

1 Gaia revisited: atmospheric carbon dioxide as a symptom of a vegetation ozone driven climate  
2 cycle

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14 **List of abbreviations**

15 AMOC: Atlantic midocean circulation.

16 BCA: Black carbon aerosol.

17 CCN: Cloud condensation nuclei.

18 CFC: Chlorofluorocarbon.

19 CESM2: Atmospheric chemistry model over time with a focus on secondary organic  
20 aerosols.

21  $E_{\text{FOS}}$ : Global emission of  $\text{CO}_2$  from the burning of fossil energy sources.

22  $E_{\text{LUC}}$ : Global emission of  $\text{CO}_2$  related to land use change.

23  $E_{\text{SLGHG}}$ : Increase in northern Pacific ocean  $\text{CO}_2$  emission caused by short lived greenhouse  
24 gas accumulation.

25  $E_{\text{SOD}}$ : Global ocean  $\text{CO}_2$  emission caused by stratospheric ozone depletion

- 26  $E_{\text{SOW}}$ : Ocean CO<sub>2</sub> emission following from a decrease in the solubility of ocean water due  
27 to seasonal ocean warming related to the inclination of the Earth axis.
- 28  $G_{\text{ATM}}$ : Global atmospheric increment in atmospheric CO<sub>2</sub>.
- 29  $G_{\text{ATM-Here}}$ : Global atmospheric growth in CO<sub>2</sub> concentration, calculated here as  $ML_{\text{YI}} - ML_{\text{YD}}$ .
- 30  $G_{\text{ATM-WCRP}}$ : Global atmospheric growth in CO<sub>2</sub> concentration as calculated by the World  
31 Climate Research Program.
- 32  $G_{\text{CB}}$ : Global Carbon Budget.
- 33  $G_{\text{CE}}$ : Global Carbon Equation.
- 34  $IE$ :  $E_{\text{FOS}} + E_{\text{LUC}}$ .
- 35  $INP$ : Ice Nucleating Particles.
- 36  $IPCC$ : International Panel on Climate Change.
- 37  $MCB$ : Marine Cloud Brightening.
- 38  $ML_{\text{YD}}$ : Yearly decrease in atmospheric CO<sub>2</sub> concentration as measured by the NOAA at  
39 Mauna Loa and calculated herein.
- 40  $ML_{\text{YI}}$ : Yearly increase in atmospheric CO<sub>2</sub> concentration as measured by the NOAA at  
41 Mauna Loa and calculated herein.
- 42  $NOAA$ : National Ocean and Atmospheric Administration.
- 43  $NPCE$ : Northern Pacific Carbon Equation.
- 44  $O_{\text{SE}}$ : Global seasonal ocean CO<sub>2</sub> emission as hypothesized by the WCRP.
- 45  $O_{\text{SU}}$ : Global seasonal ocean CO<sub>2</sub> uptake as hypothesized by the WCRP.
- 46  $S_{\text{CHWAB}}$ : Self-lofting chemically reactive arctic winter generated air bubble.
- 47  $S_{\text{CHWOM}}$ : Self-lofting chemically reactive warm ocean generated air mass.
- 48  $S_{\text{LAND}}$ : Global uptake of atmospheric CO<sub>2</sub> through land sinks.
- 49  $S_{\text{LUC}}$ : Surface area of land use change.
- 50  $SOA$ : Secondary organic aerosols.

- 51  $S_{\text{OCEAN}}$ : Global uptake of atmospheric  $\text{CO}_2$  through ocean sinks.
- 52  $S_{\text{TOTAL}}$ :  $S_{\text{LAND}} + S_{\text{OCEAN}}$ .
- 53  $U_{\text{BDD}}$ : Ocean uptake in atmospheric  $\text{CO}_2$  related biological drawdown.
- 54 UKESM1: Atmospheric chemistry model over time
- 55  $U_{\text{SIM}}$ : Ocean uptake in atmospheric  $\text{CO}_2$  caused by the seasonal melt increasing ocean  
56 water solubility due to the inclination of the Earth axis.
- 57  $U_{\text{SLGHG}}$ : Increase in northern Pacific ocean  $\text{CO}_2$  uptake caused by short lived greenhouse gas  
58 accumulation.
- 59 VOC: Volatile Organic Carbon.
- 60 WCRP: World Climate Research Program.

61

## 62 **Abstract**

63

64 A synoptic analysis of the history of climate science introduces all potential climate  
65 drivers. It is demonstrated that the Keeling curve can be seen to reflect yearly ocean  $\text{CO}_2$  emission  
66 minus yearly ocean  $\text{CO}_2$  uptake. It is also shown that the global carbon equation (GCE), at the  
67 foundation of present day climate models is not an equilibrium. This disequilibrium pertains to the  
68 neglect of the influence of industrial pollutants on the seasonal behaviour of the Northern Pacific.  
69 It is also demonstrated that the errors for carbon emission and uptake on land are too important to  
70 allow for the construction of the GCE.

71 Statistics are used to demonstrate that the increase in atmospheric  $\text{CO}_2$  is determined by industrial  
72 pollutant emission. A significant determination coefficient between ocean behaviour and  
73 pollutants emission is interpreted as caused by ozone accumulation in the Arctic night. With the  
74 return of light, in spring, the short lived GreenHouse Gas (GHG) is envisioned to become  
75 relatively warm and self-lofting. The trajectory of this slowly ascending mixture is influenced by,

76 and influencing the existing pressure systems. Part of its heat is incorporated in the Northern  
77 Pacific Ocean that releases its gasses. This mixture enriched with the ocean emitted gasses ascends  
78 to the stratosphere at the equator where it causes stratospheric ozone depletion, which in turn  
79 results in global warming. This new interpretation of the climate cycle through ozone accretion in  
80 in the tropospheric arctic night and ozone depletion by ocean emitted gasses in the stratosphere of  
81 the low latitudes is substantiated by a reconstruction of the Keeling curve. This reconstruction  
82 consists of the superposition of six functions that determine the uptake and emission of CO<sub>2</sub> by  
83 the Northern Pacific Ocean. A regression analysis between the monthly Keeling curve and present  
84 reconstruction indicates that the two are identical.

85 Finally, it is explained that the climate cycle in the Pleistocene represents a response system  
86 between vegetation composition and stratospheric ozone driven by natural pollutants and  
87 stemming from the Mesozoic era when flowering plants made their appearance.

88 The present aberrantly high atmospheric CO<sub>2</sub> concentrations recorded at Moana Loa are  
89 interpreted as resulting from the reiteration of the natural warming phase of the Pleistocene  
90 climate cycle through industrial pollutants.

91

92 **Key words:** ozone, ocean degassing, emission, uptake, particulates, aerosols, global warming,  
93 Mauna Loa, Keeling, self-lofting.

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95

## 96 **Introduction**

97

98 Efforts at describing what drives temperature on Earth date from the era of enlightenment  
99 when Herschel [1] counted sunspots to become "acquainted with certain symptoms or indications  
100 ... of the temperature of the seasons we are likely to have". These early scientists are presented

101 today as giants on who's shoulders we are standing, but science history tends, in retrospective, to  
102 narrow down earlier discoveries to a path validating present day views.  
103 As our new model requires multidisciplinary insights, this introduction will also serve to  
104 familiarize the reader with earlier neglected concepts, now newly braided into a pollutant driven  
105 climate model for system Earth. This historical review cannot be done without remarking upfront  
106 that the vigilance the CO<sub>2</sub> driven model required, generated information of immeasurable value,  
107 on which the present study leans heavily, indeed.

108

### 109 **Early considerations**

110

111 The work of Foote [2] represents such an early discovery, as she is suggested to have been  
112 first at demonstrating the warming effect of atmospheric CO<sub>2</sub>. Foote [2] referring to carbonic acid  
113 gas did conclude that “an atmosphere of that gas would give our Earth a high temperature”. This is  
114 incontestably true, but a rapid evaluation of her early work (Appendix 1, Fig 1) demonstrates that  
115 identical rates of temperature change are obtained for both the carbonic acid and her damp air.

116

117 **Fig 1. Examples of erroneous historical validation of the CO<sub>2</sub> driven climate model for**  
118 **system earth.** Visualization of the table of Foote [2] where her error margin is highlighted as two  
119 colours because she measured temperature after 2 or 3 minutes. Some experiments were shorter  
120 than other ones and this analysis gives the rates of change for identical duration and identical  
121 measuring intervals. It is not possible to evaluate how damp her damp air was. Moreover, it seems  
122 difficult to obtain 100% pure carbonic acid gas as it is very unstable.

123

124 While the *in vitro* reproduction of a theoretical atmosphere with water vapour and “acids”  
125 certainly was revolutionary, the experiment itself had proven that damp air (78% N<sub>2</sub> ± 20% O<sub>2</sub> ±

126 1% H<sub>2</sub>O, and 1% of other gasses, at the time comprising 0.028% CO<sub>2</sub>) is as sensitive to solar  
127 irradiance as pure carbonic acid gas (100% CO<sub>2</sub>). Tyndall [3] repeated Foote's experiments and  
128 found that the most important greenhouse gas was water vapour [4].  
129 Arrhenius [5] who tried to quantify how much of infrared (heat) radiation is captured by carbonic  
130 acid gas and water vapour in Earth's atmosphere, did not really rule out water vapour as a  
131 GreenHouse Gas (GHG), but found CO<sub>2</sub> to be a more effective one. A quick clustering analysis of  
132 the data from his table II illustrates the difference in absorption coefficients for carbonic acid gas  
133 and water vapour (Appendix 2). It indicates, however, that the two compounds investigated by  
134 Arrhenius in his experiment did not differ significantly in absorption effectivity.  
135 Arrhenius [5] dispatched De Marchi's [6] conclusion that water vapour is the main climate driver  
136 on Earth, because he had calculated that the water vapour variability in the atmosphere in De  
137 Marchi's hypothesis required values beyond 100% saturation. We know now that condensation or  
138 crystallisation clouds form at a saturation beyond 100 % humidity for which they require  
139 particulates as condensation cloud nuclei (CCN) or ice nucleating particles (INP).  
140 Such particulates (CCN's or INP's) had made their spectacular appearance in the scientific  
141 perception when the Krakatau explosion of 1883 caused blue suns, twilight glow, and a wide  
142 global atmospheric belt of sky haze in the lower latitudes [7, plate XXXV] at altitudes well  
143 beyond the tropopause [8, p. 348]. The Krakatau emissions had caused unmistakable global  
144 cooling. Abbot and Fowle [9] assigned it to the worldwide volcanic haze, which they simply  
145 found to have been reducing solar irradiation.  
146 Croll's [10] early hypothesis about the occurrence of glacial epochs resulting from the obliquity of  
147 the earth's axis, its precession, and the eccentricity of the earth's orbit around the sun was also  
148 dispatched by Arrhenius [5] who argued that a certain degree of obliquity causes identical  
149 warming on both hemispheres. Arrhenius had neglected the asymmetrical continent distribution  
150 which play a role in the global heat distribution.

151 While sunspots, greenhouse gasses (GHG), including water vapour, particulates and the  
152 Milankovitch cycles, all were found to contribute in different proportions to our daily weather, it  
153 is relevant here to analyse further how carbon dioxide was found to represent the chief global  
154 temperature driver.

155

## 156 **Recent developments**

157

158 Circumstantial evidence with respect to the warming effect of GHG was brought to stage  
159 by Sagan and Mullen [11] with the “faint sun paradox”. These authors required nitrogen  
160 compounds to explain the high temperatures allowing H<sub>2</sub>O to be liquid under the low solar  
161 radiation that a young sun would represent. Later, these GHG were preferred to be CO<sub>2</sub> as “CO<sub>2</sub> at  
162 high concentrations could have performed a similar function as could other polyatomic gasses”  
163 [12]. Please note that, for these authors, any other polyatomic gas could perform a similar  
164 function, so why not water vapour or ozone?

165 In 1940, thanks to the work of John Shaw at Oxford University, it was already known that water  
166 vapour, carbon dioxide and ozone play an important part in the heat balance of the Earth, because  
167 they absorb radiation in the critical part of the spectrum. By 1965, this had even become textbook  
168 knowledge, see for instance in the Handbook of Geophysics [13, fig. 10.2].

169 In spite of the three above mentioned compounds identified as the potential cause for global  
170 warming, it wasn't water vapour, nor ozone but CO<sub>2</sub> that was found to be responsible for global  
171 warming. This may be explained by the scientific approach followed by the “U.S. Standard  
172 Atmosphere of 1962”. This work defines atmospheric composition as obeying the perfect gas law  
173 requiring ideal air to be devoid of water vapour and dust.

174 Nonetheless, instruments continuously registering particle concentrations at high altitude were  
175 already in use since 1955 and testify to a scientific interest under less ideal conditions. This

176 interest did have consequences as when air transport moved to stratospheric altitude, it caused  
177 concern. Indeed, ozone depletion through pollutant emission in the stratosphere was expected to  
178 have a climatic effect [14]. A considerable gap in academic knowledge on pollutants and aerosols  
179 had developed. This was possibly due to the cold war [15].

180 The National Report of the Carbon Dioxide Assessment Committee of 1983 [16] already depicted  
181 an increase in temperature of 3 °C associated with a doubling of the atmospheric CO<sub>2</sub>  
182 concentration, without any reference to aerosols, thus echoing the earlier Charney report [17].

183 The 1983 report [16] dealt with the impact of anthropogenic aerosols in the following way:

184 “Although anthropogenic aerosols are particularly noticeable in regions near and downwind of  
185 their sources, there does not appear to have been a significant long-term increase in the aerosol  
186 level in remote regions of the globe **other than possibly the Arctic**. The climatic impact of  
187 changes in anthropogenic aerosols, if they occur, cannot currently be determined. One cannot even  
188 conclude that possible future anthropogenic changes in aerosol loading would produce worldwide  
189 heating or cooling, **although carbon-containing arctic aerosols definitely causes local**  
190 **atmospheric heating.**”

191 The increase of the tropospheric ozone concentration in the northern hemisphere was found to  
192 stem from photochemical reactions of nitrogen oxides and hydrocarbons emitted by high-flying jet  
193 aircraft [18] and little was said about the properties of ozone as a greenhouse gas, while in the  
194 1983 report [16, table 4.1] tropospheric ozone was found influence surface temperature change  
195 stronger than other constituents.

196 In this same table water vapour was also found to have a strong effect on surface  
197 temperature.

198 The use of water has tripled since the 1960s and the industrial use has at least quadrupled [19] but  
199 water isn't considered in terms of global temperature. Presently, agricultural irrigation raises  
200 serious concern because the daytime cooling from irrigation is overshadowed by nocturnal



201 warming from cloud cover [20]. These studies relate to agriculture only, but industrial cooling  
202 causes water to leave earth surface as hot damp polluted air that rises rapidly to stratospheric  
203 altitudes. These emissions form and pollute both, condensation and the crystallisation clouds  
204 (cirrus), but are not mentioned anywhere in terms of climate drivers.

205 So, in spite of the broad spectrum of potential greenhouse gas mentioned above, earliest  
206 efforts, as we will see below, focussed immediately on CO<sub>2</sub>.

207

### 208 **Modern situation**

209

210 As early as 1966, the oil industry appointed Lovelock to study a geotechnically enhanced  
211 biological pump that would naturally draw down CO<sub>2</sub> back into the oceans. This early symptom-  
212 control-driven science led to the first steps at improving our understanding of System Earth [21]  
213 but it left us with the expectation that the oceans are CO<sub>2</sub> sinks, or rather bins taking up CO<sub>2</sub> at any  
214 temperature.

215 By 1974, Lovelock and Margulis [12] had introduced “life” (photosynthesis) as a  
216 geological force capable of stabilizing the early atmospheric temperature and composition to the  
217 levels we experience today. This is the well-known Gaia hypothesis, requiring the seeding of  
218 condensation clouds with dimethyl sulphides to dim or brighten the ocean surface [21]. Ayers and  
219 Caine [22] tried to corroborate this Gaia hypothesis by quantifying how algal blooms would  
220 provide a feedback loop between life, marine cloud seeding, temperature, and carbon dioxide  
221 concentrations, but they concluded that the modelling of climate change in this way was  
222 incomplete. Nonetheless the search for what constitutes condensation cloud nuclei (CCN)  
223 continued and half a decade ago Zheng et al. [23] summarized that, apart from the sulphides  
224 envisioned by Lovelock et al. [21] numerous other particulates, organic particles, sea salts, volatile  
225 organic carbon (VOC) compounds and their mixtures, all could play a role in condensation cloud

226 growth. The unique relation between algal blooms and dimethyl sulphide or rather the unique  
227 degradation product that they provide to the troposphere as sea spray, was not required to create a  
228 response system between the carbon and the climate cycle.

229         However innovative through the introduction of feedback loops, the Gaia hypothesis has  
230 turned out to be inadequate in that it did not address the formation of the carbon sinks that  
231 permanently withdraw carbon from land, ocean and atmosphere. Moreover, the feedback loop  
232 envisioned in the Gaia hypothesis resulted in a constant cloud cover, temperature and atmospheric  
233 CO<sub>2</sub> concentration, while the Pleistocene climate cycle is characterized by an Antarctic  
234 temperature fluctuating regularly between -8 and 2 °C and a CO<sub>2</sub> concentration fluctuating  
235 between 200 and 300 ppm. These cannot be explained by the Gaia hypothesis, but the present new  
236 hypothesis will explain them.

237

### 238 **The role of ozone**

239         When the first hole in the ozone layer above Antarctica was discovered [24, 25], the focus  
240 of the scientific community shifted to the short-term health implications of ultraviolet radiation  
241 while neglecting the climate issues related to the ozone layer in the stratosphere [26]. In 1990  
242 Yung [27] already explained that sulphates and other aerosols react with ozone and it was clear to  
243 him that the CO<sub>2</sub> concentration in the atmosphere is regulated by the oceans and their temperature.  
244 Later, it was demonstrated that black carbon aerosol (BCA) can carry nitric acid (HNO<sub>3</sub>) and  
245 nitrogen dioxide (NO<sub>2</sub>) and will react with stratospheric ozone converting it to oxygen [28].  
246 Depletion of the stratospheric ozone layer had already been mentioned as having a climatic impact  
247 in 1970 [14]. These insights became more or less forgotten but were ‘rediscovered’ later on [29,  
248 30, 31, 32, 33] and they are the foundation for the present contribution.

249

### 250 **From a broad set of possibilities to CO<sub>2</sub>**

251 By the end of the last century, from sunspots, via orbital forcing, to particulates, GHG and  
252 ozone depletion, all hypotheses had been brought to the scientific arena, but in spite of the  
253 ongoing scientific debate concerning the relative role of each of the various processes at hand, the  
254 first (authorless) 1990 International Panel on Climate Change (IPCC) report [34] introduced its  
255 request to the United Nations for funding international research by stating upfront:  
256 “We are certain of the following: there is a natural greenhouse effect which already keeps the  
257 Earth warmer than it would otherwise be; emissions resulting from human activities are  
258 substantially increasing the atmospheric concentrations of the greenhouse gasses: carbon dioxide,  
259 methane, chlorofluorocarbons (CFCs) and nitrous oxide. These increases will enhance the  
260 greenhouse effect, resulting on average in an additional warming of the Earth's surface. The main  
261 greenhouse gas, water vapour, will increase in response to global warming and further enhance it.  
262 The size of this warming is broadly consistent with predictions of climate models, but it is also of  
263 the same magnitude as natural climate variability”.

264 In granting this proposal, admirable in its multinationalism, the UN deepened the gap between an  
265 objective but slow, and a biased but rapid development of our insights in System Earth. The UN  
266 funding revived the circumstantial evidence of Sagan and Mullen [11] and reinforced the  
267 homeostatic Gaia hypothesis. Simultaneously a NASA research team [35] found that aerosols  
268 from smoke and the simultaneous enhancement of solar ultraviolet radiation due to ozone  
269 depletion, fitted in the scenario of a nuclear winter, thus fully changing the context of the ozone-  
270 hole hype from a climate issue to a short-term health or peace issue. Ozone had, in spite of the  
271 earlier insights [27, 28] left the scientific arena as a potential climate driver.

272

### 273 **The role of CO<sub>2</sub>**

274 The 2007 Nobel Peace Prize, honouring the multinational cooperation of the UN-  
275 sponsored IPCC, without settling the science, had given renewed credibility to a CO<sub>2</sub>-driven

276 System Earth and by 2021, the Nobel Prize for Physics, awarded “for the physical modelling of  
277 Earth's climate, quantifying variability and reliably predicting global warming” implicitly  
278 reinforced the CO<sub>2</sub> driven climate model again. This award had anchored climate modelling  
279 (implicitly atmospheric CO<sub>2</sub> driven climate modelling) deep in time as it honoured a more than 50  
280 year old paper [36]. Manabe and Wetherald had focussed on the influence of CO<sub>2</sub>, cloudiness, O<sub>3</sub>  
281 and albedo on Earth's thermodynamics at fixed relative and absolute atmospheric humidity, but  
282 this focus wasn't at the core of the award. They did find that a doubling of CO<sub>2</sub> at a fixed relative  
283 humidity gave a temperature increase of 2 °C, but also that less than a doubling in O<sub>3</sub> gave about  
284 the same type of result (table 7), while changes in albedo affected temperature even more (table  
285 8).

286 The other recipient of the 2021 Nobel Prize for Physics was Hasselmann for his 1976 paper [37].  
287 Hasselmann did not prove that CO<sub>2</sub> was driving global climate, but he had managed to separated  
288 weather from climate as the two represent different time scales.

289 A first attempt at a hypothetical backbone of a fully CO<sub>2</sub>-driven System Earth had made its public  
290 appearance in 2007 as the Global Carbon Budget (GCB) where the atmospheric accumulation of  
291 CO<sub>2</sub> consisted of anthropogenic emission minus ocean uptake and an unidentified sink [38, fig 4].  
292 While the atmospheric increase in CO<sub>2</sub> concentration was recorded in the Keeling curve, the  
293 industrial emission was known and the required ocean sink had been established [39], in order for  
294 this GCB to be in balance an unidentified sink needed to be discovered. This unidentified sink was  
295 provided two years later [40] and by 2016, the Global Carbon Budget had fully developed into the  
296 Global Carbon Equation (GCE) [41]. This GCE is still in practice and still states that the  
297 atmospheric increase in CO<sub>2</sub> ( $G_{ATM}$ ), as recorded in the Keeling curve based on measurements by  
298 the National Ocean and Atmospheric Administration (NOAA) at Mauna Loa reflects what CO<sub>2</sub>  
299 concentration remains in the atmosphere after ocean ( $S_{OCEAN}$ ) and land sinks ( $S_{LAND}$ ) took up the  
300 bulk of the industrially ( $E_{FOS}$ ) and agriculturally ( $E_{LUC}$ ) emitted CO<sub>2</sub>.

301 It is expressed in the following way

$$302 \quad (E_{FOS}) + (E_{LUC}) = (S_{OCEAN}) + (S_{LAND}) + (G_{ATM}) \quad (0)$$

303 This equation, developed by the World Climate Research Program (WCRP) is endorsed by the  
304 IPCC and is the backbone of most climate models like the 2023 one [42].

305 Simultaneously, ozone depletion was considered by another UN funded organization, the UN  
306 Environment program that developed the Montreal protocol, in order to gain control over the  
307 industrial production of chlorofluorocarbon (CFC) and its consequences for human health. The  
308 focus on chlorofluorocarbons, and since 2016 hydrofluorocarbons, is worrisome as the radical  
309 molecules, which destroy ozone [43] in the tropo- and the stratosphere are ubiquitously present  
310 because they come with industrial pollutants [44].

311 The present contribution aims at alleviating a dangerous division of labour between the IPCC and  
312 the UN environmental program. It is dangerous in the sense that it results in blind spots for both  
313 groups. The potential climatic consequences of ozone accumulation and depletion seem to be  
314 safely salvaged with the Montreal protocol team but its focus is on health. The health issues of  
315 high atmospheric CO<sub>2</sub> concentrations, other GHG or pollutants the atmosphere seem to be safely  
316 salvaged with the IPCC but its focus is on climate.

317

### 318 **Activities of the IPCC**

319

320 The IPCC appears exhaustive in all its potential climate drivers, but their model, following  
321 the UN environmental program, underestimates both the causes and the consequences of ozone  
322 depletion. In this context, it is the role of H<sub>2</sub>O in all its phases that requires more insight.

323

### 324 **The intertwining of H<sub>2</sub>O and aerosols**

325 In the IPCC model of 2022 [45, chapter 7, fig 6], clouds are presented as a diffuse and  
326 continuous shield. This representation does not do justice to water, which, at the pressures and  
327 temperatures related to the mass of our planet, is near its triple point. This means that small  
328 pressure and temperature changes result in shifts in phase. So, instead of considering the cloud  
329 shield as a single diffuse water droplet blanket, one should conceive Earth as protected by a  
330 thermos-troposphere. This troposphere has indeed a double shield or, rather, a double sieve: an  
331 inner liquid cloud sieve and an outer solid ice sieve. While a single blanket may be capable of  
332 reflecting solar radiation back into space, the pathway of solar radiation between these two sieves  
333 can be envisioned as that of heat in a double walled thermos-bottle, keeping the Earth warm. The  
334 external solid sieve of the sphere at the boundary of the tropo- and the stratosphere is composed of  
335 featherlike ice clouds (cirrus clouds). The internal sieve, consisting of condensation clouds  
336 (stratus, stratocumulus or cumulonimbus clouds), that occupy the troposphere. Both cloud layers  
337 behave differently. While the ice clouds are part of the lower current of the stratospheric Brewer  
338 Dobson circulation cell [46] and flow toward the poles of the hemispheric winter [47], the  
339 condensation clouds are confined to three conspicuous climate cells (the Hadley, Ferrell and Polar  
340 cells) and play a more regional role.

341 Both cloud type densities are determined by the availability of both water and particulates. Water  
342 vapour, comprising 1-3% of the tropospheric gasses, isn't limiting the formation of clouds. It is  
343 the specific weight of particulates, their emission frequency and intensity that are a limiting factor  
344 to cloud formation. The altitudinal distribution of the particulate density determines what cloud  
345 type will be formed: at low altitudes condensation clouds will be formed through Cloud  
346 Condensation Nuclei (CCN) and higher up, ice clouds through Ice Nucleating Particles (INP).

347 These particulates affect Earth's kinetics through their various properties: their albedo, their  
348 chemical composition or their role at slowing down solar irradiation through the greenhouse gas  
349 (GHG) that they generate. The balance between these physical and chemical processes is difficult

350 to establish, especially as they can be influenced by seasonality and light availability at the poles  
351 because such pollutants can be sensitive to photolysis.

352

### 353 **Aerosols and earth's albedo**

354 The role of particulates as CCN's and INP's affecting the Earth's albedo continued to get  
355 attention from the IPCC. This is exemplified by their 1997 view [48] that "Tropospheric aerosols  
356 (microscopic airborne particles) resulting from combustion of fossil fuels, biomass burning and  
357 other sources have led to a negative direct forcing" by which they meant cooling. This view is  
358 maintained in their latest report 2023 [49], endorsing the freshly published view that direct  
359 negative forcing of aerosols can even be considered to mask part of the GHG global warming [50].  
360 However, in 2023, actual climate models are still found to fail at adequately accounting for the  
361 role of the dust concentration [51]. Indeed, it appears that aerosol emissions that seed  
362 condensation clouds have caused regional dimming, but cannot explain the subsequent brightening  
363 trends [52]. Similar types of uncertainties were found for aerosol/ice cloud (cirrus) interactions,  
364 where positive or negative albedo effects can be established regionally but these effects fail to be  
365 reproducible in global climate models [53].

366

### 367 **Aerosol's chemical role**

368 It is surprising in this respect that the IPCC only considered aerosols to brighten or dim  
369 solar radiation, because Liss and Lovelock [54] had already pointed at a climate driven by a  
370 cocktail of gasses and aerosols where the latter were expected to react with ozone. This is highly  
371 relevant as it has been found that ozone depletion cools the stratosphere [31] but warms the  
372 troposphere [33, 55] and that the photolysis of ozone often is exothermic [56] while winter  
373 accumulation of ozone in the Arctic can cause the short lived GHG ozone to give seasonal  
374 warming [57]. It must therefore be deduced that the IPCC, did – and does - not pay due attention

375 to the influence that aerosols and pollutants have on the short lived greenhouse gas ozone and its  
376 implication for Earth's energy budget.

