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Space-time data-driven modeling of wildfire initiation in the mountainous region of Trentino–South Tyrol, Italy

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Abstract

Wildfires are complex hazards occurring worldwide, leading to substantial economic 6 losses, fatalities, and carbon emissions. The interplay of climate change, land use alter-7 ations, and socioeconomic pressures is expected to further increase the frequency and 8 intensity of wildfires. In this context, developing reliable, dynamic prediction tools is es-9 sential for risk mitigation. This work presents a spatiotemporal wildfire prediction model 10 for the Trentino-South Tyrol region (13,600 km²) in the northeastern Italian Alps. Leverag-11 ing generalized additive models, we integrate multitemporal wildfire records (2000–2023) 12 with static and dynamic environmental controls (e.g., topography, land cover, daily pre-13 cipitation, and temperature). The resulting model predictions change dynamically over 14 space and time in response to static features, seasonal trends, and evolving meteorolog-15 ical conditions. Model outputs were evaluated using established performance metrics, 16 enabling the derivation of dynamic spatial wildfire probability thresholds. These thresh-17 olds are illustrated for varying amounts of precipitation, temperature, and different com-18 binations of static factors. Validation through multiple perspectives yielded performance 19 scores generally exceeding 0.8, confirming the model strong generalization and transfer-20 ability. To demonstrate the practical application, the model was used to hindcast past 21 wildfire initiation between 1–15 July 2022–a period marked by elevated wildfire activity. 22 By integrating static and dynamic environmental controls, this research advances the spa-23 tiotemporal prediction of wildfires in complex alpine regions, supporting the develop-24 ment of early warning systems. 25

²⁶ **Keywords:** Early warning; Space-time modeling; GAMs; Dynamic wildfire forecasting;

27 Wildfire ignition

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28 1 Introduction

Wildfires are frequent hazards posing significant threats to society and the environment. 29 On average, an estimated 350 Mha burns globally each year (Giglio et al., 2013). The 30 economic toll of wildfires is severe, with global wildfire-related losses amounting to ap-31 proximately USD\$142 billion since 2000, as reported by the international disaster database 32 (EM-DAT; Jones et al., 2022). Some of the most devastating wildfires in recent history in-33 clude the events in Australia, the United States, and Chile (Filkov et al., 2020; Maranghides 34 et al., 2023; Guerrero et al., 2024). Beyond their immediate economic and societal con-35 sequences, wildfires are a major source of carbon emissions, releasing an estimated 2–3 36 billion tons of CO_2 annually, at times surpassing 50% of emissions from fossil fuel com-37 bustion (Jolly et al., 2015; van der Werf et al., 2017). Their impacts manifest across multiple 38 scales as locally, wildfires disrupt ecosystems, threaten human settlements, and damage 39 infrastructure (Driscoll et al., 2024); regionally, they accelerate soil erosion, reducing agri-40 cultural productivity (Keeley et al., 2017); and globally, their carbon emissions contribute 41 to climate change, potentially amplifying future wildfire risk (Zscheischler et al., 2018). 42 Given the severe impacts of wildfires, there is an increasing need to minimize their harm 43 and, as such, the ability to ensure operational preparedness. In this context, reliable wild-44 fire predictions and developing wildfire early warning systems (WEWS) are essential for 45 reducing wildfire impacts and improving response efforts. 46 Wildfires occur when an ignition source interacts with available fuel under suitable 47 weather conditions, with subsequent behavior often influenced by terrain. Like other nat-48 ural hazards, wildfires are governed by multiple controls across various spatiotemporal 49 scales. At the smallest scale, fire behavior is commonly described by the "fire triangle" 50 concept (Countryman, 1972; Pyne et al., 1996), identifying heat, fuel, and oxygen as the 51 three fundamental components necessary for combustion. At the scale of a single wildfire, 52 its occurrence and propagation are driven by a complex interplay of static and dynamic 53 environmental controls (Moritz et al., 2005). Static factors such as the topography, hydro-54 climatic predisposition, and vegetation (fuel type) can determine whether an area is prone 55 to wildfires or the "where" (Faivre et al., 2014). In contrast, the timing, or the "when", is 56 governed by dynamic controls such as air temperature, precipitation, fuel moisture, and 57 anthropogenic influences—including arson or accidental ignitions—which directly affect 58 ignition likelihood and fire spread (Bessie et al., 1995; Ganteaume et al., 2013). Therefore, 59 assessing "where" and "when" wildfires may occur requires a comprehensive evaluation 60 of static and dynamic factors influencing wildfire behavior (Finney, 2005; Koutsias et al., 61 2015). 62 In this context, data-driven models are widely used to assess wildfire occurrence over 63 large areas (Martínez et al., 2009; Cao et al., 2013; Ghorbanzadeh et al., 2019; Tonini et al., 64 2020; Bjånes et al., 2021; Trucchia et al., 2022b; Yue et al., 2023). Despite its widespread use 65

⁶⁶ in fire prediction, the concept of susceptibility has been defined in various ways (Verde et al., 2010; Cao et al., 2017; Leuenberger et al., 2018). In this contribution, we adopt a

⁶⁸ definition analogous to that used for landslides and in agreement with Leuenberger et al.

(2018), describing wildfire susceptibility as the likelihood of a fire occurring within a given
 spatial unit (e.g., grid cells, catchments, administrative units) based on local predispos-

⁷¹ ing factors. In other words, it predicts "where" wildfires will likely occur (Brabb, 1984;

Guzzetti et al., 2005). In a purely spatial context, data-driven models derive statistical re-72 lationships between historical records of wildfire occurrence, i.e., wildfire inventories, as 73 static environmental factors to estimate wildfire susceptibility (Arndt et al., 2013; Oliveira 74 et al., 2012; Trucchia et al., 2022a; Zhang et al., 2021). Placing the spotlight on the spa-75 tial domain, wildfire susceptibility assessments provide valuable insights to support fire 76 management, whether as standalone analyses or as part of broader risk assessment frame-77 works and spatial planning strategies (Oliveira et al., 2021; Jappiot et al., 2009; Chuvieco 78 et al., 2010; Sebastián-López et al., 2008). Nevertheless, these approaches often disregard 79 the temporal component, excluding the assessment of critical meteorological conditions 80 that influence wildfire occurrence. 81 The temporal component of wildfire prediction is usually assessed through dynamic 82 factors, primarily focusing on daily variations in meteorological conditions, which form 83 the foundation of traditional WEWS (Chuvieco et al., 2010) and serve as tools for defin-84 ing local fuel treatments, surveillance, and suppression activities (Ager et al., 2010; Scott 85 et al., 2012; Oliveira et al., 2021). Numerous largely empirical wildfire danger indices, 86 also called fire danger rating systems, are used worldwide. A comprehensive literature 87 review of these index-based rating systems is beyond the scope of this contribution and 88 can be found in Viegas et al. (1999); Chuvieco et al. (2003); Arpaci et al. (2013); Giuseppe 89 et al. (2016); Pérez-Sánchez et al. (2017); Sirca et al. (2018); Pagnon Eriksson et al. (2023). 90 Wildfire danger indices empirically quantify the relative wildfire danger in a given area 91 based on meteorological conditions. Notable examples include the Canadian fire weather 92 index (FWI; Wagner et al., 1987), which relies on meteorological input data—wind speed, 93 air temperature, relative humidity, and precipitation—to assess daily wildfire danger. Al-94 though initially developed for pine fuel types prevalent in Canadian forests, the FWI has 95 been extensively tested worldwide (Kudláčková et al., 2024). Similarly, the McArthur's 96 forest danger index (FDI; Noble et al., 1980; Hollis et al., 2024) was developed to evaluate 97 wildfire danger and behavior in Australian eucalyptus forests, incorporating fuel avail-98 ability implicitly through atmospheric conditions. Another example is the Italian index 99 IREPI (Bovio et al., 1984), designed to assess wildfire danger during the winter–spring 100 season in northwestern Italy, which accounts for soil water loss due to evapotranspiration 101 derived from the multiple meteorological input data. Despite the utility of this fire danger 102 system and its spatially explicit features, these rating systems often overlook factors such 103 as intrinsic terrain characteristics, seasonality, and human influences, both of which are 104 critical in wildfire ignition and propagation (Chuvieco et al., 2010). 105 The integration of the spatial and temporal components for wildfire prediction in large 106 areas remains challenging. However, existing studies underscore its considerable poten-

107 tial. For instance, de Santana et al. (2021) modeled the occurrence of wildfires over a 19-108 year period by integrating static and dynamic predictors aggregated on an annual scale, 109 generating independent susceptibility maps for each year. In Verde et al. (2010); Parente 110 et al. (2016); Oliveira et al. (2021), the wildfire hazard was estimated by combining the 111 static wildfire susceptibility with a probabilistic assessment of wildfire recurrence, based 112 on projected future burned areas within each mapping unit. Bergonse et al. (2021) in-113 tegrated wildfire susceptibility with a meteorological index representing spring condi-114 tions, enabling predictions for any given year before the critical fire season. Conversely, 115 Deng et al. (2023) used a deep learning approach to predict daily wildfire ignition scores 116

by integrating dynamic weather predictors—such as temperature, rainfall, humidity, and 117 wind—with static factors. Richards et al. (2023) leveraged a spatiotemporal interpretable 118 neural network to model both wildfire ignition and spread, incorporating both static and 119 dynamic factors at a monthly resolution over the period 2001–2020. Woolford et al. (2011) 120 employed generalized additive models (GAMs) to predict wildfire occurrence during the 121 active fire season over a 24-year period, integrating both static and dynamic factors. Their 122 study emphasized the high interpretability of the chosen approach and the influence of 123 anthropogenic predictors, such as population density and proximity to infrastructure. 124

