Finite frequency inversion of cross-correlation amplitudes for ambient noise source directivity estimation

Arjun Datta[1], Shravan Hanasoge[1], and Jeroen Goudswaard[2]

[1]Department of Astronomy and Astrophysics, Tata Institute of Fundamental Research, Homi Bhabha Road, Colaba, Mumbai 400005, India.
[2]Shell India Markets Private Limited, Devanahalli Industrial Park, Mahadeva Kodigehalli, Bangalore 562 149, Karnataka, India.

Key Points:

- Seismic noise directivity estimation by waveform inversion of noise correlations
- Applied to an exploration scale data set
- Observational uncertainties incorporated in inversion for source distribution

Corresponding author: A. Datta, `arjun.datta@tifr.res.in`

Abstract

We present a new method for determining the azimuthal variation of ambient noise sources, that combines the computational speed and simplicity of traditional approaches with the rigour of waveform-inversion-based approaches to noise-source estimation. This method is based on a previously developed theoretical framework of sensitivity kernels for cross-correlation amplitudes. It performs a tomographic inversion for ambient noise sources on the Earth's surface and is suitable for small (local) scale studies. We apply the method to passive seismic data acquired in an exploration context, and account for azimuth-dependent uncertainties in observed cross-correlation amplitudes. Our inversion results correlate well with the azimuthal distribution of noise sources suggested by signal-to-noise ratio analysis of noise cross-correlation functions.

1 Introduction

The issue of heterogeneous noise sources or noise directivity has received significant attention in the field of ambient-noise seismology. Several studies have reported inaccuracies or limitations associated with the seismic measurements that can be extracted from noise cross-correlation functions (NCFs), such as surface wave dispersion (Pedersen & Krüger, 2007; Tsai, 2009; Yao & van der Hilst, 2009; Kimman & Trampert, 2010; Froment et al., 2010; Wang et al., 2016) or anelastic attenuation (Tsai, 2011). Considerable effort has gone into developing tools for the proper exploitation of ambient-noise observations that are recognized to arise from anisotropically distributed sources (Roux, 2009; Curtis & Halliday, 2010; Wang et al., 2016). Knowledge of the sources of ambient-field measurements, particularly their spatial and/or temporal variations, is now widely recognized as a prerequisite to any attempts to use these measurements to study Earth structure (e.g. Delaney et al., 2017; Lehujeur et al., 2017).

The methods by which this information is typically obtained, with varying degrees of detail, fall under three broad categories: traditional array-processing or beamforming methods (e.g. Friedrich et al., 1998; Roux et al., 2005; Gerstoft & Tanimoto, 2007; Ruigrok et al., 2011; Gal et al., 2014; Nakata et al., 2015; Liu et al., 2016; Löer et al., 2018), analysis of NCF asymmetry (Stehly et al., 2006; Yang & Ritzwoller, 2008; Ermert et al., 2016) and more recently, inversions for ambient-noise sources (Delaney et al., 2017; Ermert et al., 2017; Lehujeur et al., 2017). Inversions, by virtue of modelling and fitting waveforms (or attributes thereof) are the most rigorous, especially when based on spatially extended finite-frequency kernels.

The method we present in this paper carves a niche in this third category. Based on acoustic modelling and inverting only for noise-source directions, it is free of the need for expensive numerical simulations (Ermert et al., 2017) or ray-theoretical simplifications (Delaney et al., 2017). To the best of our knowledge it is also the first method to demonstrably account for uncertainties in observed NCF amplitudes. The simplifications of the method are that it is in 2-D (all sources and receivers required to be located on a plane) and structural variations remain unmodelled (a uniform wavespeed is assumed). At scales where the Earth's sphericity may be neglected, the former condition is not limiting because terrestrial seismic noise is known to be generated primarily on the Earth's surface (e.g. Webb, 1998). The latter restriction is also not a major impediment as amplitudes are used rather than traveltimes, so observed and synthetic waveforms can be reconciled by a judicious choice of measurement. However the restrictions do imply that our method is suitable for local-scale studies, not regional or continental scales (here the term local-scale refers to length scales that are short enough for the Earth's sphericity to be negligible). As emphasized by Lehujeur et al. (2017), local-scale studies are important candidates for noise directivity analysis, because small-aperture networks likely contain a significant fraction of measurements from interstation paths that are short compared to the seismic wavelength — measurements for which the effects of anisotropic noise sources are most pronounced (Bensen et al., 2007; Tsai, 2009; Froment et al., 2010).

We apply our method to an exploration-scale dataset acquired by Shell, which has captured azimuthally varying ambient seismic noise over a period of several months. The aim of this paper is method validation, not thorough characterization of our passive seismic data set, so we limit this study to a single frequency band. Applying the method to narrow frequency bands mitigates effects arising from neglecting (surface-wave) dispersion in the forward modelling. Our results are compared to those from NCF-asymmetry analyses.