377

### 378 **Aerosols in the past**

379

380 An aerosol-driven System Earth is nonetheless supported by our knowledge from earth  
381 history. Patterns between aerosols/dust and climate have been pointed out for the not too far  
382 geological past. Indeed, dust (loess) typically characterizes arid glacial phases of the Pleistocene  
383 [58] Antarctic whereas glacial maxima in the tropics are characterized by higher concentrations of  
384 charcoal [59]. Moreover, areas at midlatitudes, such as the Chinese Loess Plateau, display the  
385 proof of enhanced wildfires during the past eight glacial/interglacial cycles, the intensity of which  
386 has been found to be inversely correlated with global atmospheric CO<sub>2</sub> concentrations [60]. This  
387 indicates that glacial maxima are arid and fire prone and vis versa interglacial maxima are humid  
388 and fire frequency is relatively low.

389

### 390 **Natural fire as a source of pollutants**

391 Forest fires do not only produce large quantities of pollutants such as SO<sub>2</sub> [61] and NO<sub>x</sub>  
392 [62], but these high-temperature fires also carry small particulates (products of incomplete  
393 combustion being chemically reactive) to the tropopause and beyond [63]. It is therefore  
394 conceivable that, under natural circumstances, fire-generated INP form ice clouds in the  
395 tropopause. These ice clouds are part of the lower current of the stratospheric Brewer Dobson cell  
396 [46] which current transports the cirrus clouds with their potentially reactive pollutants to the  
397 poles of each hemispheric winter.

398

### 399 **Tropospheric ozone accumulation and stratospheric ozone depletion**



400 Ozone accumulates in the troposphere during the polar winter night [57, 64] and starts  
401 getting depleted with the return of the light in spring. But ozone also gets depleted in the lower  
402 stratosphere [31, 65]. The cycle of ozone, consequently, requires further scrutiny as ozone  
403 accumulation in the troposphere and its depletion in the stratosphere can cause tropospheric  
404 warming and stratospheric cooling, which stratospheric cooling in turn has been reported to result  
405 in tropospheric warming again [33].

406

#### 407 **A crucial role for the arctic heat anomaly**

408 The short lived GHG ozone has been reported to give heat accumulation in the troposphere  
409 and to result in aberrant polar vortex behaviour [55] the outer rim of these vortices forms the jet  
410 stream. Lower atmospheric warming, the northern jet-stream intensity and sea-ice cover in the  
411 Arctic correlate [66]. This relation between northern Pacific weather, jet streams and polar  
412 vortices is confirmed by past warming phase of the last Pleistocene climate cycle because there is  
413 evidence for the dry trade wind and wet storm periods on Hawaii to be related to the intensity of  
414 this same northern jet stream [67]. Reconstructions of the climate in the past provide therefore  
415 evidence that aerosols can force the arctic heat anomaly through the accumulation of short lived  
416 GHG like ozone and in their wake all the equilibrium shifts, including storm intensity and  
417 frequency, that determine the climate in the Northern Pacific Ocean. Even the frequency of the El  
418 Niño events in South (!) America could thus be explained [68].

419

#### 420 **A second hole in the stratospheric ozone layer**

421 Despite the 1987 Montreal Protocol that globally stopped the industrial use of ozone  
422 depleting substances and appeared to restore the ozone hole above the Antarctic, a second hole  
423 appeared in the ozone layer above the Arctic in 2020 [65]. Moreover, the annual mean arctic  
424 temperature amplification [69] has become increasingly clear. So time has come to determine

425 whether the sum of all emitted pollutants ( aerosols, dust, charcoal, sea spray) rather than dimming  
426 solar radiation through “negative forcing” induces global warming through ozone accumulation in  
427 the Arctic and its depletion in the stratosphere.

428

### 429 **Falsifying the Global Carbon Equation**

430

431 In the following the Global Carbon Equation will be scrutinized. First it will be  
432 demonstrated that the increase in atmospheric CO<sub>2</sub> ( $G_{ATM}$ ) as recorded in the Keeling curve is  
433 equal to the yearly increase in atmospheric CO<sub>2</sub> concentration as measured at Mauna Loa ( $ML_{YI}$ )  
434 minus decrease in atmospheric CO<sub>2</sub> concentration as measured at Mauna Loa ( $ML_{YD}$ ). Knowing  
435 the total industrial emission and total natural uptake from the literature employed by the WCRP,  
436 ocean emission ( $O_{SE}$ ) and uptake ( $O_{SU}$ ), as defined by the GCE, can be calculated and, in  
437 contradiction to the existing assumptions [70], the two are shown to be significantly different.  
438 Also, because the error margins for the land sinks ( $S_{LAND}$ ) and the emission from land use change  
439 ( $E_{LUC}$ ) are important, these errors will be argued to be too large to build on to. Moreover their  
440 relation with the global atmospheric growth, analysed statistically, will be demonstrated to either  
441 be insignificant or to have been inverted. Having established reasonable doubt, a new hypothesis  
442 is presented where the atmospheric CO<sub>2</sub> growth ( $G_{ATM}$ ) is fully generated by the Northern Pacific  
443 Ocean and expressed in the Northern Pacific Carbon Equation (NPCE).

444 Although these data-based considerations might appear to have a flavour of “climate  
445 denialism”, the conclusion will be supported by demonstrating that, instead of industrial CO<sub>2</sub>,  
446 industrial pollutants seasonally cause both an increase in ocean CO<sub>2</sub> emission and a minor increase  
447 in ocean CO<sub>2</sub> uptake, resulting in the extreme yet symptomatic atmospheric CO<sub>2</sub> growth ( $G_{ATM}$ )  
448 that has alarmed us.

449

## 450 **A new carbon equation for the Northern Pacific Ocean**

451

452           This new hypothesis will be substantiated by reconstructing the Keeling curve as the sum  
453 of six functions that determine the uptake and emission of CO<sub>2</sub> by the Northern Pacific Ocean.  
454 Five of the functions are seasonal and reflect heat flows between the tropo- and the hydrosphere  
455 for which heat flows the uptake and emission of CO<sub>2</sub> are considered a proxy. One function,  
456 following from the stratospheric ozone depletion, is global and reflects the heat flow from the  
457 strato- to the hydrosphere, causing global CO<sub>2</sub> emission. Two of the five functions pertaining to  
458 the heat flows between the tropo- and the hydrosphere, reflect the natural seasonal ocean CO<sub>2</sub>  
459 uptake and emission. Two other functions reflect how this heat flow is enhanced by pollutant  
460 emission. A fifth function reflects the biological pump taking up tropospheric CO<sub>2</sub> and  
461 transferring it to the ocean sinks. The sum of these six functions has a convincingly high  
462 probability ( $p = 0$ ) and determination coefficient (0.99896) in correlation with the CO<sub>2</sub> (Keeling)  
463 curve constructed with the monthly data of NOAA since 1961.

464           This new interpretation of the Keeling curve will be shown to be consistent with a  
465 stratospheric ozone and natural fire driven Pleistocene climate cycle and it will be argued that the  
466 conditions triggering the transition from a glacial to an interglacial maximum are presently  
467 reiterated by industrial pollutant emission at the top of an interglacial maximum.

468

469

## 470 **Material and Methods**

471

## 472 **Fact checking the interpretation of science history**

473

474 The data of Foote [2] and Arrhenius [5, table 2] were (Appendices 1 and 2) used  
475 respectively for a histogram and a principle components analysis in PAST as offered in Hammer  
476 et al. [71].

477

#### 478 **Data sets**

479

480 Analysis of hourly daily and monthly averages of atmospheric CO<sub>2</sub> concentrations from  
481 1960 to 2019, are made freely available online by the NOAA Observatory on Mauna Loa and the  
482 South Pole [72] (Appendices 3, 4, 5, 6). The yearly industrial emissions from fossil fuels ( $E_{FOS}$ )  
483 are based on Liu et al. [33, fig 1]. The yearly change in the emission caused by land use ( $E_{LUC}$ ) is  
484 also available online on the website of the Carbon Brief [73]. The yearly uptake in atmospheric  
485 CO<sub>2</sub> by land sinks ( $S_{LAND}$ ) was taken from Gasser et al. [74]. The surface of land use change was  
486 found in Winkler et al. [75, fig. 3]. The yearly ocean sink ( $S_{OCEAN}$ ) values were found in  
487 Bennington et al. [76, fig 3, blue line]. Friedlingstein et al. [42, table 1] was used to transfer all  
488 data from ppm to Gigaton (Gt) carbon. These various data sets are provided in Appendix 7.  
489 Pollutant concentrations in the atmosphere over time were found in Griffiths et al. [77, fig 13, red  
490 line] and represent various atmospheric chemistry models (Appendix 8). These models reflect  
491 various studies on anthropogenic activities having a burden on ozone formation in the troposphere.  
492 Such anthropogenic activities like industrial emission, transport and biomass burning produce  
493 numerous particulates like aerosols, sulphate, NO<sub>x</sub>, sea salt, organic and black carbon and soil  
494 dust. Two models were considered for the analysis of the probabilities and determination  
495 coefficients they displayed in comparison with what Friedlingstein et al. [42] consider industrial  
496 emission ( $IE = E_{FOS} + E_{LUC}$ ). The two models are the UKESM1-0-LL model that focuses more on  
497 volatile organic carbons (VOC), and the CESM2-WACCM model having more focus on

498 secondary organic aerosols (SOA) [77]. The compilation of these regression analyses are  
499 summarized in Table 1.

500

501 **Table 1. Comparison between O<sub>3</sub> production and industrial CO<sub>2</sub> emission**

X axis	Y axis	R <sup>2</sup> :	p:	Equation
IE	UKESM1	0.97837	1.6444E <sup>-42</sup>	Y= -11.36 x <sup>2</sup> + 227.5x - 262.8
UKESM1	CESM2	0.95461	1.1631E <sup>-33</sup>	Y= 0.0003063 x <sup>2</sup> + 1.173 x -226.8
IE	CESM2	0.9511	9.0534E <sup>-37</sup>	Y= -9.12 x <sup>2</sup> + 179.2 x - 310.6

502

503 Two functions modelling the net chemical production minus loss of O<sub>3</sub>, UKESM1-0-LL  
504 with a focus on volatile organic carbons (VOC), and CESM2-WACCM with a focus on  
505 secondary organic aerosols as shown in Griffiths et al. [77, fig. 13], are compared to the  
506 sum of fossils fuels (E<sub>FOS</sub>) [33, fig 1] plus the emission caused by land use (E<sub>LUC</sub>) [73]  
507 considered in the GCE to reflect industrial emission [73] (IE = (E<sub>FOS</sub>) + (E<sub>LUC</sub>)).

508

509 The highest determination coefficient (R<sup>2</sup>) and the lowest probabilities (p), under the assumption  
510 that industrial emission (IE = E<sub>FOS</sub> + E<sub>LUC</sub>) determines the modelled atmospheric chemistry, were  
511 used to determine which model (UKESM1 or CESM2) is best to study of the effects of aerosols  
512 and pollutants on the Northern Pacific Ocean behaviour (Table 1). UKESM1, giving slightly  
513 better determination and probabilities, was chosen for the present demonstration.

514

### 515 **The impact of pollutants and industrial CO<sub>2</sub> compared**

516

517 The polynomial regression analyses of Hammer et al. [71] were used to compare the  
518 determination coefficients (R<sup>2</sup>) and probabilities (p) between the pollutants and the yearly ocean

519 behaviour at Moana Loa, viz. yearly atmospheric CO<sub>2</sub> increase (ML<sub>YI</sub>), yearly atmospheric CO<sub>2</sub>  
520 decrease (ML<sub>YD</sub>) and their difference, being the yearly growth of the atmospheric CO<sub>2</sub> (G<sub>ATM</sub>).  
521 For each regression analysis between ocean behaviour and pollutants or emitted volumes of  
522 industrial CO<sub>2</sub>, identical orders of polynomial functions were used. The data set with all  
523 probabilities and determination coefficients of all examined polynomial regression analyses was  
524 reduced to the 50 most relevant ones and presented in Appendix 7.

525

### 526 **Two methods to calculate atmospheric CO<sub>2</sub> growth**

527

528 Two calculation methods for atmospheric increase in CO<sub>2</sub> concentration are compared, viz.  
529 the method followed by the World Climate Research Program [42] and our own method  
530 (Appendix 3, Fig 2). The WCRP evened out seasonality in the atmospheric growth in CO<sub>2</sub>  
531 concentration (G<sub>ATM-WCRP</sub>) curve following Ballantyne et al. [70]. In their method, the difference  
532 in atmospheric CO<sub>2</sub> concentration between several consecutive years  $\Delta C$  is divided by that  
533 number of years ( $\Delta t$ ) and looks as follows: (G<sub>ATM-WCRP</sub>) = ( $\Delta$  Concentration/  $\Delta t$ ) (Appendix 9).  
534 In the present contribution, the atmospheric increase in CO<sub>2</sub> concentration is calculated by  
535 subtracting the yearly atmospheric CO<sub>2</sub> decrease/ocean uptake (ML<sub>YD</sub>) from the yearly  
536 atmospheric CO<sub>2</sub> increase/ocean emission (ML<sub>YI</sub>) from the consecutive period (G<sub>ATM-Here</sub>). This  
537 was done on the basis of the selection of highest and lowest emission month of each year between  
538 1961 and 2019 as offered by the NOAA (Appendix 3).

539 The value of the yearly atmospheric decrease/ocean uptake in CO<sub>2</sub> (ML<sub>YD</sub>) was calculated by  
540 subtracting the lowest monthly atmospheric CO<sub>2</sub> concentration from that year from the value of  
541 the highest monthly CO<sub>2</sub> concentration preceding it. The yearly atmospheric CO<sub>2</sub> increase/ocean  
542 emission (ML<sub>YI</sub>) was obtained by subtracting the lowest monthly values in atmospheric CO<sub>2</sub>  
543 concentration from the highest monthly concentration following it.

544

545 **Fig 2. Details of the Keeling illustrating two ways to calculate the yearly atmospheric**  
546 **growth ( $G_{ATM}$ ).**

547 A. Method used herein :  $G_{ATM} = ML_{YI} - ML_{YD}$ . B. Method used by the WCRP  $G_{ATM} = \Delta C / \Delta t$ .

548

549 The difference between the two values ( $ML_{YI} - ML_{YD}$ ) constitutes the yearly  $G_{ATM-Here}$ .

550 It was calculated here for both the yearly difference and its 3-year average (Appendix 3 and 9).

551 Box plots of  $G_{ATM-WCRP}$  and  $G_{ATM-Here}$  were compared with the statistical PAST program [71] to  
552 establish whether they were identical or not.

553 The difference between the various calculations of  $G_{ATM}$  was also evaluated by calculating the  
554 percentage error between the various calculation methods:

555 Percentage error = averaged  $G_{ATM}$  – actual  $G_{ATM}$  (being one or three years, as calculated by the  
556 carbon group or as calculated here)/actual  $G_{ATM}$  x 100.

557

558 **Reconstruction of the Keeling curve**

559

560 **Regional seasonality**

561 Hakkarainen et al. [78] showed that seasonality approaches symmetry in the high latitudes  
562 of the southern hemisphere, unlike the northern hemisphere that is clearly asymmetrical. This was  
563 hypothesized here to reflect the near absence of industrial bias to seasonal warming or cooling in  
564 the Antarctic. We reproduced such a yearly emission diagram here for the randomly chosen period  
565 of 1990 to 1996 (Fig 3) using the monthly data sets provided by NOAA, for both Mauna Loa and  
566 The South Pole [79].

567

568 **Fig 3. Visualization of the difference in timing and yearly amplitude between the**  
569 **atmospheric CO<sub>2</sub> concentrations in different geographical localities for the randomly chosen**  
570 **1992-1996 period. (a) Mauna Loa (red line), (b) the South Pole (black line). The data are from**  
571 NOAA [79] (Appendix 5).

572

### 573 **Approximating true seasonality**

574 A second-order regression analysis of the monthly atmospheric CO<sub>2</sub> as recorded by the  
575 NOAA on Antarctica [79] (Appendix 5) served to establish the global ocean emission, expected to  
576 be free of seasonality (Fig 4). This gave a function of  $y = 8.784 \cdot x^2 \cdot 10^{-7} + 102 \cdot x + 331.9$ .

577

578 **Fig 4. Polynomial function drawn through the monthly atmospheric CO<sub>2</sub> concentrations as**  
579 **recorded by NOAA on the South Pole between 1980 and 2022 [79]. This function serves to**  
580 **establish the average seasonal amplitude in atmospheric CO<sub>2</sub> emission of the south pole by**  
581 **subtracting the monthly average from the mean determined by the function  $y = 8.784 \cdot 10^{-5} x^2 +$**   
582  **$102 \cdot x + 331.9$ .**

583 The maximum amplitude of these monthly excursions from the regression line for the month of  
584 weakest (March) and strongest atmospheric CO<sub>2</sub> concentrations (December) were summed up,  
585 averaged and divided by two. This gave an approximation of the seasonal southern hemispheric  
586 positive and negative atmospheric CO<sub>2</sub> excursions from the global ocean emission. It was varying  
587 between of 0.25 to -0.25 ppm (see Appendix 6). This amplitude was considered to represent  
588 seasonal fluctuations free of industrial enhancement.

589

### 590 **Industrial impact on ocean uptake**



591           The function of yearly industrial pollutants enhanced ocean CO<sub>2</sub> uptake (the solubility  
592 pump) for the northern hemisphere was made inversely proportional to the sea ice extent as found  
593 in Yulin et al. [98, fig 6 line 2] and also proportional to the UKESM1 [77] (Appendix 10).

594

#### 595 **Industrial impact on ocean emission**

596           The yearly seasonal Northern Pacific Ocean CO<sub>2</sub> emission was calculated based on the  
597 southern hemisphere as explained above for the yearly seasonal uptake; in order to spread this  
598 function over the correct number of months with the correct intensity, it was expressed in  
599 percentages made proportional to the average day length between 20° and 55° NL (see Appendix  
600 9 and 10) and made to reach the maximum of 0.25 ppm approximated above for the southern  
601 hemisphere.

602           The function of yearly industrial pollutants enhanced ocean CO<sub>2</sub> emission for the northern  
603 hemisphere was constructed with the expectation that the pollutant function applied here  
604 (UKESM1) is proportional to the heat produced by pollutants in the Arctic, which heat is taken up  
605 by the oceans resulting in the zone of CO<sub>2</sub> emission in the northern hemisphere as depicted by  
606 Hakkarainen et al. [81, fig 2]. This extra emission was spread in a yearly returning series of  
607 percentages over time established in comparison to the monthly Keeling curve (Appendix 10). It is  
608 stressed here that these percentages were made identical for each seasonal cycle and that they were  
609 made proportional to the UKESM1 function [77].

610

#### 611 **Stationary and cumulative functions**

612           These two functions of seasonal and industrially driven hydrosphere warming and cooling  
613 were made stationary. For the seasonality of the CO<sub>2</sub> emission and uptake, this is self-explanatory,  
614 as driven by recurrent temperatures, but for the effects of industrial pollutants this is less  
615 straightforward but explained through thermal energy being fully transformed to chemical energy.

616

### 617 **Global warming proportional to pollutant emission**

618           Also part of the solar irradiance energy warming a particular air mass will cause the  
619 airmass to be ascending in the troposphere like a droplet in a lava lamp. This ascending energy  
620 turns again into chemically energy in the stratosphere when the ozone layer gets depleted. The  
621 increase in solar radiation following from the stratospheric ozone depletion was made cumulative  
622 as the thickness of the ozone layer is transformed permanently, it was also made proportional to  
623 the pollutant emission over time as described for the UKESM1 cocktail [77] (Appendix 10).

624

### 625 **Estimating the biological pump**

626           The biological pump is estimated to reach 0.3 ppm/year, based on the increase in  
627 atmospheric CO<sub>2</sub> concentration during El Niño years when the ocean CO<sub>2</sub> emission is not  
628 functioning as a biological pump [80, fig 6]. The intensity of El Niño events were spread mildly  
629 over three months from November to January (Appendix 10). In the absence of discrete and  
630 multidisciplinary data on the increase in El Niño events, the biological pump was made to reduce  
631 linearly from 100% in 1961 to 98% in 2019.

632

633

### 634 **Summing it up and comparing to the Keeling curve**

635           The sum of all six functions (two seasonal, two seasonal-industrial, one biological and one  
636 stratospheric) represents the monthly atmospheric CO<sub>2</sub> concentration from 1961 to 2019 (684  
637 month), which was added to the atmospheric CO<sub>2</sub> concentration of January 1961. This result was  
638 compared with the Keeling curve (Appendix 8, column S) using a polynomial regression analysis.  
639 The detailed explanation of the calculation employed in Appendix 10 is detailed further on in the  
640 interpretations or in Appendix 10.

641

642

## 643 **Results**

644

### 645 **The atmospheric increase in CO<sub>2</sub> concentration**

646

647 The atmospheric increase in CO<sub>2</sub> concentration was calculated as by the World Climate  
648 Research Program [42] ( $G_{\text{ATM-WCRP}}$ ) for the months of strongest emission (from May to May) and  
649 lowest emission (from September to September) and were compared with box plots to the  
650 atmospheric increase in CO<sub>2</sub> concentration calculated in the present contribution ( $G_{\text{ATM-Here}}$ ). This  
651 was done for a period of one year, as well as for the average over three years (Appendix 9, Fig 5).  
652 Although this seems just a simple difference in methodology, it reflects a fundamental difference  
653 in interpretation. While the  $G_{\text{ATM-Here}}$  is considered to fully reflect ocean behaviour being  
654 seasonally influenced by industrial emission,  $G_{\text{ATM-WCRP}}$  is believed to represent the yearly carbon  
655 equation, where seasonality is evened out because seasonal ocean CO<sub>2</sub> uptake is expected to be  
656 equal to seasonal CO<sub>2</sub> emission.

657

### 658 **Fig 5. Box plots allowing for the comparison of six potential methods to calculate the** 659 **atmospheric increase in CO<sub>2</sub> concentration.**

660  $G_{\text{ATM-Here } 1y}$ : method used herein ( $G_{\text{ATM-Here } 1y} = \text{ML}_{Y1} - \text{ML}_{YD}$ ) for one year;  $G_{\text{ATM-Here } 3y}$ : method  
661 used here in for the average of 3 years;  $G_{\text{ATM-WCRP } \text{May } 1y}$ : method used by the WCRP ( $G_{\text{ATM-WCRP}}$   
662  $1y = \Delta C / \Delta t$ ), from May to May of the former year; (d)  $G_{\text{ATM-WCRP } \text{Sept } 1y}$ : method used by the  
663 WCRP from September to September from the former year;  $G_{\text{ATM-WCRP } \text{May } 3y}$ : method used from  
664 May to May of three years before;  $G_{\text{ATM-WCRP } \text{Sept } 3y}$ : method used by the WCRP from September

665 to September from the three years before. Black dashed rectangle: largest spread; blue dashed  
666 rectangle: largest interquartile; red dashed rectangle: lowest and highest medians.

667

668 The box plots allowing for the comparison of the six methods to calculate  $G_{ATM}$  indicate that the  
669 method employed here ( $G_{ATM-Here}$ ) gives the same results as for the methods used by the WCRP.

670 Indeed:

671 a. The median of  $G_{ATM-here}$  as calculated here for a single year is slightly inferior to  $G_{ATM-WCRP}$  for  
672 May and September (May 1y, Sept. 1y), while the median of  $G_{ATM-Here}$  for three years is slightly  
673 larger than the  $G_{ATM-WCRP}$  for September for three years but equal to the other calculations of  
674  $G_{ATM-WCRP}$  for May and September.

675 b. The interquartile range of  $G_{ATM-WCRP}$  September 1 year is larger than all others, including both  
676 calculation methods for  $G_{ATM-Here}$ .

677 c. The same is the case for the spreading of  $G_{ATM-WCRP}$  September 1 year.

678

679 Moreover, the errors between the averages of all 6 calculation methods for the increase in  
680 atmospheric  $CO_2$  concentration were also considered (Table 2). The largest error (-1.56 %) was  
681 obtained for the 3-year calculation of our  $G_{ATM-Here}$ , and the smallest (0,09 %) for the  $G_{ATM-WCRP}$   
682 from the World Climate Research Program for September for 1 year.

683

684 **Table 2. Comparison between the six calculation methods for atmospheric  $CO_2$  growth**

$G_{ATM}$	Error margin
$G_{ATM-WCRP}$ Sept 3y	1.33
$G_{ATM-Here}$ 1y	0.78
$G_{ATM-WCRP}$ May 3y	0.38

$G_{ATM-WCRP}$ Sept 1y	0.09
$G_{ATM-WCRP}$ May 1y	-0.97
$G_{ATM-Here}$ 3y	-1.56

685

686

Error margin for the average value for the atmospheric increase in CO<sub>2</sub> concentration

687

( $G_{ATM}$ ) between 1963 and 2018 calculated as  $G_{ATM-Here} = ML_{YI} - ML_{YD}$  for one or three

688

years or as  $G_{ATM-WCRP} = \Delta C / \Delta t$ , when considering the month of strongest emission (May)

689

and the month of weakest emission (September) for a period of one year and a period of 3

690

years.

691

692 It can consequently safely be said that the  $G_{ATM-Here}$  calculated here as yearly ocean CO<sub>2</sub> emission

693 minus yearly ocean CO<sub>2</sub> uptake reproduces the  $G_{ATM-WCRP}$  as calculated by the World Climate

694 Research Program (WCRP) [42]. This means that the following is true.

$$695 \quad G_{ATM-Here} = G_{ATM-WCRP} \quad (1)$$

696

### 697 **Implications for the Global Carbon Equation**

698

699 The global carbon equation (GCE) states that the total industrially emitted CO<sub>2</sub> ( $IE = E_{FOS}$

700 +  $E_{LUC}$ ) minus total up taken CO<sub>2</sub> ( $S_{TOT} = S_{OCEAN} + S_{LAND}$ ) is equal to the atmospheric increase in

701 CO<sub>2</sub> concentration ( $G_{ATM}$ ), or:

$$702 \quad IE - S_{TOT} = G_{ATM} \quad (2)$$

703 As demonstrated above,  $G_{ATM} = ML_{YI} - ML_{YD}$ , so that the GCE can also be expressed as follows:

$$704 \quad IE - S_{TOT} = ML_{YI} - ML_{YD} \quad (3)$$

705

or

$$706 \quad (IE - S_{TOT}) - (ML_{YI} - ML_{YD}) = X \quad (4)$$

707 If the Global Carbon Equation is true, than  $X = 0$ .

708 This is tested bellow with the box plots of the averages of this equation over time (Fig 6).

709

710 **Fig 6. Superimposed box and jitter plots for equation (2).**  $G_{ATM}$  calculated as  $IE - S_{TOT}$ , as  $G_{ATM}$ -  
 711 Here ( $= ML_{YI} - ML_{YD}$ ) for one and three year, measured as  $G_{ATM-WCRP} = \Delta C / \Delta t$ , for the month of May  
 712 and September for one and three years, respectively; the difference X between the six calculations  
 713 methods for  $G_{ATM}$  based on the measurements at Mauna Loa and  $G_{ATM}$  calculated as  $IE - S_{TOT}$ ; red  
 714 dashes line: Median of  $G_{ATM}$  as calculated here for one year.

715

716 In this box plot we see that X, being the difference between  $G_{ATM}$  calculated as in the GCE, as  
 717 industrial emission (IE) minus global sinks ( $S_{TOT}$ ), and  $G_{ATM}$  established with the monthly  
 718 atmospheric  $CO_2$  record at Mauna Loa, remains positive.

719 Furthermore when comparing X, for instance, to the diurnal fluctuation recorded on the 28-th of  
 720 July 2024, being more than 8 ppm  $CO_2$  ( $>16$  Gt C), these differences appear to be insignificant.