In the Alpine context, wildfires pose a significant threat to mountain forests and their 125 protective function against other natural hazards such as rockfalls, avalanches, mudflows, 126 increased run-off, and soil erosion (Maringer et al., 2016; Martin et al., 2001; Robichaud 127 et al., 2006). The region is characterized by two main fire seasons: an early spring peak in 128 March or April, driven by frost-induced dryness, and a summer peak in July and August 129 due to high temperatures (Müller et al., 2020a). To support wildfire assessment and forest 130 protection, multiple WEWS have been developed across different Alpine countries. For in-131 stance, the European forest fire information system (EFFIS; Giuseppe et al., 2016) provides 132 large-scale fire danger assessments at a spatial resolution of 8×8 km, but its applicability 133 in complex Alpine terrain is limited. National-level models, such as the one in Austria, 134 achieve resolutions of 1×1 km using meteorological inputs from weather forecasting sys-135 tems like INCA (Haiden et al., 2011), although a more recent prototype was developed, 136 achieving a spatial resolution of 100×100 m (Müller et al., 2020b). Despite these efforts, 137 wildfires in the Alps remain a significant challenge. In Italy, the average Alpine wildfire 138 covers about one hectare, but 10% of the fires exceed 10 ha, accounting for 85% of the total 139 burned area (Müller et al., 2020a). Human activities are the primary cause of ignitions, 140 responsible for about 90% of ignitions, though lightning-induced fires can contribute up 141 to 50% of ignitions in some areas during the summer season (Müller et al., 2013). Ad-142 vancements in WEWS have brought improvements, yet the spatial resolution of existing 143 models might be insufficient to capture the complex interactions between wildfire behav-144 ior and the heterogeneous Alpine topography, including narrow valleys, mountain peaks, 145 high-altitude plateaus, and their effects on weather controls (Carrega, 1995; Schunk et al., 146 2013; Müller et al., 2020b). As wildfire frequency and severity are expected to rise, there 147 is a growing need for high-resolution fire prediction models that integrate both static and 148 dynamic factors to enhance preparedness and mitigation strategies (Chuvieco et al., 2010; 149 Wastl et al., 2012). 150

This study investigates space-time wildfire initiation modeling by adapting a previ-151 ously developed framework devised for the assessment of mass movement hazards in an 152 Alpine region of northern Italy (Steger et al., 2023; Moreno et al., 2024; Steger et al., 2024; 153 Moreno et al., 2025). Following a similar modeling approach, we aim to integrate static 154 and dynamic wildfire controls using data-driven techniques to estimate wildfire initiation 155 probabilities in space and time. Furthermore, we aim to enhance the applicability of the 156 model by linking the predicted probabilities to quantitative thresholds. These analyses 157 are performed for the region of Trentino–South Tyrol (Italy), covering a period of 24 years 158 (2000-2023).159

The remainder of this paper is structured as follows: Section 2 describes the study area and its environmental characteristics. Section 3 introduces the wildfire historical records and the different static and dynamic factors used in the modeling. Section 4 details the methods, including the sampling procedure, data extraction, modeling with GAMs, validation strategies, and the thresholding approach. Section 5 presents the key results, emphasizing data sampling, model interpretation, and application. Ultimately, Sections 6 and 7 discuss the findings and conclude with an outlook on future research directions.

¹⁶⁷ 2 Study area

Our study area (shown in Figure 1) corresponds to the upper Adige River Basin, encom-168 passing the region of Trentino-South Tyrol in the northeastern Italian Alps, covering an 169 area of about 13,600 km². Its diversified landscape and climatic conditions characterize 170 the region. The mountainous topography is dominated by pronounced variations in eleva-171 tion, ranging between ~ 65 m a.s.l. (above sea level) in the southernmost part of the region, 172 close to Lake Garda to ~3900 m a.s.l. at the Ortles peak in the northwest. Due to phys-173 iographical features, Trentino-South Tyrol is highly prone to various mountain-related 174 natural hazards, including avalanches, rockfalls, and mudflows. Due to the proximity of 175 human settlements to the steep Alpine terrain, a significant portion of the infrastructure 176 is exposed to these processes. 177

The interaction of humid air masses from the northwest Atlantic, warm Mediterranean 178 influences, and dry continental air from the east shapes the regional climate. This intersec-170 tion gives rise to a pronounced seasonal cycle characterized by warm and humid summers 180 and cold and dry winters (Norbiato et al., 2009; Adler et al., 2015). The complex interplay 181 of topography and diverse atmospheric conditions of the region determines the strong 182 seasonality and spatial variations in precipitation and temperature. The annual precipita-183 tion sum varies significantly across the region, ranging from ~500 mm in the northwestern 184 rain-shaded inner valleys to over 1,700 mm in the southeastern parts, maintaining a spa-185 tial average over the study area of about 1000 mm. Similarly, mean annual temperature 186 varies substantially, from approximately +14°C in the Garda Valley to around -11°C at 187 the Ortles peak, with an overall regional average of about $+5^{\circ}$ C. The annual cycles of pre-188 cipitation and temperature are characterized by the warmest and wettest conditions dur-189 ing summer (July–August) and the coldest and driest conditions towards the late winter 190 (January–February). The most significant warming and precipitation increase occurs to-191 wards spring (April–May; Crespi et al., 2021). 192

The land cover in the region is predominantly forest (~45%), primarily consisting of coniferous trees located on hillsides. Herbaceous vegetation and heathland account for ~35%, while cultivated land, mainly concentrated in the valley bottoms, represents ~8%. The remaining land cover includes mostly rocky outcrops, water bodies, permanent snowcovered surfaces, and urban areas.

198 **3 Data**

¹⁹⁹ **3.1 Wildfire database**

²⁰⁰ The wildfire inventory for the region was obtained from the Forestry Service Department

- 201 (Abteilung Forstdienst/Ripartizione Servizio Forestale) of South Tyrol and the Forestry
- ²⁰² and Wildlife Service of Trentino (Servizio Foreste e Fauna). The databases were accessed
- in early 2024; the former was directly provided by the office in charge, and the latter can
- ²⁰⁴ be accessed at https://siat.provincia.tn.it/geonetwork/srv/ita/catalog.



Figure 1: Study area showing the spatial distribution of wildfire centroids (n = 998) across the Trentino–South Tyrol region, with symbol size and color representing the respective burned area for each event. Basemap imagery courtesy of Earthstar Geographics.

The wildfire inventory of South Tyrol is divided into two periods: records from 1999 to 2017 and those from 2017 onward. This distinction reflects changes in mapping techniques, as polygonal representations of burned areas were introduced only after 2017. Before this, fire locations were recorded as points, presumably indicating a location within the affected area. The dataset includes 608 wildfire events with attributes such as date and time of occurrence, duration, total burned area, and causes.

Similarly, the wildfire inventory of Trentino is structured into two periods: records 211 from 1966 to 1983 and those from 1984 onward. The earlier dataset comprises 1,062 events, 212 where only the burned area polygons were documented, without associated timestamps 213 and other relevant attributes. From 1984 onward, the dataset includes 3,134 wildfire 214 events, with temporal resolution improving over time. While timestamps were recorded 215 at the yearly level until 2007, only 507 events after 2007 include precise day-of-occurrence 216 information, leaving 2,627 with year-only time stamps. The recorded attributes in this 217 dataset include the occurrence time (with varying temporal resolution), total burned area, 218 and forested burned area. 219

Since the main objective of this study is to predict wildfire occurrence in space and time, the analyses focus on wildfire events for which occurrence dates are available. Only these events were extracted from the inventories to ensure the desired temporal resolution in the modeling. Further details on this methodical step are provided in Section 4.1.

224 **3.2** Environmental factors

225 Static factors

Identifying wildfire-prone areas using data-driven approaches relies on analyzing spatial
environmental factors observed at locations with and without wildfire occurrences. The
choice of predictors was guided by the numerous contributions that have explored the
influence of different static factors and their role in wildfire modeling (Ganteaume et al.,
2013; Jain et al., 2020). For this case study, we prioritized predictors that offer meaningful
insights into wildfire occurrence.

We used the NASADEM (NASA JPL, 2020) to derive standard morphological factors 232 such as *slope steepness*, which may facilitate fire spread (Chuvieco et al., 2010), and *as*-233 *pect*—decomposed into the components *eastness* and *northness*—to capture variations in 234 solar exposure and moisture retention (Müller et al., 2020b). A proxy for vegetation and 235 fuel types was represented using data extracted from the CORINE land cover and high-236 resolution map of Europe 2017 (Malinowski et al., 2020). The data were subsequently 237 reclassified into eight categories: (i) artificial surfaces, (ii) agricultural land, (iii) broadleaf 238 tree cover, (iv) coniferous tree cover, (v) herbaceous vegetation and heathland, (vi) marshes and 230 peatbogs, (vii) natural material surfaces, and (viii) permanent snow-covered surfaces and water 240 bodies. To further characterize vegetation within forested areas and measure forest struc-241 ture, we utilized the *tree cover density* dataset extracted from European Union's Copernicus 242 Land Monitoring Service information (2020). Climatic predisposition predictors, includ-243 ing total annual precipitation and mean annual temperature, averaged over a 30-year period 244 from 1981 to 2010, were also tested to characterize the regional climate and its influence on 245 wildfire occurrence (Crespi et al., 2020). Finally, to account for the anthropogenic compo-246 nent and infrastructure, buildings and roads were extracted from OpenStreetMap (OSM; 247

OpenStreetMap contributors, 2017). The Euclidean *distance to buildings* and *roads* were
 then computed to incorporate their effects on wildfire modeling. All the environmental
 factors were resampled to a 50 m spatial resolution for modeling purposes.

