1	Anomalous structure of MgCO3 liquid and the buoyancy of carbonatite melts
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8 9	[This preprint has been submitted to Earth and Planetary Science Letters]
9 10	Key Points
11	• The structure of MgCO ₃ liquid is unique among the other alkaline earth carbonates.
12	 Mg²⁺ adopts anomalously low 4-fold coordination with O₂⁻ and CO₃⁻ anions.
12	 MgCO₃ dramatically lowers carbonate liquid density and increases compressibility.
14	• Assuming MgCO ₃ -like structure, Fe-rich carbonatites remain buoyant in the mantle.
15	
16	Abstract
17	MgCO ₃ is one of the most important components of mantle-derived carbonatite melts, and
18	yet also one of the most difficult to study experimentally. Attempts to constrain its thermodynamic
19	properties are hampered by decarbonation, which occurs at only ~500 °C, far below its metastable
20	1 bar melting temperature. Molecular dynamic simulations, however, can predict the
21	thermodynamic properties of the MgCO ₃ liquid component in spite of experimental challenges.
22	Using the recently developed empirical potential model for high-pressure alkaline earth carbonate
23	liquids (Hurt and Wolf 2018), we simulate melts in the MgCO ₃ -CaCO ₃ -SrCO ₃ -BaCO ₃ system
24	from 773 to 2373 K up to 20 GPa. At 1 bar, MgCO3 liquid assumes a novel topology characterized
25	by a 4-fold coordination of the metal cation (Mg) with both the carbonate molecule and oxygen
26	ion; this is distinct from the other alkaline earth carbonate liquids in which the metal cation is in
27	~6- and ~8-fold coordination with carbonate and oxygen. With increasing pressure, MgCO ₃ liquid
28	structure becomes progressively more like that of (Ca, Sr, Ba)CO ₃ liquids with Mg ²⁺ approaching
29	6-fold coordination with carbonate groups. The novel network topology of MgCO3 liquid results
30	in a melt that is significantly more buoyant and compressible than other alkaline earth carbonate
31	liquids. Simulations of mixed MgCO3-bearing melts show that molar volume, compressibility,

32 enthalpy and heat capacity do not mix ideally with (Ca, Sr, Ba)CO₃ liquids at 1 bar, a consequence 33 of preferential metal-cation ordering in MgCO₃-bearing mixtures. As pressure increases, however, 34 mixing progressively approaches ideality with respect to molar volume, becoming nearly ideal by 35 12 GPa. The model is further applied to mantle-derived primary carbonatite melts with 36 compositions, temperatures and pressures determined by published phase equilibrium 37 experiments. The voluminous structure of liquid MgCO₃ results in a buoyant melt that inhibits a 38 density crossover with the surrounding mantle. Assuming FeCO₃ liquid also adopts the same 39 anomalous high-volume structure as MgCO₃, we predict that even the most Fe-rich 40 ferrocarbonatites would remain buoyant and be barred from sinking or stagnating in the mantle.

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42 **Keywords:** MgCO₃ melt, carbonate liquid structure, carbonatite density, alkaline earth carbonates

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1. Introduction

45 Carbonate plays an important role in partial melting of the mantle. It can lower the solidus of mantle peridotite and eclogite up to ~600 °C (e.g. Dasgupta and Hirschmann 2006; Dasgupta 46 47 and Hirschmann 2007; Hammouda 2003), producing low degree partial melts of 0.03-0.3% 48 carbonatite liquid (e.g. Green and Wallace 1988; Dalton and Wood 1993; Dasgupta and 49 Hirschmann 2006). Carbonatite liquids are highly mobile owing to their ultralow viscosities (e.g. 50 Kono et al. 2014). They are efficient agents of metasomatism (e.g. Green and Wallace 1988), are 51 important for the petrogenesis of ocean island basalts (e.g. Dasgupta et al. 2006) and impact the 52 deep carbon cycle (e.g. Dasgupta and Hirschmann 2010).

53 Of all the carbonate liquid components, MgCO₃ is one of the most geologically relevant. 54 Carbonate enters the mantle through subduction of hydrothermally altered oceanic crust (e.g. 55 Staudigel et al. 1989). Calcium from CaCO₃ partitions readily into the silicate phases of the mantle 56 and is replaced by Mg, establishing MgCO₃ as a dominant carbonate mantle component 57 (Biellmann et al. 1993). Investigations into the composition of primary carbonatite melts produced 58 by partial melting of carbonated mantle eclogite and peridotite show that MgCO₃ along with 59 CaCO₃ and FeCO₃ are the major components of interest (Dalton and Wood 1993; Dalton and 60 Presnall 1998; Hammouda 2003; Ghosh et al 2009).

61 Carbonated mantle systems have been extensively studied via phase equilibrium 62 experiments (e.g. Dalton and Wood 1993; Dalton and Presnall 1998; Hammouda 2003; Ghosh et

63 al 2009). While such experiments are immensely valuable, they have limitations. For example, 64 small variations in the starting composition can result in marked changes to the location of the 65 solidus in P-T space (Dasgupta and Hirschmann 2010). This is problematic given the significant compositional heterogeneity of carbonated mantle, both in its silicate and carbonate phases. 66 67 Compositional complexity is further compounded by wide ranges of temperatures and pressures 68 where partial melting may occur. Phase equilibrium experiments cannot be performed for every 69 relevant composition, temperature and pressure. Furthermore, experimenters often adjust 70 temperature and pressure together, making it difficult to disentangle temperature- and pressure-71 effects.

72 Thermodynamic models such as MELTS (Ghiorso et al. 2002) or THERMOCALC (Powell 73 and Holland 1988) can help resolve these issues but they rely on knowledge of the standard state 74 thermodynamic properties. For the most important carbonate crystal phases (e.g. magnesite, 75 siderite and calcite), the standard state properties are well known (e.g. Berman et al. 1985). 76 However, knowledge of the standard state thermodynamic properties of the respective liquids is 77 limited because alkaline earth carbonates and FeCO₃ decompose at temperatures lower than their 78 1 bar melting temperatures. In spite of the difficulties posed by decarbonation, some properties of 79 the CaCO₃ component have been determined experimentally (e.g. Liu and Lange 2003; Hurt and 80 Lange 2018; O'Leary et al 2015; Hurt 2018), but there is a near total absence of knowledge of the 81 standard state thermodynamic properties of MgCO₃ and FeCO₃ liquid. This is due to the very low 82 1 bar decarbonation temperatures (only 500° C for MgCO₃, Hurst 1991), which makes experiments 83 on MgCO₃- and FeCO₃-bearing liquids nearly impossible at 1 bar.

84 At present, MD simulations provide the only plausible path for obtaining the complete set 85 of carbonate liquid endmember properties needed to incorporate carbonate-silicate melting into a 86 thermodynamic modeling framework like MELTS. The empirical potential model presented in 87 Hurt and Wolf (2018) is designed specifically for simulations of alkaline earth carbonate liquids 88 at mantle conditions. In this study, we apply that model to simulate liquids in the MgCO₃-CaCO₃-89 SrCO₃-BaCO₃ quaternary system from 773 to 2373 K up to 20 GPa. The simulations supply crucial 90 constraints on the standard state thermodynamic properties of pure MgCO₃ liquid which are 91 difficult or impossible to determine experimentally.

Apart from endmember thermodynamic properties, the simulations explore liquid structure and mixing behavior in the MgCO₃-CaCO₃-SrCO₃-BaCO₃ quaternary system. The local

94 structure of carbonate liquids can be characterized by the coordination of the metal cations with 95 neighboring oxygens and carbonate molecules. In the recent experimental studies of Hurt and 96 Lange (2019) and Hurt (2018), the molar volume and compressibility of MgCO₃ is estimated by 97 systematic variations with the alkali and alkaline earth carbonates. Their estimates, however, depend on liquid structure, namely the coordination of Mg^{2+} with the carbonate groups. These 98 99 coordination numbers, which are obtainable via MD simulations, are thus necessary for correctly 100 interpreting the systematic trends in experimental density and compressibility measurements 101 among alkali and alkaline earth carbonate liquids and will allow meaningful estimates of the 102 MgCO₃ and FeCO₃ liquid components to be made based on systematics.

103 To assess whether carbonate melts are buoyant throughout the upper mantle and transition 104 zone, simulations are also performed on CaCO₃-MgCO₃ binary compositions from 1.5 to 20 GPa 105 and 1423 to 1873 K along various pressure-temperature paths approximating that of subducting 106 slabs. The simulation composition, temperature and pressure are representative of near-solidus 107 carbonatite melts formed by low-degree partial melting of carbonated eclogite and peridotite. This 108 is done in order to estimate the density of primary carbonatite melts as a function of depth and 109 determine whether any compositions of carbonate melt are sufficiently dense to either stagnate 110 upon initial melting or sink into the lower mantle. Given the likely similarity of MgCO₃ and FeCO₃ 111 liquid structures and volumes, we also explore the implications for ferrocarbonatite melts in the 112 mantle.

113

2. Methods

For this study, we use a previously published empirical potential model for the simulation of alkaline earth carbonate liquids using rigid ions and partial charges. A complete derivation and assessment of this model is available in (Hurt and Wolf 2018) where it was also applied to the simulation of liquids in the CaCO₃-SrCO₃-BaCO₃ system. This model is now applied to molecular dynamic (MD) simulations of MgCO₃-bearing liquids performed using the LAMMPS code (Plimpton, 1995).

42 different liquid compositions are simulated within the MgCO₃-CaCO₃-SrCO₃-BaCO₃ quaternary over a temperature range of 773-2373 K and up to 20 GPa. MD simulations are performed with a timestep of 1 fs and total runtimes of 120 ps. Each simulation begins with an equilibration phase of 60 ps using the canonical (NVE) ensemble with the Berendsen barostat and thermostat (Berendsen et al., 1984), which equilibrates quickly and smoothly but deviates slightly from the canonical ensemble. A second equilibration phase of 40 ps is performed under the NPT ensemble using the Nosé-Hoover barostats and thermostats (Nosé, 1984; Hoover, 1985). A third phase of 20 ps continues the NPT ensemble and constitutes the final production run.

Every simulation is checked to ensure that system energies and volumes have converged to within 0.02% and 0.27%, respectively. This level of volume convergence is achieved with a simulation size of 6860 atoms. Simulations are initialized with atomic positions and velocities corresponding to CaCO₃ liquid at 2.275 g/cm³ density (equivalent to 1 bar pressure) and 1623 K.