721 But when taking the averages of all calculation methods of X (0,395 Gt C) and comparing it to the  
 722 averages of the various parameters of the GCE (Table 3), it is clear that the imbalance X  
 723 represents a high percentage of the sinks and sources.

724

725 **Table 3. Analysis of average imbalance X in the GCE and the average value of the different**  
 726 **members of the GCE.**

Sinks & sources of $CO_2$	Average (in Gt C/year) from 1963 to 2018	% of the average X (0.395 Gt/year) of the various sinks and sources
$S_{LAND}$	2.18	18%

$S_{\text{OCEAN}}$	1.91	20%
$E_{\text{FOS}}$	5.97	6%
$E_{\text{LUC}}$	1.27	31%

727 The averaged values of all sinks and sources were calculated with Appendix 7. The  
 728 average imbalance in the GCE of X was calculated as  $((100 / \text{Average of sinks \& sources})$   
 729  $* \text{Average X})$ .

730 This demonstrates that the data for the sinks and sources of atmospheric CO<sub>2</sub> are still too tentative  
 731 to allow for the modelling of the GCE.

732

### 733 In search of X

734

735 According to the WCRP [42] seasonal ocean CO<sub>2</sub> emission ( $O_{\text{SE}}$ ) and seasonal ocean CO<sub>2</sub>  
 736 uptake ( $O_{\text{SU}}$ ) are equal and even out. Also according to the WCRP the yearly CO<sub>2</sub> emission  
 737 ( $ML_{\text{YI}}$ ) is the industrial CO<sub>2</sub> emission ( $IE = E_{\text{FOS}} + E_{\text{LUC}}$ ) plus the seasonal ocean CO<sub>2</sub> emission  
 738 ( $O_{\text{SE}}$ ). This can also be expressed as:

$$739 \quad ML_{\text{YI}} = IE + O_{\text{SE}} \quad (5)$$

740 Still according to the WCRP [42], the yearly CO<sub>2</sub> uptake ( $ML_{\text{YD}}$ ) is the total of the CO<sub>2</sub> sinks  
 741 ( $S_{\text{TOT}} = S_{\text{OCEAN}} + S_{\text{LAND}}$ ) plus the seasonal ocean CO<sub>2</sub> uptake ( $O_{\text{SU}}$ ). This can be expressed as:

$$742 \quad ML_{\text{YD}} = S_{\text{TOT}} + O_{\text{SU}} \quad (6)$$

743 As  $IE$ ,  $ML_{\text{YI}}$ ,  $ML_{\text{YD}}$ , and  $S_{\text{TOT}}$  are known,  $O_{\text{SU}}$  and  $O_{\text{SE}}$  can be calculated for each year  
 744 (Appendix 12).

$$745 \quad O_{\text{SE}} = ML_{\text{YI}} - IE \quad (7)$$

746

$$747 \quad O_{\text{SU}} = ML_{\text{YD}} - S_{\text{TOT}} \quad (8)$$

748

749 The box plots of these yearly calculations (Fig 7) indicate that the difference between the seasonal  
750 ocean emission ( $O_{SE}$ ) and the seasonal ocean uptake ( $O_{SU}$ ) is positive, significant and equal to  $X$   
751 for a calculation using  $G_{ATM-Here}$  1y. All other five calculation methods of  $X$  and  $G_{ATM}$  based on  
752 the data at Mauna Loa give similar results but are not shown here for practical reasons. This  
753 indicates that the basal paradigm of Ballantyne et al. [70] stating that, within the approach to the  
754 GCE of the WCRP,  $O_{SE}$  and  $O_{SU}$  are equal, is not true.

755

756 **Fig 7. Box plots illustrating that  $X$ , the imbalance in the GCE, is positive and reflects the**  
757 **difference between seasonal ocean  $CO_2$  emission and seasonal  $CO_2$  uptake.**

758  $ML_{YI}$ : yearly atmospheric increase in  $CO_2$  (Fig 2),  $ML_{YD}$  yearly atmospheric decrease in  $CO_2$  (Fig  
759 2),  $IE$ : Total of the industrial emission ( $E_{FOS} + E_{LUC}$ ),  $S_{TOT}$ : total of the  $CO_2$  sinks ( $S_{OCEAN} +$   
760  $S_{LAND}$ ),  $O_{SE}$ : seasonal ocean emission,  $O_{SU}$ : seasonal ocean  $CO_2$  uptake,  $G_{ATM}$ :  $G_{ATM-Here}$  1y.

761

762  **$X$  compared to land sinks and land use change over time**

763 Quite obviously, the disbalance  $X$ , in the GCE is mild, but nonetheless about 18 % of the  
764 average yearly atmospheric  $CO_2$  uptake through land sinks [74] and even 31 % of the emission  
765 through land use change as newly calibrated by the carbon brief [73].

766 The imbalance in the GCE not only represents a high percentage of the yearly land sinks and  
767 sources, but large error margins also exist in the data sets from the original registrations [74, 75,  
768 76]. It is therefore relevant here to analyse the true relation between atmospheric  $CO_2$  growth  
769 ( $G_{ATM}$ ) and land use change ( $E_{LUC}$ ) or land sinks ( $S_{LAND}$ ) (Fig 8).

770

771 **Fig 8. Counterintuitive trend between atmospheric  $CO_2$  growth ( $G_{ATM}$ ) and land  $CO_2$  sinks**  
772 **and sources.**



773 A. Correlation between atmospheric growth ( $G_{ATM}$ ) and land use change emission in  $CO_2$   
774 ( $E_{LUC}$ ). B. Correlation between atmospheric growth ( $G_{ATM}$ ) and  $CO_2$  land sinks ( $S_{LAND}$ ). C.  
775 Correlation between atmospheric growth ( $G_{ATM}$ ) and the surface area of land use change  
776 ( $S_{LUC}$ ).

777

778 A clear correlation between the development of the sinks and source over time and the  $G_{ATM}$  was  
779 not expected, but the statistical analysis between the three should have confirm the nature of the  
780 correlation hypothesized by the WCRP. This is not the case. The  $CO_2$  emission from land use  
781 change ( $E_{LUC}$ ) correlates negatively with the atmospheric  $CO_2$  growth ( $G_{ATM}$ ). This is confirmed  
782 by the positive correlation between  $CO_2$  land sink uptake ( $S_{LAND}$ ) and the same atmospheric  $CO_2$   
783 growth ( $G_{ATM}$ ). The determination coefficients are weak, but the probabilities are relevant.

784 Consequently the initial hypothesis that agriculture causes  $CO_2$  emission and that land sinks result  
785 in  $CO_2$  uptake is either unproven or incorrect. This potential paradox will be explained below and  
786 in the discussion section, but here we focus on the mathematical implications indicating that land  
787 use change rather reflects  $CO_2$  uptake, while land sinks rather reflect  $CO_2$  emission.

788 The relation distilled from GCE stating that industrial  $CO_2$  emission minus ocean and land  $CO_2$   
789 sinks uptake represent the atmospheric  $CO_2$  growth, as expressed in (9):

$$790 \quad (E_{FOS} + E_{LUC}) - (S_{OCEAN} + S_{LAND}) = G_{ATM} \quad (9)$$

791 is debatable because in our statistical analysis  $E_{LUC}$  decreases as  $G_{ATM}$  increases, while  $S_{LAND}$   
792 decrease as  $G_{ATM}$  increases, while in the original GCE,  $E_{LUC}$  had been expected to increase  
793 similarly to  $E_{FOS}$  while  $S_{LAND}$  had been expected to decrease together with  $S_{OCEAN}$ , the statistics  
794 from Figure 8 rather reflects an equation as expressed in (10):

$$795 \quad (E_{FOS} - E_{LUC}) - (S_{OCEAN} - S_{LAND}) = G_{ATM} \quad (10)$$

796 This paradox is elucidated when considering the comparison between the surface of land use  
797 change and atmospheric  $CO_2$  emission that still correlate positively (Fig 8 C). This indicates that

798 something other than CO<sub>2</sub> emission from land use change ( $E_{LUC}$ ) has caused atmospheric CO<sub>2</sub>  
799 growth ( $G_{ATM}$ ) and that this new factor contributing to atmospheric growth correlates to the  
800 surface of land use change ( $S_{LUC}$ ).

801 As was explained in the introduction, volatile organic carbon, water vapour, black carbon, aerosols  
802 and pollutants played a part in the mitigation of cooling at the glacial maximum; here we  
803 hypothesise that the surface of land use change is a proxy for the emission of such pollutants,  
804 which explains the positive correlation between agriculture ( $E_{LUC}$ ) and global atmospheric growth  
805 ( $G_{ATM}$ ).

806

### 807 **A new hypothesis**

808 Thus we can deduce that the Global Carbon Equation stating that atmospheric CO<sub>2</sub> growth ( $G_{ATM}$ )  
809 represents the difference between industrial emission (IE) and natural sinks ( $S_{TOT}$ ) is debatable  
810 because it

811 (1) leans too heavily on tentative estimates of land sinks ( $S_{LAND}$ ) and land use change emission  
812 ( $E_{LUC}$ ),

813 (2) represents a debatable hypothesis on the relation between land sinks and sources of CO<sub>2</sub> to  
814 atmospheric CO<sub>2</sub> growth,

815 (3) and requires that seasonality is evened out which is not the case.

816

817 From here on, the GCE, finding its roots in the 1979 Charney report [17], is abandoned and a new  
818 interpretation of atmospheric growth in CO<sub>2</sub> ( $G_{ATM}$ ) as yearly ocean emission in CO<sub>2</sub> ( $ML_{YI}$ )  
819 minus the yearly ocean uptake in CO<sub>2</sub> ( $ML_{YD}$ ) will be explored.

820

### 821 **Implications of the new hypothesis**

822

823 The chief implication of this new interpretation of  $G_{ATM}$  as ocean generated is that, if  
824 ocean emission ( $ML_{YI}$ ) is superior to ocean uptake ( $ML_{YD}$ ), this imbalance represents the full  
825  $G_{ATM}$  (which is positive) which can only be explained by the oceans releasing more  $CO_2$ , than that  
826 they take up. This means that the oceans are not taking up industrial  $CO_2$  and we have to conclude  
827 that industrial  $CO_2$  emission is not perceived in the troposphere. As industrial  $CO_2$  isn't perceived  
828 in the troposphere it doesn't accumulate there and consequently industrial  $CO_2$  can't cause global  
829 warming. In the absence of  $CO_2$  as the chief driver for global warming and taking into account that  
830 the surface of land use change, being a proxy for pollutant emission, plays a part in the  
831 atmospheric  $CO_2$  increase ( Fig 8), it will be investigated in the following if such pollutants can,  
832 indeed, be demonstrated to drive the warming that causes oceans to release their gasses.  
833 Having reviewed, in the introduction, the historically mentioned diversity in potential climate  
834 drivers, we now build on the predictions of Liss and Lovelock [54] who concluded in the foremost  
835 last paragraph in their paper on climate change that it was "... perhaps naïve to think that one gas  
836 acting alone (e.g. Dimethyl Sulphide) is responsible for particle formation; it is much more likely  
837 that a cocktail is involved. Many of these gases also play potentially important roles in air quality,  
838 particularly halogens in ozone destruction".

839

#### 840 **Comparing pollutant emission and ocean behaviour**

841

842 As shown in the methodology section, UKESM1 is the pollutant emission that compares  
843 best with industrial emission ( $IE = E_{FOS} + E_{LUC}$ ) (Table 1). Here we show how the yearly ocean  
844 emission ( $ML_{YI}$ ), the yearly ocean uptake ( $ML_{YD}$ ) and their difference ( $G_{ATM}$ ) correlate with the  
845 UKESM1 pollutant emission values over time (Table 4).

846

847 **Table 4. Comparison between pollutant emission and Northern Pacific Ocean behaviour.**

X-axis	Y-axis	R <sup>2</sup>	p:	Equation
UKESM1	ML <sub>YI</sub>	0.5442	3.96 E <sup>-11</sup>	0.0116 x + 7.537
UKESM1	G <sub>ATM</sub>	0.40494	7.94 E <sup>-08</sup>	0.00835 x - 2.982
UKESM1	ML <sub>YD</sub>	0.09629	1.78 E <sup>-02</sup>	0.002422 x + 10.52

848

849 The determination coefficients (R<sup>2</sup>) and probabilities of not being correlated (p) from a  
850 polynomial regression analysis between of pollutants (UKESM1) to the yearly ocean CO<sub>2</sub>  
851 emission (ML<sub>YI</sub>), yearly ocean uptake (ML<sub>YD</sub>) and their difference (G<sub>ATM</sub>) are shown  
852 here.

853

854 The R<sup>2</sup> values in Table 4 indicate that the pollutants determine 54% of the yearly ocean CO<sub>2</sub>  
855 emission (ML<sub>YI</sub>), only 9 % of the yearly ocean CO<sub>2</sub> uptake (ML<sub>YD</sub>) and 40 % of the tropospheric  
856 increase in CO<sub>2</sub> concentration (G<sub>ATM</sub>). This indicates that the pollutants have a strong impact on  
857 ML<sub>YI</sub>, and a weak impact on ML<sub>YD</sub>. In other words industrial pollutant emission influences both  
858 ocean CO<sub>2</sub> uptake as ocean CO<sub>2</sub> emission but the emission more than the uptake.

859

### 860 **Ocean behaviour over time**

861

862 When comparing monthly and yearly ocean uptake and emission over time it appears that  
863 both their intensities increased over time. While ocean uptake, on a monthly base, is more intense  
864 than ocean emission, the duration of ocean emission is longer than the duration of ocean uptake  
865 (Fig 9). A near systematic dominance of yearly ocean CO<sub>2</sub> emission over uptake follows from this  
866 emission duration. It also becomes clear from this intensity and duration analysis that for the  
867 period of 1961 to 2019, the duration in month of ocean uptake decreased in favour of ocean  
868 emission that increased in duration.

869

870 **Fig 9. Ocean CO<sub>2</sub> uptake and emission.**

871 A. Duration of ocean CO<sub>2</sub> uptake (red) and CO<sub>2</sub> emission (black) compared. B. Monthly ocean

872 CO<sub>2</sub> uptake (red) and monthly ocean CO<sub>2</sub> emission (black) compared. C. Yearly ocean CO<sub>2</sub>

873 uptake (red) and yearly ocean CO<sub>2</sub> emission (black) compared.

874

875

876 **Interpretation**

877

878 **Modern situation**

879

880 **Pollutants**

881 It has been found above (results section) that it is defensible that industrial CO<sub>2</sub>  
882 accumulation is not perceived in the troposphere and that the oceans are not taking up CO<sub>2</sub>, but are  
883 emitting it. As seen in the methods, the pollutants cocktail holding more Volatile Organic Carbon  
884 (UKESM1) gives better results in comparison to industrial emission than the cocktail holding  
885 more Secondary Organic Aerosols (CESM2) by giving the best determination coefficients (R<sup>2</sup>)  
886 and probabilities (p), so UKESM1 has been considered above (Table 4) and will also be used in  
887 the following for the following for a reconstruction of the Keeling curve.

888

889 **Temperature driving yearly ocean behaviour**

890 Seasonal CO<sub>2</sub> ocean emission, impacted by the pollutant emission, appears in earliest  
891 spring in the highest latitudes of the Northern Pacific, moves gradually towards the low latitudes  
892 until June, after which it disperses in southern direction [81, fig.2]. This emission wave is  
893 interpreted here to results from the heat anomaly for which we invoke the Le Chatelier principle,

894 stating that changes in temperature, pressure, volume, or internal concentration of a solute, each  
895 and all can shift the equilibrium state of the solute [82]. The oceans are a chemical solute with a  
896 constant pressure and volume. Hence, only temperature will shift the internal concentration of its  
897 solutes and it is temperature that drives the ocean CO<sub>2</sub> emission wave recorded in Hakkarainen et  
898 al [81, fig 2].

899

### 900 **Pollutants enhancing seasonal CO<sub>2</sub> emission**

901 This emission wave is interpreted here as induced by a temperature anomaly stemming  
902 from the Arctic because it appears in the high latitudes in spring, at polar sunrise [78, fig. 2]. The  
903 temperature anomaly is considered to following from pollution when for instance nitrogen  
904 compounds in the presence of lightening [83] cause the accumulation of the short lived  
905 greenhouse gas ozone [57]. Ozone has a clear reaction with volatile organic carbons, which  
906 reaction can be both photochemical and exothermic [84]. Because the reactions are  
907 photochemical, ozone accumulates in the arctic winter night and only starts getting depleted when  
908 the light returns in spring. The ozone enriched airmass is considered to become self-lofting in  
909 early spring, as solar irradiation makes it relatively warm, which heat causes it to ascends slowly  
910 over the Northern Pacific Ocean [85, fig. 3], while intensifying its large pressure systems. This  
911 Self-lofting CHemically reactive arctic Winter accumulated Air Bubble (SCHWAB) incorporates  
912 in its stormy course part of its heat into the ocean [86] that releases its gasses, amongst other CO<sub>2</sub>.  
913 These gasses come to enrich and react with the ascending airmass, which we can follow thanks to  
914 the science community's focus on CO<sub>2</sub>. This air mass appears in the arctic high latitudes/ low  
915 altitudes and escapes to the stratosphere in the low latitudes/high altitudes [85]. The amount of  
916 pollutants reaching the high latitudes is expected to determine the importance of the arctic heat  
917 anomaly, how much heat is incorporated in the ocean, how much gas the ocean releases and how  
918 much ozone is depleted in the stratosphere.

919

## 920 **Natural and industrial pollutants in the Arctic**

921 As explained in the introduction, natural forest fires produce massive quantities of SO<sub>2</sub>  
922 [61] and other pollutants like NO<sub>x</sub> or NH<sub>3</sub> [62]. Both forest fires and the burning of coals generate  
923 soot particles found in the atmosphere [87]. Smoke from wild fires also contributes to particulates  
924 in the stratosphere [88] as recent wild fires are reported to have brought smoke into the  
925 stratosphere of Canada [89], United States [90] and Australia [91].

926 All these fires reflect complete and incomplete combustion of organic molecules forming  
927 pollutants as alkenes [84] that occur in the natural system in cuticular waxes of plants [91]. Such  
928 molecules also contribute to (green) fossils fuels, rubbers (natural and industrial), and to plastics  
929 [93] that form pollutants. Hence, fire (natural and industrial) injects substantial amounts of partly  
930 and fully combusted pollutants in the troposphere through heat. Elevated temperatures or simple  
931 air transport [94] bring the particulates to the tropopause where they form INP to the cirrus clouds.  
932 Such aerosols are indeed reported from cirrus clouds [95]. Polluted cirrus clouds concentrate  
933 yearly in each hemispheric winter pole [47], driven by the lower current of the Brewer Dobson  
934 stratospheric circulation cell [46].

935

## 936 **Arctic tropospheric chemistry rather than albedo**

937 This interpretation of the seasonal polar enrichment of pollutants in the Arctic results in,  
938 both, the accumulation of the short lived GHG ozone and a changing polar albedo. This duality  
939 explains the uncertainties reported for the predictability of aerosol-ice cloud (cirrus) interactions  
940 [53], because these uncertainties are only albedo based. But aerosols do not only dim solar  
941 irradiance, they are also responsible for chemical polar ozone accumulation causing the SCHWAB  
942 to form. The accumulation initiates the formation of a localized heterogenic tropospheric airmass  
943 containing relative high concentrations of the short lived GHG ozone. This localized and short

944 lived GHG air mass becomes relatively warm as soon as it receives the first light beams in spring.  
945 Part of the heat generated by the ascending air mass gets incorporated in the ocean through storms  
946 and results in the emission of a broader spectrum of ocean gasses as suggested earlier by Liss and  
947 Lovelock [54] and illustrated here in Table 5.

948

949 **Table 5. Solubility indicators for pure gasses.**

<b>Gas type</b>	<b>Concentration in water at 0°C</b>	<b>Concentration in water at 60°C</b>	<b>Solubility indicator ((concentration at 0°C - concentration at 60°C)/60)</b>
<b>H<sub>2</sub></b>	0 .0019	0 .00115	0 .0000125
<b>N<sub>2</sub></b>	0 .029	0 .01	0 .000316667
<b>CH<sub>4</sub></b>	0 .04	0 .012	0 .000466667
<b>CO</b>	0 .045	0 .01	0 .000583333
<b>He</b>	0 .04	0 .0013	0 .000645
<b>O<sub>2</sub></b>	0 .07	0 .0225	0 .000791667
<b>Ar</b>	0 .1	0 .03	0 .001166667
<b>C<sub>2</sub>H<sub>6</sub></b>	0 .13	0 .02	0 .001833333
<b>C<sub>2</sub>H<sub>4</sub></b>	0 .28	0 .12	0 .002666667
<b>CO<sub>2</sub></b>	3 .45	0 .6	0 .0475
<b>H<sub>2</sub>S<sub>2</sub></b>	7	1 .5	0 .091666667
<b>Cl<sub>2</sub></b>	10	3 .2	0 .113333333
<b>SO<sub>2</sub></b>	225	25	3 .333333333
<b>NH<sub>3</sub></b>	900	180	12

950



951 The solubility indicator was calculated by subtracting a gas's concentration in water  
952 arbitrarily chosen at 60 °C, from its equally arbitrarily chosen concentration at 0 °C and  
953 dividing it by 60 ((concentration in g gas/ kg water at 0 °C ) - (concentration in g gas/ kg  
954 water at 60 °C)/60). It is summarized here for gasses at one atmosphere under air with a  
955 normal composition. Data were found in the Engineering toolbox [96]. This indicates what  
956 gasses are more soluble than CO<sub>2</sub>, and consequently have already escaped the ocean when  
957 CO<sub>2</sub> has escaped; such data are not readily available for SF<sub>6</sub>.

958

959 This ocean gas emission explains the divergence between ocean CO<sub>2</sub> uptake and emission  
960 (Fig.10): the ocean gas emission adds to the temporary GHG effect of the SCHWAB. The  
961 statistical analysis of the monthly ocean emission for the Northern Pacific Ocean by year (Fig.  
962 10), indeed, objectivizes the trend already observed in Figure 9 and shows that both monthly  
963 emission and monthly uptake increased over time, but that monthly ocean emission increased  
964 faster than monthly ocean uptake.

965

966 **Fig 10. Trends in monthly increase in CO<sub>2</sub> emission and uptake over time compared.**

967 Here we see that monthly emission increases faster than monthly uptake.

968

969 Such ocean emitted gasses, water vapour included, ascend to the stratosphere where they cause  
970 ozone depletion. Global stratospheric ozone anomalies were indeed described by Hassler et al.  
971 [31, fig.6] as related to aerosol emissions and they depicted a clear decrease in stratospheric O<sub>3</sub>  
972 percentages over time, for which there are numerous recent confirmations [32, 33, 55]. The figure  
973 6 of Hassler et al. [31] is even reminiscent of warming stripes of Hawkins [97].

974

975 **Pollutants enhancing seasonal CO<sub>2</sub> uptake.**

976 Yearly ocean CO<sub>2</sub> uptake (ML<sub>YD</sub>) is interpreted here as related to a growing volume of  
977 cold arctic water forming each year [98], because of a larger amount of industrial pollutant  
978 emission results in more heat generated each year and a larger volume of melted ice. Atmospheric  
979 CO<sub>2</sub> uptake can be seen as related to the sea-ice melting processes [99]. As the melted ice is also  
980 polluted it is expected to cause some eutrophication thus intensifying the biological pump taking  
981 atmospheric CO<sub>2</sub> down into ocean sediments sinks.

982

### 983 **Atmospheric increase in CO<sub>2</sub> concentration**

984 The atmospheric increase in CO<sub>2</sub> concentration (G<sub>ATM</sub>) results from the difference  
985 between yearly CO<sub>2</sub> ocean emission and its uptake. It has a R<sup>2</sup> in correlation to UKESM1 that is 9  
986 % higher than the average of emission and uptake. This suggests that some of the atmospheric  
987 growth in CO<sub>2</sub> concentration cannot be explained by the UKESM1 pollutant emission alone. First  
988 of all UKESM1 does not cover all the pollutants as Secondary Organic Aerosols having an albedo  
989 effect are less represented in that cocktail [77]. Also some of the atmospheric growth in CO<sub>2</sub>  
990 concentration can also be explained by the greenhouse effect of all the gasses escaping from the  
991 warming northern Pacific sea surface. In this way they contribute to further emission while they  
992 travel up to the stratosphere, as for instance water vapour, but also other gasses more soluble than  
993 CO<sub>2</sub> (Table 5). Finally this difference can also be explained by the pollutants brought directly into  
994 the stratosphere during air transport [94] and depleting the stratospheric ozone directly or  
995 depleting the stratospheric ozone directly by being produced at such high temperatures (for  
996 instance cement [100]) that the pollutants end up directly in the stratosphere. Each and all of these  
997 processes together in various ratios can be seen to contribute to the faster amplification of the  
998 Northern Pacific Ocean CO<sub>2</sub> emission over its uptake.

999

### 1000 **Duration of CO<sub>2</sub> emission versus CO<sub>2</sub> uptake in the northern Pacific Ocean**

1001 Starting 1980 (Fig 11), the duration of yearly ocean emission increased. Yulin et al, [98,  
1002 fig.4] show that changes in the rate of decrease in sea ice extent occurred by the end of the nineties  
1003 of the former century. In recording the moment of highest CO<sub>2</sub> uptake each year, the start of the  
1004 uptake season is seen to shift very mildly from April to May while yearly ocean emission starts  
1005 earlier in August instead of September (Appendix 13). This results in a slightly shorter uptake  
1006 period.

1007

1008 **Fig 11. Analysis of how the emission season became longer than the uptake season, for 60**  
1009 **years, each time considering 2 decades.**

1010 A. transition from emission to uptake moves to a later period. B. transition from uptake to  
1011 emission moves to an earlier period.

1012

1013 **A liquid and a solid shield for the Earth surface**

1014 Solid Earth's infra-red emission [101] increases the temperature from the stratosphere to  
1015 the troposphere where solar radiation accumulates because it is briefly slowed down by solid,  
1016 liquid and gaseous particulates. The infra-red emission is caused primarily by the solid Earth  
1017 surface, but also by all dust, aerosols, particulates and GHG, as a function of their absorption  
1018 potential.

1019 The Earth is shielded by two zones of water suspended in the troposphere thanks to particulates,  
1020 the outer solid zone of feathery ice clouds called cirrus formed by INP, and an inner liquid zone  
1021 consisting of condensation clouds formed by CCN. The ice shield determines how much radiation  
1022 enters and leaves the troposphere [102] but also how much pollutants reach the poles, and the  
1023 condensation cloud shield determines how much infra-red light bouncing back from solid earth is  
1024 retained in the lower troposphere. Particulates, pollutants and gas concentration in the troposphere  
1025 consequently play a crucial role in temperature on Earth. Natural and industrial modulation of the

1026 concentration of these gasses and pollutants in the troposphere are considered here to determine  
1027 the intensity, the geographical position and the composition of a gas flow between ocean and  
1028 troposphere. This gas flow, in turn is expected to determine the thickness of the stratospheric  
1029 ozone layer. The thickness of the stratospheric ozone layer determines the Pleistocene cyclicity to  
1030 which present day interglacial maximum belongs.

1031

### 1032 **Seasonal and pollutant driven ocean behaviour**

1033

1034 As demonstrated above, the existing global carbon equation (GCE) is neither in balance  
1035 nor does the statistical analysis of the relation between its components support its present  
1036 formulation. As explained above pollutants can be seen to determine the Pleistocene cyclicity.  
1037 There is consequently sufficient reason to present a new hypothesis for what determines the  
1038 Keeling curve in the Northern Pacific Ocean. It can be expressed as follows:  
1039 “Pollutants cause tropospheric winter ozone accumulation resulting in a relatively warm self-  
1040 lofting chemically reactive air bubble (SCHWAB) starting its ascend as soon as light returns to the  
1041 high latitudes. The trajectory of the SCHWAB influences and is influenced by existing pressure  
1042 systems which causes storm intensification. The storm intensifications results in the incorporation  
1043 of part of the arctic heat anomaly into the Northern Pacific Ocean. The warmed Northern Pacific  
1044 Ocean releases its gasses because the solubility of gasses depends on the temperature of the water  
1045 they are in. With this evaporation the latent heat vaporization generated by the yearly arctic heat  
1046 anomaly is fully consumed. The oceans do not warm up through the arctic heat anomaly but the  
1047 ocean emitted gasses ascend to the stratosphere where they deplete ozone. The stratospheric ozone  
1048 depletion results in stratospheric cooling and in turn global tropospheric warming and global CO<sub>2</sub>  
1049 emission (land and oceans).