²⁵¹ Dynamic factors–gridded precipitation and temperature fields

We used the high-resolution gridded daily temperature and precipitation dataset for 252 Trentino–South Tyrol provided in Crespi et al. (2021). The dataset was built upon a dense 253 network of over 200 meteorological stations, ensuring comprehensive coverage across the 254 study area. Initially, the data processing involved multiple methodical steps, including 255 quality and consistency checks, homogeneity tests, and gap-filling techniques to maxi-256 mize data completeness. Using the processed database, an anomaly-based interpolation 257 scheme was applied to project the daily fields of precipitation and mean temperature onto 258 a grid of 250 m resolution. The accuracy, estimated via leave-one-out cross-validation, 259 yielded a mean absolute error (MAE) of 1.1 mm for precipitation and 1.5°C for mean tem-260 perature, averaged across all stations and months. For each day, the dataset provides the 261 total precipitation accumulated over 24 hours, from 08:00 UTC of the previous day to 08:00 262 UTC of the observation day, and the daily mean temperature is defined as the average of 263 the daily maximum and minimum temperature. 264

²⁶⁵ 4 Method

The method implemented in this research extends a data-driven framework initially de-266 vised for the assessment of mass movement hazards (Steger et al., 2023; Moreno et al., 2024; 267 Steger et al., 2024; Moreno et al., 2025), adapting it to integrate static and dynamic wildfire 268 controls for estimating wildfire occurrence probabilities in space and time. The workflow 260 comprises five main steps: (i) filtering and sampling rulesets. Since our model is based on 270 binary data, we first need to define our target variable: the areas with wildfire occurrence 271 or wildfire presences and those without or wildfire absences. First, we defined the wild-272 fire presences by combining the different wildfire databases and filtering the individual 273 records based on the inventory attributes. This is followed by the definition of wildfire ab-274 sences, where we applied a strategic sampling procedure to define wildfire absences, con-275 sidering both spatial and temporal aspects. The subsequent (*ii*) data extraction-precipitation 276 and temperature time windows phase retrieves static and dynamic predictors, with dynamic 277 predictors grouped into time windows. In (*iii*) modeling and grid search, we implement a 278 binomial GAM to perform a grid search and select the optimal combination of cumulative 279 precipitation and temperature to describe wildfire occurrence, which is then integrated 280 with static predictors to fit the final spatiotemporal model. (*iv*) Model evaluation follows, 281 incorporating plausibility checks and cross-validation routines from multiple perspectives 282 to assess predictive performance. Finally, the predictions are linked to quantitative met-283 rics, facilitating (v) visualization and thresholding, with a hindcasting example provided to 284 demonstrate model applicability. 285

286 4.1 Filtering and sampling rulesets

287 Wildfire presences

The four wildfire inventories—two from South Tyrol and two from Trentino—were pre-288 processed before being integrated into a single dataset. For South Tyrol, we used only the 289 point-based wildfire records. These records initially spanned the period from 1999 to 2017 290 but remained consistently available through 2023. Although polygon-based records were 291 available after 2017, they were excluded while point-based records were retained to ensure 292 spatial consistency, as comparisons between points and polygons did not always follow a 293 clear pattern. Since our analysis period starts in the year 2000, any events occurring before 294 this year were removed, resulting in the filtered wildfire dataset for South Tyrol. 295

Since our model aims to predict wildfire occurrence on a daily scale, only wildfire 296 records with an associated daily timestamp were considered. In the case of Trentino, this 297 requirement led to the selection of the inventory covering the period from 1984 onward, 298 while records from earlier periods, without any timestamp, were excluded. Addition-299 ally, wildfire events lacking a recorded *day of occurrence* (i.e., those before 2007) were re-300 moved. To ensure a uniform spatial representation across the inventories in South Tyrol 301 and Trentino, wildfire occurrences in Trentino were represented as point centroids of their 302 respective burned area polygons. 303

The two inventories were merged into a single dataset, forming the preliminary wildfire sample. Events with a *burned area* smaller than 10 m² were filtered out to ensure data reliability, resulting in a final dataset for the subsequent analysis. Finally, for data extraction, we consider a buffer of 50 m around each point to enhance the spatial representativeness of each wildfire point.

309 Wildfire absences

Selecting wildfire absence data is just as crucial as the wildfire presence filtering. However, 310 defining the absences poses a greater challenge, as it requires identifying not only loca-311 tion but time periods where wildfires are unlikely to have occurred. The sampling strategy 312 is essential in shaping the model structure and is directly reflected in the modeling out-313 comes. Therefore, ensuring a well-balanced representation across spatial and temporal 314 domains is essential to prevent wrongly estimated model relationships and potentially 315 biased predictions. To this aim, we considered several rules, including (i) masking out 316 trivial terrains, (ii) implementing temporal buffers, (iii) balancing samples across years 317 and months, and (iv) masking out trivial periods that are detailed below. 318

Trivial terrains are locations where wildfire initiation is not expected. In our analysis, we defined these areas based on the available fuel that can be accounted for using the land cover map. In this way, the classes referring to *artificial surfaces*, *natural material surfaces* (*e.g.*, *bare rock*, *sand*), and *permanent snow-covered areas and water bodies* were excluded from the spatial domain as there is no fuel to be potentially burned. This resulted in a mask being used to sample space wildfire absence locations. Notably, no recorded wildfire events were located within trivial terrains.

Wildfire absences were sampled under two different schemes, including the sampling of wildfire absences at presence locations and the sampling of wildfire absences at absence locations. For the first scheme, wildfire absences at presence locations were generated by

replicating records from the filtered wildfire inventory and assigning a random observa-329 tion date within the period of analysis, i.e., 01-01-2000 to 31-12-2023. To ensure that the 330 selected dates were temporally meaningful, a temporal buffer was applied to avoid select-331 ing dates within five years of the respective wildfire occurrence. This step aimed to allow 332 sufficient time for vegetation (fuel) recovery before it may be subject to wildfire occur-333 rence, as supported in Sass et al. (2017), Bowd et al. (2021), and Mantero et al. (2023). A 334 30-day temporal buffer was also introduced to avoid overlapping meteorological informa-335 tion when extracting the dynamic predictors, as explained in Section 4.2. 336

For the second scheme, wildfire absences at absence locations were generated by randomly sampling absence locations, excluding the trivial terrains and the areas that had been burned before, according to the filtered inventory. Once absence locations were identified, they were replicated and assigned a random observation date for the specified analysis period (01-01-2000 to 31-12-2023). Similar to the first scheme, the date selection was constrained by a 30-day temporal buffer to avoid overlapping meteorological data.

The preliminary wildfire absence subset was formed by combining the wildfire absences at presence locations and those at absence locations. Subsequently, the absence subset was further refined by balancing the observation dates across the different years and months to achieve a uniform temporal representation. The results of this step were combined with the wildfire presences to obtain the modeling sample, which was subjected to a final filtering step to account for trivial periods, as described in Section 4.2.

³⁴⁹ 4.2 Data extraction–precipitation and temperature time windows

The static predictors were extracted to the modeling sample using the datasets described in 350 Section 3. Additionally, predictors regarding the occurrence date, specifically year, month, 351 and day of the year (doy) were derived from the assigned observed date in the modeling 352 sample. Data extraction was performed for the meteorological predictors such as precip-353 itation and temperature by linking each sample location and date to the high-resolution 354 gridded dataset from Crespi et al. (2021). Both predictors were retrieved up to 30 days 355 before the assigned observation date. Similar to trivial terrains, we defined trivial periods 356 based on a precipitation threshold so that observations with precipitation amounts ≥ 1.1 357 mm on the day of observation were excluded from the analysis. This filtering step was 358 necessary because including large amounts of "wet" days-presumably with no wildfire 359 occurrence—into the binary classification model would result in the model simply learn-360 ing the difference between "dry" and "wet" conditions. Hence, we applied the precipi-361 tation threshold to explicitly keep the modeling target to predict wildfire occurrence in 362 "dry" periods. 363

Finally, the precipitation and temperature predictors were structured into multiple 364 time windows for analysis. The predictor *P* represents the cumulative precipitation within 365 a given time window, while *Te* corresponds to the mean daily temperature over the same 366 period. As an analogy to what was described in Steger et al. (2023), P aims to represent 367 the medium-term preparatory precipitation, which can be seen as a proxy for the surface 368 moisture conditions, whereas Te can be seen as the short-term temperature. This approach 369 enabled a systematic evaluation of how the combination of different antecedent precipi-370 tation and temperature influences wildfire occurrence. 371

372 4.3 Modeling using GAMs and grid search

373 Theoretical background

Generalized additive models are semi-parametric regression models that extend generalized linear models (GLMs) by incorporating additive smooth functions to capture nonlinear relationships between predictor variables and the response (Wood, 2006; Bolker et al., 2009; Zuur et al., 2009). In this study, the binary response variable, representing wildfire presence or absence, follows a Bernoulli probability distribution and can be generically expressed as:

$$logit(p_i) = \beta_0 + f_1(x_{1i}) + f_2(x_{2i}) + \dots + f_n(x_{ni})$$
(1)

³⁸⁰ Where f_1 , f_2 and f_n are smooth functions to model the nonlinear effects on each of the ³⁸¹ predictors x_{1i} , x_{2i} , and x_{3i} . A logistic link function is applied to relate p_i to the effects of ³⁸² the predictor variables. Due to their high interpretability and flexibility—often described ³⁸³ in the literature as *white-box models*—GAMs have been extensively used across various ³⁸⁴ disciplines, including natural hazard modeling (Woolford et al., 2011; Pourtaghi et al., ³⁸⁵ 2016; Ríos-Pena et al., 2017; Lombardo et al., 2020; Eskandari et al., 2021; Opitz et al., 2022; ³⁸⁶ Moreno et al., 2023; Alvioli et al., 2024).

