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# Enhancing daily precipitation reconstruction: An improved version of the reddPrec R package

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## Abstract

Reconstructing high-quality daily precipitation series is essential for climate studies, hydrological modeling, and environmental applications. This work presents a new version of reddPrec, a versatile and flexible R package designed to reconstruct precipitation datasets through standard quality control, gap-filling, and grid creation procedures. The update introduces greater flexibility in spatial modeling, inclusion of dynamic covariates, and new modules for enhanced quality control and homogenization. Daily precipitation can now be predicted through machine learning approaches within a user-friendly framework, allowing users to select modeling approaches and customize settings. We demonstrate its capabilities through case studies in Switzerland and Spain, evaluating improvements in reconstruction accuracy, quality control, and homogenization. Enhanced quality control and homogenization procedures were specifically validated to ensure reliable adjustment and consistency of precipitation series. Overall, reddPrec provides a comprehensive and reliable tool for reconstructing precipitation series, supporting the creation of high-quality datasets for climate research and related fields.

*Keywords:* reddPrec, Daily precipitation, Quality control, Missing values, Homogenization, Grid

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## 1. Introduction

Precipitation data is essential for advancing climate science and supporting a wide range of research and operational applications, including climate modeling, hydrological forecasting, water resource management, ecosystem monitoring, and agriculture. Among the various sources of precipitation measurements, time series recorded by weather stations are considered the most accurate (Tapiador et al., 2012). However, these records often suffer from limitations such as missing values and inhomogeneities, which can compromise their reliability (Venema et al., 2020; Cheng et al., 2024). These challenges are further exacerbated in regions where station networks are sparse or unevenly distributed, often due to economic or geographical constraints (Hunziker et al., 2017, 2018; Bliefernicht et al., 2022). To address these issues, numerous methods have been developed to reconstruct precipitation data, aiming to complete and enhance the quality of observations in both time and space. The resulting reconstructions may take the form of station-level time series (Vicente-Serrano et al., 2010; Tang et al., 2020, 2021; Huerta et al., 2024) or gridded datasets (Serrano-Notivoli et al., 2017a; Aybar et al., 2020; Tang et al., 2022) across a range of spatial and temporal scales. In this context, the `reddPrec` R package (Serrano-Notivoli et al., 2017b) was developed to facilitate the reconstruction of daily precipitation time series.

The `reddPrec` package provides an integrated framework for reconstructing daily precipitation time series from weather station data. It offers a modular function suite that guides users through essential preprocessing steps—including quality control, gap-filling, and grid creation—to ensure that reconstructed datasets are accurate and consistent. Designed with flexibility in mind, `reddPrec` has been successfully applied in regions with dense station networks and moderately complex terrain (Serrano-Notivoli et al., 2017a; Navarro et al., 2020; Škrk et al., 2021; Bessaklia et al., 2021; Centella-Artola et al., 2023; da Silva et al., 2024; Montaña-Caro et al., 2024; Serrano-Notivoli et al., 2024), supporting a wide range of climatological and hydrological studies. However, the original version of the package had several limitations, including a basic quality control protocol, limited support for advanced spatio-temporal models, inflexible handling of dynamic covariates, and the absence of a homogenization tool. These shortcomings limited its effectiveness in more demanding settings, such as areas with sparse station coverage, complex topography, or studies requiring high temporal consistency in the data.

To address the limitations of the original implementation, the new version of `reddPrec` introduces a suite of methodological improvements aimed at increasing robustness, scalability, and flexibility in the reconstruction process. Key updates include support for user-defined machine learning models in the computation pipeline (Hothorn, 2024), integration of dynamic covariates for spatio-temporal learning (Hu et al., 2019; Ochoa-Rodriguez et al., 2019; Kossieris et al., 2024), an

41 enhanced quality control module for detecting and handling systematic errors in  
 42 time series input data (Hunziker et al., 2017, 2018), and a homogenization frame-  
 43 work (Squintu et al., 2018; Brugnara et al., 2019) tailored for daily precipitation  
 44 series. These developments enhance the applicability of the package in data-sparse  
 45 environments and heterogeneous terrain while maintaining compatibility with ex-  
 46 isting workflows for quality control, gap filling, and grid creation.

47 This paper presents the updated version of `reddPrec`, detailing its new fea-  
 48 tures and demonstrating their functionality through real-case experiments. Sec-  
 49 tion 1 briefly describes the technical basis and core functions of `reddPrec`. Section  
 50 2 describes the major improvements in the package, focusing on the expanded  
 51 modeling capabilities, enhanced quality control routines, and the integration of  
 52 a homogenization framework. Section 3 presents application experiments that  
 53 showcase the package’s performance and impact in reconstructing precipitation.  
 54 Finally, we discuss potential applications and future developments for further ex-  
 55 tending the utility of `reddPrec` in climate and hydrological research.

## 56 2. Overview of the `reddPrec` package

57 The `reddPrec` package is built around the concept of reference values (RVs),  
 58 which are local precipitation estimates generated for each day and location. These  
 59 values are derived from nearby weather stations, incorporating topographical co-  
 60 variates to capture the spatial and temporal variability of precipitation. Essen-  
 61 tially, RVs provide highly flexible, localized models that reflect the precipitation  
 62 conditions specific to each station and its surroundings.

63 RVs are computed through a combination of classification and regression func-  
 64 tions. Let  $\mathbf{X} = [x_1, x_2, \dots, x_N]$  represent the vector of covariates, where  $N$  denotes  
 65 the number of covariates, and the process is structured as follows:

- 66 1. Classification function: This function classifies each day as either dry or wet  
 67 based on predicted probabilities:

$$Y_{\text{class}}(\mathbf{X}) = \begin{cases} 1 & \text{if } f_c(\mathbf{X}) + \varepsilon_c \geq 0.5 \\ 0 & \text{if } f_c(\mathbf{X}) + \varepsilon_c < 0.5 \end{cases} \quad (1)$$

68 Where  $f_c(\mathbf{X}) \in [0, 1]$  is the predicted probability of a wet day, and  $\varepsilon_c$  is the  
 69 classification model error.

- 70 2. Regression function: This function estimates the amount of precipitation  
 71 for a given day:

$$Y_{\text{reg}}(\mathbf{X}) = f_r(\mathbf{X}) + \varepsilon_r \quad (2)$$

72 Where  $f_r(\mathbf{X}) \in \mathbb{R}_{\geq 0}$  is the predicted precipitation amount, and  $\varepsilon_r$  is the  
 73 regression error term.

74 3. RV: The final precipitation estimate is obtained by combining the dry/wet  
75 classification and the regression output:

$$RV(\mathbf{X}) = Y_{\text{class}}(\mathbf{X}) \cdot Y_{\text{reg}}(\mathbf{X}) \quad (3)$$

76 In the original version of reddPrec, RVs were constructed using data from 10  
77 nearby stations, with  $N = 3$  covariates (latitude, longitude, and elevation). Both  
78 the classification ( $f_c$ ) and regression ( $f_r$ ) models were based on generalized linear  
79 models (glm).

80 Once RVs are generated, they are utilized in the core functions of the package  
81 to:

- 82 1. Apply quality control to raw data using the `qcPrec()` function.
- 83 2. Fill missing values in time series through the `gapFilling()` function.
- 84 3. Create gridded precipitation datasets using the `gridPcp()` function, where  
85 each grid point is treated as an individual station.

### 86 3. Updates and new features

87 The latest version of reddPrec includes various additions aimed at increasing  
88 model flexibility and the use of dynamic covariates in core functions. It also  
89 provides new functions on quality control and homogenization of precipitation  
90 time series. The updates are outlined below.

#### 91 3.1. Flexibility on model foundation and dynamic covariates on RVs

92 A key parameter in the computation of RVs is the choice of the statistical  
93 model. While the previous version of reddPrec employed glm by default, this  
94 approach proved effective primarily in dense station networks and moderately  
95 challenging topographies. In regions characterized by sparse station coverage  
96 or complex terrain, the glm approach may lead to suboptimal reconstructions  
97 (Serrano-Notivoli and Tejedor, 2021).

98 To address this limitation, the updated version introduces a flexible modeling  
99 framework through the new `model_fun` argument, which allows users to spec-  
100 ify a custom machine learning model for RV computation. This enhancement  
101 broadens the applicability of reddPrec to diverse climatic and geographic con-  
102 texts by enabling a wide array of classification and regression models available  
103 in the R ecosystem. While the original glm-based method remains accessible via  
104 `learner_glm()`, the package now includes additional built-in options such as sup-  
105 port vector machines (`learner_svm()`), random forests (`learner_rf()`), extreme  
106 gradient boosting (`learner_xgboost()`), and neural networks (`learner_nn()`).  
107 Users can also define their own models by following the established structure of

108 these learner functions and tailoring the modeling parameters to their specific  
109 data and objectives.

