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# Mercury budget in global rivers at present-day: impacts from reservoirs and dams

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#### 25 Abstract

26 Many world rivers are currently polluted by mercury (Hg) compounds, leading to the bioaccumulation of 27 methylmercury (MeHg) in the food web, which poses potential health risks to humans. However, the riverine 28 Hg budgets of global scale remain poorly understood due to limited observations, complicating efficient 29 environmental governance. Here, we employ a process-driven Hg model to track its journey from sources to the 30 global ocean and assess its human exposure. Our findings indicate that ~1,500 Mg/yr of human-induced Hg are 31 released plus ~400 Mg/yr (mean) of soil erosional Hg are released to rivers. Due to the trapping effects of 32 reservoirs/dams, about 50% of the riverine Hg (~1,000 Mg/yr, mean) reaches the ocean. The different strengths 33 of the human-induced Hg releases and/or the numbers of (mega-) reservoirs/dams in regions would greatly 34 impacts the riverine Hg budgets of the corresponding regions. Human-induced Hg releases have led to an 35 accumulation rate of approximately ~1,000 Mg/yr in global reservoirs at present-day levels. This has resulted in 36 large Hg pools in reservoirs and dams, posing significant risks to future Hg pollutant restoration efforts. Our 37 model enhances the understanding of the fate of riverine Hg, providing critical information for riverine Hg 38 management and human health mitigation strategies.

#### 39 Introduction

40 River pollution by mercury (Hg), recognized as one of the most toxic heavy metals, poses significant threats 41 to global water resources and aquatic ecosystems (Mason et al. 2012, Obrist et al. 2018, Liu et al. 2021, Peng et 42 al. 2023). The riverine Hg budget is the vital part to the global Hg budget as the connections between the Hg 43 pools of land and ocean (Liu et al. 2021). Significant amounts of wastewater and Hg-containing wastes have 44 been discharged into the global rivers, leading to elevated Hg concentrations in the related streams and aquatic 45 wildlife (Eagles-Smith et al. 2009, Driscoll et al. 2013, Schoch et al. 2014). This impairs the ability of polluted 46 rivers to provide clean water and safe food for human consumption (Zhang et al. 2021). Environmental policies, 47 such as the Clean Water Act, freshwater fish consumption advisories, and the Minamata Convention, have been 48 implemented to control river Hg pollution and its impact on human health (Cain et al. 2011, Wu et al. 2022). While studies have been conducted over the past decades to monitor Hg levels in river systems (Schmeltz et al. 49 50 2011, Liu et al. 2021), our understanding of its sources, transport mechanisms, fate, and drivers at a global scale 51 remains limited.

Mercury in rivers stems from various natural and anthropogenic sources, including soil erosion, industrial wastewater (e.g., metal smelting and/or chlor-alkali production etc.), and artisanal and small-scale gold mining (ASGM) (Streets et al. 2019). At present, anthropogenic releases are the primary contributors to global riverine Hg (Kocman et al. 2017). However, relatively large uncertainty is associated with the existing releases inventories to rivers. For example, Kocman et al. (2017) estimated that approximately 1,000 Mg (300-1300 Mg) of anthropogenic Hg is discharged into global water systems annually, with contributions of 220 Mg from point sources, 40 Mg from the remobilization of contaminated systems, and 440 Mg from artisanal and small-scale 59 gold mining (ASGM). In another study, Streets et al. (2017) estimated that global Hg releases to land and water 60 reached up to 7,280 Mg in 2010 but noted the challenge in separating releases between land and water due to a 61 lack of data.

62 Riverine Hg primarily exists in the form of particulate Hg within suspended particulate matter (SPMs) in 63 rivers (Schuster et al. 2011), with only a small fraction being dissolved when the total Hg concentration are in relatively higher level than the natural conditions (Lu et al. 1987, Kocman et al. 2010, Liu et al. 2021), thus the 64 65 riverine Hg are largely influenced by the sediment fluxes (Liu et al. 2021, Peng et al. 2023). The nature sources, especially the soil erosion, serve as the important sources to the riverine Hg concentrations and fluxes (Liu et al. 66 2020), and showing impacts to the Hg concentrations in the SPMs in the rivers (Kocman et al. 2010, Peng et al. 67 68 2023). Human activities, particularly water management practices, profoundly impact the processes governing 69 riverine Hg transport in contemporary times, especially through the construction of reservoirs and dams at a 70 global scale. Reservoirs have substantial efficacy in trapping SPMs (Vörösmarty et al. 2003), leading to the 71 potential sedimentation of riverine Hg. For example, the Three Gorges Dam in China has resulted in the settling 72 of 100-200 Mg of riverine Hg in the bedload (Peng et al. 2023) and has reduced the downstream flux of riverine 73 Hg by more than 70% (Liu et al. 2020). However, the overall impact of reservoirs and dams on river Hg transport 74 at a global scale remains unknown.

75 Here, we develop a process-based global Hg model (MOSART-Hg-wrm) based on the MOdel for Scale 76 Adaptive River Transport (with water management module, MOSART-wrm) included in the Community Earth 77 System Model (CESM2) to simulate the transport and fate of Hg in global rivers at present-day (Peng et al. 78 Under Review). We consider different release scenarios to bracket the uncertainty of the emission inventories 79 and to quantify the response of river Hg levels to different emission sources. The MOSART-Hg-wrm model 80 accounts for soil Hg erosion, the transport of Hg, and its sedimentation within river systems and corresponding 81 watersheds (Peng et al. Under Review). We develop a new model capacity to track the trapping effect of 82 reservoirs and dam systems on Hg based on the MOSART water management module (wrm) (Zhou et al. 2020). 83 The riverine Hg observation records reported in the literature are used to evaluate the model results. Our 84 objectives are (1) to develop a riverine Hg model at the present-day with human-induced releases and effects of 85 reservoirs/dams; (2) to comprehensively understand the processes governing the transport of Hg from its sources 86 to sinks (such as oceans and terrestrial lakes); and (3) to identify the accumulation rate of the Hg pools in the 87 global reservoirs/dams.