1050

1051 When expressing this new hypothesis in an equation, the Northern Pacific Carbon  
1052 Equation (NPCE), the atmospheric increase in CO<sub>2</sub> concentration can be re-defined as:

$$1053 \quad G_{ATM} = (E_{SOW} + E_{SLGHG} + E_{SOD}) - (U_{SIM} + U_{SLGHG} + U_{BDD}) \quad (11)$$

1054 Where:

1055 E<sub>SOW</sub>: Ocean CO<sub>2</sub> emission following from seasonal ocean warming. This function is stationary as  
1056 it relates to the inclination of the Earth axis.

1057 E<sub>SLGHG</sub>: Increase in ocean CO<sub>2</sub> emission caused by the heat generated by the (Short Lived)  
1058 greenhouse gas in the SCHWAB. This increase in ocean emission is made proportional to  
1059 pollutant emission. It is stationary as with this evaporation the latent heat generated by the yearly  
1060 arctic heat anomaly is fully consumed.

1061 E<sub>SOD</sub>: Ocean CO<sub>2</sub> emission caused by stratospheric ozone depletion. Stratospheric ozone  
1062 depletion follows from ocean emitted GHG ascending to the stratosphere because of the warming  
1063 ocean. This stratospheric ozone depletion gives tropospheric warming and is made proportional to  
1064 pollutant emission,. This function is made cumulative as stratospheric ozone depletion is  
1065 irreversible.

1066 U<sub>SIM</sub>: Ocean uptake in atmospheric CO<sub>2</sub> following from the seasonal ice melt increasing ocean  
1067 water solubility (related to the inclination of the Earth axis). This function is stationary as it  
1068 relates to the inclination of the Earth axis.

1069 U<sub>SLGHG</sub>: Increase in ocean uptake in atmospheric CO<sub>2</sub> following intensified ocean cooling caused  
1070 by the short lived GHG in the arctic heat anomaly. This increase in ocean uptake is made  
1071 proportional to pollutant emission. It is stationary as the ocean cooling and hence the CO<sub>2</sub> uptake  
1072 stops as soon as the ice starts accreting again and cold water isn't added to the ocean anymore.

1073 U<sub>BDD</sub>: Ocean uptake in atmospheric CO<sub>2</sub> related to biological CO<sub>2</sub> draw down off the coast of Peru  
1074 because of algal blooms. It is made to weaken slightly over time from 100 to 98% as presently too  
1075 little data exist on the frequency of El Niño events. It is inferred from the ocean emission CO<sub>2</sub>

1076 taking place during El Niño years, assuming that during normal La Niña years, the ocean takes up  
1077 what the El Niño CO<sub>2</sub> emission anomaly emits during El Niño events [80].

1078 The best impression of how these functions intertwine can be obtained when zooming in on a  
1079 shorter period. Here the earliest 1961-1964 period is shown as the cumulative, the stationary  
1080 functions and their sum still fit into one readable diagram (Fig 12).

1081

1082 **Fig 12. Details of the 6 functions that determine the uptake and emission of CO<sub>2</sub> by the**  
1083 **northern Pacific Ocean and their sum.**

1084 (1) Black: ocean CO<sub>2</sub> uptake following from an increase in ocean water solubility due to ocean  
1085 cooling related to seasonal ice melt (U<sub>SIM</sub>). (2) Red: increase in ocean CO<sub>2</sub> uptake following from  
1086 an increase in ocean water solubility due to ocean cooling related to seasonal melt driven by the  
1087 short lived GHG (U<sub>SLGHG</sub>). (3) Blue: ocean CO<sub>2</sub> uptake through biological draw down (U<sub>BDD</sub>). (4)  
1088 Dark green: ocean CO<sub>2</sub> emission following from a decrease in ocean water solubility due to  
1089 seasonal ocean warming (E<sub>SOW</sub>). (5) Turquoise: increase in ocean CO<sub>2</sub> emission following from a  
1090 decrease in ocean water solubility due to ocean warming related to seasonal tropospheric heat  
1091 anomaly stemming from the short lived GHG accumulation in the Arctic (E<sub>SLGHG</sub>). (6) Purple:  
1092 CO<sub>2</sub> emission from cumulative ocean heat uptake throughout the year because of ozone depletion  
1093 in the stratosphere, (E<sub>SOD</sub>). (7) Apple green: monthly sum of all these functions (Sum 1-6).

1094

1095 When adding the atmospheric CO<sub>2</sub> concentrations measured at Mauna Loa by the NOAA in May  
1096 1961 as in Equation 9 (Appendix 10) in the following way:

1097 
$$G_{ATM\ ML-MONTH} = ((E_{SOD} + E_{SLGHG} + E_{SOW}) - (U_{SLGHG} + U_{BDD} + U_{SM})) + 316,89\ ppm \quad (12)$$

1098 and comparing the data set for each month as offered by the NOAA [72] to the function expressed  
1099 with the above equation (Fig. 13) with a regression analysis generated by the statistical program

1100 PAST [71], it is found that they are practically identical. Indeed the probability that the two are  
1101 not identical is 0 ( $p = 0$ ), and they have a determination coefficient of 0.99896.

1102

1103 **Fig 13. Comparing the monthly reconstruction of the Northern Pacific Carbon Equation**  
1104 **(NPCE) based on the assumption it is pollutant driven, and the monthly atmospheric CO<sub>2</sub>**  
1105 **concentrations recorded by NOAA at Mauna Loa.**

1106 A. Function (black) expressing the reconstructed pollutant driven north Pacific CO<sub>2</sub> emission  
1107 behaviour ( $G_{ATM\ ML-MONTH}$ ) compared to (red) the monthly recording of the CO<sub>2</sub> concentration at  
1108 Moana Loa as placed online by the NOAA [72]. B. The regression analysis comparing both  
1109 functions. Details of the functions are given in Appendix 10.

1110

1111 The various components of the Northern Pacific Carbon Equation are presented in detail as  
1112 macro's in Appendix 10 and are explained in the following.

1113

1114 **Seasonally versus industrially driven ocean behaviour**

1115 Seasonality manifests itself nearly identically in both hemispheres. True seasonal ocean  
1116 emission through warming ( $E_{SOW}$ ) and seasonal ocean uptake through seasonal arctic melting and  
1117 hence ocean cooling ( $U_{SIM}$ ), determined by inclination of the earth's axis, are approximately  
1118 comparable on both hemispheres. These are “approximately comparable” as differences in  
1119 continent configuration on both hemispheres may give minor differences. As explained in the  
1120 methods this seasonality accounts for approximately 0,25 ppm of yearly seasonal CO<sub>2</sub> emission  
1121 and -0,25 ppm of yearly seasonal CO<sub>2</sub> uptake.

1122

1123 Industrialisation manifests itself chiefly on the northern hemisphere through the accumulation of  
1124 pollutants during the polar night when photolysis is impeded. In the northern hemisphere

1125 industrial pollutant emission first causes an increase in CO<sub>2</sub> uptake (and other gasses, many being  
1126 more soluble than CO<sub>2</sub>, see Table 5). This uptake follows from a mild intensification of the  
1127 solubility pump due to extra ocean cooling related to extra seasonal arctic ice melt initiated by the  
1128 winter night accumulation of the short lived GHG ( $U_{SLGHG}$ ) and possibly other pollutants also  
1129 behaving like GHG. On the northern hemisphere industrial pollutant emission also manifests itself  
1130 later on as a decrease in ocean water solubility resulting in ocean CO<sub>2</sub> emission. This emission  
1131 follows from slow incorporation in the northern Pacific of the heat generated by the ascending  
1132 SCHWAB ( $E_{SLGHG}$ ).

1133

#### 1134 **The ascend of ozone depleting substances to the stratosphere**

1135 The above considerations explain how, yearly, industrial pollutants form a self-lofting  
1136 chemically reactive arctic winter night generated air bubble (SCHWAB) enriched with ocean  
1137 emitted GHG-pollutant cocktail, ascend through the troposphere.. CO<sub>2</sub> and CH<sub>4</sub>, being the focus  
1138 of climate scientists are participating in the SCHWAB and reveal how the SCHWAB escapes to  
1139 the stratosphere through the Hadley cell at the equator [85, fig 3].  
1140 El Niño events are also accompanied by an increase in atmospheric CO<sub>2</sub> concentration half way to  
1141 the troposphere [80, fig. 6]. This indicates that the arctic heat anomaly can also induce accidental  
1142 variations in atmospheric interconnectivity [68] that do not only result in an ascending SCHWAB  
1143 but also a Self-lofting CHemically reactive warm Ocean generated air Mass (SCHWOM)  
1144 ascending through the troposphere.

1145 The solubility of gasses at various temperatures (Table 5) indicates that when CO<sub>2</sub> is emitted,  
1146 H<sub>2</sub>S<sub>2</sub>, Cl<sub>2</sub>, SO<sub>2</sub> and NH<sub>3</sub> are also emitted. In this respect sulphur dioxide (SO<sub>2</sub>) and chlorine (Cl<sub>2</sub>)  
1147 are extremely relevant as their emissions into the stratosphere cause ozone loss when the chlorine  
1148 (Cl<sub>2</sub>) loading of the stratosphere is high [103]. Sulphur hexafluoride (SF<sub>6</sub>) also needs mentioning,  
1149 in spite of not having been able to insert it in the solubility table (Table 5), as it is both a very



1150 potent GHG and it is reacting with stratospheric ozone [104]. Atmospheric SF<sub>6</sub> has been  
1151 increasing since 1997 and the asymmetrical distribution pattern between the northern and the  
1152 southern hemisphere [105, fig. 1 red or blue circles] indicates it reflects the same type of processes  
1153 as those described here for CO<sub>2</sub>. These temporary chemical ocean warming effects are considered  
1154 here to cause the global stratospheric ozone depletion that give stratospheric cooling and global  
1155 tropospheric warming [33]. This tropospheric warming explains the gradual increase in global  
1156 ocean CO<sub>2</sub> emission (E<sub>SOD</sub>) and was made proportional to the UKESM1 function [77].

1157

### 1158 **The solubility and the biological pump (U<sub>BDD</sub>)**

1159 The relation between the ocean sinks (S<sub>OCEAN</sub>) [76] and the yearly atmospheric growth  
1160 (G<sub>ATM</sub>) at Mauna Loa as conceived herein (Fig 14) shows that the two are co-varying as expected  
1161 by the WCRP. This would indicate that as the northern hemisphere gets warmer the ocean CO<sub>2</sub>  
1162 sinks become larger. This echoes the WCRP hypothesis that industrial CO<sub>2</sub> is taken up by the  
1163 oceans, but as we have seen above and as will be developed in the discussion, industrially  
1164 emitted CO<sub>2</sub> escapes to the stratosphere and presently the Northern Pacific Ocean emit more CO<sub>2</sub>  
1165 than that it takes up.

1166 Nonetheless, in the present temperature driven model more industrial CO<sub>2</sub> emission does means  
1167 that more cold water gets produced at the poles which colder water is expected to take up more  
1168 CO<sub>2</sub>. This effect is feeding both, the chemical CO<sub>2</sub> solubility pump and the biological pump  
1169 because cold water simply contains more dissolved gasses which will serve the biological CO<sub>2</sub>  
1170 drawn down. Moreover the ice has accumulated pollutants, and as the ice melts it causes  
1171 eutrophication which can also be expected to induce a stronger biological draw down of CO<sub>2</sub>. The  
1172 solubility and the biological pump are consequently impossible to disentangle.

1173 Nonetheless, considering the late scientific focus on the weakening of the Atlantic MidOcean  
1174 Circulation [106] and knowing that the El Niño southern Oscillation [107] is part of the general

1175 ocean ventilation process, the biological pump was treated separately as the biological draw down  
1176 ( $U_{BDD}$ ) taking place in upwellings areas. As explained in the Methods [80] it was made to reach  
1177 0,3 ppm of  $CO_2$  yearly starting in november and ending in january and to decrease slightly over  
1178 time from a 100% to 98 % to acknowledge the potential slowing down of the AMOC.

1179

#### 1180 **Fig 14. Ocean $CO_2$ Sinks ( $S_{OCEAN}$ ) and yearly northern Pacific $G_{ATM}$ compared.**

1181 Regression analysis comparing the ocean sinks ( $S_{OCEAN}$ ) [76] and yearly atmospheric growth  
1182  $G_{ATM}$  as measured at Mauna Loa.

1183

#### 1184 **Fitting the NPCE in the Pleistocene climate cycle**

1185

1186 During a Pleistocene cooling phase outgoing heat is larger than incoming heat and during a  
1187 Pleistocene warming phase incoming heat is larger than outgoing heat. Ozone depletion in the  
1188 stratosphere is known to cause cooling in the stratosphere and warming at the earth surface [55].  
1189 When projecting a sawtooth pattern as suggested by Maslin and Brierly [108] on the four last  
1190 Pleistocene climate cycles [109], each cycle can be typified by slow cooling (approximately  
1191 100,000 years) and rapid warming (approximately 10,000 years), inversely accompanied by a  
1192 slow increase and a rapid decrease in dust concentration (Fig 15). When introducing the of the  
1193 stratospheric ozone cycle in a Pleistocene climate cycle, slow cooling represents slow  
1194 stratospheric ozone accumulation (more accretion than depletion), causing earth surface cooling  
1195 while rapid warming represents rapid stratospheric ozone depletion (more depletion than  
1196 accretion) causing surface warming. The  $CO_2$  concentration in the ocean being directly linked by  
1197 the Le Chatelier principle to water temperature, this means that stratospheric ozone accumulation  
1198 causes oceans to take up more  $CO_2$  than they release, while stratospheric ozone depletion causes  
1199 oceans to release more  $CO_2$  to the atmosphere. For the Pleistocene the focus is on dust and

1200 charcoal (aerosols/ particulates), as both are inversely proportional to CO<sub>2</sub> and temperature [59].  
1201 We examine here how organic compounds (VOC'S) and pollutants that forest emanate or produce  
1202 through (partial) combustion [110] play a part in these shifting ozone balances.

1203

1204 **Fig 15. Sawtooth pattern projected on the four last climate cycles.**

1205 Blue lines: temperature anomaly ( $\Delta t$  in C°), green lines: CO<sub>2</sub> atmospheric concentration (in ppm),  
1206 red lines: dust concentration(in ppm), grey vertical bar: glacial maximum (after Maslin and Bierly  
1207 [108] and Petit et al. [109]).

1208

1209 Earth, in the absence of all its shields, is considered to be cooling. This means that it emits heat,  
1210 which is a heat flow coming from its core [111]. So, in order for the troposphere to momentarily  
1211 have a constant temperature during a glacial maximum or an interglacial maximum, incoming  
1212 temperature has to be slightly higher than the outgoing temperature. This is not detailed any  
1213 further here, because the cooling of the earth is expected to be undetectably slow at the very high  
1214 time resolution we are considering here, but it is added *pro memori*.

1215

1216 **The Pleistocene cooling phase**

1217 Starting at an interglacial maximum and going toward the next glacial maximum, the  
1218 global kinetic imbalance (OUT > IN) is interpreted as driven by photosynthesis taking H<sub>2</sub>O and  
1219 CO<sub>2</sub> out of the troposphere into vegetation and sinks, but leaving O<sub>2</sub> behind (CO<sub>2</sub> + H<sub>2</sub>O =>  
1220 C<sub>6</sub>H<sub>12</sub>O<sub>6</sub> (organic molecule) + O<sub>2</sub>). This leads to an ever more oxygen rich and inflammable  
1221 troposphere. In the early cooling phase, the air is warm and very damp, and forms condensation  
1222 clouds rather than ice clouds, fire is rare. During the early cooling phase, the tropopause, which  
1223 altitude can be seen as thermally defined [112], has just expanded, forest fires have approached  
1224 their minimum and cirrus cover is minimal. This allows for the heat reflected by both the

1225 condensation clouds and the earth surface during the early cooling phase to escape rapidly from  
1226 the troposphere. Because forest fires are scarce, the SCHWAB is weak and little ocean emitted  
1227 gasses accompanied it to the stratosphere. Water, CO<sub>2</sub> and potentially reactive molecules like  
1228 VOC, NO and CO [113, 114] are incorporated in the vegetation and its sinks and the stratospheric  
1229 ozone accumulation at the tropic is faster than its break down at the stratospheric poles. The  
1230 stratospheric ozone layer is thickening, it is warming and the troposphere is cooling.

1231

### 1232 **The glacial maximum: mitigation of cooling**

1233 As the climate cools, vegetation adapts and concentrates in the low latitudes and altitudes  
1234 where gymnosperms occupy a larger proportion of the vegetation cover. While gymnosperms have  
1235 a regeneration cycle through fire, evergreen angiosperms like *Lithocarpus*, peaking in the tropics  
1236 during the glacial maximum, not only have a co-evolution with fire [115] but also distinguish  
1237 themselves by having relatively high concentrations of inflammable alkanes. So as the climate  
1238 cools fire frequency increases due to increasing O<sub>2</sub> concentrations related to the ongoing  
1239 photosynthesis, but also increases due to the intensification of the aridity related to the colder  
1240 climate. CO<sub>2</sub> and H<sub>2</sub>O are have been buried as organic molecules in sinks. The self-lofting forest  
1241 fires inject incompletely combusted organic molecules and other pollutants into the tropopause  
1242 [116] where they form IPN to the cirrus clouds. The cooling to warming threshold is attained  
1243 when the SCHWAB fed by the pollutants from the forest fires, is large enough to cause the storms  
1244 that will incorporate heat into the oceans that will release their gasses. This is when the ocean  
1245 starts degassing, a.o. CO<sub>2</sub>, but much more importantly chlorine (Cl<sub>2</sub>), sulphur dioxide (SO<sub>2</sub>) or  
1246 ammoniac (NH<sub>3</sub>) being much more soluble and reactive with ozone than CO<sub>2</sub>. The spectacular  
1247 reactivity with ozone of gasses like C<sub>2</sub>H<sub>2</sub>/Ar/O<sub>2</sub>, being slightly less soluble in water than CO<sub>2</sub>, do  
1248 need mention in this context. Indeed warming conditions just beyond those allowing for the

1249 present ocean CO<sub>2</sub> emission conditions can result in C<sub>2</sub>H<sub>2</sub>/Ar/O<sub>2</sub> degassing which deplete ozone  
1250 very rapidly [117].

1251 The glacial maximum threshold can consequently be seen to be related to the accumulation of  
1252 gymnosperm and *Lithocarpus* generated volatile organic carbon at the high latitudes during the  
1253 arctic winter causing winter ozone accumulation [57], and in its wake strong ocean gas emissions.  
1254 The intense ocean gas emission results in a slower ozone accretion at the equatorial stratosphere  
1255 then its depletion at the polar stratosphere, this stratospheric ozone thinning introduces the  
1256 Pleistocene warming phase.

1257

### 1258 **The interglacial maximum: mitigation of warming**

1259 As the climate becomes warm and more humid, the vegetation dominated by the fire prone  
1260 gymnosperms and *Lithocarpus* retreats to high altitudes and latitudes giving way to the chiefly  
1261 deciduous angiosperms forests. Natural fire is less frequent in these deciduous forests and  
1262 produces the fewer ice nucleating particle transported to the poles by the lower current of the  
1263 stratospheric Brewer Dobson circulation cell. The troposphere is at its warmest, so it is at its  
1264 highest and as surface expansion is the square of the altitude increase, the tropopause expansion,  
1265 comprising the cirrus clouds expansion has also contributed in a less dense cirrus cloud deck  
1266 allowing solar radiation (heat) to escape the troposphere again. During this phase less ozone  
1267 accumulates in the high latitudes, the SCHWAB is smaller and less ocean gasses reach the  
1268 stratosphere. Stratospheric ozone accretion at the equator is more important and frequent than its  
1269 depletion at the poles. The stratospheric ozone layer becomes thicker and Earth grows back to its  
1270 natural cooling state where vegetation captures and sequesters the water and carbon dioxide that  
1271 had partly warmed Earth.

1272

### 1273 **A vegetation/stratospheric-ozone response system**

1274 In the natural Pleistocene climate cycle, fire frequency determines the quantity and the  
1275 quality of the pollutants. These pollutants are envisioned to have these pollutants determine the  
1276 size and temperature of the SCHWAB, which determines the degree of the storm intensification  
1277 which determines the nature and the quantity of the ocean released gasses that are reacting with  
1278 ozone as they ascend into the stratosphere. Various papers [84, 117, 118 fig. 1, 119] testify to this  
1279 process.

1280 Two CO<sub>2</sub> air masses at two different altitudes and two different latitudes, indicate that it takes  
1281 about one year for the SCHWAB to migrate from the Arctic get to the top of the Hadley cell [85,  
1282 fig.6]. Under natural climatic stasis circumstances, as the interglacial maximum we have reached  
1283 presently should have been, one would not expect the new air mass to already presents itself while  
1284 the old one is still escaping the troposphere.

1285 As explained in the paragraphs on mitigation of cooling and mitigation of warming, the pollutant  
1286 formation in the natural Pleistocene climate cycle reflects a response system with the vegetation  
1287 composition. Cooling circumstances reduce the area of the chiefly deciduous angiosperms while  
1288 favouring the expansion of the conifer forests having a regenerative cycle through fire and thus  
1289 forming the pollutants driving the ozone cycle.

1290 This interpretation of the Keeling curve echoes the response system of Lovelock [12] between  
1291 algal blooms and cloudiness, known as the Gaia theory, but here the drivers are pollutants from  
1292 fires and the response system is between the stratospheric ozone and the vegetation composition.

1293

#### 1294 **A complex stacked function in the Pleistocene**

1295 It is argued here that vegetation has been driving the climatic feedback mechanisms. This  
1296 model is partly complementary to the model of Ellis and Palmer [120] leaning on the albedo of the  
1297 polluted ice intensifying Milankovitch eccentricity cycles. But during the Pleistocene the climate

1298 cycles become slightly longer each time and do not really match the eccentricity cycles but can be  
1299 seen as modulating them (Fig 16).

1300 Quite clearly, when comparing the Milankovitch cycles of Berger and Yin [121] to the past four  
1301 Pleistocene climate cycles [109], obliquity does modulates temperature, as a high obliquity gives  
1302 warming and a low obliquity gives cooling, whereas precession does the contrary but again these  
1303 cycles have periods being divergent from the vegetation driven ones and are superimposed on the  
1304 latter.

1305 Three sunspot related cycles (unnamed ~9.7-ka; 'Heinrich-Bond' ~6.0-ka; Hallstatt ~2.5-ka), have  
1306 also been demonstrated to play a minor but periodic role in Earth's climate forcing during the  
1307 Pleistocene [122], these reflect functions that are also superimposed on the other ones.

1308

1309 **Fig 16. Comparison between the four last climate cycles and the Milankovitch cycles.**

1310 A. Blue: temperature, Green: CO<sub>2</sub>, red: dust. Red shading: high obliquity, blue shading: low  
1311 obliquity (after Petit et al. [109]). B. Milankovitch cyclicities after Berger et al. [121].

1312

1313 The Pleistocene temperature curve, composed of numerous superimposed functions is strongly  
1314 serrated and also requires events of higher and more irregular frequencies to explain the observed  
1315 patterns. Historical evidence of volcanism, being irregular can also be seen to determine the  
1316 temperature and CO<sub>2</sub> patterns from the Pleistocene. Indeed eruptions are known to have an impact  
1317 on climate [9]. Insight in volcanic intensity and frequency supports this [123].

1318 Meteorites produce high concentrations of dust in the stratosphere as was measured with the  
1319 Chelyabinsk meteorite [124]. Meteorites strike Earth's atmosphere annually, but bigger impacts  
1320 occur at century or millennium frequency [125]. Meteorites consequently, can cause unexpected  
1321 serration to the Pleistocene temperature curve and in its wake, the atmospheric CO<sub>2</sub> concentration.

1322 Finally, grazers may have impacted the length of the cooling period by rapidly promoting the fire  
1323 prone grasslands over forested areas [126]. This complex co-evolutionary process offers an  
1324 explanation for the lengthening of the cooling period. Thus grazing known to suppress forest  
1325 growth can be considered to have resulted in the lengthening of the cooling period of each  
1326 Pleistocene climate cycle.

1327

### 1328 **Industry mimicking the Pleistocene warming phase**

1329 What was sketched above is a natural Pleistocene climate cycle, but industrial emissions  
1330 producing pollutants can have the same effect [14]. Considering the influence that the Brewer  
1331 Dobson Cell has on stratospheric gas mixing [127], industrially emitted gasses are considered here  
1332 to either impact ozone indirectly through transport to the poles [46] or through direct hot injection  
1333 in the stratosphere. A relation between ozone depletion and climate is not contested in the IPCC  
1334 reports, but since 2007, the causality of relation has often been inverted in the scientific literature,  
1335 in spite of earlier work [31] demonstrating the covariance between aerosols and lower  
1336 stratospheric ozone depletion.

1337

### 1338 **Where are we now within the Pleistocene climate cycle?**

1339

1340 In the past 140 000 years, there were 110,000 years of cooling and 20,000 years of  
1341 warming [128] while the interglacial maximum had a duration of approximately 10,000 years,  
1342 comprising about 5000 years of climatic stasis and 5000 years probably representing a slow start  
1343 of cooling. The amplitude of the temperature during the Pleistocene has been of 10 to 12 °C, and  
1344 the amplitude of the CO<sub>2</sub> concentration was in the range of 100 to 120 ppm. This indicates that (on  
1345 average!) cooling took place by 0.1 degree Celsius and 1 ppm CO<sub>2</sub> every 1000 years, while  
1346 warming was of approx. 0.5 °C and 5 ppm /1000 years. Warming in the natural cycle was



1347 consequently about 5 times faster than its cooling. During the long Pleistocene cooling phase  
1348 minor uptake imbalance events were only required to occur once every 1000 years. Similarly,  
1349 during the short warming phase ocean CO<sub>2</sub> emission only requires a modest imbalance every 200  
1350 years.

1351 In the past 60 years, ocean emission as measured at Moana Loa by the NOAA was of  
1352 approximately 100 ppm CO<sub>2</sub>. As the recent excursion of the atmospheric CO<sub>2</sub> has the amplitude  
1353 of a glacial cycle, and considering that it took place more than 300 x faster than during a natural  
1354 warming phase, it cannot be said that it is accidental.

1355 In considering the Pleistocene climate cycles (Fig 15), it is clear that we are at the top of an  
1356 interglacial, and in scrutinizing numerous glacial records of oxygen isotope data and pollen  
1357 records at both high and low latitudes [128] there is good reason to consider that the interglacial  
1358 maximum has been passed and the cooling cycle was on its way. It is therefore defensible that  
1359 present aberrantly high atmospheric CO<sub>2</sub> concentrations indicate that anthropogenic  
1360 industriousness has reiterated the mitigation of cooling conditions from the glacial maximum of  
1361 the Pleistocene climate cycle, but now on top of an interglacial maximum, thus bringing us in a  
1362 climatic no-mans land, at a rate of change never experienced before.

1363

1364

## 1365 **Discussion**

1366

1367 Because the above findings are in several aspects in contrast to commonly believed  
1368 hypotheses, it is essential to discuss here the various seemingly inconsistent aspects, viz. the fate  
1369 of industrially emitted CO<sub>2</sub>, the role of light-absorbing particles, the role of GHG, the oceanic pH,  
1370 the role of volcanic material and sea spray, the ozone cycle, the limits of the global carbon pump,

1371 the presently highlighted thermos conditions of planet Earth, the role of the southern hemisphere  
1372 and the fragility of our insights on carbon dioxide emission due to land use change.

