

This manuscript has not yet undergone peer-review. Subsequent versions of this manuscript may have slightly different content. If accepted, the final version of this manuscript will be available via the 'Peer-reviewed Publication DOI' link on the right-hand side of this webpage. Please feel free to contact the corresponding author.

SC-PREC4SA: A serially complete daily precipitation dataset for South America

Adrian Huerta^{1,2*}, Roberto Serrano-Notivoli³, and Stefan Brönnimann^{1,2}

¹Institute of Geography, University of Bern, Switzerland

²Oeschger Centre for Climate Change Research, University of Bern, Switzerland

³Departamento de Geografía y Ordenación del Territorio, Instituto Universitario de Ciencias Ambientales (IUCA), Universidad de Zaragoza, Zaragoza, Spain

*corresponding author: Adrian Huerta (adrhuerta@gmail.com)

ABSTRACT

This study introduces the serially complete precipitation dataset for South America (SC-PREC4SA), a daily precipitation dataset (1960-2015) designed to address observational gaps and ensure temporal consistency across diverse climates. The raw dataset underwent quality control, gap-filling, and homogenization procedures. Applied robust quality control highlighted common but also overlooked issues, enhancing data reliability. Gap-filling achieved a mean accuracy of 70 % (60 %) in the prediction on wet/dry days (wet-day magnitude). These metrics highlight the reliability of the gap-filling process, particularly in mixed climates, where station networks are sparse. The homogenization algorithm, focused primarily on wet days, effectively reduced inhomogeneities while preserving precipitation variability across South America. By integrating a unified framework and multiple outputs from 7799 stations, SC-PREC4SA provides a robust dataset that captures daily precipitation patterns with high to moderate accuracy and consistency. It offers a valuable resource for climate research, hydrological modeling, and water resource management, addressing longstanding challenges in precipitation data availability and quality for South America.

Background & Summary

Precipitation is a fundamental climate parameter integral to the global water and energy cycles^{1,2}, with applications across fields such as hydrology, agriculture, climate science, and water resource management. Access to long-term, high-quality precipitation data is essential for analysis and modeling in these areas. While some regions (e.g., Europe and North America) are covered by dense weather station networks, others (e.g., Africa and South America) have sparse and uneven coverage due to economic and geographical constraints. Gridded precipitation data, such as radar, satellite, reanalysis, and merged products³, are often preferred as they offer spatial continuity, particularly in data-sparse regions. However, these gridded datasets rely on precipitation station data for assimilation and bias correction to enhance their quality, as station data remains the most reliable source for precipitation measurement. Developing serially complete station datasets is therefore crucial, both to improve gridded datasets and to ensure long-term data quality and continuity⁴⁻⁷.

South America, a vast region with significant meridional reach and prominent orography, encompasses diverse climate and weather patterns, ranging from tropical to extra-tropical zones^{8,9}. Its climate is shaped by the Andes Cordillera (the longest continental mountain with an average height ≈ 4000 m) and the Amazon rainforest (largest rainforest on earth), which play crucial roles in humidity supply and vapor transport across the continent¹⁰. This diversity is exemplified by the extreme contrast between Colombia's equatorial regions—among the wettest globally¹¹—and Chile's Atacama Desert, the driest place on Earth¹². Such complexity leads to strong spatial-temporal precipitation variability, with marked seasonal patterns in some areas and little to none in others^{13,14}. South America is also highly susceptible to extreme events that profoundly impact socio-economic activities, energy demand, and public health. Intense rainfall can lead to flooding and landslides, especially in Andean regions, where lower-income communities in informal housing are particularly vulnerable^{15,16}. Given these dynamics, it is essential to understand South America's climate variability, yet South America's precipitation weather station data coverage is limited and inconsistent¹⁷⁻¹⁹. Differences in data density and management practices across National Meteorological and Hydrological Services (NMHSs), combined with challenging terrain, highlight the need for comprehensive continental datasets to better address cross-boundary climate and extreme weather events.

Developing serially complete datasets of daily precipitation for South America is a relatively under-researched area. Existing serial datasets are mostly limited to specific regions, such as the Central Andes²⁰ (Bolivia and Peru) and Patagonia²¹ (Argentina and Chile), or have been developed as intermediate products for country-scale gridded precipitation datasets in Brazil²², Chile²³, Ecuador²⁴, and Peru²⁴⁻²⁶. Globally, while more serial datasets are available²⁷⁻²⁹, these are typically gridded products that often

38 lack access to the underlying station time series data. Furthermore, few global datasets include serial station data alongside the
39 gridded outputs^{7,30}. While global initiatives exist, they frequently face limitations, including incomplete data collection and
40 variable processing approaches. Built largely from international sources, global datasets may lack detailed local information,
41 especially in countries with restricted data-sharing policies^{31,32}. Additionally, most global and regional datasets apply only
42 one or two of the essential steps, such as quality control, gap-filling, or homogeneity adjustments, rather than an integrated
43 approach. Creating a unified framework that incorporates all these procedures is essential for fully supporting researchers and
44 users with different needs. Thus, developing a high-quality, serially complete dataset with a comprehensive framework for
45 South America is a critical challenge that remains to be addressed.

46 In this study, we present a serially complete dataset of daily precipitation for South America (SC-PREC4SA), spanning from
47 1960 to 2015. This dataset was developed through a consolidated framework that integrates four key processes: unification,
48 quality control, gap-filling, and homogenization. This approach provides researchers and users with long-term, quality-
49 controlled data, along with gap-filled and homogenized time series outputs (each process with its output). As the first dataset of
50 its kind in the region, SC-PREC4SA offers an unprecedented opportunity to comprehensively study precipitation patterns in
51 South America based on observational data.

52 **Methods**

53 **Overview**

54 The methodology to produce SC-PREC4SA consists of four major procedures (Figure 1): unification, quality control, gap-filling,
55 and homogeneity. First, we unified the obtained raw database because multiple stations may share the same data or location
56 (raw_obs). We then used quality control processes to guarantee that the precipitation data was of the highest possible quality
57 (qc_obs). This process employs both a regular process and an enhanced protocol. Next, the gap-filling process was carried
58 out, which involved using auxiliary datasets in conjunction with statistical learning models to provide precipitation predictions
59 (mod_pred and bc_pred) and associated errors (err) for each day and station location of qc_obs. In this step, we created two
60 databases based on the type of prediction involved to fill the gaps (obs_mod and obs_bc). Following that, obs_mod and obs_bc
61 are homogenized independently (detection and adjustment) to ensure temporal variability and reduce potential inhomogeneities
62 of the previously applied process (hmg_obs_mod and hmng_obs_bc). At last, we generated multiple databases as outputs, each
63 of which reflected the results of the key procedures. In the following sections, we present further information and specifics
64 concerning these outputs, as well as the processes used.

65 Finally, due to the complex climate variability and topography in South America, instead of dividing the continent by
66 countries (Figure 2a) we used ecological regions (ecoregions) that depict a more accurate regionalization. The ecoregions
67 were as follows (Figure 2b): Northern Andes (NAS), Peruvian/Atacaman Deserts (PAD), Central Andes (CAS), Southern
68 Andes (SAS), Amazonian-Orinocan Lowland (AOL), Eastern Highlands (EHL), Gran Chaco (GCH), Pampas (PPS) and
69 Monte-Patagonian (MPN). We obtained this ecoregion classification from Griffith et al.³³, although most of them belong to
70 classification level I, we also decided to add one that belongs to classification level II (PAD). This was done to differentiate
71 how extremely arid that region was¹². In addition, we reclassified a small part of EHL within AOL to ensure better spatial
72 discretization.

73 **Data**

74 **Precipitation raw data**

75 The raw precipitation database used in this study belongs to different sources that come from NMHSs of South America (Figure
76 2a) and global databases such as Latin America as the Climate Assessment & Dataset (LACA&D) and Global Historical
77 Climatology Network Daily (GHCNd). In particular, we collected stations from seven NMHSs (Argentina, Bolivia, Brazil,
78 Chile, Colombia, Ecuador, and Peru). To fill the gap for other countries (Curacao, French Guiana, Guyana, Paraguay, Surinam,
79 Suriname, Trinidad, and Tobago, Uruguay, and Venezuela), LACA&D and GHCNd were also used although some stations
80 were also present for the other countries (Supplementary Table 1). Therefore, we compiled a large set of precipitation data
81 representing 15165 potential time series for the 1960 - 2015 period. Due to restrictions on South American NMHSs, the raw
82 data from some sources cannot be distributed with this publication. Readers who wish to obtain the primary data should contact
83 each agency or institution previously mentioned. It is important to note that while a substantial portion of the raw data is openly
84 accessible, several data series remain restricted and can only be accessed upon request (see Data Records section). Researchers
85 are referred to revise the data provided by each institution via their official webpage and for further data requests contact each
86 agency or institution individually.

87 South America has a very sparse and uneven spatial distribution of stations. Most (less) stations are found in the western and
88 eastern (central) parts of the continent. From an ecoregion perspective (Figure 2b, c), the three highest (lowest) dense ecoregions
89 were EHL, NAS, and CAS (MPN, GCH, and PAD). Regarding the temporal availability of observations (Supplementary Figure
90 1), it is noticed that at the South American scale, the amount of data increased from the 1960s to the 1980s, followed by a

91 decrease until 2015. This increase and decrease pattern was also seen at the ecoregion scale, particularly in NAS, AOL, EHL,
92 GCH, PPS, and GCH. Only PAD, CAS, SAS, and MPN presented a continuous increase from the 1960s.

93 **Auxiliary datasets**

94 In this study, we employed two auxiliary datasets: ERA5-Land precipitation and Digital Elevation Model (DEM)-derived
95 topographic covariables. These were exclusively used for the gap-filling procedure, and are detailed below.

96 The ERA5-Land³⁴ is an upgraded version of ERA5 designed specifically for land surface applications. It has a finer spatial
97 resolution of 9 km compared to 31 km and 80 km for ERA5 and ERA-Interim, respectively. A triangular mesh-based linear ap-
98 proach is used to interpolate precipitation for ERA5-Land from ERA5³⁵. We downloaded the daily aggregated ERA5-Land pre-
99 cipitation (1960 - 2015) from https://developers.google.com/earth-engine/datasets/catalog/ECMWF_ERA5_LAND_DAILY_AGGR
100 (accessed 25 October 2024).

101 The DEM-derived topographic covariables were provided by Amatulli et al.³⁶. They established a suite of topographic
102 covariables based on 7.5 arc-second Global Multi-resolution Terrain Elevation (GMTED2010³⁷) data at different spatial
103 resolutions on a global scale. Here, we used the 1 km spatial derived variables such as elevation, slope, aspect cosine, aspect
104 sine, aspect eastness, aspect northness, roughness, topographic position index, terrain ruggedness index, vector ruggedness
105 index, first-order partial derivative (E-W slope), second-order partial derivative (E-W slope), first-order partial derivative (N-S
106 slope), second-order partial derivative (N-S slope), profile curvature and tangential curvature. DEM-derived topographic
107 covariables were downloaded from <https://www.earthenv.org/topography> (accessed 25 October 2024). Furthermore, we
108 determined other variables such as latitude, longitude, and distance to the ocean at the same spatial resolution. As a result, we
109 used 19 topographic covariables.

110 **Methodology**

111 **Unification**

112 The collected database is the result of merging many sources. As a consequence, duplicated or overlapped stations (similar
113 locations or data) might be identified. Small location discrepancies among such stations may be due to differences in precision
114 in reporting the latitude and longitude coordinates or may reflect different nearby measurement sites³⁸. Therefore, some criteria
115 must be applied to eliminate duplicated stations in both the subsequent procedures and the final dataset. In this work, we used a
116 similar unification approach from previously constructed datasets⁷, but with few variations.

