic seismic waveforms using a 1-D velocity model
 163 developed for the DPRK to conceptually represent a realistic monitoring scenario where closely
 164 spaced and different underground sources are recorded by a single seismic station (Ford et al.,
 165 2009; Figure 1). We use a wave-number integration algorithm (Herrmann, 2013) to calculate
 166 synthetics for double-couple, compensated linear vector dipole (CLVD), pure explosion, and
 167 implosion point-sources recorded at 100 km distance (Figure 1A). We calculate only the down-
 168 going Green's function components to suppress strong free-surface effects on the waveform.
 169 This lets us simplify and focus our analysis on waveform differences due to source mechanism
 170 alone. We filter the 40-samples-per-second synthetic waveforms between 0.5 to 5 Hz (Tibi,
 171 2021) and normalize each trace by its respective maximum amplitude prior to alignment. We
 172 also concatenate the vertical (Z) and radial (R) components into a composite time series before
 173 warping and alignment to mimic practice in signal detection. Note that for a pure isotropic
 174 explosion source, there is no tangential (T) motion generated and thus that component is not
 175 considered in the synthetic analysis because discrimination would be trivial.

176 At 100-km source-receiver distance, we select a double-couple earthquake waveform as
 177 the main signal to align to (EQ1, Figure 1A). Each waveform has a P-wave arrival near 15 seconds
 178 after the respective event origin time, but to highlight phase arrivals of interest, we cut the
 179 waveforms from 15 to 40 seconds for both R and Z components. We warp every signal to EQ1
 180 and calculate DTW_{dist} , D_x and D_y (Figure 1B, C). We observe that for both DTW and ESA, the
 181 distance between EQ1 and itself is zero (expected) and the explosion and implosion events have
 182 greater phase and amplitude distances than alignment to the earthquake or CLVD events.
 183 Because the Green's function is the same for each waveform, this synthetic experiment suggests
 184 that differences due to source mechanism can be inferred via phase or amplitude distance
 185 information from DTW and ESA between the full waveforms. It is important to keep in mind that
 186 the source mechanism and velocity model are kept simple to illustrate how warping distances
 187 can distinguish underground event types.
 188



189
 190 *Figure 2. A) Map of DPRK and surrounding region showing seismic stations used in the cluster analysis. Average*
 191 *epicenter of the NK1 - NK6 tests, collapse event, and naturally occurring earthquakes is denoted by the red star. B)*
 192 *Event epicenters near Mount Mantap. Several earthquakes (EQ) are < 1 km distance apart and thus may be plotted*
 193 *on top of one another. C) Example vertical-component waveforms of an explosion, collapse and earthquake event*
 194 *recorded at station MDJ (network IC). National Earthquake Information Center body-wave magnitudes (m_b) tabulated*

