Preprint

GelCam: Visualizing Episodic Sinking Particle Flux via a Polyacrylamide Gel-Based Sediment Trap

Yixuan Song^{*1}, Melissa Omand¹, Colleen A. Durkin², Margaret L. Estapa³, and Ken O. Buesseler⁴ ¹Graduate School of Oceanography, University of Rhode Island, Narragansett, RI, USA ²Monterey Bay Aquarium Research Institute, Moss Landing, CA, USA ³School of Marine Sciences, Darling Marine Center, University of Maine, Walpole, ME, USA ⁴Woods Hole Oceanographic Institution, Woods Hole, MA, USA

*Corresponding author: Yixuan Song, <u>yixuan.song@uri.edu</u>

Please note that the manuscript has not undergone peer-review and is not accepted for publication at this time. Subsequent versions of this manuscript may have slightly different content. If accepted, the final version of this manuscript will be available via the Peer-reviewed Publication DOI link on the right-handside of this webpage. Please feel free to contactany of the authors; we welcome your feedback on our contribution to the literature.

GelCam: Visualizing Episodic Sinking Particle Flux via a Polyacrylamide Gel-Based Sediment Trap

3	Yixuan Song ^{*1} , Melissa Omand ¹ , Colleen A. Durkin ² , Margaret L. Estapa ³ , and Ken O. Buesseler ⁴
4	¹ Graduate School of Oceanography, University of Rhode Island, Narragansett, RI, USA
5	² Monterey Bay Aquarium Research Institute, Moss Landing, CA, USA
6	³ School of Marine Sciences, Darling Marine Center, University of Maine, Walpole, ME, USA
7	⁴ Woods Hole Oceanographic Institution, Woods Hole, MA, USA
8	
9	*Corresponding author: Yixuan Song, <u>yixuan.song@uri.edu</u>
10	Running Head: GelCam: Visualizing Episodic Particle Flux
11	Keywords: Biological carbon pump, sediment traps, marine snow, fecal pellets, aggregates,
12	particles
13	

14

15 ABSTRACT

While particle-intercepting traps remain a dominant method for quantifying the contribution of 16 sinking particles to the biological carbon pump, fluxes are typically integrated over days to months. 17 Observations of time-varying particle flux over shorter durations are very limited. To this end, we 18 prototyped a camera system called "GelCam" which captures a rapid time-lapse image sequence 19 of particles that settle into a polyacrylamide gel layer located at the base of a sediment trap tube. 20 21 Here, we describe the system design, post image processing, and results from nine deployments during the EXPORTS campaigns in 2018 and 2021 in the North Pacific and Atlantic ocean. 22 Because wave-driven oscillations of the surface-tethered traps produced a lateral motion of the 23 24 settled particles, we applied a cross-correlation method for tracking individual particles over time. All particles were subsequently classified into one of six categories based on visual traits, and then 25 quantified into a particulate organic carbon (POC) flux. Using this image-based approach, we are 26 able to distinguish differences in depth-based transfer efficiency among groups, and detect a diel 27 variation in fecal pellet flux. Additionally, the GelCam resolved flux events on timescales shorter 28 than days, allowing for the investigation of covariance among different particle types over short 29 timescales. Paired with the direct recovery of samples to quantify carbon content and high 30 resolution particle images, this approach will enhance our ability to resolve sinking "events" that 31 32 occur episodically and may be missed when integrating over a traditional trap deployment.

33

34 INTRODUCTION

The ocean sink provides a net uptake of approximately $2.8 \cdot 10^{15}$ g of carbon annually, which 35 accounts for more than 25% of anthropogenic CO₂ emissions (Friedlingstein et al., 2023). Among 36 processes of marine carbon sequestration, sinking particles transport ~ 10^{15} g/yr of carbon from 37 the surface ocean to the deep ocean (Middelburg, 2011; Muller-Karger, 2005; Henson et al., 2011). 38 39 Knowing the biological carbon pump is critical for predicting climate change effects. There are a few challenges to overcome before we can accurately model the ocean carbon uptake process. 40 First, the carbon transport is driven by complex biophysical and chemical processes (Nowicki et 41 al., 2022). Photosynthesis by phytoplankton (Basu and Mackey, 2018), consumption and 42 production by zooplankton and fish (Cavan et al., 2017), and sinking of marine particles (De La 43 Rocha and Passow, 2007) contribute to the high uncertainty in modeling the carbon transport. 44 Additionally, aggregation (Burd and Jackson, 2009; Alldredge and Gotschalk, 1988), 45 disaggregation (Briggs et al., 2020; Song et al., 2023), and microbial degradation occurring on 46 particles (Collins et al., 2015), alter the particle size and contribute to the high variability in the 47 sinking velocity of marine particles (McDonnell and Buesseler, 2010). Moreover, spatial and 48 temporal variability further complicates the modeling of global biogeochemical cycles (Piao et al., 49 2013; SierraI et al., 2007). These challenges along with observational limitations in remote areas 50 and deep sea environments, substantially constrains our understanding of carbon flux within the 51 52 ocean.

53 Sediment traps provide a direct measurement of sinking marine particles, as the downward carbon 54 flux can be retrieved from the total organic matter captured over a period of time within a fixed 55 collection area of a tube or a funnel (Buesseler et al., 2007). Traditional sediment trap 56 measurements provide an integrated carbon flux sample collected over time scales ranging from days to months (Buesseler et al., 2007). Post analyses are necessary for most traditional sediment trap measurements. For example, Muller and Suess (1979) applied elemental combustion analysis to measure the organic carbon content in sediment trap samples. Durkin et al. (2021) used microscopic imaging techniques to analyze particles preserved in the polyacrylamide gel. In addition to being labor-demanding, the traditional sediment trap method is limited in its ability to assess episodic variations on timescales of hours.

Classifying sinking particle types is important for accurately predicting the carbon flux because 63 each particle type contributes to carbon transport in its unique way. For example, zooplankton 64 fecal pellets have been observed to play a major role in controlling carbon flux in some studies 65 (e.g., (Cavan et al., 2015; Turner, 2002)). More specifically, Durkin et al. (2021) and Steinberg et 66 al. (2023) emphasized the role of salp fecal pellets in the Subarctic Northeast Pacific Ocean. 67 Additionally, the importance of large marine snow particles is highlighted in other studies, by 68 which aggregation enhances the transport of organic matter (Alldredge and Silver, 1988; Fowler 69 70 and Knauer, 1986). Contributions of phytoplankton (Durkin et al., 2016), mini fecal pellets (Gowing and Silver, 1985), and physical subduction (Omand et al., 2015) that includes particles 71 cannot be ignored as well. However, the composition of sinking particle flux varies in space and 72 73 time, leading to challenges in their classification. Developing a time-resolved particle classification system for sediment trap samples is needed. Using an imaging system, such as 74 75 optical sediment traps, not only simplifies sample processing but enables the evaluation of flux 76 components across different particle types.

To better understand and constrain temporal variability in particle fluxes and identify the ecological sources that drive these variations, we have developed a novel, low-cost, and easy-todeploy instrument, named "GelCam". This instrument allows for a rapid evaluation without

4/57

extensive post analysis, and chronological diagnosis of particle identities. Here we demonstrate 80 the GelCam design and analyze episodic particulate organic carbon fluxes from ten datasets 81 collected from two NASA EXPORTS (EXport Processes in the Ocean from Remote Sensing) 82 field campaigns. The image post-processing contains a particle tracking algorithm and particle 83 classification, which provides the arrival time and morphology of each particle. The GelCam with 84 85 the framework of imaging processing can be applied to quantify the particle flux using two methods, with one based on cumulative particle areas and the other one based on the particle 86 tracking and classification results. Here we model the particulate organic carbon from particle 87 images and compare the fluxes between GelCam and other methods. We also present the diel 88 changes of fecal pellet fluxes and investigate flux events attributable to different particle types. 89 We finally discuss the benefits of using the GelCam to extract quantitative carbon flux and 90 facilitate modeling of the biological carbon pump. 91

92

93 MATERIALS AND PROCEDURES

94 GelCam design

The GelCam is a time-lapse camera system designed to pair with a particle-intercepting sediment trap and it has two iterations that have already been deployed. In the current design, we used a cylindrical trap with an outer diameter of 127 mm, and a total length of 700 mm (Figure 1 (A)). A lid closure method used a bungee that was mounted to either a horizontal bar or to the side wall of the tube (since 2021). The open bottom of the tube was attached to the camera system with opposed flanges and an o-ring face seal. Before sealing the camera to the tube, a shallow container with a transparent bottom was filled with a ~ 10 mm thick layer of polyacrylamide gel and placed at the base of the trap, and served to retain deposited particles.



103

Figure 1. Schematic diagram of the GelCam system: (A) overview of the sediment trap with theGelCam attached to the bottom, and (B) detailed cross-section view within the enclosure.

106

The GelCam housing was composed of an optically clear one inch thick acrylic top lid, a cylindrical acetal resin (Delrin) tube and a acetal resin bottom lid (Figure 1(A)). The housing length was 177.8 mm during EXPORTS 2018. An upgrade in 2021 increased the length to 203.2 mm to provide more inner space and a larger field of view (FOV). The top lid contained a potted LED ring with a width of 4 mm and an outer diameter of 120 mm using clear epoxy resin, located on the upper surface to eliminate direct light or reflections and allowing only diffuse, scattered light to enter the camera. Brackets fabricated by 3D printing techniques and designed to secure

the batteries and electronics, were integrated within the housing. We used 7 packs of 12 V 114 (nominal) stacks, each consisting of 4 CR123A batteries, yielding 134 Watt-hrs. The batteries 115 were used to power a Arduino Nano micro controller, a Raspberry Pi Zero WiFi single-board 116 computer, a Raspberry Pi camera module, and the LED ring. The low-power Arduino was used 117 to duty cycle the R-Pi Zero, enabling us the sustain weeks-long deployments with photos every 118 119 20 minutes. The camera was mounted facing upwards toward the gel container with a viewing axis aligned with the central axis of the sediment trap, enabling us to image nearly the entire base 120 of the gel-filled cup. 121

The camera module had a sensor resolution of 2592×1944 pixels. The lens used in the 2018 122 expedition provided a FOV of approximately 55.2 mm \times 41.4 mm which translated to a pixel 123 resolution of 21.3 μ m per pixel. The taller housing used since 2021 expanded the FOV but 124 provided a smaller pixel resolution of 25.5 μ m per pixel. We calibrated the cameras before each 125 deployment, with an optical depth trained on the deepest layer of the gel where the particles 126 127 accumulated. The Arduino Nano was programmed to duty cycle the R-Pi at 20 minute intervals. However due to uncertainties in the timer built into the ATMEGA328P micro controller, the actual 128 time interval varied by deployments within a few minutes. Fortunately, we also integrated a 129 130 highly accurate real time clock, which logged the time stamp for each image. Later versions of this system have replaced the Arduino-based micro controller for a more accurate and lower 131 132 power option. Images were archived on a 64GB SD card within the R-Pi Zero. This configuration 133 had sufficient power to take photos at 20 minute intervals for 8 days. The operational depth rating of the housing was 300m. 134