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3. Results

133 3.1 Equation of state of MgCO₃ liquid

Hurt and Wolf (2018) previously showed that CaCO₃, SrCO₃ and BaCO₃ liquids have remarkably similar properties, with nearly parallel compression curves, 1 bar molar volumes that increase systematically with cation radius and congruent thermal expansion coefficients. Such systematic variation in molar volume and congruent thermal expansions, are corroborated by the experiments of Hurt and Lange (2019) and Hurt (2018).

139 To investigate whether the properties of MgCO₃ liquid follow the trends observed among 140 CaCO₃, SrCO₃ and BaCO₃, the thermodynamic properties of MgCO₃ liquid have been calculated 141 from the simulations (complete details and results are listed in Table A.1). Following Hurt and Wolf (2018), we attempt to fit the MgCO₃ simulation results with a temperature-dependent 3rd 142 143 order Birch-Murnaghan equation of state (EOS), but the resulting models show a markedly poor 144 fit to the data (Table 1). This is further exemplified by the large error bars on the fitted parameters. 145 Unlike (Ca, Sr, Ba)CO₃ liquids, MgCO₃ liquid is highly compressible at pressures <2 GPa, and its rapid change in molar volume cannot be accommodated by a 3rd order Birch-Murnaghan, thus 146 requiring us to adopt a 4th order Birch-Murnaghan EOS (details of the EOS are available in 147 148 Appendix A).

The full set of fitted parameters for this EOS is given in Table 1. This 4th order Birch-Murnaghan EOS model recovers the simulated molar volumes in the 1100 – 2000 K temperature range up to 20 GPa with RMS volume residuals of 0.27 cm³/mol (0.81%). The 1100 K compression curve predicted by this best-fit EOS model is shown in Fig. 1 along with the raw simulated molar volume data as a function of pressure up to 12 GPa. The simulated compression curves of CaCO₃, SrCO₃ and BaCO₃ liquids are also shown for comparison, together with the expected compression 155 behavior for MgCO₃ liquid (shown as a dotted line) assuming it follows the systematic trend of 156 the other alkaline earth carbonates (discussed in detail below). Clearly, the model predicts a curve 157 for MgCO₃ that is radically different from the other three alkaline earth carbonate components. 158 The 1 bar molar volume is near that of SrCO₃ and decreases dramatically with pressure, implying 159 a 1 bar compressibility far-exceeding that of the other alkaline earth carbonate liquids. The 160 compression curve rapidly flattens out and at pressures > 5 GPa, where it begins to parallel the 161 other alkaline earth carbonates. Liquids derive many of their thermophysical properties from their 162 average atomic structure, and thus we explore the pressure-dependent atomic structure of MgCO₃ 163 liquid to discover the root cause for its unique behavior among the alkaline-earth carbonates.

164 3.2 Distinct structure of MgCO₃ liquid

165 Though liquids lack long-range order, they do possess a short-range order that can be 166 described using a pair distribution function (pdf). For a particular pair of atomic types (*i* and *j*), 167 pdf curves give the average density of *j* as a function of distance from *i*. Liquid structure can also 168 be described by coordination numbers for a given atomic pair. Average coordination numbers are 169 calculated as weighted averages of the pdf curves, $g_{ij}(r)$:

$$CN_{ij} = 4\pi \int_0^{r_1} r^2 g_{ij}(r) \rho_j \, dr$$
 Eq. 1

170 where *r* is the atomic separation distance, ρ_j is the average atomic number density of atom *j*, and 171 r_1 is the maximum cutoff radius located at the first minimum in the PDF, representing the outer 172 boundary of the first nearest-neighbor peak.

Hurt and Wolf (2018) used the same empirical potential model to simulate (Ca, Sr, Ba)CO₃ liquids, and demonstrated that they have remarkably similar liquid structures and coordination numbers for the metal-oxygen (M-O), carbon-carbon (C-C) and metal-carbon (M-C) pairs. At ambient pressure, C-C and M-C coordination numbers are the same for each liquid at 10.9 and 5.9 respectively. M-O coordination numbers are similar for the three endmember liquids, increasing systematically from Ca to Ba (Ca-O: 6.9, Sr-O: 7.3 and Ba-O: 7.7).

Based on the new simulations performed in this study, we analyze the structure of MgCO₃ liquid. The 1 bar results are summarized in Fig. 2, which shows a distinctly different atomic structure and coordination environment for MgCO₃ as compared to the other alkaline earth carbonate liquids, visible in terms of C-C, M-O and M-C pairs. At 1 bar, the Mg-O and Mg-C coordination is tetrahedral (4), while the C-C coordination (11.6) is nearly ideal closest-packing,
compared to CaCO₃-SrCO₃-BaCO₃ liquids (10.9). The most significant of these differences is that
the M-C coordination, which is consistently 6 for the other alkaline earth carbonate liquids, is only
4 for MgCO₃—the same as in alkali carbonates (e.g. Roest et al 2017).

187 To determine whether the unique structure of MgCO₃ liquid persists at high pressure, 188 isothermal compression simulations are analyzed at 1100 K between 1 bar and 12 GPa. M-C and 189 M-O coordination numbers are calculated at each pressure and shown in Fig. 3. As pressure 190 increases, the Mg-O coordination number increases linearly from 4 (at 1 bar) to 5.6 (at 12 GPa). 191 In the high-pressure regime, M-O coordination number adopts systematic behavior among the 192 alkaline earth carbonates: the coordination number is ~10 for Ba-O, ~9 for Sr-O, ~8 for Ca-O and 193 ~6 for Mg-O. Thus, at pressure, the Mg-O coordination number is consistent with what might be 194 expected from periodic table systematics. For the M-C coordination numbers, CaCO₃-SrCO₃-195 BaCO₃ liquids average around 6 at 1 bar and then increase modestly to ~6.8 by 12 GPa. The Mg-196 C coordination starts off at 4 at 1 bar and converges rapidly to values consistent with the other 197 alkaline earth carbonates. Thus, simulations indicate that the liquid structure of MgCO₃ becomes 198 progressively more like that of the other alkaline earth carbonates with compression, becoming 199 nearly indistinguishable by 12 GPa.

200 3.3 Non-ideal mixing for MgCO₃ liquid at low pressure

Hurt and Wolf (2018) simulated alkaline earth carbonate liquids in three binary systems: CaCO₃-SrCO₃, CaCO₃-BaCO₃ and SrCO₃-BaCO₃. It was found that molar volumes and molar isobaric heat capacities and compressibility all mix ideally across a wide range of temperatures and pressures. To test whether such ideal mixing behavior extends to MgCO₃-bearing liquids, simulations of mixed liquids are performed in the MgCO₃-CaCO₃-SrCO₃-BaCO₃ quaternary system at 1 bar and 1100 K. For any property of interest, we define the mixing quantity as the deviation from a compositionally-weighted average of the endmembers:

$$\Delta Z_{mix} = Z - \sum Z_i X_i$$
 Eq. 2

where *Z* is the molar quantity of interest (like volume or enthalpy) for the liquid mixture, X_i is mol fraction of each endmember, and Z_i is the quantity for each pure endmember at the specified temperature and pressure. The ideality of mixing in terms of both volume and enthalpy is assessed for simulated Mg-carbonate binaries and visualized in Fig. 4 (with direct simulation resultsreported in Appendix Table C.2).

213 Properties determined from MD simulations always contain uncertainties due to random 214 Afluctuations and system size limitations. We can reduce the effects of this noise and better 215 characterize non-ideal mixing by fitting a simple sub-regular mixing model to the simulated ΔV_{mix} 216 and ΔH_{mix} values:

$$\Delta Z_{mix} = \left(w + \Delta w \left[X - \frac{1}{2}\right]\right) \cdot X[1 - X]$$
 Eq. 3

217 where ΔZ_{mix} is the mixing property for either volume or enthalpy, X is mol fraction of MgCO₃, 218 and w and Δw are constants which have been fitted using least squares minimization. The best-fit 219 values of w and Δw at 1100 K and 1 bar for all three binaries are given in Appendix Table C.3. 220 Fig. 4 depicts ΔV_{mix} in panel a and ΔH_{mix} in panel b for each simulated mixture as a function of 221 MgCO₃ content along with the best-fit sub-regular solution model (Eq. 3). Note that, for ΔH_{mix} , 222 simulated total energy values are used instead of enthalpy because they are equal at 1 bar and 223 because simulated enthalpy has larger errors than total energy due to fact that enthalpy calculations 224 fold in fluctuations on both pressure and volume. Fig. 4a shows significant non-ideal volume of mixing for all binaries, especially along the CaCO₃-MgCO₃ binary where ΔV_{mix} is as large as -225 $0.87 \pm .11$ cm³/mol. Volumes of mixing are nearly all predicted to be negative. For all binaries, 226 227 the mixing behavior is slightly asymmetric; the magnitude of ΔV_{mix} reaches a maximum at 60-75 228 mol % MgCO₃. The non-ideal mixing of volumes in MgCO₃-bearing liquids means that volumetric 229 derivative properties such as compressibility also cannot mix ideally.

230 The case of enthalpy of mixing (ΔH_{mix}) is closely analogous. Unlike liquids in the CaCO₃-231 SrCO₃-BaCO₃ system, enthalpy does not mix ideally among MgCO₃-bearing mixtures. Fig. 4b shows that ΔH_{mix} is uniformly negative or statistically consistent with zero across the binary for 232 every mixture. The peak magnitude of ΔH_{mix} increases systematically from CaCO₃ (31 KJ/mol) 233 234 to BaCO₃ (~12 KJ/mol). Within each binary, the magnitude of ΔH_{mix} peaks at 60-75 mol% MgCO₃. Interestingly, the BaCO₃-MgCO₃ system exhibits the lowest ΔV_{mix} but the highest 235 236 ΔH_{mix} ; this suggests that the mechanism responsible for non-ideal mixing in volumetric properties 237 is not the same as the one driving non-ideal mixing in enthalpic properties.

Hurt and Wolf (2018) found that simulations predict isobaric heat capacities to mix ideally according to mol fraction in CaCO₃-SrCO₃-BaCO₃ liquids. Simulations in this study of MgCO₃- bearing binary liquids from 1100 to 2000 K enabled calculation of the isobaric 1 bar heat capacity. For $CaMg(CO_3)_2$, $SrMg(CO_3)_2$ and $BaMg(CO_3)_2$, the heat capacity is estimated to be 146, 140 and 139 J/mol, which are lower than their ideal values by 7%, 11%, and 9%, respectively. Heat capacities are thus systematically lower (9% on average) than expected for MgCO₃-bearingsystems.