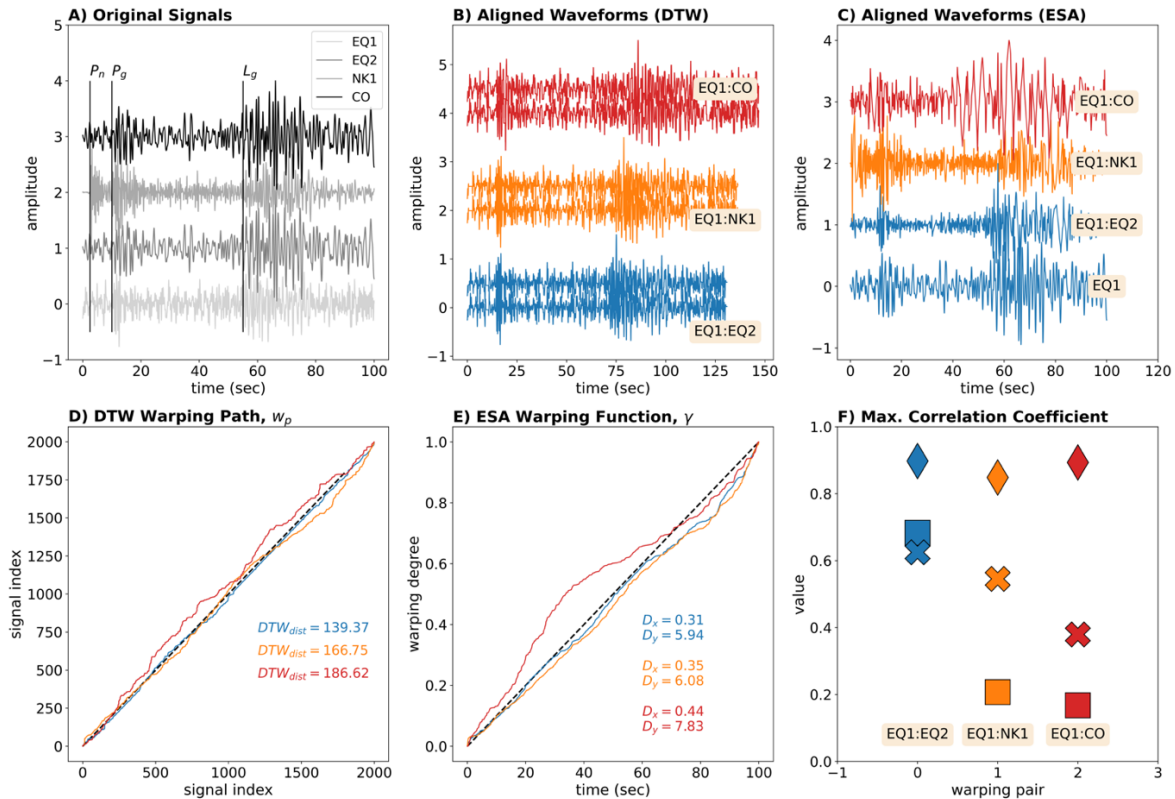
195 on the righthand side of the plot. Filter passband is between 0.5 – 5 Hz. A group velocity of 3.6 km/s is assumed to
196 estimate the L_g arrival whereas the PREM earth model is used to calculate P_n and P_g arrival times.
197

198 **Real Data Analysis: DPRK**

199 For the observational dataset, we use waveforms analyzed in Tibi (2021). This dataset
200 contains six declared explosions, fifteen nearby naturally occurring earthquakes and one
201 collapse event following the 2017 declared North Korean nuclear test (Figure 2). All events are
202 within a 10-km epicentral distance of one another. Three regional broadband stations are
203 selected from the IC, IU, and KS seismic networks. We download waveforms up to 15 minutes
204 after the respective origin time of the events from the Incorporated Research Institutions for
205 Seismology (IRIS) database to ensure P_n , P_g , and L_g phases are captured, and we filter signals
206 below 10 Hz. Due to different station start and end times, not all events are recorded; also,
207 stations with non-emergent phases are excluded from analysis.

208 Regional seismic stations recorded the declared 2006 (NK1), 2009 (NK2), 2013 (NK3),
209 2016 (NK4, NK5) and 2017 (NK6) DPRK nuclear tests, one collapse event following NK6, and
210 explosion-induced aftershocks or isolated, natural seismicity (Figure 2B). We use a subset of the
211 stations used in Tibi (2021) to assess DTW and ESA performance on each event pair combination
212 in this set of closely spaced events (Figure 2B). For any station, the signal time window spans
213 five seconds before theoretically expected P_n and ~ 40 seconds after the L_g arrival (Figure 2C).

214 We select data from a single station (MDJ, vertical component) to demonstrate how
215 nonlinear alignment between dissimilar signals may result in larger amplitude (D_y) or phase (D_x ,
216 DTW_{dist}) distance. We select two earthquakes (EQ1 and EQ2; Figure 3A), an explosion from the
217 2006 declared nuclear test (NK1; Figure 3A) and the collapse event following the largest
218 declared nuclear test for this exercise (CO; Figure 3A). P_n , P_g and L_g phases are readily
219 identifiable on all waveforms. Because the explosions have different yields and are not exactly
220 co-located (Myers et al., 2018), nonlinear warping must address differences from both the
221 source mechanism and Green's functions. Selecting EQ1 has the main signal to align to, we
222 show how the waveforms must be warped to accomplish this using DTW (Figure 3B) and ESA
223 (Figure 3C) approaches. The warping functions (and the phase or amplitude distances that are
224 computed after alignment) are given in Figure 3D and 3E. We note that in all cases, if two
225 dissimilar event type waveforms are aligned to one another (e.g., an earthquake to an
226 explosion), then a *larger* D_y , D_x and DTW_{dist} is indeed observed and the warping function
227 deviates significantly from the one-to-one diagonal line, which would be the warping path
228 between two identical signals (Figure 3D, 3E). We also calculate the cross-correlation coefficient
229 (CC) before and after alignment and note that whereas DTW and ESA both increase CC (an
230 exception being between EQ1 and EQ2 using ESA, but the differences is < 0.1), DTW (with no
231 constraints on the warping path) can increase CC by as much as 0.7 units. Such stellar
232 alignment, however, comes at the cost of appreciable waveform stretching (i.e., Figure 3B).
233