2 Data and preliminary analysis

Our data consist of one week (6-12 Nov 2016) of continuous ambient-field recordings taken by a network of 289 three-component, 5 Hz geophones placed over an area

$\sim$ 24 km $\times$ 28 km (Figure 1). The data, stored in consecutive 10-minute intervals, have been downsampled from 4 ms to 20 ms sampling interval. For each 10-min window, vertical-component NCFs of 50 s duration are computed following standard processing techniques (Bensen et al., 2007), including running absolute-mean normalization and spectral whitening after filtering the raw data in the frequency band $0.1-0.5$ Hz. The results are then averaged over a 24-hour period to obtain NCFs for each day. Clear travelling waves are observed in record sections of the NCFs (Figure A.1).

We first attempt to characterize azimuthal variations in the strength of the daily NCFs. This is a type of NCF-asymmetry analysis that requires normalizing for different interstation distances (in each azimuth bin), since NCF amplitudes are expected to decay with distance (Cupillard & Capdeville, 2010). Ignoring the effects of anelastic attenuation, one way to perform this normalization is to multiply observed amplitudes by the inter-receiver distance (e.g. Stehly et al., 2006). This formalism rests on the implicit assumption that the data exhibit the canonical geometrical spreading rate of surface waves in a laterally homogeneous medium. We test this assumption in our data and find that there are many azimuths for which the observed NCF energies (E) are poorly described by the canonical decay rate. Sorting our daily NCFs into azimuth bins of $4°$, regardless of absolute receiver location and inter-receiver distance r, we fit $1/r$ curves to the observed scatter in NCF energies plotted as a function of r (Figure A.2). Each NCF yields measurements for two (radially opposite) directions, because energies in the positive and negative branches of the waveform are measured separately. Figure 2 shows a polar plot of the χ^2 misfit comprising deviations from the $1/r$ approximation, for each azimuth bin. Given the significant deviations from theory at some azimuths, we infer that a normalization such as that of Stehly et al. (2006) would produce biases and lead to erroneous interpretation of dominant source directions. The method we propose in this paper is free of this potential bias because it does not require presuming an amplitude decay rate, but instead uses the observed amplitudes to quantify data errors in an inverse problem.

3 Methodology and synthetic tests

Hanasoge (2013) introduced sensitivity kernels for NCF amplitudes based on analytical solutions to an acoustic-wave equation in homogeneous 2-D media. In this method NCFs are forward modelled using a power spectrum $P(\omega)$ of spatially uncorrelated sources and their spatial-amplitude distribution $s(\mathbf{x})$ in a 2-D plane, along with a homogeneous wavespeed. For the inverse problem, the method takes as input the NCF energy in a time window of interest (denoted by A), either on the positive or the negative branch, and computes source-amplitude kernels which define the sensitivity of this measurement to $s(\mathbf{x})$. Given a network of receivers, the method lends itself to a fast Hessian-based tomographic inversion for source distribution. A misfit functional is defined which penalizes the logarithmic ratio of A^{obs} of A^{syn} for all receiver pairs, and a misfit kernel K is obtained as the sum of individual source-amplitude kernels weighted by their misfits. Except for details of the measurement and misfit, this is analogous to, for instance, classical (earthquake-based) finite frequency traveltime tomography with banana-doughnut kernels (e.g. Dahlen et al., 2000).

Hanasoge (2013) used synthetic tests to show the suitability of the misfit kernels for tomographic inversion. In this study, we take the straightforward next step of actually performing these inversions. We choose a 100 km $\times$ 100 km modelling domain with a uniform grid spacing of 0.25 km and a uniform wavespeed of 2.1 km/s, ensuring a sampling of 16 grid points per shortest wavelength ($\approx$ 4.4 km). The chosen wavespeed is estimated from the real data by picking the maxima of envelopes on NCF record sections (e.g. Figure S1) and averaging positive and negative branch values. Since we work with amplitudes, there is a trade off between the strength of sources and their distances from the receiver network, which we do not endeavour to resolve. We aim only to constrain the relative azimuthal distribution of noise sources. To this end, we parameterize our spatial source distribution as a ring of sources surrounding the network (see Figure 3a). The ring comprises M spatial Gaussians centered every $10°$, which form a basis set for model expansion (see Supplement), and its radius must be sufficiently larger than the size of the network. In this paper we present results for a 50 km-wide ring but have verified with synthetic tests as well as real data, that the size of the ring has a negligible impact on the inferred source directions as long as the ring is wider than the array aperture. The model vector, of length M, simply contains the basis-function coefficients. We choose $M = 36$, thereby working with a highly tractable 36-parameter model space. Since individual kernels for each receiver pair are computed, the gradient and Hessian of the misfit are both available, allowing us to invert using a standard quasi-Newton scheme (see Appendix B). The model update is obtained separately for positive- and negative-branch measurements and the two results are averaged. Our inverse method is validated by several synthetic tests, performed using the same network configuration as the real dataset.