245 3.4 Structural ordering and non-ideality in MgCO₃ mixtures

As noted above, liquid properties (e.g. molar volume, compressibility, enthalpy and heat capacity) do not mix ideally in the MgCO₃-binary compositions as they do in the CaCO₃-SrCO₃-BaCO₃ liquid system at 1 bar. The non-ideal source of this mixing should likely manifest as some structural anomaly evident in the pdf curves of mixtures (i.e. there should be some noticeable difference between the pdf curves of pure MgCO₃ compared to an MgCO₃-bearing mixture). Such a difference might help reveal the mechanism driving non-ideal mixing behavior.

252 Within the MgCO₃-CaCO₃-SrCO₃-BaCO₃ quaternary there are six binaries. Various 253 compositions within each binary were simulated at 1100 K and 1 bar and pdf curves were 254 calculated for every atomic pair. The full range of pdf curves for M-O, M-C and C-C pairs for each 255 binary composition is available in Appendix Fig. C.1. The striking uniformity of pdf curves 256 demonstrates that M-O, M-C and C-C pairs in binary liquids retain the same structure and 257 coordination as in the pure liquids. This means that, at 1 bar, Mg²⁺ maintains its 4-fold coordination 258 with carbonate and oxygen independent of bulk composition. Similarly, if small amounts of 259 CaCO₃, SrCO₃ and BaCO₃ are mixed into MgCO₃ liquid, Ca²⁺, Sr²⁺ and Ba²⁺ ions maintain their 260 octahedral coordination with carbonate. While the Mg-O and Mg-C coordination in mixtures is 261 the same as in pure MgCO₃ liquid (i.e. Mg²⁺ maintains a coordination of 4 with carbonate and 262 oxygen atoms) there are some notable differences between the pdf curves of pure and mixed 263 MgCO₃.

Simulations reveal that the key structural difference between pure and mixed MgCO₃ liquids is metal cation ordering. In pure MgCO₃ liquid at 1 bar, the primary Mg-Mg coordination shell is actually composed of two partially overlapping shells. The inner shell contains 3 Mg atoms at an average distance of 3.7 Å, and the outer shell has an additional 9 Mg atoms at an average distance of 5.3 Å. When MgCO₃ is mixed with another carbonate, this primary coordination shell is split, with the inner shell preferentially filled by other metal cations. Figure 5 shows the Mg-Mg and Mg-Ca pdf curves in pure MgCO₃ and MgCa(CO₃)₂; these curves show that Mg²⁺ vacates the 271 inner coordination shell in mixed compounds and is replaced by Ca^{2+} . This kind of cation ordering 272 was not observed in CaCO₃-SrCO₃-BaCO₃ liquids. Such ordering in MgCO₃-bearing mixtures can 273 explain why some properties of MgCO₃-bearing liquids do not mix ideally, while CaCO₃-SrCO₃-274 BaCO₃ liquids do. This behavior is likely related to cation ordering in the calcite-dolomite-275 magnesite series, which produces the highly ordered dolomite structure, where cations of mixed 276 sizes are accommodated with minimal lattice strain (Vinograd 2006). Given that V_{mix} is found to 277 be either negative or insignificant, cation ordering likely allows the liquid to pack space more 278 efficiently for liquid Mg-carbonate mixtures.

279 3.5 Ideal mixing of MgCO₃ liquid at high pressure

280 Though its structure and compressive properties are distinct at low pressures, MgCO₃ 281 liquid increasingly resembles other alkaline earth carbonates at higher pressures. This is apparent 282 in the compression curves, which become nearly parallel between 5 and 10 GPa as shown in Fig. 1. 283 To further investigate, binary mixture simulations of CaMg(CO₃)₂, SrMg(CO₃)₂ and BaMg(CO₃)₂ 284 are also performed at 1100 K up to 12 GPa, assessing whether volumes and enthalpies mix ideally 285 at pressure. The values of ΔV_{mix} calculated from the resulting simulations are $-0.02 \pm .03$, $-0.13 \pm$ 286 .03 and -0.11 \pm .03 cm³/mol for CaMg(CO₃)₂, SrMg(CO₃)₂ and BaMg(CO₃)₂ respectively. These 287 ΔV_{mix} values are dramatically smaller than at 1 bar (which range between -0.34 and -0.91 cm³/mol). 288 Fig. 6 shows the compression behavior of ΔV_{mix} at 1100 K, normalized to the pressure-dependent 289 absolute volume difference between CaCO₃ and MgCO₃ liquid, $\Delta V_{MgCO3-CaCO3}$. This pressure-290 dependent normalization is necessary since the molar volume of the MgCO₃-CaCO₃ binary 291 mixture falls between the volumes of the pure endmembers, and thus at pressures where pure 292 CaCO₃ and MgCO₃ volumes converge, ΔV_{mix} will appear small; normalizing to the volume 293 difference between the endmembers corrects for this. As seen in the figure, ΔV_{mix} approaches 0 as 294 pressure increases, reflecting increasingly ideal mixing.

At low pressures, it was previously shown that metal-cation ordering provides a mechanistic explanation for non-ideal mixing of at least some of the properties in MgCO₃-bearing liquids. Since ΔV_{mix} diminishes to nearly 0 at 12 GPa, we might expect to see little difference between Mg-Mg pdf curves in pure MgCO₃ versus MgCO₃ mixtures at that pressure. The inset on Figure 6 shows, in fact, that the structural difference between pure and mixed MgCO₃ liquids progressively lessens with pressure (up to 12 GPa), supporting the notion that metal cation ordering
 correctly explains non-ideal mixing in volumetric properties among MgCO₃-bearing liquids.

302 In contrast, the simulations indicate that non-ideal mixing of enthalpy does not disappear 303 at high pressure like ΔV_{mix} . At 12 GPa, ΔH_{mix} for CaMg(CO₃)₂, SrMg(CO₃)₂ and BaMg(CO₃)₂ is 304 calculated to be -2.9, -8.2 and -11.7 kJ/mol respectively. Normalizing ΔH_{mix} to the difference in 305 enthalpy between the endmembers would yield a more meaningful comparison (since absolute 306 values of enthalpy increase with pressure). This result gives a normalized ΔH_{mix} of -1.2, -2.1 and -1.9% at 12 GPa compared to -1.1, -1.8 and -4.5% at 1 bar for CaMg(CO₃)₂, SrMg(CO₃)₂ and 307 308 $BaMg(CO_3)_2$, respectively. $BaMg(CO_3)_2$ is the only mixture that undergoes a significant reduction 309 in ΔH_{mix} with pressure, while the normalized ΔH_{mix} for the other binary liquids remain unchanged. 310 This suggests that cation ordering in MgCO₃-bearing liquids is not solely responsible for non-ideal 311 mixing of enthalpies.

312

4. Discussion

313 4.1 Fourfold cation coordination in MgCO₃ liquid at 1 bar

314 Previous molecular dynamic studies have found that most alkaline earth carbonate liquids 315 (CaCO₃, SrCO₃, BaCO₃) adopt similar 1 bar atomic structures, with metal cations in octahedral 316 coordination with carbonate anions and 7-8 fold coordination with oxygen (Hurt 2018, Vuilleumier 317 2014). A thorough search of the literature reveals, to the best of our knowledge, that this study is 318 the first to demonstrate that MgCO₃-bearing melts are unique among those of the alkaline earth carbonates, with Mg^{2+} in tetrahedral coordination with both oxygen and carbonate anions. This 319 320 structural anomaly is all the more surprising, given that Mg-bearing carbonate crystals (magnesite, 321 dolomite, huntite, and northsite) all possess 6-fold oxygen coordination.

In silicate systems, Mg²⁺ is known to enter into 4-fold coordination with oxygen in glasses 322 323 and melts across a wide range of compositions (e.g. Roy 1950; Waseda and Toguri 1977; Kubicki 324 et al 1992; Shimodo et al 2008; Trcera et al 2009; Cormier and Cuello 2013; Morizet et al 2017). In silicate crystals, Mg²⁺ is mostly in 6-fold coordination; however, two silicate crystal phases, 325 spinel and åkermanite (Ca₂Mg[Si₂O₇]), host Mg²⁺ in fourfold coordination (e.g. Kroecker and 326 Stebbins 2000). Thus, while Mg²⁺ is mostly found in 6-fold coordination in solid silicate phases, 327 328 it commonly enters into 4-fold coordination upon fusion, opening the door for similar behavior in 329 carbonate systems.

Unfortunately, it is difficult to assess the coordination of Mg²⁺ in carbonate melts from past 330 analytical studies. Detailed structural information on Mg²⁺ coordination in MgCO₃-bearing 331 332 carbonate melts/glasses is largely unavailable. This is in part because melt structure is commonly 333 determined via experimental analyses (such as XANES or EXAF) of glasses, and while silicate 334 melts readily quench to glass, there is only one entirely carbonate system that is capable of 335 vitrification: K₂CO₃-MgCO₃ (e.g. Datta et al 1964; Ragone et al 1966). This system is doubly 336 unique because it is the only known MgCO₃-bearing carbonate liquid that forms a stable melt at 1 337 bar (Ragone et al 1966). K₂CO₃-MgCO₃ glass structures have been studied by both infrared (Datta et al 1964; Genge et al 1995) and Raman spectroscopy (Sharma and Simons 1980; Genge et al 338 1995). Neither Genge nor Sharma comment on Mg^{2+} coordination in their spectroscopic studies. 339 However, Datta et al (1964) posited that the low refractive index and density of the K₂CO₃-MgCO₃ 340 341 glass (compared to the respective crystals) may result from a low (4-fold) Mg coordination. Datta 342 also recognized the significance of the fact that the only known MgCO₃-bearing carbonate liquid 343 that is stable at 1 bar is also the only carbonate capable of quenching to a glass, implying its 344 structure must somehow be unique among alkaline earth carbonate liquids; Datta thus hypothesized that it could be related to a decrease in Mg²⁺ coordination number. Our MD 345 346 simulations suggest that MgCO₃ liquid does in fact assume a unique topology characterized by a low (4-fold) coordination of Mg^{2+} , which corroborates the suppositions of Datta et al. (1964) on 347 348 the structure of MgCO₃-bearing carbonate melts.

