



# Limitations of $M_w$ and $M$ Scales: Compelling Evidence Advocating for the Das Magnitude Scale ( $M_{wg}$ )—A Critical Review and Analysis

Ranjit Das<sup>1</sup> · Arindam Das<sup>2</sup>

Received: 8 October 2024 / Accepted: 22 December 2024  
© The Author(s), under exclusive licence to Indian Geotechnical Society 2025

**Abstract** Precise determination of earthquake size is crucial for various geoscientific and engineering applications. The Moment Magnitude ( $M_w$ ) scale, introduced by Kanamori in 1977, was a significant advancement. Kanamori (1977) advocated use of  $M_w$  for large earthquakes ( $\geq 7.5$ ). (Hanks and Kanamori in J. Geophys. Res. 84:2348–2350) later extended the  $M_w$  scale named as  $M$  scale by considering close coincidence of three equations. The use of the moment magnitude scale  $M$  scale for magnitudes below 7.5 is not appropriate, as Eq. (1) from Purcaru and Berckhemer (Purcaru and Berckhemer in Tectonophysics 49:189–198, 1978) was specifically derived for  $M_s$  values in the range of  $M_s \lesssim 7.0$ . Furthermore, the  $M$  scale has not been validated globally for magnitudes below 7.5; its validation is limited to Southern California. Additionally, the  $M$  or  $M_w$  scale is based on surface waves and may not be applicable for all earthquake depths. Furthermore, Gutenberg and Richter (Gutenberg and Richter in Bull Seismol Soc Am 46:105–145, 1956) recommended using body waves, rather than surface waves, for the development of a magnitude scale as surface waves do not represent the earthquake source. To address these shortcomings, the Das Magnitude scale ( $M_{wg}$ ) has been introduced in recent literature (Bulletin of Seismological Society of America, Das et al. (Das et al. in Bull Seism Soc Am 109:1542–1555, 2019); Natural Hazard, 2023), incorporating global data during 1976–2006 with 25,708 events with observed seismic moments ( $M_0$ ) and body wave magnitudes

( $m_b$ ), in line with the recommendations of Gutenberg and Richter (Gutenberg and Richter in Bull Seismol Soc Am 46:105–145, 1956). Recent seismological literature (Gasperini and Lolli, (Gasperini and Lolli in Bull Seismol Soc Am, 2024)) has inaccurately critiqued Das et al. (Das et al. in Bull Seism Soc Am 109:1542–1555, 2019), stating that the  $M$  scale is adequate and suggesting that certain foundational assumptions of the  $M_{wg}$  scale are inappropriate. However, our detailed analysis demonstrates that the  $M_{wg}$  scale is firmly grounded in robust scientific evidence and constructed on sound principles. It is important to note that all magnitude scales, including  $M_w$ ,  $M$ , and  $M_e$ , are developed using linear least squares methods. Therefore, if the fundamental assumptions underlying the  $M_w$ ,  $M$ , and  $M_e$  scales are considered valid, there is no justifiable reason to challenge the foundational assumptions of the  $M_{wg}$  scale.  $M_{wg}$  uses body-wave magnitude instead of surface waves and is applicable to all depths, making it more suitable for a wider range of earthquakes.  $M_{wg}$  is a better measure of energy release compared to  $M_w$ , providing a more accurate representation of earthquake strength. The  $M_{wg}$  scale demonstrates a closer correspondence with observed  $m_b$  and  $M_s$  values at a global level compared to the  $M$  scale. The average difference between observed  $m_b$  and  $M$  is  $-0.31 \pm 0.30$ , whereas the difference between observed  $m_b$  and  $M_{wg}$  is significantly smaller at  $0.008 \pm 0.33$  (Das et al. (Das et al. in Bull Seism Soc Am 109:1542–1555, 2019)). The uncertainty associated with the development of  $M_{wg}$  is limited due to its simplicity, while  $M_w$  or  $M$  involve a constant term and multiple substitutions, potentially introducing additional uncertainty.  $M_{wg}$  was developed and validated using global datasets, ensuring its applicability to a diverse range of seismic events. The  $M_{wg}$  scale significantly reduces the statistical differences with  $m_b$  and  $M_s$  compared to the larger discrepancies observed with  $M$ , offering a more reliable framework for

✉ Ranjit Das  
ranjit.das@ucn.cl

<sup>1</sup> Computer Science and Engineering Department, Universidad Católica del Norte, Avenida Angamos, 0610 Antofagasta, Chile

<sup>2</sup> Civil Engineering Department, Indian Institute of Technology Bombay, Bombay, India

understanding regional energy budgets. In this manuscript, we critically compare  $M_w$ ,  $M$ , and  $M_{wg}$  scales. We revalidated the  $M$  scale using a comprehensive global dataset, finding that it significantly deviates from observed  $m_b$  ( $<5.5$ ) and  $M_s$  (6–8) within their respective applicable ranges. Furthermore, revalidation of the  $M$  scale concludes that it is not appropriate to be below 7.5.

**Keywords** Seismic moment magnitude scale · Body-wave magnitude ( $m_b$ ) · Surface-wave magnitude ·  $M_{wg}$  Scale · Seismic moment

## Introduction

Accurate measurement of earthquake magnitudes is fundamental for all seismological and engineering studies. Over the years, several magnitude scales have been developed, each with its own methodology and scope. The Richter scale, introduced in 1935, was one of the first methods used to quantify earthquake size based on the amplitude of seismic waves. However, its applicability was limited to specific regions. Conventional magnitude scales, including  $m_b$ ,  $M_s$ , and  $M_L$ , encounter saturation issues when the earthquake's rupture dimension exceeds the wavelength of the seismic waves used for magnitude determination. For large earthquakes, rupture dimensions often exceed the wavelengths of seismic waves commonly employed for magnitude estimation. This phenomenon, known as saturation, introduces systematic errors in assessing the energy released, hindering accurate magnitude determination. Traditional magnitude scales, such as  $M_L$  and  $m_b$ , based on short-period (1 Sec) seismic waves, typically saturate around magnitude 7.0 ([1], [2]). While the surface wave magnitude ( $M_s$ ) extends the saturation limit to approximately 8.3, it remains inadequate for characterizing the largest earthquakes. To address these limitations, the moment magnitude scale ( $M_w$ ) was introduced by Kanamori [3] in seismic literature to understand the earthquake size without saturation. As Kanamori [3] defined  $M_w$  scale for large earthquakes ( $\geq 7.5$ ), therefore, Hanks and Kanamori [1] extended the  $M_w$  scale to measure lower, medium and large earthquakes and in this way  $M$  scale (1979) has been the standard for decades, offering a means to quantify the energy released during an earthquake. Despite its widespread use, the  $M$  scale is not without its shortcomings. Given the extensive use of the  $M$  scale as global scale in seismological and geophysical literature, it is crucial to assess its suitability for accurate evaluation and interpretation in terms of energy. The term  $M_w$  is commonly used by GCMT (Global Centroid Moment Tensor) for estimating moment magnitude. However, according to Hanks and Kanamori [1], the correct term should simply be  $M$  if one intends to

measure earthquakes across all ranges ( $\geq 3$ ), including small, medium, large, and very large events.

The inadequacy of surface waves in characterizing deeper earthquakes prompted Gutenberg and Richter [4] to propose a magnitude scale based on body waves. Subsequently, Kanamori [5] introduced a magnitude scale derived from seismic moment ( $M_0$ ) for more accurate representation of seismic events for all depths. The complex nature of earthquakes precludes a singular metric for their quantification, necessitating the development of diverse magnitude scales, such as  $M_e$  [6] and  $M_{wg}$ , [7–10] to comprehensively characterize seismic events. The  $M_w$  scale aims to offer a more accurate representation of earthquake magnitudes by incorporating a broader frequency range and minimizing regional biases. Through a detailed comparison of the  $M_w$ ,  $M$  and  $M_{wg}$  scales, including an analysis of their respective strengths and limitations, this study highlights their importance in seismology.

## Background of $M_w$ Scale

### *Kanamori's Introduction of the $M_w$ Scale (1977)*

In 1977, Kanamori introduced an earthquake magnitude scale that revolutionized the way seismic events are quantified and understood. This scale, commonly referred to as the moment magnitude scale ( $M_w$ ), was designed to overcome the limitations of earlier magnitude scales such as the Richter scale (1935), body wave magnitude ( $m_b$ ), surface wave magnitude ( $M_s$ ), and local magnitude ( $M_L$ ). The development of the  $M_w$  scale involved several key steps:

1. *Empirical Relationship:* Kanamori [3] utilized the empirical relationship between Gutenberg and Richter energy and magnitude:
  - a.  $\log E_s = 1.5 M_s + 11.8$ . This relationship was derived from the Gutenberg-Richter equation for seismic energy:
  - b.  $\log E = 5.8 m + 2.4$ . and the substitution  $m = 2.5 + 0.63 M$ .
2. *Development of Moment Magnitude Scale  $M_w$ :* Kanamori [3] expressed Gutenberg-Richter energy equation in terms of seismic moment using the equation:
  - a.  $E/M_0 = (\Delta\sigma)/2\mu = 5 \times 10^{-5}$ , where  $\sigma$  represents earthquake stress drop and  $\mu$  is the shear modulus. By substituting  $E/M_0 = (\Delta\sigma)/2\mu = 5 \times 10^{-5}$ , into the Gutenberg-Richter energy equation, Kanamori derived:  $\log M_0 = 1.5 M_s + 16.1$ . Kanamori [3] reversed the equation  $\log M_0 = 1.5 M_s + 16.1$  to

express surface wave magnitude ( $M_s$ ) in terms of seismic moment ( $M_0$ ) and termed this as the moment magnitude ( $M_w$ ). The validation of the  $M_w$  scale was conducted by comparing it with observed surface wave magnitudes ( $M_s$ ) (Table 2 of [3]). Kanamori asserted that the  $M_w$  scale could be seen as a natural extension of  $M_s$ , as outlined in Table 2 of Kanamori [3]. This comparison demonstrated that the  $M_w$  scale provided a consistent and reliable measure of earthquake magnitude that aligned well with observed  $M_s$  values earthquakes.  $M_w$  scale of Kanamori [3] was derived for large earthquakes (magnitude  $\geq 7.5$ ), addressing the saturation problem of earlier magnitude scales for such events. However, recognizing the need for a magnitude scale that could accurately measure smaller, medium, large and large earthquakes, Kanamori emphasized the importance of developing a scale appropriate for a broader range of seismic events. This led to the further refinement of the  $M_w$  scale in a later study by Hanks and Kanamori [1].

### Hanks and Kanamori's Refinement (1979)

The moment magnitude scale ( $M_w$ ) was further updated by Hanks and Kanamori in 1979 to extend the initial work by Kanamori [3]. To achieve this, Hanks and Kanamori [1] compared the Eqs. (4–6) from their study and found that these equations closely coincided. Based on this comparison, they proposed the extended moment magnitude scale  $M$  as  $2/3 \log M_0 - 10.7$ . In the abstract of Hanks and Kanamori [1], they explicitly stated that the  $M$  scale is developed based on the close coincidence between their Eqs. (4–6) which lead them to define  $M$  scale as  $2/3 \log (M_0) - 10.7$  and stated that  $M$  is uniformly valid for  $3.0 \lesssim M_L \lesssim 7.0$ ,  $5 \lesssim M_s \lesssim 7.5$ , and  $M_w \gtrsim 7.5$ .

Furthermore the derivation of the  $M$  scale involved incorporating the ratio  $E/M_0 = (\Delta\sigma)/2\mu = 5 \times 10^{-5}$  into the Gutenberg-Richter energy magnitude equation as performed in Kanamori [3]:

$$\log E = 1.5M_s + 11.8$$

where  $M_s$  represents the surface wave magnitude. Hanks and Kanamori [1] reversed the resultant equation and termed the  $M_s$  as  $M_w$ . Kanamori [3] clearly stated that the constant value  $\Delta\sigma$  is strictly applicable for shallow and large earthquakes, and the shear modulus  $\mu$  ( $3-6 \times 10^{11}$ ) is strictly applicable for crustal earthquakes. Thus these limitations of  $M_w$  in the constant also applies in  $M$  scale of Hanks and Kanamori [1]. To enhance the accuracy of the  $M_w$  scale to lower and intermediate earthquakes, Hanks and Kanamori compared the equation  $\log M_0 = 1.5M_s + 16.1$  with the

Eq. (1) by Purcaru and Berckhemer [11]. Additionally, they compared it with the California local equation within the range  $3.0 \leq M_L \leq 7.0$ . However, significant spatial variations in local magnitude were observed, indicating the need for further study (e.g., [6, 12, 13], Aditya et al. [14], Amit et al. [15], Amit, [16]). The validation of  $M$  scale [1] was conducted utilizing local and surface wave magnitudes of Southern Californian seismicity (Tables 1 and Table 2 of [1]).

### Limitations and Recent Critiques

$M$  scale is based on surface waves and hence not appropriate for deep and intermediate earthquakes. Recent investigations suggest that the  $M$  scale developed by Hanks and Kanamori [1] is not adequate on a global level, particularly for smaller and medium earthquakes. This discrepancy may arise from the limited data range used by Purcaru and Berckhemer [11], use of surface wave magnitude in equation  $\log M_0 = 1.5 M_s + 16.1$ , use of constant value  $5 \times 10^{-5}$  in the Gutenberg Richter energy equation and the localized validation in Southern California. This localized approach raises questions about the universal applicability of the  $M$  scale across diverse tectonic settings. The details limitations we are discussing below.