387 Grid search and model fit

In this contribution, GAMs were applied in two distinct phases: (i) to identify the optimal time windows P and Te that best describe wildfire initiation, and (ii) to integrate static factors with the identified best P and Te windows to estimate wildfire probabilities in space and time. The analyses were conducted using the comprehensive R package mgcv (version 1.9-1; Wood, 2017).

The selection of optimal time windows for *P* and *Te* was conducted through a system-393 atic grid search. This procedure involved a pairwise comparison of different P and Te time 394 window combinations using a binomial GAM. This model was fitted to the previously de-395 fined modeling sample, as explained in 4.1, incorporating iteratively the different combi-396 nations of *P* and *Te*, and the *doy* as predictors. The grid search evaluated 30 time windows 397 for *P* and 10 for *Te*, resulting in 300 unique combinations. Each combination underwent a 398 10-fold random cross-validation with 10 repetitions to ensure robustness, yielding 300,000 399 iterations. The performance of the model was recorded for each iteration—based on well-400 known performance metrics addressed in Section 4.4—allowing for a systematic compar-401 ison to identify the optimal pair of time windows that best describe wildfire initiation. 402

The selected optimal time windows *P* and *Te* were integrated with the static predictors 403 to develop the spatiotemporal wildfire initiation model. The model fitting process was it-404 erative, starting with a simple baseline configuration and progressively incorporating ad-405 ditional predictors. Each newly introduced predictor was assessed by visually inspecting 406 partial effect plots and formally inspecting p-values, confidence intervals, and the overall 407 model fitting performance. Once the final model fit was obtained, the relative contribution 408 of each predictor was evaluated using a permutation-based variable importance analysis. 409 This assessment indicates how much each predictor contributes to the model by randomly 410 permuting its values and measuring the resulting change in the model prediction error. 411

Specifically, the analysis was based on the deviance explained—a widely used measure 412 of goodness of fit derived from the likelihood function. The importance of each predictor 413 was quantified by comparing the deviance explained by the final model fit against a series 414 of models, each with a single predictor permuted. These permutations disrupt the origi-415 nal relationship between the predictor and the response; therefore, a greater reduction in 416 the deviance explained indicates a more substantial contribution of each permuted predic-417 tor (Schlögl et al., 2025). In this study, permutation feature importance was summarized 418 using the mean and standard deviation to quantify both the average effect and the associ-419 ated uncertainty. For the interpretation and plausibility of the model relationships, partial 420 effect plots were generated, illustrating how individual predictors influence wildfire oc-421 currence probabilities while holding all other predictors constant at their average values. 422 Additionally, interaction effects between predictors were visualized using contour plots, 423 providing further insights into the complex relationships governing wildfire dynamics. 424

425 **4.4 Model evaluation**

For a comprehensive evaluation of the model performance, we leveraged multiple validation strategies using the R package *sperrorest* (version 3.0.5; Brenning, 2012). The assessment included k-fold random cross-validation (RCV), k-fold spatial cross-validation (SCV), temporal cross-validation (TCV) based on years and months, and leave-one-factorout cross-validation (FCV) using the land cover data. These approaches ensured a robust model generalization and transferability assessment across different spatial, temporal, and environmental conditions.

Random cross-validation was implemented by partitioning the dataset into indepen-433 dent training and testing sets. We applied a ten-fold RCV with ten repetitions, result-434 ing in 100 iterations (Brenning, 2012). While RCV provides a general measure of model 435 predictive performance, it may overlook spatial heterogeneity, potentially leading to over-436 optimistic results if the model performs inconsistently across different subregions. To ad-437 dress this, we incorporated SCV, which estimates the spatial transferability of the model 438 by evaluating its performance in spatially distinct partitions. The spatial partitioning was 439 achieved using a k-means clustering approach, ensuring alignment with the ten-fold, ten-440 repetition setup used in RCV. 441

To further assess model generalization and transferability, we applied TCV and FCV. 442 The former was conducted by iteratively excluding wildfire observations from either 443 one month (leave-one-month-out) or one year (leave-one-year-out) and then evaluating 444 model predictions on the omitted data. This approach allowed us to test how well the 445 model performs when transferred to unseen time periods at the scale of years and months. 446 Similarly, FCV was performed using the five nontrivial land cover classes—agricultural 447 land, broadleaf tree cover, coniferous tree cover, herbaceous vegetation, and heathland, and marshes 448 and peatbogs— to define independent training and testing datasets, ensuring that the model 449 performance was not overly dependent on specific land cover conditions. 450

The model performance for each validation approach was quantified using the area under the receiver operating characteristic curve (AUROC) computed for independent testing sets. The ROC curve represents the classifier performance across varying cut-off thresholds, with AUROCS values ranging from 0.5, i.e., random classification to 1, i.e., perfect discrimination (Hosmer et al., 2013). Higher AUROC values indicated stronger
 performance, providing an objective measure to compare different validation strategies
 and assess the model robustness across spatial, temporal, and land cover variations.

458 4.5 Thresholding and visualization

In this research, we adapted an extension of empirical rainfall thresholds—traditionally 459 employed in landslide early warning systems (Brunetti et al., 2010) — to wildfire predic-460 tion. Conventional landslides thresholds operate on the premise that slope failures be-461 come probable when specific rainfall parameters exceed critical values. Recent advance-462 ments in Steger et al. (2024) showed that spatial dynamic landslide probability thresholds 463 can be derived by incorporating both dynamic meteorological controls and static terrain 464 conditions. Building upon this methodological framework, we developed analogous spa-465 tially explicit thresholds for wildfire prediction, enabling the reclassification of dynamic 466 probability maps to directly visualize threshold exceedance patterns across varying spa-467 tiotemporal contexts. 468

In our wildfire analysis, we employ heatmap visualizations to represent wildfire proba-469 bility scores for increasing amounts of the meteorological predictors P and Te. To enhance 470 the practical application of this results, we established three distinct threshold derived 471 from the ROC curve, transforming the continuous probability scores (0–1) into meaning-472 ful discrete classes. These thresholds represent specific combinations between correctly 473 classified wildfires (true positive rate or sensitivity; TPR) and wrongly classified absences 474 (false alarms or false positive rate; *FPR*), thereby providing a more nuanced categoriza-475 tions of wildfire likelihood. The *TPR95* threshold establishes a high sensitivity benchmark, 476 ensuring that 95% of observed wildfires occur in areas and periods exceeding this prob-477 ability value. While this threshold effectively captures most wildfire-prone conditions, it 478 comes at the expense of generating a high proportion of false alarms. On the other hand, 479 the TPR35 threshold implements a more stringent criterion, identifying only 35% of wild-480 fire occurrences. This higher-probability threshold substantially reduces false alarms but 481 excludes 65% of actual wildfire events, capturing only the most extreme wildfire-prone 482 conditions. 483

Positioned between these two extremes, we implemented an optimal OPT threshold 484 based on the Youden index, which maximizes the sum of correctly classified wildfire and 485 non-wildfire locations and periods. This threshold corresponds to the point on the ROC 486 curve with maximum vertical distance from the diagonal line representing a random clas-487 sification (Fluss et al., 2005; Ruopp et al., 2008). The OPT threshold provides the most 488 balanced compromise between the classification of wildfire and non-wildfire conditions. 489 Utilizing these three thresholds, our final dynamic probability map delineates four dis-490 tinct wildfire likelihood classes, with *OPT* as the central reference point. 491

The resulting dynamic maps directly visualize wildfire threshold exceedance across the study region. To further evaluate the practical utility of the methodical approach, we conducted a comprehensive hindcasting analysis. This exercise allowed us to reconstruct and examine the spatiotemporal evolution of wildfire probability thresholds during the period of 1–15 July 2022, when several wildfire events affected the study area.