135 Cruises and sampling platform

Gelcams were deployed in the subarctic North Pacific (NP) and North Atlantic (NA) during two 136 EXPORTS cruises. The former cruise conducted three deployments aboard the *R/V Roger Revelle* 137 near Ocean Station Papa (50°N, 145°W) during three repeating sampling cycles identified as 138 Deployment 1, 2, and 3, from August 15 to September 5 2018 (Table 1) (other NASA EXPORTS 139 papers referred the deployment duration as "epoch"). The deployment plan is presented by Siegel 140 141 et al. (2021) and sediment trap details are presented by Estapa et al. (2021). On a surface-tethered trap (STT) array, we installed the GelCam at three near-surface traps, corresponding to 95, 145, 142 and 195 m in Deployment 1. Weather-induced damage required subsequent repairs, causing a 10 143 m increase in depth during Deployment 2 and 3. The second EXPORTS campaign conducted 144 measurements on the R/V James Cook on May 5-8, 2021 near the Porcupine Abyssal Plain 145 Sustained Observatory site (49°N, 16°W). Durkin et al. (2021) described the preparation of the 146 sediment trap tube and polyacrylamide gel in detail. 147

During EXPORTS NP, all GelCam cameras were functional, at all three depths across three 148 deployment durations, obtaining 9 total records. GelCams deployed during Deployment 1 and 3 149 collected more than 400 images each within more than six days of deployments. Gelcams during 150 Deployment 2 recorded about 360 images. During EXPORTS NA, housing failures resulted in 151 152 only one successful record collected at 125 m depth. This measurement recorded 169 images within four days. Additionally, the camera system infrequently encountered a compression issue 153 while transmitting the data to the SD card, leading to anomalous images being saved. These 154 corrupted image files could not be used for any image processing. A manual check identified 25 155 images with quality issues out of 3,846. 156

Table 1. Summary of sampling locations. *The North Atlantic deployment applied a hardware

Deployment	Cruise	Location	Dates	Depths	# of images
NP Deployment 1 STT 1, 2, 3	R/V Revelle	Subarctic North Pacific 50.1°N, 145.1°W	Aug. 15-21 2018 Deployment 1	95 m 145 m 195 m	445 410 443
NP Deployment 2 STT 1, 2, 3	R/V Revelle	Subarctic North Pacific 50.4°N, 145.1°W	Aug. 24-28 2018 Deployment 2	105 m 155 m 205 m	369 368 341
NP Deployment 3 STT 1, 2, 3	R/V Revelle	Subarctic North Pacific 50.6°N, 144.9°W	Aug. 31-Sep. 5 2018 Deployment 3	105 m 155 m 205 m	442 446 413
NA Deployment	R/V James Cook	Subarctic North Atlantic 48.9°N, 15.0°W	May 5-8 2021	125 m	169*

upgrade of a larger FOV and eliminated the bungee mount bar.

159

160 Imaging processing procedures

161 Color decomposition to remove ambient light variations

An image processing pipeline was developed for the GelCam images. Each raw image (see 162 example in Figure 2(A)) shows the LED near the outer edge of the gel cup (large white circle), the 163 bar holding the bungee (EXPORTS 2018 only), and darkened areas of the background where the 164 clear tube was mounted to the STT frame and near the top at the lid attachment. Here we created 165 a static mask to exclude the LED ring, the bungee mount bar, and other bright background objects 166 from the FOV. Areas of very bright pixels represented the LED ring near the edge and the orange 167 bungee in the middle. At the same time, less bright areas of the bungee mount bar were outlined 168 and then masked out by two fitted straight lines. During daytime, ambient sunlight caused portions 169 of the background to appear in a green-blue hue, even at depths up to 205 m. By decomposing the 170 images into the green, blue, and red channel (Figure 2(B) and (C)), we found that the red channel 171 was highly effective at filtering out the ambient light, since at the depth of the traps, all of the 172

ambient red wavelengths have been absorbed by seawater. We then converted the red channelimage to a gray scale for region of interest (ROI) identification.

The temporal variations in the blue and green channel were very effective at illustrating the day-175 and-night cycle in ambient sunlight. Since we were interested in diurnal variation of particle flux, 176 it was crucial that we account for any variations in particle detection likelihood or contouring that 177 178 might arise due to the variations of background light. As particles occupied less than 7.5% of the field of view, the blue and green pixel intensity (averaged over each image) provided a reliable 179 representation of the overall ambient light changes over time (Figure 3). The variations were very 180 low during nighttime, while noticeable peaks were observable around mid-day. A strong 181 correlation was found between the lighting variations within the green and blue channels and 182 concurrent measurements of surface light levels obtained from a Photosynthetically Active 183 Radiation (PAR) sensor deployed on a WireWalker (Suppl. Mat., Figure S3). The example 184 sequence of a single particle over time shown in Figure 3 highlights the RGB color variations 185 between night and day, with the background color turning green during the daylight hours. The 186 day-night pattern was not evident in the normalized red channel, and so by using only the red 187 channel for particle edge detection, we markedly reduced the sensitivity to background light 188 189 variations. Selection of the red channel also had the effect of eliminating the color variation associated with the trap tube components that were far away from the illumination source, such 190 191 as the tube bracket and top lip (see the darker concentric rings in Figure 2(A) compared to (C)).



Figure 2. Pipeline of image pre-processing procedures, with (A) original captured image, (B) decomposed images in green and blue channels, (C) red channel image, (D) masked grayscale image after background removal, and (E) binarized image with particles outlined. The semitransparent mount bar with the orange bungee did not exist during EXPORTS 2021.

197

192

198 Background removal

The next step in image preparation was masking of the static parts of the image that contained strongly contrasting tube parts close to the illumination source, such as the bungee bar and knot (Figure 2(C)). Here we used a proper orthogonal decomposition (POD) method, which has been proven effective in particle image velocimetry (PIV) (Mendez et al., 2017) and particle tracking

(Song and Rau, 2022). In our current configuration, we observed particle movements and 203 rotations driven by local circulations and small-scale flows acting on the sediment trap. Such 204 movements made the application of POD-based background removal possible. We implemented 205 this by applying a static mask first to all images in the red channel, followed by a singular value 206 decomposition. Subsequently, we eliminated the first principal component and reconstructed the 207 208 image frames. Figure 2 (D) shows the resulting background-removed product of the masked redchannel image in (C). Using this method, particles were successfully preserved while the 209 background was eliminated. 210



Figure 3. Example day-and-night cycle represented by the average pixel intensities in green (solid line) and blue (dotted line) channels, combined with one example tracking series at a time interval of roughly 6 hours (16 frames) during Deployment 3 STT 1. See Figure S3 in supplement materials for correlations with PAR signals.

216

217 **Particle contouring**

Next we developed a double-threshold method to convert particles in the grayscale image to binary form. First, we computed the Otsu threshold (Otsu, 1979) for each image frame after the background removal. We then averaged these thresholds to establish a global black-and-white cutoff. The first round of binarization was mostly effective, but we found that using a global

threshold did not completely outline particles with looser structures, such as aggregates, which 222 had segmented shapes. To remedy this, we set an additional but lower threshold by subtracting 223 three standard deviations of the pixel intensities from the global threshold. The lower cut-off 224 value retained relatively dimmer pixels, often bridging between aggregate segments, although it 225 also included some background noise. Here we retained pixels that passed the lower threshold if 226 227 they connected with pixels that had passed the first global threshold. Lastly, we applied an area filter with a 50 pixels minimum (ESD (equivalent spherical diameter) ~ 170 μ m), to exclude 228 noise features and small particles that lacked sufficient morphological details for classification. 229 We compiled a table assigning each particle a unique index and summarizing the images and 230 morphology of particles in all image frames. Figure 2 (E) shows an example of the particles that 231 232 passed our binarization method, outlined in red.

233 Particle tracking

234 We implemented a particle tracking method to continuously monitor the particle motion within the gel and measure the particle flux. The surface tethered sediment trap experienced fairly strong 235 vertical and rotational wave-driven motions during deployment. For the first two deployments in 236 the North Pacific, the gel cup was freely able to rotate within the tube, and this appeared to 237 enhance the rotational movements of the gel. Thus, the particles experienced horizontal motion 238 relative to the camera throughout the deployments (See supplementary videos). The high 239 viscosity of the gel helped to keep these motions fairly smooth, with strong auto-correlation 240 between the particle motions that were fairly straightforward to track by eye and with our bulk 241 242 motion routines. We first utilized particle image velocimetry (PIV) to capture the bulk movement between two consecutive image frames. We used the open-source PIVlab code (Thielicke and 243

Stamhuis, 2014) to implement a single PIV pass to detect the velocities associated with the bulk 244 motion. A square evaluation window with a size of 320×320 ($x \times y$) pixels was selected based 245 on the maximum distant particles seemed to travel within the 20 min image interval. This window 246 size also ensured sufficient bright pixels for cross correlation in the first few frames. We set the 247 window overlap at 50% in both x and y directions. Following this, a two-dimensional Gaussian 248 regression model (Nobach and Honkanen, 2005) provided a sub-pixel estimation. Additionally, 249 we applied a universal outlier detection after the single pass to eliminate inaccurate vectors 250 (Westerweel and Scarano, 2005). The PIV analysis yielded a bulk velocity field of 15×11 vectors. 251 To approximate particle movement between frames, we applied a two-dimensional interpolation 252 method to each particle's centroid position. 253

Next, we performed another cross-correlation method to precisely track each particle's movement 254 between adjacent frames. Here we created two interrogation windows, w_1 and w_2 , from two 255 successive frames. We defined the first interrogation window, w_1 , as a rectangular box around a 256 257 particle indexed *i* after the contouring. Each side of the window had a minimum distance of 8 pixels from the particle's edge. The second interrogation window, w_2 in the following frame, was 258 centered on the guessed position, based on the particle's estimated centroid after interpolating the 259 260 PIV vectors. We expanded this window size by 16 pixels for large particles (if the longest particle dimension was greater than 16 pixels), and by 8 pixels if the particle was small (less than 16 261 pixels). We then conducted a cross correlation search within w_2 to find the best match. The 262 location that generated the best correlation, with a correlation coefficient > 0.7 and a distinct peak 263 in the correlation plane, was allocated as position *i*' in the next frame. If this location matched a 264 particle *j* in the second frame (j = i'), we linked the two particle indices *i* and *j* in our tracking 265 series. Conversely, we considered particles without matches as new arrivals. Particle images in 266

Figure 3 illustrate one tracking example of a long fecal pellet that applied this algorithm. Despite challenges including translation, deformation, re-orientation, and background lighting variations, our method effectively identified the same particles over consecutive images. To verify, we randomly selected and manually examined 500 particle sequences (out of 43,286 total).

