The moment duration scaling relation for slow rupture arises from transient rupture speeds

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12 Key Points:

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- A Burridge-Knopoff model with only two dimensionless parameters; the homogeneous stress on a fault and a velocity strengthening friction term.
- The simplicity of the model allows for both numerical and analytical calculations of moment versus duration scaling relationships during fault slip.
- Moment versus duration scaling relation for slow events arises from transient rupture speeds.

¹⁹ Plain language summary

Observations have shown that the duration of earthquakes is related to the seis-20 mic moment through a power law. The power law exponent is different for regular 21 earthquakes and slow aseismic rupture, and the origin of this difference is currently 22 debated in the literature. In this letter, we introduce a minimal mechanical friction 23 model that contains both slow and regular earthquakes, and demonstrate that the dif-24 ferent power laws emerge naturally within the model because the propagation speed 25 of slow earthquakes decays as a power law in time whereas the propagation speed of 26 regular earthquakes remains fairly constant. 27

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28 Abstract

The relation between seismic moment and earthquake duration for slow rupture fol-29 lows a different power law exponent than sub-shear rupture. The origin of this dif-30 ference in exponents remains unclear. Here, we introduce a minimal one-dimensional 31 Burridge-Knopoff model which contains slow, sub-shear and super-shear rupture, and 32 demonstrate that different power law exponents occur because the rupture speed of 33 slow events contains long-lived transients. Our findings suggest that there exists a 34 continuum of slip modes between the slow and fast slip end-members, but that the 35 natural selection of stress on faults can cause less frequent events in the intermediate 36 range. We find that slow events on one-dimensional faults follow $\bar{M}_{0,\rm slow,1D} \propto \bar{T}^{0.63}$ 37 with transition to $\bar{M}_{0,\text{slow},\text{1D}} \propto \bar{T}^{\frac{3}{2}}$ for longer systems or larger prestress, while the sub-shear events follow $\bar{M}_{0,\text{sub-shear},1D} \propto \bar{T}^2$. The model also predicts a super-shear scaling relation $\bar{M}_{0,\text{super-shear},1D} \propto \bar{T}^3$. Under the assumption of radial symmetry, the 38 39 40 generalization to two-dimensional fault planes compares well with observations. 41

42 **1** Introduction

Over the last decades, an increasing number of slow slip events on faults have 43 been reported (Bürgmann, 2018). A measure that is viewed as a key to unravelling 44 the mechanism of slow and fast rupture is the relation between seismic moment M_0 45 and slip event duration T. Regular fast earthquakes have long been known to follow 46 a moment duration scaling relation of $M_0 \propto T^3$. Ide et al. suggested that slow events 47 follow a unified scaling relation $M_0 \propto T$ (Ide, Beroza, Shelly, & Uchide, 2007). They 48 suggested that the linear relation between moment and duration for slow events can 49 be explained in two ways: (1) the average slip is proportional to the fault length as for 50 fast propagation, and the stress drop is constant for all events, which gives the relation 51 $M_0 \propto T$. (2) the slip amount is constant for all events, and the fault area increases 52 linearly with time $L^2 \propto T$, which results in $M_0 \propto T$. Peng and Gomberg (2010) 53 elaborated on the ideas of Ide et al. (2007) and reached a different conclusion; that 54 rupture should span a continuum between fast and slow velocity end-members. Later 55 studies have reported on a variety of scalings between moment and duration ranging 56 from $M_0 \propto T$ to $M_0 \propto T^2$ (Aguiar, Melbourne, & Scrivner, 2009; Frank, Rousset, 57 Lasserre, & Campillo, 2018; Gao, Schmidt, & Weldon, 2012; Ide, Imanishi, Yoshida, 58 Beroza, & Shelly, 2008). 59