*Limitations of Purcaru and Berckhemer's Equation:* Purcaru and Berckhemer [11] explicitly stated that their derived equation  $\log M_0 = 1.5 M_s + 16.1$  is reliable exclusively below  $M_s \lesssim 7.0$ , emphasizing its accuracy within the range  $5.0 \leq M_s \leq 7.0$ . However, Hanks and Kanamori [1] asserted that Purcaru and Berckhemer [11] developed the equation  $\log M_0 = 1.5 M_s + 16.1$  for the range  $5.0 \leq M_s \leq 7.5$  (Page 2348 of [1], stating, "Which is remarkably coincident with the  $M_0$ - $M_s$  relationship empirically defined by Purcaru and Berckhemer [11] for  $5 \lesssim M_s \lesssim 7.5$ :  $\log M_0 = 1.5 M_s + 16.1 (\pm 0.1)$ " consequently, the  $M$  scale proves inadequate for measuring earthquake sizes below magnitude 7.5. However,  $M$  scale is applicable for large earthquakes ( $\geq 7.5$ ).

*Applicability of the Unsaturated  $M$  Scale:* The unsaturated  $M$  scale holds global applicability for large earthquakes, as its foundation lies in equations such as  $\log E_s = 1.5M_s + 11.8$  [17] and  $\log M_0 = 1.5 M_s + 16.1$ . To extend support to smaller earthquakes within the  $M$  scale, Hanks and Kanamori [1] incorporated the relationship between  $\log M_0$  and  $M_L$  for Southern California [18], focusing on the magnitude range  $3 \lesssim M_L \lesssim 7.0$ . Consequently, the  $M$  scale can be deemed suitable exclusively for Southern California within the range  $3 \lesssim \text{magnitude} \lesssim 7.0$ , raising concerns about its universal applicability. It's essential to recognize that different regions worldwide possess distinct tectonic environments and geological settings. Numerous studies have underscored notable spatial variations in local magnitude (e.g., [6, 12, 13]). Utilizing the  $M$  scale ( $2/3 \log M_0 - 10.7$ ) for

smaller magnitude earthquakes on a global scale could potentially compromise earthquake hazard evaluations, early warning systems, and other seismic studies.

**Validation of the  $M$  Scale:** The development of a new method or scale requires a crucial step: validation. Typically, this involves the comparison of observed and estimated magnitudes, a standard practice in seismological literature ([1, 3], Ekström et al. [19]). Kanamori [3] validated the  $M_w$  scale ( $\geq 7.5$ ) on a global level (see Table 1 of [3]), while Hanks and Kanamori [1] focused their validation specifically on the tectonic conditions in Southern California, as detailed in Tables 1 and 2 of their publication. However, the limited scope of validation for the  $M$  scale by Hanks and Kanamori [1], confined to Southern California, raises questions about its applicability beyond this specific region for magnitudes  $\lesssim 7.5$ . The absence of validation across diverse global tectonic settings suggests potential limitations in the universal suitability of the  $M$  scale, especially when applied beyond the region for which it was derived and validated. An analysis by Das et al. [8] further highlights significant discrepancies between different magnitudes ( $m_b$ ,  $M_s$ ,  $M_c$ ) and  $M$  scale on a global level (see Figs. 1, 2, and 3 of [8]). The differences between observed  $m_b$  and  $M_s$  values compared to  $M_{wg}$  and  $M$  are statistically significant, as evidenced by null hypothesis testing.

**Applicability for Shallow Earthquakes:** Since surface waves are inadequate for measuring the size of deeper earthquakes, Gutenberg and Richter [4] recommended creating a magnitude scale based on body waves. Likewise, Kanamori [5] introduced an alternative magnitude scale equation, defined in terms of  $M_o$ , for assessing deep earthquakes. Kanamori [3] clearly stated that the constant used in the development of the  $M_w$  scale is applicable only to shallow and large earthquakes, specifically within the context of crustal earthquakes. Kanamori [5] introduced a distinct equation tailored for measuring earthquake size in the case of intermediate and deeper earthquakes, acknowledging the limitation of the  $M$  or  $M_w$  scales primarily designed for shallow earthquakes. This differentiation highlights a fundamental inadequacy within the  $M_w$  or  $M$  scales.

In the evaluation of earthquake sizes across various depths — shallow, intermediate and deeper — two key equations come into play [1, 5]: Equation (i) for  $M = 2/3 \log M_o - 10.7$  (Based on Surface Wave, applicable for shallow depths) and Equation (ii) for  $m_w = \log M_o / 2.4 - 10.1 / 2.4$  (Based on Body waves, applicable for all depths). The existence of these distinct equations, each associated with specific depth ranges, emphasizes the nuanced nature of earthquake size assessment based on depth considerations. This delineation prompts the recognition of an inadequacy within the  $M$  scale, tailored predominantly for shallow seismic events. Consequently, there arises a compelling rationale for embracing a new earthquake scale that seamlessly captures seismic events across all depth ranges—shallow, intermediate, and deeper.

**Derivation from Non-Direct Measurements:** The formula for  $M$  scale ( $2/3 \log M_o - 10.7$ ) was not derived from direct measurement of observed seismic moment. Instead, this scale was derived by substituting a constant term ( $E_s / M_o = 5 \times 10^{-5} = \Delta\sigma / 2\mu$ ) into the Gutenberg energy equation  $\log E_s = 1.5 M_s + 11.8$ . It is important to note that  $\Delta\sigma$  generally vary from a few bars to 125. The variability of  $\Delta\sigma$  is significant, and as a result, the assumption of constant stress drop cannot be made (Percaru and Berckhemer [11], [8, 9]). Consequently, depending on the constancy value, the  $M$  value for a given earthquake will change significantly. Furthermore, it is worth noting that the value of constancy ( $E_s / M_o = 5 \times 10^{-5}$ ) suggested by Kanamori [3] is only applicable for shallow earthquakes.

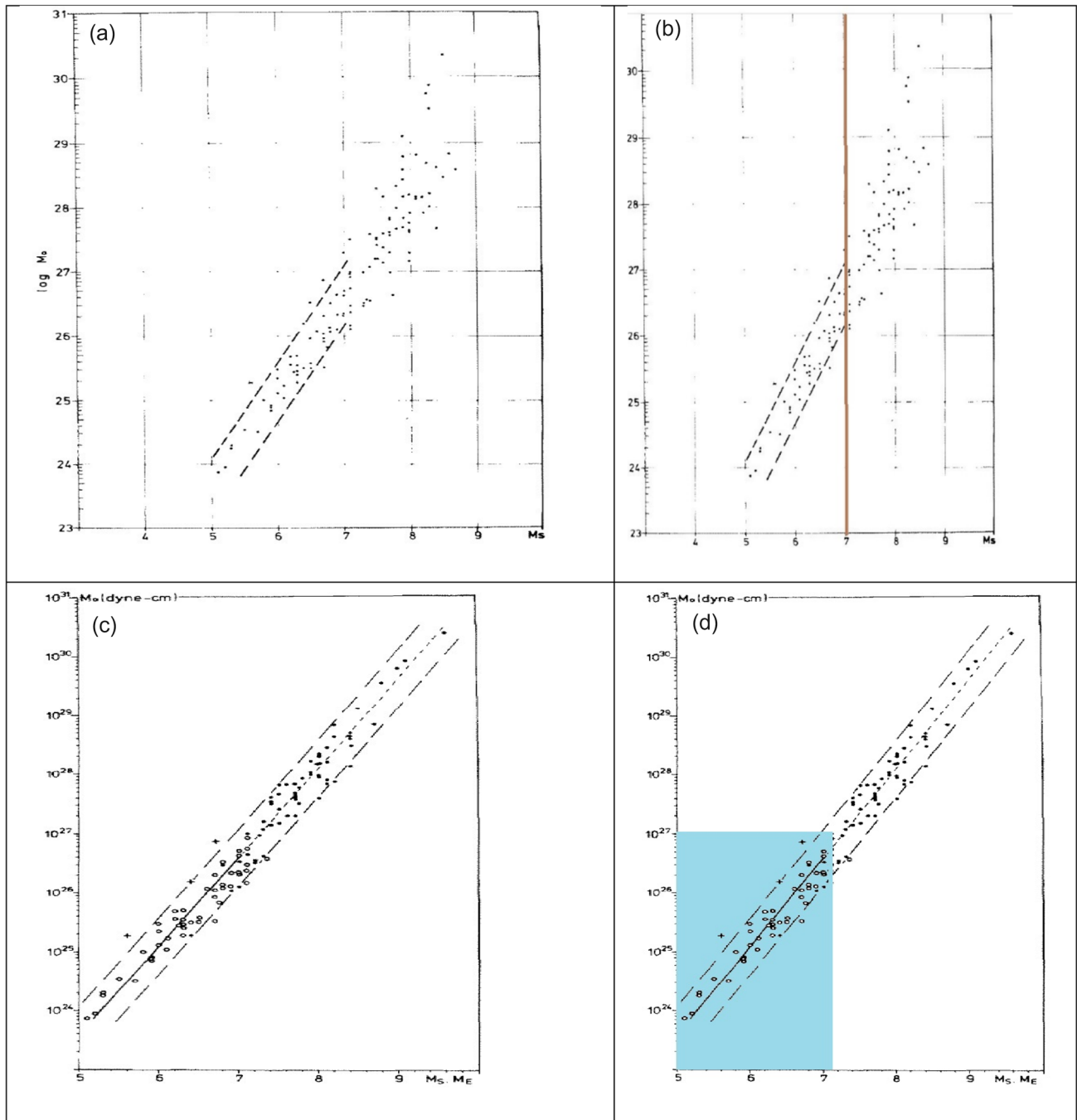
Choy and Boatwright [6] used a different constant value ( $E_s / M_o = 1.6 \times 10^{-5}$ ). Bormann and Di Giacomo [20] indicated that there can be significant variations in the constant value, ranging from  $(-7 \Theta_k - 3)$ , while Kanamori used a value of  $\Theta_k = -4.3$ . Therefore, there are significant limitations in the  $M_w$  or  $M$  scales due to the adoption of a fixed  $E_s / M_o$  value of  $5 \times 10^{-5}$ .

Recent studies by Gasperini and Lolli [21] have claimed that the  $M$  scale is adequate for measuring earthquake magnitudes. However, they failed to address the limitations of the  $M$  scale as outlined in this study. Their assertion of adequacy is based on the introduction of a misprint, which lacks scientific evidence and is irrelevant to the broader discussion. Notably, Eq. (1) of Purcaru and Berckhemer [11] explicitly demonstrates that the  $M_s$  scale becomes unreliable for magnitudes exceeding 7.0.

Our comprehensive analysis, detailed in Appendix A, underscores the limitations of the  $M$  scale and demonstrates the strength of the Das scale ( $M_{wg}$ ) in addressing these shortcomings. The claims by Gasperini and Lolli [21], which fail to adequately consider the well-documented constraints of the  $M$  scale, lack substantive evidence to counter our findings. By providing rigorous revalidation using global datasets and transparent methodologies, our study establishes a scientifically robust case for the adoption of the Das scale ( $M_{wg}$ ) as a superior tool for earthquake magnitude assessment.