234
 235 *Figure 3. Demonstration of how nonlinear warping algorithms may be able to distinguish underground event types.*
 236 *A) Vertical component waveforms recorded at station MDJ filtered between 0.5 and 3 Hz. EQ1 and EQ2 are m_b 2.5*
 237 *and 3.4, respectively. NK1 = declared nuclear explosion test on 2006/10/09 (m_b 4.3) and CO = collapse event following*
 238 *declared nuclear explosion NK6 on 2017/09/03 (m_b 4.0) B) Waveforms of each event warped to match EQ1 using*
 239 *dynamic time warping (DTW). C) Waveforms of each event warped to match EQ1 using elastic shape analysis (ESA).*
 240 *D) The warping paths between each signal pair in panel B. The dynamic time warping distances (DTW_{dist}) are*
 241 *tabulated in the lower right corner. E) The warping function through time for each signal pair in C. Here, both phase*
 242 *(D_x) and amplitude (D_y) distances are given in the lower right corner. For both D) and E) the diagonal dashed line*
 243 *signifies what the warping path would be if no distortion between signals was needed for alignment. F) The cross-*
 244 *correlation coefficient before (squares), after DTW (diamonds) and after ESA (crosses) alignment.*

245
 246 We next compute DTW and ESA for each pair across all stations and apply hierarchical
 247 clustering analysis to the DTW and ESA distances obtained. We show the analysis for MDJ in
 248 Figure 4. We also report CC between each pair to see where a particular distance metric may
 249 align with empirical signal similarity (Figure 4D). We calculate condensed matrix representations
 250 of D_x , D_y , and DTW_{dist} and show the results graphically using a dendrogram (Figure 4E, F). We
 251 observe more structure in the D_y and DTW_{dist} matrices, based off a qualitative comparison to the
 252 cross-correlation matrix. When an earthquake is warped to an explosion (or vice versa), we
 253 generally observe a higher phase or amplitude distance (Figure 4A, B) and a lower CC score
 254 (Figure 4C). In contrast, the collapse event is not as easy to discern without additional
 255 information.

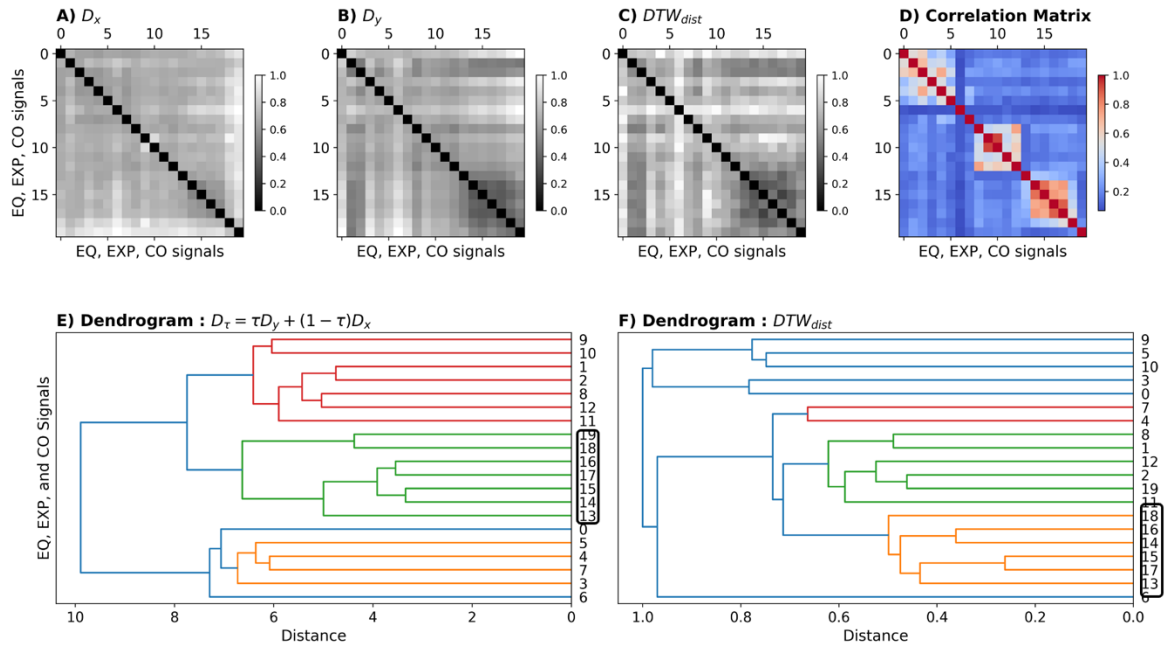
256 To assess whether we can achieve better classification between earthquake, explosion
257 and collapse events using both phase and amplitude information, we form a simple linear
258 combination of D_x and D_y following Tucker et al., (2014), using a weighting coefficient, τ . Since D_y
259 appears to have more structure than D_x , we weight D_y more in the below formulation,