1373

1374 **Where did the industrially emitted CO<sub>2</sub> go?**

1375

1376 The present analysis demonstrates that oceans are emitting CO<sub>2</sub> and not taking it up. While  
1377 the G<sub>ATM</sub>, as it is envisioned in the present paper, records the difference between ocean CO<sub>2</sub>  
1378 emission and ocean CO<sub>2</sub> uptake, the IPCC considered it to reflect the Industrial Emission (IE)  
1379 minus the Total Sinks. Within the IPCC hypothesis expressed by the Global Carbon Equation,  
1380 seasonal ocean emission (O<sub>SE</sub>) and uptake (O<sub>SU</sub>) were expected be equal and hence neglectable. In  
1381 Figure 9 we have seen that the duration of ocean CO<sub>2</sub> uptake decreases, which, in the light of a  
1382 paradigm that states that ocean uptake follows directly from increased industrial emission is  
1383 difficult to conceive. Increased yearly uptake for a shorter seasonal period of time is explained  
1384 here, as modelled, as following from an increase in melting speed related to the yearly arctic heat  
1385 anomaly. We also see in Figure 9 that the period of emission increases in duration, this is  
1386 explained here, by higher temperatures of ocean surface water related to the SCHWAB becoming  
1387 larger. When explaining the atmospheric increase in CO<sub>2</sub> concentration, as argued in this paper,  
1388 merely through first a tropospheric local and then a stratospheric global heat anomaly, the  
1389 question arises why is the industrial CO<sub>2</sub> emission not perceived in the Keeling curve?  
1390 This can be explained by the study of El Niño events that cause a CO<sub>2</sub> concentration anomaly in  
1391 the mid troposphere at 6 km altitude [79]. This indicates that the CO<sub>2</sub> produced at sea surface and  
1392 relatively high sea surface temperature, rises rapidly to 6 km. This relatively high sea surface  
1393 temperature represents a anomaly of *less than 1 C°* [129]. Knowing that El Niño years are  
1394 accompanied with a CO<sub>2</sub> increase of 1 ppm four months after the event, and knowing the CO<sub>2</sub>

1395 rises to the mid troposphere, estimated at 6 km, the ascending speed of that CO<sub>2</sub> holding air mass  
1396 is 2 m/ hour.

1397 At 0,7 m from a domestic chimney employing coal, lowest temperatures are 178 C° [130]. The  
1398 exhaust temperature of gas vehicle varies between 150 and 600 C° [131] while the analysis of the  
1399 European waste heat potential summarizes industrial categories that start a 100 C° and end at >  
1400 1000 C° [132]. Industrial and domestic CO<sub>2</sub>, are consequently entering the troposphere at  
1401 “temperature anomalies” of minimally 60 C° and can be expected to rise at 120 m/hour.

1402 Consequently, while the primary condition for the global carbon equation (being that land, ocean  
1403 and atmosphere form a closed system) is incontestably true, this does not necessarily imply that  
1404 land, ocean and troposphere are also a closed system. Indeed, CO<sub>2</sub> at low temperatures anomalies  
1405 already escapes from the troposphere to the stratosphere [85], which means that gasses entering  
1406 the troposphere at minimally 60 x that temperature cannot be expected to accumulate in the  
1407 troposphere, but reach into the stratosphere, while the cirrus clouds sieve out the coarser ozone  
1408 sensitive particulates.

1409 This corroborates the NPCE and indicates that what is measured at Moana Loa is not industrial  
1410 CO<sub>2</sub> accumulation, but a flow of gasses escaping the warming Northern Pacific Ocean. This yearly  
1411 flow of relatively warm air mass is denounced by the daily CO<sub>2</sub> measurements at Mauna Loa. It  
1412 must consequently be concluded that the Keeling curve reflects the acceleration of ocean  
1413 warming, which we showed in our model is proportional to industrial pollutant emission. This  
1414 means that oceans are presently degassing at a speed and acceleration higher than at a  
1415 glacial/interglacial transition, and, in the wake of this acceleration, so does the stratospheric ozone  
1416 depletion causing the ocean emission to accelerate.

1417

1418 **Where are the ocean gasses coming from?**

1419

1420 The formation of coastal and oceanic carbon sinks [133] gives an imbalance in the global carbon  
1421 budget calculated in this way, by sequestration the carbon into shelf and ocean sediments. This  
1422 imbalance can be solved by considering that the mid-ocean ridges have heightened nutrient flows  
1423 when active [134] and constantly feed the oceans with gasses. The gasses are transferred at each  
1424 transition from a glacial to an interglacial maximum from the oceans to the atmosphere when  
1425 natural ozone formation at the poles and its depletion in the stratosphere causes the ocean surface to  
1426 warm and emit the gasses that have been accumulating in the ocean during a Pleistocene cooling  
1427 phase.

1428

### 1429 **Polyatomic molecules and other light-absorbing aggregates**

1430

1431 The words of Lovelock and Margulis [12]: “CO<sub>2</sub> at high concentrations could have  
1432 performed a similar function as could other polyatomic gasses” refer to the fully hypothetical  
1433 models developed for the early Archean, when it was assumed that the young Sun was faint. By  
1434 now the faint Sun paradox is found to be questionable for the early Archean period [135]. The  
1435 higher rotation velocity of the early Earth could have caused Earth to be warmer [136]. These  
1436 decisions, rather than results, pertain to a hypothesized the deep past. We are now focussing on the  
1437 past 2.5 million years where the troposphere is chiefly composed of N<sub>2</sub> and O<sub>2</sub>, and where of all  
1438 trace elements, H<sub>2</sub>O, with an average concentration of 0.4% is the most frequently occurring  
1439 polyatomic molecule in the atmosphere after which CO<sub>2</sub> follows far behind, as it is measured in  
1440 parts per million, and CH<sub>4</sub> is even more scarce as it is measured in parts per billion. Without any  
1441 doubt all electrons within all atoms of these molecules when excited by light and when falling  
1442 back to a lower energy level will, depending on its atomic composition, emit more or less infra red  
1443 light and as such participate to the earth’s kinetic budget, but the most frequent more than other  
1444 compounds. Water vapour is the most frequent and the width of the absorption spectrum of water

1445 vapour clearly encompasses that of other polyatomic molecules in the troposphere. Furthermore  
1446 pollutants also occur frequently in the troposphere, alkenes, for instance, represent an important  
1447 fraction of the particulate matter [137] and they are present in higher concentrations than methane.  
1448 Particulate matter, for instance, displaying high concentrations in the Ganges valley can be  
1449 expected to cause radiative absorption [138]. As explained in the introduction: all solid  
1450 particulates, all gasses, and all liquids cause radiative absorption. The hypothetical choice of CO<sub>2</sub>  
1451 for a debatable problem in the Archean, should be treated with a correct sense of proportionality,  
1452 instead of elevating CO<sub>2</sub> to the chief kinetic driver for System Earth.

1453

### 1454 **Ocean pH**

1455

1456         Given that the oceans, on a yearly basis, take up less CO<sub>2</sub> than that they are releasing, CO<sub>2</sub>  
1457 uptake cannot be causing the oceans to decrease in pH, and consequently the cause for this  
1458 changing pH must be sought for elsewhere. Indeed, the alkalinity changes measured lately in the  
1459 ocean [139] appear to be in contradiction with the present ocean degassing model. To understand  
1460 this change in pH of the oceans, it is important to return to the theoretical basis of the alkalinity of  
1461 solutes. The Le Chatelier principle states that changes in temperature, pressure, volume, or internal  
1462 concentration of a solute, each and all can shift the equilibrium state. The oceans are a chemical  
1463 solute, and neither the pressure nor the volume of the surface water are changing considerably.  
1464 This leaves temperature to determine the pCO<sub>2</sub> and its pH in ocean surface water.

1465

### 1466 **Hawaii as an atypical upwelling area**

1467

1468         The focus on the temperature dependency of pH and pCO<sub>2</sub> in ocean systems, requires a  
1469 new interpretation for the covariance between the atmospheric CO<sub>2</sub> recorded at Mauna Loa

1470 Observatory and the nearby ocean station ALOHA for which the study of Dore et al. [140] set the  
1471 tone. In analysing how the conclusion that “Increasing concentration of atmospheric carbon  
1472 dioxide leads to higher dissolution of carbon dioxide in the ocean, thus lowering the ocean pH”  
1473 [141] came about, we return to the early study [140] that showed atmospheric  $\text{CO}_2$ , ( $G_{\text{ATM}}$ ) and  
1474  $\text{pCO}_2$  in ocean water are inversely proportional but increase both, while the pH is proportional to  
1475 the  $G_{\text{ATM}}$  but decreases [140, fig. 1]. This decreasing pH is interpreted as caused by the uptake of  
1476 atmospheric  $\text{CO}_2$ , but can also be assigned to the intensification of seasonal storms carrying cold  
1477 nutrient rich water to the ocean surface and thus stimulating the calcium carbonate production  
1478 ( $\text{CaCO}_3$ ) which drives the ocean water pH down [142]. When considering the pH at in situ  
1479 temperature [140, fig. 3], we see it decrease from spring equinox to autumn equinox, thus clearly  
1480 reflecting a temperature signal. It consequently is what happens in between autumn equinox and  
1481 spring equinox that is relevant to understand the pH and the  $\text{pCO}_2$  registered at the ALOHA  
1482 ocean site. In this context Calil and Richards [143] show that Hawaii is an atypical situation for  
1483 the whole Northern Pacific as it is influenced by upwelling and that starting September the ocean  
1484 is cooling again. The increasing intensity of such cooling upwelling settings over time has been  
1485 associated to the recent climate conditions [144]. The sea water temperature, the  $\text{pCO}_2$  and the pH  
1486 at the ALOHA ocean site, consequently reflect upwelling intensification while the overall  
1487 atmospheric  $\text{CO}_2$  signal is determined by the Northern Pacific as a whole.

1488

### 1489 **Volcanic particulates and sea spray**

1490

1491 When the Krakatau explosion of 1883 caused global cooling, Abbot and Fowle [9]  
1492 highlighted volcanic glass and sulphur as potential components of the haze that obstructed solar  
1493 radiation through “volcanic glass and sulphur”. The petrology of dust that fell upon the mirror of  
1494 their telescope had indicated the presence of “volcanic glass, quartz, feldspar and kaolin” but this

1495 composition raised some doubt at the time as to its origin, which sulphur did not. By now [e.g.,  
1496 145] such particulates are known to be normal volcanic products and excellent ice nucleating  
1497 particles [146], which again, considering the altitude of the event, does not mean that the  
1498 temporary cooling came from those particulates. There is consequently no reason to associate  
1499 cooling to sulphur clouds, neither in the tropo- or the stratosphere as has since then been  
1500 propagated. Sea spray containing sulphur [23] is typical of the storm of a Pleistocene warming  
1501 phase, while sulphur in the natural stratospheric presence of hydrochloric acid will deplete ozone  
1502 [103].

1503 Volcanoes are all different, depending on their genesis and their position relative to the oceans.  
1504 Pering et al. [147], for instance, demonstrate that, for the La Fossa crater, Vulcano Island, and the  
1505 North East Crater of Mt. Etna, the sulphur production varied between 4% and 1% of their water  
1506 production. Volcanoes are also producing ash clouds, the composition of which is primarily  
1507 siliceous [148]. Volcanoes, injecting both water and ash in the atmosphere, are the way, par  
1508 excellence, to form ice nucleating particles (INP) [149]. In the light of water as a source of free  
1509 radicals or chlorine hydrates, it is crucial not to be tempted to jump to the conclusion that water  
1510 vapour will result in our long awaited high tech solution to the mitigation of warming, on the  
1511 contrary.

1512 In a comparable way marine cloud brightening (MCB) through dimethyl sulphides from sea spray  
1513 will only dim or brighten Earth's surface as cloud condensation nuclei (CCN), but have no  
1514 chemical cooling effect. In fact effort have even been made at reducing sulphur emission [150].  
1515 So, it cannot really be the dimethyl sulphides that will be dimming solar radiation, but rather it is  
1516 the altitude at which such particle resides, given their density and origin, that will determine its  
1517 role in our climate cycle. All particulates will come to increase the density of either of the walls of  
1518 our thermos-sphere (the condensation cloud or the ice cloud wall) and dimethyl sulphides can be

1519 seen as depleting ozone [27], while our knowledge of the stratosphere composition, its chemical  
1520 behaviour and its vertical circulation patterns are insufficient to try chemical experiments with it.

1521

## 1522 **The ozone cycle**

1523

1524 Despite the 1987 Montreal protocol aiming at the global reduction of the infamous CFCs, a  
1525 global increase in atmospheric concentrations of CFC-11 or CF6 was observed in the last decade  
1526 [151]. Moreover, there is convincing evidence for the direct participation of polar stratospheric  
1527 cirrus in the polar stratospheric ozone depletion via sunlight-giving photochemical reactions in  
1528 spring, for which notably hydrogen oxide radicals ( $\text{HO}_x$ ), nitrogen oxides ( $\text{NO}_x$ ) and sulphur are  
1529 responsible [152]. This is in strong support of satellite-observed warming pattern that resembles  
1530 closely the atmospheric distribution of chlorofluorocarbons (CFCs), as highlighted in a review on  
1531 climate models and radiative forcing [153]. Finally as mentioned above, ozone reacts with free  
1532 radicals and water in the damp phase can dissociate into H and OH, both being very reactive free  
1533 radicals (Pers. Comms. Prof. Wijbrans, July 11-th 2024). Considering the importance such free  
1534 radicals have for our climate, and seeing how faithfully the emission over time of for instance SF6  
1535 follows the fluctuations we also observe for the ocean emitted  $\text{CO}_2$ , it is of crucial importance that  
1536 we learn to separate between natural and industrial occurrences of free radicals.

1537 Asymmetric ozone distribution patterns comparable to those for  $\text{CO}_2$  or  $\text{CH}_4$  [84, fig.6] in the  
1538 troposphere can be witnessed in Bednarz et al. [118, fig. 2]. In analysing the tropospheric  
1539 behaviour of ozone [118, fig. 1], we see it appears in the winter in the northern high latitudes,  
1540 expand in spring and retrieve in summer. Ozone even react with water vapour [112], thus  
1541 supporting our new model where a self-lofting chemically reactive arctic winter night generated  
1542 air bubble (SCHWAB) appears early in the year in the northern high latitudes, and comes to  
1543 intensify the existing atmospheric interconnectivity, by lowering the Aleutian Low and



1544 heightening the Hawaiian High. The SCHWAB is influenced by, and influencing the existing  
1545 pressure systems, to finally escape to the stratosphere at the equator with the ascending currents of  
1546 the Hadley cell.

1547 Lightning frequency has increased drastically in the Arctic over the past decade [154]. This  
1548 increase follows from higher aerosol concentrations in the atmosphere [155]. The predominant  
1549 mechanism of ozone formation in the upper troposphere is lightning-induced precursors such as  
1550 oxides or pollutants like nitrogen ( $\text{NO}_x$ ), carbon monoxide (CO), and hydrocarbons (HC) [82].  
1551 While lightning frequency is highest in the arctic summer [156], one should keep in mind that  
1552 ozone depletion requires bromides, which only become available as soon as (sea) ice starts melting  
1553 [157]. Lightning in the absence of bromides can explain the winter accumulation of the short  
1554 lasting GHG ozone in the arctic winter night, as it is sensitive to photolysis requiring light.

1555 The intensification of industrial and agricultural dust and pollutants like for instance the volatile  
1556 organic carbons, or isoprene produced by the palm oil industry [158] can be expected to intensify  
1557 the formation of the SCHWAB over time. It explains why the period of ocean  $\text{CO}_2$  emission  
1558 became longer than the period of ocean  $\text{CO}_2$  uptake, starting a month earlier, and ending a month  
1559 later, thus indicating that the SCHWAB became larger and more sensitive to light, with time.

1560

### 1561 **Funnel shaped hemisphere**

1562

1563 The lower current of the stratospheric Brewer Dobson circulations concentrates cirrus clouds each  
1564 year at the pole of the hemispheric winter [46]. These cirrus clouds are formed by ice nucleating  
1565 particles of diverse industrial origin [47] and can be envisioned to get transported in this way to  
1566 the high latitudes. This lower current functions like a funnel, each time occupying a smaller  
1567 latitudinal circumference. This means that it concentrates all pollutants that reached freezing  
1568 altitudes towards the poles, which poles being the coldest heaviest atmospheric point on earth will

1569 attract these pollutant laden ice clouds . This explains why the UKESM1 emission [77], reflecting  
1570 industrial activities distributed randomly over the northern hemisphere are giving good results in  
1571 correlation with the  $G_{ATM}$  measured at Mauna Loa.

1572

### 1573 **The limits of the biological carbon pump**

1574

1575 The modelling of an algal bloom driven System Earth, as is the Gaia hypothesis, was  
1576 unsuccessful [22]. Nonetheless phytoplankton is still considered as a means for the restoration of  
1577 the natural carbon cycle by drawing atmospheric carbon down to the ocean sinks through the  
1578 biological pump [e.g., 159]. Many authors still expect that boosting phytoplankton production  
1579 through Fe seeding will draw down the  $CO_2$  concentration in the atmosphere and mitigate  
1580 warming, but it has often been tried to validate this hypothesis, though without success [e.g., 160].  
1581 The Banda Sea upwelling area sheds light on this vexing situation [161]. Parts of these  
1582 phytoplankton blooms do reach the ocean bottom as faecal pellets holding both inorganic  $CaCO_3$   
1583 [162] and organic carbon of zooplanktonic origin. The Banda Sea is such an upwelling area where  
1584 cold and nutrient-rich upwelling water causes organic carbon drawdown through a more elaborate  
1585 food chain favouring the sedimentation of the refractory (= resistant to decay) zooplankton  
1586 exoskeletons and eggs. These upwelling waters are feeding each consecutive trophic level with a  
1587 specific speed and duration adapted to the life cycle of these species. Any imbalance in these  
1588 equilibria developed over time will have a so-called harmonica effect. Nutrient supply and the  
1589 intercorrelated pH and temperature of marine systems are even responsible for the Harmful Algal  
1590 Blooms (HAB) [163] inducing the  $CO_2$  transport into the troposphere.

1591

### 1592 **Albedo or chemically driven global warming**

1593

1594 In 2023, direct negative forcing through aerosols raised alarm. Aerosols from aviation  
1595 were even considered to mask part of the GHG global warming [50]. This suggested that the air  
1596 transport industry could mitigate global warming in a similar way as Marine Cloud Brightening  
1597 (MCB) would do. But clouds reflect on both sides and most infrared radiation comes from the  
1598 earth surface which means that the denser the clouds, the more intense the reflection. Moreover,  
1599 pollutants from (air transport) industry participate in both lightening intensification and in the  
1600 shedding of molecules participating in the chemical formation of the winter ozone accumulation.  
1601 In its latest report [49] the IPCC supported the view that particulates induced albedo could mask  
1602 the effect of GHG [50] and prepared to lean towards geoengineering, but, in contrast with studies  
1603 from the 1970 [14], they only focussed on clouds albedo and neglected the chemically induced  
1604 arctic heat anomaly.

1605

### 1606 **Southern hemisphere**

1607

1608 The southern hemisphere displays the same atmospheric increase in CO<sub>2</sub> concentration as  
1609 the northern hemisphere but only experiences mild seasonal CO<sub>2</sub> emission fluctuations. This is  
1610 explained here because industrialisation is less prolific in the southern hemisphere and has less  
1611 influence on tropospheric ozone. Furthermore, the Antarctic consists of a continent which, with its  
1612 low thermal mass, is slow at warming up. The heat anomaly from the winter ozone accumulation  
1613 is consequently chiefly occurring in the northern hemisphere, while the heat anomaly following  
1614 from ozone depletion in the stratosphere is experienced globally.

1615

### 1616 **Land use change emission, land use change surface**

1617

1618           The relation between  $G_{ATM}$  and land use change was established above in the results  
1619 through the new hypothesis that land use change surface ( $S_{LUC}$ ) increase is responsible for an  
1620 increase in the production of aerosols and pollutants. Here we see (Fig 17) that land sinks ( $E_{LUC}$ )  
1621 increase with increasing land use change surface.  $E_{LUC}$  instead of emitting  $CO_2$  to the troposphere,  
1622 is found here to form a sink, which is explained here as being transported to hydrosphere through  
1623 consumption and thus disappearing from the troposphere. This explains, for instance, the extreme  
1624 eutrophication of the North Sea [e.g. 164] .

1625

1626   **Fig 17. Trend analyses between land sinks ( $S_{LAND}$ ) and land use change emission (  $E_{LUC}$ ) or**  
1627 **surface of land use change ( $S_{LUC}$ ).**

1628

1629           The difficulties at separating between land  $CO_2$  sinks ( $S_{LAND}$ ) and land use change  $CO_2$   
1630 emission ( $E_{LUC}$ ) have been reported by their developers [74, 75]. The evaluation of these two  
1631 parameters is more complex than commonly believed. Indeed, Hakkarainen et al. [80, fig. 3] show  
1632 that the tropical rainforests belts are emitting  $CO_2$ . In the most recent publications of June 2024 of  
1633 a the European Centre for Medium-Range Weather Forecasts [165] we also see very clearly that,  
1634 on land, in the summer there is  $CO_2$  emission in the southern hemisphere. This is also true when  
1635 looking at the areas where the Brazil tropical forest did not undergo burning and cutting [166] and  
1636 this holds also for Africa [167]. The response of tropical trees to rising temperatures is a key  
1637 uncertainty limiting our ability to predict  $CO_2$  biosphere-to-atmosphere feedbacks in a warming  
1638 world. Indeed tropical trees show that the ratio between photosynthesis and respiration intensity is  
1639 affected by high temperatures for some taxa more than for others [168]. These considerations can  
1640 explain the paradoxical correlations observed in our statistical analysis of the relation between  
1641 carbon sinks and sources on land. Alternatively, the  $CO_2$  emission of tropical forests [80, fig.3]

1642 can also be explained by fungi in soils because fungal respiration rates also increase at increasing  
1643 temperatures [169].

1644 Present paradoxical considerations are mentioned here to urge a reassessment of the impact  
1645 of agriculture on CO<sub>2</sub> but even more so on organic aerosols. Indeed, agriculture according to  
1646 present analysis is expected to result in more aerosols, volatile organic carbons (pollutants), and as  
1647 demonstrated here, more ocean CO<sub>2</sub> emission. The use of natural land sinks to store CO<sub>2</sub> is  
1648 welcomed of course, but it is advised to pay attention to the taxa used to these purposes and the  
1649 latitudes where this is done. Also respiration of forest soils requires further study.

1650

### 1651 **Stratospheric ozone depletion reflected in ocean CO<sub>2</sub> emission only**

1652

1653 In the present reconstruction of Keeling curve (Fig 12, 13, Appendix 10) the influence of  
1654 stratospheric ozone depletion was only made to influence ocean CO<sub>2</sub> emission ( $E_{SOD}$ ) but not the  
1655 increase in uptake through the solubility pump. It is therefore highlighted here that present study  
1656 reflects the northern hemisphere until 2019 where the Arctic was hardly affected by stratospheric  
1657 ozone depletion, in mild contrast to the Northern Pacific Ocean at Maua Loa [170, fig. 11 a] which  
1658 was affected by it. Recent results even indicate that stratospheric ozone depletion is most  
1659 conspicuous in the tropics [171, fig. 4 e] while it is absent in the high northern and moderate in the  
1660 high southern latitudes. This negative ozone accumulation anomaly in the stratospheric tropics is  
1661 in support of present Pleistocene climate model where the ozone cycle starts in the high  
1662 tropospheric latitudes, comes to feed the low latitude stratospheric ozone layer and gets depleted  
1663 in the high stratospheric latitudes.

1664

### 1665 **Angiosperms taking the floor**

1666

1667            Since their diversification in the Cretaceous, the angiosperms are considered to have  
1668 outcompeted the gymnosperms that retrieved to colder areas [172]. This overtaking is  
1669 accompanied with cooling [172]. The cooling from the Mesozoic to the Cenozoic is typically  
1670 associated with a decrease in atmospheric CO<sub>2</sub> concentrations, accompanied by an increase in  
1671 amplitude and possibly periodicity [173, fig1]. It is suggested here that our Pleistocene vegetation  
1672 driven cyclicality in temperature and atmospheric CO<sub>2</sub> concentration, reflects the end member of co-  
1673 evolution between vegetation composition and the troposphere that started in the Mesozoic. This  
1674 would imply that the very fast and even accelerating capacity of angiosperms at drawing down  
1675 both carbon dioxide and water from the troposphere has had bearings on the stratospheric ozone  
1676 cycle, allowing for more ozone accretion at the stratospheric low latitudes than depletion at the  
1677 stratospheric high latitudes.

1678

#### 1679 **Limitations of this study**

1680

1681            The present study demonstrates that the Keeling curve is a symptom of temperature driven  
1682 System Earth, which temperature is determined by a cocktail of industrial pollutants injected in  
1683 the tropopause. The pollutant values taken from Griffiths et al. [77] had a very low resolution  
1684 which required that the monthly values were made linear between two consecutive years  
1685 (Appendix 10), such important parameters should be more detailed and provided by industry on a  
1686 regular base. Also, in spite of Griffiths et al. [77] looking far into the future, predictions aren't  
1687 developed in the present study, as pollutant production depends on a Montreal protocol type of  
1688 regulations which feasibility belongs to the political domain which is far beyond the scope of this  
1689 paper. Also the bulk chemical transformation taking place in the arctic winter night was only  
1690 touched upon, but requires further investigation.

1691 In the present study the contribution of the biological pump to the solubility pump is potentially  
1692 underestimated and deserves further scrutiny.

1693 With respect to the stratospheric ozone depletion, the turnover speed of the polluted Brewer  
1694 Dobson Cell will also contribute to the thickness of the stratospheric ozone layer and influence  
1695 earth surface irradiance, but this was not considered here either.

1696 So while this study is supported by good statistical results, it is far from perfect and its chief  
1697 quality is that it brings our knowledge on the strato-, tropo-, hydro- and biosphere together in one  
1698 model for the Northern Pacific Ocean. The Atlantic ocean, the southern Pacific , the southern  
1699 hemisphere, the Canadian Rockies, New Zealand, all function differently as in these divergent  
1700 sites vegetation and lithosphere contribute differently to the atmospheric CO<sub>2</sub> concentration. The  
1701 only thing that all these localities share is the global warming following from the stratospheric  
1702 ozone depletion.

1703 The ideal study consists of a full kinetic global budget for System Earth. In such a kinetic study,  
1704 the turnover speed of the conveyer belt and all its tributaries is required. Estimates in the  
1705 contribution of the formation of Pacific mid and deep water to the global kinetic ocean budget are  
1706 still tentative and divergent. Also, complex functions describing the statistical presence and  
1707 absence of clouds reflecting the solar radiation allowed by the thickness of the stratospheric ozone  
1708 layer are required to make a more detailed kinetic model, but this has not been developed for  
1709 System Earth yet.

1710

1711

## 1712 **Conclusions**

1713

1714 Our introduction aimed at being exhaustive with respect to all the hypothesis brought to  
1715 the scientific arena until now and has made it plausible that a CO<sub>2</sub> driven climate model has been

1716 favoured prematurely over other models, like the pollutant-stratospheric ozone driven climate  
1717 model highlighted here. Independently, these early decisions initiated the international UN  
1718 scientific co-operation that enabled the present study through the wealth of data on which this  
1719 paper leans heavily.