## Development of Das Magnitude Scale

Kanamori [3] introduced the moment magnitude scale,  $M_w$ , to address the shortcomings of existing magnitude scales like  $M_L$  and  $M_s$ . This scale was validated for magnitudes  $\geq 7.5$  through comparisons with observed  $M_s$  values, as shown in Table 1 and Table 2 of Kanamori's study. Later, Hanks and Kanamori [1] extended the  $M_w$  scale to lower magnitudes using a similar approach. They validated the extended scale by comparing  $M$  estimates with observed  $M_s$  and  $M_L$



**Fig. 1** We reproduce the plots in Figs. 1 and 4 of Percaru and Berckhemer [11]: **a** Original Fig. 1A of Percaru and Berckhemer [11], **b** Vertical brown line clearly shows that  $\log M_0$  and  $M_s$  shows linear relation, **c**: Original Fig. 4 of Percaru and Berckhemer [11]. **d**: High-

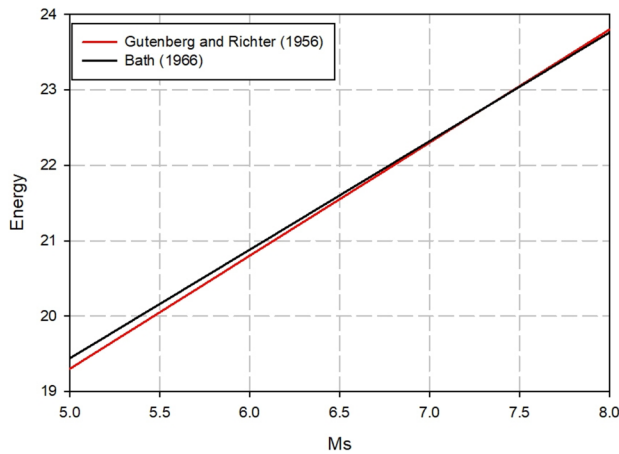
lighted rectangle denotes the  $\log M_0$  and  $M_s$  data as shown in Fig. 4 of Percaru and Berckhemer [11]. Figures 1 and 4 of Percaru and Berckhemer [11] are reproduced with permission from @elsevier 1978

values for Southern California, as detailed in Tables 1 and 2 of their research. However, surface wave magnitudes ( $M_s$ ) are not suitable for intermediate and deeper depth earthquakes, which is a critical limitation raised by Gutenberg and Richter [4]. They suggested using body waves to derive the magnitude scale. To address this, Kanamori [5] proposed

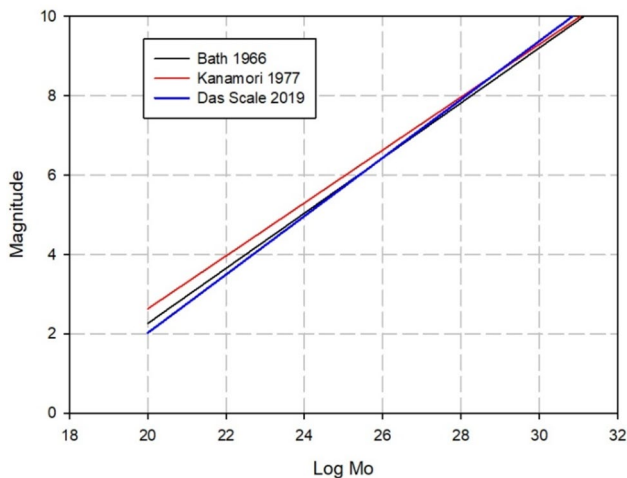
a separate magnitude scale,  $m_w = (\log M_0 - 10.4)/2.4$ , based on the linear equation  $\log M_0 = 2.4 M_B + 10.1$ , specifically for deeper earthquakes.

In summary, the current limitations and critiques of the  $M$  scale strongly warrant the development of a new magnitude scale. It is therefore inappropriate to claim that there is no room





**Fig. 2** Comparison of energy magnitude equations of Gutenberg and Richter [4] with Bath [23]. The figure illustrates the logarithmic relationship between earthquake magnitude and energy release as defined by Gutenberg and Richter's equation and Bath's equation



**Fig. 3** Comparison of Kanamori  $M_w$  scale, Bath [23]  $M_{w-bath}$ , and Das Scale. The figure illustrates the logarithmic relationship between earthquake magnitude and energy release as defined by Gutenberg and Richter's equation and Bath's equation

for the development of a new earthquake magnitude scale, as recent literature suggests [21]. The development of  $M_L$ ,  $M_s$ ,  $m_b$ ,  $M_w$  and  $M$ , further demonstrates the ongoing evolution and refinement of earthquake magnitude scales. Each iteration aimed to address limitations in the previous scale, reflecting the scientific community's commitment to a more accurate understanding of earthquakes. The very existence of the Choy and Boatwright [6] energy equation ( $M_e$ ) serves as a testament to the need for further development. It emerged because the  $M_w$  and or  $M$  scale, while valuable, is not the ultimate answer. Therefore, ongoing research into a new earthquake magnitude scale based on body waves as suggested by Gutenberg and Richter [4] demonstrates

a healthy scientific response to the complexities of measuring events.

Given the well-documented limitations of existing magnitude scales like  $M$  and  $M_w$ , there is a critical need for a scale that incorporates global tectonic factors and is based on body waves for more accurate earthquake measurement. This necessity has driven the development of the Das Magnitude Scale ( $M_{wg}$ ) by Das et al. [8], offering a more comprehensive approach to seismic analysis.

Das et al. [8] adopted a meticulous approach to developing the  $M_{wg}$  scale. They began by assembling a vast instrumental dataset encompassing global seismicity. This dataset included 25,708 seismic moments directly observed from the Global Centroid Moment Tensor (CMT) database and  $m_b$  magnitudes from the International Seismological Centre (ISC) database. Notably, their study incorporated data for shallow, intermediate, and deeper earthquakes.

Recognizing the limitations of  $M_w$  or  $M$  scales, Das et al. [8] aimed to create the  $M_{wg}$  scale. This new scale aligns closely with existing magnitude scales within their respective applicable ranges. To ensure consistency with established techniques used by Hanks and Kanamori [1] and earlier works, Das et al. [8] derived a straightforward least-squares fitting relationship between seismic moment ( $M_0$ ) and  $m_b$  magnitude ( $\log M_0 = 1.36 m_b - 17.24$ ) for the magnitude range of 3.5–7.0 and the timeframe of 1976–2006.

It is important to note that the linear relationship between seismic moment and any magnitude ( $M_s/m_b/M_L$ ) is fundamental in the development of any new magnitude scale. In the case of the  $M_w$  scale [3], this linear relationship ( $\log M_0 = 1.5 M_s + 16.1$ ) was derived through several substitutions and use of constant terms. For development of the energy magnitude ( $M_e$ ), the linear relationship between  $\log M_0$  and surface wave magnitude  $M_s$  was derived directly from the observed dataset, as performed by Das et al. [8].

In developing the  $M_{wg}$  scale, Das et al. [8] derived a linear relationship between  $\log M_0$  and body wave magnitude ( $m_b$ ) using observed seismic moment data, providing a more direct and accurate correlation based on actual measurements. This relationship, expressed as  $\log M_0 = 1.36 m_b + 17.24$ , offers an advantage over the  $M$  scale, which is based on the equation  $\log M_0 = 1.5 M_s + 16.1$ . Unlike the  $M$  scale, which was derived through several steps and involved constant terms ( $E/M_0 = 5 \times 10^{-5}$ ), Das et al.'s approach is more straightforward and precise. Das Scale was derived from simple calculation between  $\log M_0$  and  $m_b$  following the theoretical recommendation of Kanamori and Anderson [22].

The core principle behind Das et al.'s [8] approach lies in the relationship between  $m_b$  and  $M_0$ . If  $m_b$  is used in the equation, then  $M_0$  should also be constrained. By treating  $M_0$  independently, it can be placed on the left side of their equation, leading to the formula  $M_{wg} = \log M_0 /$

1.36–12.68. This formulation provides a magnitude value that avoids saturation, a potential issue with other scales. Developing a magnitude scale ( $M_{wg}$ ) based on body waves instead of surface waves has been a long-term demand initially raised by Gutenberg and Richter [4].

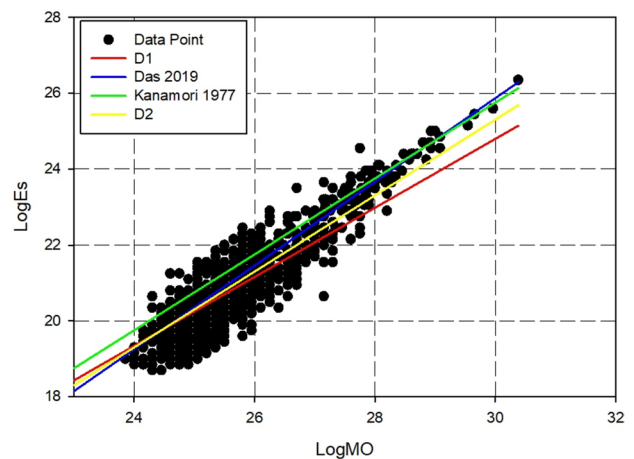
Furthermore, the correlation of seismic energy with the  $M_{wg}$  scale is explained through the well-established Gutenberg energy equation ( $\log E_s = 1.5 M_s + 11.8$ ). Finally, the validation of the  $M_{wg}$  scale follows the procedure outlined by Kanamori [3], ensuring adherence to established practices.

Through a comprehensive analysis, the advantages of the  $M_{wg}$  scale over the  $M_w$  and  $M$  scales became evident:

- Hanks and Kanamori [1] applied Eq. (1) from Purcaru and Berckhemer (11) within the range  $5.0 \lesssim M_s \lesssim 7.5$  for development of  $M$  scale. However, Purcaru and Berckhemer (11) specifically formulated their equation for the range  $5.0 \leq M_s \lesssim 7.0$ , acknowledging the reliability of  $M_s$  only up to 7.0. This reveals a significant limitation of the  $M$  scale, which restricts its applicability to seismic events with  $M_s$  values  $\leq 7.5$ . In contrast, the  $M_{wg}$  scale does not have this restriction.
- In contrast to  $M$ , which is limited to shallow earthquakes, the utilization of  $m_b$  in the  $M_{wg}$  scale allows for the measurement of earthquakes across all depths. This broader scope enhances the applicability and accuracy of the  $M_{wg}$  scale.
- In contrast to  $M_w$  or  $M$  which encounters limitations due to the assumption of a constant term ( $E/M_0 = 5 \times 10^{-5}$ ) in the Gutenberg Energy Equation ( $\log E_s = 1.5 M_s + 11.8$ ),  $M_{wg}$  avoids such constraints. By utilizing 25,708 instrumentally recorded global seismic moments,  $M_{wg}$  overcomes issues related to constancy, ensuring a more accurate representation of seismic magnitudes while accommodating regional and tectonic variations.
- Unlike  $M$ , which primarily relies on surface wave magnitudes,  $M_{wg}$  is derived from observed  $M_0$  and  $m_b$ , capturing the initial cycles of P waves. This approach allows  $M_{wg}$  to establish a closer correlation with both low- and high-frequency spectra of seismic signals. Consequently,  $M_{wg}$  offers enhanced insights into seismic damage potential, making it invaluable for assessing the broader impact of earthquakes beyond their magnitude alone.
- In contrast to  $M$ ,  $M_{wg}$  shows a remarkable alignment in applicable ranges with observed  $m_b$  and  $M_s$  values across all earthquake depths. Comparisons with  $M_E$ ,  $M$ , and observed radiated energy consistently underscore  $M_{wg}$ 's superior agreement. Notably, 75% of estimated energy values using  $M_{wg}$  closely match observed radiated energy values, surpassing  $M$ . This heightened

accuracy in energy magnitude representation significantly enhances the utility of the  $M_{wg}$  scale in earthquake size estimation.

- Enhanced Alignment:  $M_{wg}$  emerges as a superior magnitude scale compared to  $M$ , particularly in terms of alignment with old magnitude scales ( $M_L$ ,  $M_s$ ) and seismic event representation across various magnitude ranges. Figures 1, 2, and 3 of Das et al. [8] vividly illustrate the close alignment of  $M_{wg}$  with observed magnitude scales, contrasting with  $M$ . Notably,  $M_{wg}$  exhibits minimal deviation from observed old magnitude scales, statistically emphasizing its superior accuracy over  $M$ . The disparity between  $M_{wg}$  and observed magnitude scales is notably lower than that between  $M$  and the observed scales, with  $M$  tending to overestimate smaller and intermediate earthquakes. This discrepancy is evident in Fig. 3 and Table 2 of Das et al. [8], highlighting  $M_{wg}$ 's reliability over  $M$ , especially for smaller and medium earthquakes at the global level. In terms of seismic radiated energy  $E_s$ ,  $M_{wg}$  proves more reliable than  $M$ , as depicted in Fig. 4 of Das et al. [8] and Fig. 1 of Das et al. [9]. Further analysis reveals significant disparities between observed  $m_b$  and estimated  $M$ , whereas  $M_{wg}$  demonstrates a considerably smaller average difference when compared to observed  $m_b$ . Similarly,  $M_{wg}$  showcases a closer alignment with observed  $M_s$  compared to  $M$ . The robust connection of  $M_{wg}$  with  $M_E$  as addressed in Das et al. [8], further underscores its enhanced accuracy in representing seismic energy.
- The  $M_{wg}$  Scale: A More Accurate Measure of Earthquake Energy Release



**Fig. 4** Radiated Energy as a Function of Seismic Moment This figure depicts the radiated energy ( $E_s$ ) of a global earthquake dataset plotted against its seismic moment ( $M_0$ ). The radiated energy values are estimated using both the Moment Magnitude ( $M$ ) scale (represented by the Green solid line) and the proposed  $M_{wg}$  scale (represented by the blue solid line). Both estimations are derived from the Gutenberg-Richter energy equation:  $\log(E_s) = 1.5 * M_s + 11.8$

The most reliable method for understanding energy connection for a magnitude scale is to employ energy magnitude relations (e.g.,  $\text{Log } E = 1.5 M_s + 11.8$ ). One of the foundational energy magnitude relations is Gutenberg Energy magnitude relation  $\text{Log } E_s = 1.5 M_s + 11.8$ . However, Bath [23] derived an alternative equation relating energy and magnitude as:

$$\text{Log } E_s = 12.24 + 1.44 M_s \quad (1)$$

Both equations exhibit nearly similar relationships (above  $M_s$  6.5) between energy and magnitude (Fig. 2).

Above discussion indicates  $M_w$  and or  $M$  is a more precise energy representation for large and very large earthquakes (Fig. 3). However, Choy and Boatwright [6] demonstrated that  $M_w$ , in conjunction with the Gutenberg-Richter equation, overestimates radiated energy for magnitude values in the smaller and medium magnitude ranges.