349 4.2 Comparison of MgCO₃ to other alkaline earth carbonate liquids

350 From the simulations of this study, it appears that MgCO₃ has a liquid structure that is 351 distinct from the other alkaline earth carbonates and properties that deviate strongly from those 352 predicted on the basis of systematic variations (see Hurt 2018; Hurt and Lange 2019). Hurt and 353 Lange (2019) explored systematic variations in the 1 bar molar volumes of alkali and alkaline earth 354 carbonate liquids. On the basis of systematic trends, two different estimates of MgCO₃ liquid molar volume were made; one estimate assumes that Mg^{2+} is in 6-fold coordination with carbonate and 355 356 oxygen and has a liquid structure that is like the other alkaline earth carbonates. The other estimate, based on preliminary results from this study, assumes Mg^{2+} is in 4-fold coordination with oxygen 357 358 and carbonate and has a liquid structure that is more like an alkali carbonate. Under the 6-fold 359 assumption, Hurt and Lange predict a 1 bar MgCO₃ liquid molar volume of $34.4 \pm .6$ cm³/mol. In 360 the case of a 4-fold Mg-O coordination, Hurt and Lange (2019) estimated that MgCO₃ would have 361 a 1 bar liquid molar volume of \sim 40 cm³/mol at 1100 K.

362 Our simulations clearly predict 4-fold coordination for Mg²⁺ and a liquid structure that is more like alkali carbonate than typical alkaline earth carbonates (a full discussion of similarities 363 364 between MgCO₃ and alkali carbonate liquid structure is available in Appendix B). In accordance 365 with the estimates of Hurt and Lange (2019), our simulations predict a larger 1 bar molar volume 366 than what might be expected from systematic variations in the molar volumes of alkaline earth 367 carbonate liquids. The fitted Birch-Murnaghan EOS model (Appendix Eq. A.1-A.4) predicts a 1 368 bar MgCO₃ liquid volume of 44.21±.09 cm³/mol at 1100 K. However, it's likely that the actual 369 volume is somewhat lower since the empirical potential model of Hurt and Wolf (2018) typically 370 slightly overestimates 1 bar volumes. For CaCO₃, SrCO₃ and BaCO₃ liquids, the model requires a 371 -0.35, -0.33 and -0.75 GPa pressure correction, respectively, to bring simulated 1 bar volumes into 372 agreement with the experimental projections of Hurt and Lange (2019). An analogous -0.6 GPa 373 pressure correction would bring the MgCO₃ simulations into agreement with the alkali-like 374 structural predictions from Hurt and Lange (2019).

375 Hurt (2018), provides analogous systematic trend equations for the compressibility of 376 alkali and alkaline earth carbonate liquids based on a positive linear relationship between liquid 377 compressibility and volume of fusion. Their model equation predicts a compressibility for MgCO₃ liquid at 1100 K of $\sim 23\pm1$ 10⁻² GPa⁻¹ assuming it fits into the alkali carbonate trend and $\sim 18\pm6$ 10⁻ 378 ² GPa⁻¹ if it fits into the alkaline-earth trend. Given the similarity of the MgCO₃ liquid structure to 379 380 Li_2CO_3 (Appendix Fig. B.1), the alkali carbonate trend might actually provide a better prediction 381 (for more details regarding similarities between MgCO₃ and alkali carbonates, see Appendix B). 382 These estimates of MgCO₃ liquid compressibility differ dramatically from other alkaline earth 383 carbonate liquids (i.e. CaCO₃, SrCO₃ and BaCO₃) which have relatively modest 1 bar/1100 K compressibilities of 4.33 - $5.97 \cdot 10^{-2}$ GPa⁻¹ Our MgCO₃ liquid simulations predict a bulk modulus 384 (K_T) of 6.15±.15 GPa, which is equivalent to a compressibility (β_T) of 16.3 \cdot 10⁻² GPa⁻¹ at 1100 K. 385 386 broadly consistent with the elevated values predicted by Hurt (2018).

387 4.3 Comparison to previous thermodynamic studies of MgCO₃ liquid properties

As noted above, the structure of MgCO₃ liquid becomes more like the other alkaline earth carbonates as it is subjected to progressively higher pressures. By 12 GPa, the M-O and M-C coordination numbers of MgCO₃ liquid converge to values in line with CaCO₃, SrCO₃, and BaCO₃ liquid (Fig. 2). This pressure-induced transformation of liquid structure is also evident in the
 compression curve of MgCO₃ liquid, which begins to parallel the other three alkaline earth
 carbonate liquids at 5-10 GPa (Fig. 1).

394 Our predictions of the standard state thermodynamic properties of MgCO₃ liquid stand in 395 stark contrast to previous estimates made by Kang et al. 2016. These calculations fit a 396 thermodynamic model to the congruent melting curve of MgCO₃, which is bracketed by four 397 experimentally determined points (~1858 K at 3 GPa & 1883 K at 3.6 GPa from Irving and Wyllie 398 1975, and 2183 K at 8 GPa & 2363 K at 15 GPa from Katsura and Ito 1990). This approach arrives 399 at a molar volume of MgCO₃ liquid of \sim 32 cm³/mol and a compressibility of 2.6 \cdot 10⁻² GPa⁻¹ at 400 1100 K and 1 bar. While the MgCO₃ properties proposed by Kang et al. differ markedly from our 401 simulation results, Kang et al. arrived at their values using a method that is statistically 402 problematic. The estimates of standard state thermodynamic properties of MgCO₃ liquid (and 403 FeCO₃ as well) proposed by Kang et al. were calculated by fusion curve analysis. The fusion curve 404 of MgCO₃ is currently bracketed by 4 experimental points (and only 3 for FeCO₃). However, 5 405 thermodynamic parameters are fitted based on only 4 experimental observations. Since the number 406 of fitted parameters exceeds the number of constraining observations, the reported thermodynamic 407 solution is highly under-constrained and non-unique, allowing alternative sets of thermodynamic 408 properties to recover the observables just as well.

409 We can see some of the resulting issues directly in the fitted thermodynamic properties 410 proposed by Kang et al. (2016); for example, the thermal expansions of MgCO₃ and FeCO₃ liquid 411 are identical to the thermal expansion of their respective crystal phases. The thermal expansion of 412 a liquid is generally expected to be substantially greater than that of the crystal, however (e.g. 413 Lange 1997). CaCO₃, SrCO₃ and BaCO₃ liquids, for example, have a thermal expansion ~3 times 414 greater than their respective crystal phases (Hurt and Lange 2018). Kang et al. also give a 415 temperature dependence of compressibility ($\delta\beta_T/\delta T$) of MgCO₃ liquid that is lower, by an order of 416 magnitude, than those of $CaCO_3$, $SrCO_3$ and $BaCO_3$ liquids (Hurt et al. 2018). For these reasons, 417 we strongly recommend the standard state MgCO₃ liquid thermodynamic properties derived from 418 this study.

419 4.4 The density of primary carbonatite melts

420 Natural carbonatite liquids have been shown to be effective agents of mantle metasomatism
421 and scavengers of trace elements (e.g. Blundy and Dalton 2000), and thus the mobility and ascent

422 rates of primary carbonatite melts are particularly geologically important. Ascent rate depends 423 mostly on melt viscosity and density contrast. Viscosity has been measured at temperatures and 424 pressure representative of the mantle for calcitic and dolomitic liquids (e.g. Kono et al. 2014). The 425 densities of some important carbonate liquid components (e.g. CaCO₃) at pressure have been 426 constrained by 1 bar liquid density experiments (e.g. Liu and Lange 2003; Hurt and Lange 2018) 427 and liquid sound speed measurements (e.g. O'Leary et al. 2015; Hurt 2018) as well as by molecular 428 dynamic simulations (e.g. Vuilleumier et al 2014). However, primary carbonatite melts generated 429 in the mantle by low-degree partial melting contain significant amounts of MgCO₃ and FeCO₃ 430 (Dalton and Wood 1993; Dalton and Presnall 1998; Ghosh et al. 2009).

While the densities of primary carbonatite melts along a carbonated mantle liquidus are mostly unknown, it may be estimated by simulations from this study. Using the compositions of low-degree primary carbonatite melts generated in phase equilibrium experiments of carbonated peridotite and eclogite (Dalton and Wood 1993; Dalton and Presnall 1998; Hammouda 2003; Ghosh et al. 2009), simulations at temperature and pressure are performed for simplified experimental compositions by projecting them into the CaCO₃-MgCO₃ binary. Complete details on these calculations are available in Appendix D.

438 The simulated density results are listed in Table 2 and depicted as a function of depth in 439 Fig. 7. The simulations indicate that $CaCO_3$ -MgCO₃ binary liquids will be relatively low in density 440 due to the open network topology of the MgCO₃ melt component. Thus, primary carbonatite melts 441 maintain a significant density contrast with the surrounding mantle rocks throughout the entire 442 upper mantle and transition zone. Even the densities of melts from Hammouda (2003), which 443 represent a cold slab subduction PT path (and are relatively enriched in denser CaCO₃ and FeCO₃ 444 components) never approach that of the mantle. Given the ultralow viscosity of these kinds of 445 liquids, any positive density contrast will swiftly drive liquid ascent; any degree of carbonate melt 446 will thus quickly escape from a descending slab, inhibiting the subduction and storage of carbonate 447 into the lower mantle.