497 5 Results

⁴⁹⁸ 5.1 Wildfire inventory and modeling sample

In South Tyrol, the initial dataset contained 608 recorded wildfires. After applying the 499 temporal filter, eight events occurring before the year 2000 were excluded, resulting in a 500 final dataset of 600 wildfires. In Trentino, the initial inventories comprised 4,196 wildfires 501 in total. We did not use the first inventory—comprising 1,062 events from the period from 502 1966 to 1983—due to missing temporal information. From this subset of 3,164 wildfires, 503 we kept only those occurring within our period of interest (2000–2023), leaving 946 wild-504 fire events. Additionally, we filtered out the records without a daily timestamp, resulting 505 in a sample size of 507 wildfires. The two processed inventories were merged, yielding 506 a combined dataset of 1,107 wildfire records. These records met the criteria of having a 507 recorded occurrence date between 2000 and 2023 and an associated daily timestamp. Fi-508 nally, events with a burned area smaller than 50 m² were filtered out, resulting in a final 509 dataset of 998 wildfires, which we used in subsequent analysis. 510



Figure 2: Temporal distribution of wildfires (n = 998) in the study area between 2000 and 2023. Panel (a) shows the annual wildfire counts, distinguishing between events occurring in the warmer months (May–October) and colder months (November–April). Panels (b) and (c) display the monthly distribution of wildfires, with a distinction based on the respective burned area.

Although a detailed characterization of the wildfire regime is not the primary focus of this research, we find it valuable to illustrate some general patterns observed in the

inventories. The temporal distribution of wildfires across the study area reveals several 513 notable peaks, as shown in Figure 2a. The most prominent years were 2007 and 2022, each 514 registering over 100 wildfire events each year. Other years with elevated activity—each 515 exceeding 50 wildfires—include 2012, 2015, 2017, and 2019. Monthly trends, in Figures 516 2b–c, show two seasonal peaks in wildfire occurrence: one in early spring and another 517 in mid-to-late summer. While both South Tyrol and Trentino follow this bimodal pat-518 tern, South Tyrol tends to experience more wildfires during the summer peak, whereas 519 Trentino shows a more pronounced peak during the early spring. Regarding weekly pat-520 terns, approximately 30% of wildfires occurred during weekends (i.e., Saturday and Sun-521 day), while the remaining 70% took place on weekdays. Between 2000 and 2023, the study 522 area experienced an annual average of 41 wildfires, burning approximately 31 hectares 523 per year. Wildfires occurred throughout the year and were recurrent phenomena both in 524 the warmer months (May–October) and in the colder ones (November–April). Although 525 some years show a clear predominance of wildfires in one period over the other, the over-526 all distribution between warm and cold months remains relatively balanced. In terms of 527 wildfire size, the majority of wildfires were relatively small: 84 out of the 998 events ex-528 ceeded 1 ha, and only one surpassed 90 ha. Regarding wildfire causes, available only for 529 events in South Tyrol. Roughly 50% of the events have no reported cause. Among the iden-530 tified causes, lightning accounted for 23%, cigarette-related ignitions for 11%, while the re-531 maining causes included poorly maintained electrical infrastructure, agricultural burns, 532 recreational activities, and arson. 533



Monthly wildfire presence and absence distribution

Figure 3: Data sampling results. Bar plots show the monthly distribution of the sample data, wildfire presences (in red) and absence observations (in gray) before (a) and after (b) excluding trivial periods from the dataset.

⁵³⁴ The final modeling dataset was derived from a combination of wildfire presence and

absence samples, totaling 5,989 observations. This included 998 wildfire presences and 535 4,991 absence observations, respectively, maintaining an approximate 1:5 ratio. To avoid 536 an uneven temporal representation, absence observations were uniformly distributed 537 across years (2000–2023), with proportional representation across months, ensuring that 538 each day of the year had an equal chance of being selected (Figure 3a. At this stage, the 539 sample reflected only the spatial and temporal structure of wildfire presence and absence 540 samples without considering whether the respective observation experienced precipita-541 tion. We applied a precipitation-based threshold to account for trivial periods, specifically 542 removing observations with the precipitation on the day of observation (i.e., day 0) equal 543 to or exceeding 1 mm. This filtering step led to excluding 1,735 observations, reducing the 544 sample size to 4,254 records. The impact was more pronounced in the absence data, with 545 nearly 30% of those records excluded, dropping from 4,991 to 3,398. This shows that a sig-546 nificant portion of the absence samples occurred on "dry" periods, with approximately 547 70% of these experiencing no measurable precipitation (Figure 3b. 548

⁵⁴⁹ Wildfire presence data was also affected, with 142 events—around 15% of the initial ⁵⁵⁰ wildfire sample—removed. Most of these excluded wildfire records occurred during the ⁵⁵¹ summer months (June–August), indicating that a notable number of wildfires were asso-⁵⁵² ciated with precipitation ≥ 1 mm on the day of occurrence. This observation underscores ⁵⁵³ the complexity of wildfire-weather interactions during the wetter summer period, where ⁵⁵⁴ precipitation does not prevent ignition.

555 5.2 Optimal time windows and modeled relationships

The final modeling dataset was used in a repeated 10-fold RCV with 10 repetitions to identify the optimal time windows for representing the meteorological predictors P and *Te* in wildfire prediction (Figure 4).



Figure 4: Results of the grid search for optimal Te-P time window combinations. The plot shows the median AUROC values obtained for each combination of Te and P, using a 10-fold RCV with ten repetitions. The best-performing combination (yellow star) yielded a mean AUROC of 0.82 and corresponds to the mean temperature on the observation day (Te_0) combined with the cumulative precipitation over the 29 days preceding the event (P_{28}).

Among the tested combinations, the pairing of Te_0 and P_{28} led to the highest model 559 performance, achieving a median AUROC of 0.82. In this configuration, Te_0 denotes the 560 mean daily temperature on the observation date, while P_{28} corresponds to the cumulative 561 precipitation over the preceding 29-day period. Overall, model performance was highest 562 when short temperature windows (e.g., Te_0 to Te_4) were paired with medium precipitation 563 windows (e.g., P_{10} to P_{30}). In contrast, lower AUROCs were observed when longer tem-564 perature windows (e.g., Te_5 to Te_{10}) were combined with short or medium precipitation 565 windows (e.g., P_0 to P_{15}). Based on these results, Te_0 and P_{28} were selected as the optimal 566 dynamic predictors and were subsequently integrated with static predictors to develop 567 the final wildfire dynamic model. 568



Figure 5: Partial effect plots. Panels (a)–(h) display the partial effects of continuous predictors on the response scale, with the center white lines representing the mean estimated effects, and the red bands indicating the 95% confidence intervals. Panel (i) shows the effect of the categorical predictor *land cover* on the linear scale, where red dots represent the mean estimated effect and error bars indicate the 95% confidence intervals. Lastly, panel (j) illustrates the predictor *year*, which was modeled as a random effect, but excluded during the final prediction phase.

Figure 5 illustrates the partial dependence of selected predictors on wildfire probabil-569 ity, with the influence of all other predictors held constant. The dynamic predictors Te_0 and 570 P₂₈ indicated that the highest wildfire probabilities occurred under conditions of elevated 571 Te_0 and limited P_{28} . In contrast, wildfire probabilities were the lowest when P_{28} was abun-572 dant and Te_0 was relatively low. The *doy* revealed a seasonal trend, with increased wildfire 573 probabilities around *doy*₁₀₀, corresponding to early spring (March–April), and lower prob-574 abilities around doy_{250} , which aligns with the late summer (August–September). Among 575 topographic variables, Aspect showed a clear pattern: slopes facing south (135°–225°) were 576 associated with a higher likelihood of wildfire occurrence compared to other orientations. 577 In terms of vegetation structure, *tree cover density* revealed that moderate values (20–70%) 578 were linked to increased wildfire probability, whereas sparse (<20%) and dense (>70%) 579 tree cover were linked to reduced probabilities. The distance to buildings, used as a proxy for 580 anthropogenic influence, displayed a clear decreasing trend in wildfire probability with 581 increasing distance from built-up areas. The mean annual temperature, indicative of gen-582 erally warmer regions, exhibited a relatively linear relationship, where hotter areas were 583

related to a higher likelihood of wildfire occurrence. Likewise, total annual precipitation, 584 serving as a proxy for overall dryness, revealed a relatively linear trend with drier areas 585 being more prone to wildfires. The *land cover*, when using *agricultural land* as the reference 586 category, several classes—including deciduous tree cover, coniferous tree cover, herbaceous veg-587 etation and heathland, and marshes and peatbogs—were significantly associated with higher 588 wildfire probability. Finally, year was incorporated into the model fit as a random effect 589 to account for the interannual variability in the reported wildfire reporting. Then, it was 590 subsequently excluded from the predictions to ensure that inconsistencies did not sys-591 tematically influence the resulting wildfire probabilities in the temporal distribution of 592 reported events. 593

⁵⁹⁴ 5.3 Variable importance and model evaluation

Permutation variable importance analysis revealed positive contributions from all selected 595 predictors to the model deviance explained. The 29-day accumulated precipitation (P_{28}) 596 emerged as the predominant predictor, followed by temperature at the observation day 597 (Te_0) . Among the static environmental factors, tree cover density and land cover demon-598 strated relatively high variable importance, while the dynamic *doy* and topographic vari-599 able aspect exhibited moderate influence. The remaining predictors, while still contribut-600 ing positively to model deviance explained, displayed comparatively lower relative im-601 portance. 602



Figure 6: Permutation variable importance results. The red dots denote the mean portion of deviance explained, and the error bars denote the associated standard deviation.