110 In addition to expanding modeling flexibility, the new version of reddPrec in-  
111 troduces support for dynamic covariates (predictors that change their values each  
112 time step) in RVs construction through the `dynam_cov` argument. Previously, the  
113 package supported only static covariates, limiting its adaptability to real-time or  
114 temporally varying predictors. The new implementation allows the incorporation  
115 of dynamic variables, such as radar and satellite-based products, or outputs from  
116 atmospheric models, which can vary over time. When used alongside static co-  
117 variates, these dynamic inputs may enable more context-sensitive and accurate  
118 reconstructions, particularly in environments where precipitation processes are  
119 influenced by rapidly changing atmospheric conditions and with scarce observed  
120 data.

### 121 *3.2. Enhanced quality control*

122 Accurate precipitation reconstruction requires input data to be free from sig-  
123 nificant errors, inconsistencies, or outliers. In the previous version of reddPrec,  
124 quality control was limited to standard spatial-based routines on RVs, which pri-  
125 marily focused on detecting isolated suspicious values. The updated version, how-  
126 ever, significantly complements this aspect by introducing a comprehensive and  
127 modular quality control system, called enhanced quality control.

128 The enhanced quality control approach addresses recurring data quality issues  
129 that may go undetected by standard methods. Originally developed by Hunziker  
130 et al. (2017, 2018), this technique comprises visual tests designed to allow users  
131 to remove or flag problematic periods in a time series. In the new reddPrec, these  
132 tests have been automated by introducing a classification scheme that evaluates  
133 the overall quality of each station. Rather than flagging isolated time series peri-  
134 ods, stations are categorized according to their test levels. The enhanced quality  
135 control tests include:

- 136 • **Truncation:** Truncation is identified when heavy precipitation episodes are  
137 systematically reduced in frequency above a certain threshold. Here, the  
138 maximum boundary of a time series is determined as the daily precipitation's  
139 maximum moving window value. This boundary is then assessed based on  
140 its persistence over time:
  - 141 – Level 0: No truncation is detected (the maximum boundary persists  
142 for less than 3 years).
  - 143 – Level 1: A constant maximum boundary lasts between 3 and 5 years.
  - 144 – Level 2: The maximum boundary persists for more than 5 years.

- 145 • Small gaps: Small gaps refer to periods of unreported precipitation events  
146 that result in a reduction of frequency in low precipitation ranges. For this  
147 test, the total counts of precipitation values are computed in five ranges  
148 (0–1, 1–2, 2–3, 3–4, and 4–5 mm) for each year. The percentage of years  
149 with zero counts in these ranges is used to define the quality level:
  - 150 – Level 0: No small gaps (0%; years show at least one value in any range).
  - 151 – Level 1: Small gaps persist in at least 20% of consecutive years.
  - 152 – Level 2: Small gaps extend for more than 20% of consecutive years.
- 153 • Weekly cycle: The weekly cycle examines the occurrence of wet days to de-  
154 tect significant differences between days of the week. For each day, the prob-  
155 ability of precipitation is computed by dividing the number of wet days by  
156 the total count of records. A two-sided binomial test (using a 95% confidence  
157 level) then determines which days show significantly different probabilities:
  - 158 – Level 0: No atypical weekly cycle (similar precipitation probabilities  
159 across the week).
  - 160 – Level 1: At least two days present an atypical probability.
  - 161 – Level 2: More than two days show atypical probabilities, or one day  
162 exhibits an extremely different probability (a difference of more than  
163 10%).
- 164 • Precision and rounding patterns: This test assesses the consistency in deci-  
165 mal precision across the time series. First, the unique decimal patterns  
166 (sorted in descending order) are computed for each year. The mode, rep-  
167 resenting the most dominant decimal pattern, is then identified, and the  
168 proportion of the time series displaying this pattern is calculated:
  - 169 – Level 0: The dominant decimal pattern is present in more than 70% of  
170 the time series, indicating coherent precision.
  - 171 – Level 1: The dominant pattern is observed in between 50% and 70%  
172 of the data.
  - 173 – Level 2: There is no dominant decimal pattern, suggesting inconsisten-  
174 cies in reporting precision.

175 To improve reproducibility and transparency, these enhanced quality control  
176 procedures can be applied automatically using the `eqc_Ts()` function, which com-  
177 putes the quality levels as defined above. However, the thresholds and criteria can  
178 be customized based on user needs. In addition, summary plots for visual inspec-  
179 tion can be generated with the `eqc_Plot()` function, offering an intuitive overview

180 of the quality classification for each station as it was originally established in Hun-  
181 ziker et al. (2017, 2018).

182 By incorporating these robust quality control checks into the core functions of  
183 reddPrec, the updated package can ensure that reliable and consistent data are  
184 used in both the construction of RVs and the final precipitation estimates, thereby  
185 significantly enhancing the overall robustness and reproducibility of precipitation  
186 reconstructions.

### 187 3.3. Homogenization

188 Inhomogeneities (e.g., changes in station location, instrumentation, and obser-  
189 vation techniques) in precipitation time series can significantly affect statistical  
190 reconstructions and climatological analyses. To address this issue, the new ver-  
191 sion of reddPrec incorporates a homogenization framework specifically designed  
192 for daily precipitation datasets.

193 This procedure builds upon methods previously applied at global and continen-  
194 tal scales (Squintu et al., 2018; Brugnara et al., 2019), with targeted adaptations  
195 for daily precipitation. It combines the strengths of both relative and absolute  
196 approaches. Relative homogenization uses comparison between stations, while  
197 absolute homogenization uses individual station data to detect/adjust changes.  
198 Although absolute methods generally have lower detection power than relative  
199 ones (Venema et al., 2012), they serve as a critical fallback when relative test-  
200 ing is not feasible (such as in scarce-station networks). The entire procedure is  
201 implemented in the `hmgTs()` function and follows three main stages:

- 202 • Break detection: A combination of statistical tests and intercomparison  
203 of their results is used to detect breakpoints. Five univariate breakpoint  
204 tests are applied: Student’s t-test, Mann-Whitney, Buishand R, Pettitt,  
205 and the Standard Normal Homogeneity Test. In the relative approach,  
206 the algorithm identifies up to `neibs_max` well-correlated stations (corre-  
207 lation  $>$  `cor_neibs`) within a specified radius distance (`thres`). These  
208 nearby stations are used to construct difference series (target minus neigh-  
209 bor) under three temporal aggregations (annual, April–September, and Oc-  
210 tober–March) and two indices (PRCPTOT: total precipitation; R1mm: num-  
211 ber of wet days). A breakpoint is confirmed in a given year if it is detected  
212 in at least `perc_break` percent of the significant difference series (p-value  $<$   
213 0.05), with a tolerance of  $\pm 1$  year. If fewer than `neibs_min` nearby stations  
214 are found, the algorithm defaults to the absolute approach.
- 215 • Adjustment: A quantile-matching technique is used to adjust detected breaks,  
216 following the methodology of Squintu et al. (2018). While originally devel-  
217 oped for temperature data, the method has been adapted here for precipi-  
218 tation. Dry values (i.e., when `wet_day` = 0) are excluded from correction.



219 Wet values are transformed using both square root and logarithmic functions  
220 to approximate a normal distribution before adjustment. In the relative ap-  
221 proach, adjustment factors are computed from both the target and nearby  
222 series, assuming the post-break period is correct and applying corrections  
223 retrospectively. In the absolute case, adjustments rely solely on the target  
224 series and function similarly to a quantile mapping procedure.

- 225 • Quality control of adjustments: Because daily precipitation adjustments  
226 can influence the extreme tails of the distribution, corrected values can  
227 be constrained to not exceed a specified threshold difference (controlled  
228 via `apply_qc`) compared to the original data. This constraint can be ap-  
229 plied either to all values or to values above a defined precipitation threshold  
230 (`mm_apply_qc`). This step preserves the correction of extremes while pre-  
231 venting the generation of implausibly large values.

232 The homogenization workflow is fully integrated into the reddPrec ecosystem  
233 and operates seamlessly alongside the quality control and gap-filling modules. It  
234 is important to note that `hmgTs()` requires gap-filled input data and is mainly  
235 designed to correct inhomogeneities in the gap-filled data. The primary outputs  
236 are the homogenized time series and the corresponding break year information.  
237 By incorporating homogenization as a core module, the updated reddPrec en-  
238 sures temporal consistency in the input data, leading to more robust and reliable  
239 precipitation reconstructions.

## 240 4. Demonstration experiments

241 The following examples illustrate the performance and practical utility of the  
242 newly introduced features in the updated version of reddPrec. Each experiment  
243 showcases a specific enhancement and demonstrates its impact on daily precipi-  
244 tation reconstruction.