The only common carbonate component heavy enough to potentially affect a density crossover in the mantle is FeCO₃ (which is a common component in mantle-derived carbonatite melts, e.g. Hammouda 2003). Though our work does not directly address iron partitioning in carbonate-rich melts, we can use our model to approximate a minimum threshold iron-content that a hypothetical ferrocarbonatite must contain in order to affect a density crossover within the 453 mantle. While FeCO₃ liquid was not directly simulated in this study, it's likely that FeCO₃ shares 454 the volumetric properties of MgCO₃, since Fe²⁺ and Mg²⁺ have the same valence and similar ionic 455 radii, and thus can frequently occupy the same atomic sites within crystal lattices and liquid 456 pseudo-lattices. Assuming that MgCO₃ liquid molar volume is a reasonable proxy for FeCO₃, the 457 FeCO₃ concentration necessary for affecting a density crossover can be calculated as a function of 458 depth for a ferrocarbonatite in the MgCO₃-CaCO₃-FeCO₃ system. For the purposes of this 459 calculation, it will be assumed that the molar volumes of all three components mix ideally and that $X^{Mg} = X^{Ca}$. The MgCO₃ & FeCO₃ liquid volumes are approximated at temperature and pressure 460 using the thermodynamic properties presented this study (Table 1) in conjunction with Eqs. A.1-461 462 A.4. The partial molar volume of the $CaCO_3$ liquid component is accurately estimated by combining experimental and theoretical constraints, using a 3rd order Birch-Murnaghan EOS with 463 464 properties taken from the 1 bar molar volume model of Hurt and Lange (2019), the compressibility 465 results of Hurt (2018) and the K' results from Hurt and Wolf (2018). The partial molar volume of 466 the MgCO₃, FeCO₃ and CaCO₃ components were calculated at P-T conditions along an average 467 carbonated peridotite solidus given by Dasgupta and Hirschmann (2010):

$$T(^{\circ}C) = 0.0238 \times P^{3} - 2.2084 \times P^{2} \times 73.7991 \times P + 830.3808$$
Eq. 4

468 where P is in GPa. Using this approach, Fig. 8a shows the minimum requisite FeCO₃ concentration 469 needed for a ferrocarbonatite to stagnate in the mantle gradually decreases through the upper 470 mantle from a mol fraction of ~ 1 in the crust, to 0.5 (35 wt.% FeO) at the top of the transition zone. 471 Throughout the transition zone, the threshold iron content is ~0.65 mol fraction FeCO₃. Such iron 472 concentrations far exceed what's observed in phase equilibrium experiments on carbonated 473 peridotite/eclogite (which generally produce carbonatite melts with <10 wt. % FeO). Fig. 8b shows 474 the density of a melt with an approximate composition of a natural ferrocarbonatite (Thompson et 475 al. 2002) which was calculated as a function of depth using the technique described above. As 476 evident, its density remains consistently lower than the mantle through the transition zone.

477 It's difficult to assess whether significant volumes of such highly Fe-enriched (>40 mol% 478 FeCO₃) ferrocarbonatites are actually produced in the mantle. After an exhaustive search of the 479 literature, the ferrocarbonatite in Swartbooisdrif, Namibia (Thompson et al. 2002) appears to be 480 the only natural ferrocarbonatite yet identified with an unambiguously magmatic origin. 481 Thompson gives its composition as approximately Fe_{0.26} Mg_{0.22} Ca_{0.52} CO₃. We find it interesting 482 that the Swartbooisdrif ferrocarbonatite has a composition nearly identical to that of Ankerite, 483 $Ca(Mg,Fe)(CO_3)_2$. If FeCO₃ behaves similarly to MgCO₃, we might expect an equilibrium liquid 484 composition resulting from partial melting that is roughly equal parts Fe and Mg to be energetically 485 favorable within the FeCO₃-MgCO₃-CaCO₃ ternary. Given the lack of natural samples, the 486 energetic and volumetric properties predicted by our simulations, and the measured Fe-contents 487 found in melting experiments on carbonated mantle materials, we find that ferrocarbonatites are 488 unlikely to have sufficient Fe-enrichment to stagnate in the mantle owing to their large relative 489 buoyancy.

490

5. Conclusions

491 We perform classical MD simulations of MgCO₃-bearing liquids using the empirical 492 potential model of Hurt and Wolf (2018) at mantle pressures and temperatures. We find that 493 MgCO₃ liquid assumes a novel topology characterized by four-fold coordination of Mg²⁺ with both the carbonate molecule and O²⁻. Such a structure is markedly different from that of the other 494 495 alkaline earth carbonate liquids in which the metal cation is in 6-fold coordination with the 496 carbonate molecule and 7-8 fold coordination with oxygen atoms. The liquid structure and 497 resulting thermodynamic properties of MgCO₃ are found to be more similar to that of Li₂CO₃ than 498 other alkaline earth carbonate liquids, resulting in a dramatically more buoyant and compressible 499 liquid than previous studies have suggested. The voluminous structure and subsequent low density 500 of MgCO₃-rich liquids bolsters the density contrast of MgCO₃-bearing carbonatite melts in the 501 mantle, increasing their already significant rates of ascent and preventing even relatively dense 502 ferrocarbonatite melts from achieving a density crossover with the surrounding mantle.

503 The simulations have further suggested that MgCO₃ molar volume, thermal expansivity, 504 compressibility, enthalpy and heat capacity do not mix ideally with (Ca, Sr, Ba)CO₃ liquids. Non-505 ideal mixing of volumetric properties stems from localized metal cation ordering around the Mg²⁺ ion. We find that it is energetically favorable to surround tetrahedral Mg-CO₃ complexes with 506 507 octahedral Ca-, Sr-, Ba-carbonate complexes. This effect gradually diminishes with pressure, 508 where MgCO₃ molar volume is found to mix near-ideally with CaCO₃-SrCO₃-BaCO₃ liquids by 509 12 GPa. On the other hand, non-ideal mixing of enthalpy in MgCO₃-bearing systems persists even 510 at high pressure.

511 The novel topology of MgCO₃ liquid indicated by our simulations opens questions on 512 another important component in mantle-derived carbonatite melts: FeCO₃. Given the similar ionic

radii of Fe²⁺ and Mg²⁺, it's possible that Fe²⁺ also assumes a 4-fold coordination with $(CO_3)^{2-}$ and 513 514 O²⁻ in FeCO₃ liquid, producing a similar liquid structure to MgCO₃; the result would be a FeCO₃ 515 component that is significantly less dense than what has been expected by other studies (e.g. Kang 516 and Schmidt 2016, 2017). Such a voluminous FeCO₃ component would inhibit ferrocarbonatite 517 melts generated within subducting slabs from ever stagnating or sinking in the mantle, due to their 518 low relative densities. Multi-angle energy dispersive X-ray diffraction experiments on MgCO₃-519 and FeCO₃-bearing melts are needed in order to verify the liquid structure of these important 520 carbonate components.

521

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Table 1. The thermodynamic properties of MgCO₃ liquid as derived from the simulations. These are fitted parameters of both a 3rd and 4th order Birch-Murnaghan EOS (Eqs. A1-A4) that have been fitted to the simulations of MgCO₃ liquid from 773 to 2000 K up to 20 GPa by the least squares method. $V_{Tr,0}$ is the molar volume at 1100 K and 1 bar; α is the thermal expansion; $K_{Tr,0}$ is the bulk modulus at 1 bar and 1100 K and $\delta K_0 / \delta T$ is its temperature dependence; *K*' and *K*'' are first- and second-pressure derivatives of the bulk modulus. The root mean square of the pressure residuals of the 4th order fit is 0.02 GPa and of the volume residuals is 0.27 cm³/mol.

Property	Best F	Fit (1σ)
	3 rd order EOS	4 th order EOS
RMS Press. residual (GPa)	0.11	0.02
V _{Tr,0} (cm ³ /mol)	46.0(2)	44.21(9)
$\alpha (10^{-4} \cdot K^{-1})$	0.004(13)	1.07(14)
$K_{Tr,0}$ (GPa)	0.9(20)	6.15(15)
$\delta K_0 / \delta T$ (10 ⁻³ · GPa/K)	-0.1(18)	-1.37(35)
Κ'	33(75)	2.41(1)
K"		3.16(1)

Table 2. Simulated densities of (Ca, Mg)CO₃ binary liquids approximating compositions of carbonatite melts from phase equilibrium experiments of carbonated mantle (Dalton and Wood 1993, Dalton and Presnall 1998, Hammouda 2003, and Ghosh et al. 2009). The pressures and temperatures of the simulations reflect those of the phase equilibrium experiments. Corrected density adjusts the simulated density to account for small concentrations of FeCO₃ and SiO₂ present in the liquid produced in the original phase equilibrium experiments.

MgCO ₃	Press.	Temp.	Gram	Sim.Vol.	Sim. p	Corrected p
(mol%)	(GPa)	(K)	formula wt.	(cm ³ /mol)	(g/cm^3)	(g/cm^3)
		Da	lton and Wood (1	993)		
13.6	1.5	1423	98.22	39.19	2.506	2.507
17.4	1.5	1423	97.69	39.23	2.490	2.497
22.5	2.2	1523	96.95	38.28	2.532	2.545
28.7	2.5	1523	96.05	37.76	2.544	2.558
31.6	2.8	1573	95.63	37.47	2.552	2.573
35.1	3.0	1573	95.11	37.09	2.564	2.585
		Dalt	on and Presnall (1998)		
40.9	3.0	1518	93.64	36.59	2.559	2.561
44.0	3.5	1543	93.15	35.94	2.592	2.592
45.8	4.0	1563	92.86	35.44	2.620	2.618
49.2	5.0	1603	92.33	34.54	2.673	2.669
50.9	6.0	1653	92.05	33.77	2.726	2.717
55.5	7.0	1703	91.33	33.03	2.765	2.756
			Hammouda (200	3)		
10.5	6.0	1523	98.43	34.91	2.820	2.871
10.9	7.0	1273	98.36	33.61	2.927	2.957
10.8	10.0	1373	98.38	32.23	3.052	3.086
14.1	10.0	1473	97.86	32.47	3.014	3.068
			Ghosh et al (2009	9)		
87.1	12.5	1873	86.35	29.53	2.925	3.071
86.8	13.5	1823	86.40	29.12	2.967	3.096
61.1	15.0	1873	90.45	29.62	3.053	3.133
76.1	17.5	1833	88.08	28.30	3.112	3.182

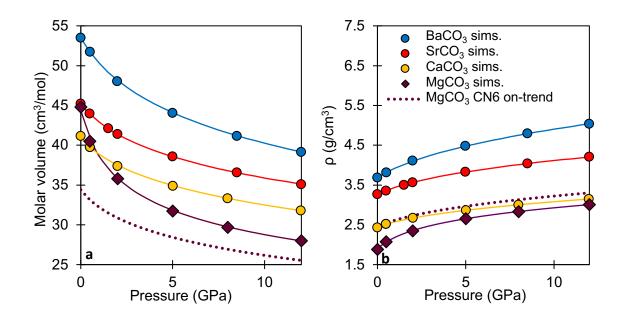


Figure 1. Compression curve of MgCO₃ liquid strongly deviates from expectations based on the other alkaline-earth carbonates. Panel a shows simulated molar volumes at 1100 K as a function of pressure for MgCO₃, CaCO₃, SrCO₃ and BaCO₃ liquids. Points show simulation results, while solid lines represent best-fit Birch-Murnaghan EOS for each liquid (3rd order for CaCO₃, SrCO₃ and BaCO₃ and 4th order for MgCO₃). The dotted line represents the expected compression curve of MgCO₃ if its volumetric properties were in-line with periodic trends observed among the other alkaline earth carbonates (i.e. if Mg-C coordination was 6-fold rather than 4). As evident, CaCO₃, SrCO₃ and BaCO₃ form nearly parallel curves that increase systematically from Ca to Ba, while MgCO₃ liquid deviates dramatically, especially at low pressure. Panel b shows the same compression curves in terms of density.