The shape of slip patches can also influence the observed scaling. Ben-Zion 60 (2012) argued that fractal slip patches can result in a scaling relation $M_0 \propto T^2/\log(T)$ 61 because the average displacement is approximately constant rather than proportional 62 to the rupture dimension. Bounded propagation can also play an important role (Ben-63 Zion, 2012; Gomberg, Wech, Creager, Obara, & Agnew, 2016). Gomberg et al. (2016) 64 65 suggested that the scaling relation between moment and duration is the same for slow and fast events, but that a transition occurs between a two-dimensional scaling and 66 a one-dimensional scaling when the rupture propagation switches from unbounded to 67 bounded in one direction. Assuming the fault displacement can be approximated using 68 dislocation theory, this results in a transition from T^2 to T. They suggest that there 69 should be a bimodal but continuous distribution of slip modes, and that a difference 70 in scaling relations alone does not imply a fundamental difference between fast and 71 slow slip. The above mentioned theoretical considerations implicitly assume constant 72 rupture velocity. However, this contradicts observations by Gao et al. (2012) that 73 show that the average rupture speed for slow events decreases with increasing seismic 74 moment, which is a strong indication of transient rupture speeds. 75

Slow slip events emerge in numerical models with rate-and-state friction. Colella, 76 Dieterich, and Richards-Dinger (2011) simulated a Cascadia-like subduction zone using 77 rate-and-state friction. They analyzed a large number of slip events, and found that 78 the seismic moment M_w scales as $M_w \propto T^{1.5}$ for $M_w \leq 5.6$ with a transition to $M_w \propto T^2$ for $M_w > 5.6$. Shibazaki, Obara, Matsuzawa, and Hirose (2012) modeled the 79 80 subsuction zone of southwest Japan with rate-and-state friction. For slow events, they found a scaling $M_0 \propto T^{1.3}$. Liu (2014) used rate-and-state friction on a 3D subduction 81 82 fault model and found a scaling $M_0 \propto T^{1.85}$. Romanet, Bhat, Jolivet, and Madariaga 83 (2018) highlighted the role of interactions between faults. They argue that the scaling 84 relationships of slow slip events and earthquakes emerge from geometrical complexities 85 86 due to interactions between fault segments. The moment duration scalings have not only been addressed using rate-and-state friction. Ide (2008) introduced a brownian 87 walk model for slow rupture, where the assumption is that there is a random expansion 88 and contraction of the fault area, so that its radius can be described as a Brownian 89 walk with a damping term. This model predicts $M_0 \propto T$ for large T. 90

⁹¹ Here, our goal is to answer the following two questions: (1) Is there a separation ⁹² of two distinct classes (Ide et al., 2007), or is there a continuum of possible scaling ⁹³ relations between the fast and slow end-members (Peng & Gomberg, 2010)? (2) Can a ⁹⁴ difference in $M_0 - T$ scaling relations alone be attributed to different physical mecha-

nisms behind slow and fast rupture? We address both (1) and (2) through a Burridge-95 Knopoff type model with Amontons-Coulomb friction with a velocity strengthening 96 friction term that has previously been shown to contain a large variety of rupture 97 phenomena, including sub-shear, super-shear and slow propagation (Thøgersen et al., 98 2019). Velocity strengthening friction has been shown to be a generic feature of dry 99 friction (Bar-Sinai, Spatschek, Brener, & Bouchbinder, 2014), and has been reported 100 in Halite shear zones at low slip speeds or large confining pressures (Shimamoto, 1986). 101 The friction law can also be interpreted as a transition from a dry contact to a lubri-102 cated sliding regime with increasing velocity (a Stribeck curve) under the additional 103 assumption that the transition from dry to contact to lubricated sliding occurs at a 104 small sliding speed (Gelinck & Schipper, 2000; Olsson, Åström, De Wit, Gäfvert, & 105 Lischinsky, 1998). 106

For homogeneously stressed faults, the model can be reduced to only two dimen-107 sionless parameters $\bar{\tau}$ and $\bar{\alpha}$ representing the prestress and a velocity strengthening 108 friction term, respectively. The advantage of such approach is that the simplicity of the 109 model allows us to calculate moment duration scaling relations both through numer-110 ical simulations and through analytical calculations. Through numerical simulations, 111 we demonstrate that there exists a continuum of rupture modes between the slow 112 and fast end-members, but that the most likely selection of $\bar{\tau}$ in nature produces two 113 distinct classes separating sub-shear and slow rupture velocities. Through analytical 114 calculations, we show that the scaling relation for slow fronts arises due to long-lived 115 transients in the rupture velocity. Such transient rupture velocity has been observed 116 in nature through a dependence on the average rupture speed on the seismic moment 117 for slow fronts (Gao et al., 2012). In addition, the model predicts a separate scaling 118 for super-shear rupture not previously reported in the literature. 119