- 117 • Criteria 1: If the distance between two stations (or more) is less than 10 m, we calculated the correlation and mean
118 absolute error with the five nearby stations. The selected station is the one with the highest correlation and lower mean
119 absolute error.
- 120 • Criteria 2: If the distance between two stations (or more) is larger than 10 m but smaller than 25 km, we calculated the
121 correlation, mean absolute error, and the percentage of similar data (excluding values below 0.5 mm). If the correlation
122 exceeds 0.999, the mean absolute error is less than 0.1, and the percentage of identical data exceeds 50 %; the station
123 with a longer period is kept. At least two of the conditions should be fulfilled in order to remove a station.
- 124 • Criteria 3: The station elevation values are replaced by the nearest grid cell of a 250 m resolution DEM. This process
125 was made due to concerns regarding the accuracy of the raw elevation data. To ensure station location, we examined the
126 surrounding area. If a station has more than 50% negative DEM values within a 500 m, it is removed. This is done to
127 ensure that the stations are located above sea level (continental area).

128 The data generated at this stage represents the raw_obs (Figure 1).

129 **Quality Control**

130 After the station unification, stations are quality-controlled (QC) using two strategies: standard and enhanced QC. The first
131 approach detects (mostly) single suspicious values, while the second addresses recurring data quality issues that may remain
132 undetected by standard QC processes^{18,20}.

133 **Standard** Previous precipitation QC research provided the basis for the automatic standard QC checks^{6,29,39}. Although these
134 QC checks were used globally (or in large regions) with no climate-specific criteria, we decided to establish one based on the
135 percentage of wet days (> 0.1 mm). This means that some QC steps were applied differently depending on the type of climate
136 (arid or wet; Supplementary Figure 2). The standard QC steps are as follows:

- 137 • Repetition nonzero check (SQC-01): Daily records were flagged if constant values exceeding 10 mm/d persist for more
138 than four days. If constant values were discovered within a month, all values were flagged.

- 139 • Repetition zero check (SQC-02): Detection of suspicious zeros records was performed by comparing the frequency of
140 zeros of each year with the one obtained on a normal scale (average). A year is flagged if the frequency of zeros is greater
141 (or lower) than the normal frequency of zeros plus (minus) six times the interquartile range of the normal frequency of
142 zeros. The frequency of zeros was only computed if 85% of data was available; and, if the percentage of wet days is
143 above 5%.
- 144 • Subsequent month duplicated records check (SQC-03): Duplicated daily records in the subsequent months (up to eleven)
145 were identified by calculating a correlation and the number of equal values (excluding values below 0.5 mm). The criteria
146 for the temporal correlation coefficient and the number of days with equal values are set at 0.3 and 10, respectively. If the
147 conditions are met, both the 10 days (target) and duplicates are flagged.
- 148 • Subsequent year-month duplicated records check (SQC-04): Duplicated daily records in the same month for the
149 subsequent years (up to eleven) were identified by calculating a correlation and the number of equal values (excluding
150 values below 0.5 mm). The criteria for the temporal correlation coefficient and the number of days with equal values are
151 set at 0.3 and 10, respectively. If the conditions are met, both the 10 days (target) and duplicates are flagged.
- 152 • Z-score-based outlier check (SQC-05): Daily records were flagged if its difference from the daily-normal mean was
153 larger than nine sample standard deviations in stations with a percentage of wet days above 5%. The daily-normal mean
154 is calculated from data within a 15-day window centered on all available years (at least 10 years). For stations with a
155 percentage of wet days below 5%, the difference was set three times higher. This step was repeated three times.
- 156 • Spatiotemporally isolated value check (SQC-06): A daily record was flagged if it was extreme both in space and time. To
157 meet these conditions, the percentile difference i) with the five nearby stations within a radius of 400 km (space) and ii)
158 with the previous and next day (time) must be greater than the 99.99th percentile.
- 159 • Unique or full dry records check (SQC-07): Stations with fewer than 15 unique values or more than 99.5% dry records (<
160 0.5 mm/d) are flagged. This step was only conducted in stations with a percentage of wet days above 5%.

161 After setting any standard QC-flagged observation as missing, we defined two criteria to select the best stations in terms of
162 the amount of available data: time series i) with at least 10 years of data for each day of the year, and ii) with at least 5 years of
163 continuous data (a year is full if it has 70% of data). The selected stations were used for the enhanced QC.

164 **Enhanced** The enhanced QC process was developed in Hunziker et al.^{18,20}. They created a comprehensive set of tests
165 that inspect often overlooked issues (truncations, small gaps, asymmetric rounding patterns, and measurement precision
166 inconsistencies, among others) through data visualization techniques that let users manually correct errors or remove specific
167 periods of the dataset. Nevertheless, by increasing the number of weather stations and the size of the study area, the data
168 visualization approach is impractical. In this regard, we automatized the tests by proposing a classification level to describe
169 how good a station was. Therefore, instead of flagging time series periods, we flagged the stations based on the enhanced QC
170 test levels. The enhanced QC tests are as follows:

- 171 • Truncation (EQC-01) is when heavy precipitation episodes are truncated or noticeably reduced in frequency above a given
172 threshold. Because there is no preceding algorithm for truncation, it is defined here as when the maximum boundary
173 of a time series lasts for a set length of time (years). The maximum boundary is computed as the daily precipitation's
174 maximum moving window value. Thus, based on the length of years:
 - 175 – Level 0: no truncation (a constant maximum boundary lasts less than 3 years).
 - 176 – Level 1: a constant maximum boundary lasts longer than 3 years but less than 5 years.
 - 177 – Level 2: a constant maximum boundary lasts more than 5 years.
- 178 • Small gaps (EQC-02) can be seen as unreported precipitation events that result in a gap or a frequency reduction in values
179 below a specific threshold. To define the small gaps, we calculated the total count of values in five precipitation ranges
180 from 0-1, 1-2, 2-3, 3-4, and 4-5 mm (not including the values in the limits) for each year. Therefore, considering the
181 number of years with zero counts:
 - 182 – Level 0: no small gaps (years with at least one value in any of the precipitation ranges).
 - 183 – Level 1: small gaps in at least 5 consecutive years.

184 – Level 2: small gaps in more than 5 consecutive years.

185 • Weekly cycles (EQC-03) are characterized by the occurrence of wet days that significantly differ between the days of the
186 week. To compute the weekly cycles, first, for each day of the week, the probability of precipitation is calculated by
187 dividing the total number of wet days by the total counts of values. Later, the number of wet days is tested by a two-sided
188 binomial test (95% confidence level). Based on how many days were significant, we define:

189 – Level 0: no atypical weekly cycle (similar probability between the days of the week).

190 – Level 1: at least two days present an atypical probability (significant test).

191 – Level 2: more than two days present an atypical probability (significant test) or one day presents an extremely
192 different probability (more than 10%).

193 • Precision and rounding (EQC-04) patterns depict inconsistencies in the frequency of decimal values in the time series.
194 As there is no absolute correct frequency of decimals, we decided to measure how similar the decimal patterns are in the
195 time series. A decimal pattern is interpreted as the list of unique decimal values observed, sorted in descending order. In
196 this way, the decimal pattern for each year is computed first, followed by the selection of the most dominating pattern
197 (mode). Based on how much (in percentage) this dominating pattern represents the time series, we define:

198 – Level 0: coherent precision and rounding pattern (similar decimal pattern in more than 70% of the time series).

199 – Level 1: a similar decimal pattern in less than 70% but more than 50% of the time series.

200 – Level 2: different decimal patterns (no dominant pattern).

201 Preliminary experiments showed that the automatic enhanced QC could characterize the described issues in both high- and
202 low-quality time series (Supplementary Figures 3 and 4). However, some additions should be made before the application in
203 South America as a whole. This is due to the variety of climates (from wet to extremely arid) and inherent issues with the
204 precision and scale of daily precipitation values (Supplementary Figure 2). For that purpose, we set to level 0 the application of
205 EQC-01, EQC-02, and EQC-04 (EQC-03) in stations with a percentage of wet days below 15% (5%). Additionally, EQC-02
206 can be impacted by the length of the decimal patterns (EQC-04) so it is more probable to find small gaps in time series with
207 fewer decimals or with full integers; an issue that is usual in South America. Therefore, we also set level 0 for the application
208 of EQC-02 if the dominant decimal pattern of fewer decimals (less than 5 decimals) is more than 25% in the time series.

209 Each enhanced QC test indicates that level 0 represents the stations with the fewest quality issues. Ideally, we would select
210 only those stations at level 0 after applying the enhanced QC process. However, allowing some quality issues in subsequent
211 analyses to retain a greater number of stations is a reasonable trade-off. This approach can broaden spatial coverage, which is
212 particularly beneficial in South America, where station density is limited, while still maintaining acceptable data quality. Based
213 on this, we only flagged any station at level 2 in EQC-01, EQC-02, or EQC-03. Therefore, the selected stations in the enhanced
214 QC that intersect the ones in standard QC represent the dataset (qc_obs) used for the following process (Figure 1). Finally, it
215 should be mentioned that we did not discard the other stations as they remain valid, particularly for the gap-filling process.

216 **Gap-filling**

217 The key process for gap-filling is based on the concept of reference values (RVs) developed by Serrano-Notivoli et al.^{40,41}. RVs
218 are estimated independently for each day and location using the available data from nearby stations and topographic covariables
219 (latitude, longitude, and elevation). This means that RVs are local models in space and time. Thus, this framework enables the
220 building of highly flexible models that can represent local precipitation conditions.

221 RVs employ a hybrid modeling approach to predict daily precipitation. The logic is to first use a classification model
222 to predict whether a day is wet or dry, and then apply a regression model only to the wet days to estimate the amount of
223 precipitation. As a result, for each location and day, the RV is based on two predicted values: (i) a binomial prediction (BP)
224 of the probability of occurrence of a wet day and (ii) a magnitude prediction of precipitation (MP), in the case where a wet
225 day is predicted. The combination of these two values ($RV = MP$ if $BP > 0.5$, else $RV = 0$) produces the estimated RV and its
226 associated uncertainty (standard error) for each day and location. Generalized linear models (glm) are used as the modeling
227 foundation in RVs.

228 Previous research demonstrated the usefulness of RVs in dense-stations and moderately-challenging-terrain regions⁴²⁻⁴⁴.
229 However, some additions and changes must be made before implementing the framework in South America:

- 230 • Use of machine learning approaches as the modeling foundation in RVs. Besides glm, we tested support vector machines
231 (svm), random forests (rf), extreme gradient boosting (xgboost), and neural network (nn) models. Preliminary experiments
232 (Supplementary Figure 5) in small areas evidenced that machine learning models were better than glm. In addition,
233 it was found that glm, rf, and xgboost were more applicable over larger areas (arid to wet) without many fine-tuning
234 adjustments. On the other hand, nn and svm required adjustments to parameters for different areas, making them less
235 generalizable. Although rf and xgboost yielded similar mean efficiency, we opted for xgboost due to the greater number
236 of stations with higher efficiency. Hence, we used xgboost as the modeling foundation for the gap-filling process. It must
237 be pointed out that we used the default xgboost model^{45,46} (*nthread* = 1, *nrounds* = 5) as the hyperparameter tuning only
238 offered slight improvements.
- 239 • Use virtual stations to enhance the density of the original station network^{47,48}. Virtual stations come from the ERA5-
240 Land and were particularly useful for the early period of the dataset. These time series were not directly used, but a
241 bias-correction version (quantile mapping) with the closest station to the grid point. Similarly, the topographic covariables
242 were also obtained. So, we obtained a virtual station for each station, and they were fed into the gap-filling process as
243 real stations.
- 244 • Use of more than three topographic covariables in the classification and regression modeling. Besides, latitude, longitude,
245 and elevation, we also employed 16 variables listed in section 3.1. To reduce both autocorrelation and dimensionality, we
246 remove the autocorrelated (> 0.8) variables first and then retain the first principal components. Based on Horn's parallel
247 analysis⁴⁹, we noted that four components explained more than 65% of the variance (Supplementary Figure 6). In this
248 fashion, we used latitude, longitude, elevation, and the first four principal components as independent variables in the
249 gap-filling process.
- 250 • Use of an iterative framework of RVs. This was done to take advantage of those stations that did not have a common
251 period at the beginning of the gap-filling process^{47,48,50}, particularly in the early period. For the gap-filling procedure,
252 we employed up to three cycles in which we searched for nearby stations within i) 175 km, ii) 275 km, and iii) 650 km. It
253 should be noted that, despite the enormous distance ratio, we limited the number of stations to a minimum of 8 and a
254 maximum of 16 stations for a target station.