In some cases, two particles overlapped and crossed each other (28% of the total). The overlap 271 272 issue became increasingly problematic in the later deployment stage as more particles appeared in the FOV. In these cases, we flagged cases where two or more particles, such as i_1 and i_2 , 273 corresponded with the same match *j* in the next frame. Moreover, we also ran our tracking 274 algorithm in a reverse chronological order to detect "separation" events, the reverse of 275 overlapping. In addition to 500 randomly chosen particle sequences, we manually examined 186 276 overlapping cases. After applying the algorithm above, the tracking methods successfully 277 detected and differentiated the overlapping particles in all but two cases. This result convinced us 278 that our approach to distinguish overlapping particles was successful. 279

280 The final challenge we needed to resolve were cases where the particles moved past a masked feature such as the bungee bar. The bar, shown in Figure 2 (A), resulted in particles leaving and 281 the re-entering the FOV over durations that were longer than the consecutive frame interval. Here, 282 283 we extended the duration of predicted centroids over 12 hours using the PIV vectors. Unpaired particles were defined as those appearing in the last of the tracking series without any matches 284 285 since then. We compared the distance between the centroid of a newly arrived particle and 286 predicted locations of unpaired particles on their estimated trajectories, based on the new arrival time. If the distance was less than 32 pixels, this newly arrived particle would undergo the same 287 cross correlation method with other unpaired particles. Based on the same criteria described above, 288

two particles were considered as a match if the correlation coefficient was high with a unique peakin the correlation plane.

Not all particles were traceable from their arrival through the entire deployment to the last image, 291 and some, such as swimmers (organisms that actively entered the trap and were not part of the 292 passive flux) were discounted from the particle flux counts. Some particles drifted off-frame, while 293 294 others abruptly disappeared or had an intensity that faded below the threshold. To deal with this, we used a labeling system that indicated whether the particle disappeared at an edge, faded in FOV, 295 or was manually identified as a swimmer. Of the 43,286 tracked particles, 10,169 were tracked 296 from arrival to the final frame, 7,439 moved out of the FOV, 132 were categorized as large 297 swimmers. The rest of the tracking sequences were labeled as particles faded in FOV, weak cross-298 correlations, and small areas. Additionally, some particles were found to enter the FOV from the 299 sides, suggesting that they could be re-entering the image after arriving and leaving at an earlier 300 time. Fourteen percent of the total particles were flagged as a potential re-entry. Detailed diagnoses 301 302 with labels can be found in Table S1 in the supplemental materials.



303

Figure 4. Passively sinking particle classes with example sub-images, including (A) aggregates,
(B) large loose fecal pellets, (C) long fecal pellets, (D) dense detritus, (E) large salp pellets, and
(F) rhizaria.

307

308 Particle classification

Following (Durkin et al., 2021), we classified each unique particle into one of six categories, as 309 summarized in Figure 4. These particle categories were developed based on detailed microscope 310 images of the same gels as those presented here. However, due to the image quality and color 311 variations inherent in the GelCam prototype, the machine learning tools (Amaral and Durkin, 312 313 2024) that were developed for the microscope images were not optimal for these ROIs. Because of the relatively small number of GelCam images involved, and quantity of classified images 314 315 needed to train a supervised model, we decided to manually assign each particle an identity based 316 on their unprocessed colored sub-images. Aggregates had irregular shapes and loose structures and could have been clusters of phytoplankton cells and other detritus. Large loose fecal pellets 317 318 were produced by large zooplankton (e.g. pterpods) and were often elongated with dense coloring. Their shapes were more regular than aggregates and composed of one or a few distinct sub units. 319 320 These might be partially degraded long fecal pellets as well. Long fecal pellets were thin and elongated with smooth edges. Due to the lower pixel resolution, we classified most small ellipsoid 321 particles as dense detritus. Following the study that has been published on the presence of salps 322 323 (Bruland and Silver, 1981), we categorized their pellets as large and golden-colored, with nearly rectangular shapes. We also detected a few rhizarians, primarily Phaeodaria. Large zooplankton 324 were occasionally observed kicking and struggling in the gel layer, and were excluded as 325

swimmers. We differed from the classification suggested by Durkin et al. (2021), in that we did

not identify any phytoplankton cells or mini pellets, since these usually had a size of less than 50

 μm in ESD (Gowing and Silver, 1985), and were smaller than our detection limit.

329

330 ASSESSMENT

331 Modeling carbon fluxes



Figure 5. Cumulative particle areas as a function of time for selected samples in NP. Panel (A) 333 consists of the number of white pixels from binarized images (dotted gray line) and the lighting-334 corrected cumulative areas (red solid line) during Deployment 1 STT 1. Panel (B) shows 335 cumulative particle areas using the white pixel method (thick lines) and using the tracking-based 336 method (thin lines) for Deployment 2 STT 1 & 3. Thin lines only focus on passively sinking 337 particles classified following Figure 4. Panel (C) uses the white pixel method and compares the 338 cumulative particle areas during Deployment 3. See Figure S5 for comparisons of all 339 measurements. 340

Tracking "new" particle arrivals using the timelapse method enables new insights into carbon 342 flux variability, but it also raises new challenges. Because this is a new method, with imaging 343 tools that are still undergoing optimization, we tried two different methods for establishing an 344 image-based carbon flux proxy. The first - hereafter referred to as the "white pixel method" -345 summed up the area of all white pixels in the binarized image. The white-pixel-method aims to 346 quickly obtain the bulk particle accumulated areas using a single black-and-white threshold. This 347 method is sensitive to the day-night variations in ambient light. Figure 5 shows the time-varying 348 percentage FOV coverage, calculated through cumulative particle area divided by the entire FOV. 349 Brighter ambient light conditions tended to cause an apparent reduction in size, or result in some 350 particles dropping below the detection limit, which resulted in a drop of white pixel area during 351 daytime (see gray line, Figure 5(A)). A linear correction was developed that utilized the difference 352 in light intensity relative to the nighttime blue channel levels. The percentage of pixels that were 353 "quenched" by the brighter background were then computed based on the interpolated estimate 354 355 and the captured particle areas. We correlated the excessive blue light intensity with the percentage attenuation, allowing for a corrected white pixel number at each daytime time stamp 356 (see red line, Figure 5(A)). Detailed steps for this method can be found in the supplemental 357 358 materials.

While the background lighting correction did appear to correct the daytime dips in cumulative area computed via the "white pixel" method, we still observed a sharp reduction in area about 70 hours after deployment during Deployment 1 (Figure 5(A)). Review of the raw images revealed that this occurred because the trap lid bungee broke suddenly during a storm. It obscured part of the camera view, causing a flattening/decrease in the curve. Consequently, the corrected curve was composed of two nearly monotonic line segments. Similarly, deployments during Deployment 2 and 3 showed monotonic accumulations (Figure 5 B and C), despite some spikes caused by the arrival of large particles or swimmers. Overall, cumulative particle area flux decreased with depth (STT 1 was shallowest (95 - 105 m), STT 3 was deepest (195 - 205 m)).

The second method, hereafter referred to as the "tracking-based method", used a conservative 368 version of the particle tracking and classification. We did not sum up all tracking series at their 369 370 apparent time of arrival. Otherwise particles could be counted more than once due to the interruption in the tracking sequences, which was also a common issue applying the particle 371 tracking. Rather, we implemented a frame-based calculation. Given a specific image frame with 372 time stamp, we first found all identified particles that had been contoured. Next, we used tracking 373 sequences to trace each particle to its state upon arrival. Using the size and shape upon arrival is 374 intended to exclude any potential degradation or other morphological changes over time. Next, 375 following the labeling system described above, we added all particles that left the FOV before 376 this time stamp and subtracted particles that came into the FOV from the edges. This operation 377 378 efficiently corrected the cumulative particle areas influenced by the gel movement. It is likely that this method could underestimate the particle size for those arriving during the day. However, 379 the particle size of individual tracked particles did not show a distinct diurnal pattern. 380

Although it requires more processing time, an advantage of the tracking-based method over the white pixel method is that it allows an estimate of the time of arrival of particular particle types/sizes. The particulate organic carbon of each new particle was determined following the equation $C = A \cdot V^B$ listed in Durkin et al. (2021), where *C* is the carbon per particle (*mg*), *A* is the scaling coefficient, *V* is the particle volume (μm^3), and *B* is a unitless exponent. The exponent *B* characterizes the porosity of a given type of particle as its size increases. Table S2 in supplemental materials lists shapes, volumes, and modeling parameters of each particle type. We followed 388 Durkin et al. (2021) closely to model the POC flux. It should be noted that dense detritus were 389 poorly resolved in our images relative to the microscope images in that paper. This particle type 390 could be a mixed group of aggregates and fecal pellets.

After excluding air bubbles and swimmers, we extracted particle areas from the tracking series 391 and summed all six classified categories. Figure 5(B) compares the two methods of quantifying 392 393 cumulative particle areas in Deployment 2, where the symbolized curves represent the classification-based integrated areas and the others are the white-pixel-based cumulative particle 394 areas. The two methods showed great agreement with each other, but the approach based on 395 tracking and classification generated more smooth responses. Random spikes were eliminated 396 because the tracking series excluded air bubbles and swimmers. These symbolized curves did not 397 show any valley, because we only extracted particle areas at the time of arrival. In other words, 398 using the particle tracking automatically eliminated ambient light induced area variations. 399 Moreover, this approach accounted for particles leaving and entering the unmasked area, which 400 401 also explained the small deviation from using white pixels only. See Figure S5 for measurements of all stations using both methods. 402



404

Figure 6. Modeled particulate organic carbon (POC) flux by each type during EXPORTS NP. POC
fluxes are also compared with measured data in formalin-poisoned trap tubes (Estapa et al., 2021)
and modeled ones based on microscopic images (Durkin et al., 2021).

408

The tracking based method was used to model the net POC flux for each GelCam deployment 409 categorized according to particle type (stacked bars, Figure 6). These results are compared to the 410 total carbon fluxes measured directly in the bulk flux tubes (open circles) and modeled POC 411 fluxes using microscope images and classification (crosses). The details of these carbon flux 412 methods can be found in (Estapa et al., 2021) and (Durkin et al., 2021) respectively. In general, 413 the GelCam-modeled POC flux had fairly good agreement with the other methods in the 414 shallower traps (STT 1 at 95m and STT 2 at 155m), and tended to underestimate the flux relative 415 to the other measurements at STT 3 (at 205m). We suspect that a primary reason for this 416

discrepancy at depth is that the GelCam could not resolve small particles. Particles with an equivalent diameter smaller than 100 μ m could contribute up to 46% to the total POC at STT 3 during the North Pacific EXPORTS cruise (Durkin et al., 2021), whereas the detection limit was 170 μ m. Additionally, visual sensors have a difficult time detecting nearly transparent or disaggregating materials as particles became further degraded with depth.