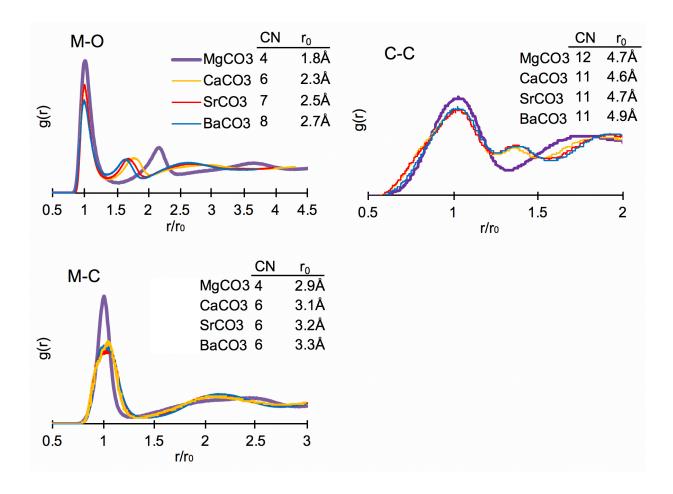


Figure 2. Liquid structure of MgCO₃ is distinct from CaCO₃, SrCO₃ and BaCO₃. Pair distribution functions are calculated from liquid simulations at 1 bar and 1100 K for atomic pairs between carbon (C), oxygen (O) and metal cations (M: Mg, Ca, Sr, Ba). The M-O, C-C and M-C pairs shown in panels a, b and c, respectively. For visual comparison, interatomic separation distance (r) has been normalized to the position of the first peak (r_0). Average coordination number (CN) has been calculated for each pair in the embedded tables.

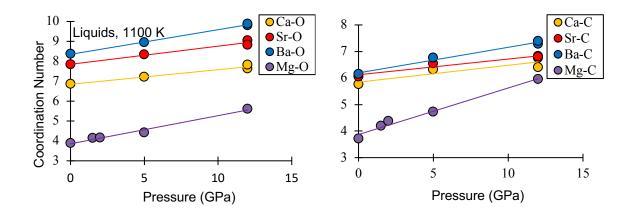


Figure 3. The structure of MgCO₃ liquid becomes progressively more like (Ca, Sr, Ba)CO₃ liquids at high pressure. The simulated evolution of coordination for M-O and M-C pairs (in panels a and b), is shown for pure MgCO₃, CaCO₃, SrCO₃ and BaCO₃ liquids as a function of pressure up to 12 GPa. In MgCO₃, the coordination of M-O and M-C pairs starts at 4 at 1 bar and approaches ~6 at 12 GPa (while other alkaline earths' coordination evolution is more modest, rising 2-4 times slower).

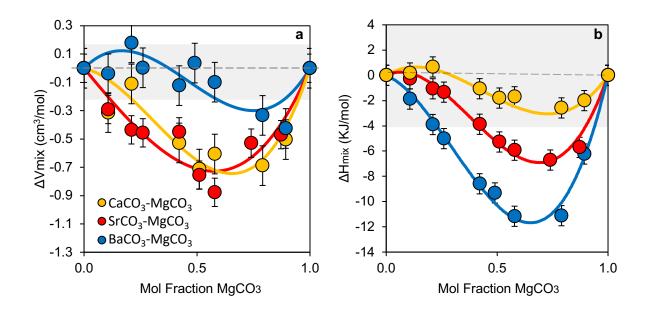


Figure 4. Non-ideal mixing dominates volumetric and enthalpic properties in MgCO₃-bearing systems at 1 bar. The ΔV_{mix} (a) and ΔH_{mix} (b) of MgCO₃-bearing binary liquids at 1100 K and 1 bar are shown, with non-zero values indicating non-ideal mixing. The shaded regions reflect the range of ΔV_{mix} and ΔH_{mix} for the ideal mixing of (Ca, Sr, Ba)CO₃ liquids calculated by Hurt and Wolf (2018). Both ΔV_{mix} and ΔH_{mix} are, for the most part, negative and reach a minimum at MgCO₃ molar concentrations of 0.6 to 0.8. Points show simulation results and lines are the best-fit sub-regular solutions (Eq. 3).

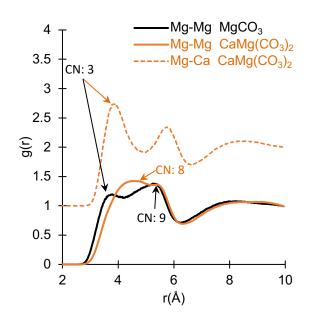


Figure 5. Mg-Mg pdf curve in pure MgCO₃ liquid are significantly altered upon mixing due to Mg^{2+} vacating the inner coordination shell in favor of Ca^{2+} , Sr^{2+} or Ba^{2+} ions. The pdf curves are generated from simulations of pure MgCO₃ liquid and a mixture, $CaMg(CO_3)_2$, at 1100 K and 1 bar. The inner coordination shell (CN=3) in pure MgCO₃ disappears in the mixture and is filled with Ca^{2+} . This same effect occurs in MgCO₃ mixtures with the other alkaline earth cations (Sr²⁺ and Ba²⁺). This likely results from an energetic penalty for Mg²⁺ to be surrounded by other Mg²⁺ ions; the preferential ordering observed here results from surrounding the Mg-CO₃ tetrahedral complexes with the (Ca, Sr, Ba)²⁺ octahedral complexes, and might be favored as a mechanism to handle cation size mismatch that relaxes stresses in the atomic structure.

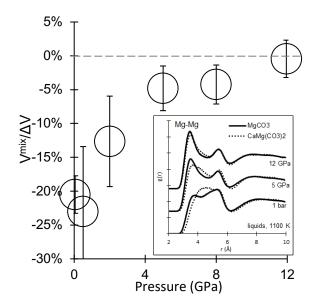


Figure 6. Mixing of molar volumes for CaCO₃ and MgCO₃ liquid approach ideality with increasing pressure due to a progressively diminishing cation ordering effect. V_{mix} is shown as a function of pressure normalized to ΔV_{Ca-Mg} (which is the difference between pure CaCO₃ and MgCO₃ liquid at the given pressure). As pressure increase, the magnitude of non-ideality approaches 0. The figure inset shows corresponding liquid Mg-Mg pdf curves of pure MgCO₃ (solid lines) and CaMg(CO₃)₂ (dotted lines) at 1 bar, 2 and 12 GPa. The disparity between the two curves reflects the degree of cation ordering in the liquid. As pressure increase, the mixed CaMg(CO₃)₂ curve conforms to that of pure MgCO₃.

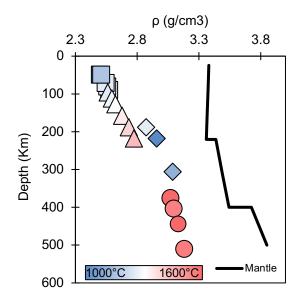


Figure 7. The densities of primary carbonatite melts maintain a high negative density contrast with the surrounding mantle. The simulated densities of MgCO₃-CaCO₃ binary liquids representing primary carbonatite melts are corrected for minor FeCO₃ and SiO₂ concentration and are shown as a function of depth. The simulation temperatures, pressures and compositions are based on the phase equilibrium experiments of Dalton and Wood (1993) (squares), Dalton and Presnall (1998) (triangles), Hammouda (2003) (diamonds) and Ghosh et al. (2009) (circles). Coloring reflects simulation temperature. Mantle density as a function of depth is taken from PREM (Dziewonski and Anderson, 1981).

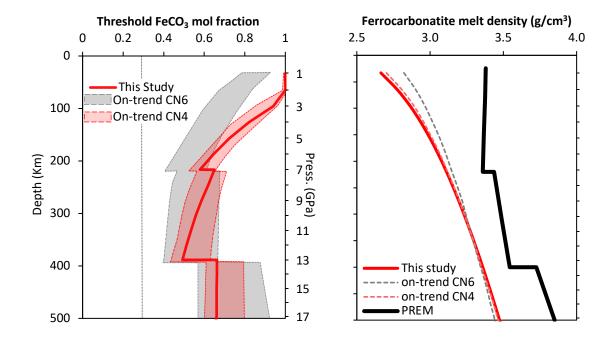


Figure 8. Panel a shows the threshold mole percent FeCO₃ necessary for a hypothetical ferrocarbonatite (with composition (Fe_x, Mg_{0.5'(1-x)}, Ca_{0.5'(1-x)})CO₃) to achieve neutral buoyancy in the mantle as a function of depth. Three different estimates are depicted: the red line shows results from this study while the gray- and red-shaded regions represent estimated volumetric properties from Hurt (2018), which assume that both Mg²⁺ and Fe²⁺ are in 6-fold and 4-fold coordination with carbonate respectively. The dotted vertical dotted line shows the iron concentration of the only natural, unambiguously-magmatic ferrocarbonatite found in the literature (Thompson et al. 2002). Panel b shows the estimated densities of this natural ferrocarbonatite composition (Fe_{0.26} Mg_{0.22} Ca_{0.52} CO₃) calculated for the three models (red lines are simulations from this study and the optimal on-trend CN6 and CN4 models from Hurt 2018 are shown in gray- and red-dashed lines) along with the density of PREM (Dziewonski and Anderson, 1981).