The model exhibited robust predictive capability across all validation routines pre-603 sented in Figure 7. AUROC scores generally exceeded 0.80, corresponding to excellent dis-604 *crimination* as defined by Hosmer et al. (2013). The two 10-fold cross-validation strategies, 605 RCV and SCV, yielded median AUROC values of 0.83 and 0.81, respectively. As antici-606 pated, SCV—which employs k-means clustering to account for spatial autocorrelation— 607 resulted in a slightly lower performance and a broader interquartile range (IQR) than 608 RCV. 609 The leave-one-out cross-validation routines, namely TCV for years and months, and 610

FCV for land cover, as shown in Figures 7c–d, yielded mean AUROC values of 0.82, 0.79, and 0.79, respectively. The lower performance score across specific folds (years, months,

and land cover classes) likely reflects discrepancies between the local conditions driving 613 wildfire occurrence in those subsets and the patterns learned from the remaining data 614 used for the training. TCV at the annual level, in general, demonstrates high temporal 615 transferability, with comparatively lower performance observed in 2002, 2016, and 2018, 616 and higher AUROC values in 2008, 2012, and 2014. On a monthly scale, October exhibits 617 the lowest scores, while the winter months, December–February, yield the highest predic-618 tive performance. FCV indicates AUROC values consistently exceeding 0.75 across all land 619 cover classes except for the *agricultural land*, which showed a lower score of approximately 620 0.72. 621



Figure 7: Summary of the predictive model performance. Panel (a) presents the 10-fold RCV and 10-fold SCV results with ten repetitions. Panels (b–c) display the results of the TCV across individual years and months, while panel (d) shows the FCV across different lad cover classes.

5.4 Spatial dynamic thresholds and predictions

The *Te–P* heatmap in Figure 8 provides a visualization of the relationship between temperature and precipitation, and their association with wildfire probabilities. In this analysis, static factors and *doy* were excluded from the predictions. The heatmap reveals that the highest wildfire probabilities are predicted under conditions of elevated Te_0 combined with reduced P_{28} . This relationship is highlighted by the spatial distribution of the modeling samples—with their respective Te_0 and P_{28} values—within the heatmap. Wildfire occurrences (denoted by crosses) are predominantly clustered in regions of higher probability, while absence observations (denoted by points) concentrate in areas with lower probability scores.

The selected thresholds act as decision boundaries distinguishing between areas with 632 high and low predicted wildfire probabilities. In other words, predictions below a given 633 threshold indicate low wildfire probability, while those above suggest high wildfire prob-634 ability scores. For instance, 95% of the wildfire occurrences exceed the TPR95 (green dot-635 ted line) threshold, but a low proportion of correctly classified absences (TNR of 41%) fall 636 below this threshold, implying relatively high false alarms (FPR of 59%). Conversely, the 637 more conservative *TPR35* (blue dotted line) threshold minimizes false alarms (FPR of 5%) 638 by correctly classifying 95% of absences (TNR of 95%), but identifies only 35% of wildfire 639 occurrences. The OPT threshold (red solid line) provides a balanced trade-off, achiev-640 ing a TPR of 77% while maintaining an FPR of 25% (TNR of 75%). To put it in another 641 way, this threshold optimizes the identification of wildfire occurrences while keeping an 642 acceptable level of correct absence classification. 643



Figure 8: *Te–P* heatmap showing the predicted wildfire probability in relation to temperature and precipitation inputs. The plot illustrates the combined effect of *Te* and *P* on the predicted probability of wildfire occurrence, assuming average values for *doy* and all static predictors (i.e., *aspect, tree cover density, total annual precipitation, mean annual temperature, distance to buildings,* and *land cover*). Crosses represent the 859 wildfires, while points indicate the 3,398 absence observations. The curves correspond to the selected probability thresholds: *TPR95* (green), *OPT* (red) and *TPR35* (blue).

⁶⁴⁴ The heatmap and thresholds highlight the relationship between dynamic predictors,

prediction scores, and the link to the ROC curve. Furthermore, the probability thresholds
demonstrate a potential application of our model for early warning systems, analogous
to established landslide probability thresholds, and provide flexibility in implementation
based on specific wildfire management objectives. It is important to note, however, that
accurate visualization of these heatmaps also requires the integration of static factors.

- ⁶⁵⁰ Figure 9 illustrates how seasonal variations and changes in static environmental condi-
- tions influence the Te-P relationship with wildfire probability. Specifically, we performed a comparative analysis of the Te-P heatmaps across different combinations of *doy*, *aspect*,
- a comparative analysis of the *Ie–P* heatmaps across different combinations of *doy, asp* and *tree density cover*, with the remaining predictors being excluded.



Figure 9: *Te*–*P* heatmaps showing the predicted wildfire probability for increasing amounts of *Te* and *P* and associated *OPT* threshold for different combinations of *doy* (doy_{30} and doy_{240}), *aspect* (225° and 0°) and *tree density cover* (30% and 100%).

Panels (a, c) versus panels (b, d) demonstrate the influence of the seasonal variation through *doy*. It is observed that substantially lower short-term temperature Te_0 is required to exceed the *OPT* wildfire probability threshold on a winter day (doy_{30}) compared to a summer day (doy_{240}), where higher amounts of Te_0 are necessary to reach equivalent predicted probabilities. By contrasting Figure 9a–b versus Figure 9c–d, we investigate the role of topographic orientation and vegetation structure. The former set of panels represents ⁶⁶⁰ conditions with a south-facing *aspect* of 225° and a relatively low *tree cover density* of 30%, ⁶⁶¹ whereas the latter reflects a north-facing *aspect* of 0° and dense forest cover (100%). The ⁶⁶² heatmaps reveal that substantially higher Te_0 and lower P_{28} are necessary to achieve high ⁶⁶³ wildfire probabilities and exceed the *OPT* threshold under conditions of dense vegetation ⁶⁶⁴ and in north-facing slopes. Such analyses can be extended to any combination of predic-⁶⁶⁵ tors in the model and offer valuable insights into the spatiotemporal dynamics of wildfire ⁶⁶⁶ occurrence.



Figure 10: Hindcasting example of wildfire occurrence during the period between 1–15 July 2022. The illustrations display the spatial distribution of the predicted wildfire probability across the entire study area using the three selected probability thresholds: *TPR95*, *OPT*, and *TPR35*.

To further demonstrate the model potential applications, we conducted a hindcast analysis covering the period of 1–15 July 2022—a timeframe chracterzied by elevated wildfire activity across the study region. The resulting model predictions, initially expressed as continuous probabilities, were subsequently reclassified using the predefined thresholds

(TPR95, OPT, and TPR35). Figure 10 presents four representative snapshots at 5-day inter-671 vals from this period, illustrating the temporal evolution of predicted wildfire probability 672 thresholds. These sequential images reveal a pronounced increase in predicted wildfire 673 occurrence as the date approaches mid-July, with substantial expansion of areas exceed-674 ing the *OPT* and *TPR35* thresholds. The spatial patterns emerging from these predic-675 tions show notable regional variability, with specific zones—particularly the inner valley 676 bottoms—consistently exceeding both the mid and high probability thresholds through-677 out the analysis period. This spatial heterogeneity reflects differences in the underlying 678 environmental characteristics across the landscape, including topography, infrastructure, 679 vegetation, land cover, and meteorological conditions. A pronounced transition is ob-680 served between 10 and 15 July, during which several inner-valley areas surpass the high-681 est threshold, indicating high probability scores. The complete temporal overview of the 682 hindcast is available as an animated GIF (*Trentino–SouthTyrol.GIF*), offering a dynamic 683 visualization of how predicted wildfire probability thresholds evolved throughout this 684 two-week window. 685

686 6 Discussion

In this study, we implemented an interpretable spatiotemporal modeling framework that 687 integrates static environmental and dynamic meteorological factors to predict wildfire oc-688 currence. The model demonstrates strong predictive capabilities, with AUROC values 689 generally exceeding 0.80, highlighting its effectiveness in capturing key drivers of wild-690 fire initiation. These included static landscape features, daily temperature and precipita-691 tion fields, and a seasonal proxy. The following discussion addresses critical aspects of 692 the study, including reflections on model interpretability and flexibility, considerations 693 related to wildfire processes and drivers, and the applicability of the model in early warn-694 ing. 695

606 6.1 Model flexibility and interpretability

The sampling strategy played a key role in constructing the modeling dataset. As the 697 quality and representativeness of the input data heavily influenced the model outcomes, 698 careful design of the sampling process is as relevant as the modeling phase itself. While 699 many data-driven studies emphasize comparing algorithms, evaluating data preparation 700 and sampling strategies is often underappreciated. Generating the modeling sample in-701 volves filtering the wildfire presences and sampling the absence observations due to the 702 large disproportion between the two. The selection of wildfire presences—conditioned on 703 the availability of historical records—enables event filtering based on the attributes in the 704 wildfire inventory, e.g., date of occurrence, and burned area. In contrast, sampling absence 705 observations, especially in the spatiotemporal domain, offers broader flexibility. In this re-706 search, the absence sampling was guided by set of structured rules: as detailed in Section 707 4.1, we (a) masked out trivial terrain based on land cover data, (b) implemented tempo-708 ral buffers in locations that experienced wildfires before, (c) ensured temporal balance by 709

sampling across different years and months, and (d) excluded trivial periods. These steps
 collectively aimed to reduce sampling biases and improve the model generalization.