245 The primary dataset used for these experiments is the Swiss National Basic  
246 Climatological Network (Swiss NBCN), which provides some of the highest-quality  
247 and longest continuous precipitation records available (Begert, 2007; Fülle-  
248 mann et al., 2011). Its comprehensive spatial coverage, spanning a range of elevations  
249 and climate zones across Switzerland, makes it an ideal testbed for evaluating  
250 the methodological improvements. Additional datasets were also used in specific  
251 cases, as described in the respective experiment subsections.

### 252 4.1. Gap-filling and grid creation

253 This experiment tested the original reddPrec approach (glm) with two newly  
254 integrated model options (rf and xgboost) for gap-filling and grid creation. To fa-  
255 cilitate evaluation, we used two complementary metrics: the Matthews correlation

256 coefficient (mcc) and the refined index of agreement (dr). The mcc quantifies per-  
257 formance in dry/wet day classification, while dr assesses the accuracy of wet-day  
258 precipitation amounts (Chicco and Jurman, 2023; Willmott et al., 2012). These  
259 metrics together enable a comprehensive assessment of both categorical and con-  
260 tinuous components of precipitation modeling. The mcc and dr range from -1 (no  
261 agreement) to +1 (perfect agreement), with 0 indicating no better than random  
262 classification and no predictive skill, respectively (Appendix A).

263 For the gap-filling task, models were trained using 15 nearby stations with  
264 seven static covariates (latitude, longitude, elevation, and the first four principal  
265 components of multiple topographic features) over the period 2010–2015. Figure  
266 1 presents leave-one-out cross-validation results in terms of mcc and dr. Both  
267 rf and xgboost outperformed glm in overall performance, with the most notable  
268 improvements observed in the classification component. Specifically, mcc values  
269 showed an average increase of approximately 0.05, with most results clustered  
270 at 0.7 and 0.85. In contrast, improvements in dr were more modest (slightly  
271 below 0.05), and dispersion remained similar across models. Among the machine  
272 learning methods, rf exhibited better efficiency in this setting.

273 We then applied the same modeling approaches to grid creation. Unlike gap-  
274 filling, which estimates missing values at known station locations, grid creation  
275 generates precipitation values at ungauged points, treating each grid cell as a  
276 virtual station. In this experiment, we produced daily grids at  $0.009^\circ$  resolution  
277 ( $\approx 1$  km) over Switzerland for an extreme precipitation event that occurred be-  
278 tween July 24–28, 2014. The same set of static covariates was used, with two  
279 dynamic covariates added (Figure C.9): MODIS Aqua/Terra surface reflectance  
280 bands 1 and 2 (Vermote and Wolfe, 2021). Model outputs were evaluated against  
281 the RhydchprobD product (Frei and Isotta, 2019), comparing spatial patterns  
282 and point-to-grid values. RhydchprobD, constructed through conditional simula-  
283 tion based on Gaussian Random Fields, served as a robust benchmark due to its  
284 use of approximately ten times more observation stations, including those from  
285 neighboring countries.

286 Figures 2 and 3 show the modeled spatial fields and scatterplots of predicted  
287 versus observed values at the nearest station-grid point. A visual inspection sug-  
288 gested that all models captured spatial variability well compared to RhydchprobD.  
289 Wet to dry areas were realistically distributed, particularly on July 24 (north to  
290 south) and July 27 (southwest to northeast). The models also captured high  
291 rainfall magnitudes on July 26 (northeast area) and July 28 (southwest area), al-  
292 beit with lower intensity than RhydchprobD. While spatial patterns were similar  
293 across models, rf and xgboost better preserved observed values than glm (Figure  
294 3). The mcc and dr values were consistently close to 1 for both rf and xgboost.  
295 In contrast, the glm exhibited some days with values near 0.5, reflecting lower  
296 accuracy. These results demonstrate that rf and xgboost provide improved mod-

297 eling of both wet/dry day classification and wet-day precipitation amounts. The  
298 lower extremes of magnitude and slightly reduced coherence in our model outputs,  
299 compared to RhydchprobD, likely result from using fewer input stations. Nonethe-  
300 less, this highlights the robustness of reddPrec under limited data scenarios and  
301 complex terrain.

302 To further test the flexibility of the grid-creation workflow, we applied the  
303 same approaches to an extreme rainfall event in eastern Spain. This case focuses  
304 on the torrential rainfall that occurred in the province of Valencia on October  
305 29, 2024, which led to widespread flooding and infrastructure disruption. Unlike  
306 the Swiss case, the Spanish station network was denser but unevenly distributed,  
307 offering a realistic test of modeling under operational constraints (Figure C.10).  
308 Daily precipitation data from 443 AEMET (and other sources) stations were used  
309 to generate  $0.009^\circ$  ( $\approx 1$  km) grids over the affected region. The same dynamic  
310 covariates were included, alongside three static covariates (latitude, longitude,  
311 elevation). Model outputs were compared against a reference grid produced by  
312 the Consejo Superior de Investigaciones Científicas (CSIC), which used the same  
313 station data (Beguería et al., 2024) through a universal kriging approach using  
314 (rescaled) coordinates and elevation as covariates.

315 Figure 4 shows the spatial precipitation fields and the predicted vs. observed  
316 scatterplots for each model and the CSIC reference. Visually (Figure 4a), all mod-  
317 els reproduced the spatial structure and localized intensity of the event, capturing  
318 accumulations up to 700–800 mm (maximum 24h precipitation recorded was 772  
319 mm (AEMET, 2024)). In terms of value preservation (Figure 4b), all models  
320 performed well (mcc and dr near 1), though rf and xgboost once again achieved  
321 tighter fits (i.e., less spread). The CSIC product showed a nearly perfect match,  
322 likely due to its modeling technique (nugget effect), which aims to replicate exact  
323 station values.

324 These experiments highlight the versatility and strength of reddPrec in han-  
325 dling both gap-filling and grid creation tasks across diverse geographic and data  
326 scenarios. By integrating advanced machine learning techniques and dynamic  
327 covariates, the package can deliver accurate reconstructions even under diverse  
328 conditions, such as dense to sparse or uneven station networks and high-intensity  
329 precipitation events. While default model configurations already yield solid re-  
330 sults, performance can be improved through hyperparameter optimization, tai-  
331 lored covariate selection, and domain-specific user input. Overall, reddPrec offers  
332 a flexible and scalable precipitation modeling framework, supporting research ap-  
333 plications and operational needs in climate and hydrological monitoring.

#### 334 *4.2. Enhanced quality control*

335 To ensure the reliability of precipitation series before modeling or reconstruc-  
336 tion, reddPrec introduces an enhanced quality control (QC) framework. This

337 feature extends the standard QC procedures by incorporating both visual and  
338 automatic diagnostic tools to detect subtle or non-obvious data issues.

339 We first illustrate the functionality of the enhanced QC with `eqc_Plot()`,  
340 which provides an integrated diagnostic overview for individual time series. Fig-  
341 ure 5 shows an example output from this function for two contrasting stations:  
342 one with high-quality data and another exhibiting significant quality deficiencies.  
343 Each panel includes time series plots of the full precipitation range, along with  
344 a focused view of low precipitation values (0 to 5 mm) to highlight anomalies in  
345 the light-rain regime. In addition, four key diagnostic components support the  
346 enhanced QC checks:

- 347 • **Truncation:** A smoothed line plot highlights periods dominated by constant  
348 heavy precipitation values, which may indicate sensor or recording issues.  
349 This commonly occurs when rain gauges are not emptied promptly, causing  
350 overflow and truncation of extreme events. Such censoring can lead to un-  
351 derestimation of total rainfall during intense events and may distort analyses  
352 of extremes, including intensity-duration-frequency curves, risk assessments,  
353 and infrastructure planning (e.g., flood defenses).
  
- 354 • **Small gaps:** A set of colored lines displays the yearly counts of precipitation  
355 values within sub-millimeter intervals (e.g., 0–1 mm, 1–2 mm, etc., exclud-  
356 ing integers), helping to detect systematic omissions of light precipitation.  
357 These gaps may arise from observer errors (missed or rounded records),  
358 instrument limitations (e.g., tipping buckets failing to register drizzle), or  
359 processing steps (e.g., aggregation or flagging). Such omissions bias rainfall  
360 statistics by underestimating light precipitation frequency and totals, ulti-  
361 mately affecting wet-day counts, rainfall intensity distributions, and climate  
362 trend analyses.
  
- 363 • **Weekly cycle:** A bar plot shows the fraction of wet days for each day of the  
364 week, with annotations above each bar indicating the count of wet days and  
365 a dotted line representing the overall weekly average. Statistically unusual  
366 days are highlighted in red, revealing weekly cycle patterns. Those patterns  
367 are often linked to human activity, such as missed observations on weekends  
368 or holidays. Automatic stations may also introduce cycles due to mainte-  
369 nance schedules or data recording intervals. These artificial patterns can  
370 distort analyses of rainfall frequency, persistence, or seasonality.
  