For further understanding the  $M_w$  [3] scale's accuracy in representing earthquake energy release on a global scale for large earthquakes, we utilized radiated energy data from two sources. The first dataset, compiled by Choy and Boatwright [6] and another from Purcaru and Berckhemer [11]. The dataset of Purcaru and Berckhemer [11] encompasses a broader historical period, ranging from 1857 to 1977, and includes large earthquakes. Figure 4 indicates that for large and very large earthquakes, characterized by higher  $\log M_0$  values, the observed energy ( $\log E_s$ ) closely matches the  $M_w$  scale (green line). In contrast, for smaller and medium-sized earthquakes, with lower and medium  $\log M_0$  values, the  $M_w$  scale tends to overestimate the actual energy released. This limitation of  $M_w$  scale is well-documented, as Choy and Boatwright [6] highlighted similar overestimations when using the Gutenberg Energy equation. Figure 4 demonstrates that the  $M_{wg}$  scale shows a stronger overall correlation between the observed energy ( $\text{Log } E_s$ ) and the  $M_{wg}$  scale (blue line) across all earthquake sizes, encompassing small, medium, and large events.

The overestimation of radiated energy seen with the  $M$  scale (in the range  $< 7.5$ ) can be reduced by using the  $M_{wg}$  scale. Since the  $M_{wg}$  scale is derived from a more direct measure of the seismic moment ( $M_0$ ), which closely reflects the actual energy released during an earthquake, it provides a stronger correlation with energy compared to the  $M_w$  scale. While the  $M_w$  scale is widely adopted, it tends to be less accurate for smaller and intermediate earthquakes because it depends on long-period surface waves (not directly connected to source).

## Validity of the Moment Magnitude Scales ( $M$ , $M_w$ , $M_{wg}$ ) in Global Tectonics and its Alignment with Existing Magnitude Scales ( $m_b$ , $M_s$ )

One of the fundamental criteria in the development of magnitude scales is that it should align as closely as possible with the older magnitude scales in their applicable ranges. Different magnitude scales such as  $m_b$ ,  $M_s$ ,  $M_w$ ,  $M$  and  $M_{wg}$  should closely coincide to ensure consistency in measuring the radiated energy from earthquakes. The close coincidence between  $M$  and previous magnitude scales (e.g.,  $M_s$ ,  $M_L$ ) is crucial because it ensures continuity and consistency in the seismic size in terms of energy. Kanamori [3] compared  $M_w$  with global  $M_s$  (Tables 1, 2 of [3]). When Kanamori [3] founds  $M_w$  and  $M_s$  are closely related then he asserted  $M_w$  is a natural continuation of  $M_s$ . Hanks and Kanamori [1] also compared  $M$  with  $M_L$  and  $M_s$  (see Table 1 and Table 2 of [1]). Choy and Boatright [6] also compared  $M_e$  with  $M_w$  and  $M_s$ . The similarities between older magnitude scales (within their applicable range) and  $M$  demonstrate how closely these scales align with the measurement of seismic energy. For instance, using significantly different  $M_s$  and  $M$  values in energy calculations ( $\text{Log } E_s = 1.5 M_s + 11.8$ ) can produce different energy estimates for the same earthquake. Therefore, when a new magnitude scale is developed, it is crucial that it closely aligns with existing older scales within their applicable ranges to ensure consistency and accuracy in seismic hazard assessments.

The body wave magnitude is effective up to magnitudes near 7.0 [1], and the surface wave magnitude is applicable up to around 8.3 [1]. However, a comprehensive scale is needed that can accurately reflect seismic activity across the entire spectrum of earthquake sizes, from smaller to medium and higher magnitudes. In deriving,  $M_w$  scale, Kanamori [3] compared  $M_w$  with the existing  $M_s$  scale using global data for range  $\geq 7.5$ . This comparison was crucial to establish whether  $M_w$  could reliably replace  $M_s$ , providing a consistent measure of earthquake magnitude. Hanks and Kanamori [1] further validated the  $M$  scale by comparing it with the local magnitude scale ( $M_L$ ) and Surface wave magnitude ( $M_s$ ) using data from Southern California. These comparisons were essential to determine if  $M$  closely coincided with the traditional scales. Das et al. [8] statistically and numerically demonstrated how  $M_{wg}$  scale is closely connected with observed existing magnitude scales ( $m_b$ ,  $M_s$ ). The  $M_w$  or  $M$ , while effective for large earthquakes, encounters limitations when applied to smaller and medium events. Its applicability range is skewed towards higher magnitudes, hindering direct comparisons and accurate size estimation for smaller and medium earthquakes. To address this issue again, we revalidated the scale using global datasets from the period 1976–2014 obtained from the Global Centroid Moment Tensor (GCMT) and the International Seismological Centre



(ISC). In our analysis, we considered 39,792  $m_b$  values from ISC and corresponding seismic moments from GCMT, as well as 21,256  $M_s$  values from ISC and GCMT.

Our investigations clearly demonstrate (Fig. 5), as previously established by Das et al. [8], that  $M$  deviates significantly from  $m_b$  and  $M_s$  magnitudes within their applicable ranges. When comparing the alignment of amplitude-based scales like  $m_b$  and  $M_s$  with  $M$  and  $M_{wg}$ , the  $M_{wg}$  scale shows statistically significant closeness to  $m_b$  and  $M_s$ . As Fig. 5 shows,  $M$  aligns well with larger magnitudes (around 7.5 and above), which aligns with the range for which it was originally developed. However, for smaller and medium earthquakes,  $M$  deviates significantly from the observed trends of  $m_b$  and  $M_s$  data (refer to Figs. 5a, 2b).

All discussion reinforces the notion that the  $M$  scale is most appropriate for measuring large earthquakes exceeding 7.5, as initially suggested by Kanamori [3]. The  $M$  scale exhibits a tendency to significantly overestimate seismic moment values for earthquakes with surface wave magnitudes ( $M_s$ ) below 6.5 and body wave magnitudes ( $m_b$ ) across their applicable range, typically up to 7.0 [1]. However, the  $M_{wg}$  scale offers a potential solution by mitigating this overestimation for  $M_s$  values below 6.5 and  $m_b$  values up to 6.0.

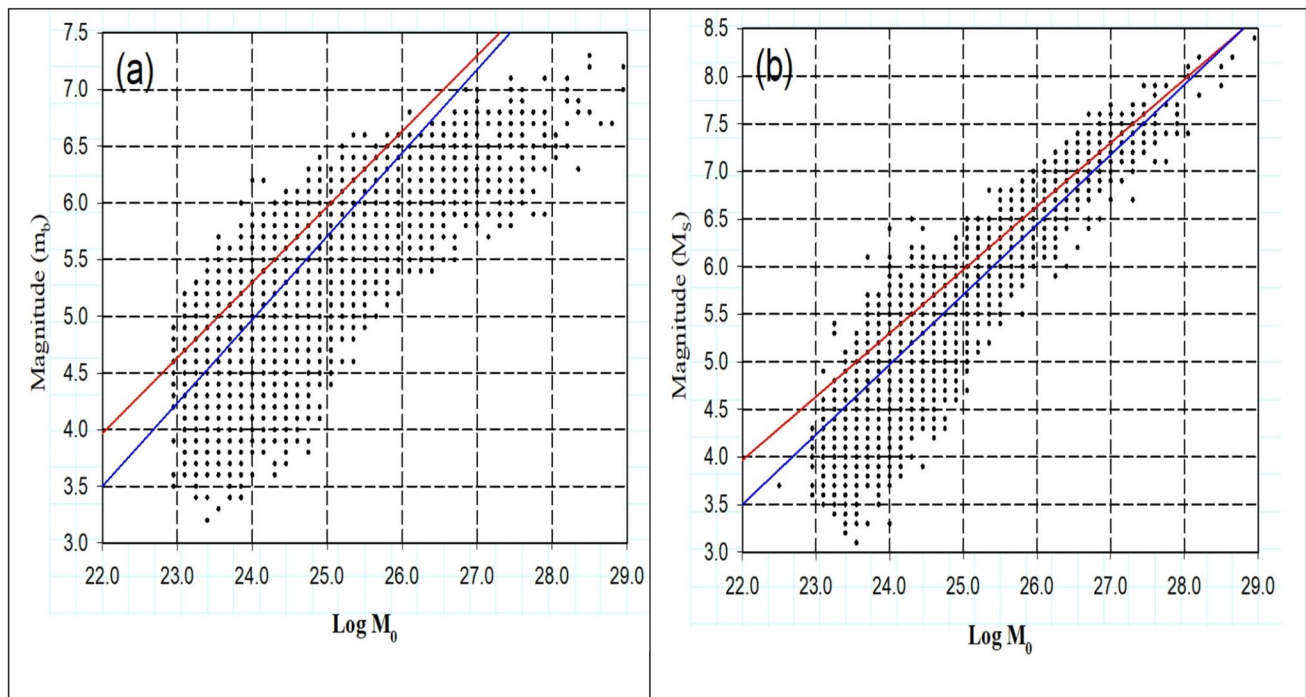
To statistically assess the differences between  $M$  and  $M_{wg}$  estimates, a null hypothesis ( $H_0$ ) is established, assuming both datasets share the same mean and variances. This hypothesis is rejected at confidence levels of 0.05 and 0.025 for  $m_b$  values between 4.5 and 5.5, and  $M_s$  values between

5.5 and 7.5 [8]. This rejection indicates statistically significant differences between  $M$  and  $M_{wg}$  estimates within these magnitude ranges. Consequently, the  $M_{wg}$  scale emerges as a single, uniformly valid magnitude scale applicable for earthquakes with magnitudes exceeding 4.5 [8].

Furthermore, the validation extends beyond  $m_b$  and  $M_s$ . To assess the similarity of  $M_{wg}$  and  $M$  with other observed magnitude scales, a globally distributed dataset of 1,361 energy magnitude ( $M_e$ ) events from 1995 to 2007 was utilized [8]. A comparison of  $M_e$ ,  $M$ , and  $M_{wg}$  revealed a noteworthy pattern: 53% of  $M_{wg}$  values were closer to  $M_e$  than  $M$ . This finding strengthens the case for  $M_{wg}$ , suggesting it aligns more closely with another independent measure of earthquake energy [8].

The strong correlation between  $M_{wg}$  and  $m_b$ ,  $M_s$ , and  $M_e$  indicates that  $M_{wg}$  is a more reliable measure than  $M$ . This is because  $M_{wg}$  closely aligns in the applicable ranges with scales ( $m_b < 7.0$ ,  $M_s < 8.3$ ) derived from direct amplitude measurements. While  $m_b$  and  $M_s$  are suitable for certain earthquake magnitudes,  $M_{wg}$  is particularly effective for large earthquakes and demonstrates a high correlation with  $m_b$  and  $M_s$  within their respective applicable ranges. Therefore, the close correspondence between  $M_{wg}$  and these other scales (in their applicable ranges) is critically important.

Hanks and Kanamori [1] derived the equation for  $M$  ( $M = 2/3 * \log M_0 - 10.7$ ) using the CGS (centimeter-gram-second) system, reporting results to one decimal



**Fig. 5** Comparison of  $M$  scale (red line) and  $M_{wg}$  (blue line) with observed **a:**  $m_b$  and **b:**  $M_s$

place. Kanamori [3] derived  $M_w$  by giving equation  $\log W_0 = 1.5 M_s + 16.1$ , where  $W_0 = M_0 / (2 \times 10^2)$ . The difference in the constant term between the two equations ( $M_w = (\log W_0 - 16.1) / 1.5$ ,  $M = 2/3 \log M_0 - 10.7$ ) leads to a rounding error of up to 0.03. Notably,  $M = 2/3 \log M_0 - 10.7$  is widely used and serves as a standard in global databases, including GCMT (Global Centroid Moment Tensor) and other regional seismic monitoring agencies. Thus, it is inaccurate to assert that  $M_w = (\log W_0 - 16.1) / 1.5$  and  $M = 2/3 \log M_0 - 10.7$  yield identical estimates. This discrepancy arises from the constant term difference, and it is crucial to note that both forms of  $M$  will provide the same estimates only when using the alternate definition of  $M = 2/3 \log M_0 - 10.73333$ . The selection of the constant term in the  $M$  scale, specifically 10.7, is crucial for maintaining precision and consistency with the Eqs. (3, 4, 5, 6, and 7) established by Hanks and Kanamori [1]. It is important to note that all the equations used in their work adhered to a consistent level of precision, which is why the constant term in the  $M$  scale was also set to 10.7 to match this precision.

To shed light on the discrepancies between  $M_e$  and  $M$ , we will scrutinize the equations formulated by Choy and Boatright [6] and the Gutenberg-Richter energy equation, both expressed within the CGS system.

$$\log E = 1.5 M_s + 11.4 \quad (\text{with } E \text{ in ergs}) \quad (2)$$

$$\log E = 1.5 M_s + 11.8 \quad (\text{with } E \text{ in ergs}) \quad (3)$$

Equation (2) can be redefined to derive a new magnitude expression, denoted as  $M_e$  (Choy and Boatright, 2006):

$$M_e = \frac{\log E - 11.4}{1.5} \quad (4)$$

$$M_e = \frac{\log E - 11.4 + .4 - .4}{1.5} \quad (5)$$

$$M_e = M + 0.27 \quad (6)$$

Equation (6) shows that  $M$  and  $M_e$  are not equivalent, as they are derived from different physical principles. This distinction was emphasized by Choy and Boatright (2006) and Bormann and Giacomo [20]. Choy et al. [24] also pointed out that there is no inherent reason for  $M_e$  and  $M$  to be the same. Since the  $M$  scale is defined based on the ratio  $E_s/M_0 = 5 \times 10^{-5}$ , it differs from  $M_e$ . The disparity between  $M_e$  and  $M$  stems from their basis in distinct fundamental properties of seismic events.