The modeling approach considers the integration of static and dynamic factors to ac-712 count for the complex interplay of wildfire occurrence. An advantage of this framework is 713 the interpretability of the model outcomes, which we briefly highlight in the following dis-714 cussion. The optimal time window analysis identified that the combination yielding the 715 highest predictive performance includes Te_0 , the mean daily temperature on the obser-716 vation day, and P_{28} , the cumulative precipitation over the preceding 29 days (see Figure 717 4). While this combination achieved the highest median AUROC, it is noteworthy that 718 performance differences among alternative time window configurations were relatively 719 small, with AUROC values ranging from 0.77 to above 0.81. This reinforces the impor-720 tance of precipitation and temperature as dynamic predictors, while also indicating that 721 the model results are relatively stable and not highly sensitive to changes in the selected 722 time windows. 723

Regarding the modeled relationships illustrated in Figure 5 Aspect depicts high prob-724 abilities in slopes facing south (135°–225°). In the northern hemisphere, south-facing 725 slopes receive more solar radiation during the day than others, and this translates to lower 726 average moisture levels and thus increasing wildfire probabilities, as several studies indi-727 cate (Alexander et al., 2006; Oliveira et al., 2013; Oliveras et al., 2009). The tree density cover 728 reveals that areas with very sparse tree cover (<20%), may lack sufficient fuel to sustain 729 wildfires, while densely forested areas (>70%) appear to be more resilient to wildfire ini-730 tiation due to forest microclimate effects—tendency to be cooler, more humid and with 731 less sun exposure. In contrast, regions with intermediate and fragmented tree density 732 cover (20–70%) exhibit the highest wildfire probabilities, likely due to a combination of 733 fuel availability and increased edge effects (Song et al., 2017; Rotbarth et al., 2025). 734

The *total annual precipitation* was used as a proxy to represent the spatial variability 735 of long-term precipitation regimes across the study area. The results indicate that drier 736 regions are more susceptible to wildfire occurrence, whereas wetter areas are associated 737 with lower wildfire probabilities. Similarly, the *mean annual temperature* captures broader 738 spatial thermal gradients, revealing that relatively warmer zones are generally more prone 739 to wildfires than colder ones. These drier and warmer areas tend to exceed the criti-740 cal probability thresholds earlier under specific temperature and precipitation conditions, 741 compared to their wetter and cooler counterparts. 742

Although the wildfire inventory lacks the specific data on the fire causes for our study 743 area, previous research in comparable alpine regions has identified human activity as a 744 dominant ignition source (Arpaci et al., 2014; Müller et al., 2020b; Arndt et al., 2013; Müller 745 et al., 2020a). In our analysis, distance to buildings was used as a proxy to account for an-746 thropogenic influence. The results indicate that areas closer to infrastructure—and thus 747 more accessible to human activity—are associated with higher wildfire probability. Ini-748 tially, distance to roads was also considered an additional proxy; however, including both 749 predictors in the model fit introduced redundancy, leading to the loss of statistical signif-750 icance in one of them. Therefore, we retained only *distance to buildings*, given its broader 751 spatial representation of human influence. 752

The *land cover*, identified as the second most influential predictor, served as a proxy for fuel availability. Compared to the reference category *agricultural land*, all other classes ex-

hibited association with higher wildfire probability. Overall, agricultural land appeared to 755 be the least prone to wildfire occurrence. Among the remaining classes, *marshes and peat*-756 bogs were slightly less preferred regarding wildfire occurrence. Both *coniferous tree cover* 757 and *broadleaf tree cover* showed nearly identical regression coefficients, suggesting a com-758 parable influence on wildfire probability—consistent with the findings reported in Müller 759 et al. (2020a). The herbaceous vegetation and heathland class displayed a slightly higher coef-760 ficient, pointing to increased fire likelihood. This supports the conclusions in Oliveira et al. 761 (2013), where it was reported that shrublands, grasslands, and sparsely vegetated areas 762 are more prone to ignition due to their intrinsic flammability and ease of combustion. 763

The predictor *doy*, used as a proxy for seasonal effects, is expected to capture the tempo-764 ral distribution of wildfire occurrences throughout the year. However, since the model al-765 ready incorporates two dynamic predictors, P_{28} and Te_0 , which inherently exhibit seasonal 766 variability, we interpret the effect of *doy* as capturing residual seasonality not explained by 767 these two meteorological predictors. Notably, the model reveals elevated wildfire prob-768 ability in early spring, a period typically associated with rapid warming and snowmelt. 769 This suggests that the selected meteorological predictors may be less effective in fully cap-770 turing the seasonal dynamics influencing fire occurrence during this transitional period. 771 Conversely, doy indicates comparatively lower fire probabilities during summer, which 772 we interpret as evidence that P_{28} and Te_0 already account for much of the seasonal vari-773 ation during that time. In practical terms, this implies that the wildfire probability may 774 be higher in early spring than in late summer for a given location and under the same 775 precipitation and temperature conditions. 776

These results are a sample of the complex interplay between dynamic meteorological and static environmental factors in determining wildfire probability, providing quantitative insights into the spatiotemporal dynamics of wildfire occurrence across the landscape. Also, the analysis shows that identical meteorological conditions can produce markedly different wildfire probabilities depending on the underlying environmental context, highlighting the importance of incorporating dynamic and static predictors in wildfire assessment.

784 6.2 Limitations and future perspectives

Among all the predictors, the 29-day cumulative precipitation P_{28} and the 1-day mean tem-785 perature Te_0 emerged as the most influential ones in the model, as shown in Figure 6. Both 786 P_{28} and Te_0 act as proxies for fuel moisture, so that drier fuels, resulting from prolonged 787 dry and warm periods, are more prone to wildfire ignition than moist fuels. Hence, lower 788 cumulative precipitation and high temperature levels reflect higher wildfire occurrence 789 probabilities. However, limitations arise from the use of these two predictors. The use 790 of absolute values in precipitation and temperature may mislead the results, since equal 791 levels of these two meteorological predictors can have different impacts depending on 792 the local climatologies. For instance, an area experiencing usually high temperatures and 793 anomalously low precipitation—relative to the long-term averages—may exhibit height-794 ened wildfire probabilities compared to an area with equal static settings for which such 795 dynamic conditions are typical. In our model, we included total annual precipitation and 796 mean annual temperature averaged over 30 years to partly account for these climatologies. 797

Still, the variable importance assessment indicated they were among the least influential 798 predictors. We believe that the use of anomalies in the precipitation and temperature, 799 rather than absolute values, could provide the means to directly account for the clima-800 tologies and improve the overall interpretability of the model. Future perspectives of the 801 model could benefit from integrating indices such as temperature anomalies (e.g., daily 802 deviations from the climatological means) or counts of days above-below specific temper-803 ature thresholds. In terms of precipitation widely recognized drought indicators such as 804 the standardized precipitation index (SPI; McKee et al., 1993) or the standardized pre-805 cipitation evapotranspiration index (SPEI; Vicente-Serrano et al., 2010) could be imple-806 mented. These indices offer standardized, temporally dynamic representations of hydro-807 logical stress and have proven valuable in capturing preconditions favorable for wildfire 808 occurrence (Turco et al., 2017, 2018; Smith et al., 2023). 800

In a similar context, one of the most relevant, yet underrepresented aspects in this 810 model is the role of compound events, particularly droughts and heatwaves. Compound— 811 simultaneous occurrence—hot and dry extreme events can trigger and exacerbate cascad-812 ing processes such as wildfires (Richardson et al., 2022). Atmospheric heatwaves, defined 813 as prolonged and consecutive periods of anomalously high temperatures, have been con-814 sistently linked to increased wildfire activity (Barriopedro et al., 2023). Similarly, "hot 815 drought" conditions (characterized by warm temperatures and simultaneous meteoro-816 logical drought) represent elevated fire-weather types. Recent studies have identified two 817 heat-induced fire-weather types, "heatwave" and "hot drought", which collectively ac-818 count for nearly 50% of wildfires in the Mediterranean region (Ruffault et al., 2020; Santos 819 et al., 2024). These events are marked by significantly elevated temperatures, low humid-820 ity, and dry fuels, often with anomalous values compared to surrounding days. Including 821 compound drought and heatwave indices in such a model would allow a more realistic 822 characterization of wildfire scenarios (Lemus-Canovas et al., 2025), especially in light of 823 expected increases in their frequency under global warming (Felsche et al., 2024). 824

As is the case for the meteorological predictors, further model improvements could 825 involve expanding some static predictors into dynamic representations. For example, *land* 826 *cover*, used as a proxy for fuel type, could be replaced or complemented by dynamic fuel 827 maps that reflect seasonal changes in vegetation structure and flammability. Similarly, the 828 anthropogenic influence—currently represented by static proximity to infrastructure— 829 could be much further elaborated through time-varying proxies such as mobility data, 830 population fluxes, or seasonal tourism patterns (Pittore et al., 2023; Renner et al., 2018). 831 Another limitation pertains to uncertainty in the wildfire inventory. In South Tyrol, the 832 dataset is point-based, yet the meaning of these points, whether they represent ignition lo-833 cations or random points within burned polygons, is unclear. In Trentino, wildfire points 834 were generated from polygon centroids, assuming that ignition occurs at the spatial center 835 of the fire. This spatial uncertainty could affect the accuracy of predictor extraction, par-836 ticularly for sensitive predictors. Furthermore, another point of improvement could be 837 adding the potential wildfire spread in addition to the occurrence probability. Incorpo-838 rating fire spread in a second-stage model would require a potentially different modeling 830 framework and additional predictors (e.g., wind speed and direction, slope steepness; 840 Povak et al., 2018; Linn et al., 2007; Jellouli et al., 2022). Such an extension could sig-841 nificantly enhance the predictive utility of the framework, particularly for operational or 842

⁸⁴³ early-warning applications.