422 The particle-classified GelCam derived fluxes were primarily composed of aggregates, large loose fecal pellets, long fecal pellets, and dense detritus during EXPORTS-NP. Long fecal pellets were 423 dominant at STT 1, contributing 67% to the total POC fluxes. Aggregates, large loose fecal pellets, 424 and dense detritus were less abundant but still contributed a significant part comparing to rhazaria. 425 At STT 3, aggregates were the most abundant, and contributed up to 70% of the total carbon. Salp 426 pellets were observed as one of the dominant particle types in the microscope analysis of the gels 427 (Durkin et al., 2021; Steinberg et al., 2023). Similarly, we found that large salp pellets dominated 428 the GelCam-based POC calculations for some deployments. For example, two huge salp pellets 429 430 constituted 79% of the total carbon during Deployment 2 STT 2.

Transfer efficiency was highly variable across particle types. Here we defined T_{100} as transfer 431 efficiency, which is the ratio of carbon flux at 100 m below the reference depth (defined as the 432 433 depth of STT 1) and carbon flux at the reference depth (i.e., POC_{STT3}/POC_{STT1}). The transfer efficiency in the NP were 14%, 9.6%, and 6.0%, respectively. The percentage was lower than the 434 435 sediment trap measurements (Estapa et al., 2021) due to the low modeled POC fluxes at STT 3. Missing small particles could be one main reason. Long fecal pellets and large loose fecal pellets 436 had very high attenuation (> 97.0%), showing a good agreement with the microscope analysis 437 (Durkin et al., 2021). During Deployment 2 STT 3, we did not see any large loose fecal pellets and 438 439 only 1.3% carbon of long fecal pellets remained compared to STT 1. Aggregates had the highest transfer efficiencies, which varied from 30% to 68%. One likely reason would be the conversion
from other particle types and aggregation. We did not calculate the transfer efficiency of large salp
fecal pellets due to the low sample number.

443

444 Episodic fluxes of fecal pellets

In contrast to traps that use only a final accumulation to estimate a time-integrated carbon flux, the 445 GelCam provides a unique opportunity to visually examine the time series of particle fluxes during 446 the trap deployment. One of the more striking patterns that emerged from the particle arrival and 447 tracking analysis was the observation of diurnal fluctuations in the arrivals of fecal pellets across 448 deployments and stations in the shallowest trap. Particle fluxes of aggregates (Figure 7(B), green 449 line) and fecal pellets (Figure 7(B), blue line) at STT 1 were binned into two-hour time intervals 450 and averaged into local solar time bins (operational time zone). A total of 13 days across three 451 452 deployments were used for averaging the hourly fluxes. The fecal pellet fluxes included large loose and long fecal pellets only. We did not include the dense detritus, rhizaria or salp fecal pellets in 453 the analysis. It should be noted that we also disregarded the fluxes after the mount bungee broke 454 during Deployment 1 but kept the data before the interruption. The average fecal pellet flux 455 reached a maximum of up to 3 mmol C m⁻² d⁻¹ occurring around 11 PM to 12 AM. After the daily 456 maximum, the carbon fluxes from fecal pellets reduced quickly and reached less than 1 mmol C 457 m⁻² d⁻¹. In contrast, the aggregate fluxes did not show any distinct diurnal patterns. Because of the 458 459 very strong attenuation with depth of the fecal pellet classes, we did not have enough samples to implement a statistical analysis in STT 2 and STT 3. We also observed diurnal changes in the flux 460

of fecal pellets in the North Atlantic dataset (STT 1 at 125 m; see Figure S6). The GelCam observed
the most fecal pellets during the second night and a consistent flux of aggregates over three days.

463



464

Figure 7. (A) acoustic backscattering strength from the 150 KHz ADCP with a depth range of 25 m to 500 m, shown in decibel (db). The dashed black line represents the deployment depth of GelCam at 105 m. (B) Hourly fluxes of fecal pellets in blue and aggregates in green at STT 1 during EXPORTS-NP. POC flux was averaged within a two-hour time interval. Mean values are displayed as center lines. The upper and lower boundaries of the patches represent 75th and 25th percentile of the samples.

471

The flux increase in fecal pellets occurring at night likely resulted from diurnal migration of 472 zooplankton that swam to depths shallower than the trap in the upper ocean to feed at night (Cyr 473 and Pace, 1992; Haney, 1988; Mackas and Bohrer, 1976). In Figure 7(A) the acoustic 474 backscattering signal analyzed from the 150 KHz acoustic Doppler current profiler (ADCP) 475 indicated higher near-surface migrator activity during the nighttime compared to day (see the 476 477 anomaly of acoustic backscatter in Figure 1 from Maas et al. (2021)). Another interesting finding is that the time varying flux of fecal pellets in Figure 7 shows a skewed trend between the fecal 478 pellet flux increase and decrease. The data suggests a sharp drop-off after midnight, with a 479 minimum at around 6 a.m. (local) and a steady slow increase over the day. 480

481

482 Flux events across particle types

Flux events, sometimes called "pulse" events, have high rates of particle flux within a short time 483 484 frame. To better model the biological carbon pump, it is critical to quantify these flux events (Smith et al., 2018). Bauerfind et al. (1994) observed a sharp increase of fecal pellet flux after the 485 486 maximum diatom sedimentation. However, this and other studies (e.g., (Bauerfeind et al., 1994) and (Cao et al., 2024)) have conducted observations of flux events over a long time scale (from 487 weeks to months). The use of GelCams enables measurements of flux events over timescales 488 from hours to days. Among the GelCam deployments described in this manuscript, Deployment 489 3 from the North Pacific campaign (Aug. 31 - Sep. 5, 2018) provided the most complete and best 490 491 quality record at all three depths. Here we examine the flux patterns of discrete particle types from the NP Deployment 3 as a function of time and depth (Figure 8). In order to highlight the 492 temporal trends over multiple days and de-emphasize the day-and-night variation, we smoothed 493

the flux time series using a moving average of 73 frames, equivalent to approximately 24 hours.
We note that while this approach helped to highlight multi-day trends, it reduced the pattern of
diel variations highlighted in Figure 7.

The flux of aggregates showed a decrease over depth consistent with the traditional Martin Curve pattern due to particle remineralization (Figure 8(A)). On average, the flux was 52% (±14% at STT 2), and 27% (±6% at STT3) of the flux at STT 1 (105 m). These ratios stayed fairly constant, even as the flux roughly tripled at all depths, varying from a minimum of 0.14 up to 1.0 mmol C $m^{-2}d^{-1}$ at STT 1 over the six day deployment. The time series of the deeper fluxes was thus correlated with the shallowest trap STT1 ($R^2 = 0.92$ (p < 0.05) and $R^2 = 0.84$ (p < 0.05) at STT 2 and STT 3, respectively).

The flux of large loose and long fecal pellets also showed an increase between the early and later 504 stages of the deployment Figure 8(B) and (C)). The depth attenuation of large loose pellets was 505 on average 46% between STT 2 and STT 1, and nearly 100% by STT 3 with only 8 unique 506 particles of this type observed at 205 m. The temporal variations of large loose pellets were 507 correlated between STT 1 and STT 2 ($R^2 = 0.31$, p < 0.05) and uncorrelated ($R^2 = 0.03$, p < 0.05) 508 by STT3. The correlation between the shallower traps, which both showed a distinctive increase 509 510 on Sep. 3 and peak on Sep. 4, lent confidence that we were resolving real short-term variations in flux that could be tracked across depth (compare the solid and short-dashed lines in Figure 511 512 8(B)).

The temporal variability in long fecal pellets at STT1 showed a similar pattern, with a gradual increase from 0.33 mmol C m⁻²d⁻¹, and a peak of 4.3 mmol C m⁻²d⁻¹ around 6 p.m. on Sep. 4 (Figure 8(C)). We observed a very strong depth attenuation, with a mean loss of 9.4% between STT 1 and STT 2, and 0.20% by STT 3. The time series of POC fluxes between STT 1 and STT 2 were weakly correlated with $R^2 = 0.26$ (p < 0.05), while the higher attenuance at STT 3 caused a lower correlation of $R^2 = 0.13$ (p < 0.05). Interestingly, we also saw a correlation between long fecal pellets and large loose pellets with R^2 of 0.39 (p < 0.05) at STT 1. It is plausible that there was cross-over with the labeling of these two categories and they were derived from the same origins, or that the distinct groups of animals that created the pellets were correlated with fluctuations of abundance, and thus pellet production rates.

The dense detritus category was perhaps the most confounding in terms of temporal variations 523 (Figure 8(D)). POC flux was overall fairly low, with a range between 0.03 and 0.12 mmol C 524 $m^{-2}d^{-1}$ at STT1. This was the only category where we observed a sustained period where the flux 525 at STT2 exceeded that of STT1. With the data we have, it is only possible to speculate about the 526 reason(s) for this reversal. For example, it could have been due to disaggregation of large particles 527 into the dense detritus category between STT1 and STT2. Alternatively, there may have been a 528 source of dense pellets located between the two traps, horizontal advection of patches, or a 529 temporal lag that could have contributed to temporary inversions in the expected flux attenuation 530 profile. By STT3, we observed that dense detritus were attenuated by 35% compared to the 531 shallowest depth. 532

Additionally, we noticed a strong correlation between aggregates at STT 1 and dense detritus at greater depths ($R^2 = 0.91$ (p < 0.05) and $R^2 = 0.50$ (p < 0.05), respectively for STT2 and STT3). One likely reason was that the category of dense detritus were actually small aggregates or fragments. We only observed 6 rhizaria and 36 salp pellets across three depths during the six day deployment in Deployment 3. Thus correlations based on these categories were less suggestive. Tables S3 and S4 in supplemental materials show a full list of coefficients of determination between depths or particle types.





Figure 8. Time variations of POC fluxes across particle types during NP Deployment 3. Fluxes
at STT 1, 2, and 3 are represented by solid, dotted, and dashed lines, respectively.

543

544 **DISCUSSION**

545 GelCams provide a new tool for observing short time variations of particle fluxes occurring 546 within the duration of a typical surface tethered trap deployment. GelCams, when used alongside other flux measuring instruments, can be used to estimate carbon flux, and resolve episodic events and diel variations among different particle types. The modeled POC fluxes from the GelCams generally agreed with ground-truth measurements (Estapa et al., 2021; Durkin et al., 2021; Buesseler et al., 2020), although the current design cannot resolve small particles like the mini pellets described in other studies (Durkin et al., 2021). The simple, low cost, and open source design suggests that GelCams could be reproduced and attached to other platforms such as neutrally buoyant sediment traps.

Sediment traps equipped with a GelCam offer several other advantages. By taking a continuous series of photographs of the trap contents, we can observe if and when swimmers arrive, and any disruptions or changes that occur to the particles during the time spent in the trap before recovery. By tracking each individual particle over time, we can qualitatively observe if the particles are undergoing morphological changes, or if the gel is effective at preserving their size and other characteristics.