A comprehensive analysis of 1385 worldwide energy data points revealed that 87% of the data deviated from the observed  $M$  values, underscoring the intrinsic dissimilarity

between  $M$  and  $M_e$ . Moreover, 21% of observed  $M_e$  data deviates by 0.5 or more magnitude units from  $M$ . The maximum difference between  $M$  and  $M_e$  can reach up to 1.2 magnitude units. For instance, the earthquake on September 9, 2005, was reported with an  $M$  of 7.4, while the  $M_e$  was reported as 6.2.

The fundamental equation for the  $M_{wg}$  scale, as introduced by Das et al. [8], links seismic moment ( $\log M_0$ ) with body wave magnitude ( $m_b$ ), differing from the traditional relationship between  $\log M_0$  and Surface wave magnitude ( $M_s$ ) established by Kanamori [3]. The equation is as follows:

$$\log M_0 = 1.36 m_b + 17.24 \quad (7)$$

Considering Eqs. (5) and (7), the relationship between energy magnitude ( $M_e$ ) and  $M_{wg}$  can be formulated in terms of the slowness parameter ( $\Theta$ ) as follows [20]:

$$M_e = \frac{1.36}{1.5} M_{wg} + \frac{5.44 + \Theta}{1.5} \quad (8)$$

From Eq. (8), it becomes evident that the  $M_{wg}$  scale is not identical to the energy magnitude ( $M_e$ ) because both scales are based on different physical properties. This distinction emphasizes the importance of understanding and acknowledging the unique characteristics inherent in each seismic magnitude scale.

Upon scrutinizing 1385 observed radiated energy data, it is revealed that among these earthquakes, 730 events exhibited a closer alignment of  $M_{wg}$  with  $M_e$  compared to  $M$ . This implies that 53% of the data with  $M_{wg}$  demonstrated a closer agreement with  $M_e$  than  $M$  [8]. The same conclusion can be reached by using the dataset from Choy and Boatright [6] as demonstrated in Das et al. [8]. Hence, it is crucial to emphasize that  $M_{wg}$  exhibits a superior agreement in terms of energy magnitude ( $M_e$ ).

## Linearity Issue in Magnitude Scale Development

Most earthquake magnitude scales rely on a technique called linear least squares regression. This method assumes a straight-line relationship between two variables, such as seismic moment ( $M_0$ ) and a specific magnitude (e.g.,  $M_s$ ,  $m_b$ ,  $M_w$ ,  $M$  etc.). Kanamori [3] utilized a linear least squares relationship between  $\log M_0$  and  $M_s$  to develop the  $M_w$  scale. Hanks and Kanamori [1] also employed the linear least square method between  $\log M_0$  and  $M_s$  for the development of  $M$  scale. Linear least square relationship was the standard practice for developing relationships between  $\log M_0$  and magnitudes (e.g., Thatcher and Hanks [18], [1, 25]). Hanks and Kanamori [1] used linear least square relationship between  $\log M_0$  and  $M_L$ . As  $M_L$  and  $m_b$  are

based on 1 s amplitudes, therefore, Hanks and Kanamori [1] considered the applicable range for  $M_L$  and  $m_b$  up to 7.0.

Similarly, the energy magnitude scale ( $M_e$ ) was derived by Choy and Boatwright [6] using linear least squares regression. More than 90% of the seismic literature advocates the linear relation between  $\log M_0$  and magnitudes. However, a recent study by Gasperini and Lolli [21] has raised concerns about the validity of using linear regression between  $\log M_0$  and body wave magnitude ( $m_b$ ). It is important to note that Gasperini and Lolli [21] overlooked the fact that most magnitude scales (e.g.,  $M_w$ ,  $M$ ,  $M_e$ ,  $M_{wg}$ ) were based on linear least squares regression. While these authors (Gasperini and Lolli) conducted in their earlier studies (e.g., [25]) linear regression between  $M_w (=2/3 \log M_0 - 10.7)$  and  $m_b$ , they raised questions about the linear regression performed by Das et al. [8] between  $\log M_0$  and  $m_b$ . Therefore, we will briefly explain this issue for the benefit of the reader. While deriving the  $M$  scale, Hanks and Kanamori [1] utilized a linear equation between  $\log M_0$  and  $M_s$  and in similar way Choy and Boatright [6] utilized linear least square equation between  $\log M_0$  and  $M_s$  for the range 5–8.6. Following the historical practices, Das et al. [8] utilized a linear equation between  $\log M_0$  and  $m_b$  for the range 3.0–7.0. Hanks and Kanamori [1] advocated to use  $m_b$  up to 7.0 and  $M_s$  up to 8.3. Note that  $m_b$  and  $M_L$  are measured on 1 s and are considered applicable up to 7.0. Hanks and Kanamori [1] utilized the linear least square relation between  $\log M_0$  and  $M_L$  in the range 3.0–7.0. However, a recent study by Gasperini and Lolli [21] specifically targets the validity of linear least squares regression for  $M_{wg}$  scale, while overlooking its established use in developing other magnitude scales, such as those proposed by Hanks and Kanamori [1] and Choy and Boatright [6].

Furthermore, Gasperini and Lolli themselves utilized linear relationships between  $M_w$  and  $m_b$  (e.g., [25]), considering  $m_b$  values greater than 7.0. Since the  $M_w$ ,  $M$ , and  $M_e$  scales are fundamentally based on linear least squares, applying linear least squares in the development of the  $M_{wg}$  scale is appropriate. Additionally, Das et al. [8] clearly outlined the rationale for adopting linear least squares in the formulation of the  $M_{wg}$  scale.

To support nonlinear relationships between  $\log M_0$  and  $m_b$ , Gasperini and Lolli [21] mischaracterized the findings of Kanamori and Anderson [22]. Therefore, we will clarify this issue with direct reference to Kanamori and Anderson's [22] original work. From theoretical considerations utilizing the viscoelastic model proposed by Haskell [26], Kanamori and Anderson [22] proposed three distinct slopes for the relationship between  $\log M_0$  and  $M_s$  (20 s). This conceptual framework applies across the entire magnitude spectrum, encompassing very small, small, medium, large, and very large earthquakes. Therefore, it is affirmed that employing a

single linear relationship is not suitable for representing the complete magnitude range (–infinity to infinity, e.g., 0–8). According to Kanamori and Anderson [22], the expected scaling coefficient between the logarithm of scalar seismic moment ( $M_0$ ) and  $M_s$  is approximately 1 for very small earthquakes, roughly 3/2 not for very large earthquakes (i.e., applicable for medium and large ones), and about 3 for the very large earthquakes (see abstract of Kanamori and Anderson [22] “*Except for very large earthquakes, the relation  $M_s \sim (2/3) \log M_0 \sim 2 \log L$  is established by the data*”, “*For very large earthquakes  $M_s \sim (1/3) \log M_0 \sim \log L \sim (1/3) \log E$  s. For very small earthquakes  $M_s \sim \log M_0$ ,  $\sim 3 \log L \sim \log E$  s.*”; page 1083 of Kanamori and Anderson [22], “*We have discussed only surface-wave magnitude but the same arguments hold for body-wave magnitude except that characteristic times are now to be compared with 1 s rather than 20 s. In this case, only very small earthquakes can be expected to scale as  $m_b \sim \log L$  3 and more earthquakes can be expected to satisfy the  $m_b \sim \log L$ .*”). Kanamori and Anderson [22] further assert that similar arguments apply to the relationship between  $\log M_0$  and  $m_b$  (1 s), indicating that a single linear relation is not viable for the entire range of body wave magnitudes, encompassing very small, small, medium, large, and very large earthquakes.

In the context of  $\log M_0$  and  $m_b$ , with a slope of around 3 for very large earthquakes ( $> 8.0$ ), around 3/2 not for very large earthquakes (i.e., medium and large earthquakes), and around 1.0 for very small earthquakes. The applicability of the linear relationship with a slope of 3 for very large earthquakes is not observed in  $m_b$  (1 s). Thus, justifying a linear relationship between  $\log M_0$  and  $m_b$  in the range of  $3.5 \leq m_b \leq 7.0$  aligns with the theoretical model of Kanamori and Anderson [22], as magnitudes less than  $m_b < 3$  can be considered very small earthquakes. It is, nevertheless, considered inappropriate to establish linear relationships across the entire magnitude range, covering very small, small, medium, large, and very large earthquakes, based on the considerations of Kanamori and Anderson [22].

Local magnitude  $M_L$  and teleseismic body wave magnitude  $m_b$  are determined from the amplitude of waves with a period of about 1.0 s. Thatcher and Hanks (1972) derived linear relationships between  $\log M_0$  and  $M_L$  for the magnitude range 3–7.0. Consequently, in deriving the moment magnitude ( $M$ ) scale, Hanks and Kanamori [1] opted to utilize the relationship of Thatcher and Hanks (1972) between  $\log M_0$  and  $M_L$  for magnitudes ranging from 3 to 7.0. Several authors (e.g., [25, 27]) also admitted that relationship between moment and  $m_b$  is also linearly related lower than the minimum current  $m_b$  determination. A large number of studies (e.g., [28], Borman and Yadav [2], [25, 27],) have explored the linear regression between  $M$  and  $m_b$ . Bormann et al. [28] established a linear relationship between

$M$  and  $m_b$  within the magnitude range 4.4–7.6. Borman and Yadav [2] also performed linear relationship between  $M$  and  $m_b$  within the magnitude range 4.4–7.6. Similarly, Scordilis [29] derived a linear relationship between  $M$  and  $m_b$  for the magnitude range  $3.5 \leq m_b \leq 6.2$ . Das et al. [30] also established a linear relationship within the magnitude range 2.9–6.5. Gasperini et al. [25] demonstrated a linear relationship between  $M$  and  $m_b$  greater than 7.0, as evident in their Figs. 1 and 2. The development of a linear relationship (in the range of medium and large, and small  $m_b$  earthquakes) between  $\log M_0$  and  $m_b$  aligns with the theoretical considerations of Kanamori and Anderson [22]. However, there is no physical reasoning why a non-linear relationship should exist between  $\log M_0$  and  $m_b$  in the medium, large and smaller  $m_b$  ranges.

Das et al. [8] established a linear relationship between  $\log M_0$  and  $m_b$ , grounded in the IASPEI [31] rule and the theoretical framework presented by Kanamori and Anderson [22], aligning with the findings of various researchers [1], Borman et al. [28], Wason et al. [32], [25], see page 932 of [25], “Almost all papers cited above used linear regression laws between  $M_w$  and  $m_b$ ”, [9, 33, 34].

### Critical Analysis of Gasperini and Lolli [21]

Recently, Gasperini and Lolli [21] claimed that the  $M$  scale is adequate, basing their argument on an alleged misprint in the caption of Table 1 in Purcaru and Berckhemer [11]. By introducing the misprint issue Gasperini and Lolli [21] distorted the original definition of ME of Purcaru and Berckhemer [11]. The misprint issue has been thoroughly discussed in the Appendix. It is scientifically and technically incorrect to assert that the  $M$  scale is adequate across all magnitudes and depths; this issue needs to be addressed to prevent bias in seismic energy budgeting. Gasperini and Lolli [21] questioned the validity of the basic assumptions of  $M_{wg}$  scale, specifically the application of linear least squares regression. However, as discussed above, we have demonstrated how linear least squares regression is indeed appropriate for developing magnitude scales. Note that linear least square regression was the core in building the  $M$ ,  $M_w$ ,  $M_c$  scales.

Gasperini and Lolli [21] argued that one could derive an infinite number of magnitude scales based on their chosen criteria, and they introduced hypothetical lines D1 (equation 15 of Gasperini and Lolli [21]) and D2 (equation 16 of Gasperini and Lolli [21]), claiming that these produce lower Absolute Average Deviation (AAD) and standard deviations compared to the  $M_{wg}$  scale. They attempted to imply that derivation of the  $M_{wg}$  scale is just one among countless possibilities. However, this argument is based on incorrect facts (e.g., use constant term  $-4.7$  instead of correct value of  $-4.8$ ) overlooks the critical limitations of the  $M$  and or  $M_w$  scale, particularly for measuring smaller and

medium earthquakes, especially those below magnitude 7.5. Das and colleagues along with many researchers have demonstrated that the  $M$  scale is inadequate for these cases due to its inherent assumptions and constraints. The development of the  $M_{wg}$  scale specifically addressed these shortcomings, making it a more accurate and physically consistent alternative. Thus, the suggestion that an infinite number of magnitude scales can be defined dismisses the importance of addressing these known deficiencies and ignores the need for empirically validated and practically useful seismic measures.