$$D_\tau = \tau D_y + (1 - D_x)\tau \quad (5)$$

262
263 To optimize the weighting coefficient τ , we randomly set aside 50% of the signals as
264 training data (Tucker et al., 2014) and employ a Leave-One-Out (LOO) cross-validated knn
265 classifier for varying τ levels ($0 \leq \tau \leq 1$ in an increment of 0.1) in expression (5). We set the
266 value of knn to three because we have exactly three signal types to cluster. Our metric for
267 classification accuracy is the percentage of true predictions returned by knn, based off the signal
268 type labels we assign. We found that τ values between 0.1 – 0.9 classify signal types to the
269 ~70% accuracy level, and we do not observe appreciable changes between τ values to the
270 hundredths decimal point. We only have twenty labeled signals to work with at MDJ, so using a
271 larger signal database (for a given station) could give us more robust statistics. Using either
272 nonlinear alignment methods suggests that the explosion waveforms are separated from
273 earthquake waveforms, but the collapse event may group with either the earthquake (DTW) or
274 explosion (ESA) population. The results shown for MDJ do not dramatically change if the signal
275 envelope is used, or if the R and T component waveforms are concatenated to Z.

276 We report which ground motion component, signal type, window length and frequency
277 passband that best distinguishes explosion from non-explosion signals for our full set of regional
278 stations in Table 1. The criterion for choosing a particular passband is an improvement in
279 classification accuracy. We list how the LOO cross validation with knn performed for D_x , D_y and
280 DTW_{dist} separately to see if any one distance is superior to another (Table 1). For stations less
281 than 300 km from the source, knn classification using D_y performs comparably to DTW_{dist} and
282 slightly better than D_x . Stations at greater distances (>300 km) from the source do not record
283 waveforms with qualities that are sufficient for this type of analysis (i.e., SNR > 3). This may
284 impact classification accuracy as it is essentially the same between all distance metrics at SEO2
285 (Table 1). Note also that the relative time window accordingly widens to accommodate larger P
286 to L_g separation time. For the τ analysis at station MDJ, we attempted to find the best linear
287 combination of D_x and D_y that improved classification accuracy, but ultimately found that D_y by
288 itself is superior, which may not be the case for every dataset. Moreover, there is no reason to
289 stick to a linear relationship between the phase and amplitude distance matrices; this
290 assumption may be relaxed in future work. For most of the signals analyzed here, there is little
291 improvement in accuracy between the signal waveforms compared to the envelope functions
292 (for the frequency bands considered here), but this topic deserves further exploration.

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Figure 4. Cluster analysis at station MDJ. Indices along the axes in (A)-(D) are as follows: 0 – 12 (earthquake), 13 – 18 (explosion) and 19 (collapse). A) Phase distance, B) amplitude distance, C) DTW distance, and D) maximum cross-correlation coefficient between every earthquake, explosion and collapse signal pair. The diagonal of each symmetric distance matrix is zero or one. E) and F) show the dendrogram trees from hierarchical cluster analysis for ESA and DTW, respectively. In E), a value of 0.5 is used for the coefficient τ (Eq.5). The indices corresponding to explosions are enclosed by the black rectangle