1

Appendices

2 A. Equations of state used to model simulation data

3 The simulated molar volumes of MgCO₃ liquid were used to fit the parameters of the EOS below
4 (adapted from Angel 2000):

5
$$P = 3K_{T,0}f_E(1+2f_E)^{\frac{5}{2}}\{1+c3f_E+c4f_E^2\}$$
 Eq. A. 1

6
$$c3 = \frac{3}{2}(K'-4)$$
, $c4 = \frac{3}{2}\left[K_{T,0}K'' + (K'-4)(K'-3) + \frac{35}{9}\right]$

where $K_{T,0}$ is the isothermal bulk modulus at 1 bar and simulation temperature (*T*), c3 and c4 are the 3rd and 4th order EOS terms, which are written in terms of the fitted parameters *K*' and *K*'' (the 1st and 2nd pressurederivatives of the bulk modulus). $K_{T,0}$ is a function of temperature according to:

10
$$K_{T,0}(T) = K_{Tr,0} + \frac{\delta K_0}{\delta T} (T - 1100 K) \text{ Eq. A. 2}$$

11 $\left(\frac{\delta K_0}{\delta T}\right)$ is the fitted 1 bar temperature dependence of the bulk modulus and $K_{T,\theta}$ is the fitted 1 bar bulk modulus 12 at the reference temperature. The Eulerian finite strain, $f_{E_{\tau}}$ is defined as:

13
$$f_E = \left[\binom{V_{T,0}}{V_{T,P}}^{2/3} - 1 \right] / 2 \quad \text{Eq. A. 3}$$

14 $V_{T,P}$ is the simulation volume at *T* and pressure (*P*). $V_{T,0}$ is the volume at 1 bar and *T* and can be calculated 15 from the fitted 1 bar volume at 1100 K ($V_{Tr,0}$) and the fitted thermal expansion (α). $V_{T,0}$ thus becomes:

16
$$V_{T,0} = V_{Tr,0} e^{\alpha(T-1100 K)}$$
 Eq. A.4

17 The parameter α is the fitted 1 bar thermal expansion coefficient, assumed to be temperature-independent.

18 **B.** Comparison of MgCO₃ to alkali carbonate liquids

MD simulations of MgCO₃ liquid indicate that, at ambient pressure (where MgCO₃ liquid is unstable) Mg²⁺ assumes a 4-fold coordination with both the carbonate molecule and oxygen much like alkali carbonate liquids, Li₂CO₃, NaCO₃ and K₂CO₃ (e.g. Roest et al. 2017; Kohara et al. 1999). This stands in stark contrast to CaCO₃-SrCO₃-BaCO₃ liquids in which metal cations are in 6-fold coordination with carbonate and 7- to 8-fold coordination with oxygen (Hurt and Wolf 2018). The similarity between MgCO₃ and alkali carbonate liquid structure is well-illustrated by a comparison of pdf curves for MgCO₃ liquid versus Li₂CO₃ liquid, as shown in Figure B.1 (Roest et al. 2017). This resemblance in liquid structure (and
physical properties) is not unprecedented; it's partially explained by the Mg-Li diagonal relationship, in
which the compounds of certain elemental pairs (Li-Mg, Be-Al, B-Si and C-P) that are diagonally adjacent
on the Periodic Table, share some properties in common (e.g., Feinstein 1984). Thus, it's expected that
liquid MgCO₃ volumetric properties should be similar to that of Li₂CO₃.

Alkali carbonates generally have a larger volume than their alkaline earth counterparts, not only because of the additional metal cation per formula unit but because the partial molar volume of $(CO_3)^{2-}$ is greater, ~38 cm³/mol versus 31 cm³/mol (Hurt and Lange 2019). The partial molar volume of $(CO_3)^{2-}$ is greater in alkaline earth carbonate liquids because of changes in network topology, not because the size of the carbonate molecule actually increases (See Hurt and Lange 2019 for a full discussion).

35 *C. Structure and thermodynamics of MgCO*₃-carbonate mixtures

The thermodynamic and structural properties of MgCO₃-bearing liquids is explored through molecular dynamic simulations. The conditions and basic results for each run are summarized in Table C.1, including simulations of pure MgCO₃ as well as binary mixtures of MgCO₃ with CaCO₃, SrCO₃ and BaCO₃ liquids. The detailed results of binary mixtures and associated excess properties are reported in Table C.2, and the best-fit asymmetric sub-regular solution model parameters are presented in Table C.3.

The associated atomic structures of these compositional binary melts are presented in Figures C.1 and C.2. The liquid structures are shown for all pure endmember melts as well as 50:50 mixtures along each MgCO₃-bearing binary in Figure C.1. The complete set of pdfs for each liquid for M-O, M-C, and C-C pairs are shown for each of these liquids. To better understand the liquid structure with respect to carboncarbon pairs, the pdfs of MgCO₃, CaCO₃, and CaMg(CO₃)₂ liquid are shown in Figure C.2, along with a theoretical weighted average of the endmembers.

47 **D.** Description of density simulation corrections

The simulation results presented in the main text highlight the novel topology of MgCO₃ liquid and provide estimates of its thermodynamic properties and mixing behavior which is important for future integration of the MgCO₃ liquid component into thermodynamic modeling software such as MELTS or THERMOCALC. However, the pair potential model can also be directly applied to geologic systems by providing property estimates of natural carbonatite liquids at relevant temperatures and pressures. In order to estimate the density of primary carbonatite melts generated by low-degree partial melting of carbonated mantle, simulations are performed in the CaCO₃-MgCO₃ liquid system. The simulated liquid compositions, temperatures and pressures are based on the results of phase equilibrium experiments carbonated mantlepartial melting experiments.

57 To constrain the composition of carbonatite melts produced by partial melting of the carbonated mantle, 58 Dalton and Wood (1993) and Dalton and Presnall (1998) conducted phase equilibrium experiments on 59 carbonated peridotite systems at pressures ranging from 1.5 to 7 GPa and temperatures from 1423 to 1703 60 K. Hammouda (2003) and Ghosh et al. (2009) conducted similar experiments for carbonated eclogite in the 61 6-17.5 GPa pressure range. They produced carbonate melts with > 92 wt. % CaCO₃ and MgCO₃ for the peridotite-derived melts and > 88 wt. % for the eclogite-derived. The remaining wt. % is mostly FeCO₃ and 62 63 SiO₂. We estimate the densities of the primary carbonate melts of Dalton and Wood (1993), Dalton and Presnall (1998), Hammouda (2003) and Ghosh et al. (2009) by simplifying them to CaCO₃-MgCO₃ binary 64 compositions and simulating them at their respective temperature and pressure conditions. The simulated 65 CaCO₃-MgCO₃ melts range in density between 2.49 to 3.11 g/cm³ at 1.5 and 17.5 GPa respectively; the full 66 results are presented in Table 2. Since the minor $FeCO_3$ and SiO_2 components are omitted in these 67 68 simulations, the raw simulated densities have been corrected to account for their absence.

An approximate correction to the simulated densities may be made by assuming FeCO₃ has the same partial molar volume as the MgCO₃ melt component and integrating it into the density calculation. The density may be further corrected by accounting for the SiO₂ liquid component of the carbonatite melt by using the volumetric data from Lange and Carmichael (1987) and Kress et al. (1988) and assuming the SiO₂ volume mixes ideally into the carbonate melt. Table 2 gives the simulated melt composition, gram formula weight, molar volume and simulated density along with the FeCO₃-SiO₂-corrected density.

- 76 **Table C.1.** The complete simulation results are shown below. Total energy and Enthalpy are relative to
- their respective values for MgCO₃ at 1100 K and 0 GPa. 1 σ error on temperature is 2 K, 0.02 GPa for pressure, 0.1 cm³/mol for volume, 0.2 KJ for total energy and 0.9 KJ for enthalpy.

	Temp.	Press.	Mol. Vol.	Tot. Energy	Enthalpy
	(K)	(GPa)	(cm ³ /mol)	(KJ/mol)	(KJ/mol)
			MgCO ₃		
	773	0.00	43.10	-49	-49
	1100	0.00	44.48	0	0
	1499	0.00	46.46	61	61
	2000	0.00	47.98	136	136
	1099	0.49	40.52	-2	17
	900	1.99	35.39	-32	-
	1101	2.01	35.84	-2	70
	1500	2.02	36.85	58	132
	1869	2.98	35.89	112	219
	2273	3.58	35.97	170	299
	1101	4.99	31.74	2	160
	1101	5.00	31.74	3	161
	1501	5.00	32.66	61	224
	1100	8.00	29.71	8	245
	2274	8.05	31.96	173	430
	1099	12.01	28.03	17	353
	1500	11.99	28.67	72	416
	1099	15.01	27.14	25	432
	2000	14.95	28.41	147	571
	1100	19.99	26.07	39	560
Aol frac. N	AgCO ₃	CaCO ₃ -MgC	O3 Simulations	1	
0.11	1100	0.00	41.06	196	195
0.12	1424	1.49	39.19	239	298
0.14	1424	1.49	39.28	236	294
0.15	1423	1.54	39.23	232	291
0.15	1372	9.98	32.23	223	544
0.17	1423	1.46	39.14	227	284
	1523	2.21	38.28	235	319
0.2			12.06	10	43
	1099	0.00	43.06	43	45
0.2	1099 1101	$\begin{array}{c} 0.00\\ 0.00\end{array}$	43.00	4 <i>3</i> 174	173
0.2 0.21					
0.2 0.21 0.21	1101	0.00	41.62 38.35	174	173
0.2 0.21 0.21 0.22	1101 1523	0.00 2.20	41.62	174 230	173 314
0.2 0.21 0.21 0.22 0.22	1101 1523 1523	0.00 2.20 2.20	41.62 38.35 38.35	174 230 230	173 314 314
0.2 0.21 0.21 0.22 0.22 0.22	1101 1523 1523 1523	0.00 2.20 2.20 2.49	41.62 38.35 38.35 37.76	174 230 230 222	173 314 314 316
0.2 0.21 0.21 0.22 0.22 0.26 0.28	1101 1523 1523 1523 1573	0.00 2.20 2.20 2.49 2.83	41.62 38.35 38.35 37.76 37.47	174 230 230 222 224	173 314 314 316 329
0.2 0.21 0.21 0.22 0.22 0.22 0.26 0.28 0.29	1101 1523 1523 1523 1573 1573	0.00 2.20 2.20 2.49 2.83 2.52	41.62 38.35 38.35 37.76 37.47 37.70	174 230 230 222 224 216	173 314 314 316 329 310
$\begin{array}{c} 0.2 \\ 0.21 \\ 0.22 \\ 0.22 \\ 0.26 \\ 0.28 \\ 0.29 \\ 0.32 \end{array}$	1101 1523 1523 1523 1573 1525 1573	0.00 2.20 2.20 2.49 2.83 2.52 2.84	41.62 38.35 38.35 37.76 37.47 37.70 37.33	174 230 230 222 224 216 217	173 314 314 316 329 310 322
$\begin{array}{c} 0.2 \\ 0.21 \\ 0.21 \\ 0.22 \\ 0.22 \\ 0.26 \\ 0.28 \\ 0.29 \\ 0.32 \\ 0.32 \end{array}$	1101 1523 1523 1523 1573 1525 1573 1572	0.00 2.20 2.20 2.49 2.83 2.52 2.84 3.03	41.62 38.35 38.35 37.76 37.47 37.70 37.33 37.09	174 230 230 222 224 216 217 216	173 314 314 316 329 310 322 328
$\begin{array}{c} 0.2\\ 0.21\\ 0.21\\ 0.22\\ 0.22\\ 0.26\\ 0.28\\ 0.29\\ 0.32\\ 0.32\\ 0.35\\ \end{array}$	1101 1523 1523 1523 1573 1575 1573 1572 1518	0.00 2.20 2.20 2.49 2.83 2.52 2.84 3.03 2.98 3.01	41.62 38.35 38.35 37.76 37.47 37.70 37.33 37.09 36.71 36.99	174 230 230 222 224 216 217 216 200 208	173 314 314 316 329 310 322 328 309 319
$\begin{array}{c} 0.2\\ 0.21\\ 0.21\\ 0.22\\ 0.22\\ 0.26\\ 0.28\\ 0.29\\ 0.32\\ 0.32\\ 0.35\\ 0.35\\ \end{array}$	1101 1523 1523 1523 1573 1575 1573 1572 1518 1573	0.00 2.20 2.20 2.49 2.83 2.52 2.84 3.03 2.98	41.62 38.35 38.35 37.76 37.47 37.70 37.33 37.09 36.71	174 230 230 222 224 216 217 216 200	173 314 314 316 329 310 322 328 309