255 From 1960 to 2015, we produced RVs for each station and day under the abovementioned settings. The RVs framework
256 generated two primary outputs: the predicted RV (*mod_pred*) and error (*err*). However, we also obtained a bias-correction
257 version of the predicted RV named *bc_pred*. This bias correction (quantile mapping) was conducted in the final stage (of
258 the iterative framework), using the original time series to enhance the model estimation. Therefore, we built two databases
259 (Figure 1) by filling in the missing values of *qc_obs* with *mod_pred* (*obs_mod*) and *bc_pred* (*obs_bc*). These serially-complete
260 databases were used for the following process.

261 **Evaluation metrics** After calculating *mod_pred* and *bc_pred*, we compared them to days with available observations in
262 *qc_obs* to assess the gap-filling framework's efficiency. This evaluation is leave-one-out cross-validation because the RVs were
263 predicted without using the station's observation. For this purpose, we employed a variety of metrics that evaluate both the
264 continuous and categorical nature of precipitation. The continuous metrics were: the refined index of agreement (*dr*), mean
265 absolute error (*mae*), root mean squared error (*rmse*), normalized *mae* (*nmae*), and normalized *rmse* (*nrmse*). The categorical
266 metrics were: accuracy, precision, recall, F-measure (*f1*), balanced accuracy (*bcc*), and G-mean (*g_mean*). Despite the amount
267 of metrics, it should be stated that we focus the evaluation on two key metrics: *dr* and *bcc* (see Technical Validation section).

268 **Homogeneity**

269 Non-climatic factors (changes in station location, instrumentation, and observation techniques) can impact measurements. Time
270 series must be homogenized to reduce inhomogeneities and provide more reliable observations⁵¹. The literature provides a
271 range of homogenization algorithms, particularly for high-density networks with strongly correlated data^{51,52}. As a result,
272 homogenization performance worsens significantly when applied in sparse networks, with erroneous corrections likely due to
273 the low signal-to-noise ratio⁵³. In this view, the homogenization approach must be tailored to the characteristics of the study
274 area and climate variable.

275 In this study, we used a similar homogenization strategy based on previously applied approaches on global and continental
276 scales^{54,55}. Therefore, we used an automatic algorithm for both detection and adjustment without the use of metadata
277 information. In addition, relative and absolute approaches are combined for situations in which relative homogenization can not
278 be performed. The absolute test, which has a lower power of detection than the relative tests⁵¹, is thus intended as a backup test
279 for when a relative test is hardly possible⁵⁵.

280 **Detection** To ensure high confidence in breakpoint detection, we used a combination of different statistical tests and
281 intercomparison of their results. Five univariate breakpoint tests were applied⁵⁶: Student's, Mann-Whitney, Buishand-R,
282 Pettit, and Standard Normal Homogeneity Test. We opted for univariate rather than multivariate tests because they are more
283 straightforward to apply and are less likely to include spurious break detections⁵⁷. Depending on the availability of nearby
284 stations for a target time series:

- 285 • **Relative:** The algorithm searches for up to eight well-correlated (correlation > 0.6) nearby stations within a 1000 km
286 radius. Later, the five tests are applied to difference series (target minus nearby) created with three different temporal
287 aggregations (annual, April-to-September, and October-to-March) and two indices (PRCPTOT: total precipitation; R1mm:
288 number of wet days). Overall, we applied up to 54 different combinations of the approach. Finally, the breakpoint is
289 assigned using the pairwise comparison approach⁵⁸: a breakpoint is set to a certain year if it is found in at least the 7% of
290 the number of difference time series that are significant (p-value < 0.05), using a tolerance of ± 1 years.
- 291 • **Absolute:** The absolute approach is used if the algorithm detects fewer than four nearby stations or none at all. As a
292 result, the five tests are only used in the actual PRCPTOT and R1mm series at the three temporal aggregations. Here, it is
293 used up to nine distinct combinations. Finally, the breakpoint is assigned similarly to the relative approach.

294 **Adjustment** We adapted the quantile-matching technique outlined in Squintu et al.⁵⁴ to correct the inhomogeneous time
295 series. This method applies adjustments of varying magnitudes based on the value being corrected. Particularly, this approach
296 offers a more robust correction for extreme records, as it tailors adjustments to reflect the intensity of the observed values rather
297 than applying a uniform correction across all dates⁵⁹. It should be mentioned that this algorithm was created for temperature
298 data, therefore, we made some changes to be used for precipitation. Dry values (< 0.1 mm) are not corrected, and wet values
299 were transformed twice (square root and log) before the algorithm execution to force a normal distribution. Following the
300 correction, values were reversed to provide the actual precipitation values. Based on this consideration, the correction was
301 applied in two ways:

- 302 • **Relative:** The adjustment factor was computed using the target and nearby time series of the detection stage. It is assumed
303 that data after the break is correct, thus, the correction is backward. The correction is performed if there is a detected
304 break year, otherwise, the original data is kept.
- 305 • **Absolute:** The adjustment factor was computed using the target time series. This can be seen as an application of quantile
306 mapping as there are no nearby stations. Data before the break year is corrected based on the quantiles of the sample
307 after the break year. Similarly, the correction is only if there is a detected break year.

308 The homogeneity framework was applied after gap-filling to: (i) detect inhomogeneities caused by the gap-filling approach,
309 and (ii) because the method was more reliable when there were no gaps in the time series^{5,47,48,60}. We apply the homogenization
310 to each ecoregion only once. In addition, adjusted values were set to not exceed one unit difference of the root cubic difference
311 with the raw data. This was done to decrease the influence of the adjustment in the extreme tails^{61,62}. It still keeps the extreme
312 adjustment while preventing the creation of extremely excessive values, especially in extremely arid and wet areas. Finally,
313 two new databases have been created: hmg_obs_mod and hmg_obs_bc (Figure 1), which are homogenized versions of the
314 gap-filled databases obs_mod and obs_bc, respectively.

315 Data Records

316 The set of data generated in SC-PREC4SA consists of three key components. For rapid access, the data are divided into
317 different repositories and are stored in a figshare collection⁶³(<https://doi.org/10.6084/m9.figshare.c.5959863>).

- 318 • **SC-PREC4SA metadata:** A file (.csv) that provides information about each station. The file contains the following
319 information (*headers*): station code (*ID*), name (*NAME*), longitude in decimal degrees (*LON*), latitude in decimal degrees
320 (*LAT*), elevation from sources (*ALTs*), elevation from DEM in meters sea above level (*ALT*), country (*COUNTRY*), source
321 (*SOURCE*), ecoregion (*ECOREGIONS*).
- 322 • **SC-PREC4SA data:** Files (.csv) for each station that include the nine daily precipitation (mm/day) outputs (Figure
323 1) from 1960 to 2015. Each file contains the following information (*headers*): time step (*time_step*), raw time series
324 (*raw_obs*), quality-controlled time series (*qc_obs*), gap-filling model prediction (*mod_pred*), gap-filling model prediction
325 with a bias correction (*bc_pred*), error of gap-filling model prediction (*err*), quality-controlled time series plus mod_pred
326 (*obs_mod*), quality-controlled time series plus bc_pred (*obs_bc*), homogenized time series of obs_mod (*hmg_obs_mod*);
327 and, homogenized time series of obs_bc (*hmg_obs_bc*). Due to the number of stations, files are subdivided by ecoregions
328 compressed folders (.zip).

- SC-PREC4SA gap-filling metrics: A file (.csv) provides information about each station's gap-filling evaluation metrics. The file contains the following information (*headers*): station code (*ID*), station ecoregion (*ECOREGIONS*), type of gap-filling model (*mod_pred* or *bc_pred*; *MOD*), pairwise number of dates used to calculate the metrics (*n_data*), refined index of agreement (*dr*), mean absolute error (*mae*), root mean squared error (*rmse*), normalized mean absolute error (*nmae*), normalized root mean squared error (*nrmse*), accuracy (*accuracy*), precision (*precision*), recall (*recall*), F-measure (*f1*), balanced accuracy (*bcc*), G-mean (*g_mean*). In addition, the percentage of wet days was added (*wet_day*).

It should be pointed out we can not share raw data from Ecuador, Bolivia neither Peru due to data-sharing restrictions. Therefore, we set as missing data (*NA*) the *raw_obs* and *qc_obs* columns in each file that belongs to these countries in the SC-PREC4SA data repository.

Finally, we will upload the data generated with the other RV foundation models tested in the gap-filling procedure (glm and rf). Those will also be available in the main repository. The purpose of providing these different versions is for further research (see Usage Notes section).

Technical Validation

We report the suitability of SC-PREC4SA by exploring three key procedures: quality control, gap-filling, and homogenization. First, we summarise the results of the applied quality control. Second, we use statistical indicators to evaluate the gap-filling model's efficiency. Lastly, we assess the impact of homogenization by measuring the magnitude of breaks/adjustments as well as the temporal variability of PRCPTOT and R1mm.

Quality control

Following the unification process, we obtained a total of 14345 stations that underwent both standard and enhanced quality control (QC) procedures.

The results of the standard QC process, summarized in Table 1 by ecoregion, are expressed in terms of the number of flagged daily records and their corresponding percentages. Overall, flagged data accounted for less than 0.15% of the total dataset in South America, indicating a minimal portion of flagged values. Most issues stemmed from suspected zero values, duplicate records, and outliers, as identified in steps SQC-02, SQC-03, SQC-04, and SQC-05. At the ecoregion level, MPN exhibited the highest percentage of flagged data, despite having fewer stations (Supplementary Table 2). However, as expected, ecoregions with a higher station density showed greater flagged daily data.

From a temporal perspective (Figure 3), the proportion of flagged data per year was consistent with the overall average, with minor fluctuations across QC steps. However, pre-1965 data showed elevated percentages, peaking at 0.3–0.4 % in 1960 and 1963. This can be attributed to the frequent use of repeated zero values (SQC-02), indicating that early South American precipitation data may include a substantial number of false zeros.

To address systematic data quality issues undetected by standard QC, we implemented enhanced QC procedures. Table 2 presents results from Level 2 of EQC-01 to EQC-04, considered the "worst-case" scenario in terms of quality issues. On average, about 30 % of stations exhibited previously undetected issues, primarily due to small gaps (EQC-02) and precision/rounding patterns (EQC-04). These issues affected more than 15 % of the time series across South America (Supplementary Figure 7).