301 Discussion

302 *Method Sensitivity and Comparison*

303 We conceptually showed how nonlinear warping distances can distinguish dissimilar
304 signal types using synthetics, and when we applied advanced clustering on the actual data, we
305 saw that modest classification accuracies can be achieved. This result may stem from complex
306 wave propagation that is generally not captured when using a 1-D or laterally homogenous
307 earth model but is certainly present in the real Earth. Differences between source types can be
308 deduced from narrowband filters, as is commonly used in P_g/L_g ratio analyses (Pyle and Walter,
309 2021; Tibi et al., 2023). Similarly, the selection of an appropriate filtering passband, minimum
310 SNR, and consequently, time window length was central to our analysis. For MDJ, we analyzed
311 several frequency passbands to capture lower-frequency L_g (< 1 Hz) or higher frequency P (> 2
312 Hz). All signal types from this dataset were present at MDJ, and our clustering approach is most
313 accurate for this station. However, we acknowledge that this is a small, imbalanced dataset and
314 future work should target a labeled database of diverse source types with varying SNR so that
315 the DTW/ESA framework can be further assessed in comparison to neural network classifiers
316 (e.g., Eggertsson et al., 2024; Maguire et al., 2024) or other discrimination approaches.
317 Recently, focal depth discriminants have been developed at local-to-regional distances that use
318 differential magnitudes or spectral amplitude ratios (i.e., R_g/S_g , P_g/S_g) between mine blasts and
319 earthquakes (Koper et al., 2024). In this study, the depth was held constant in the synthetic
320 experiment (1 km) and was 5-km and shallower for the DPRK dataset. There is opportunity to
321 extend the nonlinear alignment framework for depth discrimination as well since focal depth
322 differences can influence body and surface wave excitation (Zhang et al., 2002).

323 *Optimal Signal Separation and Monitoring Implications*

324 Previous studies have shown that for ESA, either amplitude or phase distance can perform
325 better for a particular signal classification application (Tucker et al., 2014) and linear
326 combinations of them can provide better results than using either one alone. Using MDJ as an
327 example, we saw that D_y by itself and the joint combination of D_y and D_x had a greater
328 classification accuracy than D_x (Table 1, Figure 4). Why might that be? It could be due to the
329 higher relative P_n and P_g amplitudes on the explosion waveforms, which are present across a
330 wide frequency band. Alternatively, it may be due to the complexity of phase information
331 between these source types, which suggests phase-based metrics alone may not offer a simple
332 interpretation. We are also interested in the relationship between cross-correlation coefficient
333 and phase or amplitude distance, as well. We showed that a given distance matrix has an
334 inverse relationship to the similarity matrix, consistent with our hypothesis that any two highly
335 similar signals will have a smaller distance between them in phase or amplitude spaces (Figures
336 1 and 4). ESA could be adapted to assist empirical cross-correlation signal detection by

337 extending the correlation range of templates to account for small differences in source
338 mechanism or Green's function. One of the biggest shortcomings of standard correlation
339 detectors is the curation of an optimal template library and appropriate detection statistic
340 (Gibbons, 2022). We believe automatic event screening could leverage one (or more) signal
341 distance spaces to address this issue.

342

343 **Conclusion**

344 We have shown that nonlinear alignment techniques such as DTW or ESA have potential to
345 discriminate signal types, with special consideration to frequency content, time-window, and
346 component analyzed. Low magnitude events may be difficult to classify due to their lower SNR
347 when regional distance stations are used. The potential advantage of the discriminant method
348 we presented here is that one can use the full waveform, increasing the available time
349 bandwidth product. Future directions include examining the transportability of this
350 discriminant method using a larger regional dataset, systematic evaluation of how monitoring
351 distance, frequency passband, or how varying SNR influences results.

352 **Data and Resources**

353 To replicate our workflow, the *Computer Programs in Seismology* (CPS) software must be
354 compiled (installation here: <https://www.eas.slu.edu/eqc/eqccps.html>). Earthquake, collapse,
355 and explosion waveforms are freely accessible through the IRIS data web-service (last accessed
356 on May 10, 2024). Maps are made using the PyGMT software (Uieda et al., 2023).