1720 The atmospheric increase in CO<sub>2</sub> concentration ( $G_{ATM}$ ), commonly believed to reflect the  
1721 tropospheric accumulation of industrial CO<sub>2</sub> that was not taken up by the oceans and land sinks as  
1722 expressed in the Global Carbon Equation of the WCRP [42] is demonstrated here to reflect the  
1723 difference between the Northern Pacific Ocean CO<sub>2</sub> emission and its uptake. By replacing  $G_{ATM}$   
1724 with the difference between the northern Pacific CO<sub>2</sub> emission and uptake, the Global Carbon  
1725 Equation proves to be out of balance by 0.395 Gt C. The further analysis of this imbalance  
1726 indicates that it reflects the difference between Northern Pacific Ocean seasonal emission and its  
1727 uptake, as interpreted by the World Climate Research Program. Although minimal, it does  
1728 represent 25% of the average emission CO<sub>2</sub> through land use change and 15 % of the average CO<sub>2</sub>  
1729 uptake of land sinks. This imbalance chiefly reflects the difficulties at estimating the CO<sub>2</sub> land  
1730 sinks and land sources. The complexities in establishing the CO<sub>2</sub> emission through land use  
1731 change ( $E_{LUC}$ ) and land sinks ( $S_{LAND}$ ) are shown to be of fundamental nature and are significant  
1732 enough to urge new studies.

1733 Because the increase in atmospheric CO<sub>2</sub> as recorded in the Keeling curve is fully explained by  
1734 the northern Pacific emission and uptake in CO<sub>2</sub>, industrial CO<sub>2</sub> emission is not detected in the  
1735 troposphere. This absence is explained by the heat of industrial emission causing CO<sub>2</sub> to self-loft  
1736 out of the troposphere. In the absence of CO<sub>2</sub> as the GHG causing global warming, industrial  
1737 pollution is considered as a driver for it, because pollutants are expected to be sieved out of the  
1738 self-lofting industrial emission by forming ice clouds (cirrus) in the tropopause. The pollutants  
1739 forming the ice clouds are brought to the Arctic by the lower current of the stratospheric Brewer  
1740 Dobson cell. In the arctic winter night, these chemically diverse particulates often transform into



1741 the short lived GHG ozone, which in the absence of light allowing for photolysis, accumulates.  
1742 This ozone rich chemically reactive mixture becomes self-lofting as soon as it gets illuminated in  
1743 early spring.  
1744 The Keeling curve was hypothesized here to reflect the yearly ascend of the self-lofting  
1745 chemically reactive arctic winter generated air bubble (SCHWAB) at Mauna Loa , but also its  
1746 ozone depleting consequences in the stratosphere. To help the readers imaging, we envision a lava  
1747 lamp principle where the trajectory of the SCHWAB is influenced by, and influencing the existing  
1748 pressure systems, to finally escape to the stratosphere at the equator.  
1749 This was tested by modelling, under the following assumptions: (1) CO<sub>2</sub> uptake depends on  
1750 seasonal melting, (2) CO<sub>2</sub> uptake also results from enhanced melting through ozone accumulation  
1751 in the arctic winter night, (3) seasonal CO<sub>2</sub> uptake is also related to the intensity of La Niña  
1752 (biological pump), (4) seasonal ocean CO<sub>2</sub> emission takes place in the Northern Pacific Ocean, (5)  
1753 the Northern Pacific Ocean CO<sub>2</sub> emission is also influenced by the incorporation arctic heat  
1754 anomaly into the Northern Pacific, and (6) CO<sub>2</sub> emission is also caused by the buffering of the  
1755 SCHWAB by the ozone in the stratosphere, resulting in global stratospheric cooling and global  
1756 tropospheric warming. The cumulative result of this modelling exercise was found to reproduce  
1757 the Keeling curve.  
1758 The present societal focus on industrial CO<sub>2</sub> emission is an attempt to deal with a symptom. While  
1759 aerosols are being presented lately as causing cooling and thus masking the true GHG warming, it  
1760 is demonstrated here that the chemical warming effect of these aerosols widely surpasses their  
1761 effect as cloud brightening seeds. This is especially true as the troposphere with an outer solid ice  
1762 cloud shield and an inner condensation cloud shield functions as a thermos-troposphere.  
1763  
1764 Present new hypothesis, where climate change is driven by industrial pollutant emission, is  
1765 corroborated by a fire dust-driven Pleistocene response system between vegetation composition

1766 and the stratospheric ozone layer, thus reminiscent of the Gaia hypothesis of Lovelock [12]. The  
1767 cooling phase of the Pleistocene climate cycle, going from the interglacial to the glacial  
1768 maximum, is mitigated by the increasing frequency of (gymnospermous) forest fires producing  
1769 natural pollutants that result in a natural SCHWAB unleashing a broad spectrum of ocean emitted  
1770 gasses depleting the ozone that accumulated in the lower stratosphere during the Pleistocene  
1771 cooling phase. This causes the stratospheric ozone layer to become thinner, and results in  
1772 stratospheric cooling and tropospheric warming. The warming phase, from the glacial to the  
1773 interglacial maximum, is mitigated by the return to warm and humid circumstances reducing the  
1774 fire frequency by restoring the dominance of the vegetations rich in deciduous (chiefly  
1775 angiosperm) trees over the vegetations having a fire regenerating cycle (chiefly gymnosperms).  
1776 This shift in vegetation type ratios allows for the SCHWAB to ascend to the stratosphere without  
1777 unleashing a broad spectrum of ocean emitted gasses, thus resulting in more ozone accumulation  
1778 in the stratospheric tropics than ozone depletion in the stratospheric poles. This causes the ozone  
1779 layer to become thicker, and results in stratospheric warming and tropospheric cooling.

1780

1781 The present study can be seen to build onto the introductory sentence of the first (1990)  
1782 IPCC report [34] in the following way:

1783 “We are certain of the following: there is a natural greenhouse effect which already keeps the  
1784 Earth warmer than it would otherwise be; pollutant, volatile organic carbons in particular,  
1785 (organic) dust, black carbon and particulate matter emissions resulting from human activities are  
1786 substantially causing a yearly arctic heat anomaly increasing the atmospheric concentrations of the  
1787 greenhouse gasses: carbon dioxide, methane, chlorofluorocarbons (CFCs), nitrous oxide, water  
1788 vapour, and others gasses which solubility at increasing water temperature is higher than that of  
1789 CO<sub>2</sub>. These increases reduce the greenhouse effect in the lower stratosphere resulting on average  
1790 in an additional warming of the Earth's surface.”

1791

1792

### 1793 **Acknowledgements**

1794

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2296

## 2297 **Figure captions**

2298 **Fig 1. Examples of erroneous historical validation of the CO<sub>2</sub> driven climate model for**  
2299 **system earth.** Visualization of the table of Foote [2] where her error margin is highlighted as two

2300 colours because she measured temperature after 2 or 3 minutes. Some experiments were shorter  
2301 than other ones and this analysis gives the rates of change for identical duration and identical  
2302 measuring intervals. It is not possible to evaluate how damp her damp air was. Moreover, it seems  
2303 difficult to obtain 100% pure carbonic acid gas as it is very unstable.

2304 **Fig 2. Details of the Keeling illustrating two ways to calculate the yearly atmospheric**  
2305 **growth ( $G_{ATM}$ ).** (a) Method used herein :  $G_{ATM} = ML_{YI} - ML_{YD}$ . (b) method used by the WCRP  
2306  $G_{ATM} = \Delta C / \Delta t$ .

2307 **Fig 3. Visualization of the difference in timing and yearly amplitude between the**  
2308 **atmospheric CO<sub>2</sub> concentrations in different geographical localities for the randomly chosen**  
2309 **1992-1996 period.** (a) Mauna Loa (red line), (b) the South Pole (black line). The data are from  
2310 NOAA [79] (Appendix 5).

2311 **Fig 4. Polynomial function drawn through the monthly atmospheric CO<sub>2</sub> concentrations as**  
2312 **recorded by NOAA on the South Pole between 1980 and 2022 [79].** This function serves to  
2313 establish the average seasonal amplitude in atmospheric CO<sub>2</sub> emission of the south pole by  
2314 subtracting the monthly average from the mean determined by the function  $y = 8.784 * x^2 E-07 +$   
2315  $102 * x + 331.9$ .

2316 **Fig 5. Box plots allowing for the comparison of six potential methods to calculate the**  
2317 **atmospheric increase in CO<sub>2</sub> concentration.** (a) Here 1y: method used herein ( $G_{ATM} = ML_{YI} -$   
2318  $ML_{YD}$ ) for one year; (b) Here 3 y: method used here in for the average of 3 years; (c) May 1 y:  
2319 method used by the WCRP ( $G_{ATM} = \Delta C / \Delta t$ ), from May to May of the former year; (d) Sept. 1y:  
2320 method used by the WCRP ( $G_{ATM} = \Delta C / \Delta t$ ), from September to September from the former year;  
2321 (e) May 3y: method used by the WCRP ( $G_{ATM} = \Delta C / \Delta t$ ), from May to May of three years before;  
2322 (d) Sept. 3y: method used by the WCRP ( $G_{ATM} = \Delta C / \Delta t$ ), from September to September from the  
2323 three years before. Black dashed rectangle: largest spread; blue dashed rectangle: largest  
2324 interquartile; red dashed rectangle: lowest and highest medians.



2325 **Fig 6. Superimposed box and jitter plots for equation (2).**  $G_{ATM}$  calculated as  $IE-S_{TOT}$ , as  $G_{ATM-}$   
2326  $_{Here}$  ( $=ML_{YI}-ML_{YD}$ ) for one and three year, measured as  $G_{ATM-WCRP} = \Delta C/\Delta t$ , for the month of May  
2327 and September for one and three years, respectively; the difference X between the six calculations  
2328 methods for  $G_{ATM}$  based on the measurements at Mauna Loa and  $G_{ATM}$  calculated as  $IE-S_{TOT}$ ; red  
2329 dashes line: Median of  $G_{ATM}$  as calculated here for one year.

2330 **Fig 7. Box plots illustrating that X, the imbalance in the GCE, is positive and reflects the**  
2331 **difference between seasonal ocean CO<sub>2</sub> emission and seasonal CO<sub>2</sub> uptake.**  $ML_{YI}$ : yearly  
2332 atmospheric increase in CO<sub>2</sub> (Fig 2),  $ML_{YD}$  yearly atmospheric decrease in CO<sub>2</sub> (Fig 2),  $IE$ : Total  
2333 of the industrial emission ( $E_{FOS} + E_{LUC}$ ),  $S_{TOT}$ : total of the CO<sub>2</sub> sinks ( $S_{OCEAN} + S_{LAND}$ ),  $O_{SE}$ :  
2334 seasonal ocean emission,  $O_{SU}$ : seasonal ocean CO<sub>2</sub> uptake,  $G_{ATM}$ :  $G_{ATM-Here}$  1y.

2335 **Fig 8. Counterintuitive relation between atmospheric CO<sub>2</sub> growth ( $G_{ATM}$ ) and land CO<sub>2</sub>**  
2336 **sinks and sources.** A. Correlation between atmospheric growth ( $G_{ATM}$ ) and land use change  
2337 emission in CO<sub>2</sub> ( $E_{LUC}$ ), B. Correlation between atmospheric growth ( $G_{ATM}$ ) and CO<sub>2</sub> land sinks  
2338 ( $S_{LAND}$ ), C. Correlation between atmospheric growth ( $G_{ATM}$ ) and the surface area of land use  
2339 change ( $S_{LUC}$ ).

2340 **Fig 9. Ocean CO<sub>2</sub> uptake and emission.** A. Duration of ocean CO<sub>2</sub> uptake (red) and CO<sub>2</sub>  
2341 emission (black) compared, B. Yearly ocean CO<sub>2</sub> uptake (black) and yearly ocean CO<sub>2</sub> emission  
2342 (red) compared, C. Monthly ocean CO<sub>2</sub> uptake (black) and monthly ocean CO<sub>2</sub> emission (red)  
2343 compared.

2344 **Fig 10. Comparison between the trend in monthly increase in CO<sub>2</sub> emission and increase in**  
2345 **CO<sub>2</sub> uptake over time.** Here we see that monthly emission increases faster than monthly uptake.

2346 **Fig 11. Analysis of how the emission season became longer than the uptake season, for 60**  
2347 **years, each time considering 2 decades.** A. transition from emission to uptake moves to a later  
2348 period, B. transition from uptake to emission moves to an earlier period.

2349 **Fig 12. Details of the 6 functions that determine the uptake and emission of CO<sub>2</sub> by the**  
2350 **northern Pacific Ocean and their sum ( $\sum_{1-6}$ ).** (1) Ocean CO<sub>2</sub> uptake following from an increase  
2351 in ocean water solubility due to ocean cooling related to seasonal ice melt ( $U_{SIM}$ ), (2) increase in  
2352 ocean CO<sub>2</sub> uptake following from an increase in ocean water solubility due to ocean cooling  
2353 related to seasonal melt driven by the short lived GHG ( $O_3$ ) ( $U_{SLGHG}$ ), (3) ocean CO<sub>2</sub> uptake  
2354 through biological draw down ( $U_{BDD}$ ), (4) ocean CO<sub>2</sub> emission following from a decrease in ocean  
2355 water solubility due to seasonal ocean warming ( $E_{SOW}$ ), (5) increase in ocean CO<sub>2</sub> emission  
2356 following from a decrease in ocean water solubility due to ocean warming related to seasonal  
2357 tropospheric heat anomaly stemming from the short lived GHG ( $O_3$ ) accumulation in the Arctic  
2358 ( $E_{SLGHG}$ ), (6) Purple: CO<sub>2</sub> emission from cumulative ocean heat uptake throughout the year  
2359 because of ozone depletion in the stratosphere, ( $E_{SOD}$ ), (7) Apple green: monthly sum of all these  
2360 functions (Sum 1-6).

2361 **Fig 13. Comparing the monthly reconstruction of the Northern Pacific Carbon Equation**  
2362 **(NPCE) based on the assumption it is pollutant driven, and the monthly atmospheric CO<sub>2</sub>**  
2363 **concentrations recorded by NOAA at Mauna Loa.**

2364 A. Function (black) expressing the reconstructed pollutant driven North Pacific CO<sub>2</sub> emission  
2365 behaviour ( $((E_{SOD} + E_{SLGHG} + E_{SOW}) - (U_{SLGHG} + U_{BDD} + U_{SM})) + 320 \text{ ppm}$ ) compared to (red) the  
2366 monthly recording of the CO<sub>2</sub> concentration at Moana Loa as placed online by the NOAA [72]. B.  
2367 The regression analysis comparing both functions. Details of the functions are given in Appendix  
2368 10. The various components of the NPCE are explained further below.

2369 **Fig 14. Ocean CO<sub>2</sub> Sinks ( $S_{OCEAN}$ ) and yearly northern Pacific  $G_{ATM}$  compared.**

2370 Regression analysis comparing the ocean sinks ( $S_{OCEAN}$ ) [76] and yearly atmospheric growth  
2371  $G_{ATM}$  as measured at Mauna Loa.

2372 **Fig 15. Sawtooth pattern projected on the four last climate cycles.** Blue lines: temperature  
2373 anomaly ( $\Delta t$  in  $C^\circ$ ), green lines:  $CO_2$  atmospheric concentration (in ppm), red lines: dust  
2374 concentration(in ppm).

2375 **Fig 16. Comparison between the four last climate cycles and the Milankovitch cycles.** Green:  
2376  $CO_2$ , blue: temperature, red: dust, red shading: high obliquity, blue shading: low obliquity.

2377 **Fig 17. Trend analyses between land sinks ( $S_{LAND}$ ) and land use change  $CO_2$  emission ( $E_{LUC}$ )**  
2378 **or surface of land use change ( $S_{LUC}$ ).**

2379

## 2380 APPENDICES

2381 **Appendix 1** Intro The table of Foote serving for Fig 1

2382 **Appendix 2** Intro Table of Arrhenius analysed

2383 **Appendix 3** M&M Explaining how G ATM was calculated in here (Serving for Fig 2)

2384 **Appendix 4** M&M G ATM from ppm to Gt C and the duration of emission and uptake in month

2385 **Appendix 5** M&M Monthly average atmospheric  $CO_2$  concentration as measured by on the South  
2386 Pole compared to ML and the regression analysis smoothing the curve.

2387 **Appendix 6** M&M Calculation of the average seasonnal amplitude in the southern hemisphere

2388 **Appendix 7** M&M Sink and source of  $CO_2$  as found in literature plus pollutants and surface of  
2389 land use change

2390 **Appendix 8** M&M Regression analysis with the informartion for the Tables 1 and 4 on top.

2391 **Appendix 9** M&M Six calculation methods for atmospheric growth compared.

2392 **Appendix 10** M&M Reconstruction of the Keeling explained.

2393 **Appendix 11** M&M Average duration of a day length Northern Hemisphere.

2394

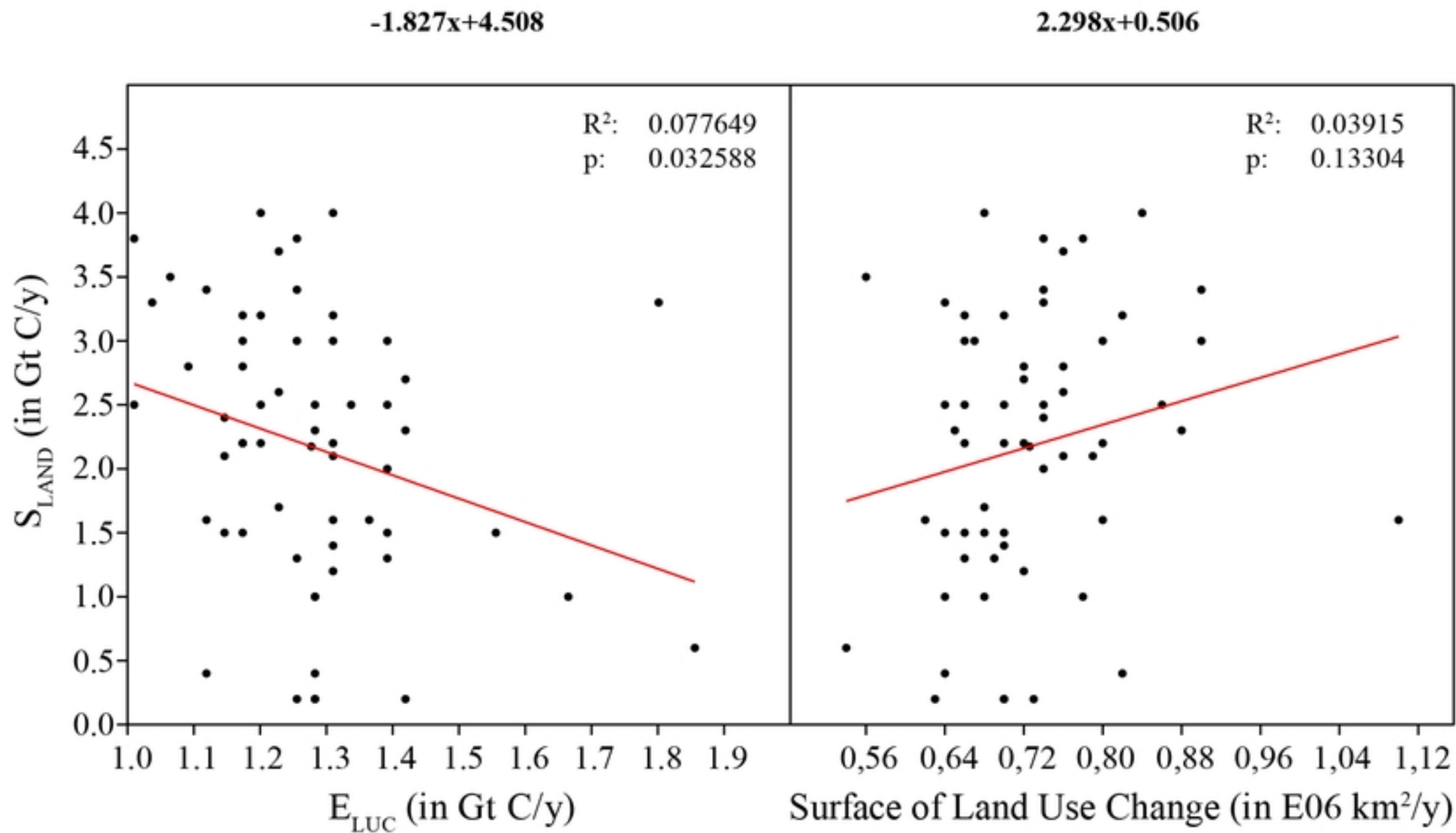


Fig 17



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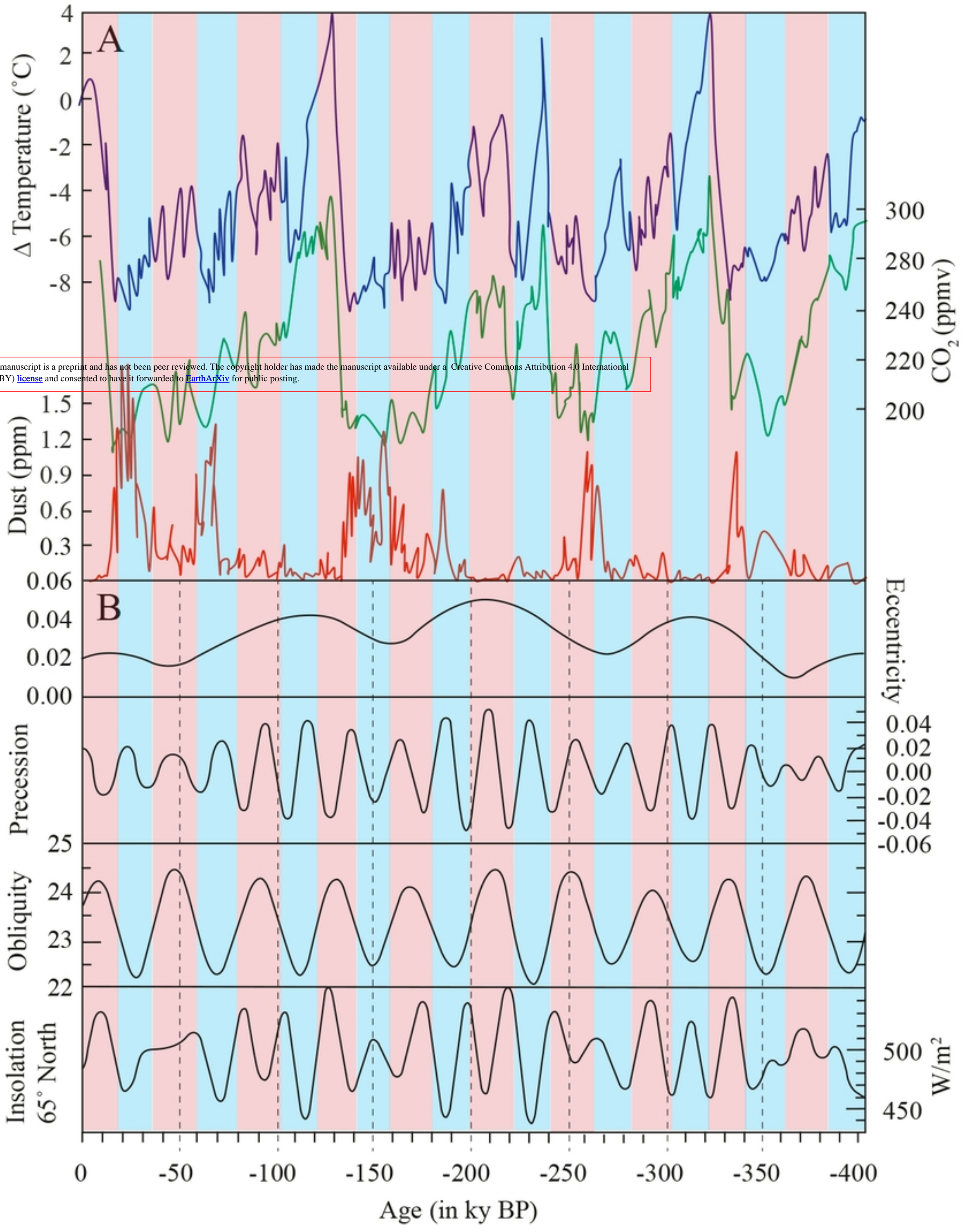