2 A one-dimensional Burridge-Knopoff containing slow and fast rupture

¹²² We solve the one-dimensional Burridge-Knopoff model (Burridge & Knopoff, ¹²³ 1967) with Amontons-Coulomb friction with a linear velocity strengthening term. The ¹²⁴ dimensionless equation of motion for a chain of N blocks can be written as (a detailed ¹²⁵ derivation can be found in the supplementary information)

$$\ddot{\bar{u}}_i - \bar{\bar{u}}_{i-1} - \bar{\bar{u}}_{i+1} + 2\bar{\bar{u}}_i + \bar{\alpha}\dot{\bar{u}}_i - \bar{\tau}^{\pm} = 0, \quad \forall i \in [0, N],$$
(1)

where \bar{u} is the dimensionless displacement

$$\bar{\tau}^{\pm} = \frac{\tau/\sigma_N \mp \mu_k}{\mu_s - \mu_k} \tag{2}$$

¹²⁷ is the dimensionless prestress where σ_N is the normal stress, τ is the initial shear ¹²⁸ stress, and μ_s and μ_k are the static and dynamic coefficients of friction, respectively. ¹²⁹ \pm denotes the sign of the block velocity. For positive velocities, we only need to ¹³⁰ consider $\bar{\tau}^+$, but negative velocities can occur in a small subset of our simulations. In ¹³¹ such situations, we need to prescribe the relation between μ_s and μ_k , which we set to ¹³² $\mu_s = 2\mu_k$, so that $\bar{\tau}^- = \bar{\tau}^+ + 2$. In the rest of the paper we will use $\bar{\tau}$ as a reference ¹³³ to $\bar{\tau}^+$. The second dimensionless parameter

$$\bar{\alpha} = \frac{\alpha}{\sqrt{\rho E}} \tag{3}$$

¹³⁴ is a viscous term, where ρ is the density, E is the elastic modulus, and α is a velocity ¹³⁵ strengthening term with units [Pa s/m]. $\bar{\alpha}$ can range from 0 to infinity, where $\bar{\alpha} = 0$ re-¹³⁶ covers the ordinary Amontons-Coulomb friction without viscosity. $\bar{\tau}$ has an upper limit ¹³⁷ of 1, where the prestress equals the static friction threshold. For $\bar{\tau} < 0$, steady state ¹³⁸ propagation does not occur (Amundsen et al., 2015). This corresponds to a prestress



Figure 1. (a) We solve the one-dimensional Burridge-Knopoff model with Amontons-Coulomb friction with velocity-strengthening dynamic friciction for homogeneously loaded faults. V is the driving velocity, K is the driving spring constant, m is the block mass, k is the spring constant, and f_i is the friction force on block i. The friction law is given by a static friction coefficient μ_s , and a dynamic friction coefficient μ_d plus a velocity strengthening term αv (b). To obtain the seismic moment and duration for a given fault length we fix the block at position N. The model can be written in dimensionless units with only two parameters: $\bar{\tau}$ representing the prestress on the fault, and $\bar{\alpha}$ representing the velocity strengthening friction term. This simple model produces a large variety of slip, including, slip pulses, cracks, sub-Rayleigh rupture, super-shear rupture, slow rupture, and arresting fronts (c). The colorbars show the rupture length \bar{L} of arresting fronts, and the steady state rupture speed $\bar{v}_{c,\infty}$ for given $\bar{\tau}$ and $\bar{\alpha}$ (adapted from Thøgersen et al. (2019))

below the dynamic friction level. We thus simulate $\bar{\tau} \in [10^{-7}, 1]$ and $\bar{\alpha} \in [10^{-3}, 10]$. The boundary conditions assuming triggering through soft tangential loading (small driving velocity V and driving spring stiffness K) are given by $\bar{u}_{-1} = \bar{u}_0 + 1 - \bar{\tau}$. The rightmost block is fixed so that $\bar{u}_N = 0$. Blocks start to move once the static friction threshold is reached, which in dimensionless units can be written as