#### 88 The global riverine Hg budget

Our modeling effort results in a global budget of riverine Hg, including its sources, transport, and fate (Fig. 1). Anthropogenic Hg releases approximately 7,300 Mg/yr (mean) into land and water (Streets et al. 2017), with about 20% of these entering rivers by our simulation, translating to roughly 1,500 (mean, range from 1,100-2,000) Mg/yr entering river systems. Of these human-induced Hg releases to rivers, 30%-47% originate from ASGM. The gross Hg inputs to rivers, consisting of erosional and human-induced Hg releases, total approximately 1,900 Mg/yr, with around 30% of this amount contributed by erosional soil Hg fluxes (~400
Mg/yr).

96 Besides, Atmospheric Hg deposition has limited short-term impacts on erosional Hg fluxes through changes 97 in soil Hg concentration in the watershed, indirectly increasing erosional Hg by increasing soil Hg by time (see 98 Uncertainty analysis, *Deposition Experiment*). The long-term impacts of atmospheric Hg deposition on the 99 watershed are accounted for in this simulation through the incorporation of soil Hg concentration changes from 100 pre-industrial to the present-day by using the present-day soil Hg concentration dataset. Due to the barrier effect 101 of dams and reservoirs, 40-50% of riverine Hg settles with sediments in these reservoirs. Riverine Hg from 102 ASGM sources has a higher likelihood of reaching the oceans because reservoirs and dams predominantly align 103 with industrial Hg sources (e.g., lower hydropower prices encourage more industrial activities).

Less than 20 Mg/yr of Hg flux is exchanged between rivers and the atmosphere due to the limited air-water
 exchange area of rivers. Only 110 Mg/yr of Hg sedimentation occurs in the riverbed, which could be resuspended

106 by the wash effect from flooding pulses (mostly during wet seasons or extreme precipitation events). Ultimately,

 $\sim$  1,000 (777-1,146) Mg/yr of riverine Hg is exported to the ocean, with clear spatial distribution across continents,

- 108 notably higher riverine Hg fluxes export to the coasts in Asia and South America than in other regions.
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**Fig. 1 Pathways (direct and indirect) and budget** [unit: Mg/yr, mean (range)] **of Hg in global rivers**. POWERGEN = power generation, INDS = industry, ASGM = artisanal and small-scale gold mining.

113 Hg inputs to the global rivers

The inputs of Hg to rivers include nature sources (soil erosion, dissolved phase Hg in water, volcano eruption etc.) and human-induced releases, with the latter as the predominant source at the present-day (Fig. 1,

116 2). Due to the complex of simulating the other natural sources, only soil erosion is considered in this study. We 117 consider the Baseline Scenario [based on the revised the Arctic Monitoring and Assessment Programme (AMAP) 118 emission scenario, see Method] wherein the gross anthropogenic Hg releases to rivers is 1,500 Mg/yr, 119 representing the present-day conditions [with ~1,000 Mg/yr global riverine Hg budget by Liu et al. (2021)] (see 120 Method, Table S1). We categorize anthropogenic Hg releases into two categories to align with the AMAP 121 inventory (Steenhuisen et al. 2019, Steenhuisen et al. 2022): Industrial and ASGM (Fig. 2). The spatial 122 distributions of anthropogenic Hg sources in North America are based on the Streets inventory (Streets et al. 123 2019) with other continents on the AMAP inventory (see Methods). Hence, the higher industrial Hg releases are 124 mostly located in East Asia, South Asia, West Europe, and part of South Africa, which are align with the 125 atmospheric Hg emissions (Fig. 2a). And the ASGM Hg releases are mostly located in South and Central America, 126 Southeast Asia, and South-Central Africa (Fig. 2b). 127 The contribution of soil Hg erosion to riverine Hg is relatively small (Fig. 2c), influenced by factors such 128 as precipitation, vegetation, soil characteristics, slope, and runoff (Dutta 2016, Peng et al. 2023, Peng et al. Under 129 Review). In general, erosional Hg contributes 420 (400-440) Mg/yr to rivers, with a spatial distribution similar 130 to that of the pre-industrial era (Peng et al. Under Review). Regions adjacent to plateaus, such as those near the Qinghai-Tibet Plateau in East Asia, South Asia and Southeast Asia, as well as areas near mountains, such as the 131 132 west coast of South America near the Andes Mountains, show higher erosional Hg fluxes. The model also

133 simulates significantly higher Hg erosion flux in the Amazon watershed at present-day (39.2 Mg/yr) than in the

134 pre-industrial era (26.2 Mg/yr) due to deforestation in the region (Peng et al. Under Review). Similarly, the

135 deforestation of the rainforest also causes higher Hg fluxes from soil erosion in Myanmar and Bangladesh

136 (Southeast Asia) (Feinberg et al. 2024).







#### 140 Impacts from reservoirs/dams to riverine Hg (riverine processes of Hg)

We find that reservoirs/dams can significantly influence global gross riverine Hg budget by reducing the transport flux by 40-50% (the sedimentation of riverine Hg in river channel are not included) (Fig. 1, 3). The Hg form may change after settled into the bed of reservoirs, such as methylation and/or to dissolved phase, but in this study, we assume the Hg fluxes are only impacted by resuspension at this stage due to the relatively smaller

145 fluxes related the biological activities and other activities (Liu et al. 2019). Water management activities, such 146 as reservoirs and dams, play a significant role in reducing riverine Hg fluxes and SPM in downstream rivers 147 (Syvitski et al. 2005, Syvitski et al. 2022). This is apparent in the spatial pattern of riverine Hg fluxes along 148 channels (Fig. 2). The increasing trends of fluxes downstream can be interrupted by reservoirs and dams, which 149 trap Hg input from upstream and decrease the flux significantly, and the fluxes continue increasing after passing 150 the dams. We conducted paired sensitivity experiments with no water management (Water Management 151 *Experiment*, see Method) (Table S3). The role of reservoirs and dams can be evaluated based on the differences 152 between the paired control experiment. For example, one paired model run includes No Human-induced Hg Release with and without Reservoir or NHHR w/wt R. The other paired model run includes Baseline Hg-induced 153 154 Hg Releases with and without Reservoir or BHHR w/wt R (Fig. 3). 155 The fluxes of riverine Hg delivery to the ocean are decreased by 41% from 417 Mg/yr (NHHR wt R) to 248