Fig 16

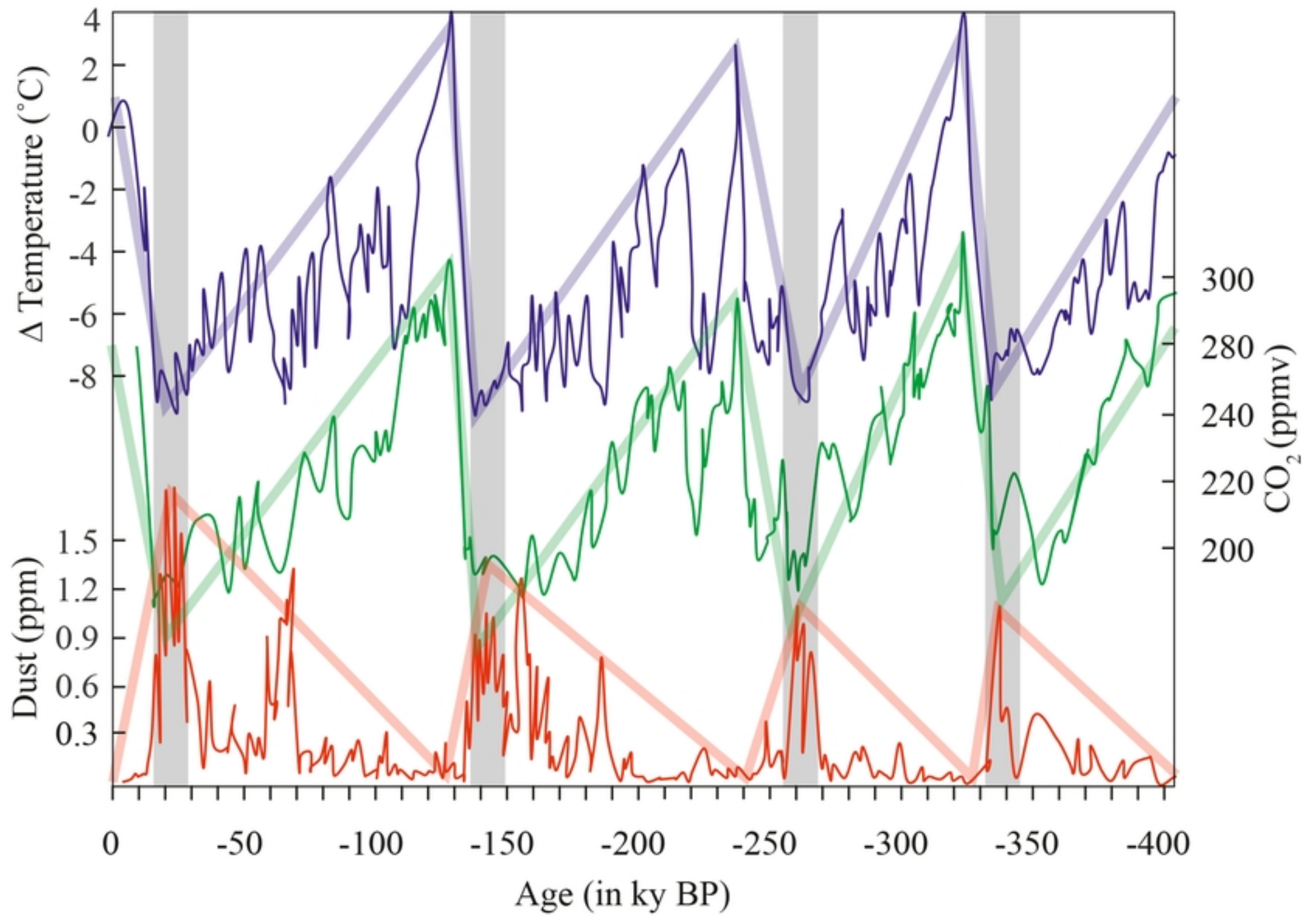


Fig 15

$$1.88x - 0.09806$$

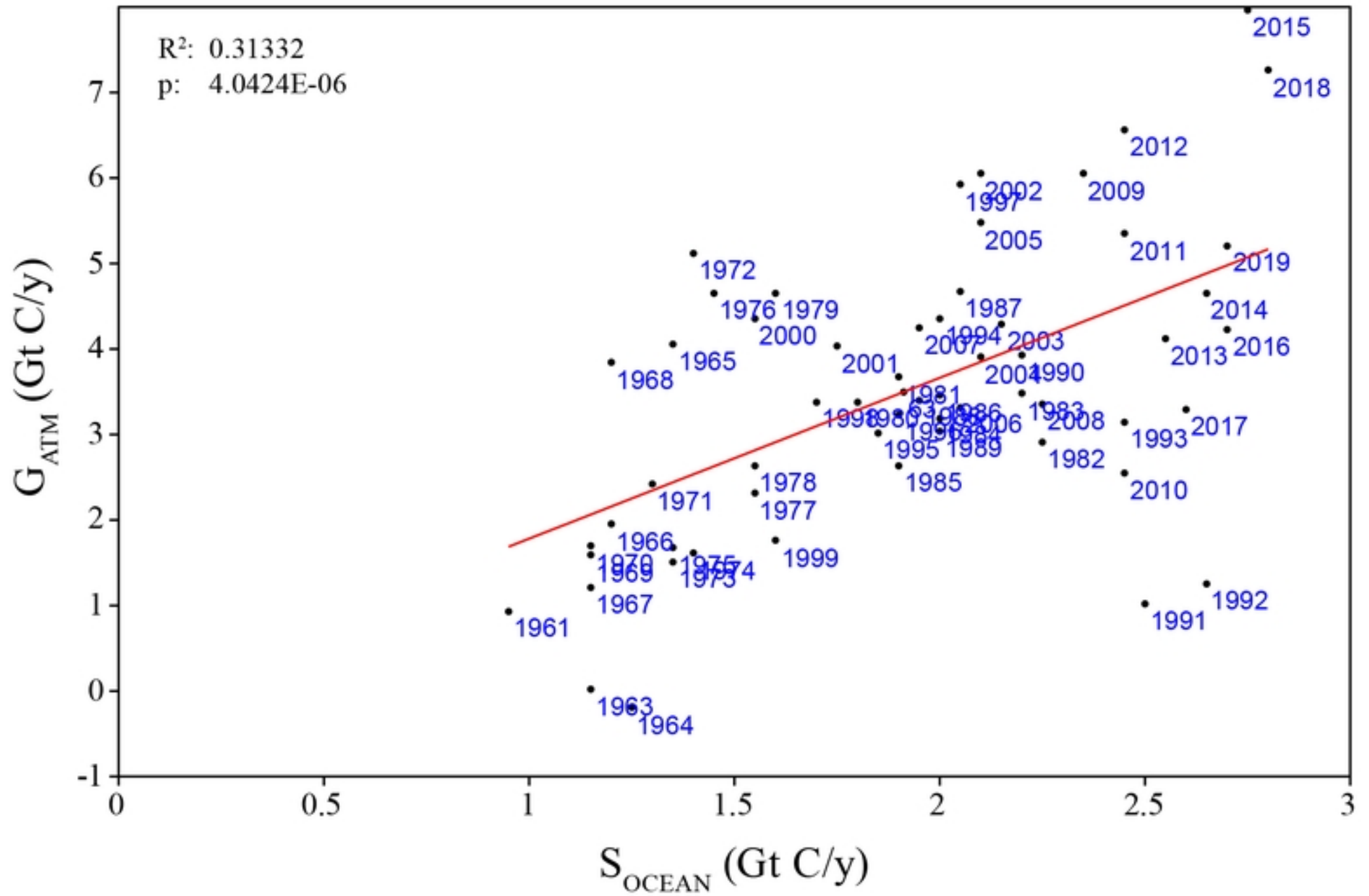


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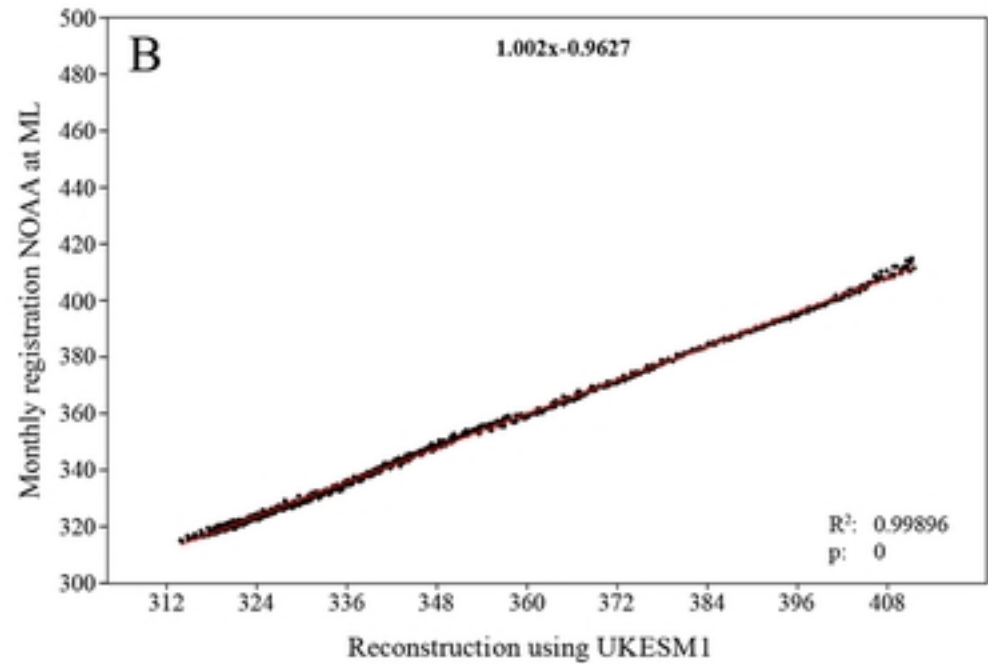
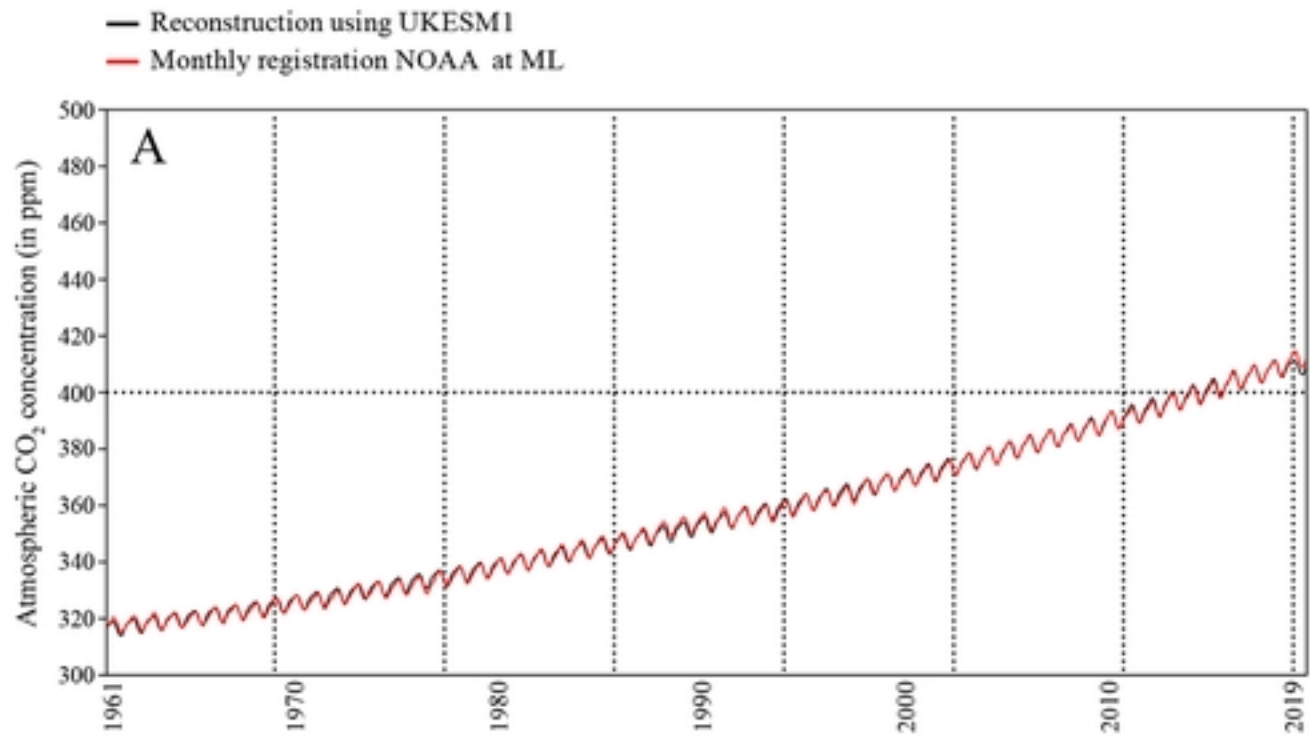


Fig13



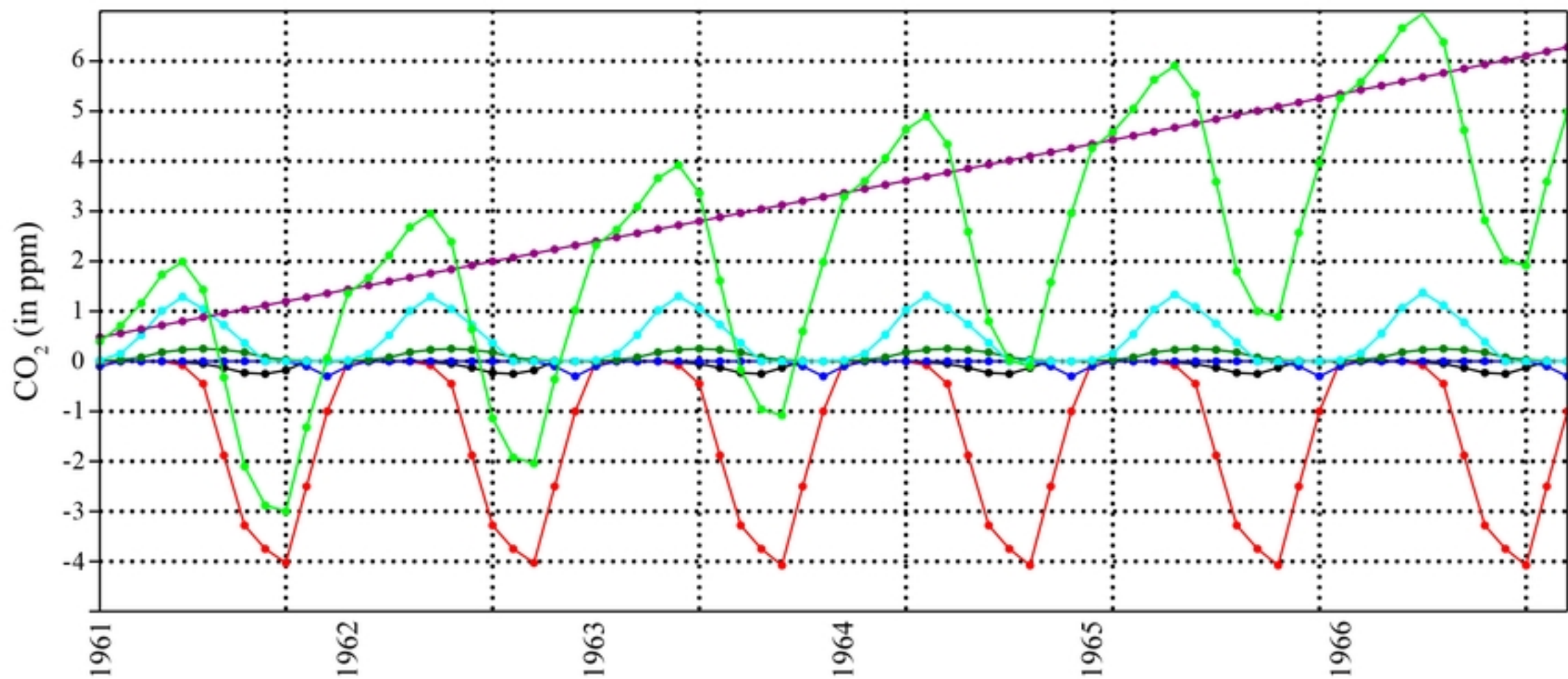
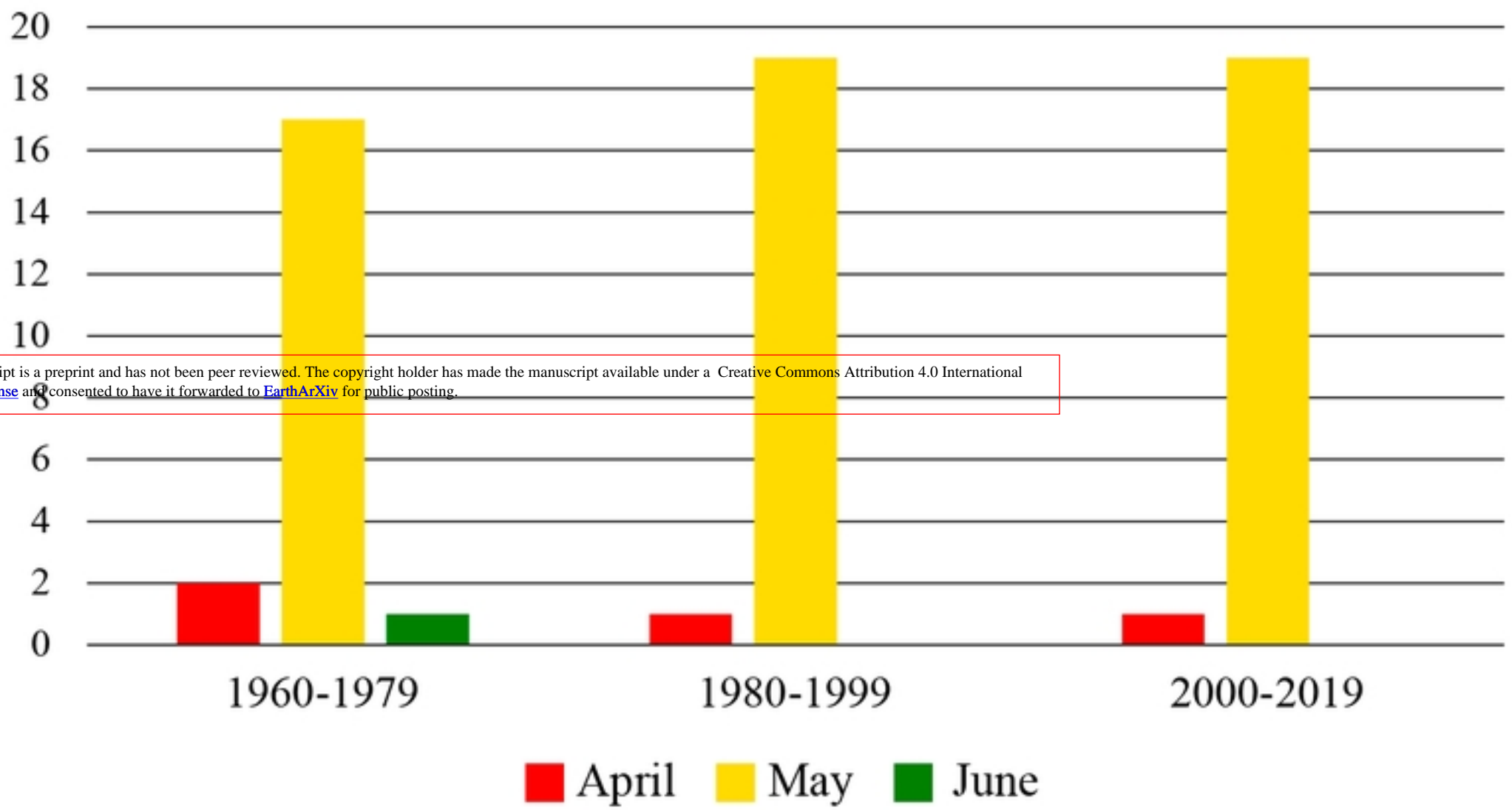


Fig 12

# A Transition from emission to uptake



# B Transition from uptake to emission

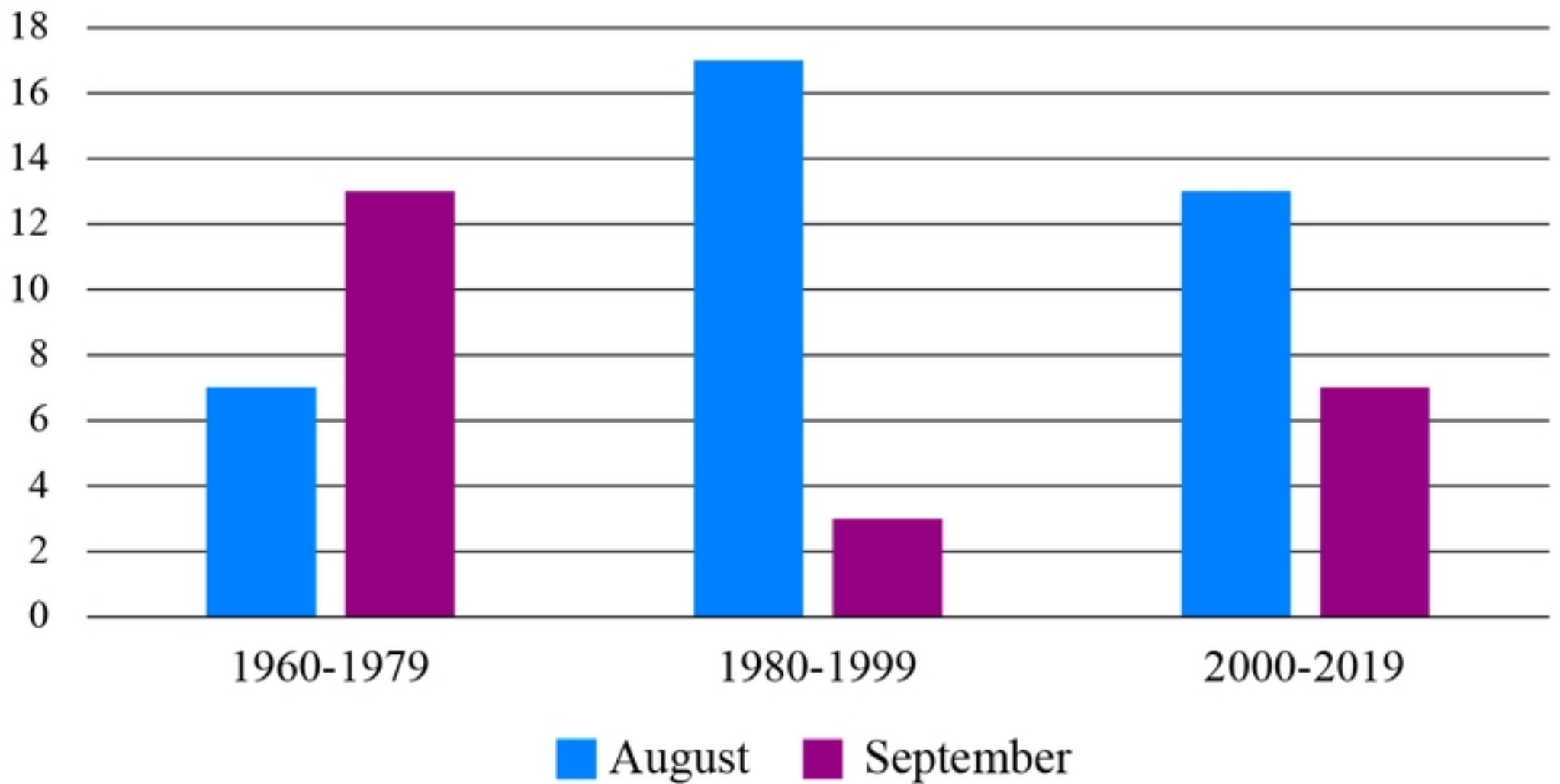


Fig 11

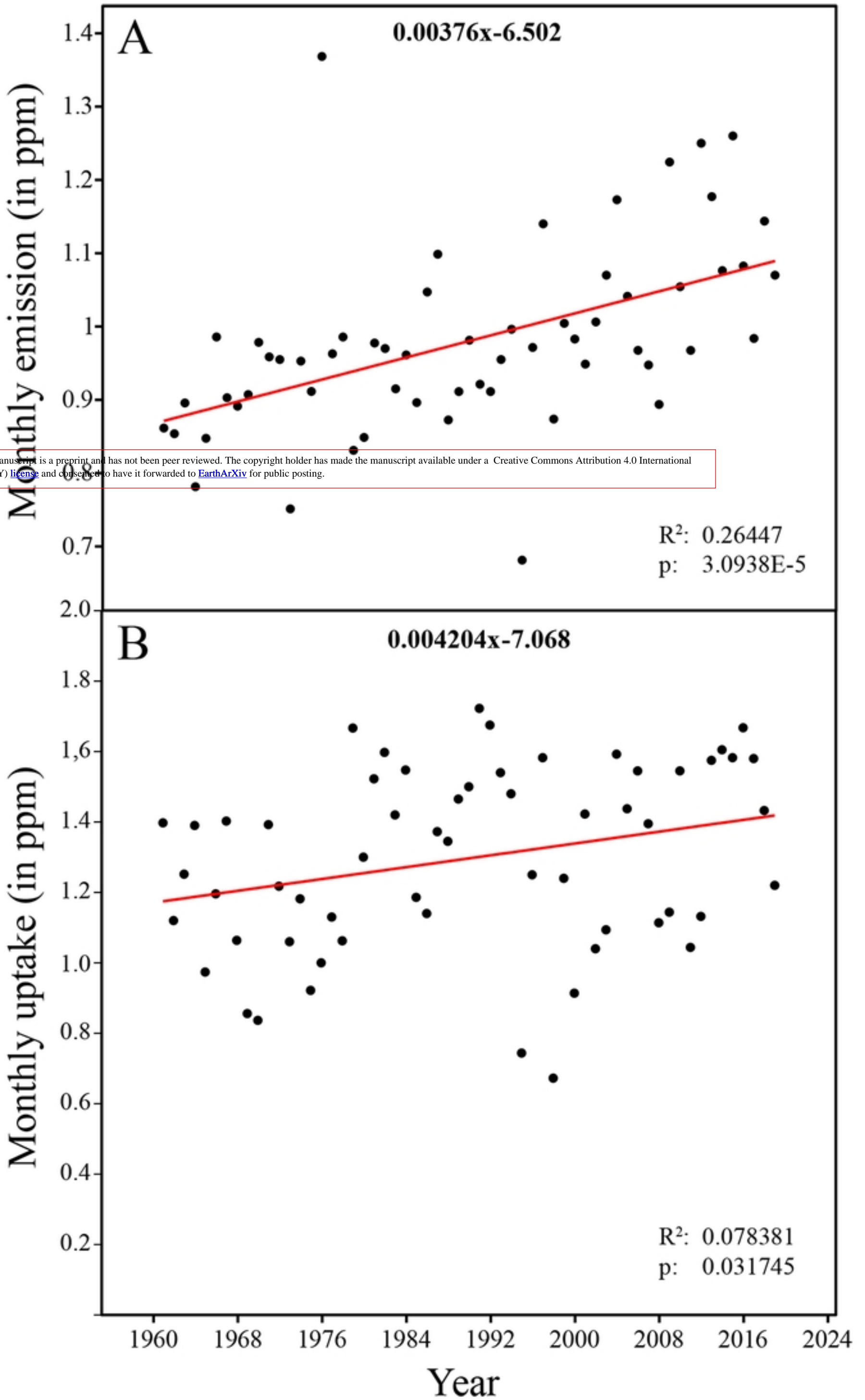


Fig 10

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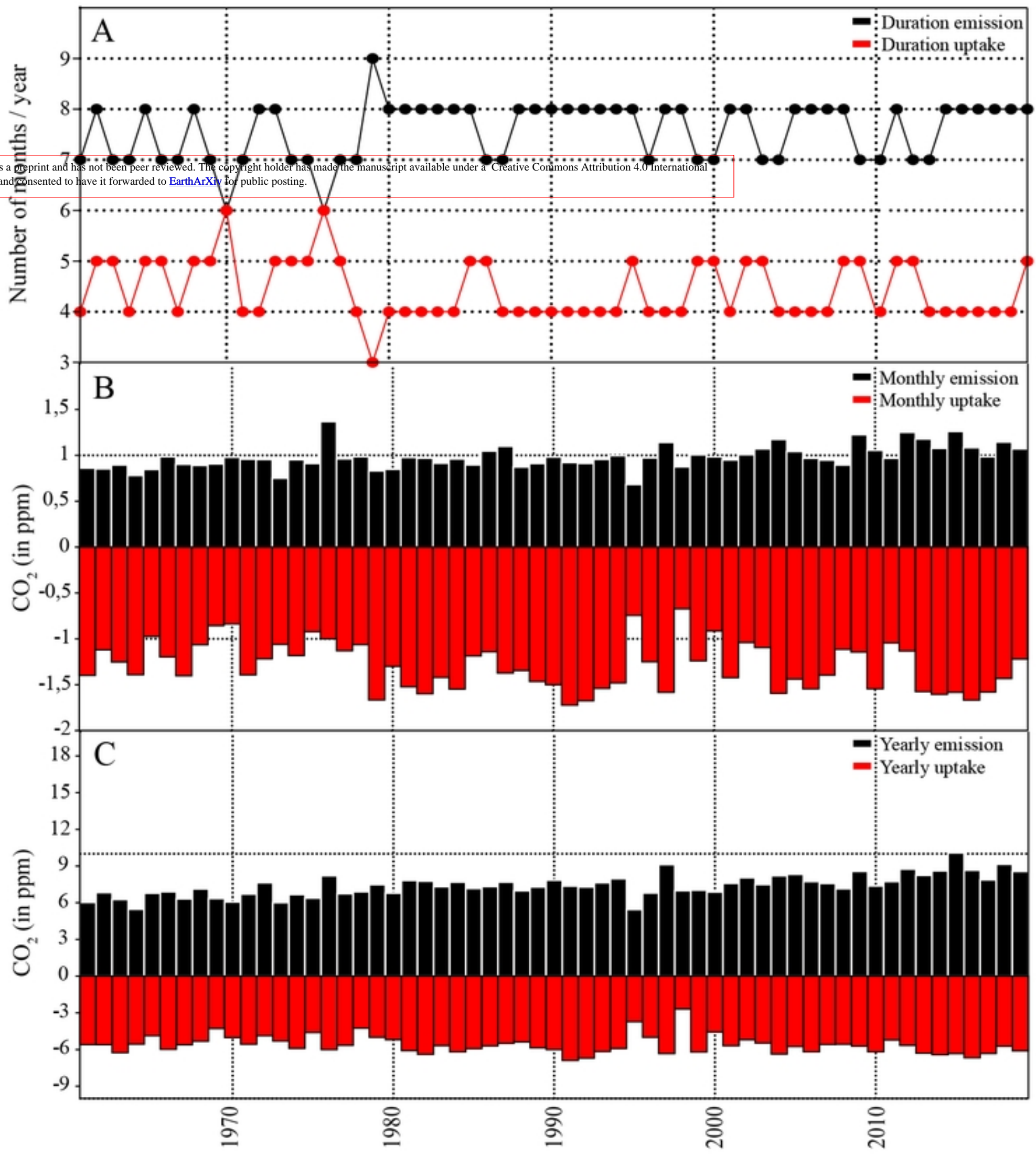