At the ecoregion level, every region except PAD exhibited at least ≈ 20 % of stations with undetected issues. Regions such as SAS and MPN (and to a lesser extent CAS, EHL, and PPS) had over 40 % (30 %) of stations affected. The PAD region showed no issues due to its specific wet-day percentage conditions. Overall, the enhanced QC results underscore the presence of significant quality issues in South American precipitation data, potentially impacting processes like gap-filling, homogenization, and previous analyses.

The automatic enhanced QC findings for CAS align closely with previous research. For example, Hunziker et al.²⁰ reported that approximately 40 % of observations were unsuitable for calculating monthly temperature means and precipitation sums due to quality issues. Our study identified similar problems in 34.74 % of precipitation time series. The discrepancy may be due to differences in methods; Hunziker et al.²⁰ manually applied additional tests beyond the four primary tests developed here. These additional tests, which often require visual inspection, were excluded due to challenges in automating such analyses.

The thresholds for the enhanced QC (and standard QC) were designed to account for South America's diverse climates. For example, in extremely arid areas, the lack of wet days poses challenges for reliably identifying patterns in the time series. While Hunziker et al.²⁰ did not encounter such issues in CAS (where wet-day percentages range from 10–70 %), these challenges are more pronounced in other parts of South America (Supplementary Figure 7).

To balance data quality and spatial coverage, we flagged only time series classified as Level 2 in EQC-01, EQC-02, or EQC-03. This approach prioritized retaining a larger number of stations while ensuring adequate spatial representation across South America (Supplementary Figure 7). As a result, 7799 stations (*qc_obs*) were retained, providing comprehensive coverage in the region. While this represents a reduction in the number of stations, the improvement in data quality significantly

381 outweighs the drawback of fewer observations. Notably, this number of stations is still greater than those in existing global
382 datasets^{7,27,28}.

383 **Gap-filling**

384 The results of the gap-filling framework were evaluated using statistical metrics computed for each output: `mod_pred` (model
385 prediction) and `bc_pred` (bias-corrected prediction). The evaluation focused on the metrics `dr` (continuous) and `bcc` (categorical),
386 as these metrics provide a standardized and intuitive measure of model performance (in both > 0.5 means good results). The `dr`
387 metric was chosen for its broad applicability and resistance to counterbalanced errors, as noted in a previous study⁶⁴. The `bcc`
388 metric was selected to ensure fair representation of imbalanced and balanced classification classes⁶⁵, particularly critical for
389 South America, where extremely arid and wet regions coexist with semi-arid, semi-wet, and mixed climates. By using `dr` and
390 `bcc`, the framework effectively addresses both continuous and categorical aspects of precipitation modeling.

391 The summarized results for `dr` and `bcc` by ecoregion (Table 3) revealed that the gap-filling framework performed relatively
392 well across South America, with both metrics exceeding 0.5 on average. This indicates the model's capability to capture
393 general precipitation patterns and reliably classify wet and dry days, despite the climatic diversity. Among the ecoregions, the
394 SAS region exhibited the highest performance, with `dr` and `bcc` values consistently above 0.75, while NAS, AOL, and PAD
395 demonstrated the lowest performance, with metric values approaching 0.5 for `dr` and 0.7 for `bcc`.

396 On average, the bias-corrected predictions (`bc_pred`) improved regression performance, as indicated by higher `dr` values
397 compared to `mod_pred`. However, this improvement was not uniform, with the MPN ecoregion showing no evident enhancement.
398 The classification results remained consistent between `mod_pred` and `bc_pred`, as the bias correction only altered wet day
399 estimates. Despite improving overall agreement, the bias correction introduced larger errors, as reflected in higher `mae` and
400 `rmse` values (Supplementary Table 3). This suggests that while the bias correction reduces systematic bias, it may overcorrect
401 certain high-precipitation days, leading to larger variability in errors. The observed disparity between `mod_pred` and `bc_pred`
402 underscores the rationale for providing two gap-filling outputs, allowing users to balance overall accuracy (`dr`) against precision
403 in individual predictions (`mae` and `rmse`).

404 Visual analysis of `dr` and `bcc` at the station level (Figure 4a) further validated these findings. The framework performed
405 better in semi-arid, semi-wet, and mixed environments compared to extremely arid or wet regions (Figure 4b). This trend was
406 more pronounced in the classification (`bcc`) than in regression (`dr`), particularly in stations with wet day percentages ranging
407 from 10–50 % in `bcc` or 5–30 % in `dr`. Notably, the bias correction aligned with these patterns, as `dr` values improved within
408 the 5–30 % wet-day range. However, some stations in extreme climates (arid) fell below acceptable regression performance
409 thresholds ($dr < 0.5$).

410 When compared to previous gap-filling studies on a South American scale, similar patterns were observed. Albeit not from
411 a climate diversity perspective (arid to wet), Tang et al.⁷ reported better performance in regions with dense station networks
412 and lower performance in sparse networks. This study also identified higher metric values in dense areas such as SAS and
413 southern EHL. However, significant differences emerged regarding the Andes Cordillera (CAS), where Tang et al.⁷ reported
414 lower performance. This discrepancy may arise from differences in the frameworks: Tang et al.⁷ relied heavily on the temporal
415 correlation of ERA5 for gap-filling, whereas this study employed flexible local models that relax the need for spatiotemporal
416 correlation, addressing known issues with ERA5 precipitation in complex terrains like the Andes⁶⁶.

417 Overall, the gap-filling framework demonstrated strong performance across diverse climates and terrains in South America,
418 effectively addressing challenges related to station sparsity and climatic variability. Nevertheless, limitations remain, particularly
419 in capturing extreme precipitation events, highlighting the need for further refinement and improvement.

420 **Homogeneity**

421 Following the gap-filling process, two datasets were constructed: `obs_mod` and `obs_bc`, representing observations filled with
422 model predictions and bias-corrected predictions, respectively. These datasets underwent a homogenization procedure, resulting
423 in two additional datasets: `hmg_obs_mod` and `hmg_obs_bc`.

424 The table 4 displays the main results of the applied homogenization procedure: detection and adjustment. At the South
425 American scale, approximately 75 % of the time series employed a relative detection approach, leveraging correlations with
426 nearby stations, while the remaining 25 % relied on the absolute approach due to sparse station networks or challenging
427 terrain. Ecoregions such as SAS, EHL, and PPS showed higher applicability of the relative approach, while areas like CAS and
428 GCH required the absolute method. This reliance on the absolute approach in some regions underscores its utility in sparsely
429 populated or complex terrains where relative detection methods struggle.

430 Breakpoints were detected in nearly all time series, with more than 95 % of the series presenting statistical significance.
431 PAD exhibited slightly lower rates of breakpoints (≈ 90 %). The temporal distribution of breakpoints varied across ecoregions
432 (Supplementary Figure 8), with substantial distribution between 1970 and 2010. However, some regions, such as NAS and
433 AOL, displayed frequent breakpoints in the 1970s, whereas GCH and PPS showed peaks in the 2000s. The prominence of
434 breakpoints in indices like `PRCPTOT` and `R1mm` suggests that certain inhomogeneities were introduced during the gap-filling

stage that may not have been fully accounted for in the gap-filling evaluation. Approximately 48 % of daily data in both obs_mod and obs_bc datasets is synthetic, contributing to these patterns.

The adjustments applied through the quantile-matching method varied based on precipitation deciles, with mean values ranging between 2 and 4.5 mm across South America. Most adjustments fell within a -0.5 to 0.5 range (difference between roots of cubic), though extreme values near ± 1 were observed (Supplementary Figure 9), particularly in stations that belong to extremely arid or wet ecoregions. Despite the similarity in the mean adjusted magnitude, we observed slightly higher peaks on the adjusted distribution in PAD, EHL, AOL, PPS, and MPN in obs_mod. The evidence that there were fewer adjusted values in obs_bc could be attributed to the fact that it was bias-corrected (quantile mapping) before the adjustment. This outcome is also supported by a slightly higher number of breakpoints in obs_mod rather than obs_bc (Supplementary Figure 8).

Homogenization impacts on PRCPTOT and R1mm indices were explored at both ecoregional and continental scales (Figures 5 and 6). As expected, we noted a higher impact in PRCPTOT rather than R1mm due to the homogeneity focused on (magnitude) wet days rather than on dry days. Although significant breakpoints were detected in the R1mm indices in the precipitation time series, these were not adjusted. So, in general, the adjusted wet day magnitude was done to follow the wet day distribution. Ecoregions like PAD and SAS exhibited similar patterns in both gap-filled and homogenized datasets, while AOL and EHL demonstrated notable differences. The variability of PRCPTOT demonstrates not only the impact of homogenization and gap-filling but also the distribution of missing data (Supplementary Figure 1).

Regardless of the variability in PRCPTOT (and R1mm) in the different datasets, these revealed important climatological features, such as high and low precipitation years associated with extreme El Niño or La Niña events⁶⁷. Evidence of a climate shift in the 1970s was particularly pronounced in NAS and AOL^{68,69}, where a marked increase in precipitation was observed. Some regions, such as PAD and EHL, displayed clear fluctuations, implying increasing or decreasing trends, although detailed trend analysis was not performed. These results suggest that some important features found in previous work are presented in the reconstructed data. Nevertheless, more in-depth analysis is required for a better understanding.

In summary, the homogenization process primarily impacted wet days, aligning the datasets with known climatological features. Despite its limitations, including challenges in dry/wet day corrections and variations in PRCPTOT and R1mm, the results provide a robust foundation for understanding precipitation variability in South America. The availability of multiple datasets reflects the inherent uncertainty in a complex region such as South America. The plurality of datasets is an advantage as they provide different views of the same variable

Usage Notes

The SC-PREC4SA database is a very useful dataset for a variety of applications. The single database would simplify access to various datasets, hence improving research. Combining data from several sources into a single repository simplifies analysis, maintains consistency, and facilitates regional study and decision-making. This dataset not only provides different outputs targeted to researchers and practitioners, but it also shares the procedures used to create them. This results in a more consistent, traceable, and reproducible resource, increasing its usefulness for scientific and practical purposes.

Despite rigorous quality control, discrepancies in observation time continue to be an issue. Stations frequently have inconsistent or shifting reporting timings¹⁸, and most databases, including this one, lack the hourly data and metadata required to solve this. While existing methods provide possible solutions²⁹, they might alter precipitation intensities. Users should be cautious when interpreting results.

The gap-filling process in this dataset primarily relies on ERA5-Land data. Other gridded precipitation datasets, such as satellite-based products, could also be used. Satellite products were not used in this study because they have mostly available since the 2000s, but their potential to improve precipitation estimation represents an important opportunity. This challenge will be investigated in future studies to improve the accuracy of gap-filled data, particularly in the past and sparse areas.

The dataset was developed using the xgboost model for gap-filling. However, other versions based on glm and rf are also available, allowing users to combine the best estimations from different models using a multi-strategy merge technique. Although this method has been used in previous datasets^{6,7}, we chose not to employ a multi-strategy merging framework, in this case, to ensure uniformity in error and estimation over the entire region utilizing a single model foundation. This results in a more consistent approach to the dataset's application.

The homogenization process used on the dataset may affect the gap-filling results (regression part), particularly for extreme precipitation indices. Although, no abrupt changes were seen in the homogenized PRCPTOT and R1mm mean time series. It is worth noting, however, that we did not test other extreme indices, which may be more sensitive to homogenization processes. Homogenization is especially difficult in locations with sparse station data, where a lack of nearby stations might restrict the detection of inhomogeneities and result in non larger (or too large) adjustments in time series. In locations with low station density, the uncertainties imposed by homogenization may be more noticeable⁵³. These problems highlight the importance of careful interpretation of homogenized data, particularly when working with sparse datasets.