156 Mg/yr (NHHR w R). If we consider the anthropogenic Hg releases, reservoirs/dams reduce Hg fluxes to ocean 157 by 50% from 2,021 Mg/yr (BHHR wt R) to 1,014 Mg/yr (BHHR w R). The difference between the two pairs of 158 model runs reflect the large spatial variabilities in the impact of reservoirs (Fig. 3bd). For the setting with baseline 159 anthropogenic Hg release, the riverine Hg exports of Yangtze, Yellow, Parana, and Niger River have been reduced by 77%, 75%, 63% and 61%, respectively. This variability in reduction largely reflects the number of 160 161 reservoirs or/and the portion of mega reservoirs along these rivers. Indeed, some rivers show limited impacts 162 from water management due to less reservoirs and/or less mega reservoirs along the river, such as Congo, Amazon, and Ganges-Brahmaputra River, which have been reduced by only <1%, 1%, and 17%, respectively. 163 164 As for the no human-induced Hg release scenarios, the pattern is similarly to the above (Fig. 3bd).

165 Reservoirs regulate river flooding and consequently the seasonal fluctuations of Hg fluxes, even though human-induced Hg releases to rivers are constant across different seasons in the estimations. For example, flood 166 pulses driven by increased runoff and larger soil erosional fluxes during the wet season contrast with reduced 167 168 fluxes during the dry season, thereby altering soil Hg erosion fluxes. The differences between monthly global 169 Hg exports have been reduced from 22.0-52.5 Mg/month (NHHR wt R) to 12.5-29.3 Mg/month (NHHR w R) 170 with the amplitude reduced by 45% from 30.6 to 16.8 Mg/month. The monthly Hg export also decreased from 171 150.3-196.6 to 74.6-94.3 Mg/month under the BHHR w/wt R with amplitude reduced by 58% from 46.3 to 19.7 172 Mg/month. The damping effects of dams are more pronounced in the Northern Hemisphere (Fig. 3ef), attributed to the larger landmass, and numbers of rivers and reservoirs as compare to the Southern Hemisphere (Lehner et 173 174 al. 2011, Mulligan et al. 2020). Furthermore, this effect is amplified by the larger anthropogenic Hg releases, 175 leading to significant differences in riverine Hg fluxes during summer in the Northern Hemisphere (Fig. 3e). In 176 conclusion, the reservoirs/dams shrink the seasonal changes of riverine Hg.



Fig. 3 Impact of reservoir and dams on riverine Hg transport. a) Riverine Hg fluxes in global rivers under *No Human-induced Hg Release without Reservoir scenario (NHHR wt R)*, b) Differences of riverine Hg fluxes between *No Human-induced Hg Release with/without Reservoir (NHHR w/wt R)*, c) Riverine Hg fluxes in global rivers under *Baseline Human-induced Hg Releases without Reservoir scenario (BHHR wt R)*, d) Differences of riverine Hg fluxes between *Baseline Human-induced Hg Releases with/without Reservoir scenario (BHHR wt R)*, d) Differences of riverine Hg fluxes between *Baseline Human-induced Hg Releases with/without Reservoir (BHHR wt R)*, e) Monthly riverine Hg fluxes to ocean in Northern Hemisphere, f) Monthly riverine Hg fluxes to ocean in Southern Hemisphere.

#### 184 **River Hg fluxes and budgets**

185 The modeled global riverine Hg discharge to the ocean/lakes amounts to ~1,000 Mg/yr in Baseline scenario 186 (also the BHHR w R) (Fig. 1, 2). The top ten largest riverine Hg exporters contribute 51% (506 Mg/yr) of the global discharge, with the highest river as the Amazon (202 Mg/yr), followed by the Congo (87 Mg/yr), Yangtze 187 (52 Mg/yr), and Ganges-Brahmaputra (51 Mg/yr). These large rivers dwarf the export from smaller rivers in 188 189 these regions (less than 30 Mg/yr), especially the smaller ones along the north coast of South America, the Indian 190 Peninsula, the eastern China, and Southeast Asia. The denser river networks in East Asia, South America, and 191 Central Africa could result in higher riverine Hg fluxes due to the intersection of the rivers within the large 192 watershed. For example, the Ganges River and the Brahmaputra River share the same estuary.

193 The modeled riverine Hg fluxes exhibit large spatial variations and seasonal variabilities (Fig. 4 and S1). Generally, the flux of Hg in the most river increases downstream along the channel, reaching its maximum at 194 195 the river mouths, if there are no major dams. This pattern reflects the continuous input of Hg from wastewater 196 and soil erosion in their watersheds and the high transport efficacy of SPM by rivers (Peng et al. 2024). For 197 example, the riverine Hg fluxes keep increasing in the upstream of Three Gorges Dam (0-24 Mg/yr) and this 198 trend resume after Three Gorges Dam in the channels of Yangtze River (4-52 Mg/yr) (Fig.2). The rivers like 199 Panama and Mekong River with increased riverine Hg fluxes along the channels and reach maximum at the river 200 mouth, while this trend is less influenced by dams.