- 371 • **Precision and rounding:** A bar plot displays the yearly frequency of decimal  
372 digits (0–9), allowing users to detect changes in measurement resolution or  
373 rounding patterns over time. Rounding may stem from observer practices  
374 (e.g., recording only to the nearest millimeter, tenth of a millimeter, or full

375 integers) or instrumentation constraints. This can distort rainfall intensity  
376 distributions, misrepresent light and moderate rainfall events, and introduce  
377 long-term biases that affect hydrological modeling and trend detection.

378 Based on these diagnostics and considering Figure 5, we can easily notice the  
379 contrast between the two time series. In the high-quality series (`raw_data_00`),  
380 the precipitation patterns appear continuous and consistent across the entire pe-  
381 riod with no evident truncation. The small gaps plot shows a good even dis-  
382 tribution of light precipitation across different sub-millimeter ranges and years,  
383 suggesting minimal omissions. The weekly cycle is practically flat with no anom-  
384 alous days, and the distribution of decimal frequencies remains balanced over time,  
385 indicating stable measurement precision.

386 On the other hand, the low-quality series (`raw_data_01`) displays multiple  
387 issues. A truncation period is evident (1980 to 1990), where fixed high values  
388 (approximately 20 mm) dominate the time series, possibly due to gauge overflow.  
389 The small gaps plot reveals years (1970 to 2010) with substantial drops in light  
390 precipitation, pointing to systematic underreporting of drizzle or light rain. The  
391 weekly cycle is irregular, with Sundays showing a significantly high wet-day frac-  
392 tion, hinting at observer-related biases. Finally, the rounding and precision plot  
393 shows an important shift around 1990, indicating changes in instrumentation or  
394 data handling practices.

395 This comparison highlights how enhanced QC tools can uncover subtle but  
396 critical issues that would otherwise go unnoticed, supporting more robust data  
397 preparation and improving the reliability of downstream analyses. However, when  
398 dealing with larger datasets, the visual approach may not be practical. In such  
399 cases, automation of the enhanced QC is proposed through the use of a level-  
400 criteria definition with the `eqc_Ts()` function.

401 To illustrate `eqc_Ts()`, we compared two datasets with prior knowledge of  
402 their quality: Switzerland and Aragón (Spain). For Switzerland, we used the Swiss  
403 NBCN, while for Aragón, precipitation data were collected from public sources  
404 (National Meteorological Service of Spain – AEMET) and automatic hydrological  
405 monitoring systems.

406 The quality level distribution for Switzerland and Aragón, displayed in Figure  
407 6, reflects clear differences between the two datasets. In Switzerland, almost all  
408 stations were classified as level 0, indicating consistently high data quality. Only a  
409 few stations were flagged as level 1, mainly due to minor issues such as truncation  
410 effects. In contrast, the Aragón dataset showed a broader spread across levels 0,  
411 1, and 2. Many stations were flagged at level 2, suggesting more frequent issues  
412 across all enhanced QC tests, especially small gaps and precision/rounding pat-  
413 terns. The mixture of data sources in Aragón likely contributes to this variability.  
414 In general, these results demonstrate how the enhanced QC framework can detect

415 both station-level problems and broader network inconsistencies, helping guide  
416 better data selection and preparation for modeling.

417 While this experiment demonstrates the effectiveness of the enhanced QC  
418 framework, it is important to recognize its limitations. The automatic classifi-  
419 cation can distinguish clear cases of high- or low-quality series but may be less  
420 sensitive to intermediate-quality conditions. Moreover, the performance of the  
421 enhanced QC can be influenced by the prevailing climate regimes, whether the  
422 region is predominantly wet, dry, or transitional, as highlighted by Huerta et al.  
423 (2024). Nevertheless, the flexible design, allowing users to adjust level defini-  
424 tions, ensures adaptability across different datasets, climate regimes, and project  
425 needs. Beyond its role in supporting reconstruction workflows, the enhanced QC  
426 opens up promising opportunities for broader applications, such as the systematic  
427 evaluation of citizen weather station networks or quality control in multi-source  
428 precipitation datasets. By improving our ability to identify and address data in-  
429 consistencies, enhanced QC contributes to building more reliable, inclusive, and  
430 resilient climate data infrastructures for future research and decision-making.

### 431 *4.3. Homogenization*

432 To complement the enhanced quality control, reddPrec introduces a homoge-  
433 nization function designed to detect and adjust hidden inhomogeneities in daily  
434 precipitation series. The function implements a multi-test detection strategy com-  
435 bined with a quantile-based adjustment procedure, offering a flexible and auto-  
436 mated workflow for improving dataset consistency.

437 Testing homogenization algorithms on daily precipitation remains particularly  
438 challenging. Unlike air temperature data, for which benchmark datasets and  
439 well-established evaluation frameworks are available, precipitation series are more  
440 variable, discontinuous, and lack widely accepted benchmarks at daily timescales.

441 To address this gap, we constructed a corrupted version of the Swiss NBCN  
442 dataset by introducing controlled random artifacts into the original series. This  
443 synthetic corruption adds obvious inhomogeneities, such as shifts in bias, variance,  
444 and frequency, while preserving a known ground truth for evaluation. Full details  
445 of the corruption procedure are provided in Appendix B. Using this corrupted  
446 dataset, we evaluated the reddPrec homogenization module exclusively on the  
447 corrupted series (1960-2015), focusing on three main components: break detection,  
448 adjustment performance, and trend preservation.

449 First, break detection performance was assessed using two metrics: Break De-  
450 tection Accuracy (BDA) and Timing Accuracy (TA). Both metrics range from 0  
451 to 1, where values near 1 indicate more accurate detection (Appendix A). Specif-  
452 ically, a BDA close to 1 means a high proportion of true breaks were correctly  
453 identified, while a TA near 1 indicates that all true break points were detected  
454 within the specified tolerance window. Second, for adjustment performance, we

455 evaluated how well the method corrected the corrupted precipitation series by  
456 computing the mcc and dr metrics, comparing the adjusted series with the origi-  
457 nal uncorrupted data. Third, for trend preservation, we examined whether the  
458 homogenization procedure retained the short-term trends by calculating linear  
459 slopes on the PRCPTOT and R1mm indices. The agreement between the origi-  
460 nal and homogenized trends was assessed using the mcc-slope and dr-slope metrics  
461 (mcc and dr applied to slopes).

462 For this evaluation, we applied the `hmg_Ts()` function with the following  
463 settings: `neibs_min = 2`, `neibs_max = 12`, `cor_neibs = 0.5`, `wet_day = -1`,  
464 `perc_break = 22`, `apply_qc = 0.5`, and `mm_apply_qc = 0.1`, allowing the method  
465 to automatically define adjustment periods. Although setting `wet_day = -1` (in-  
466 cluding adjustment of zero-precipitation values) is not advisable when working  
467 with real datasets, it was adopted here to correct the artificially introduced inho-  
468 mogeneities in the corrupted series. In operational settings, it is crucial to ensure  
469 that real precipitation series undergo thorough quality control and are free from  
470 obvious errors before homogenization. The outcomes of the evaluation, including  
471 break detection accuracy, adjustment performance, and trend preservation, are  
472 summarized in Figure 7.

473 In the break detection evaluation, the homogenization method demonstrated  
474 strong performance in identifying the majority of artificial breakpoints, as indi-  
475 cated by a high BDA value (above 0.7). The TA results further confirmed the  
476 method’s precision, with detected breakpoints falling within a narrow window  
477 around the true dates, supported by the  $\pm 1$ -year tolerance criterion for break-  
478 point agreement.

479 Regarding adjustment performance, the method showed notable success in  
480 correcting the corrupted series. Both mcc and dr scores improved significantly  
481 following homogenization, indicating enhanced agreement with the original un-  
482 corrupted precipitation characteristics. In particular, mean mcc and dr values  
483 exceeded 0.7, reflecting effective adjustment. However, a higher variance was ob-  
484 served in mcc compared to dr, suggesting greater uncertainty in correcting wet/dry  
485 day frequencies than precipitation amounts.

486 In terms of short-term trend preservation, the method achieved moderate to  
487 high success. The linear slopes calculated from the PRCPTOT and R1mm indi-  
488 cators exhibited moderate and high mcc values (around 0.7 and 0.5, respectively),  
489 reflecting good agreement in the directionality of trends (i.e., correct identification  
490 of positive and negative trends). Meanwhile, moderate dr values (approximately  
491 around 0.5) indicated that although the general magnitude of trends was captured,  
492 discrepancies remained regarding their exact strength.