560 We tried two different methods to model the time-resolved particle flux from the image sequences. The white pixel method and the tracking-based methods agreed with each other in terms of the 561 cumulative particle areas. We can choose between the two depending on the processing time and 562 563 required data products. The white pixel method yielded a quick method to estimate the bulk particle flux, but was affected by background ambient light variations between day and night. We 564 proposed a simple method to remove this effect using the red channel of the RGB image, which 565 yielded a time series with a more monotonic increase in cumulative particle area. We can simply 566 extract the bulk flux information and time varying fluxes in particle areas, although the carbon 567 flux will require a further calibration incorporating particle transparency properties. This method 568 569 is similar to the attenuance-flux approach employed by Estapa et al. (2024) using optical sediment

traps, except the attenuance flux approach sums the attenuance values of each pixel, while the 570 "white pixel method" here used the binarized images so each pixel had a value of unity (i.e., the 571 572 method does not account for light intensities and transparency of particles). While more computationally- and labor-intensive, the particle tracking method with classification was less 573 affected by the day-night ambient light variation, and provided information about the specific 574 575 time of arrival of individual particles. These classifications allowed us to examine the relative importance of each particle type in the bulk flux, and short-term patterns (such as diel variation) 576 and detailed temporal correlations between particle types and depths. The GelCam framework 577 thus generates a large data set that allows for both individual case studies and statistical analysis. 578 Recovery and subsequent analysis of the gel samples under an imaging microscope and via 579 extraction for molecular or mineral analysis can reveal even more biological details and help with 580 ground truthing (see (Durkin et al., 2021)). 581

The GelCam image series emphasizes the important role of fecal pellets in carbon flux during the 582 583 EXPORTS North Pacific campaign. In agreement with previously published results (Durkin et al., 2021), long fecal pellets were the dominant particle type in the traps near the surface. The 584 long fecal pellets and large loose fecal pellets showed rapid flux attenuation, with roughly 91% 585 586 of the POC flux removed between the top two traps. Aggregates and dense detritus had higher transfer efficiencies than fecal pellets, where more complex physical and biochemical processes 587 could significantly influence the results. Microbial degradation potentially played a more 588 589 important role (Stephens et al., 2024), as small aggregates or detritus could be fragments of degraded fecal pellets. Additionally, the particle type of dense detritus could include aggregates 590 or fecal pellets in small size. Distinguishing between them was a challenge in the GelCam images, 591 592 which had a resolution approximately four times lower than the 115× microscope images

published in Durkin et al. (2021). Functionally, these particle types may represent particles in a 593 more advanced stage of microbial degradation compared to the other pellet categories. The 594 595 particles that were categorized as aggregates were possibly highly degraded detritus with a loose structure. It also seems plausible that the aggregates and dense detritus in the deeper traps 596 originated as fecal pellets at shallow depths. Moreover, physical processes including aggregation 597 598 (Burd and Jackson, 2009; Kiørboe, 2001; Alldredge and Silver, 1988) and disaggregation (Briggs et al., 2020; Song et al., 2023) which can result in morphological changes as the particle sink over 599 time, likely contributed to the uncertainty in the transfer efficiency of any particular particle class. 600 Compared to traditional sediment trap instruments, GelCam provided unique insights into the 601 variability of carbon flux during the trap deployments, providing a time-series with 15 minute 602 resolution. With this record, we were able to observe diel variations in fecal pellet fluxes in the 603 shallowest traps in both the North Pacific and North Atlantic. The diel flux peaked at midnight, 604 likely due to the migration of zooplankton and fish to graze above the trap depth. This finding 605 606 was consistent with other diel vertical migration evidence from nets and bio-acoustics (Maas et al., 2021), showing the swarming of plankton species to the surface ocean at night. A comparison 607 using aggregates did not show any distinct diurnal changes. Analysis of 14 cumulative days of 608 609 fecal pellet flux also revealed a temporal asymmetry, with a rapid three-fold drop in pellet flux between midnight and 6 a.m. (local) and then a gradual increase over the rest of the day. This 610 611 asymmetrical pattern could be attributable to feeding or defecation behaviors, and further studies 612 are required to elucidate the drivers and if similar patterns are observed in other locations.

The temporal correlations we presented in Tables S3 and S4 between trap depths are without any lag. These analyses thus assume that the processes that dictate these short timescale changes in flux are occurring simultaneously across depth. In practice however, if particles are sourced from

above the shallowest trap, and take some time to sink past and into deeper traps, we might expect 616 a time lag reflecting the vertical distance between the traps divided by the sinking speed. Previous 617 studies using particle abundance proxies such as optical backscattering (Dall'Olmo and Mork, 618 2014; Briggs et al., 2011) and laser optical plankton counters (Petrik et al., 2013; Jackson et al., 619 2015) have shown a descending front of particles reaching successively deeper depths over time. 620 621 We expected to potentially see similar results in the timing of flux events with the GelCam. Overall however, these records did not reveal obvious time lags, and so the correlations were 622 analyzed with zero lag. 623

An exception to this may be qualitatively evident in the GelCam timeseries from STT 1 and 3 624 from Deployment 3 EXPORTS-NP (see Figure 5(C) and Figure 8(A)). In Figure 5(C), we 625 observed that the total cumulative particle areas all had a noticeable slope shift, with an increase 626 at STT 1 (105 m) around 75 hours after the deployment while the slope of STT 3 (205 m) did not 627 increase until 96 hours. This shift in slope provided a bulk sinking speed of roughly 100 m/day. 628 629 The particle classifications suggest that the increase was partially driven by aggregates (Figure 8(A)). Dividing the vertical distance between STT 1 and 3, by the time lag, we achieve a roughly 630 estimated sinking speed of 50 m/day for aggregates. The estimated sinking speeds of bulk 631 632 particles or aggregates fall within the range of another study on the same cruise (Romanelli et al., 2024) and previous studies (McDonnell and Buesseler, 2010; Fowler and Small, 1972; Trull et 633 634 al., 2008; Ploug et al., 2008). The absence of this pattern for other particle types could be 635 indicative of faster sinking speeds, though our data was not sufficient to make a conclusive assessment of this hypothesis. We anticipate that future GelCam studies will allow more robust 636 observations of these dynamics. 637

The GelCam system is a promising tool for providing high temporal resolution flux, but it also 638 has a number of limitations that could be addressed with future technical improvements. A 639 primary challenge was the low pixel resolution and variability in illumination, which can be 640 improved in the future. In this study, we could not determine cumulative mini fecal pellets, as 641 their size was smaller than one pixel. Additionally, particle classification had high uncertainty for 642 643 small particles that just passed the detection threshold. For example, long fecal pellets, or fibers could be misidentified due to inadequate details. Durkin et al. (2021) noted aggregates as one 644 dominant population in the Subarctic North Pacific, but many aggregates were lost during the 645 identification process because of their loose structures and transparent exopolymer particles (TEP) 646 (Passow, 2002; Azetsu-Scott and Passow, 2004). Inevitably, we could not account for the 647 contributions of organic carbon from TEP. 648

Salp fecal pellets were found to be abundant during EXPORTS-NP (Durkin et al., 2021; 649 Stamieszkin et al., 2021), but we could only resolve ones with large sizes due to over-saturated 650 colors in small particles. We observed very few large salp pellets but these large particles could 651 substantially affect the modeling of POC fluxes (e.g., Deployment 2 STT 2). Since we only 652 deployed one GelCam at each depth, with a sampling area of $< 30 cm^2$, the limitation could cause 653 654 high variability of rhizaria and large salp pellets, collected from different sediment trap tubes. These rare occurrences compared to large numbers of other particles suggests that the deployment 655 656 of larger numbers of GelCams and other particle-intercepting sediment traps could improve 657 collection statistics. Fortunately the low cost of GelCam could make it possible in the future. Additionally, smaller salp pellets were usually classified into aggregates or detritus. We could 658 minimize this uncertainty by resolving particle color in the next development iteration. 659

We only resolved 2D images of particles from the single-camera imaging system. While we could 660 predict thickness of particles that retained regular shapes, such as long fecal pellets, particles prone 661 662 to flattening within the gel would have more biased thickness predictions. These particles, such as especially delicate aggregates, tend to flatten and extend their projection areas when resting on a 663 platform, thus leading to the overestimation of thickness based on assumptions of spherical shapes. 664 In addition to uncalibrated 3D morphological information, 2D images resulted in the overlap of 665 sub-images. For future improvements, capturing an additional view for 3D volume reconstruction, 666 or establishing an empirical relationship for aggregate thickness, would be advantageous. 667

Hardware upgrades will enhance the GelCam design in the future. For instance, in EXPORTS-668 NA, we encountered images with a blue hue, but new GelCam models will send fixed parameters 669 to the Pi camera to ensure calibrated white balance. Upgrading to more advanced camera modules 670 with attached lenses will also generate a higher pixel resolution. Although it is still challenging 671 to detect mini fecal pellets ($< 50 \ \mu m$), these improvements can reduce the detection limit of the 672 673 imaging system. We are also exploring the integration of GelCam with Langrangian floats. As Siegel and Deuser (1997) demonstrated, horizontal bulk flow is normally orders of magnitude 674 greater than the sinking velocity, so replacing the surface tethered traps (STT) with neutrally 675 676 buoyant sediment traps (NBST) will lead to a more precise flux measurement within the water column. Furthermore, incorporating a swimmer exclusion device (e.g. the "labyrinth of doom" 677 678 (Coale, 1990)) could still be favorable. Although we could exclude them during the post 679 processing, the swimming zooplankton brought more uncertainty in PIV analysis and they may potentially interact with other particles. 680

The white pixel method can be further developed by building an empirical relation to the POC fluxes. This idea is not new, as Bishop et al. (2016) estimated the carbon fluxes by empirically

35/57

relating thickness of particles to their projection areas and converting the particle volume into the 683 carbon content. Similarly, Estapa et al. (2024) empirically correlated the light attenuance of 684 particles with the POC fluxes. We did not apply any of these empirical relations, because particles 685 were illuminated from the side (not behind as required for an attenuance estimate), particle areas 686 were corrected based on lighting, and particles preserved in the gel layer did not become fully 687 flattened due to gravity. Building a reliable relation requires an extensive datasets collected by 688 GelCam and other similar instruments featuring upward facing cameras and side lighting (i.e., 689 quasi-darkfield illumination). Alternatively, we could utilize the other tracking-based method for 690 more accurate POC flux estimation than an empirical relation. The tracking-based and white-691 pixel-based methods are interchangeable depending on the specific purpose. 692

Implementing an automated particle classification process will be beneficial, to reduce the labor 693 of processing sub-images manually. Leveraging machine learning methods will provide 694 significant insights into the model of carbon export(Irisson et al., 2022; Trudnowska et al., 2021; 695 696 Amaral and Durkin, 2024). For example, Davies et al. (2017) applied a deep convolutional neural network (CNN) to classify particles, including larvae, copepod, and oil droplets, in an in situ 697 imaging system. However, many classifiers are still constrained by the limited size of available 698 699 particle image datasets (Irisson et al., 2022). We believe machine learning methods and GelCam will have a mutually beneficial relationship in the future. On one side, machine learning will 700 701 enhance the particle identification capabilities of the GelCam. More importantly, as a low-cost 702 but reliable instrument, the GelCam records a wide range of particle types, which will generate a larger database of particle images and improve the accuracy of machine learning classifiers. 703 Future work can expand the GelCam in more studies of the biological carbon pump. The 704

705 polyacrylamide gel layer maintains particle structure during sample collection, therefore we can

incorporate more GelCams in various directions to capture the 3D shape. By removing the gel 706 layer from the sediment trap, we can track marine particles without interfering with natural 707 biological processes. This adjustment would enable visual quantification of microbial 708 degradation rates acting on fecal pellets and aggregates. In other words, this modified instrument 709 will potentially allow observations of time varying morphological changes or even fragmentation. 710 711 Additionally, continuing to deploy GelCams with traditional sediment traps at different locations, seasons, and depths will not only yield estimates of a bulk particle sinking velocity but also 712 provide more insights into the varying settling behaviors across different particle types. 713

714 ACKNOWLEDGMENTS

Data collected through this project was supported by NASA Grants 80NSSC17K0662 and
80NSSC21K0015.