201 Rivers with higher anthropogenic releases and/or erosion in their watersheds can easily maintain higher Hg 202 fluxes in their channels, such as the Amazon, Congo, Yangtze, and Ganges-Brahmaputra Rivers (Fig. 2 and 4). 203 For example, the highest riverine Hg flux in the Amazon River (201.7 Mg/yr near its mouth) is caused by the 204 effect of ASGM releases and erosional Hg flux to the watershed (152.5 and 32.2 Mg/yr, respectively). Similar 205 situation applies to the Congo River. In contrast, the higher riverine Hg fluxes of Yangtze (52 Mg/yr at mouth) 206 and Ganges-Brahmaputra River (51.2 Mg/yr at mouth) are caused by the higher industrial releases (100.7 and 207 37.8 Mg/yr, respectively) and erosional Hg flux (from Himalaya to ocean; 23.6 and 49.6 Mg/yr. respectively) in these regions. Indeed, the watersheds of these rivers include the most industrialized regions in eastern China 208 209 (Fig. 2a). Interestingly, some smaller rivers with relatively small watersheds, such as Kapuas River (14.6 Mg/yr 210 at mouth), could also have relatively higher riverine Hg fluxes due to their intense anthropogenic Hg releases 211 (Baiq 2018).

212 The riverine Hg fluxes estimated by model agree well with the observation of Hg levels in river mouths 213 (Fig. 4b). The observational dataset includes 44 rivers from different continents spanning five orders magnitudes with a geo mean flux of  $1 \times 10^{3.99 \pm 3.9}$  kg/yr. The model ( $19 \times 10^{3.90 \pm 3.7}$  kg/yr) reproduces the observations with a 214 high coefficient of determination ( $R^2 = 0.97$ , p<0.01). The model results agree well with observation records for 215 216 many major rivers, such as Yangtze (model vs. observation: 52 vs. 48 Mg/yr), Ob' (3.2 vs. 2.4 Mg/yr), and 217 Hudson (0.3 vs. 0.33 Mg/yr). Although the model has performed well in simulating erosion of small rivers (Tan 218 et al. 2022), due to the coarse anthropogenic releases the model has higher bias for smaller rivers, such as Kolyma 219 (model vs. observation: 4.7 vs. 1.1 Mg/yr), Mackenzie (2.6 vs. 0.6 Mg/yr) and Yukon (3.8 vs. 0.9 Mg/yr). These 220 rivers exhibit relatively lower riverine Hg fluxes compared to larger rivers, and they appear to be much more 221 sensitive to the local anthropogenic Hg releases. For instance, discrepancies in releases in large watersheds may 222 be mitigated by the canceling effect among grids with the same watershed, but such effect is less likely to occur 223 in smaller watersheds. Furthermore, the model agrees with observations taken along the river channels, which is 224 available for a few major rivers. For example, the model (24 Mg/yr) closely resembles the observation (26 Mg/yr) 225 upstream of the Three Gorges Dam (TGD), as it does at the immediate downstream of the TGD (4.7 and 4.4 226 Mg/yr) (Liu et al. 2020).

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Fig. 4 Riverine Hg fluxes in global rivers. a) Spatial patterns with zoomed regions of major rivers; b) Model results versus observation records in continents.

#### 231 Impacts of Human-induced Hg Releases

232 We consider five scenarios to cover the uncertainty of anthropogenic Hg releases from industrial and ASGM 233 sources to the rivers at present-day: (1) Low Release (1,100 Mg/yr), (2) Moderate Release (1,500 Mg/yr) (similar 234 level with Baseline Scenario), (3) High Industrial Release (2,000 Mg/yr, ASGM contribute 600 Mg/yr), (4) High 235 ASGM Release (2,000 Mg/yr, ASGM contribute 850 Mg/yr), (5) No Release (0 Mg/yr) (Fig. 1, 5, Table S1). The 236 Low Release scenario represents the unique period during the COVID-19 pandemic when industrial and ASGM 237 releases are significantly reduced due to lockdown measures (Wu et al. 2021) (Fig. 5a). The Moderate Release 238 scenario depicts a more realistic post-Minamata Convention scenario with reduced anthropogenic Hg releases 239 (Liu et al. 2021) (Fig. 5b). The High Industrial Release scenario mirror the condition around the year 2000, with relatively higher fraction of wastewater as mismanaged (Amos et al. 2014) (Fig. 5c). The High ASGM Release 240 241 scenario represents the early phase of the Minamata Convention, when the industrial releases declined rapidly 242 but the ASGM was still growing due to poor environmental management in the mining countries (Cheng et al. 243 2022, Prescott et al. 2022, Peng et al. 2023) (Fig. 5d). The No Release scenario is an idealized scenario with 244 only natural sources at present-day. The different sectors contributions of Hg sources among scenarios also 245 facilitate an assessment of the sensitivity of riverine Hg to anthropogenic releases. The results indicate that the anthropogenic Hg releases have a significant impact on the dynamics of riverine 246

247 Hg (Fig. 5e-g). The input of Hg to rivers is increased by higher anthropogenic Hg releases, while the contribution 248 of soil erosion remains stable under different scenarios. The fate of anthropogenic Hg releases, including riverine 249 inputs and exports to ocean, also varies among scenarios. For instance, despite comparable levels of total releases 250 between High Industrial Release and High ASGM Release scenarios, the former exhibits a lower export fraction 251 (Fig. 5f). With the Moderate Release scenario as a reference, only 28% of the increased industrial releases are 252 ultimately transported to the ocean for the High Industrial Release scenario (Table S2), but 54% exported for 253 the High ASGM Release scenario. 254 Higher industrial releases contribute to elevated riverine budgets in rivers of East Asia and Southeast Asia, while increased ASGM releases notably amplify riverine Hg budgets in the Amazon River Watershed, Central 255 256 Africa, and Southeast Asia. This is attributed to the difference of water management among regions. For instance, 257 the Amazon River, with active ASGM activities in its watershed, has fewer reservoirs/dams and/or fewer mega 258 reservoirs/dams, and subsequently higher export fractions. In contrast, the Yangtze River has higher industrial 259 releases in its watershed and impact from reservoirs and dams (Fig. 3b). Indeed, regions with thriving industries 260 often implement more extensive water management practices to provide water and hydroelectricity for reservoirs (Chen et al. 2016). This shift in water management could result in lower riverine Hg export fractions of industrial 261 releases (Fig. 5g). 262

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Fig. 5 Impacts of river Hg budget associated with anthropogenic Hg releases to rivers. a-d) Anthropogenic Hg releases under different release scenarios, e) Annual inputs and exports in global rivers under different release scenarios, f) Monthly inputs and exports in global rivers under different release scenarios, g) Riverine Hg discharges to the ocean in selected rivers under different release scenarios.