488 Furthermore, uncertainty values were also calculated in the created dataset, which can be important for improving
489 precipitation data analysis. These uncertainty estimations allow for more robust results since they account for spatial and
490 temporal variations. This approach is consistent with prior efforts that employed uncertainty quantification to improve the
491 reliability of precipitation analyses and extreme event evaluations^{41,42}. Users can improve their interpretations of precipitation
492 patterns by incorporating these uncertainty estimates, especially in areas with scarce or inconsistent data.

493 Finally, it is crucial to note that an update to SC-PREC4SA is not currently planned. Nonetheless, because the development
494 of SC-PREC4SA is part of the ANDEX program, there are some initiatives⁷⁰. ANDEX proposes solutions to make high-quality
495 information available throughout all Andean countries that meet their objectives. Policies and strategies for collecting data and
496 establishing observational networks are proposed.

497 Code availability

498 SC-PREC4SA was constructed using the R (v4.1.2) programming language. The entire code used is freely available at GitHub
499 (<https://github.com/adrHuerta/obs-prec-sa2>) under the GNU General Public License v3.0.

500 References

- 501 1. Rodell, M. *et al.* The observed state of the water cycle in the early twenty-first century. *J. Clim.* **28**, 8289–8318,
502 <https://doi.org/10.1175/JCLI-D-14-00555.1> (2015).
- 503 2. L'Ecuyer, T. S. *et al.* The observed state of the energy budget in the early twenty-first century. *J. Clim.* **28**, 8319–8346,
504 <https://doi.org/10.1175/JCLI-D-14-00556.1> (2015).
- 505 3. Sun, Q. *et al.* A review of global precipitation data sets: Data sources, estimation, and intercomparisons. *Rev. Geophys.* **56**,
506 79–107, <https://doi.org/10.1002/2017RG000574> (2018).
- 507 4. Vicente-Serrano, S. M., Beguería, S., López-Moreno, J. I., García-Vera, M. A. & Stepanek, P. A complete daily
508 precipitation database for northeast Spain: reconstruction, quality control, and homogeneity. *Int. J. Climatol.* **30**, 1146–
509 1163, <https://doi.org/10.1002/joc.1850> (2010).
- 510 5. Woldesenbet, T. A., Elagib, N. A., Ribbe, L. & Heinrich, J. Gap filling and homogenization of climatological datasets in the
511 headwater region of the Upper Blue Nile Basin, Ethiopia. *Int. J. Climatol.* **37**, 2122–2140, <https://doi.org/10.1002/joc.4839>
512 (2017).
- 513 6. Tang, G. *et al.* SCDNA: A serially complete precipitation and temperature dataset for North America from 1979 to 2018.
514 *Earth Syst. Sci. Data* **12**, 2381–2409, <https://doi.org/10.5194/essd-12-2381-2020> (2020).
- 515 7. Tang, G., Clark, M. P. & Papalexiou, S. M. SC-Earth: a station-based serially complete earth dataset from 1950 to 2019. *J.*
516 *Clim.* **34**, 6493–6511, <https://doi.org/10.1175/JCLI-D-21-0067.1> (2021).
- 517 8. Garreaud, R. D., Vuille, M., Compagnucci, R. & Marengo, J. Present-day South American climate. *Palaeogeogr.*
518 *Palaeoclimatol. Palaeoecol.* **281**, 180–195, <https://doi.org/10.1016/j.palaeo.2007.10.032> (2009).
- 519 9. Espinoza, J. C. *et al.* Hydroclimate of the Andes part I: main climatic features. *Front. Earth Sci.* **8**, 64, <https://doi.org/10.3389/feart.2020.00064> (2020).
- 520 10. Junquas, C., Li, L., Vera, C., Le Treut, H. & Takahashi, K. Influence of South America orography on summertime
521 precipitation in Southeastern South America. *Clim. Dyn.* **46**, 3941–3963, <https://doi.org/10.1007/s00382-015-2814-8>
522 (2016).
- 523 11. Mejía, J. F. *et al.* Towards a mechanistic understanding of precipitation over the far eastern tropical Pacific and western
524 Colombia, one of the rainiest spots on Earth. *J. Geophys. Res. Atmospheres* **126**, e2020JD033415, <https://doi.org/10.1029/2020JD033415> (2021).
- 525 12. Schween, J. H., Hoffmeister, D. & Löhnert, U. Filling the observational gap in the Atacama Desert with a new network of
526 climate stations. *Glob. Planet. Chang.* **184**, 103034, <https://doi.org/10.1016/j.gloplacha.2019.103034> (2020).
- 527 13. Ferreira, G. W. & Reboita, M. S. A new look into the South America precipitation regimes: observation and forecast.
528 *Atmosphere* **13**, 873, <https://doi.org/10.3390/atmos13060873> (2022).
- 529 14. Bazzanella, A. C., Dereczynski, C., Luiz-Silva, W. & Regoto, P. Performance of CMIP6 models over South America. *Clim.*
530 *Dyn.* **62**, 1501–1516, <https://doi.org/10.1007/s00382-023-06979-1> (2024).
- 531 15. Poveda, G. *et al.* High impact weather events in the Andes. *Front. Earth Sci.* **8**, 162, <https://doi.org/10.3389/feart.2020.00162>
532 (2020).
- 533
- 534

- 535 **16.** Ozturk, U. *et al.* How climate change and unplanned urban sprawl bring more landslides. *Nature* **608**, 262–265,
536 <https://doi.org/10.1038/d41586-022-02141-9> (2022).
- 537 **17.** de los Milagros Skansi, M. *et al.* Warming and wetting signals emerging from analysis of changes in climate extreme
538 indices over South America. *Glob. Planet. Chang.* **100**, 295–307, <https://doi.org/10.1016/j.gloplacha.2012.11.004> (2013).
- 539 **18.** Hunziker, S. *et al.* Identifying, attributing, and overcoming common data quality issues of manned station observations.
540 *Int. J. Climatol.* **37**, 4131–4145, <https://doi.org/10.1002/joc.5037> (2017).
- 541 **19.** Condom, T. *et al.* Climatological and hydrological observations for the South American Andes: in situ stations, satellite,
542 and reanalysis data sets. *Front. Earth Sci.* **8**, 92 (2020).
- 543 **20.** Hunziker, S. *et al.* Effects of undetected data quality issues on climatological analyses. *Clim. Past* **14**, 1–20, <https://doi.org/10.5194/cp-14-1-2018> (2018).
- 545 **21.** Aguayo, R. *et al.* PatagoniaMet: A multi-source hydrometeorological dataset for Western Patagonia. *Sci. Data* **11**, 6,
546 <https://doi.org/10.1038/s41597-023-02828-2> (2024).
- 547 **22.** Xavier, A. C., Scanlon, B. R., King, C. W. & Alves, A. I. New improved Brazilian daily weather gridded data (1961–2020).
548 *Int. J. Climatol.* **42**, 8390–8404, <https://doi.org/10.1002/joc.7731> (2022).
- 549 **23.** Boisier, J. P. *et al.* CR2MET: A high-resolution precipitation and temperature dataset for hydroclimatic research in Chile.
550 In *EGU general assembly conference abstracts*, 19739, <https://doi.org/10.5281/zenodo.7529682> (2018).
- 551 **24.** Fernandez-Palomino, C. A. *et al.* A novel high-resolution gridded precipitation dataset for Peruvian and Ecuadorian water-
552 sheds: Development and hydrological evaluation. *J. Hydrometeorol.* **23**, 309–336, [https://doi.org/10.1175/JHM-D-20-0285.](https://doi.org/10.1175/JHM-D-20-0285.1)
553 **1** (2022).
- 554 **25.** Aybar, C. *et al.* Construction of a high-resolution gridded rainfall dataset for Peru from 1981 to the present day. *Hydrol.*
555 *Sci. J.* **65**, 770–785, <https://doi.org/10.1080/02626667.2019.1649411> (2020).
- 556 **26.** Huerta, A., Lavado-Casimiro, W. & Felipe-Obando, O. High-resolution gridded hourly precipitation dataset for Peru
557 (PISCOp_h). *Data Brief* **45**, 108570, <https://doi.org/10.1016/j.dib.2022.108570> (2022).
- 558 **27.** Schamm, K. *et al.* Global gridded precipitation over land: A description of the new GPCP First Guess Daily product.
559 *Earth Syst. Sci. Data* **6**, 49–60, <https://doi.org/10.5194/essd-6-49-2014> (2014).
- 560 **28.** Funk, C. *et al.* The climate hazards infrared precipitation with stations—a new environmental record for monitoring
561 extremes. *Sci. data* **2**, 1–21, <https://doi.org/10.1038/sdata.2015.66> (2015).
- 562 **29.** Beck, H. E. *et al.* MSWEP V2 global 3-hourly 0.1 precipitation: methodology and quantitative assessment. *Bull. Am.*
563 *Meteorol. Soc.* **100**, 473–500, <https://doi.org/10.1175/BAMS-D-17-0138.1> (2019).
- 564 **30.** Tang, G., Clark, M. P. & Papalexiou, S. M. EM-Earth: the ensemble meteorological dataset for planet earth. *Bull. Am.*
565 *Meteorol. Soc.* **103**, E996–E1018, <https://doi.org/10.1175/BAMS-D-21-0106.1> (2022).
- 566 **31.** Menne, M. J., Durre, I., Vose, R. S., Gleason, B. E. & Houston, T. G. An overview of the global historical climatology
567 network-daily database. *J. atmospheric oceanic technology* **29**, 897–910, <https://doi.org/10.1175/JTECH-D-11-00103.1>
568 (2012).
- 569 **32.** Lewis, E. *et al.* GSDR: a global sub-daily rainfall dataset. *J. Clim.* **32**, 4715–4729, [https://doi.org/10.1175/JCLI-D-18-0143.](https://doi.org/10.1175/JCLI-D-18-0143.1)
570 **1** (2019).
- 571 **33.** Griffith, G. E., Omernik, J. M. & Azevedo, S. H. Ecological classification of the Western Hemisphere. *Unpubl. Rep.* **49**,
572 http://ecological-regions.info/html/sa_eco.htm (1998).
- 573 **34.** Muñoz-Sabater, J. *et al.* ERA5-Land: A state-of-the-art global reanalysis dataset for land applications. *Earth system*
574 *science data* **13**, 4349–4383, <https://doi.org/10.5194/essd-13-4349-2021> (2021).
- 575 **35.** Wu, X., Su, J., Ren, W., Lü, H. & Yuan, F. Statistical comparison and hydrological utility evaluation of ERA5-Land and
576 IMERG precipitation products on the Tibetan Plateau. *J. Hydrol.* **620**, 129384, [https://doi.org/10.1016/j.jhydrol.2023.](https://doi.org/10.1016/j.jhydrol.2023.129384)
577 **129384** (2023).
- 578 **36.** Amatulli, G. *et al.* A suite of global, cross-scale topographic variables for environmental and biodiversity modeling. *Sci.*
579 *data* **5**, 1–15, <https://doi.org/10.1038/sdata.2018.40> (2018).
- 580 **37.** Danielson, J. J. & Gesch, D. B. Global multi-resolution terrain elevation data 2010 (GMTED2010), [https://doi.org/10.](https://doi.org/10.5066/F7J38R2N)
581 **5066/F7J38R2N** (2011).
- 582 **38.** Applequist, S., Durre, I. & Vose, R. The Global Historical Climatology Network Monthly Precipitation Dataset, Version 4.
583 *Sci. Data* **11**, 633, <https://doi.org/10.1038/s41597-024-03457-z> (2024).