The hypothetical lines D1 and D2, introduced by Gasperini and Lolli [21] in an attempt to invalidate the Das magnitude scale, were not derived following the standard procedures typically used for developing a magnitude scale. Specifically, the hypothetical line D2 exhibits a significant difference of 0.4 magnitude units (m.u.) across the higher magnitude range. This discrepancy arises from the use of a constant term of 4.8 instead of 4.3, as originally proposed by Kanamori [3]. The 0.4 m.u. difference for large and very large earthquakes ( $\geq 7.5$ ) undermines the close alignment with  $M_s$  data, which served as the basis for the development of the  $M_w$  scale by Kanamori [3].

D1 and D2 hypothetical lines did not adhere to the established methods used for scales such as  $M_w$ ,  $M_c$ , and  $M_{wg}$ . There are key differences between the hypothetical lines and the original magnitude scales (i.e.,  $M_w$ ,  $M$ ,  $M_c$  and  $M_{wg}$ ). Kanamori's  $M_w$  scale, derived from surface wave magnitude, is primarily applicable to shallow earthquakes. Das's  $M_{wg}$  scale, on the other hand, utilizes body wave magnitudes, allowing it to function across the entire earthquake depth range. Additionally, the  $M$  scale was developed and validated for a specific region (Southern California), whereas the  $M_{wg}$  scale is formulated based on global data, offering a more comprehensive and consistent application. Importantly, the development of the  $M_{wg}$  scale addressed the limitations of the  $M$  scale. While energy comparison played a role in the  $M_{wg}$  scale's development, it wasn't the sole focus. Thus, hypothetical lines D1 and D2 don't meet the essential criteria for a new magnitude scale.

Furthermore, the evaluation criteria for D1 and D2 did not include any statistical tests, such as a null hypothesis, to determine their significance. Assessing their adequacy solely based on Absolute Average Deviation (AAD) and standard deviation does not provide a thorough comparison. In contrast, during the development of the  $M_{wg}$  scale, statistical significance was assessed using a null hypothesis to validate the results. The absence of rigorous statistical testing in the work of Gasperini and Lolli [21] undermines the validity of their arguments.

As illustrated in Fig. 4, the hypothetical lines D1 and D2 introduced by Gasperini and Lolli [21] significantly deviate in large magnitude ranges from the  $M$  and  $M_{wg}$  scales. While these lines align well with smaller and medium earthquakes, their divergence becomes pronounced for larger events. Although the hypothetical lines D1 and D2 appear reasonable in the small to medium magnitude range, they show



significant deviations in the large and very large magnitude ranges. A fundamental criterion for a new magnitude scale is its alignment with existing scales (eg,  $m_b$ ,  $M_s$ ). Both D1 and D2 fail to meet this criterion, exhibiting significant deviations from observed  $M_s$  magnitudes ( $>7.0$ ).

Furthermore, the constancy issue inherent in Hanks and Kanamori's [1] scale persists in D1 and D2. However, it's important to recognize that similar arguments could be extended to other magnitude scales, such as  $M_w$ ,  $M$  and  $M_e$ . If Kanamori [3] and Hanks and Kanamori [1] had used Bath's equation (Eq. 1) instead of Gutenberg Richter energy equation  $\log E = 1.5 M_s + 11.8$ , the formula for the moment magnitude ( $M_w$ ) would have been:

$$M_{w-Bath} = \log Mo / 1.44 - 11.625 \quad (9)$$

This illustrates that, just as with the  $M$  scale, it is theoretically possible to derive an infinite number of magnitude scales based on different mathematical approaches or assumptions. Therefore the same argument ("In summary, our exercise teaches us that assumption and arguments leading to  $M_w$  are scientifically unjustified because  $M_w$  is only one of the infinite moment magnitude scales that could be defined and it is not even the best, based on the criteria proposed by its authors", [21]) put forward by Gasperini and Lolli [21] can also be applied to the  $M_w$  and or  $M$ . The same argument can also be applied to  $M_e$  [6] scale if one were to use Bath's equation instead of the Gutenberg Energy equation. Since Kanamori [5] stated that both Bath's equation and the Gutenberg-Richter energy equation are equivalent, it is feasible to use Bath's equation in the development of either the  $M_w$  or  $M_e$  scales. This illustrates that the choice of equations and assumptions in deriving magnitude scales can lead to various formulations, making the argument about the infinite possibilities of Gasperini and Lolli [21] equally valid for  $M_w$ ,  $M$ ,  $M_e$  and  $M_w$  scales. Given the limitations of the  $M_w$  and or  $M$  scales in accurately representing all magnitudes and depths, it is essential to address these issues and consider the development of a new, more comprehensive scale  $M_{wg}$ .

### Comments on Gasperini and Lolli [21]

- (i) Gasperini and Lolli [21] incorrectly used  $\log_{10}(\Delta\sigma/2\mu) = -4.7$ , whereas the correct value, as derived from Choy and Boatwright [6], is  $-4.8$ , based on the equation  $E_s = 1.6 \times 10^{-5} Mo$  (pls see abstract of Choy and Boatwright 1995). The implications of this error are discussed in points: a) *Inaccuracies in Eq. 12 of Gasperini and Lolli [21]*: The value  $-4.7$  used in Eq. (12) of Gasperini and Lolli [21] directly affects the computed relationship between  $M$  and  $M_E$ . Correcting this value to  $-4.8$  shifts the results and changes the interpretation of when  $M$  and  $M_E$  coincide. b) *Hypothetical Line D2 (equation 16 in G&L's Analysis) is incorrect*: The incorrect value for  $\log_{10}(\Delta\sigma/2\mu)$  directly affects the derivation of Line D2 ( $M_{wcb}$  as per [21]), which was calculated based on the erroneous scaling relationship. As a result, this line does not accurately represent the intended relationship, leading to flawed conclusions. c) *Impact on Results and Discussion and Conclusion*: The miscalculation propagates through Gasperini and Lolli [21] analysis, affecting their interpretation of  $M_e$  and  $M$ . The incorrect scaling introduces bias, which undermines the validity of their arguments. Therefore, Table 1 and Fig. 1 of Gasperini and Lolli [21] are incorrect. d) *Incorrect comparisons of  $M_{wg}$  and  $M_{wcb}$* : The incorrect value directly affects the computed  $M_{wcb}$  and related comparisons with other magnitude scales. These comparisons, highlighted in the discussion and conclusions, are, therefore, invalid. Thus, main conclusions of the paper of Gasperini and Lolli [21] are incorrect and invalid.
- (ii) *Incorrect Representation of Eq. (1) of Purcaru and Berckhemer [11]*  
 Gasperini and Lolli [21] incorrectly stated that Eq. (1) of Purcaru and Berckhemer [11] was developed in the range  $5 \leq M_s \leq 7.5$ . In their statement, Gasperini and Lolli [21] write, "whereas Purcaru and Berckhemer [11] from earthquakes with  $5 \leq M_s \leq 7.5$  recorded all over the world obtained  $\log_{10} Mo = 1.5M_s + 16.1$ ." However, Purcaru and Berckhemer [11] explicitly stated that their Eq. 1 is developed to earthquakes with magnitudes up to  $M_s \lesssim 7.0$  (Page 189 of Purcaru and Berckhemer [11] "In the range of moderate to large earthquakes ( $M_s \lesssim 7$ ,  $Mo \lesssim 10^{27}$  dyn.cm), where  $M_s$  is a reliable measure, the relation between  $\log Mo$  and  $M_s$  is linear and average relation see Fig. 4:  $\log Mo = (16.1 \pm 0.1) + 1.5 M_s$ —(1) is established to give the best fit with the observed data"). Therefore, Gasperini and Lolli's reference to  $M_s \leq 7.5$  inaccurately extends the range specified by Purcaru and Berckhemer [11].
- (iii) *Mischaracterization of Purcaru and Berckhemer [11] regarding a printing mistake*: Gasperini and Lolli [21] suggest the presence of a printing error in Purcaru and Berckhemer's work, without providing supporting references. This claim is unfounded and, if accepted, could distort the understanding of the  $M_E$  scale as originally defined by Purcaru and Berckhemer [11]
- (iv) *Misunderstanding the applicability of the  $M_w$  scale*: Gasperini and Lolli [21] fail to distinguish between the  $M$  scale and  $M_w$  scale, which differ by 0.03, and do not discuss the limitations of the  $M_w$  scale for

global events below  $M_s$  7.5. Kanamori [3] and Hanks and Kanamori [1] outlined the range of applicability for the  $M_w$  scale, with Kanamori [3] specifically noting that the  $M_w$  scale was designed as a natural continuation of the  $M_s$  scale for great earthquakes. However, the  $M_w$  scale is not suitable for smaller and medium earthquakes, as illustrated by evidence in Das et al. [8] and further explored in [35, 36], which demonstrate that the  $M_w$  scale overestimates smaller and medium earthquakes.

- (v) *Inconsistency in the relationship between  $\text{Log } M_0$  and  $m_b$* : While Gasperini and Lolli acknowledged a linear relationship between  $M_w$  and  $m_b$  in their earlier publications (e.g., [27]), they later suggest nonlinear relationships between  $\text{Log } M_0$  and  $m_b$  for the Das et al. [8] model. However, we have demonstrated that a nonlinear relationship is not suitable for the 3.0 to 7.0 magnitude range, although nonlinear relations are appropriate for a broader range of magnitudes. We have provided a detailed justification for why a linear relationship is used in the specific range discussed.
- (vi) *Criticism of  $m_b$  values up to 7.0*: Gasperini and Lolli [21] criticized the use of  $m_b$  values up to 7.0 in Das et al. [8], despite employing similar values in their own work (e.g., [27]). This inconsistency may confuse readers. To clarify, we have explained above the rationale behind adopting a linear relationship between  $\text{Log } M_0$  and  $m_b$  for values up to 7.0, based on the foundational work of Kanamori and Anderson [22].
- (vii) *Misinterpretation of Kanamori and Anderson [22]*: Gasperini and Lolli [21] claim that Kanamori and Anderson [22] noted the linearity of the relationship between  $\text{Log } M_0$  and  $m_b$  only for magnitudes below 5.0. However, Kanamori and Anderson [22] discussed

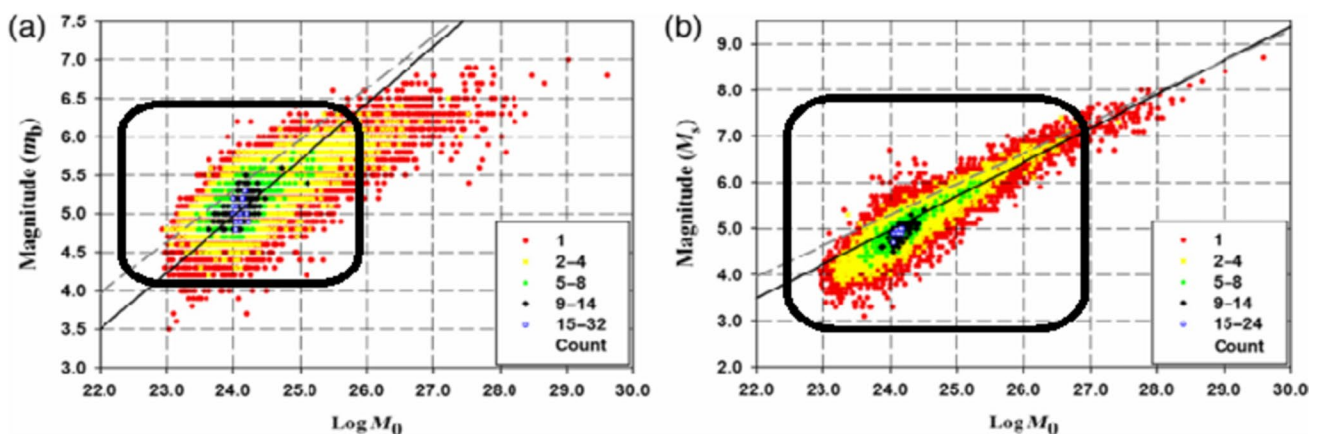
the relationship between  $\text{Log } M_0$  and  $M_s$ , not  $m_b$ . They did not specify a range of  $m_b \leq 5.0$ , which Gasperini and Lolli incorrectly reference. This misinterpretation further misrepresents the work of Kanamori and Anderson [22].

### Why the Old Standard $M_w$ And or $M$ Isn't Enough: The Need for $M_{wg}$

- (1) *Improved Consistency of the  $M_{wg}$  Scale in Reducing Discrepancies with observed  $m_b$  and  $M_s$* : The differences between the  $M$  scale and the other observed scales,  $m_b$  and  $M_s$ , are significantly greater (Null Hypothesis test was performed) compared to the differences between the  $M_{wg}$  scale and  $m_b$  and  $M_s$  (Figs. 1, 2, 3 of [8], Fig. 5, Fig. 6). The average difference between observed  $m_b$  and estimated  $M_{wg}$  is  $0.008 \pm 0.33$ . The average difference between observed  $m_b$  and estimated  $M$  is  $-0.31 \pm 0.30$ . The average difference between observed  $M_s$  and estimated  $M$  is  $-0.43 \pm 0.28$ , whereas the average difference between observed  $M_s$  and estimated  $M_{wg}$  is  $-0.12 \pm 0.26$ .

The  $M_{wg}$  scale's ability to reduce these discrepancies is crucial for achieving a more consistent representation of energy levels across different magnitude scales. Note that this parameter was critical when designing the  $M$  or  $M_w$  scale.