#### 269 Sink of Hg in Reservoirs/Dams

Increased Hg releases from human activities lead to higher Hg accumulation rates in reservoirs, the major
 sink for riverine Hg (Fig. 6). Globally, Hg accumulation rates in reservoirs rise approximately 6-fold from 14
 Mg/month in the *No Release scenario* (representing pre-industrial era) to 84 Mg/month in the *Baseline Release*

273 scenario (representing present-day). The contribution of different sectors (e.g., industrial vs. ASGM sources) to

- Hg releases has a smaller effect on accumulation rates compared to riverine Hg fluxes, as shown by similar rates
- 275 in the High Industrial Release (106 Mg/month) and High ASGM Release (104 Mg/month) scenarios (Fig 6a).
- 276 Due to the dominance of human-induced releases over natural sources, the overall increase in accumulation
- shows no apparent seasonal variation, remaining relatively consistent month to month (Fig. 6a).





Fig. 6 Reservoir storages of riverine Hg. a) Storage and monthly accumulation (Mg/mon) of Hg in global
reservoirs, b) final storages of Hg during 10 years of simulation in large reservoirs/dams, c) monthly
accumulation rates (Mg/mon) of Hg storages in global reservoirs/dams under *Baseline Release Scenario*.

Reservoirs exhibit significant regional variabilities in the rates of Hg accumulation, influenced by reservoir density and watershed Hg releases (Fig. 6c). In East and South Asia, higher reservoir densities combined with substantial industrial Hg releases (Fig. 5) result in elevated Hg accumulation rates. In contrast, the US, despite its high reservoir density, has relatively lower Hg releases, leading to low to moderate accumulation rates. Reservoirs in parts of South Africa and South America exhibit high accumulation rates, consistent with elevated Hg releases from ASGM activities (Fig. 5). Meanwhile, Southern Europe, with a a high density of reservoirs but lower regional Hg releases, shows relatively lower Hg accumulation rates.

289 Several specific reservoirs illustrate the dominance of human-induced Hg releases on accumulation patterns.

290 For example, the Owen Falls Reservoir Group (upstream of the Nile), the High Aswan Dam (Nile), the

291 Sanmenxia Reservoir (Yellow River), the Three Gorges Dam (Yangtze River), and the Macagua Dam (Caroní

292 River) show substantial increases in Hg accumulation between *No Release* to *Low Release* scenarios, with even

293 larger increases from Low to Moderate or High Release levels (Fig. 6b). Sectoral contributions also play a critical

294 role: while reservoirs like the Sanmenxia and the Three Gorges experience greater accumulation under the *High* 

295 Industrial Release scenario, those such as the Owen Falls Reservoir Group see much higher accumulation rates

296 under High ASGM Release scenario than the High Industrial Release one.

#### 297 Implications

298 Anthropogenic Hg releases significantly augment riverine Hg delivery from upstream to downstream and 299 export into inland lakes and/or oceans (Liu et al. 2021, Yu et al. 2021). However, measuring riverine Hg fluxes 300 is challenging due to rapid changes in runoff parameters and considerable variability in Hg concentrations. The 301 lack of historical data also hinders the proper evaluation of the trends of riverine Hg export and the contribution 302 of anthropogenic Hg releases, posing management challenges, particularly in international rivers such as the 303 Mekong, Jordan, Indus, Niger, and Nile Rivers (Campbell 2007, Campbell 2009). Our estimation provides 304 insights into understanding the regional-based contribution of anthropogenic Hg releases to downstream rivers 305 and oceans. The Hg concentrations in the rivers have strong connection with the fish Hg (Scudder et al. 2009, 306 Blanchfield et al. 2021), hence this simulation could help understand the Hg pollution status in the rivers and 307 coastal environment, which is vital for fishing and aquaculture. Furthermore, the settled Hg in reservoirs creates 308 a suitable environment for Hg methylation, with conditions similar to nutrient-rich wetlands that support 309 biological activity and methylation. Our results provide valuable insights for future studies to better understand 310 the amounts and changes of Hg pools in reservoirs, which are important for estimating and observing the 311 methylation in the reservoirs.

312 Our assessments emphasize the critical role of reservoirs as significant reservoirs in global Hg 313 biogeochemistry, despite the uncertainty surrounding their masses. The presence of settled Hg pools within 314 global reservoirs introduces complexities to ongoing environmental governance efforts on a global scale. 315 Although water management benefits agriculture and related industries, Hg settled in reservoir bottoms and 316 riverbeds could potentially act as potential Hg sources to rivers (Liu et al. 2020, Peng et al. 2023). For instance, 317 the observation in Yangtze River noticed that the settled Hg pool of the Three Gorges Dam shift from sink to the 318 sources of riverine Hg response to the decreased riverine Hg fluxes from upstream and tributaries (Peng et al. 319 2023). Our model serves as a scientific tool to guide riverine Hg management strategies and identify effective 320 pathways for reducing Hg releases, especially those impacting human health. Furthermore, our estimation of 321 accumulation rates in global reservoirs can aid in monitoring Hg pools during restoration projects and in 322 identifying potential risks associated with reservoir or dam management such as reservoirs/dams demolition.