$$\bar{u}_{i-1} + \bar{u}_{i+1} - 2\bar{u}_i \ge 1 - \bar{\tau} \tag{4}$$

¹⁴⁴ Moving blocks restick if the velocity changes sign. The system is sketched in Figure 1a. ¹⁴⁵ This model has previously been used to determine the steady state rupture velocity ¹⁴⁶ which includes sub-shear, supershear, and slow rupture, as well as an arresting region ¹⁴⁷ at low $\bar{\tau}$ and intermediate $\bar{\alpha}$ (Thøgersen et al., 2019). The steady state front speed ¹⁴⁸ $\bar{v}_{c,\infty}$ can be found exactly when $\bar{\alpha} = 0$ (Amundsen et al., 2015). For $\bar{\alpha} > 0$ we can use ¹⁴⁹ the empirical expression (Thøgersen et al., 2019)

$$\bar{v}_{c,\infty} \approx \frac{1 - e^{-\frac{\bar{\tau}}{\bar{\alpha}(1-\bar{\tau})}}}{\sqrt{1-\bar{\tau}^2}}.$$
(5)



Figure 2. One-dimensional seismic moment $\overline{M}_{0,1D}$ and event duration \overline{T} obtained from simulations. The color of the markers show the average front speed $\langle \overline{v}_c \rangle$. The origin of the four different scaling exponents $\{2(1 - \beta), \frac{3}{2}, 2, 3\}$ is discussed in detail in the text. (a) In the limit of small $\overline{\tau}$, there is a separation in two distinct scalings for fast and slow events. (b) For large $\overline{\tau}$ the model predicts super-shear rupture, which has a different scaling exponent than regular sub-shear earthquakes. (c): For intermediate $\overline{\tau}$, the central part of the diagram is populated. (d) Results from the entire range of $\overline{\tau}$ and $\overline{\alpha}$ show that moment duration can exhibit a continuum of slip modes in between the slow and fast end-members.

¹⁵⁰ 3 Moment duration scaling relations

¹⁵¹ 3.1 Model results

¹⁵² We run simulations until all blocks have ruptured or all blocks have stopped. ¹⁵³ We have performed 1120 simulations to obtain the moment duration diagram with ¹⁵⁴ $N \in 5 \times 2^{\{0,7\}}$. In dimensionless units the zeroth order moment for rupture propagation ¹⁵⁵ along a line is

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$$\bar{M}_{0,1\mathrm{D}} = \langle \bar{u} \rangle \bar{L},\tag{6}$$

where $\langle \bar{u} \rangle$ is the average displacement on a fault of length \bar{L} . We run the simulations until all blocks are immobile, or until the average velocity reaches 0.1% of the steady state slip speed $\bar{\tau}/\bar{\alpha}$ (Thøgersen et al., 2019). The seismic moment and the duration are measured when 99% of the total displacement has been reached.

Figure 2 shows the measured $\overline{M}_{0,1D}$ and event duration \overline{T} for all simulations. If 160 the stress drop is small compared to the absolute shear stress, as is often found for 161 faults (Shearer, Prieto, & Hauksson, 2006), τ should often lie close to the dynamic 162 level, which corresponds to $\bar{\tau} \simeq 0$. For low values of $\bar{\tau}$, the arresting region in $(\bar{\tau}, \bar{\alpha})$ 163 gives rise to a separation of these scaling relations, so that fast and slow rupture 164 fall into two distinct lines in the moment duration diagram (Figure 2a). This is in 165 line with the ideas of Ide et al. (2007). This separation occurs because steady state 166 propagation at small $\bar{\tau}$ and intermediate $\bar{\alpha}$ is forbidden (Figure 1a). If we include also 167 larger prestress values we obtain a continuum of slip modes in the moment duration 168 diagram (Figure 2d), in line with the suggestions of Peng and Gomberg (2010). The 169 model also predicts a second scaling relations for super-shear rupture, which is found 170 at large $\bar{\tau}$, that has not previously been reported (Figure 2b). 171

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3.2 Origin of scaling relations - analytical calculations

The simplicity of the model allows an analytical treatment of several aspects which helps explain the various scaling relations between seismic moment and event duration. We summarize the analytical predictions for slip, front speed and event duration, and explain why the different scaling relations appear. A detailed derivation is given in the supplementary information.