- 584 **39.** Hamada, A., Arakawa, O. & Yatagai, A. An automated quality control method for daily rain-gauge data. *Glob. Environ.*
585 *Res* **15**, 183–192, http://www.airies.or.jp/journal_15-2eng.html (2011).
- 586 **40.** Serrano-Notivoli, R., de Luis, M. & Beguería, S. An R package for daily precipitation climate series reconstruction.
587 *Environ. modelling & software* **89**, 190–195, <https://doi.org/10.1016/j.envsoft.2016.11.005> (2017).
- 588 **41.** Serrano-Notivoli, R., de Luis, M., Saz, M. Á. & Beguería, S. Spatially based reconstruction of daily precipitation
589 instrumental data series. *Clim. Res.* **73**, 167–186, <https://doi.org/10.3354/cr01476> (2017).
- 590 **42.** Serrano-Notivoli, R., Beguería, S., Saz, M. Á., Longares, L. A. & de Luis, M. SPREAD: a high-resolution daily gridded
591 precipitation dataset for Spain—an extreme events frequency and intensity overview. *Earth Syst. Sci. Data* **9**, 721–738,
592 <https://doi.org/10.5194/essd-9-721-2017> (2017).
- 593 **43.** Škrk, N. *et al.* SLOCLIM: a high-resolution daily gridded precipitation and temperature dataset for Slovenia. *Earth Syst.*
594 *Sci. Data* **13**, 3577–3592, <https://doi.org/10.5194/essd-13-3577-2021> (2021).
- 595 **44.** Centella-Artola, A. *et al.* A new long term gridded daily precipitation dataset at high-resolution for Cuba (CubaPrec1).
596 *Data Brief* **48**, 109294, <https://doi.org/10.1016/j.dib.2023.109294> (2023).
- 597 **45.** Chen, T. & Guestrin, C. Xgboost: A scalable tree boosting system. In *Proceedings of the 22nd acm sigkdd international*
598 *conference on knowledge discovery and data mining*, 785–794, <https://dl.acm.org/doi/10.1145/2939672.2939785> (2016).
- 599 **46.** Chen, T. *et al.* Extreme Gradient Boosting [R Package Xgboost Version 1.2. 0.1]. In *Proc. ACM SIGKDD Int. Conf. Knowl.*
600 *Discov. Data Min.*, 13–17, <https://cran.r-project.org/web/packages/xgboost/index.html> (2020).
- 601 **47.** Huerta, A. *et al.* PISCOeo_pm, a reference evapotranspiration gridded database based on FAO Penman-Monteith in Peru.
602 *Sci. data* **9**, 328, <https://doi.org/10.1038/s41597-022-01373-8> (2022).
- 603 **48.** Huerta, A. *et al.* High-resolution grids of daily air temperature for Peru-the new PISCOt v1. 2 dataset. *Sci. data* **10**, 847,
604 <https://doi.org/10.1038/s41597-023-02777-w> (2023).
- 605 **49.** Horn, J. L. A rationale and test for the number of factors in factor analysis. *Psychometrika* **30**, 179–185, <https://doi.org/10.1007/BF02289447> (1965).
- 607 **50.** Gonzalez-Hidalgo, J. C., Peña-Angulo, D., Brunetti, M. & Cortesi, N. MOTEDAS: a new monthly temperature database for
608 mainland Spain and the trend in temperature (1951–2010). *Int. J. Climatol.* **35**, 4444–4463, <https://doi.org/10.1002/joc.4298>
609 (2015).
- 610 **51.** Venema, V. K. *et al.* Benchmarking homogenization algorithms for monthly data. *Clim. Past* **8**, 89–115, <https://doi.org/10.5194/cp-8-89-2012>
611 (2012).
- 612 **52.** Guijarro, J. A. *et al.* Homogenization of monthly series of temperature and precipitation: Benchmarking results of the
613 MULTITEST project. *Int. J. Climatol.* **43**, 3994–4012, <https://doi.org/10.1002/joc.8069> (2023).
- 614 **53.** Gubler, S. *et al.* The influence of station density on climate data homogenization. *Int. journal climatology* **37**, 4670–4683,
615 <https://doi.org/10.1002/joc.5114> (2017).
- 616 **54.** Squintu, A. A., van der Schrier, G., Brugnara, Y. & Klein Tank, A. Homogenization of daily ECA&D temperature series.
617 *Int. journal climatology* **39**, 1243–1261, <https://doi.org/10.1002/joc.5874> (2018).
- 618 **55.** Brugnara, Y., Good, E., Squintu, A. A., van der Schrier, G. & Brönnimann, S. The EUSTACE global land station daily air
619 temperature dataset. *Geosci. Data J.* **6**, 189–204, <https://doi.org/10.1002/gdj3.81> (2019).
- 620 **56.** Hurtado, S. I., Zaninelli, P. G. & Agosta, E. A. A multi-breakpoint methodology to detect changes in climatic time
621 series. An application to wet season precipitation in subtropical Argentina. *Atmospheric Res.* **241**, 104955, <https://doi.org/10.1016/j.atmosres.2020.104955> (2020).
- 623 **57.** Lund, R. B., Beaulieu, C., Killick, R., Lu, Q. & Shi, X. Good practices and common pitfalls in climate time series
624 changepoint techniques: A review. *J. Clim.* **36**, 8041–8057, <https://doi.org/10.1175/JCLI-D-22-0954.1> (2023).
- 625 **58.** Caussinus, H. & Mestre, O. Detection and correction of artificial shifts in climate series. *J. Royal Stat. Soc. Ser. C: Appl.*
626 *Stat.* **53**, 405–425, <https://doi.org/10.1111/j.1467-9876.2004.05155.x> (2004).
- 627 **59.** Brugnara, Y., McCarthy, M. P., Willett, K. M. & Rayner, N. A. Homogenization of daily temperature and humidity series
628 in the UK. *Int. J. Climatol.* **43**, 1693–1709, <https://doi.org/10.1002/joc.7941> (2023).
- 629 **60.** Tomas-Burguera, M., Vicente-Serrano, S. M., Beguería, S., Reig, F. & Latorre, B. Reference crop evapotranspiration
630 database in Spain (1961–2014). *Earth Syst. Sci. Data* **11**, 1917–1930, <https://doi.org/10.5194/essd-11-1917-2019> (2019).
- 631 **61.** Gutmann, E. *et al.* An intercomparison of statistical downscaling methods used for water resource assessments in the
632 United States. *Water Resour. Res.* **50**, 7167–7186, <https://doi.org/10.1002/2014WR015559> (2014).

- 633 **62.** Berg, P. *et al.* Robust handling of extremes in quantile mapping—“Murder your darlings”. *Geosci. Model. Dev.* **17**,
634 8173–8179, <https://doi.org/10.5194/gmd-17-8173-2024> (2024).
- 635 **63.** Huerta, A., Serrano-Notivoli, R. & Brönnimann, S. A serially complete daily precipitation dataset for South America
636 (SC-PREC4SA), <https://doi.org/10.6084/m9.figshare.c.7588178> (2024).
- 637 **64.** Cinkus, G. *et al.* When best is the enemy of good—critical evaluation of performance criteria in hydrological models.
638 *Hydrol. Earth Syst. Sci.* **27**, 2397–2411, <https://doi.org/10.5194/hess-27-2397-2023> (2023).
- 639 **65.** Thölke, P. *et al.* Class imbalance should not throw you off balance: Choosing the right classifiers and performance metrics
640 for brain decoding with imbalanced data. *NeuroImage* **277**, 120253, <https://doi.org/10.1016/j.neuroimage.2023.120253>
641 (2023).
- 642 **66.** Imfeld, N. *et al.* Summertime precipitation deficits in the southern Peruvian highlands since 1964. *Int. J. Clim.* **39**,
643 4497–4513, <https://doi.org/10.1002/joc.6087> (2019).
- 644 **67.** Cai, W. *et al.* Climate impacts of the El Niño–southern oscillation on South America. *Nat. Rev. Earth & Environ.* **1**,
645 215–231, <https://doi.org/10.1038/s43017-020-0040-3> (2020).
- 646 **68.** Carvalho, L. M., Jones, C., Silva, A. E., Liebmann, B. & Silva Dias, P. L. The South American monsoon system and the
647 1970s climate transition. *international J. Climatol.* **31**, 1248–1256, <https://doi.org/10.1002/joc.2147> (2011).
- 648 **69.** Jacques-Coper, M. & Garreaud, R. D. Characterization of the 1970s climate shift in South America. *Int. J. Climatol.* **35**,
649 2164–2179, <https://doi.org/10.1002/joc.4120> (2015).
- 650 **70.** IANIGLA-CONICET & CR². Observatorio de Nieve en los Andes de Argentina y Chile. <https://observatorioandino.com>.
651 [Accessed 04-12-2024].

652 **Acknowledgements**

653 AH acknowledges support from Swiss Government Excellence Scholarships for Foreign Scholars (ESKAS-Nr: 2023.0404).
654 RSN is supported by grant RYC2021-034330-I funded by MCIN/AEI/10.13039/501100011033 and by "European Union
655 NextGenerationEU/PRTR". The authors are grateful for the availability of precipitation data from global databases (LACA&D
656 and GHCNd), as well as seven NMHSs (BDHI-Argentina, SENAMHI-Bolivia, INMET-Brazil, DMC-Chile, IDEAM-Colombia,
657 INAMHI-Ecuador, and SENAMHI-Peru). The creation of SC-PREC4SA has been achieved as part of the ANDEX program
658 (<https://www.andex-rhp.org/>), which is a Regional Hydroclimate Project (RHP) of the GEWEX Hydroclimatology Panel (GHP).

659 **Author contributions statement**

660 AH developed the dataset methodology and creation in consultation with RSN and SB. RSN collected the raw observed dataset.
661 AH prepared the data, conducted the experiments, and wrote the first draft of the manuscript. All authors were involved in
662 discussions with regard to data development and quality, and all reviewed the manuscript.

663 **Competing interests**

664 The authors declare no competing interests.

Figures & Tables

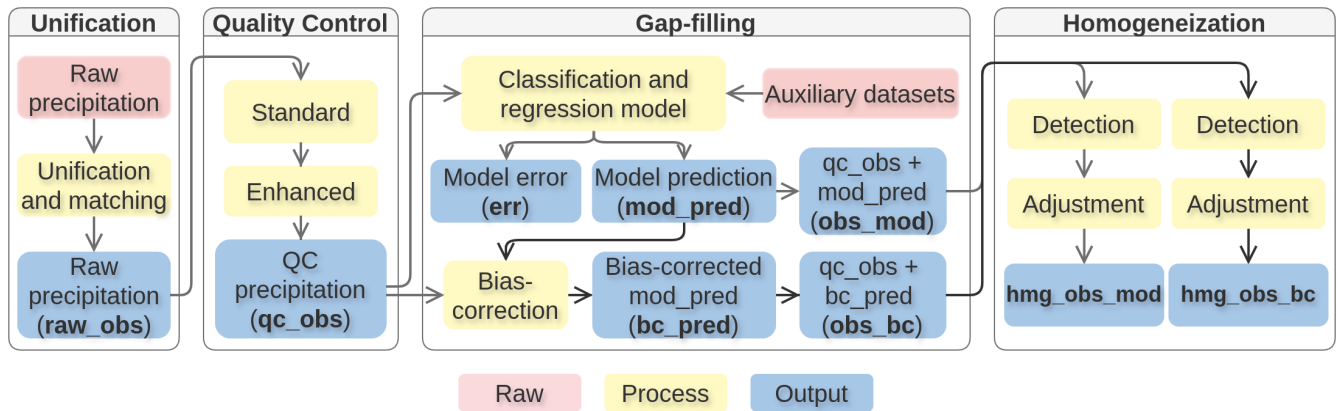


Figure 1. Schematic overview of the development of a serially complete dataset of daily precipitation for South America (SC-PREC4SA). Raw data, related processes, and main output files are specified.