To better tailor the model toward dry conditions, we applied a threshold-based filter-844 ing to exclude periods characterized by excessive precipitation. However, the downside of 845 this approach is that it removed a significant number of summer wildfire records, many 846 likely linked to convective storms and lightning activity. These events are particularly im-847 portant in alpine regions, where afternoon thunderstorms are frequent and lightning is 848 a known natural ignition source. As a result, our model may underrepresent lightning-849 induced fires, specifically those occurring during so-called "wet" periods. In future devel-850 opments, we recommend including classified atmospheric circulation patterns (Lemus-851 Canovas et al., 2019) as part of the filtering rule to better account for lightning-induced 852 events. 853

Wildfires are more complex than other natural hazards, which often have natural phys-854 ical triggers. A wildfire can only occur if a source of ignition is present, such as lightning as 855 a natural source, or the much more frequent (in)direct human influence (Chuvieco et al., 856 2003; Müller et al., 2020a). While our model includes a proxy for anthropogenic influence, 857 *distance to buildings*, it does not explicitly account for lightning or human influence as trig-858 gers. As such, the model outputs should not be interpreted as a pure relative probability 859 of wildfire initiation, but rather as the probability of an area experiencing a wildfire given 860 that ignition occurs (human or lightning). This distinction is crucial, particularly when 861 using the model to inform practitioners. In this context, the spatial predictions reflect the 862 relative predisposition of the landscape to wildfires under specific static terrain and dy-863 namic meteorological conditions. Future studies should explore more direct proxies of 864 human activity, such as forest trail density, proximity to power lines, proximity to cable 865 cars, mobility datasets, and investigate multi-scale dynamic factors influencing ignition 866 potential (Arpaci et al., 2014; Chuvieco et al., 2010). 867

6.3 Considerations on early warning

The most relevant model outcome, the wildfire prediction scores, can be illustrated as 860 continuous probability surfaces (e.g., ranging from 0 to 1) or discrete categorized classes. 870 While the continuous outputs retain the full granularity of the model, categorized maps 871 are often preferred in decision-making contexts for their intuitive representation of dif-872 ferent danger levels. However, deriving classes from statistical properties of the data dis-873 tribution, such as quantiles, geometry, and standard deviation, can obscure the practi-874 cal meaning and limit the interpretation for end-users. To address this, we adopted the 875 classification strategy proposed in Steger et al. (2023), which categorizes raw probabili-876 ties based on their association with key performance metrics such as the TPR and false 877 alarms, which hold particular relevance in early warning. Among the evaluated thresh-878 olds, the optimal cut-off *OPT* provided a balanced trade-off, capturing 77% of observed 870 wildfires while limiting the false alarms to 25%. Alternative threshold such as the *TPR95* 880 (high sensitivity) and the *TPR35* (low false alarms) allow the approach to be tailored to 881 different risk priorities. Such priorities can be seen in a context where a false alarm may 882 lead to unnecessary allocation of resources, whereas a missed alarm may have more severe 883 consequences—such as failing to issue timely warnings and inadequate preparedness for 884 wildfire events. However, as our results show, prioritizing high sensitivity (95%) substan-885

tially increases false alarms (59%), whereas minimizing false alarms (e.g., 5%) comes at the cost of reduced wildfire classification rates (TPR = 35%).

Other approaches for threshold selection are particularly valuable when dealing with 888 imbalanced datasets—such as those with few wildfire observations and many absences— 889 where standard metrics like accuracy may become less informative (Saito et al., 2015). In 890 such contexts, the thresholding can be guided based on the precision-recall curve, where 891 the thresholds are chosen based on the recall (TPR, proportion of actual wildfires correctly 892 classified) and precision (the proportion of predicted wildfire locations that correspond 893 to actual wildfires). This can provide a more meaningful assessment of predictive perfor-894 mance, especially for early warning applications targeting rare but high-impact hazards 895 like wildfires (Patton et al., 2023). In any case, the threshold selection should always be 896 done accordingly to the end-user needs, whether the concern may be the cost of the false 897 alarms or if the tolerance is low to accept almost any alarm. 898

The temporal analysis showcased in this study confirms the model capacity to reflect 899 dynamic changes in wildfire probabilities in response to evolving meteorological condi-900 tions. While originally applied for hindcasting, the model can be extended to nowcasting 901 and forecasting applications, provided that the input levels for precipitation and tempera-902 ture do not substantially deviate from the fitting dataset. Notably, the model also enables 903 exploratory what-if scenarios, a valuable asset in the context of climate risk prepared-904 ness and adaptation planning. For example, users can assess the implications of a drop 905 in cumulative precipitation P_{28} under anomalously high temperatures Te_0 for given static 906 settings and a given seasonal context. Additional scenario-building can explore the role of 907 fuel-related predictors such as changes in *tree cover density* or transitions between *land cover* 908 classes, to evaluate their influence on wildfire probability. This narrative-driven simula-909 tions offer a flexible tool to assess potential wildfire outcomes under hypothetical future 910 climatic conditions and support more informed and forward-looking decision-making in 911 wildfire risk management. 912

Our model delivers wildfire predictions at a relatively high spatial resolution (50 913 m) and daily temporal resolution, offering significantly finer granularity than existing 914 global and European early warning systems, such as EFFIS (current resolution: 9 km; 915 Di Giuseppe et al., 2025) and the former global early warning system for wildland fire 916 (https://gfmc.online/gwfews/index-12.html; de Groot et al., 2006), suspended on 2021. 917 This high-resolution capability is particularly advantageous in complex Alpine environ-918 ments like our study area, where wildfires tend to be small-scale and spatially heteroge-919 neous. Nevertheless, benchmarking to better understand potential benefits and limita-920 tions in model outcomes is still relevant for future perspectives. 921

922 7 Conclusion

In this research, we employed GAMs to predict wildfire initiation in space and time, ac-923 counting for static terrain attributes and dynamic meteorological conditions in the region 924 of Trentino–South Tyrol, Italy. We captured the meteorological drivers of wildfire occur-925 rence based on a combination of 29-day cumulative precipitation (P_{28}) , 1-day mean tem-926 perature (Te_0) , and day of the year (doy) factor. These dynamic predictors were comple-927 mented by static proxies representing fuel type and structure, meteorological predisposi-928 tion, topographic features, and anthropogenic influence. Particular emphasis was placed 929 on implementing a representative sampling scheme and generating interpretable outputs, 930 including the visualization of the modeled relationships and variable importance. Model 931 predictions were further linked and classified using combinations of true positive rates 932 and false alarms to derive thresholds that support practical applications in the realm 933 of early warning. The model showed strong predictive performance, with AUROC val-934 ues generally exceeding 0.80 under a rigorous multi-validation framework accounting for 935 spatial and temporal variability. Beyond predictive skill, we illustrated the model utility 936 by hindcasting a two-week period of elevated wildfire activity in July 2022. The flexible 937 framework can be adapted to other natural hazards governed by the interplay of static and 938 dynamic factors. We see our approach as a contribution to the broader development of 939 data-driven solutions that characterize the occurrence of natural hazards through the joint 940 assessment of static controls and time-varying triggers. Crucially, we emphasize that such 941 models should prioritize input data quality, carefully designed sampling strategies, and 942 the interpretability of results to ensure their value in scientific and operational decision-943 making contexts. 944

945 8 CRediT authorship contribution statement

Mateo Moreno: Writing – original draft, Visualization, Validation, Methodology, Formal
analysis, Data curation, Conceptualization. Stefan Steger: Writing – review & editing,
Supervision. Laura Bozzoli: Data curation, Writing – review & editing. Stefano Terzi:
Writing – review & editing. Andrea Trucchia: Writing – review & editing. Cees van
Westen: Writing – review & editing, Supervision. Luigi Lombardo: Writing – review &
editing, Supervision.

952 9 Financial support

⁹⁵³ This research has been supported by the EO4MULTIHA project (4000141754/23/I-DT), ⁹⁵⁴ funded by the European Space Agency.

955 10 Code and data availability

The modeling procedure was conducted in R. The scripts are available at the repository https://github.com/mmorenoz/Wildfire_EarlyWarning. The daily precipitation and temperature fields are available in Crespi et al. (2020).

959 11 Acknowledgement

The research that led to these results is related to the EO4MULTIHA project (https://eo4so-960 ciety.esa.int/projects/eo4multihazards/, which received funding from the European Space 961 Agency (ESA). We thank the Faculty of Geo-information Science and Earth Observation 962 (ITC) – University of Twente for covering the open-access publication fees. We thank Dr. 963 Serkan Girgin for his support in using the CRIB platform. We thank Dr. Paolo Fiorucci 964 and Dr. Marj Tonini for their valuable feedback. Finally, we thank the Office for Forest 965 Planning of the Autonomous Province of Bolzano, especially Alessandro Andriolo, for 966 providing and revising the wildfire data. 967

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