Postdoctoral support for YS was provided by Oceankind foundation. We wish to thank Dr. Nils
Haentjens, Dr. Vinicius Amaral, Roger Patrick Kelly, Sean O'Neil, Dr. Alyson Santoro, Nicola
Paul, and the EXPORTS team for assistance with data collection, their helpful discussions and
constructive recommendations. We are also grateful to captains and crew of the R/V Revelle and
R/V James Cook.

722

723 CONFLICT OF INTEREST

724 None declared.

725

726 **REFERENCES**

- Alldredge, A. L. and Gotschalk, C. (1988). In situ settling behavior of marine snow1. *Limnology and Oceanography*, 33(3):339–351.
- Alldredge, A. L. and Silver, M. W. (1988). Characteristics, dynamics and significance of marine
- snow. *Progress in Oceanography*, 20(1).
- Amaral, V. J. and Durkin, C. A. (2024). A computer vision-based approach for estimating carbon
 fluxes from sinking particles in the ocean. *bioRxiv*.
- Azetsu-Scott, K. and Passow, U. (2004). Ascending marine particles: Significance of transparent
- exopolymer particles (TEP) in the upper ocean. *Limnology and Oceanography*, 49(3):741–748.
- Basu, S. and Mackey, K. (2018). Phytoplankton as Key Mediators of the Biological Carbon Pump:
- Their Responses to a Changing Climate. *Sustainability*, 10(3):869.
- Bauerfeind, E., Bodungen, B., Arndt, K., and Koeve, W. (1994). Particle flux, and composition of
 sedimenting matter, in the Greenland Sea. *Journal of Marine Systems*, 5(6):411–423.
- 560 southenting match, in the Greemand Sea. *Southat of marine Systems*, 5(0). 411-425.
- Bishop, J. K. B., Fong, M. B., and Wood, T. J. (2016). Robotic observations of high wintertime
- carbon export in California coastal waters. *Biogeosciences*, 13(10):3109–3129.
- 741 Briggs, N., Dall'Olmo, G., and Claustre, H. (2020). Major role of particle fragmentation in
- regulating biological sequestration of CO2 by the oceans. *Science*, 367(6479):791–793.
- 743 Briggs, N., Perry, M. J., Cetinic, I., Lee, C., D'Asaro, E., Gray, A. M., and Rehm, E. (2011). High-
- ⁷⁴⁴ ′ resolution observations of aggregate flux during a sub-polar North Atlantic spring bloom. *Deep*
- *Sea Research Part I: Oceanographic Research Papers*, 58(10):1031–1039.
- 746 Bruland, K. W. and Silver, M. W. (1981). Sinking rates of fecal pellets from gelatinous zooplankton
- 747 (Salps, Pteropods, Doliolids). *Marine Biology*, 63(3):295–300.

- 748 Buesseler, K. O., Antia, A. N., Chen, M., Fowler, S. W., Gardner, W. D., Gustafsson, O., Harada,
- K., Michaels, A. F., Rutgers van der Loeff, M., Sarin, M., Steinberg, D. K., and Trull, T. (2007).
- An assessment of the use of sediment traps for estimating upper ocean particle fluxes. *Journal of*
- 751 *Marine Research*, 65(3):345–416.
- 752 Buesseler, K. O., Benitez-Nelson, C. R., Roca-Mart'1, M., Wyatt, A. M., Resplandy, L., Clevenger,
- S. J., Drysdale, J. A., Estapa, M. L., Pike, S., and Umhau, B. P. (2020). High-resolution spatial and
- temporal measurements of particulate organic carbon flux using thorium-234 in the northeast
- Pacific Ocean during the EXport Processes in the Ocean from RemoTe Sensing field campaign.
- *Elementa: Science of the Anthropocene*, 8(1).
- Burd, A. B. and Jackson, G. A. (2009). Particle Aggregation. *Annual Review of Marine Science*, 1(1):65–90.
- Cao, J., Liu, Z., Lin, B., Zhao, Y., Li, J., Wang, H., Zhang, X., Zhang, J., and Song, H. (2024).
 Temporal and vertical variations in carbon flux and export of zooplankton fecal pellets in the
 western South China Sea. *Deep Sea Research Part I: Oceanographic Research Papers*,
 207:104283.
- Cavan, E. L., Henson, S. A., Belcher, A., and Sanders, R. (2017). Role of zooplankton in
 determining the efficiency of the biological carbon pump. *Biogeosciences*, 14(1):177–186.
- Cavan, E. L., Le Moigne, F. A. C., Poulton, A. J., Tarling, G. A., Ward, P., Daniels, C. J., Fragoso,
- G. M., and Sanders, R. J. (2015). Attenuation of particulate organic carbon flux in the Scotia Sea,
- Southern Ocean, is controlled by zooplankton fecal pellets. *Geophysical Research Letters*,
 42(3):821–830.
- Coale, K. H. (1990). Labyrinth of doom: A device to minimize the "swimmer" component in
 sediment trap collections. *Limnology and Oceanography*, 35(6):1376–1381.

- Collins, J. R., Edwards, B. R., Thamatrakoln, K., Ossolinski, J. E., DiTullio, G. R., Bidle, K. D.,
- Doney, S. C., and Van Mooy, B. A. S. (2015). The multiple fates of sinking particles in the North
- Atlantic Ocean. *Global Biogeochemical Cycles*, 29(9):1471–1494.
- Cyr, H. and Pace, M. L. (1992). Grazing by Zooplankton and Its Relationship to Community
- 575 Structure. *Canadian Journal of Fisheries and Aquatic Sciences*, 49(7):1455–1465.
- Dall'Olmo, G. and Mork, K. A. (2014). Carbon export by small particles in the Norwegian Sea.
- *Geophysical Research Letters*, 41(8):2921–2927.
- Davies, E., Brandvik, P., Leirvik, F., and Nepstad, R. (2017). The use of wide-band transmittance
- imaging to size and classify suspended particulate matter in seawater. *Marine Pollution Bulletin*,
 115(1-2):105–114.
- De La Rocha, C. L. and Passow, U. (2007). Factors influencing the sinking of POC and the efficiency of the biological carbon pump. *Deep Sea Research Part II: Topical Studies in Oceanography*, 54(57):639–658.
- 784 Durkin, C. A., Buesseler, K. O., Cetinic, I., Estapa, M. L., Kelly, R. P., and Omand, M. (2021). A
- 785 Visual' Tour of Carbon Export by Sinking Particles. *Global Biogeochemical Cycles*, 35(10).
- 786 Durkin, C. A., Van Mooy, B. A. S., Dyhrman, S. T., and Buesseler, K. O. (2016). Sinking
- phytoplankton associated with carbon flux in the Atlantic Ocean. *Limnology and Oceanography*,
 61(4):1172–1187.
- Estapa, M., Buesseler, K., Durkin, C. A., Omand, M., Benitez-Nelson, C. R., Roca-Mart'ı, M.,
 Breves,
- E., Kelly, R. P., and Pike, S. (2021). Biogenic sinking particle fluxes and sediment trap collection
- rg2 efficiency at Ocean Station Papa. *Elementa: Science of the Anthropocene*, 9(1).

- Estapa, M. L., Durkin, C. A., Slade, W. H., Huffard, C. L., O'Neill, S. P., and Omand, M. M. (2024).
- A new, global optical sediment trap calibration. *Limnology and Oceanography: Methods*,
 22(2):77–92.
- Fowler, S. W. and Knauer, G. A. (1986). Role of large particles in the transport of elements and
- ⁷⁹⁷ organic compounds through the oceanic water column. *Progress in Oceanography*, 16(3):147–194.
- Fowler, S. W. and Small, L. F. (1972). Sinking Rates of Euphausiid fecal pellets. *Limnology and Oceanography*, 17(2):293–296.
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Bakker, D. C. E., Hauck, J.,
- 801 Landschutzer," P., Le Quer' e, C., Luijkx, I. T., Peters, G. P., Peters, W., Pongratz, J.,
- Schwingshackl, C., Sitch, S., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., Anthoni, P.,
- Barbero, L., Bates, N. R., Becker, M., Bellouin, N., Decharme, B., Bopp, L., Brasika, I. B. M.,
- 804 Cadule, P., Chamberlain, M. A., Chandra, N., Chau, T.-T.-T., Chevallier, F., Chini, L. P., Cronin,
- M., Dou, X., Enyo, K., Evans, W., Falk, S., Feely, R. A., Feng, L., Ford, D. J., Gasser, T., Ghattas,
- J., Gkritzalis, T., Grassi, G., Gregor, L., Gruber, N., Gurses, O., Harris, I., Hefner, M., Heinke, J.,
- Houghton, R. A., Hurtt, G. C., Iida, Y., Ilyina, T.," Jacobson, A. R., Jain, A., Jarn'ıkova, T., Jersild,
- A., Jiang, F., Jin, Z., Joos, F., Kato, E., Keeling, R. F., 'Kennedy, D., Klein Goldewijk, K., Knauer,
- J., Korsbakken, J. I., Kortzinger, A., Lan, X., Lef' evre, N., Li, H., Liu, J., Liu, Z., Ma, L., Marland,
- 810 G., Mayot, N., McGuire, P. C., McKinley, G. A., Meyer, G., Morgan, E. J., Munro, D. R., Nakaoka,
- S.-I., Niwa, Y., O'Brien, K. M., Olsen, A., Omar, A. M., Ono, T., Paulsen, M., Pierrot, D., Pocock,
- K., Poulter, B., Powis, C. M., Rehder, G., Resplandy, L., Robertson, E., Rodenbeck, C., Rosan, T.
- M., Schwinger, J., S" ef' erian, R., Smallman, T. L., Smith, S. M., Sospedra-Alfonso, R., Sun, Q.,
- Sutton, A. J., Sweeney, C., Takao, S., Tans, P. P., Tian, H., Tilbrook, B., Tsujino, H., Tubiello, F.,
- van der Werf, G. R., van Ooijen, E., Wanninkhof, R., Watanabe, M., Wimart-Rousseau, C., Yang,