Fig 9



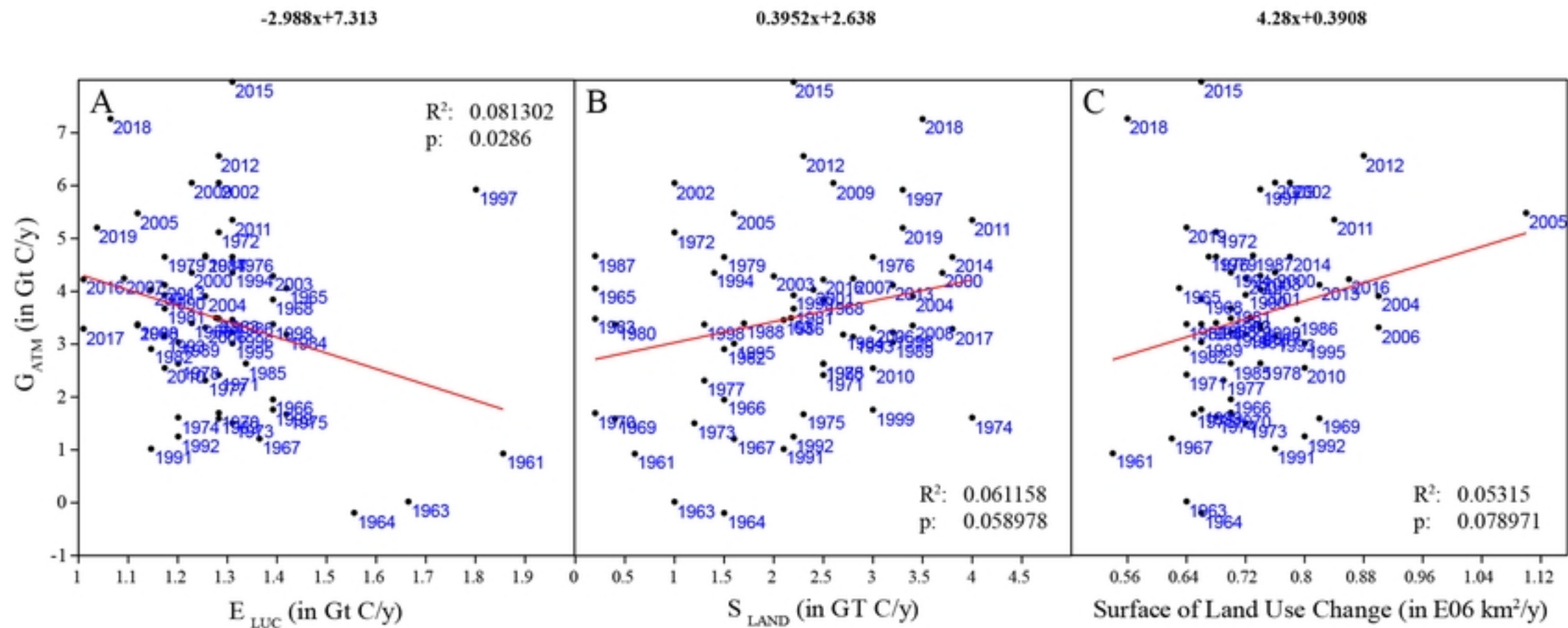


Fig 8

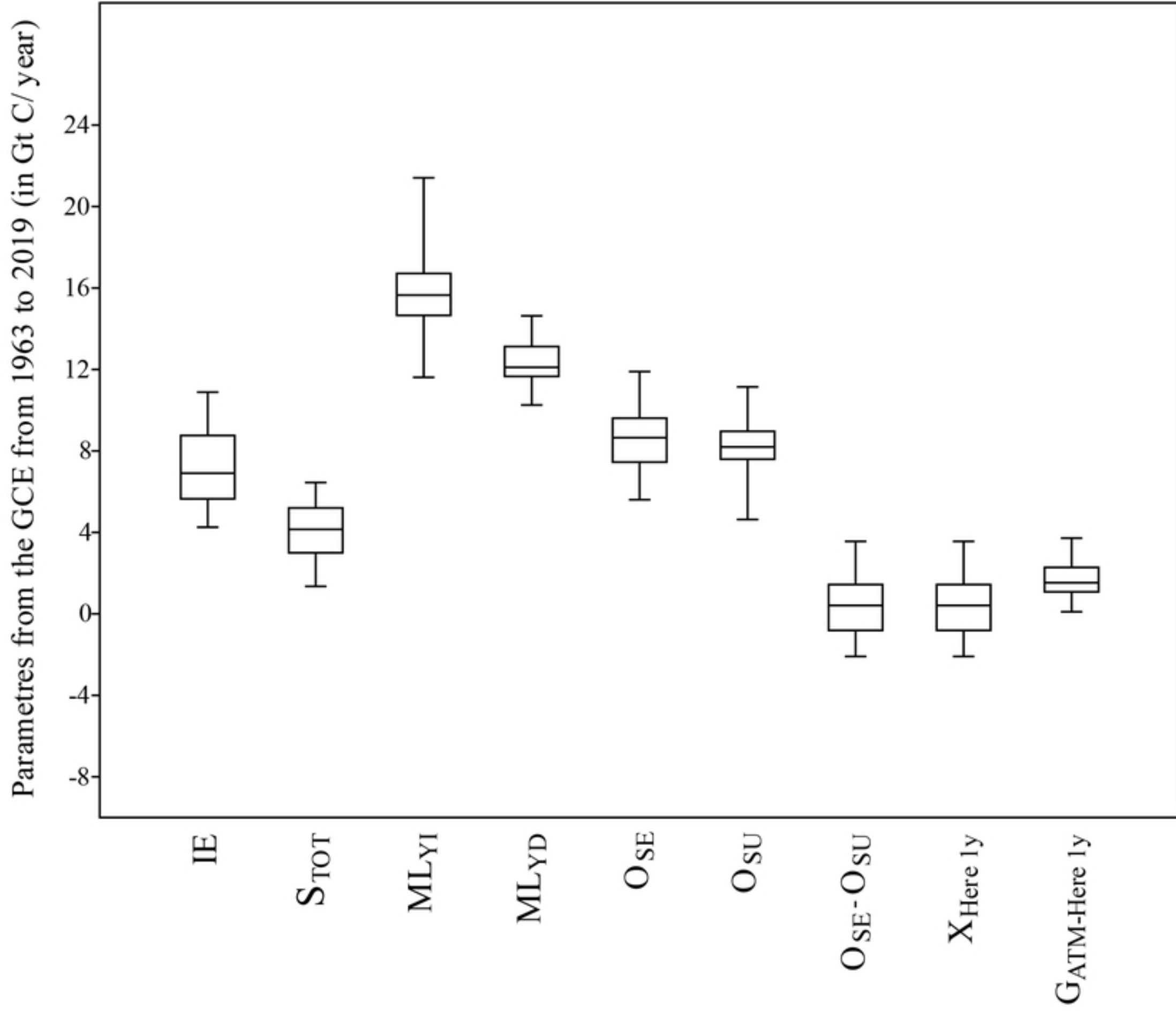


Fig 7

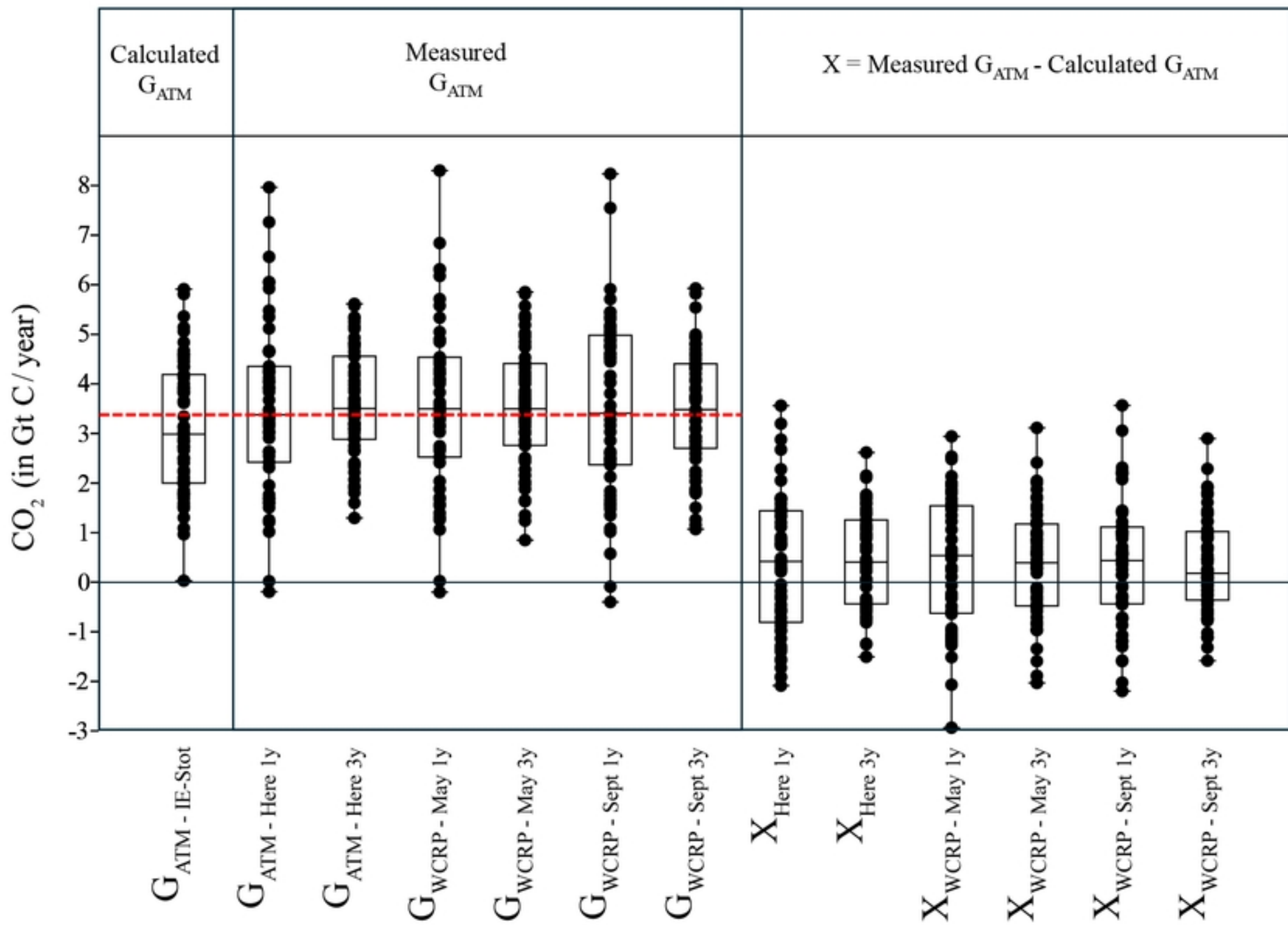


Fig 6

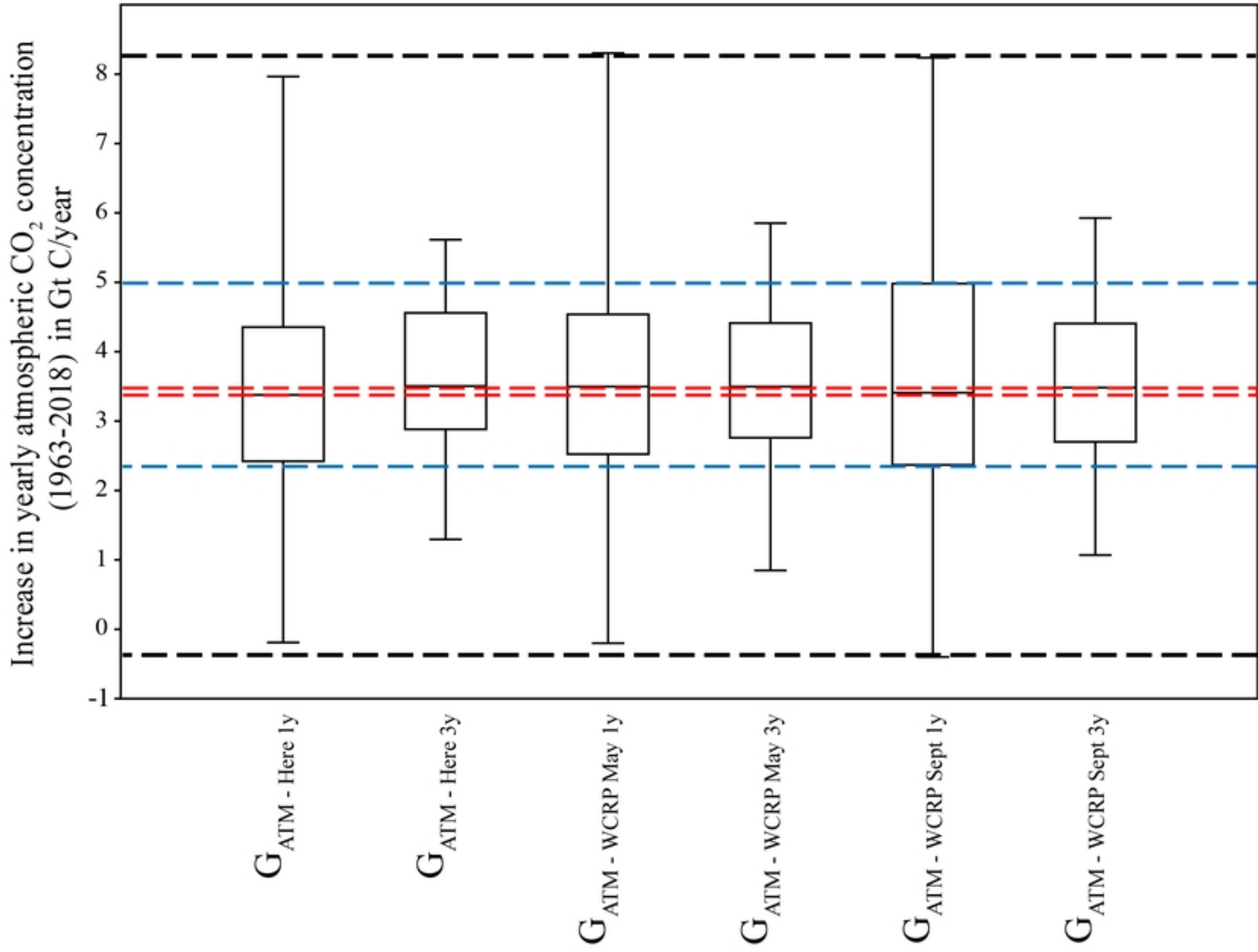


Fig 5



— Atmospheric CO<sub>2</sub> concentration measured on the South Pole (NOAA)  
— Global warming emission inferred from polynomial regression analysis ( $8.784 \cdot 10^{-5} x^2 + 0.102x + 331.9$ )

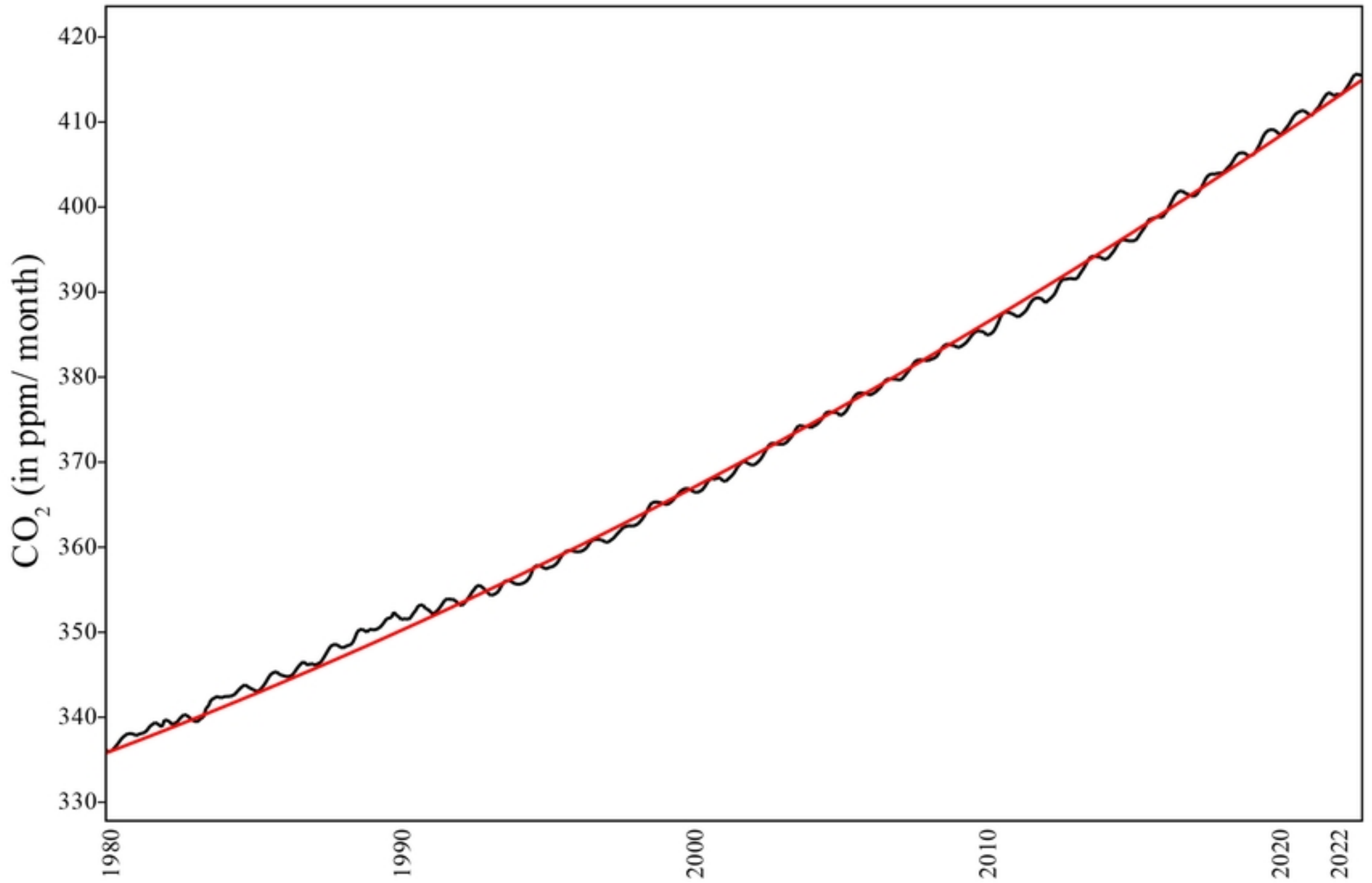


Fig 4

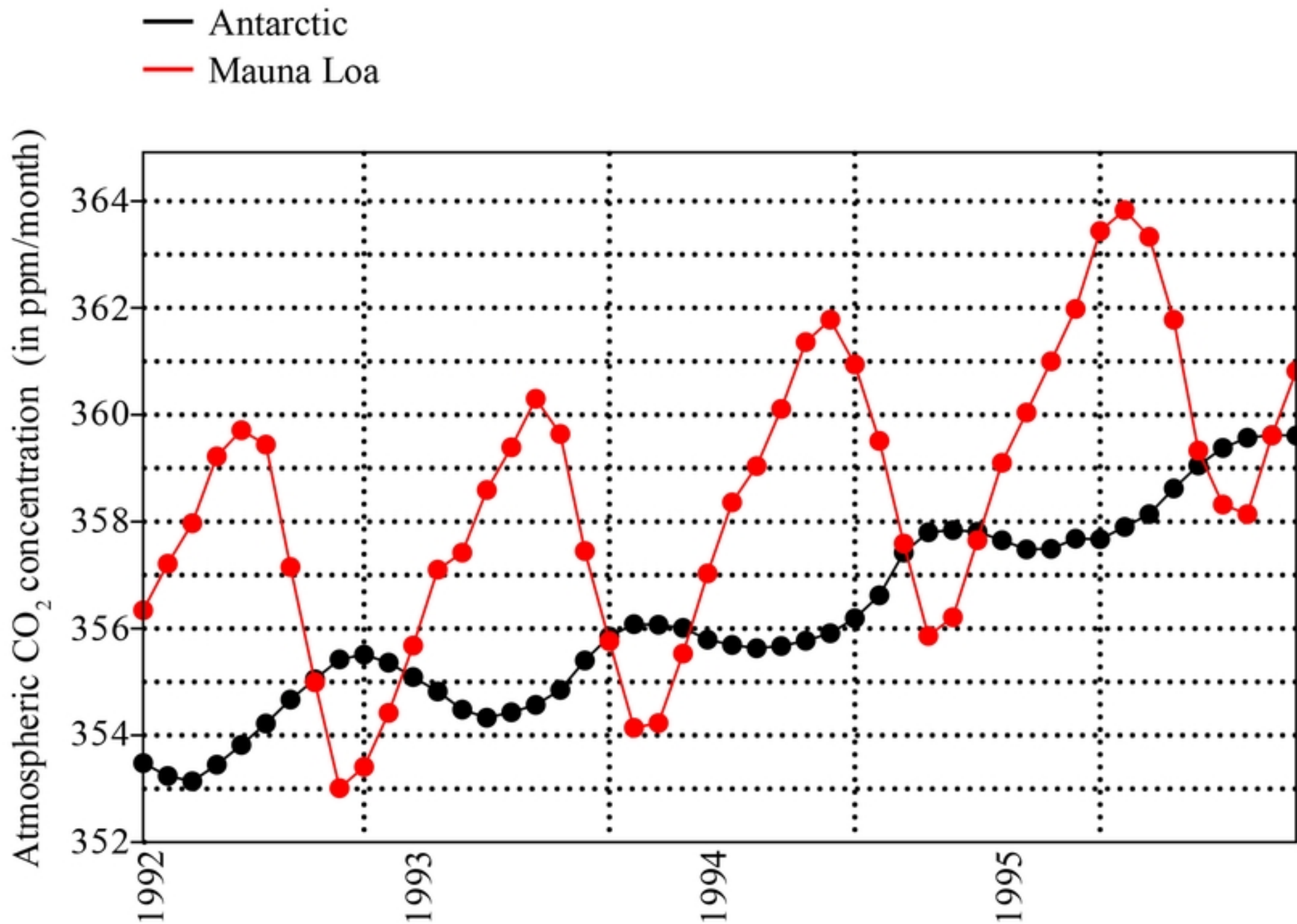


Fig 3

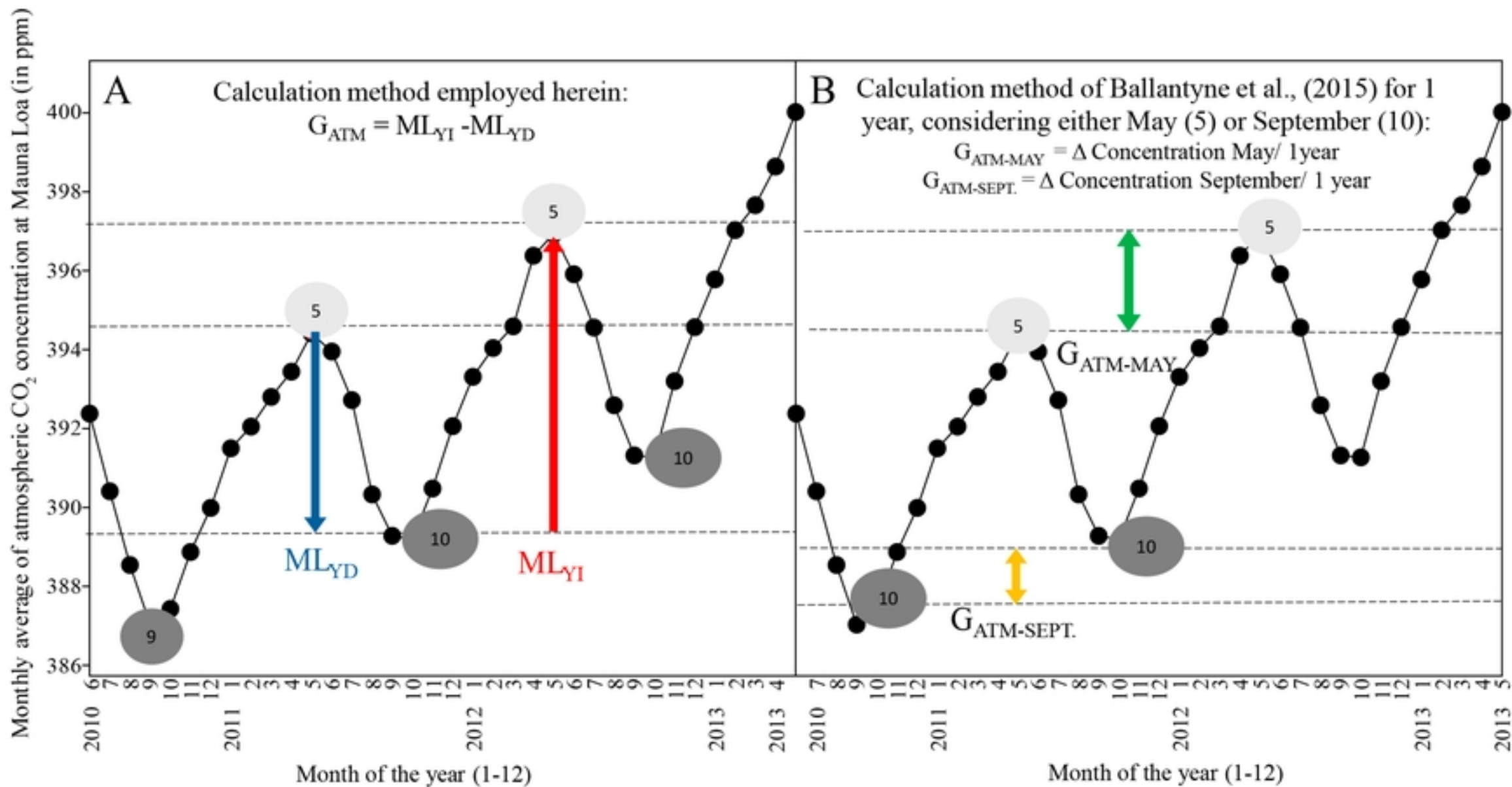


Fig 2

Max & minimal warming rates in the 1856 experiment of Foote (in °C/mn)

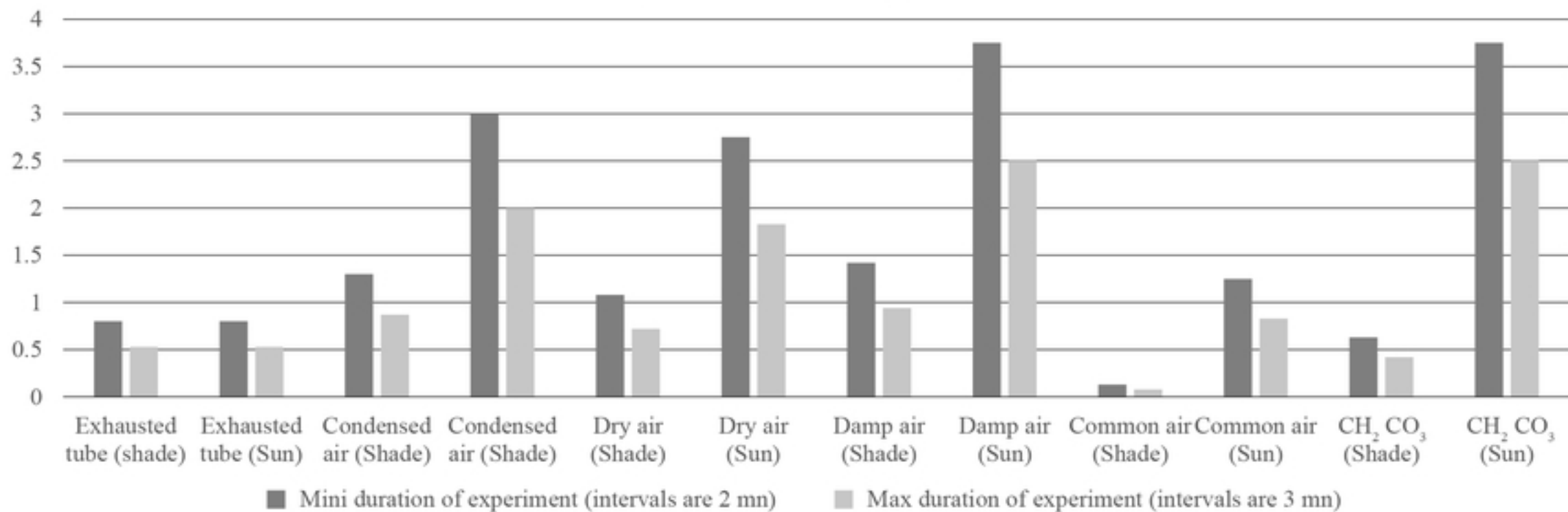


Fig 1