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432

#### 433 Method

#### 434 Model platform

435 The MOdel for Scale Adaptive River Transport for Mercury under WateR Management (MOSART-Hg-436 wrm) is based on the MOSART-Hg model used in an accompanying study (Peng et al. Under Review), with 437 an addition of the water management module integrated into the MOSART-Hg-wrm model (Voisin et al. 438 2013, Zhou et al. 2020) (Table S3). The model incorporates soil Hg erosion as an input of riverine Hg from 439 natural and legacy sources and considers the contribution of anthropogenic releases. It accounts for the 440 downstream transport of riverine Hg or its settling in riverbeds, while also incorporating the trapping effects 441 of dams on riverine Hg. Implemented within the Community Earth System Model version 2.1.3 (CESM2 442 version 2.1.3), MOSART-Hg-wrm incorporates features developed from the Energy Exascale Earth System 443 Model (E3SM) (Zhou et al. 2020), including the sediment transport module (Li et al. 2022) and erosion 444 module (Tan et al. 2018, Tan et al. 2022). These modifications enable the simulation of riverine Hg dynamics 445 under contemporary environmental conditions.

446 The model accounts for water management practices, represented by reservoir constructions globally, 447 allowing the determination of reservoir trap efficiency (Vörösmarty et al. 2003, Voisin et al. 2013, Zhou et 448 al. 2020), albeit without considering reservoir management actions such as storage adjustments. The model 449 facilitates the determination of trapping efficiency by leveraging input datasets of global dams through the 450 calculation method outlined by Vörösmarty et al. (2003). Consequently, the sediment trapping processes can 451 be effectively ascertained. This methodology finds broad application in hydrology and biogeochemical 452 models, exemplified by its utilization in models like Global Nutrient Export from Watersheds 2 (NEWS 2) (Mayorga et al. 2010). Data on global reservoirs are sourced from the Global Reservoir and Dam (GRanD) 453 454 database (Lehner et al. 2011), initialized on a grid basis by Zhou et al. (2020). This dataset encompasses dam 455 locations, reservoir capacities, and major functions for over 4,200 dams worldwide (Table S3). The trapping 456 efficiency is calculated as followed (Vörösmarty et al. 2003, Li et al. 2022):

457 
$$e_{trap} = 1 - \frac{0.05}{\Delta t_{local}^{0.05}}$$
(1)

458 where  $\Delta t_{local}^{0.05}$  is the increase of local water residence time due to the reservoir [years], estimated as the effective 459 reservoir storage capacity divided by the mean annual inflow from the reservoir upstream.

460 The simulations span the period 1995-2014, with the initial five years dedicated to make the river 461 systems into steady state (i.e., spin-up), and the subsequent five years are considered for analysis. The spatial 462 resolution matches that of previous studies, with the Community Land Model 5.0 for Mercury (CLM5-Hg) 463 and MOSART-Hg-wrm running at resolutions of  $0.9^{\circ} \times 1.25^{\circ}$  and  $0.5^{\circ} \times 0.5^{\circ}$ , respectively. The former has 464 been extensively evaluated in simulating erosional Hg dynamics (Peng et al. Under Review). We include soil Hg datasets from Wang et al. (Wang et al. 2019) and the Liu et al. (Liu et al. 2023). Climate data, obtained 465 from the GSWP3 dataset, drive the model at a  $0.5^{\circ}$  horizontal resolution, spanning the same temporal range 466 as previous studies (1995-2014) with a six-hour time resolution. Climate dataset parameters include 467 Precipitation (mm H<sub>2</sub>O/sec), Incoming Solar Radiation (W/m<sup>2</sup>), Temperature (K), Pressure (Pa), Winds (m/s), 468 469 Humidity (kg/kg), and Downward Longwave Radiation (W/m<sup>2</sup>).

#### 470 Anthropogenic Hg release for rivers

471 We use an inverse estimation process to establish a plausible range for global Hg releases to rivers. 472 Existing inventories includes that developed by Kocman et al. (2017) (hereafter Kocman inventory), which 473 suggests a global release of Hg to rivers as 1,000 Mg/yr (lower bound). The other inventory is developed by 474 Streets et al. (2017) (hereafter referred to as Streets inventory), which has a global release of Hg to the land 475 and rivers as 7,280 Mg/yr, but the proportion release to land and water are unknown. As only global total 476 releases are provided by these studies, the spatial distribution of Hg sources is based on atmospheric Hg 477 release inventories, namely AMAP inventory (Steenhuisen et al. 2019, Steenhuisen et al. 2022) and Streets Atmospheric inventory (Streets et al. 2019). Indeed, the anthropogenic Hg emissions/releases to atmosphere, 478 479 land, and water have similar spatial distribution as they share the major contributing sources (Streets et al. 480 2017). With the MOSART-Hg-wrm model, the Hg release to rivers can be constrained by the global amount 481 of Hg exported to the ocean, ~1,000 Mg/yr based on river monitoring data at river mouths (Liu et al. 2021). Preliminary model simulation indicates that direct driving the MOSART-Hg-wrm with the Kocman 482 483 inventory generates too low Hg export to the ocean.

484 The Hg release from sources to rivers are categorized as direct and indirect release. Direct release 485 involves Hg discharge into rivers via urban sewage pipe networks, while indirect release entail Hg initially 486 deposited into the soil, subsequently mobilized by erosional processes, and transported to rivers. The 487 sediment yield serves as a crucial parameter for determining indirect release. In this study, the sediment yield 488 dataset represents the mean sediment yield in 2010, estimated by MOSART-Hg-wrm under the No Emission 489 scenario. We also find that altering the portion of direct and indirect release to rivers does not affect the 490 calculated yield.