493 Overall, these experiments confirmed that the homogenization procedure is  
494 effective not only in detecting and adjusting hidden inhomogeneities but also in  
495 retaining the broader precipitation patterns and short-term trends, even in the

496 presence of artificially introduced disturbances.

497 To further explore the impact of homogenization, we extended the analysis  
498 by comparing long-term trend slopes using both the homogenized corrupted data  
499 and the real Swiss NBCN data over the full period. Additionally, we applied  
500 the homogenization function (using the same parameters as above, except setting  
501 `wet_day = 0`) directly to the real dataset to assess potential changes. While  
502 the Swiss NBCN dataset is often assumed to be reliable, this exercise allowed  
503 us to examine potential inherent inhomogeneities and evaluate whether applying  
504 homogenization would reveal significant changes. Results from these comparisons  
505 are displayed in Figure 8.

506 Figure 8a reveals a notable impact on long-term trends after adjusting the  
507 corrupted data. Although the corrupted series were effectively corrected overall,  
508 the magnitude of the original trends was more difficult to replicate exactly (dr  
509 close 0.0). This effect was mainly observed in stations where artificial corruption  
510 was introduced (red points), while uncorrupted stations (blue points) remained  
511 practically unchanged by the homogenization, as expected (dr and mcc close to  
512 1).

513 Considering the homogenization of the original Swiss NBCN dataset (Figure  
514 8b), we observed some minor adjustments, with slight changes in long-term trends  
515 compared to the original data (dr and mcc close to 1). These effects were more  
516 pronounced for the magnitude of PRCPTOT trends than for R1mm trends.

517 These extended findings highlight the intrinsic difficulty of precisely recovering  
518 original long-term trends after correcting local inhomogeneities in daily precipita-  
519 tion series. Even moderate adjustments made to correct breakpoints can introduce  
520 noticeable impacts on long-term trends. Nevertheless, the application of homog-  
521 enization to the real Swiss NBCN dataset confirmed that the method does not  
522 introduce major false corrections: stations without significant inhomogeneities re-  
523 mained largely unchanged. This supports the reliability of the homogenization  
524 procedure in practice. Further investigation into the homogenization of high-  
525 quality networks would be valuable, but this lies beyond the scope of the present  
526 study.

527 In summary, the homogenization approach implemented in `reddPrec` demon-  
528 strates strong potential for detecting and adjusting hidden inhomogeneities in  
529 reconstructed daily precipitation series while preserving key precipitation pat-  
530 terns and trends. Nevertheless, some limitations were identified: although the  
531 correction of localized breaks was generally effective, the precise reconstruction of  
532 long-term trends proved more difficult, highlighting the sensitivity of daily pre-  
533 cipitation to small adjustments. It is important to note that this homogenization  
534 procedure is specifically designed for reconstructed precipitation datasets, where  
535 small inconsistencies are not only from physical measurement errors but may  
536 arise from the reconstruction process itself. In this sense, the homogenization of



537 reconstructed precipitation offers an additional "view" of precipitation variabil-  
538 ity, complementing traditional observations and reconstructions, similar to the  
539 framework proposed in Huerta et al. (2024). Thus, reddPrec's homogenization  
540 tool contributes a novel perspective to the growing efforts to refine and enhance  
541 precipitation datasets for climatic and hydrological studies.

## 542 **5. Future developments, limitations, and conclusions**

543 In this work, we presented reddPrec as a versatile tool for reconstructing daily  
544 precipitation series, featuring advanced quality control, gap-filling, homogeniza-  
545 tion, and grid creation procedures. The method offered is highly flexible and easy  
546 to use. However, future research is needed to incorporate other climate variables  
547 (such as air temperature), broadening its scope. Adding uncertainty quantifica-  
548 tion tools for the enhanced quality control and homogenization would help assess  
549 result reliability and strengthen the package. Furthermore, expanding reddPrec  
550 to support additional programming languages (Python and Julia) would increase  
551 its usability and reach, enabling a broader user base and facilitating integration  
552 with various platforms.

553 In conclusion, reddPrec provides a robust tool for reconstructing precipita-  
554 tion data, offering flexibility for climate research. While challenges remain, the  
555 package represents a significant step toward creating high-quality, reproducible  
556 precipitation datasets. Continued development will expand its capabilities and  
557 make it even more valuable for a diversity of fields.

## 558 6. CRediT authorship contribution statement

559 **Adrian Huerta:** Writing – original draft, Writing – review & editing, Visu-  
560 alization, Software, Methodology, Investigation, Formal analysis, Data curation,  
561 Conceptualization. **Stefan Brönnimann:** Writing – review & editing, Supervi-  
562 sion. **Martín de Luis:** Writing – review & editing, Supervision. **Santiago Be-**  
563 **guería:** Writing – review & editing, Supervision. **Roberto Serrano-Notivoli:**  
564 Writing – review & editing, Visualization, Supervision, Methodology, Conceptu-  
565 alization.

## 566 7. Declaration of competing interest

567 The authors declare that they have no known competing financial interests or  
568 personal relationships that could have appeared to influence the work reported in  
569 this paper.

## 570 8. Software and data availability

571 The stable version of reddPrec is available on CRAN ([https://doi.org/10.](https://doi.org/10.32614/CRAN.package.reddPrec)  
572 [32614/CRAN.package.reddPrec](https://doi.org/10.32614/CRAN.package.reddPrec)) and the latest available on GitHub ([https://](https://github.com/rsnotivoli/reddPrec)  
573 [github.com/rsnotivoli/reddPrec](https://github.com/rsnotivoli/reddPrec), last access: 10 Jun 2025). The data and  
574 code of the demonstration examples are available at Figshare (Huerta, 2025) and  
575 Github (<https://github.com/adrHuerta/examples-reddPrec>, last access: 10  
576 Jun 2025).

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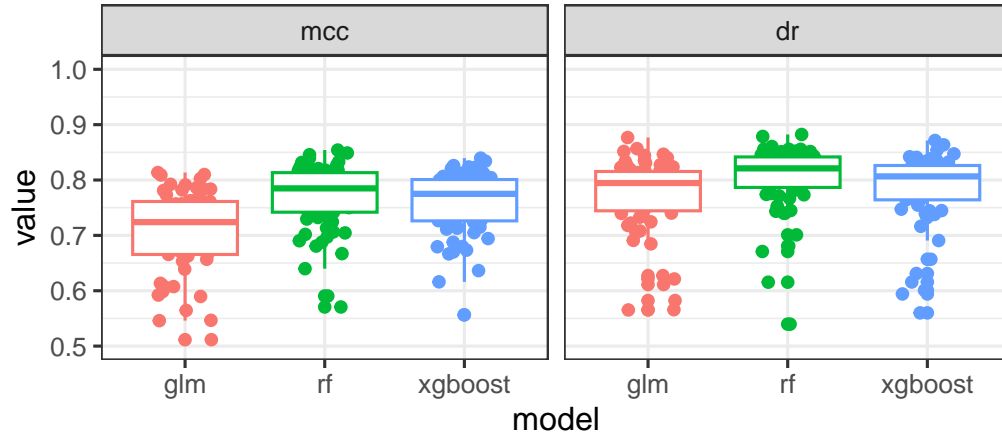


Figure 1: Boxplots of gap-filling evaluation metrics (mcc: Matthews correlation coefficient, dr: refined index of agreement) for precipitation data in Switzerland (2010–2015) from 69 stations. The performance of three gap-filling methods—linear model (glm), random forest (rf), and extreme gradient boosting (xgboost)—is compared based on these metrics. In both plots, the maximum value is one, meaning perfect prediction.

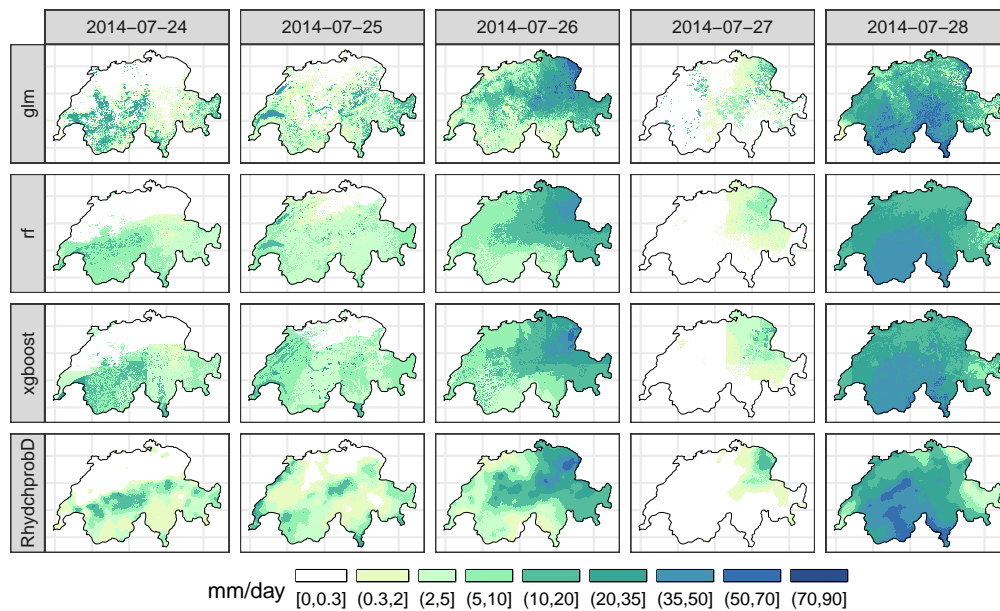


Figure 2: Daily precipitation fields ( $0.009^\circ \approx 1 \text{ km}$ ) for an extreme precipitation event (July 24–28, 2014) in Switzerland. The comparison includes three gridding methods—generalized linear model (glm), random forest (rf), and extreme gradient boosting (xgboost)—along with the median value of the RhydchprobD product. The glm, rf, and xgboost grids were constructed using 69 stations.