First, we can determine the average slip on a fault. If the stress is at the dynamic level after rupture (the stress drop equals $\bar{\tau}$), we can calculate $\langle \bar{u} \rangle$ exactly

$$\langle \bar{u} \rangle = \frac{\bar{\tau}\bar{L}^2}{3} + \frac{(1-\bar{\tau})\bar{L}}{2}.$$
(7)

Equation 7 is derived for soft tangential loading, and we stress that a different boundary conditions could lead to different dependencies between \bar{L} , $\langle \bar{u} \rangle$ and $\bar{\tau}$. Combining equation 7 with equation 6 we find that the seismic moment can be written as

$$\bar{M}_{0,1D} = \frac{\bar{\tau}\bar{L}^3}{3} + \frac{(1-\bar{\tau})\bar{L}^2}{2},\tag{8}$$

which only depends on the prestress $\bar{\tau}$ and the length of the fault \bar{L} . To obtain the moment duration scaling relation we need to determine $\bar{L}(\bar{T})$, and thus have to combine equation 8 with information about the rupture propagation and the afterslip (i.e. the amount of slip after the propagation has stopped).

A key observation on the rupture propagation is shown in Figure 3. While fast fronts exhibit short transients and quickly reach the steady state propagation speed given by equation 5, slow rupture contains long transients where the propagation speed decays. In the figure, we have illustrated this effect as a decay in the average rupture speed $\langle \bar{v}_c \rangle$ with increasing seismic moment \bar{M}_0 . This result is in line with observations by Gao et al. (2012).



Figure 3. (a) Average propagation speed $\langle \bar{v}_c \rangle$ as a function of seismic moment for $\bar{\tau} \in [10^{-7}, 10^{-3}]$ and $\bar{\alpha} \in [10^{-3}, 10]$. Yellow markers show fast fronts while blue show slow fronts. Grey lines show predictions for $\bar{\alpha} = 10$ and $\bar{\tau} \in [10^{-5}, 10^{-3}]$ from equation 14. The prediction for $\bar{\tau} = 0$ follows $\langle \bar{v}_c \rangle \propto \bar{M}_{0,\mathrm{LD}}^{-\frac{\beta}{2(1-\beta)}}$. (b) This transient velocity can be approximated by subtracting the steady state front velocity $v_{c,\infty}$ from equation 5 and scaling with the initial rupture velocity $v_{c,0}$ found from equation S32. The dashed line shows $(\bar{t}\bar{v}_{c,0})^{-\beta}$ with $\beta = 0.6852$. (c) The same fit when the steady state is not subtracted.



Figure 4. Moment duration scaling examples. Dashed lines show the evolution in time of the seismic moment. (star) shows the final value of the duration and the moment, while (circle) marks the point when the front reaches the end of the fault (i.e without afterslip) for $N = 5 \times 2^{\{0,7\}}$. The solid lines show the predictions of moment versus duration discussed in the text. The top three curves use the slow scaling (equation 17 and 18), while the bottom two use the fast scaling relation (equation 11). The colored regions highlight the different scaling exponents discussed in the text.

¹⁹³ 3.2.1 Fast regime

The short transients in the fast regime indicate that we can approximate the rupture length by

$$\bar{L} = \int_0^{t_{\text{rupture}}} \bar{v}_c(t') dt' \approx \bar{v}_{c,\infty} \bar{t}_{\text{rupture}},\tag{9}$$

where $t_{rupture}$ is the time it takes for a rupture front to reach the end of the fault. The time it takes to arrest completely depends upon the existence of a backward propagating arresting front. If we assume that this backward propagation occurs at roughly the same speed as the forward propagation we obtain

$$\bar{T} \approx \frac{2\bar{L}}{\bar{v}_{c,\infty}} \tag{10}$$

200 so that

$$\bar{M}_{0,1D} \approx \frac{\bar{\tau}}{3} \left(\frac{\bar{v}_{c,\infty} \bar{T}}{2} \right)^3 + \frac{1 - \bar{\tau}}{2} \left(\frac{\bar{v}_{c,\infty} \bar{T}}{2} \right)^2.$$
(11)