491

The grid-based anthropogenic Hg releases to rivers are calculated by the following equations:

492 
$$EI_{ASGM}^{THg} = AMAP_{ASGM}^{THg} \times Sed_{yld} \times Coef_{ASGM}^{I}$$
(2)

493 
$$EI_{INDS}^{THg} = AMAP_{INDS}^{THg} \times Sed_{yld} \times Coef_{INDS}^{I}$$
(3)

494 
$$EI_{POWERGEN}^{THg} = AMAP_{POWERGEN}^{THg} \times Sed_{yld} \times Coef_{POWERGEN}^{I}$$
(4)

495 where  $EI_{ASGM}^{THg}$ ,  $EI_{INDS}^{THg}$ ,  $EI_{POWERGEN}^{THg}$  represent the indirect release by ASGM, industrial, or power-generation 496 sectors, respectively (unit: g/km<sup>2</sup>/yr);  $AMAP_{ASGM}^{THg}$ ,  $AMAP_{INDS}^{THg}$ ,  $AMAP_{POWERGEN}^{THg}$  represent the spatial 497 distribution of release from different sector (unit: g/km<sup>2</sup>/yr);  $Sed_{yld}$  represents the sediment yield (unit: 498 kg/m<sup>2</sup>/s); and  $Coef_{ASGM}^{I}$ ,  $Coef_{INDS}^{I}$ ,  $Coef_{POWERGEN}^{I}$  represent the coefficient of Hg emission associated with 499 erosion (unit: m<sup>2</sup>s/kg).

$$ED_{ASGM}^{THg} = AMAP_{ASGM}^{THg} \times Coef_{ASGM}^{D}$$
(5)

501 
$$ED_{INDS}^{THg} = AMAP_{INDS}^{THg} \times Coef_{INDS}^{D}$$
(6)

$$502 ED_{POWERGEN}^{THg} = AMAP_{POWERGEN}^{THg} \times Coef_{POWERGEN}^{1D} (7)$$

where  $ED_{ASGM}^{THg}$ ,  $ED_{INDS}^{THg}$ ,  $ED_{POWERGEN}^{THg}$  represent the direct release for varied sectors (unit: g/km<sup>2</sup>/yr);  $AMAP_{ASGM}^{THg}$ ,  $AMAP_{INDS}^{THg}$ ,  $AMAP_{POWERGEN}^{THg}$  represent the spatial distribution of release (unit: g/km<sup>2</sup>/yr); and  $Coef_{ASGM}^{D}$ ,  $Coef_{INDS}^{D}$ ,  $Coef_{POWERGEN}^{D}$  represent the coefficient of direct Hg emission (taken as unity).

506 
$$E_{total}^{THg} = EI_{ASGM}^{THg} + EI_{INDS}^{THg} + EI_{POWERGEN}^{THg} + ED_{ASGM}^{THg} + ED_{INDS}^{THg} + ED_{POWERGEN}^{THg}$$
(8)

where  $E_{total}^{THg}$  represents the total anthropogenic Hg release to rivers, which is eventually used to drive the MOSART-Hg-wrm model. The coefficients in equations (2) to (7) are adjustable to suit the requirements of various scenarios, generating a series of scenarios: *Low Emission, Moderate Emission, High Industrial Emission*, and *High ASGM Emission* with global emission ranging from 1,100 to 2,000 Mg/yr (Table S1).

We find a total anthropogenic emission of  $\sim$ 1,500 Mg/yr best matching the river Hg export to the ocean (*Moderate Emission*). Our estimations suggest that ASGM release account for 30-47% of the total anthropogenic Hg release to rivers (Table S1), which aligns with the assumption of the Kocman inventory that 40% of the release originate from ASGM. To better match the river Hg export dataset, we adopt the spatial pattern of AMAP inventory except for the North America, where the Streets Atmosphere inventory is used, as the former generates too high river Hg export in this region.

#### 517 Experiment Design

518 The primary experiment (the baseline scenario) conducted in this study employs the *Moderate Emission* 

with hybrid spatial distribution suggested by AMAP and Streets Atmosphere inventories. This experiment
facilitates a realistic representation of the present-day scenario by closely aligning with observational records.
A series of grouped experiments are devised to evaluate uncertainties stemming from different processes:
No Reservoir vs. With Reservoir (i.e., driven by *NHHR* and *BHHR*), Inventory (spatial distribution following
AMAP or Streets Atmosphere inventories), Soil Hg (soil Hg concentrations following Wang or Liu dataset),
and Deposition (no atmospheric deposition or not) (Table S3).

#### 525 Uncertainty Analysis

526 The MOSART-Hg-wrm model has uncertainties comparable to those in previous studies using the MOSART-Hg model (Peng et al. Under Review). We adopt a same model resolution as our previous study 527 528 focusing on the pre-industrial era (Peng et al. Under Review). We find that the model with a higher resolution 529  $(0.5^{\circ} \times 0.5^{\circ})$  performs better in simulating erosion flux in single topography units than a coarser resolution 530  $(1^{\circ}\times 1^{\circ})$ , however, the difference can be largely mitigated by using specific scaling factors. To balance computational cost and simulation accuracy, we use a resolution of  $0.9^{\circ} \times 1.25^{\circ}$ . In general, the accuracy of 531 532 the erosion processes and riverine sediment/Hg delivery processes are validated against observations in 533 previous studies (Tan et al. 2018, Li et al. 2022). For instance, the model has better performance in simulating global sediment yield compared to alternative models such as the RUSLE model, with a 59% lower 534 discrepancy relative to observations (Tan et al. 2018). Additionally, sediment fluxes in the MOSART-Hg 535 536 demonstrate similar performance to the BQART model (Peng et al. Under Review).

537 The water management component (reservoirs/dams) accounts for sediment and riverine Hg trapping 538 effects, alongside their influences on flow processes, which is supported by empirically validated 539 relationships between reservoir properties and river parameters (Vörösmarty et al. 2003) (see Method). The 540 theoretical basin trapping by this method agree well with the observed values, such as the Nile (observed 541 100% vs. theoretical 99%), Kizil Irmak (observed 98% vs. theoretical 95%) and Krishna River (observed 75% vs. theoretical 70%) (Vörösmarty et al. 2003). The water management module has also been 542 543 independently validated in previous studies (Zhou et al. 2020, Li et al. 2022). Notably, sediment flux estimations in MOSART-Hg-wrm exhibit high performance compared to previous estimations (Syvitski et 544 545 al. 2022) (Fig. S2). A similar level of model accuracy is expected for the Hg-trapping effects of reservoirs 546 and dams.