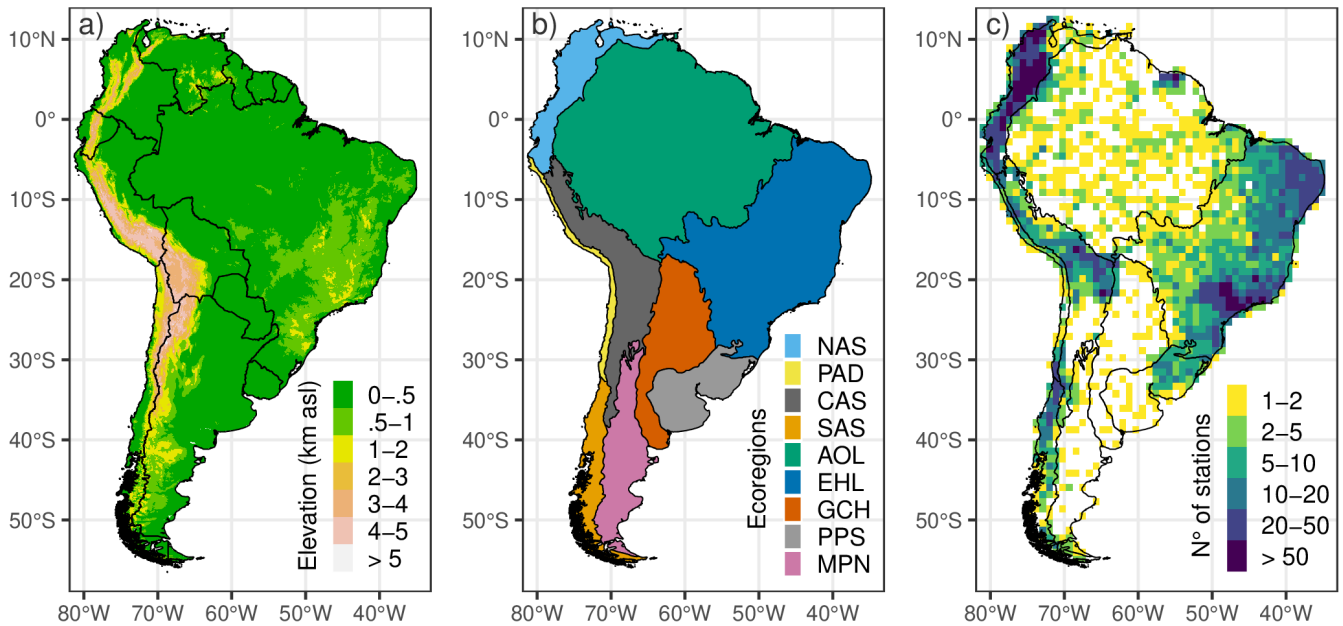


Figure 2. (a) Study area of contiguous South America displaying the elevation and countries as black lines. (b) Ecoregions of contiguous South America: Northern Andes (NAS), Peruvian/Atacaman Deserts (PAD), Central Andes (CAS), Southern Andes (SAS), Amazonian-Orinocan Lowland (AOL), Eastern Hihglands (EHL), Gran Chaco (GCH), Pampas (PPS) and Monte-Patagonian (MPN). (c) The number of raw collected stations in a grid size of 0.9° from 1960 to 2015; black lines represent the ecoregions.

Standard Quality Control	Ecoregions									South America
	NAS	PAD	CAS	SAS	AOL	EHL	GCH	PPS	MPN	
SQC-01	154 0	0 0	62 0	22 0	57 0	90 0	0 0	9 0	0 0	394 0
SQC-02	9134 0.02	730 0.03	1461 0.01	1460 0.03	4382 0.05	18266 0.04	0 0	365 0.02	366 0.05	36164 0.03
SQC-03	15538 0.04	0 0	1102 0.01	852 0.02	2512 0.03	3066 0.01	62 0.01	0 0	60 0.01	23192 0.02
SQC-04	22220 0.06	56 0	2658 0.02	1038 0.02	3890 0.05	6529 0.01	118 0.02	62 0	60 0.01	36631 0.03
SQC-05	7919 0.02	684 0.03	6736 0.06	3319 0.07	1144 0.01	19196 0.04	323 0.05	205 0.01	725 0.1	40251 0.04
SQC-06	450 0	25 0	136 0	33 0	126 0	699 0	10 0	26 0	15 0	1520 0
SQC-07	1819 0	196 0.01	1445 0.01	235 0	440 0.01	849 0	64 0.01	74 0	520 0.07	5642 0
SQC	57234 0.15	1691 0.07	13600 0.13	6959 0.14	12551 0.15	48695 0.1	577 0.08	741 0.03	1746 0.24	143794 0.13

Table 1. Summary of standard quality control (SQC) process. The number of flagged data and percentage (compared to the total daily data) are displayed for each quality control step by ecoregion. The last column (row) shows the same, but encompassing the entire South America area (SQC steps). The bottom right corner displays the unified results.

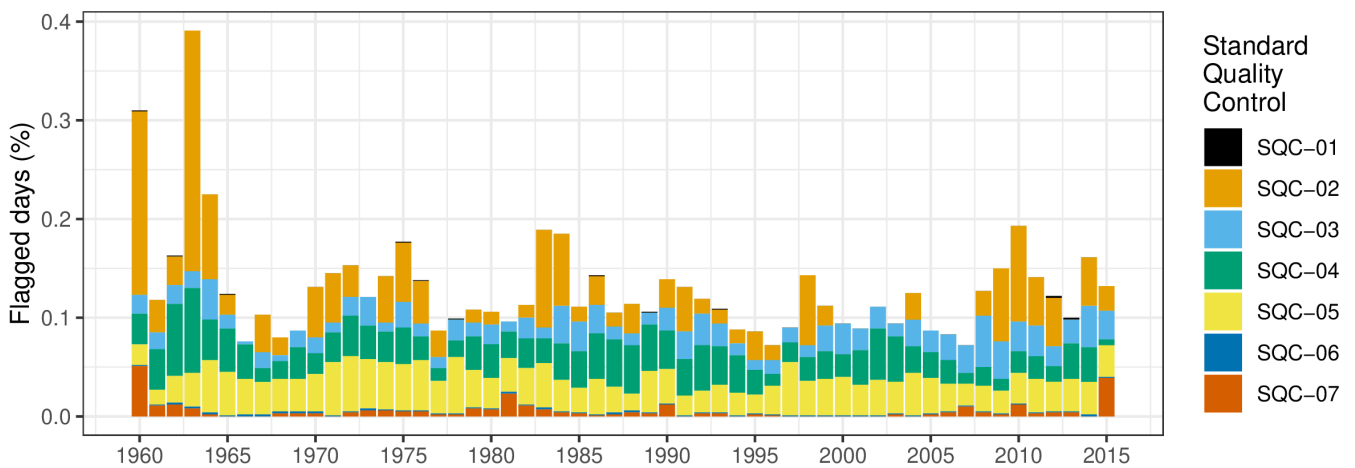


Figure 3. Flagged daily data (%) by each standard quality control step from 1960 to 2015 in South America.

Enhanced Quality Control	Ecoregions									South America
	NAS	PAD	CAS	SAS	AOL	EHL	GCH	PPS	MPN	
EQC-01	58	0	91	20	10	39	8	2	8	236
	2.02	0	11.67	5.13	1.27	0.83	13.56	0.89	12.9	2.34
EQC-02	249	0	144	40	169	1340	17	63	3	2025
	8.65	0	18.46	10.26	21.53	28.42	28.81	28	4.84	20.07
EQC-03	46	0	6	7	20	26	1	3	2	111
	1.6	0	0.77	1.79	2.55	0.55	1.69	1.33	3.23	1.1
EQC-04	261	0	124	158	80	848	9	38	13	1531
	9.07	0	15.9	40.51	10.19	17.99	15.25	16.89	20.97	15.18
EQC	560	0	271	186	225	1678	26	79	18	3043
	19.46	0	34.74	47.69	28.66	35.59	44.07	35.11	29.03	30.16

Table 2. Summary of enhanced quality control (EQC) process. The number of stations (Level = 2) and percentage (compared to the total stations) are displayed for each quality control step by ecoregion. The last column (row) shows the same, but encompassing the entire South America area (EQC steps). The bottom right corner displays the unified results.

Statistical Metrics	Model output	Ecoregions									South America
		NAS	PAD	CAS	SAS	AOL	EHL	GCH	PPS	MPN	
dr	mod_pred	0.53	0.58	0.61	0.74	0.53	0.63	0.64	0.68	0.63	0.62
	bc_pred	0.58	0.61	0.64	0.75	0.57	0.66	0.65	0.7	0.63	0.64
bcc	mod_pred	0.7	0.68	0.73	0.81	0.71	0.74	0.73	0.76	0.68	0.73
	bc_pred	0.7	0.68	0.72	0.81	0.71	0.74	0.72	0.75	0.68	0.72

Table 3. Summary of gap-filling evaluation metrics: refined index of agreement (dr) and balanced accuracy (bcc). The mean value is displayed for each metric by model output (model prediction without [mod_pred] and with bias-correction [bc_pred]) and ecoregion. The last column displays the mean value at the South American scale. In bold when the statistical metric is best depending on the model output.