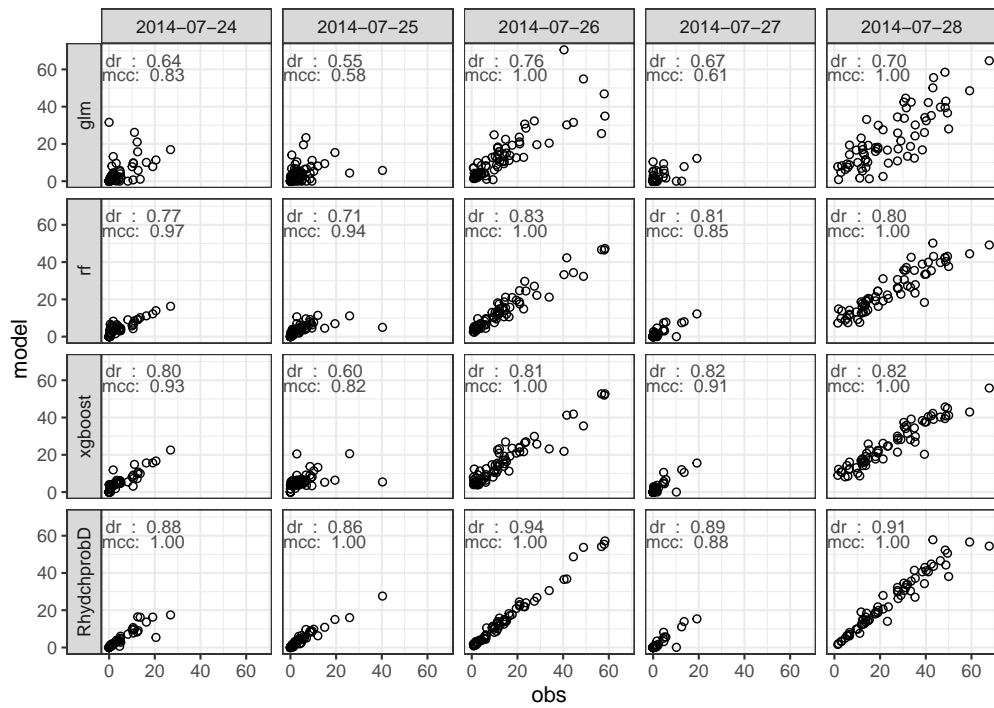


Figure 3: A scatterplot comparing station values with the nearest grid for an extreme precipitation event (July 24–28, 2014) in Switzerland. Each plot displays the evaluation metrics (mcc: Matthews correlation coefficient, dr: refined index of agreement).

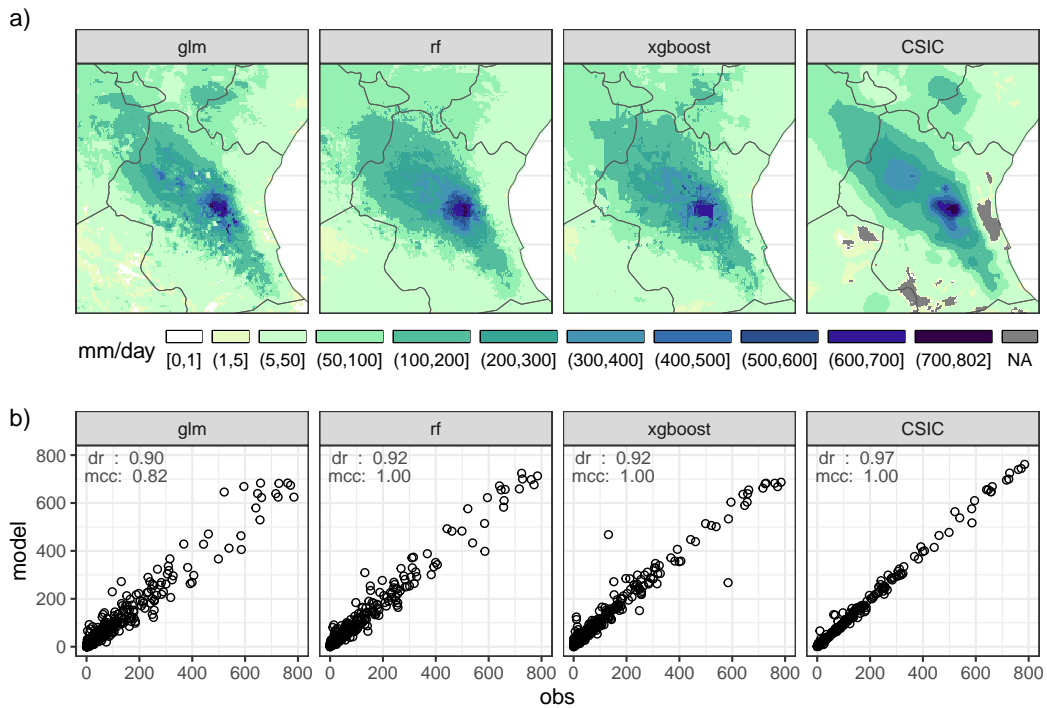


Figure 4: (a) Daily precipitation fields ( $0.009^\circ \approx 1$  km) for an extreme precipitation event (October 29, 2024) in Valencia, Spain. (b) A scatterplot comparing station values with the nearest grid values with evaluation metrics (mcc: Matthews correlation coefficient, dr: refined index of agreement) is displayed for each plot. The comparison includes three gridding methods—generalized linear model (glm), random forest (rf), and extreme gradient boosting (xgboost)—along with fields developed by the Consejo Superior de Investigaciones Científicas (CSIC). The glm, rf, and xgboost grids were constructed using data from 443 stations.