This relation implies that there is a separate scaling for sub-shear $(\bar{\tau} \to 0)$ and supershear rupture $(\bar{\tau} \to 1)$ in our simulations:

$$\bar{M}_{0,\text{sub-shear},1D} \propto \bar{T}^2$$
 (12)

203 and

$$\bar{M}_{0,\text{super-shear},1D} \propto \bar{T}^3.$$
 (13)

Note that we also predict that $\overline{M}_{0,\text{sub-shear},1D}$ will transition to a \overline{T}^3 scaling for large \overline{L} if $\overline{\tau}$ is small but nonzero. The moment duration, in the fast regime using equation 11, is shown in the two bottom lines of Figure 4. This figure also shows numerically obtained values for the moment duration. The agreement between the numerical simulations and equation 11 is good.

209 3.2.2 Slow regime

For slow fronts, $\bar{v}_c(\bar{t})$ is transient. To obtain an approximation for $\bar{v}_c(\bar{t}, \bar{\alpha}, \bar{\tau})$, we plot $\bar{v}_c(\bar{t})$ for a selection of $\bar{\tau}$ and $\bar{\alpha}$ in Figure 3. All curves collapse when we subtract the steady state front velocity $v_{c,\infty}$ and scale with the initial rupture velocity $v_{c,0}$ given in equation S32. We can then write down

$$\bar{v}_{c,\text{slow}} \approx (\bar{v}_{c,0} - \bar{v}_{c,\infty})(\bar{v}_{c,0}\bar{t})^{-\beta} + \bar{v}_{c,\infty}$$

$$(14)$$

Figure 3 shows that this relation fits well with simulations for small $\bar{\tau}$, and we measure empirically the exponent $\beta \approx 0.6852$. To obtain a parametric equation for \bar{M}_0 and \bar{T} , we need to find $\bar{T}(\bar{L})$. \bar{T} has two main components; the time it takes to rupture a fault of length \bar{L} , $\bar{t}_{rupture}$, and the time it takes for all motion to stop. Unlike for fast fronts, the arresting phase in the slow regime is not governed by a backward propagating arresting front, but rather a slow exponential decay in velocity. We denote this time $\bar{t}_{afterslip}$, and define

$$\bar{T} = \bar{t}_{\text{rupture}} + \bar{t}_{\text{afterslip}}.$$
(15)

 $\bar{t}_{rupture}$ can be found from equation 14

$$\bar{L} = \int_0^{t_{\text{rupture}}} \bar{v}_c(\bar{t}') \mathrm{d}\bar{t}' \tag{16}$$

$$=\frac{(\bar{v}_{c,0}-\bar{v}_{c,\infty})}{(1-\beta)\bar{v}_{c,0}^{\beta}}\bar{t}_{\text{rupture}}^{1-\beta}+\bar{v}_{c,\infty}\bar{t}_{\text{rupture}},\tag{17}$$

The afterslip time can also be found analytically, and the detailed calculation is given

²²³ in the supplementary information. The result is

$$\bar{t}_{\text{afterslip}} = \log(100) \frac{2\bar{L}^2 \bar{\alpha}}{\pi^2} \tag{18}$$

where log(100) indicates that we take the time when 99% of the slip has been accumulated (which is necessary because the afterslip is exponentially decaying). The prediction of seismic moment versus duration can then be found using equation 8 for the seismic moment, equation 17 for $\bar{t}_{rupture}$ (this has to be solved numerically for nonzero $\bar{\tau}$), and equation 18 for $\bar{t}_{afterslip}$, with $\bar{T} = \bar{t}_{rupture} + \bar{t}_{afterslip}$. The excellent agreement between the analytical approach and the numerical simulations is demonstrated in Figure 4.