357 **Declaration of Competing Interests**

358 The authors acknowledge that there are no conflicts of interest recorded.

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369 **References**

- 370 Alvizuri, C., and C. Tape (2018). Full moment tensor analysis of nuclear explosions in North
371 Korea. *Seis. Res. Lett.*, 89(6), 2139–2151. <https://doi.org/10.1785/0220180158>.
372
- 373 Anderson, K. R. and J. E. Gaby (1983). Dynamic Waveform Matching. *Information Sciences*, 31,
374 221 – 242.
375
- 376 Ben-Menahem, A. and Singh, S.J. (198). *Seismic Waves and Sources*, Springer Verlag, New York,
377 NY.
378
- 379 Chun, K-Y and P. G. Richards (2004). *Dongbei Broadband Network* [Data set]. International
380 Federation of Digital Seismograph Networks. https://doi.org/10.7914/SN/5G_2004.
381
- 382 Eggertsson, G., Lund, B., Roth, M., and Schmidt, P. (2024). Earthquake or blast? Classification of
383 local-distance seismic events in Sweden using fully connected neural networks, *Geophys.*
384 *Jour. Int.*, 236, 1728 – 1742, doi.org/10.1093/gji/ggae018.
385
- 386 Ford, S. R., D. S. Dreger, and W. R. Walter (2009). Source Analysis of the Memorial Day explosion,
387 Kimcaek, North Korea, *Geophys. Res. Lett.*, 36, L21304, doi:10.1029/2009GL040003.
388
- 389 Gibbons, S. J. (2022). The optimal correlation detector?, *Geophys. Jour. Int.*, 228, 355 – 365,
390 doi.org/10.1093/gji/ggab344.
391
- 392 Hale, D. (2013). Dynamic warping of seismic images, *Geophys.*, 78(2), S105–S115.
393
- 394 Herrmann, R. B. (2013). Computer programs in seismology: An evolving tool for instruction and
395 research, *Seism. Res. Lettr.* **84**, 1081-1088, doi:10.1785/0220110096.
396
- 397 Joshi, S.H., E. Klassen, A. Srivastava, I. H. Jermyn (2007). A novel representation for Riemannian
398 analysis of elastic curves in R^n , *Proceedings of IEEE CVPR*. pp. 1–7.
399
- 400 Kumar, U., C. P. Legendre, L. Zhao, and B. F. Chao (2022). Dynamic time warping as an alternative
401 to windowed cross correlation in seismological applications, *Seismological Research*
402 *Letters*, 93(3), 1909–1921. doi: 10.1785/0220210288.
403
- 404 Kong, Q., R. Wang, W. R. Walter, M. Pyle, and B. Schmandt (2022). Combining Deep Learning
405 With Physics Based Features in Explosion-Earthquake Discrimination, *Geophys. Res. Lett.*,
406 49, e2022GL098645. <https://doi.org/10.1029/2022GL098645>

407 Linville, L., K. Pankow, and T. Draelos, T. (2019). Deep learning models augment analyst decisions
408 for event discrimination, *Geophys. Res. Lett.*, 46, 3643–3651.
409 <https://doi.org/10.1029/2018GL081119>.
410

411 Maguire, R., B. Schmandt, R. Wang, Q. Kong, and P. Sanchez (2024). Generalization of Deep-
412 Learning Models for Classification of Local Distance Earthquakes and Explosions across
413 Various Geologic Settings, *Seismol. Res. Lett.* XX, 1–10, doi: 10.1785/0220230267.
414

415 Mikesell, T. D., A. E. Malcolm, D. Yang and M. H. Haney (2015). A comparison of methods to
416 estimate phase delays: numerical examples for coda wave interferometry, *Geophysical*
417 *Journal International*, 202, 347 – 360, doi: 10.1093/gji/ggv138.
418

419 Müller, M. (2021). Fundamentals of music processing using Python and Jupyter notebooks,
420 *Springer*, doi:10.1007/978-3-030-69808-9.
421

422 Myers, S. C., S. R. Ford, R. J. Mellors, S. Baker, and G. Ichinose (2018). Absolute locations of the
423 North Korean nuclear tests based on differential seismic arrival times and InSAR.
424 *Seismological Research Letters*, 89(6), 2049–2058. <https://doi.org/10.1785/0220180123>.
425

426 Pasyanos, M. E. and A. Chiang, (2022). Full moment tensor solutions of U.S. underground
427 nuclear tests for event screening and yield estimation, *Bull. Seis. Soc. of Amer.*, 112, 538–
428 552.
429