- (2) *Closer Alignment with Seismic Energy Radiation*: It has been demonstrated [8, 9] that the  $M_{wg}$  scale provides a better correlation with seismic energy, making it the preferred choice for use.
- (3) *Resolution of Technical Issues in the  $M$  Scale*: The technical problems associated with the  $M$  scale, arising



**Fig. 6** Comparison between  $M_{wg}$  and  $M$  Scales [8]. **a** A statistically significant deviation is observed between  $m_b$  and  $M$  compared to  $M_{wg}$ , with this deviation being statistically significant. **b** A statistically strong deviation is observed between  $M_s$  and  $M$  compared to  $M_{wg}$

from the misuse of Eq. 1 from Purcaru and Berckhemer [11] for magnitudes below 7.5, are effectively eliminated with the adoption of the  $M_{wg}$  scale.

- (4) *Based on Global Tectonics:* The  $M_{wg}$  scale incorporates a global understanding of tectonic processes, making it more suitable for evaluating earthquakes across different tectonic settings.
- (5) *Fulfills a Long-Standing Demand by Gutenberg and Richter [4]:* The  $M_{wg}$  scale addresses the recommendations of seismology pioneers Gutenberg and Richter [4], who emphasized the importance of developing magnitude scales based on direct measurements of body waves rather than surface waves (*“Determination of magnitude at large distances.—Theoretically, this should be based rather on body waves than on surface waves; unfortunately, readings of amplitude and period for body waves are less generally available”*, page 142 of [4]). This approach ensures greater accuracy in magnitude determination, particularly for large-distance events, and aligns with their call for integrating body wave data into seismic analysis.
- (6) *Addresses the limitations of the  $M_w$  Scale:* It overcomes certain shortcomings of the  $M_w$  and or  $M$  scales, such as inaccuracies in representing smaller and intermediate earthquakes, and provides a more accurate measure of seismic energy across various magnitude ranges.
- (7) *Utilizes Instrumental Observations:* The  $M_{wg}$  scale is derived from directly observed seismic data, ensuring that it is based on actual measurements rather than relying on assumptions or indirect estimations. The  $M_{wg}$  scale aligns closely with the  $m_b$  and  $M_s$  scales within their respective applicable ranges. This is significant because amplitude-based scales, which directly measure seismic signal amplitudes, also show a strong correlation with  $M_{wg}$ , making it a more versatile and beneficial scale for assessing large earthquakes beyond the limitations of  $m_b$  and  $M_s$ .

These characteristics make the  $M_{wg}$  scale a robust and versatile tool for assessing earthquake magnitudes on a global level.

## Conclusions

The earthquake magnitude scale plays a crucial role in earthquake science, making accurate knowledge of this scale essential for various seismological, geophysical, geoscientific and engineering applications. The introduction of the  $M_w$  scale by Kanamori [3] marked a significant advancement in seismology. However, the original scale was not suitable for earthquakes with magnitudes below 7.5. To

address this limitation, Hanks and Kanamori [1] extended the scale to cover magnitudes as low as 3.0 by introducing the formula  $M = 2/3 \log M_0 - 10.7$ .  $M$  and  $M_w$  are distinct magnitude scales, differing by approximately 0.03 due to variations in their definitions and precision. The GCMT or other data agency should refrain from using the term  $M_w$  for magnitudes below 7.5, as Kanamori specifically defined  $M_w$  for earthquakes with magnitudes  $\geq 7.5$ . Furthermore, the  $M$  scale, as defined by Hanks and Kanamori [1], is expressed as  $M = 2/3 \log M_0 - 10.7$ , and is intended for smaller, medium and large magnitudes. This distinction should be maintained to avoid confusion between the two scales. Despite its widespread use, the  $M$  scale has critical limitations and given that it was developed over 50 years ago, there have been substantial advancements in seismology and seismic instrumentation. These developments and along with critical limitations of  $M$  scale necessitate a thorough reevaluation of the  $M$  scale to ensure its continued accuracy and relevance.

This paper clearly demonstrates the limitations of the  $M$  and  $M_w$  scales for magnitudes below 7.5, with several reasons supporting this conclusion: 1) Kanamori [3] cautioned against using the  $M_w$  scale for earthquakes below this threshold, prompting the subsequent development of improvements by Hanks and Kanamori [1]. 2) Hanks and Kanamori [1] incorporated an equation  $\log M_0 = 1.5 M_s + 16.1$  from Purcaru and Berckhemer (11) for surface wave magnitude ( $M_s$ ), but this equation was incorrectly used for the range  $5.0 \leq M_s \leq 7.5$ . Importantly, Purcaru and Berckhemer (11) themselves acknowledged that  $M_s$  reliability only extends up to 7.0. This inherent restriction in  $M_s$  translates to potential inaccuracies in the  $M$  scale for earthquakes exceeding this magnitude threshold. The inadequacy of the  $M$  scale for earthquakes below a magnitude of 7.5 is clearly evident from the instrumental datasets, as demonstrated in Figs. 5 and 6.

The key limitations of the  $M_w$  and  $M$  scales—namely, 1) inapplicability to deeper and intermediate-depth events, 2) lack of global validation (applicable for  $M$  scale), and 3) overestimation of energy for lower and intermediate-magnitude earthquakes—can be effectively addressed by using the  $M_{wg}$  scale. Both the  $M$  and  $M_{wg}$  scales are physics-based, as they are derived from  $M_0$ . The assumptions utilized in Das et al. [8] align with the principles established by Kanamori and Anderson [22], affirming their scientific validity.

Additionally, the  $M_{wg}$  scale aligns well with existing magnitude scales ( $m_b$  and  $M_s$ ). The  $M$  scale relies on a constant value to relate moment magnitude and surface wave magnitude. This assumption, however, has been challenged since the work of Purcaru and Berckhemer [11]. In contrast, the  $M_{wg}$  scale eliminates the need for a constant value altogether. Furthermore, the  $M_{wg}$  scale's reliance on

body-wave magnitudes ( $m_b$ ) holds promise for overcoming the limitations of the  $M_w$  scale, particularly for earthquakes at deeper and intermediate depths.

This investigation provides a detailed refutation of the claims made by Gasperini and Lolli [21], specifically their assertion that the  $M$  and  $M_w$  scales are adequate. Additionally, their more recent allegation that the basic assumption mainly using linear regression between  $\log M_0$  and  $m_b$  for development of  $M_{wg}$  scale are flawed has also been found to be incorrect. Linear regression is a standard method in the development of magnitude scales, including the  $M$ ,  $M_w$ , and  $M_e$  scales, and there is no valid reason to reject its application in the case of the  $M_{wg}$  scale. Moreover, the use of linear regression between  $\log M_0$  and  $m_b$  aligns with the theoretical framework established by Kanamori and Anderson [22]. The linear relationship between  $\log M_0$  and  $M_L$  is a fundamental component of the  $M$  scale. If this relationship is correct, then a linear relationship between  $\log M_0$  and  $m_b$  should also hold, as both  $M_L$  and  $m_b$  are based on 1-s seismic waves. Based on a thorough scientific analysis and critical evaluation, it is evident that the conclusions drawn by Gasperini and Lolli [21] are unfounded. This is primarily due to their use of an incorrect constant term,  $-4.7$ , as well as their misinterpretation of the works of Purcaru and Berckhemer [11] and Kanamori and Anderson [22].

We recommend that various international agencies, such as the ISC, NEIC, and GCMT, consider adopting the  $M_{wg}$  scale alongside existing magnitude scales to enhance the understanding of earthquake sources. It is important to note that the  $M_{wg}$  scale is based on body waves, fulfilling a longstanding demand originally emphasized by Gutenberg and Richter [4].

**Acknowledgements** We would like to extend our sincere gratitude to the Editorial Board and the reviewers for their critical and constructive reviews of this manuscript. Their insightful comments and suggestions have been invaluable in improving the clarity, depth, and overall quality of this work. Their dedication to upholding high scientific standards is deeply appreciated. Thanks to Fondecyt Iniciacion Project Grant 11200618.

**Data availability** Body- and surface-wave magnitudes of earthquakes for the entire globe from the International Seismological Centre (ISC, United Kingdom) database (<http://www.isc.ac.uk/iscbulletin/search/bulletin>, last accessed August 2018), and the moment magnitudes from the Global Centroid Moment Tensor (CMT) database (<http://www.globalcmt.org/CMTsearch.html>, last accessed October 2018) during the period 1976–2014 have been compiled in this study.

#### Declaration

**Conflict of interest** The author acknowledges that there are no conflicts of interest recorded.

## Appendix A

While criticizing Das et al. [8], Gasperini and Lolli [21] introduced an alleged misprint in the caption of Table 1 of Purcaru and Berckhemer [11] i.e.,  $M_s = 7-7.5$  without any reference or evidence. Note that this label ' $M_s = 7-7.5$ ' in the caption of Table 1 of Purcaru and Berckhemer [11] has been used in the development of ME scale by Purcaru and Berckhemer [11]. This label ' $M_s = 7-7.5$ ' has no connection in the development of  $M$  scale of Hanks and Kanamori [1]. When the original authors, Purcaru and Berckhemer [11], explicitly acknowledge in their article that they derived Eq. (1) for the magnitude range  $\lesssim 7.0$  due to the reliability of  $M_s$  up to 7.0 (*see page 189 of Purcaru and Berckhemer [11], "In the range of moderate to large earthquakes ( $M_s \lesssim 7$ ,  $\log M_0 < 10^{27}$  dyn.cm), where  $M_s$  is a reliable measure, the relation between  $\log M$  and  $M_s$  is linear and the average relation (see Fig. 4) .....is established to give the best fit with the observed data"*) and the Eq. (1) of Purcaru and Berckhemer [11] has been used in the development of  $M$  scale [1] for magnitude range  $5 \leq M_s \leq 7.5$ . Therefore, introducing alleged misprint in the caption of Table 1 of Purcaru and Berckhemer [11] i.e. ' $M_s = 7-7.5$ ' is irrelevant.

In the formulation of  $M$  scale, Hanks and Kanamori [1] leaned on the support provided by Eq. (1) from Purcaru and Berckhemer [11] to establish the global applicability of the  $M$  scale for magnitudes below 7.5 (Page 1223 of [1], "Which is remarkably coincident with  $M_0$ - $M_s$  relationship empirically defined by Purcaru and Berckhemer [11] for  $5 < M_s < 7.5$ "). However, a closer examination reveals that Purcaru and Berckhemer [11] originally defined the relationship between  $\log M_0$  and  $M_s$  specifically for the magnitude range  $\lesssim 7.0$ . Consequently, this reevaluation suggests that the  $M$  scale is inadequate for accurately representing seismic events below 7.5 on a global scale.

Gasperini and Lolli [21] argued that if  $M_s$  range is considered as  $\leq 7-7.5$  in the caption of Table 1 in Purcaru and Berckhemer [11], then the global applicability of the  $M$  scale could extend to earthquakes with magnitudes below 7.5. However, this argument is incorrect, as the validity of the  $M$  scale for magnitudes below 7.5 is solely linked to Eq. (1) in Purcaru and Berckhemer [11], which only applies up to a magnitude of 7 and has no connection to the caption of Table 1 of Purcaru and Berckhemer [11]. It is important to note that the magnitude range referenced ( $M_s = 7-7.5$ ) in the caption of Table 1 pertains exclusively to the development of the ME scale and has no relevance to the  $M$  scale development.

Das et al. [8] argued that the Eq. (1) of Purcaru and Berckhemer [11] for the magnitude range  $5.0 \leq M_s \leq 7.5$  is not reliable due to the inconsistency of the defined magnitude range (moderate-to-large earthquakes defined as  $M_s \leq 7.0$  and  $M_s = 7-7.5$ ). There is no importance of this



inconsistency issue because Purcaru and Berckhmer [11] clearly explained that they derived their Eq. (1) for  $M_s \lesssim 7.0$  (see page 189, [11]) and this equation is the key equation for M scale applicability for the range  $< 7.5$ .

Das et al. [8] identified two distinct ranges for moderate and large earthquakes in Purcaru and Berckhmer [11]: one for  $M_s \leq 7$  in the development of Eq. 1 of Purcaru and Berckhmer [11] and another for  $M_s = 7-7.5$  in the development of ME scale, highlighting an inconsistency in the definition of moderate and large earthquake ranges. Gasperini and Lolli [21] attempted to resolve this inconsistency by suggesting a misprint in the range  $M_s = 7-7.5$ . However, even if one accepts the possibility of a misprint, the inconsistency remains because two separate magnitude ranges for moderate and large earthquakes persist. Furthermore, the reader should note that even with this misprint one cannot claim M scale is adequate below 7.5 because a) adequacy of M scale below 7.5 is connected only with Eq. (1) of Purcaru and Berckhmer [11] which is applicable to  $M_s \lesssim 7$ .