We can determine the bound on the slow front scaling relation by noting that for infinitesimal $\bar{\tau}$, $\bar{v}_{c,\infty} \approx 0$ and $\bar{M}_{0,1D} \approx \frac{\bar{L}^2}{2}$. This yields

$$\bar{T}_{\bar{\tau}=0} = \bar{v}_{c,0} (1-\beta)^{\frac{1}{1-\beta}} \bar{L}^{\frac{1}{1-\beta}} + \log(100) \frac{2\bar{L}^2 \bar{\alpha}}{\pi^2}, \tag{19}$$

where the first term will dominate over the second term (negligible afterslip) for large \bar{L} because $\frac{1}{1-\beta} > 2$ for the measured $\beta = 0.6852$. We can then solve for

$$\bar{L} \approx \frac{\bar{T}_{\bar{\tau}=0}^{1-\beta}}{(1-\beta)\bar{v}_{c,0}^{1-\beta}}$$
(20)

²³⁵ which gives us

$$\bar{M}_{0,\text{slow},1\text{D},\bar{\tau}=0} \approx \frac{\bar{L}^2}{2} \propto \bar{T}^{2(1-\beta)} \tag{21}$$

with $2(1 - \beta) \approx 0.6296$. We also observe a transition to a different scaling at large

 $\overline{M}_{0,1D}$ when $\overline{\tau}$ is nonzero. To obtain the exponent in this regime, we note that in this limit the steady state rupture velocity is reached, so that

$$\bar{T} \approx \frac{\bar{L}}{\bar{v}_{c,\infty}} + \log(100) \frac{2\bar{L}^2 \bar{\alpha}}{\pi^2}.$$
(22)

For large
$$\overline{L}$$
 and nonzero $\overline{\tau}$, the afterslip will dominate, so that $\overline{L} \propto \overline{T}^{\frac{1}{2}}$. The cubic
term in equation 8 will eventually dominate, which results in

$$\bar{M}_{0,\text{slow},1\mathrm{D},\bar{L}\gg1,\bar{\tau}>0} \propto \bar{T}^{\frac{3}{2}} \tag{23}$$

This means that the moment duration scaling relation in the slow regime is expected to follow a power law with exponent $2(1 - \beta)$ with a possible transition to $\frac{3}{2}$ at large $\bar{M}_{0,1D}$

²⁴⁴ 4 Discussion

We have demonstrated that a simple Burridge-Knopoff model with Amontons-Coulomb friction is capable of reproducing the range of power law scaling relations between seismic moment and duration observed in nature. The simplicity of the model means that we can calculate the scaling relations analytically, and we find the onedimensional exponents $2(1 - \beta)$ with a transition to $\frac{3}{2}$ for large seismic moments for slow rupture, 2 for sub-shear rupture, and 3 for super-shear rupture, where β is the power law exponent of the transient slow rupture velocity.

In this letter, we aimed to address two questions. First, wether there is a sep-252 aration of two distinct classes, or is there a continuum of possible scaling relations 253 between the fast and slow end-members. We argue that the most likely value for $\bar{\tau}$ is 254 close to 0, which corresponds to shear stress at the dynamic level, or to ruptures where 255 the stress drop is small compared to the background stress like in faults (Shearer et 256 al., 2006). If this is indeed the case, the moment duration scaling should contain a 257 continuum of slip modes between the slow and fast end-members. However, because 258 large $\bar{\tau}$ would in this case be unlikely, it would result in a distinction of fast and slow 259

scalings simply because this is more likely. This would indicate that both the interpretations by Ide et al. (2007) and by Peng and Gomberg (2010) hold in the sense that there is a continuum of slip modes, but the natural variation of $\bar{\tau}$ could result in more frequent events along the end-member scalings.

In our simulations, the separation into the slow and sub-shear scaling relations occurs spontaneously under the assumption that $\bar{\tau} \approx 0$. It has also been suggested that the observed separation could be strongly affected by instrumental limitations (Agnew, 2009). In particular because the vastly different time-scales involved in aseismic and sub-shear rupture require different measurement techniques (Gomberg et al., 2016; Peng & Gomberg, 2010).