430 Pyle, M. L., and W. R. Walter (2021). Exploring the Effects of Emplacement Conditions on
431 Explosion P/S Ratios across Local to Regional Distances, *Seismol. Res. Lett.* 93, 866–879,
432 doi: 10.1785/0220210270.
433

434 Srivastava, A., E. Klassen, S. Joshi, and I. Jermyn (2011). Shape analysis of elastic curves in
435 Euclidean spaces. *IEEE Transactions on Pattern Analysis and Machine Intelligence* 33 (7),
436 1415–1428.
437

438 Stevens, J. L. and S. M. Day (1985). The physical basis of mb: MS and variable frequency
439 magnitude methods for earthquake/explosion discrimination. *Journal of Geophysical*
440 *Research*, 90, 3009–3020.
441

442 Tan, J., and C. A. Langston (2022). Shape Dynamic Time Warping for Seismic Waveform
443 Inversion, *Bulletin of the Seismological Society of America*, XX, 1–18, doi:
444 10.1785/0120220051.

445 Tibi, R. (2021). Discrimination of Seismic Events (2006– 2020) in North Korea Using P/Lg
446 Amplitude Ratios from Regional Stations and a Bivariate Discriminant Function, *Seismol.*
447 *Res. Lett.* 92, 2399–2409, doi: 10.1785/0220200432.
448

449 Tibi, R., N. Downey, and R. Brogan (2023). Testing and Design of Discriminants for Local Seismic
450 Events Recorded during the Redmond Salt Mine Monitoring Experiment, *Bull. Seismol.*
451 *Soc. Am.* 114, 906–923, doi: 10.1785/0120230193.
452

453 Tucker J. D., Wu, W., and Srivastava, A (2013). Generative models for functional data using and
454 amplitude separation, *Comp. Stats. and Data Analysis* 61, 50 – 66,
455 doi:10.1016/j.csda.2012.12.001.
456

457 Tucker, J. D., W. Wu, and A. Srivastava (2014). Analysis of signals under compositional noise with
458 applications to SONAR data, *IEEE Journal of Oceanic Engineering*, vol 29, no. 2. pp 318-
459 330.
460

461 Uieda, L., Tian, D., Leong, W. J., Schlitzer, W., Grund, M., Jones, M., et al. (2023). PyGMT: A
462 Python interface for the Generic Mapping Tools. (v0.9.0) [Software]. Zenodo.
463 <https://doi.org/10.5281/zenodo.7772533>.
464

465 Walter, W. R., D. A. Dodge, G. Ichinose, S. C. Myers, M. E. Pasyanos, and S. R. Ford (2018). Body-
466 wave methods of distinguishing between explosions, collapses, and earthquakes:
467 Application to recent events in North Korea, *Seismol. Res. Lett.* 89, 2131–2138.
468

469 Yuan, C., J. Bryan, and M. Denolle (2021). Numerical comparison of time-, frequency and
470 wavelet-domain methods for coda wave interferometry, *Geophysical Journal*
471 *International*, 226(2), 828–846, doi: 10.1093/ GJI/GGAB1.
472

473 Zhang, J., T. Lay, J. Zaslów, and W. R. Walter (2002). Source effects on regional seismic
474 discriminant measurements, *Bull. Seismol. Soc. Am.* 92, 2926–2945.

475 Tables

476 **Table 1:** Hierarchical Clustering Results for Regional Seismic Stations. Note that due to data
 477 availability issues or low SNR, not all events are included at a particular station.

Component(s)	Signal Type	Events Included	Source Receiver Distance (km)	Window Length (sec)	Filter band (Hz)	% Accuracy y (Dx)	% Accuracy y (Dy)	% Accuracy (DTWdist)
BHN	Waveform	EQ1, EQ2, EQ7, NK5, NK6, CO	344.1	100	0.8 – 8.0	33	45	40
BHZ	Waveform	EQ1, EQ2, EQ3, EQ4, EQ5, EQ6, EQ7, EQ8, EQ9, EQ10, EQ11, EQ12, EQ13, NK1, NK2, NK3, NK4, NK5, NK6, CO	371.0	100	0.5 – 5.0	65	90	90
BHN	Waveform	EQ1, EQ2, EQ3, EQ9, EQ10, EQ11	461.8	120	0.8 – 6.0	67	67	67

478

Station	SEHB	MDJ	SEO2
Netw ork	KS	IC	KS

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