0.42	1100	-0.01	41.94	126	125
0.44	1543	3.50	35.94	184	309
0.46	1563	4.00	35.44	183	324
0.51	1100	-0.01	42.06	105	104
0.51	1100	-0.01	42.06	105	104
0.51	1998	0.01	46.60	249	249
0.51	1100	0.50	40.05	104	124
0.51	1100	2.00	36.41	102	174
0.51	1799	2.98	37.15	207	317
0.51	1800	2.99	37.16	207	318
0.51	1573	4.98	34.42	176	347
0.51	1099	4.99	33.13	103	268
0.51	1603	5.02	34.54	181	354
0.51	1655	5.98	33.77	184	386
0.51	1100	8.00	31.33	108	358
0.51	1100	12.00	29.86	117	475
0.56	1703	7.03	33.03	183	414
0.58	1102	0.00	42.41	91	90
0.61	1873	14.97	29.62	205	648
0.76	1834	17.48	28.30	174	668
0.79	1099	0.00	43.06	43	43
0.87	1873	12.48	29.53	150	518
0.87	1821	13.54	29.12	145	538
		SrCO ₃ -MgCO	3 Simulations		
0.11	1100	0.01	44.89	291	291
0.21	1100	0.00	44.66	257	256
0.26	1101	-0.01	44.60	240	239
0.42	1099	0.01	44.48	185	185
0.49	1100	11.99	31.51	172	549
0.51	1100	0.00	44.11	154	154
0.51	2000	0.03	48.71	299	300
0.58	1100	-0.01	43.93	131	131
0.74	1100	0.01	44.15	78	78
0.87	1100	0.00	44.11	36	36
		BaCO ₃ -MgCO	D3 Simulations		
0.11	1100	0.00	52.52	422	421
0.26	1101	0.00	51.18	346	346
0.42	1099	0.01	49.59	266	266
0.49	2000	-0.02	53.57	373	372
0.49	1099	12.00	33.60	247	650
0.58	1100	0.00	48.19	189	188
0.79	1099	0.00	46.04	89	88
0.89	1100	0.01	45.02	45	45

81 Table C.2: The volume and enthalpy results of MgCO3-bearing binary simulations at 1100 K and 1 bar	are are
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82 shown below. Ideal volumes and enthalpies and ΔV_{mix} and ΔH_{mix} calculated according to Eq. 7 and 8 are

also listed.

Comp.	Mol. Vol.	Ideal Vol.	ΔV_{mix}	Enthalpy	Ideal H	ΔH_{mix}
Mol%MgCO ₃	(cm^3/mol)	(cm^3/mol)	(cm^3/mol)	(KJ/mol)	(KJ/mol)	(KJ/mol)
morringeog		· · · · · · · · · · · · · · · · · · ·	O_3-MgCO_3	(10,1101)	(10,1101)	(10,1101)
0.0	41.00	41.00	0.00	218.9	218.9	0.0
10.7	41.06	41.37	-0.31	195.7	195.5	0.2
21.0	41.62	41.73	-0.11	173.6	173.0	0.7
42.1	41.94	42.47	-0.53	125.6	126.7	-1.1
51.0	42.06	42.78	-0.71	105.4	107.2	-1.8
57.9	42.41	43.01	-0.61	90.5	92.2	-1.7
79.0	43.06	43.75	-0.69	43.4	46.0	-2.6
89.3	43.60	44.11	-0.50	21.5	23.5	-2.0
100.0	44.48	44.48	0.00	0.0	0.0	0.0
		SrCO	D ₃ -MgCO ₃			
0.0	45.26	45.26	0.00	326.0	326.0	0.0
10.7	44.89	45.18	-0.29	290.8	291.1	-0.3
21.0	44.66	45.10	-0.44	256.5	257.6	-1.1
26.0	44.60	45.06	-0.46	239.9	241.2	-1.3
42.1	44.48	44.93	-0.45	184.8	188.7	-3.9
51.0	44.11	44.86	-0.76	154.4	159.7	-5.3
57.9	43.93	44.81	-0.88	131.4	137.3	-5.9
74.0	44.15	44.68	-0.53	78.1	84.8	-6.7
87.2	44.11	44.58	-0.47	36.1	41.8	-5.7
100.0	44.48	44.48	0.00	0.0	0.0	0.0
			03-MgCO 3			
0.0	53.53	53.53	0.00	475.0	475.0	0.0
10.7	52.52	52.56	-0.04	422.2	424.1	-1.9
21.0	51.81	51.63	0.18	371.4	375.3	-3.9
26.0	51.18	51.17	0.00	346.4	351.4	-5.0
42.1	49.59	49.71	-0.12	266.3	274.9	-8.6
49.0	49.13	49.09	0.04	233.0	242.3	-9.3
57.9	48.19	48.29	-0.10	188.9	200.1	-11.2
79.0	46.04	46.38	-0.33	88.6	99.7	-11.1
89.3	45.02	45.45	-0.43	44.7	50.9	-6.2
100.0	44.48	44.48	0.00	0.0	0.0	0.0

	ΔV_{mix}	(cm ³ /mol)	ΔH_{mix}	(KJ/m89)
	W	Δw	W	∆w
CaCO ₃ -MgCO ₃	-2.91	-6.78	-7.33	-35.51
SrCO ₃ -MgCO ₃	-3.625	-3.77	-22.90	-62,56
BaCO ₃ -MgCO ₃	-1.345	-6.47	-40.20	-62 <u>5</u> 6 -92.95

Table C.3: Gives the coefficients fitted to Eq. 1 to describe both ΔV_{mix} and ΔH_{mix} calculated by our 86 simulations of liquids in each of the three binaries at 1100 K and 1 bar. 87

92

Figure B.1. The liquid structure of MgCO₃ is similar to that of Li₂CO₃. The liquid pdf curves of
MgCO₃ from our simulations (solid lines) are compared to those of Li₂CO₃ generated by FPMD
simulations from Roest et al. 2017 (dotted lines) at 1200 K and 15 kbar. CaCO₃ liquid at 1200 K
and 15 kbar is shown for comparison (grey dashed line). MgCO₃ liquid is more like Li₂CO₃ than
CaCO₃.

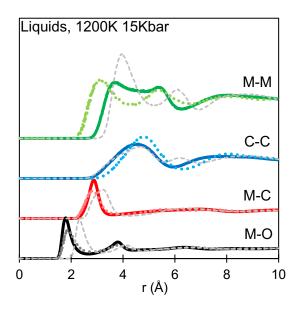
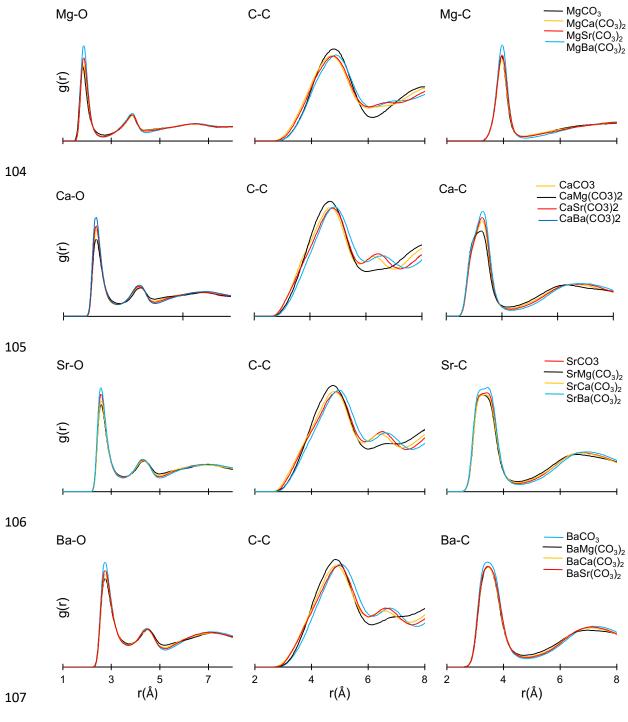
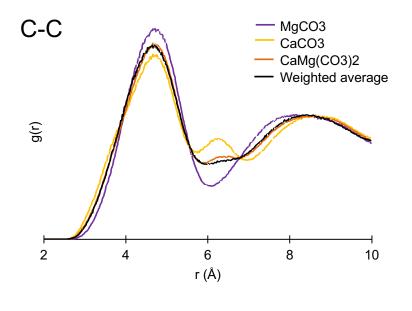


Figure C.1: PDF curves for M-O, C-C and M-C pairs in pure liquids are unchanged by mixing. The coordination of M^{2+} with oxygen atoms and carbonate groups is unaffected by the composition of the liquid. The carbon-carbon pdf curve changes when Mg^{2+} is introduced into the mix but remains a weighted average of the two endmember components (Appendix. 4).



109 Figure C.2: Structural packing of carbonate molecules varies systematically for Mg-bearing 110 carbonate mixtures. Shown below are pdf curves from simulations of C-C pairs (at 1100 K and 0 bar) for 111 pure MgCO₃ liquid (purple line), pure CaCO₃ (yellow line) liquid and an intermediate binary composition, 112 CaMg(CO₃)₂ (orange line). The black line is the C-C pdf curve of CaMg(CO₃)₂ calculated from the 113 composition-weighted average of the two end members which nearly perfectly predicts the simulated 114 structure; this demonstrates that carbonate packing is not responsible for deviations from ideality in CaCO₃-115 MgCO₃ mixtures.



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