547

Another source of uncertainties is the spatial distribution of anthropogenic Hg sources to the freshwater

548 environments. We consider two different atmospheric Hg inventories, AMAP inventory (Steenhuisen et al. 549 2019, Steenhuisen et al. 2022) and Streets inventory (Streets et al. 2019), as a proxy for the spatial 550 distribution of anthropogenic releases to rivers. The riverine Hg budgets from both scenarios (AMAP scenario and Streets scenario hereafter) exhibit a strong correlation with observations at a global scale (Fig. 551 S3). However, the different spatial distributions of Hg sources results in varied performances in different 552 regions. The Streets scenario shows better performance in simulating riverine Hg in North America, while 553 the AMAP scenario performs better in other regions, particularly due to the more detailed sector contributions 554 555 for the latter.

The soil Hg concentration datasets play a crucial role in the Hg erosion processes. To address the 556 uncertainties stemming from these processes, we conduct the Soil Hg Experiment involving different soil 557 Hg concentration datasets (referred to as Liu dataset by Liu et al. (2023) and Wang dataset by Wang et al. 558 (2019)) (see Methods). The diverse datasets exert limited impacts on global-scale riverine Hg budgets, with 559 560 the Liu dataset resulting in only a slight increase in riverine Hg export to the ocean (<10%) than using Wang 561 dataset. Furthermore, Deposition Experiment is devised to assess uncertainties arising from the change of 562 soil Hg concentrations due to atmospheric Hg deposition and the loss of soil Hg by erosion (see Table S3). 563 In this experiment, soil Hg concentrations are dynamically adjusted based on the atmospheric deposition and erosion processes. However, the disparities between the scenarios with constant and dynamic soil Hg 564 concentrations are minimal, with changes in riverine Hg export to the ocean < 2% (Fig. S4). Similarly, given 565 that the rates of gaseous Hg evaporating from rivers are substantially lower than the atmospheric Hg 566 567 deposition rates (da Silva et al. 2006), the combined uncertainties from gaseous Hg evaporating and 568 deposition are expected to remain within 2%.

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## **Appendix** Mercury budget in global rivers at presentday: impacts from reservoirs and dams

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## Figures



Figure S1 Seasonal fluxes of riverine mercury discharge over land

Figure S2 The sediment load of estimation versus other estimations (Syvitski et al. 2022). The 1950/2010 sediment load is referred to as the estimation by Syvitski et al. (2022).



Figure S3 The comparison of Model results by Inventory Experiment with observation, AMAP represents the AMAP inventory, the Streets represent the Streets inventory (see Table S3).



**Figure S4** The comparison of Model results by *Soil Hg Experiment* with observation, Wang represents the soil Hg concentration of Wang et al. (2019), the Chen represent the soil Hg concentration dataset of Liu et al. (2023) (see Table S3).



Figure S5 The comparison of Model results by *Deposition Experiment*, the Deposition represents consider the setting of Deposition in Table S3, the Without Deposition represents the setting of Without Deposition in Table S3.



## Tables

	Anthronogonia		Indirect Release (Mg/yr)			Direct Release (Mg/yr)		
Scenarios	Ha budget (Ma/yr)	Changes (Mg/yr)	ASCM	Industrial		ASCM	Industrial	
	ing budget (mg/yi)		ASGM	INDS	NDS POWERGEN		INDS	POWERGEN
No Release	0	= 0	0	0	0	0	0	0
Low Release	1100	= No Release + 1100	142	73	8	378	460	39
Moderate Release	1500	= Low Release $+$ 400	232	84	6	378	748	59
Revised AMAP Release	About 1500			$\approx$ Modera	ate Release			
High Industrial Release	2000	= Moderate Release + 500	232	343	25	378	949	74
High ASGM Release	2000	= Moderate Release + 500	367	132	9	475	949	74

#### Table S1 The sets of industrial Releases and ASGM Releases.

Scenarios	Anthropogenic Hg Releases (Mg/yr)	Inputs of Riverine Hg (Mg/yr)	Outputs of Riverine Hg (Mg/yr)	Increased Release* (Mg/yr)	Increased Outputs (Mg/yr)	Increased Export/ Increased Releases (%)
No	0	418	248	0	-	-
Low	1100	1498	777	1100	529	48%
Moderate	1500	2005	1007	400	230	58%
High Industrial	2000	2532	1146	500	139	28%
High ASGM	2000	2436	1275	500	268	54%

Table S2 Fates of Riverine mercury under different scenarios.

Caption: \* the increased Release as compared with the previous one, such as 'moderate - low = 400', 'high -moderate = 500'.

Table 55 Experiments Design	Table S	5 Ex	periments	Design
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Experiment ID	Experiment Name	Soil Hg Concentration Dataset	Anthropogenic Hg Release (Mg/yr)	Soil Hg Dynamic	Atmospheric Hg Deposition	Water Management	Spatial Distribution of Release	Release Scenario
Baseline			About 1500					Revised AMAP
Anthropogenic Release Uncertainty			0	-		Yes	AMAP2015	No
			1100					Low
			1500					Moderate
			2000					High ASGM
		Wang	2000					High Industrial
	No Release without wrm		0	Constant	No	No	-	No
Water Management	No Release with wrm					Yes		
	Baseline without wrm		- 1500			No		Revised AMAP
	Baseline with wrm					Yes		
I	AMAP					Yes	AMAP 2015	-
Inventory	Streets						Streets	
Soil Hg	Wang	Wang					AMAP 2015	Modorato
	Chen	Chen						Moderate
Donosition	Deposition	Wang		Dynamic	Yes			
Deposition	Without Deposition	Wang		Constant	No			

Caption: Wang represents the soil Hg dataset from Wang et al. (2019), Chen represents the soil Hg concentration from Liu et al. (2023), AMAP2015 represents the inventory by Steenhuisen et al. (2019), Steenhuisen et al. (2022), Streets represents the inventory of Streets et al. (2019).

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