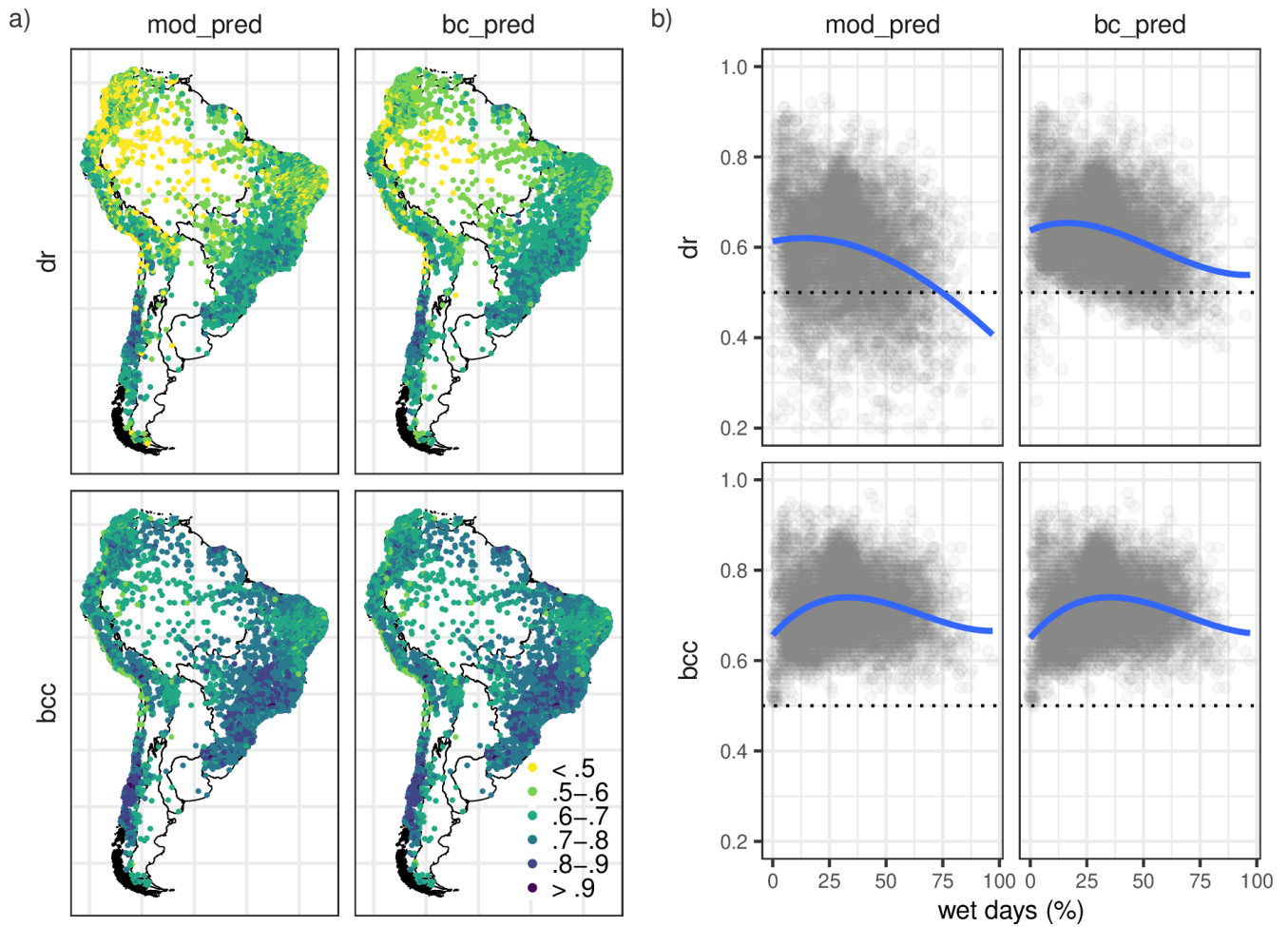


Figure 4. a) Spatial distribution gap-filling evaluation metrics: refined index of agreement (dr) and balanced accuracy (bcc). b) Relationship between evaluation metrics versus the percentage of wet days (wet days). The blue line in b) represents the trend line between both variables computed using splines. The dotted line in b) displays the 0.5 value in both metrics. Statistical metrics are divided by model output (model prediction without [mod_pred] and with bias correction [bc_pred]).

Homogenization process	Database	Ecoregions									South America
		NAS	PAD	CAS	SAS	AOL	EHL	GCH	PPS	MPN	
Relative detection (%)	obs_mod	55.34	52.31	42.63	81.9	70.49	95.47	41.03	98.09	56	74.52
	obs_bc	63.68	53.85	49.22	86.81	71.16	96.04	41.03	98.09	56	78.25
Significant detection (%)	obs_mod	99.01	90.77	97.4	100	99.16	99.55	97.44	100	96	98.96
	obs_bc	98.93	88.21	96.01	98.47	99.83	99.25	92.31	100	98	98.62
Mean adjustment (mm)	obs_mod	3.33	3.43	1.92	3.09	3.58	3.09	3.94	3.62	1.43	3.19
	obs_bc	3.37	3.69	2.15	3.08	4.38	4.21	4.03	4.2	2.01	3.74

Table 4. Summary of the homogenization process for database and ecoregions: percentage of stations where the relative detection test was performed (relative detection), percentage of stations where a significant break detection was found (significant detection), and, the mean value of the applied adjustment (mean adjustment). The last column (row) shows the same but encompasses the entire South American area.

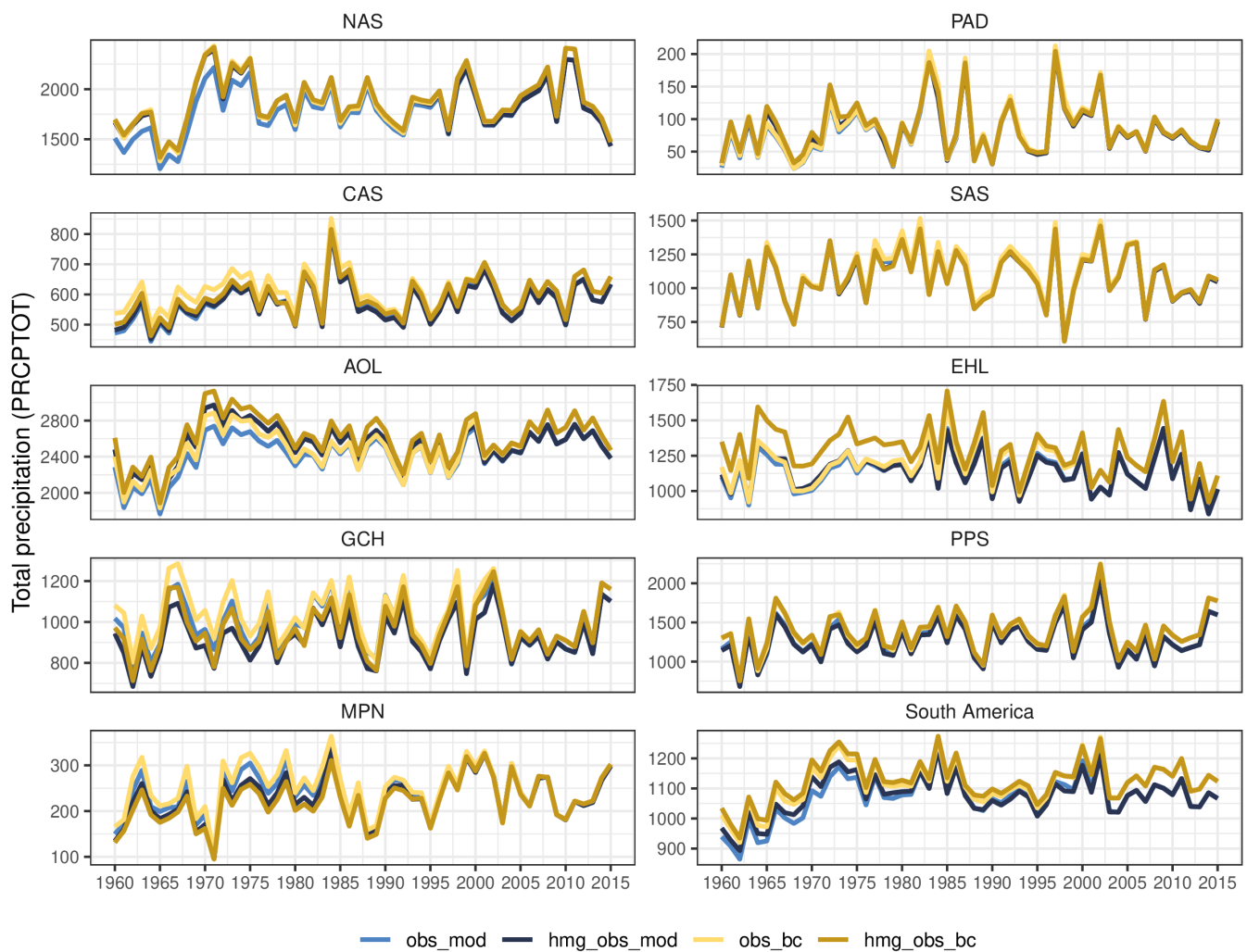


Figure 5. Mean time series of the total precipitation (mm/year) by ecoregions and South America after (observed data plus model prediction without [hmg_obs_mod] and with bias-correction [hmg_obs_bc]) and before (observed data plus model prediction without [obs_mod] and with bias-correction [obs_bc]) the homogenization process.

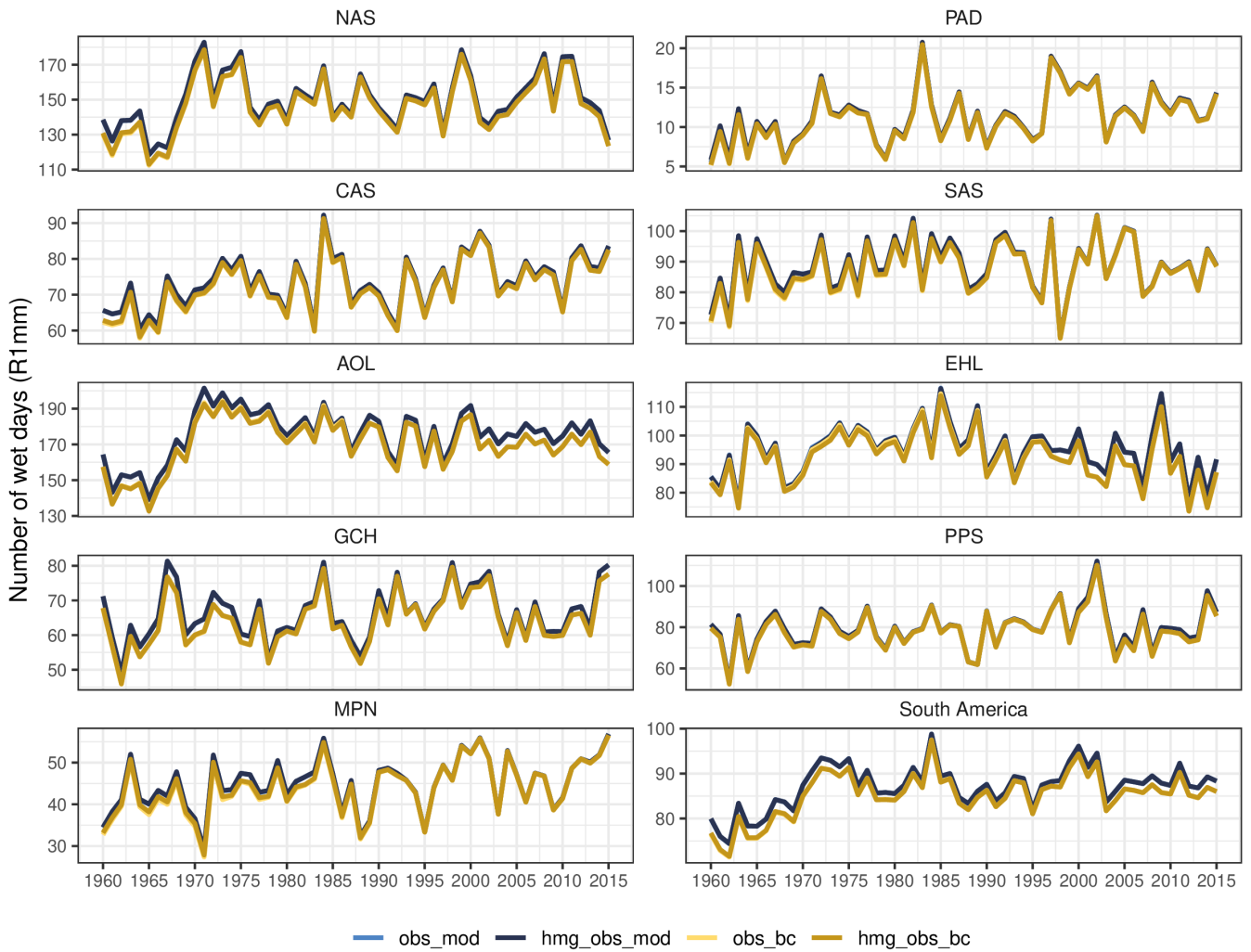


Figure 6. Mean time series of the number of wet days (days/year) by ecoregions and South America after (observed data plus model prediction without [hmg_obs_mod] and with bias-correction [hmg_obs_bc]) and before (observed data plus model prediction without [obs_mod] and with bias-correction [obs_bc]) the homogenization process.