- D., Yang, X., Yuan, W., Yue, X., Zaehle, S., Zeng, J., and Zheng, B. (2023). Global Carbon Budget
 2023. *Earth System Science Data*, 15(12):5301–5369.
- Gowing, M. M. and Silver, M. W. (1985). Minipellets: A new and abundant size class of marine
 fecal pellets. *Journal of Marine Research*, 43(2):395–418.
- Haney, J. F. (1988). Diel patterns of zooplankton behaviour. *Bulletin of Marine Science*,
 43(3):583–603.
- Henson, S. A., Sanders, R., Madsen, E., Morris, P. J., Le Moigne, F., and Quartly, G. D. (2011). A
- reduced estimate of the strength of the ocean's biological carbon pump. *Geophysical Research Letters*, 38(4).
- 825 Irisson, J.-O., Ayata, S.-D., Lindsay, D. J., Karp-Boss, L., and Stemmann, L. (2022). Machine
- Learning for the Study of Plankton and Marine Snow from Images. *Annual Review of Marine Science*, 14(1):277–301.
- Jackson, G. A., Checkley, D. M., and Dagg, M. (2015). Settling of particles in the upper 100 m of
- the ocean detected with autonomous profiling floats off California. Deep Sea Research Part I:
- 830 *Oceanographic Research Papers*, 99:75–86.
- Kiørboe, T. (2001). Formation and fate of marine snow: small-scale processes with large- scale
 implications. *Scientia Marina*, 65(S2).
- Maas, A. E., Miccoli, A., Stamieszkin, K., Carlson, C. A., and Steinberg, D. K. (2021). Allometry
 and the calculation of zooplankton metabolism in the subarctic Northeast Pacific Ocean. *Journal of Plankton Research*, 43(3):413–427.
- Mackas, D. and Bohrer, R. (1976). Fluorescence analysis of zooplankton gut contents and an investigation of diel feeding patterns. *Journal of Experimental Marine Biology and Ecology*, 25(1):77–85.

- McDonnell, A. M. P. and Buesseler, K. O. (2010). Variability in the average sinking velocity of
 marine particles. *Limnology and Oceanography*, 55(5):2085–2096.
- Mendez, M., Raiola, M., Masullo, A., Discetti, S., Ianiro, A., Theunissen, R., and Buchlin, J.-M.
 (2017).
- POD-based background removal for particle image velocimetry. *Experimental Thermal and Fluid Science*, 80.
- Middelburg, J. J. (2011). Chemoautotrophy in the ocean. *Geophysical Research Letters*, 38(24).
- 846 Muller, P. and Suess, E. (1979). Productivity, sedimentation rate, and sedimentary organic matter
- in the" oceans—I. Organic carbon preservation. Deep Sea Research Part A. Oceanographic
- 848 *Research Papers*, 26(12):1347–1362.
- Muller-Karger, F. E. (2005). The importance of continental margins in the global carbon cycle. *Geophysical Research Letters*, 32(1).
- Nobach, H. and Honkanen, M. (2005). Two-dimensional Gaussian regression for sub-pixel
 displacement estimation in particle image velocimetry or particle position estimation in particle
 tracking velocimetry. *Experiments in Fluids*, 38(4):511–515.
- Nowicki, M., DeVries, T., and Siegel, D. A. (2022). Quantifying the Carbon Export and
 Sequestration Pathways of the Ocean's Biological Carbon Pump. *Global Biogeochemical Cycles*,
 36(3).
- Omand, M. M., D'Asaro, E. A., Lee, C. M., Perry, M. J., Briggs, N., Cetinic, I., and Mahadevan,
 A.' (2015). Eddy-driven subduction exports particulate organic carbon from the spring bloom. *Science*, 348(6231):222–225.
- 860 Otsu, N. (1979). A Threshold Selection Method from Gray-Level Histograms. *IEEE Transactions*
- 861 *on Systems, Man, and Cybernetics*, 9(1).

- Passow, U. (2002). Transparent exopolymer particles (TEP) in aquatic environments. *Progress in Oceanography*, 55(3-4):287–333.
- Petrik, C. M., Jackson, G. A., and Checkley, D. M. (2013). Aggregates and their distributions
 determined from LOPC observations made using an autonomous profiling float. *Deep Sea Research Part I: Oceanographic Research Papers*, 74:64–81.
- Piao, S., Sitch, S., Ciais, P., Friedlingstein, P., Peylin, P., Wang, X., Ahlstrom, A., Anav, A.,
- Canadell," J. G., Cong, N., Huntingford, C., Jung, M., Levis, S., Levy, P. E., Li, J., Lin, X., Lomas,
- M. R., Lu, M., Luo, Y., Ma, Y., Myneni, R. B., Poulter, B., Sun, Z., Wang, T., Viovy, N., Zaehle,
- S., and Zeng, N. (2013). Evaluation of terrestrial carbon cycle models for their response to climate
- variability and to ¡scp¿ ¡scp¿ CO ¡sub¿2¡/sub¿ ¡/scp¿ ¡/scp¿ trends. Global Change Biology,
 19(7):2117–2132.
- Ploug, H., Iversen, M. H., and Fischer, G. (2008). Ballast, sinking velocity, and apparent diffusivity
- within marine snow and zooplankton fecal pellets: Implications for substrate turnover by attached
 bacteria. *Limnology and Oceanography*, 53(5):1878–1886.
- Romanelli, E., Giering, S., Estapa, M., Siegel, D., and Passow, U. (2024). Intense storms affect
 sinking particle fluxes after the North Atlantic diatom spring bloom. *bioRxiv*.
- Siegel, D. and Deuser, W. (1997). Trajectories of sinking particles in the Sargasso Sea: modeling
- of statistical funnels above deep-ocean sediment traps. *Deep Sea Research Part I: Oceanographic*
- 880 *Research Papers*, 44(9-10):1519–1541.
- Siegel, D. A., Cetinic, I., Graff, J. R., Lee, C. M., Nelson, N., Perry, M. J., Ramos, I. S., Steinberg,
- D. K.,' Buesseler, K., Hamme, R., Fassbender, A. J., Nicholson, D., Omand, M. M., Robert, M.,
- 883 Thompson, A., Amaral, V., Behrenfeld, M., Benitez-Nelson, C., Bisson, K., Boss, E., Boyd, P. W.,
- Brzezinski, M., Buck, K., Burd, A., Burns, S., Caprara, S., Carlson, C., Cassar, N., Close, H.,

- B85 D'Asaro, E., Durkin, C., Erickson, Z., Estapa, M. L., Fields, E., Fox, J., Freeman, S., Gifford, S.,
- 886 Gong, W., Gray, D., Guidi, L., Haentjens, N., Halsey, K., Huot, Y., Hansell, D., Jenkins, B., Karp-
- 887 Boss, L., Kramer, S., Lam," P., Lee, J.-M., Maas, A., Marchal, O., Marchetti, A., McDonnell, A.,
- 888 McNair, H., Menden-Deuer, S., Morison, F., Niebergall, A. K., Passow, U., Popp, B., Potvin, G.,
- Resplandy, L., Roca-Mart'i, M., Roesler, C., Rynearson, T., Traylor, S., Santoro, A., Seraphin, K.
- D., Sosik, H. M., Stamieszkin, K., Stephens, B., Tang, W., Van Mooy, B., Xiong, Y., and Zhang,
- X. (2021). An operational overview of the EXport Processes in the Ocean from RemoTe Sensing
- (EXPORTS) Northeast Pacific field deployment. *Elementa: Science of the Anthropocene*, 9(1).
- SierraI, C. A., Harmon, M. E., Moreno, F. H., Orrego, S. A., and Del Valle, J. I. (2007). Spatial and
- temporal variability of net ecosystem production in a tropical forest: testing the hypothesis of a
 significant carbon sink. *Global Change Biology*, 13(4):838–853.
- Smith, K. L., Ruhl, H. A., Huffard, C. L., Messie, M., and Kahru, M. (2018). Episodic organic
 carbon' fluxes from surface ocean to abyssal depths during long-term monitoring in NE Pacific.
- 898 *Proceedings of the National Academy of Sciences*, 115(48):12235–12240.
- Song, Y., Burd, A. B., and Rau, M. J. (2023). The deformation of marine snow enables its
 disaggregation in simulated oceanic shear. *Frontiers in Marine Science*, 10.
- Song, Y. and Rau, M. J. (2022). A novel method to study the fragmentation behavior of marine
 snow aggregates in controlled shear flow. *Limnology and Oceanography: Methods*, 20(10):618–
 632.
- Stamieszkin, K., Steinberg, D. K., and Maas, A. E. (2021). Fecal pellet production by
 mesozooplankton in the subarctic Northeast Pacific Ocean. *Limnology and Oceanography*,
 66(7):2585–2597.

- 907 Steinberg, D. K., Stamieszkin, K., Maas, A. E., Durkin, C. A., Passow, U., Estapa, M. L., Omand,
- 908 M. M., McDonnell, A. M. P., Karp-Boss, L., Galbraith, M., and Siegel, D. A. (2023). The Outsized
- Role of Salps in Carbon Export in the Subarctic Northeast Pacific Ocean. *Global Biogeochemical Cycles*, 37(1).
- 911 Stephens, B. M., Durkin, C. A., Sharpe, G., Nguyen, T. T. H., Albers, J., Estapa, M. L., Steinberg,
- D. K., Levine, N. M., Gifford, S. M., Carlson, C. A., Boyd, P. W., and Santoro, A. E. (2024). Direct
 observations of microbial community succession on sinking marine particles. *The ISME Journal*,
 18(1).
- 915 Thielicke, W. and Stamhuis, E. J. (2014). PIVlab Towards User-friendly, Affordable and
- 916 Accurate Digital Particle Image Velocimetry in MATLAB. Journal of Open Research Software, 2.
- 917 Trudnowska, E., Lacour, L., Ardyna, M., Rogge, A., Irisson, J. O., Waite, A. M., Babin, M., and
- Stemmann, L. (2021). Marine snow morphology illuminates the evolution of phytoplankton
- blooms and determines their subsequent vertical export. *Nature Communications*, 12(1):2816.
- 920 Trull, T., Bray, S., Buesseler, K., Lamborg, C., Manganini, S., Moy, C., and Valdes, J. (2008). In

situ measurement of mesopelagic particle sinking rates and the control of carbon transfer to the

- ocean interior during the Vertical Flux in the Global Ocean (VERTIGO) voyages in the North
- Pacific. Deep Sea Research Part II: Topical Studies in Oceanography, 55(14-15):1684–1695.
- Turner, J. (2002). Zooplankton fecal pellets, marine snow and sinking phytoplankton blooms. *Aquatic Microbial Ecology*, 27:57–102.
- Westerweel, J. and Scarano, F. (2005). Universal outlier detection for PIV data. *Experiments in Fluids*, 39:1096–1100.
- 928

929

930

931 Data Availability Statement

- 932 The data that support the findings of this study are openly available in in NASA's SeaBASS
- 933 archive at https://oceandata.sci.gsfc.nasa.gov/ob/getfile/dd2fe323be_EXPORTS-
- 934 EXPORTSNP_RR1813_GelCam_20180814-20180909_R1.sb and
- 935 https://oceandata.sci.gsfc.nasa.gov/ob/getfile/8a0152ccab_EXPORTS-
- 936 EXPORTSNA_JC214_GelCam_20210504-20210509_R1.sb.