The second question we aimed to address was whether a difference in $M_0 - T$ 270 scaling relations alone could be attributed to different physical mechanisms behind 271 272 slow and fast rupture. Our model contains only two dimensionless parameters, which highlights that the observed scaling relations do not necessitate complex underlying 273 mechanisms. The same friction law with different values for the coefficients and a 274 varying prestress can explain the entire range of scaling relations, and the slow scaling 275 regime arises simply because slow rupture speeds are transient. Our findings thus 276 suggest that there could be a separation in slow and fast scalings without a difference 277 in physical mechanism. To assign different physical mechanisms to slow and fast 278 rupture thus requires more observations than a different scaling relation alone. 279

To compare our results to observations on faults, it is instructive to discuss relations that would be obtained for rupture on a 2D plane. If we can assume radial symmetry, we can use the same expression for $\langle \bar{u} \rangle$ as in 1D, but $\bar{M}_{0,2D} = \langle \bar{u} \rangle \bar{L}^2$, which changes the scaling by a term \bar{L} . This changes the scaling relations to

$$\bar{M}_{0,\text{sub-shear},2D} \propto \bar{T}^3$$
 (24)

(26)

 $\bar{M}_{0,\text{super-shear.2D}} \propto \bar{T}^4$ (25)

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where $3(1-\beta) \approx 0.9444$ is the exponent that is dominant for $\bar{\tau} = 0$ at large \bar{L} . This is remarkably close to the hypothesized exponent of 1 from observations (Ide et al., 2007). The transition from $3(1-\beta)$ to 2 also indicates that a simple linear scaling relation between seismic moment and duration for slow events is not appropriate, because it is only valid at $\bar{\tau} = 0$. We find it likely that a scaling in the approximate range $M_0 \propto T$ to T^2 should be observed for slow events, depending also on the decaying exponent β . For a constant $\bar{\alpha}$, this variation in the power law exponent occurs due to changes in the stress state of the interface. This is in line with observations, where different studies have reported on scaling exponents ranging from approximately 1 to 2 (Aguiar et al., 2009; Frank et al., 2018; Gao et al., 2012; Ide et al., 2007, 2008).

 $\bar{M}_{0,\text{slow.2D}} \propto \bar{T}^{\{3(1-\beta),2\}}$

From our results in Figure 3 we are in a position to explain the observed relation between average rupture speed and seismic moment (Gao et al., 2012). A transient rupture speed with a decaying exponent β would result in a two-dimensional scaling relation $\langle \bar{v}_c \rangle \propto \bar{M_0}^{-\frac{\beta}{3(1-\beta)}}$. Gao et al. (2012) observed that slow events follow the approximate relation $\langle \bar{v}_c \rangle \propto \bar{M_0}^{-0.5\pm0.05}$, which indicates that $\beta \approx 0.6 \pm 0.025$. Using equation 26 yields a moment duration scaling relation for slow rupture following $\bar{M}_0 \propto \bar{T}^{\{1.1,1.3\}}$, which is fully consistent with their observed linear relationship between seismic moment and duration.

Here, we have assumed that propagation is not bounded. Gomberg et al. (2016) demonstrated that there will be a change from a two-dimensional scaling to a onedimensional scaling when the rupture propagation goes from unbounded to bounded in one of the directions. While we have demonstrated that different scalings can originate without such effect, a bounded system would add a number of possible transitions in moment duration, and would in principle allow for scaling relations following both the
 two-dimensional and the one-dimensional exponents.

311 5 Conclusion

Linear elasticity and Amontons-Coulomb friction with a viscous term is sufficient 312 to produce a large variety in scaling exponents between seismic moment and duration. 313 This suggests that different scaling relations for fast and slow slip events do not require 314 different or complex underlying physical mechanisms. Our findings also suggest that 315 there exists a continuum of slip modes between the slow and fast slip end-members, 316 but that the natural selection of stress on faults can cause less frequent events in 317 the intermediate range. We find that the sub-shear scaling follows $M_0 \propto T^2$ (which 318 corresponds to T^3 in 2D), while the slow scaling follows $T^{2(1-\beta)}$ (which corresponds to 319 $T^{3(1-\bar{\beta})}$ in 2D) with a transition to $T^{\frac{3}{2}}$ (T^2 in 2D) for larger seismic moments depending 320 on the prestress. $\beta \approx 0.6852$ corresponds to the power law decay in the slow rupture 321 velocity with time. The model also predicts a separate scaling for super-shear rupture 322 with $M_0 \propto T^3$ (T^4 in 2D). 323

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