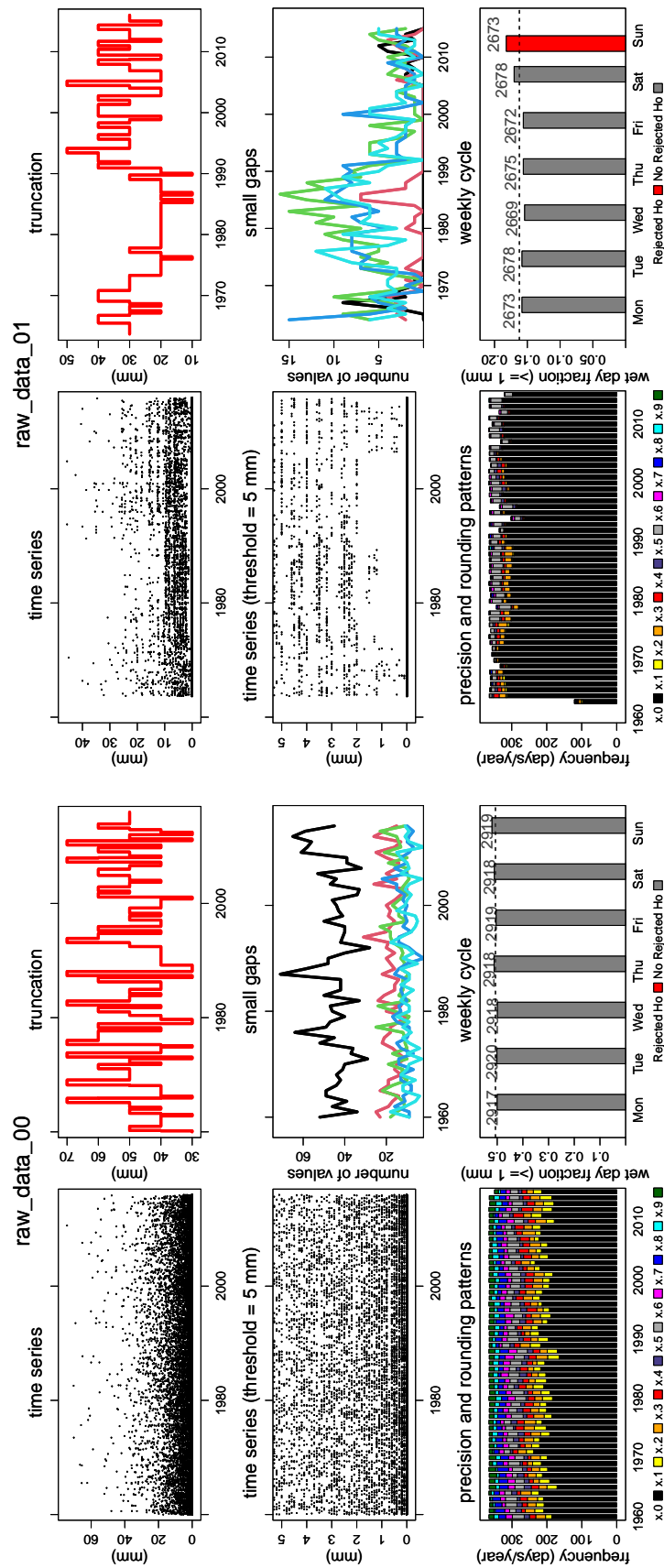


Figure 5: Example enhanced quality control plots for two time series: one from a high-quality station (raw\_data\_00) and one from a low-quality station (raw\_data\_01). In addition to the raw time series (including one with a lower precipitation range), the plots illustrate key diagnostic tests applied during quality control—truncation, small gaps, precision and rounding patterns.

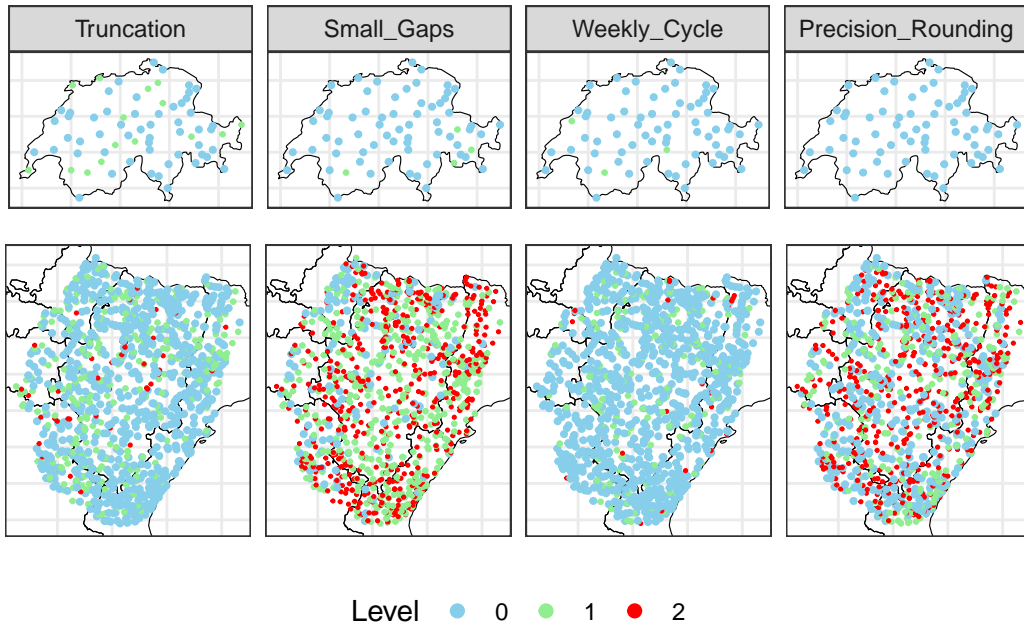


Figure 6: Application of automatic enhanced quality control to daily precipitation data for Switzerland (top) and Aragon (Spain, bottom) during 1980–2015. Each quality control test—truncation, small gaps, weekly cycles, and precision or rounding patterns—displays the assigned quality level (0, 1, or 2) for each station.

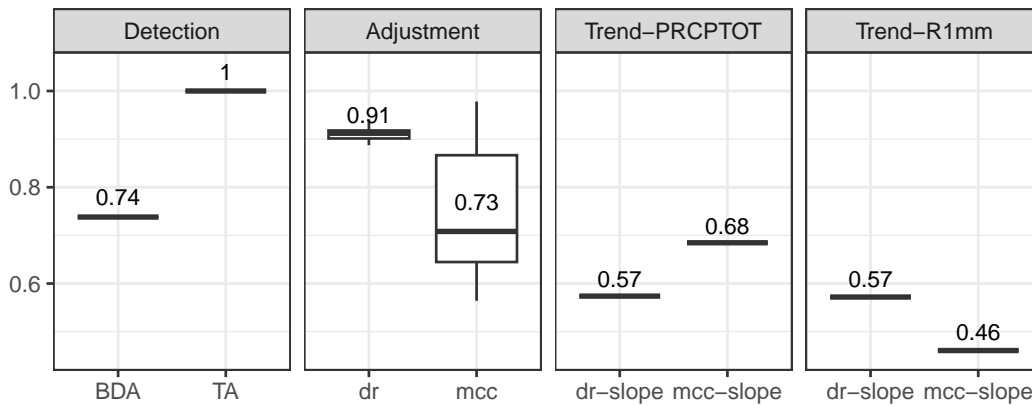


Figure 7: Summary plot for homogenization evaluation between corrupted and original Swiss NBCN: detection (BDA: break detection; and TA: Timing Accuracy), adjustment (dr: refined index of agreement; and mcc: Matthews correlation coefficient), and trend preservation in PRCP-TOT and R1mm indices.

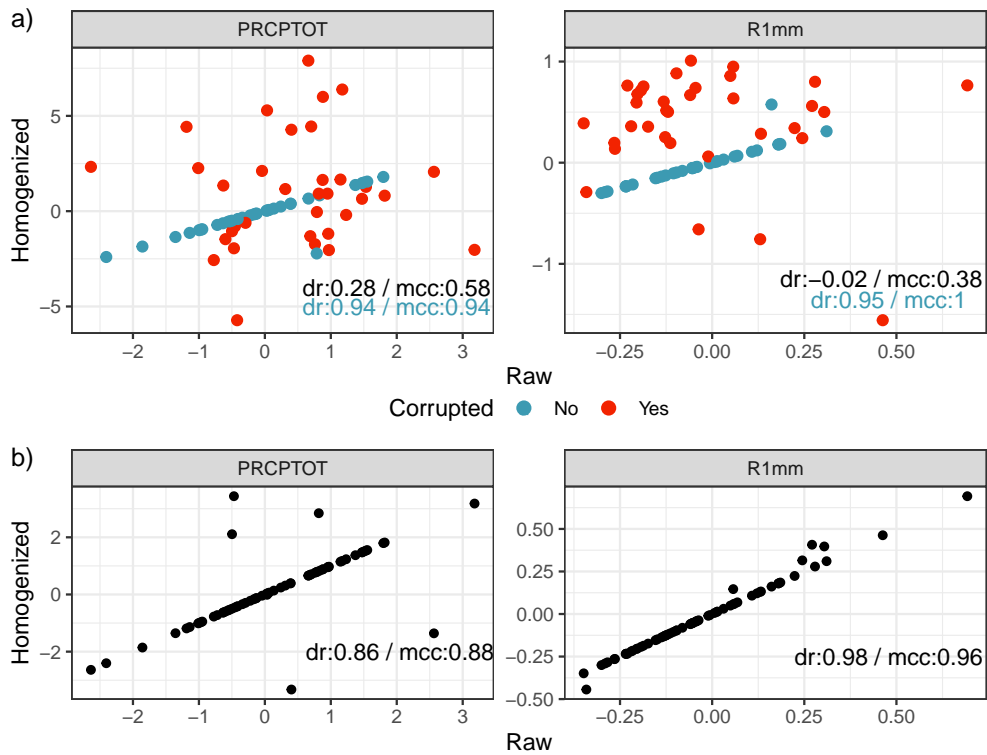


Figure 8: Scatterplot of homogenized and raw trend slopes for PRCPTOT and R1mm indices for the 1960-2015 period. (a) Display the long-term trend slopes between the homogenized corrupted and raw Swiss NBCN. (b) Display long-term trend slopes between the homogenized and raw Swiss NBCN. In the plot, the refined index of agreement (dr) and Matthews correlation coefficient (mcc) metrics are shown. In (a), dr and mcc are computed from the whole sample of points (black) and non-corrupted time series (red).