Gasperini and Lolli [21] specifically focuses on labeling “ $M_s = 7-7.5$ ” as a misprint, rather than stating that “ $M_s \leq 7$ ” may also be a misprint. This choice aligns with their defense of the applicability of the M scale in the range of  $5.0 \leq M_s \leq 7.5$ . However, the Gasperini and Lolli [21] made a significant mistake by misinterpreting an alleged misprint of “ $M_s = 7-7.5$ ” as: “ $M_s \leq 7-7.5$ ”. This alleged misprint introduces a major error because in the caption of Table 1 header of Purcaru and Berckhmer [11] discusses moderate to large earthquakes and assigns the range as “ $M_s = 7-7.5$ ”. It is important to note that the range suggested by Gasperini and Lolli [21] using “ $M_s \leq 7-7.5$ ,” cannot be considered as moderate to large earthquakes, as it would include magnitudes such as 1, 2, 3, 4, 5, 6, and 7.5. Magnitudes 1, 2, 3, 4, etc., are not classified as moderate earthquakes. Even with the inclusion of  $M_s \leq 7-7.5$  in the caption of Table 1 in Purcaru and Berckhmer [11], the essential condition of Eq. (1) of Purcaru and Berckhmer [11] remains unaltered.

In the caption of Table 1, Purcaru and Berckhmer [11] specified an  $M_s$  range of 7–7.5 to define the ME scale. However, Gasperini and Lolli [21] claimed that this range is a printing error, suggesting it should be  $M_s \leq 7-7.5$ . This assertion by Gasperini and Lolli [21] is incorrect, as altering the  $M_s$  range would fundamentally change the definition of the ME scale as originally established by Purcaru and Berckhmer [11]. For justifying the arguments for  $M_s$  range in the caption of Table 1 of Purcaru and Berckhmer [11] they also have presented statements that lack clarity and supporting evidence. One of the statements is as quoted “The clear intention of Purcaru and Berckhmer [11] to indicate the range  $M_s < 7-7.5$  is demonstrated by the parallel indication of

the same range in terms of scalar moment as  $M_0 \leq 10^{27}$ ”. Gasperini and Lolli [21] claimed that the magnitude range  $M_s \leq 7.0-7.5$  is inferred from a parallel indication, specifically  $M_0 \leq 10^{27}$  (dyn cm) in the caption of Table 1 of Purcaru and Berckhmer [11]. However, this assertion is considered incorrect, as the data in Table 2 of Purcaru and Berckhmer [11] indicates that even with  $M_0 = 10^{27}$ ,  $M_s$  can reach as high as 8.4. The representation of the range is comprehensive, taking into account both  $M_s$  and  $M_0$ . Purcaru and Berckhmer [11] explicitly mentioned this when deriving their Eq. 1, stating the condition  $M_s \lesssim 7.0$  and  $M_0 \leq 10^{27}$  (dyn.cm).

Figure 1 in Purcaru and Berckhmer [11] displays a plot of  $\log M_0$  and  $M_s$  in the range  $M_s = 5-8.6$ , incorporating a large set of data compiled from various literature sources, encompassing both ordinary and anomalous earthquakes (Fig. 1). However, they derived Eq. (1) specifically for  $M_s \lesssim 7.0$ , acknowledging the reliability of  $M_s$  up to 7.0. Purcaru and Berckhmer [11] explicitly define the range for moderate to large earthquakes using  $M_s \lesssim 7$  and  $M_0 \leq 10^{27}$  (dyn.cm). From Fig. 1a of Purcaru and Berckhmer [11], it is clearly visible that  $\log M_0 = 1.5 M_s + 16.1$  is derived using  $M_s \lesssim 7.0$  and  $M_0 \leq 10^{27}$  (Fig. 1). The broken lines of Fig. 1a of Purcaru and Berckhmer [11] clearly demonstrates that  $M_s$  data was considered up to  $\leq 7.0$  (Fig. 1). In Fig. 3 of Purcaru and Berckhmer [11], the relationship between  $M_0$  and  $E_s$  (seismic energy) for moderate to large earthquakes is illustrated using crosses as symbols. Figure 4 of Purcaru and Berckhmer [11] clearly shows the relation between  $\log M_0$  and  $M_s$  for moderate to large earthquakes for  $M_s < 7$  (Page 197 of Purcaru and Berckhmer [11], “The two broken lines indicate that  $M_0$ , corresponding to fixed  $M_s$  or  $M_E$ , varies within a factor of 6. Constant strain drop:  $\log M_0 = 16.1 + 1.5 M$  (full line)”).

Gasperini and Lolli [21] claim that there are sentences on page 189 of Purcaru and Berckhmer [11] that indicate a magnitude range up to 7.5 (“the clear intention of Purcaru and Berckhmer [11] to indicate the range  $M_s \leq 7.0-7.5$  is demonstrated by the parallel indication of the same range in terms of scalar moment as  $M_0 \leq 10^{27}$  dyn.cm and even by other sentences of the articles e.g., at pages 189 and 195”). However, there is only one sentence on page 189 that mentions the range of “7.5” (page 189 of [11], “The scatter within the linear band in Fig. 1, for  $M_s \leq 7.5$ , is therefore interpreted in terms of variable stress drop”). This sentence explains that the variation or spread of the data points within the linear band can be attributed to the concept of variable stress drop.

Based on above discussions, Fig. 1 and several statements (e.g., on page 189 “In the range of moderate to large earthquakes ( $M_s \lesssim 7$ ,  $M_0 \lesssim 10^{27}$  dyn.cm), where  $M_s$

is a reliable measure, the relation between  $\log M_0$  and  $M_s$  is linear ... is established", on page 190 "The two broken straight lines (the band) represent the linear trend of the relation  $\log M_0$ - $M_s$  for moderate to large earthquakes ( $M_s \lesssim 7$  and/ or  $M_0 \lesssim 10^{27}$  dyn.cm")) of Purcaru and Berckhemer [11], the conclusion can be drawn that Gasperini and Lolli [21] have inaccurately represented or mischaracterized the findings of Purcaru and Berckhemer [11].

## References

- Hanks TC, Kanamori H (1979) A moment magnitude scale. *J Geophys Res* 84:2348–2350
- Bormann P, Yadav RBS (2010) Reply to "Comment on a homogeneous and complete earthquake catalog for Northeast India and the adjoining region by RBS Yadav, P. Bormann, BK Rastogi, MC Das, and S. Chopra" by R. Das and HR Wason. *Seismol Res Lett* 81(2):235–240
- Kanamori H (1977) The energy release in great earthquakes. *J Geophys Res* 82:2981–2987
- Gutenberg B, Richter CF (1956) Earthquake magnitude, intensity, energy, and acceleration: (Second paper). *Bull Seismol Soc Am* 46(2):105–145
- Kanamori H (1983) Magnitude scale and quantification of earthquakes. *Tectonophysics* 93:185–199
- Choy GL, Boatwright JL (1995) Global patterns of radiated seismic energy and apparent stress. *J Geophys Res: Solid Earth* 100(B9):18205–18228. <https://doi.org/10.1029/95JB01969>
- Boudebouda A, Athmani A, Ranjit D (2024) A unified earthquake catalog for Northern Algeria based on an advanced moment magnitude scale using a robust regression method. *Pure Appl Geophys* 181(4):1117–1138. <https://doi.org/10.1007/s00024-024-03451-x>
- Das R, Sharma ML, Wason HR, Choudhury D, Gonzales G (2019) A seismic moment magnitude scale. *Bull Seism Soc Am* 109(4):1542–1555. <https://doi.org/10.1785/0120180338>
- Das R, Meneses C, Urrutia D (2023) Regression relationships for conversion of body wave and surface wave magnitudes toward Das magnitude scale  $M_{wg}$ . *Nat Hazards* 117:365–380. <https://doi.org/10.1007/s11069-023-05863-9>
- Das R, Joshi S, Meneses C, Biswas T (2023) Advanced unified earthquake catalog for North East India. *Appl Sci* 13(5):2812
- Purcaru G, Berckhemer H (1978) A magnitude scale for very large Earthquakes. *Tectonophysics* 49:189–198
- Keir D, Stuart GW, Jackson A, Ayele A (2006) Local earthquake magnitude scale and seismicity rate for the Ethiopian rift. *Bull Seismol Soc Am* 96:2221–2230
- Ristau J, Rodgers GC, Cassidy JF (2003) Moment magnitude–local magnitude calibration for earthquakes off Canada's west coast. *Bull Seismol Soc Am* 93:2296–2300
- Akash A, Shiuly A, Sau D (2024) Ground motion prediction model for Himalayan region: a comprehensive review. *Acta Geophysica* 72:2163–2191
- Shiuly A, Roy N, Sahu RB (2020) Prediction of peak ground acceleration for Himalayan region using artificial neural network and genetic algorithm. *Arab J Geosci*. <https://doi.org/10.1007/s12517-020-5211-5>
- Shiuly A (2018) Global attenuation relationship for estimating peak ground acceleration. *J Geol Soc India* 92(1):54–58. <https://doi.org/10.1007/s12594-018-0952-4>
- Richter, C. F. (1958). *Elementary Seismology*, W.H. Freeman, San Francisco, California.
- Thatcher W, Hanks TC (1973) Source parameters of southern California earthquakes. *J Geophys Res* 78:8547–8576
- Ekström G, Dziewonski AM (1988) Evidence of bias in estimations of earthquake size. *Nature* 332(6162):319–323
- Bormann P, Giacomo DD (2010) The moment magnitude  $M_w$  and the energy magnitude  $M_e$ : common roots and differences. *J Seismol* 15:411–427
- Gasperini P, Lolli B (2024) Comment on "a seismic moment magnitude scale" by Ranjit Das, Mukat Lal Sharma, Hans Raj Wason, Deepankar Choudhury, and Gabriel Gonzalez. *Bull Seismol Soc Am*. <https://doi.org/10.1785/0120230230>
- Kanamori H, Anderson DL (1975) Theoretical basis of some empirical relations in Seismology. *Bull Seism Soc Am* 65(5):1073–1095
- Bath, M (1966). *Earthquake energy and magnitude*, in *Physics and Chemistry of the Earth*, vol 7, L H. Ahrens, F Press, S. K Runcorn, and H C. Urey, Editors, Pergamon Press, New York, 117–165.
- Choy GL, McGarr A, Kirby SH, Boatwright J (2006) An overview of the global variability in radiated energy and apparent stress. Washington DC Am Geophys Union Geophys Monograph Ser 170:43–57
- Gasperini P, Lolli B, Vannucci G (2013) Body-wave magnitude  $m_b$  is a good proxy of moment magnitude  $M_w$  for small earthquakes ( $m_b < 4.5$ –5.0). *Seismol Res Lett* 84(6):932–937
- Haskell N (1964) Total energy and energy spectral density of elastic wave radiation from propagating faults. *Bull Seism Soc Am* 56:1811–1842
- Lolli B, Gasperini P, Vannucci G (2014) Empirical conversion between teleseismic magnitudes ( $m_b$  and  $M_s$ ) and moment magnitude ( $M_w$ ) at the Global, Euro-Mediterranean and Italian scale. *Geophys J Int* 199:805–828. <https://doi.org/10.1093/gji/ggu264>
- Bormann P, Liu R, Xu Z, Ren K, Zhang L, Wendt S (2009) First application of the new IASPEI teleseismic magnitude standards to data of the China National Seismographic Network. *Bull Seismol Soc Am* 99(3):1868–1891. <https://doi.org/10.1785/0120080010>
- Scordilis EM (2006) Empirical global relations converting  $M_S$  and  $m_b$  to moment magnitude. *J Seismolog* 10:225–236
- Das R, Wason HR, Sharma ML (2011) Global regression relations for conversion of surface wave and body wave magnitudes to moment magnitude. *Nat Haz* 59(2):801–810
- International Association of Seismology and Physics of the Earth's Interior (IASPEI) (2013). Summary of Magnitude Working Group recommendations on standard procedures for determining earthquake magnitudes from digital data, Available at: [http://www.iaspei.org/commissions/commission-on-seismological-observation-and-interpretation/Summary\\_WG\\_recommendations\\_20130327.pdf](http://www.iaspei.org/commissions/commission-on-seismological-observation-and-interpretation/Summary_WG_recommendations_20130327.pdf) (last accessed September 2022).
- Wason HR, Das R, Sharma ML (2012) Magnitude conversion problem using general orthogonal regression. *Geophys J Int* 190(2):1091–1096
- Das R, Wason HR, Sharma ML (2014) Unbiased estimation of moment magnitude from body- and surface-wave magnitudes. *Bull Seismol Soc Am* 104(4):1802–1811. <https://doi.org/10.1785/0120130324>
- Das R, Wason HR, Gonzalez G, Sharma ML, Chodhury D, Roy N, Salazar P (2018) Earthquake magnitude conversion problem. *Bull Seismol Soc Am* 108(4):1995–2007

35. Das R., Meneses , C., Hua, W.(2024a). Seismic and GNSS strain-based probabilistic seismic hazard evaluation for Northern Chile using das magnitude scale, Geoenvironmental Disaster, accepted
36. Das A, Das T, Chaudhuri CH, Choudhury D (2024) Integrated study on buried pipelines, co-seismic landslides, and magnitude conversion for 2023 Türkiye earthquake. Eng Geol 337:107599

Springer Nature or its licensor (e.g. a society or other partner) holds exclusive rights to this article under a publishing agreement with the author(s) or other rightsholder(s); author self-archiving of the accepted manuscript version of this article is solely governed by the terms of such publishing agreement and applicable law.

**Publisher's Note** Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.