SUPPLEMENTAL MATERIALS

Camera calibration

As the camera viewed through layers of acrylics and polyacrylamide gel, we printed a calibration target as shown in Figure S1 in order to measure the magnification ratio and reconstruct undistorted images. The calibration plate consisted of a grid of black dots used for establishing a mapping between the image and the real coordinates. Larger dots had a diameter of 2 mm with a spacing of 10 mm, while the diameter of smaller dots was 1 mm. During the calibration, the target was placed onto the top surface of the GelCam housing, which was also the bottom of the gel layer. Making sure the target was covering the entire field of view, we then captured the image of the calibration target. With a simple binarization, we computed centers of all black dots and mapped their locations with real coordinates. The distortion effects were corrected by fitting a polynomial function. As a result, we could estimate the pixel size in any location of captured images. The average magnification was 21.3 μ m per pixel during EXPORTS 2018 and 25.3 μ m per pixel during EXPORTS 2021.



Figure S1. Calibration image used for calculating magnification and correcting distortion during EXPORTS 2021.

Summary of tracking sequences and labeling system

Table S1 presents the number of binarized particles and tracking results from ten deployments. The shallowest depth, STT 1, recorded the most particles, while measurements in STT 2 observed half as many particles as STT 1. With a time interval of approximately 20 minutes, our tracking algorithm successfully monitored particles for more than ten hours (~30 frames) on average. Excluding fast-moving zooplankton, air bubbles, and other misidentified particles, the tracking duration could be further improved for particles categorized in Figure 4.

We employed a numerical labeling system to categorize various situations when the particle tracking was ceased. Specifically, the notation "-1" denoted instances when particles were leaving the FOV. Whenever no particle was successfully matched, and the best correlation coefficient was smaller than 0.5, we identified these cases using the label "-2". The label "-3" was assigned when we could not find a matched particle, particularly when the particle's area in the preceding frame was less than 100 pixels ($ESD \sim 240 \mu m$). Another distinct scenario arose when we failed to track a particle due to its insufficient average gray scale intensity lower than the global Otsu's threshold. In such cases, we marked the particle with the label "-4". Occasionally, we encountered moving zooplankton during our analysis, leading to the assignment of a "-5" flag. The notation "-6" represented all other situations where the tracking process was lost or unsuccessful. Lastly, we defined a "-7" where a particle was entering the FOV from the outside. The label "-7" was generated in a same way as "-1", when we ran the tracking algorithm in a reverse chronological

order. These designated labels enabled a systematic characterization of the diverse outcomes encountered throughout our investigation.

Table S1. Summary of particle tracking. Binarized particles represent particles contoured and preprocessed for tracking. Tracking particles include all particle categories with zooplankton and air bubbles, before performing a classification. Average tracking frames correspond to the mean duration of tracking. Each frame represents a time interval of 20 minutes approximately.

				-1	-2	-3	4-	-S-		-7
Deployme	nts	Contoured	Iracking	leaving	weak	small	dim	large	-0 others	entering
		particles	seducitices	FOV	correlation	areas	particles	swimmer	onicis	FOV
E.	STT1	118,775	13,786	3,244	3,827	1,731	937	06	1,926	2,180
NP Fnoch1	STT2	N.A.	N.A.	N.A.	N.A.	N.A.	N.A.	N.A.	N.A.	N.A.
THOOT	STT3	100,585	10,193	1,166	3,062	2,025	235	17	1,687	1,180
Ę	STT1	84,908	3,525	681	456	634	302	14	318	469
NP Enoch2	STT2	41,586	1,580	240	348	454	62	2	137	207
Thoodr	STT3	28,626	1,404	234	366	347	87	3	152	174
Ę	STT1	176,828	5,423	845	402	862	250	0	30	704
NF Fnoch3	STT2	94,347	3,328	758	571	673	93	4	296	583
CIIAOda	STT3	57,579	2,810	108	65	137	15	0	36	396
NA		32,068	1,184	163	89	300	169	2	58	100

Lighting correction of the white pixel method



Figure S2. Correction of the cumulative particle areas in white pixels. (a) linear interpolation during the nighttime, (b) linear fitting of the percentage attenuation against the blue light intensities, and (c) corrected number of white pixels.

We estimated the total area of deposited particles using grayscale images, which we derived from the red-channel image after the background subtraction. Again we applied the global Otsu's blackandwhite threshold. The number of white pixels that exceeded the global Otsu threshold provided a quick approximation of the total deposited particles. However, the presence of ambient light, especially at shallow depths, posed a challenge, as high ambient lighting intensities could lead to the underestimation of particle sizes due to light scattering effects.

To tackle this challenge, we developed a correction method based on the ambient lighting intensity. As highlighted in the section above, the predominant background illuminations were observed in the green and blue channels. Here we defined excessive blue light intensity as the difference in light intensity from the nighttime blue channel. By definition, the excessive intensity reached the maximum around noon. Based on the observation that particle areas became smaller under stronger ambient light, we assumed that the excessive intensity was linearly correlated with the percentage decrease in binarized particle area. It should be noted that the blue channel was used because the area attenuation was slightly more sensitive than the green light. To calculate the loss percentage in particle area, we interpolated linearly between the end of last night and the start of concurrent nighttime time stamps to estimate unaffected white pixel number. The loss percentage could then be computed based on the interpolated estimation and the captured particle areas. We then correlated the excessive blue light intensity with the percentage attenuation, allowing for a corrected white pixel number at each daytime time stamp. Despite occasional dips in the curve, the implementation of this correction yielded promising results quickly and efficiently, as detailed in below sessions. More detailed steps for this method can be found in supplemental materials.

Other supplemental figures and tables



Figure S3. Time series of surface PAR and average light intensities in blue (A) and green (B) channels. The average blue/green channel light intensities were normalized by the maximum average across all image frames.



Figure S4. (a) interrogation and search windows, and (b) the plane of correlation coefficients.

Table S2. Parameters that used to model the particulate carbon content of each particle type. We use the same particle types and parameters presented by Durkin et al. (2021) except dense detritus. Here *ESD* represents the equivalent spherical diameter obtained from particle projection areas, *l* represents the major axis length, while *w* is the minor axis length. Carbon mass per particle *C* (*mg*) is calculated by $C = A \cdot V^B$, where *A* is the scaling coefficient, *V* is the particle volume (μm^3), and *B* is a unitless exponent.

particle type	shape	volume	Α	В
aggregates	sphere	$\frac{\pi}{6ESD^3}$	$0.1 \cdot 10^{-9}$	0.80
large loose fecal pellets	cylinder	$\frac{\pi}{4}lw^2$	$0.1 \cdot 10^{-9}$	0.83
long fecal pellets	cylinder	$\frac{\pi}{4}lw^2$	$0.1 \cdot 10^{-9}$	1
dense detritus	sphere	$\frac{\pi}{6ESD^3}$	$0.1 \cdot 10^{-9}$	0.83
salp pellets	cuboid	$\frac{1}{4}lw^2$	$0.04 \cdot 10^{-9}$	1
rhizaria	sphere	$\frac{\pi}{6ESD^3}$	$0.004 \cdot 10^{-9}$	0.939



Figure S5. Accumulative fluxes at different deployments. Panels of (A), (B), (C), and (D) use the number of white pixels with the lighting correction method to quantify the time-integrated fluxes, while symbolized curves in (E), (F), (G), and (H) apply the particle classification results. Panel (A) also compares the cumulative fluxes of using the raw images or the ambient light correction in Deployment 1 STT 1. A zero duration on the x axis corresponded to the time when the deployment started. The significant discrepancy between (A) and (E) resulted from the storm during Deployment 1.



Figure S6. Time-varying POC fluxes of three particle types (aggregates in green, large loose and long fecal pellets in blue, dense detritus) in EXPORTS-NA. Shaded areas represent the nighttime.

Table S3. Coefficients of determination R^2 of different particle types across depths, with p-values in the bracket.

aggregates	STT 1	STT 2	STT 3
STT 1	1 (0)	0.91 (< 0.01)	0.84 (< 0.01)
STT 2	0.92 (< 0.01)	1 (0)	0.85 (< 0.01)
STT 3	0.85 (< 0.01)	0.84 (< 0.01)	1 (0)
large loose fecal pellets	STT 1	STT 2	STT 3
STT 1	1 (0)	0.31 (< 0.01)	0.03 (< 0.01)
STT 2	0.31 (< 0.01)	1 (0)	0.003 (0.28)

STT 3	0.03 (< 0.01)	0.003 (0.27)	1 (0)
long fecal pellets	STT 1	STT 2	STT 3
STT 1	1 (0)	0.26 (< 0.01)	0.13 (< 0.01)
STT 2	0.26 (< 0.01)	1 (0)	0.002 (0.33)
STT 3	0.17 (< 0.01)	0.002 (0.35)	1 (0)
dense detritus	STT 1	STT 2	STT 3
STT 1	1 (0)	0.005 (0.17)	0.07 (< 0.01)
STT 2	0.005 (0.17)	1 (0)	0.55 (< 0.01)
STT 3	0.06 (< 0.01)	0.55 (< 0.01)	1 (0)

Table S4. Coefficients of determination R^2 across particle types at the same depths, with p-values in the bracket.

STT 1	aggregates	large loose fecal pellets	long fecal pellets	dense detritus
aggregates	1 (0)	0.59 (< 0.01)	0.50 (< 0.01)	0.005 (0.16)
large loose fecal pellets	0.59 (< 0.01)	1 (0)	0.39 (< 0.01)	0.08 (< 0.01)
long fecal pellets	0.50 (< 0.01)	0.39 (< 0.01)	1 (0)	0.22 (< 0.01)
dense detritus	0.005 (0.16)	0.09 (< 0.01)	0.22 (< 0.01)	1 (0)
STT 2	aggregates	large loose fecal pellets	long fecal pellets	dense detritus
aggregates	1 (0)	0.09 (< 0.01)	0.14 (< 0.01)	0.85 (< 0.01)
large loose fecal pellets	0.09 (< 0.01)	1 (0)	0.39 (< 0.01)	0.09 (< 0.01)
long fecal pellets	0.14 (< 0.01)	0.39 (< 0.01)	1 (0)	0.12 (< 0.01)
dense detritus	0.85 (< 0.01)	0.09 (< 0.01)	0.12 (< 0.01)	1 (0)
STT 3	aggregates	large loose fecal pellets	long fecal pellets	dense detritus
aggregates	1 (0)	0.09 (< 0.01)	0.06 (< 0.01)	0.65 (< 0.01)
large loose fecal pellets	0.09 (< 0.01)	1 (0)	0.35 (< 0.01)	0.17 (< 0.01)
long fecal pellets	0.06 (< 0.01)	0.35 (< 0.01)	1 (0)	0.17 (< 0.01)
dense detritus	0.65 (< 0.01)	0.17 (< 0.01)	0.17 (< 0.01)	1 (0)