716 **Appendix A. Metrics used for evaluation**

717 To evaluate the performance of precipitation in the gap-filling and grid creation  
718 experiments, we employed two metrics:

- 719 • The Matthews correlation coefficient (*mcc*) is a balanced classification per-  
720 formance measure, particularly suitable for binary classification with imbal-  
721 anced classes (Chicco and Jurman, 2023). The *mcc* values range from -1  
722 (total disagreement) to +1 (perfect agreement), with 0 indicating no better  
723 than random classification. The *mcc* is calculated as:

$$mcc = \frac{TP \times TN - FP \times FN}{\sqrt{(TP + FP)(TP + FN)(TN + FP)(TN + FN)}}$$

724 where *TP* denotes the number of days correctly classified as wet ( $\geq 0.1$  mm),  
725 *TN* as days correctly classified as dry ( $< 0.1$  mm), *FP* as days incorrectly  
726 classified as wet, and *FN* as days incorrectly classified as dry.

- 727 • The refined index of agreement (*dr*) is a modified version of the traditional  
728 index of agreement proposed by Willmott et al. (2012). It is suited for  
729 continuous variables and aims to improve sensitivity to systematic biases  
730 and error distributions. The *dr* metric ranges from -1 (no agreement) to  
731 +1 (perfect agreement), with 0 indicating no predictive skill. The *dr* is  
732 calculated as:

$$dr = 1 - \frac{\sum_{i=1}^n |p_i - \hat{y}_i|}{2 \sum_{i=1}^n |\hat{y}_i - \bar{y}|}$$

733 where  $n$  is the number of observations,  $p_i$  is the predicted precipitation on  
734 day  $i$ ,  $\hat{y}_i$  is the observed precipitation on day  $i$ , and  $\bar{y}$  is the mean of  $\hat{y}_i$ .

735 In addition, for the homogenization experiment, we used two custom metrics:

- 736 • The break detection accuracy (*BDA*) measures the overall accuracy of break  
737 detection. The metric ranges from 0 to 1, with higher values indicating more  
738 accurate break detection. It's calculated as:

$$BDA = \frac{TP}{TP + FP + FN}$$

739 where *TP* represents the detected breaks that correspond to the true break  
740 years, *FP* the detected breaks that do not correspond to any true break,  
741 and *FN* the true breaks that were missed.

- 742 • The timing accuracy (TA) metric measures the temporal precision of break  
743 detection. It calculates the proportion of true breaks that were detected  
744 within a tolerance (here equal to  $\pm 2$  years). TA varies from 0 to 1. A value  
745 of 1 indicates that all true breakpoints were detected within the tolerance  
746 window, while values close to 0 suggest poor timing between detected and  
747 true breakpoints. It is defined as follows:

$$TA = \frac{\text{Number of true breaks detected within tolerance}}{\text{Total number of true breaks}}$$

748 **Appendix B. Construction of corrupted Swiss NBCN.**

749 To test the homogenization approach implemented in reddPrec, we created  
750 a synthetically corrupted version of the Swiss NBCN(1960-2015). The original  
751 dataset was subjected to random artificial break injections affecting 50% of the  
752 stations within the 1970-2000 period.

753 Three types of artificial changes were introduced:

- 754 • Bias shifts: a constant offset ( $\pm 2$  mm) applied to a portion of the series.
- 755 • Variance changes: multiplicative variability alterations ( $\times 1.5$ ).
- 756 • Frequency changes: increased wet-day frequency, where selected dry days (0  
757 mm) were converted to wet days using gamma-distributed values.

758 To ensure consistency with the correction strategy in the demonstration ex-  
759 ample, only dry-to-wet frequency changes were introduced. The homogenization  
760 method relies on the `wet_day` argument to determine whether 0 mm values (dry  
761 days) are included in the adjustment process. In this study, `wet_day = 0` was  
762 used, meaning dry days were treated as valid values and included in the adjust-  
763 ment, allowing effective correction of added wet-day noise. By contrast, if `wet_day`  
764 `= 1` had been used, zeros would have been excluded from the adjustment proce-  
765 dure, complicating the correction of artificially added dry days. For this reason,  
766 wet-to-dry changes were excluded from the corruption setup to maintain coherence  
767 with the correction strategy.

768 The break metadata (station, break date, break years, type of break) were  
769 recorded to allow precise evaluation of detection accuracy, adjustment effective-  
770 ness, and trend preservation.



771 Appendix C. Supplementary Figures

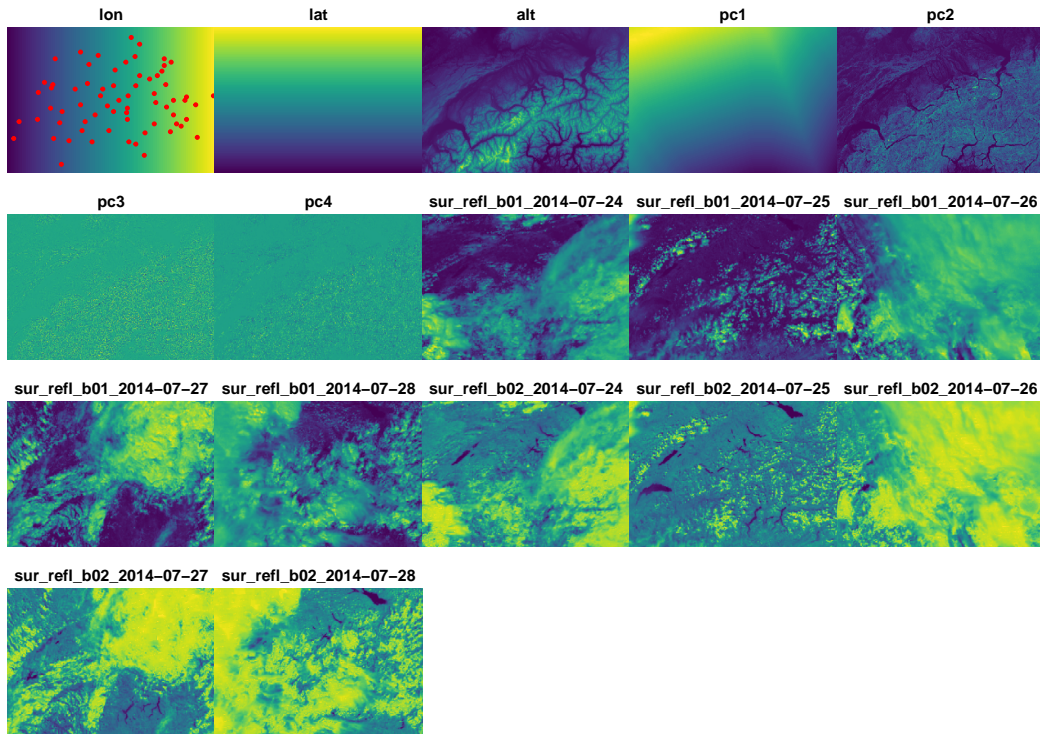


Figure C.9: Spatial covariates for an extreme precipitation event (July 24–28, 2014) in Switzerland. Static covariates include longitude (lon), latitude (lat), elevation (alt), and the first four principal components of topographical covariates (pc1, pc2, pc3, and pc4). Dynamic covariates include surface reflectance (sur\_refl) for bands 1 (b01) and 2 (b02). In the first subplot, the stations used for spatial modeling are shown as red dots. Colorbar legends were omitted.

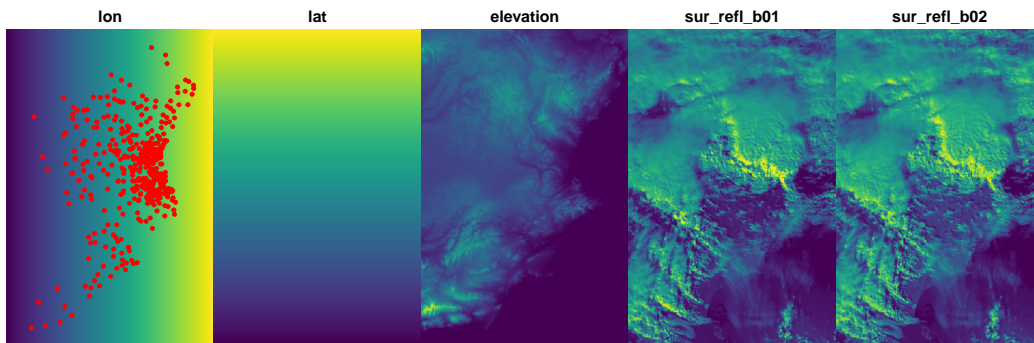


Figure C.10: Spatial covariates for an extreme precipitation event (October 29, 2024) in Valencia, Spain. Static covariates include longitude (lon), latitude (lat), elevation (elevation), and surface reflectance (sur\_refl) for bands 1 (b01) and 2 (b02). In the first subplot, the stations used for spatial modeling are shown as red dots. Colorbar legends were omitted.