We begin by using a test model that follows the parameterization used for the inverse problem (Figure 3), so all sources are constrained to lie on the same ring around the receiver network. We run a series of examples using different-size subsets of the complete receiver network and the two end-member cases are shown in Figure 3. The combination of misfit kernels from positive- and negative-branch measurements illuminates all model perturbations and the inversion recovers the true model with an accuracy that improves, as expected, with increased network coverage. With 256 receivers, the test results are essentially perfect — all source locations and amplitudes are recovered exactly. Next, we employ a test model (Figure 4) that is free of the aforementioned parameterization, i.e. it contains arbitrarily oriented sources at arbitrary distances from the receiver network. In this case it is difficult to quantify the accuracy of the model obtained by inversion (true source amplitude information is lost), but we argue that source directions and their relative strengths are recovered reasonably well. Taken together, results from these tests validate the inverse algorithm and the imaging concept advocated by this paper. In particular, the latter tests justify our choice of 50 km for the size of the ring with which we parameterize the model space.

4 Application to real data

We apply the method to our data after tapering and downsampling the observed NCFs to 0.2 s. This factor-of-10 downsampling is in accord with the highest frequency considered (0.5 Hz) and reduces the computational cost of forward modelling. Based on the work of Section 3, the modified NCFs are subjected to inversion using a 60 km $\times$ 60 km modelling domain and 256 receivers. However three important considerations remain: the choice of $P(\omega)$ for the noise sources, the choice of window size for the measurements A, and quantification of observational errors or uncertainties. We discuss each one in detail.

A judicious choice of $P(\omega)$ is required to produce synthetic NCFs that match the input observed NCFs not only in shape (frequency content) but also – in order to initialize the inversion – in order of magnitude of amplitude. Since $P(\omega)$ contributes to the synthetic NCFs through interaction with other terms in a spatial integral (Hanasoge, 2013, eq. 11), it is empirically estimated as a skew-symmetric Gaussian function (Figure A.3a). Its amplitude is determined by requiring that the initial synthetics have energies equal to an average representation of the observed NCF energies, namely the $1/r$ curve that best fits the observed NCF energies as a function of interstation distance (Figure A.3b). This is similar to the analysis of Section 2 but here, the energies are computed for complete NCFs (not separately for positive and negative branches) and all observations are taken together, regardless of azimuth. The second consideration, that of the measurement window, is important because the window needs to be narrow enough to ensure that it fits only the coherent signals in the NCFs and yet broad enough to accommodate those cases where the signals in the observed and synthetic NCFs are significantly offset from each other in time (the chosen wavespeed is likely a poor approximation in some parts of the study region). We set an 8 s window around the central arrival time dictated by the chosen homogeneous wavespeed. The taper applied to the observed NCFs exists only outside and beyond this measurement window.

Lastly, the data analysis is made meaningful by introducing data errors into the inverse problem. We consider two types of error: σ_1, based on the energy-decay analysis of Section 2, and σ_2, due to the signal-to-noise ratio (SNR) of the observed NCFs. The first error, σ_1, is simply an estimate for ΔA obtained using $A = \sqrt{E}$, and ΔE as yielded by curve-fitting in the relevant azimuth bin (e.g. Figure S2). SNR is defined as the ratio of the energy density of an NCF in a time window of interest (we use the inversion measurement window), to its energy density outside that window. Similar to the analysis of Section 2, SNR is measured separately on the positive and negative branches, yielding two values for each NCF and therefore an error estimate (σ_2) for each measurement made for source inversion. SNR values are translated into data errors by qualitatively identifying three SNR regimes (>3, $2 < \text{SNR} < 3$ and < 2) to which we assign $\sigma_2 = 5\%$, 50% and 80% respectively of the A value. We assume that errors on different measurements are uncorrelated ($\mathbf{C}_D$ in (B.12) is diagonal) and that the two errors considered are independent, i.e. $\sigma^2 = \sigma_1^2 + \sigma_2^2$. The latter assumption is based on the observation that high SNR values coincide with large departures from a $1/r$ energy decay rate (see Figures 2,6). Other sources of error, such as neglecting the topography of the region and differing station elevations, are not explicitly accounted for.

The inverse algorithm thus customized, is applied separately to the observed NCFs for each day. We achieve total misfit reductions of $\sim 70 - 80\%$, with at least 50% of measurements being fit to an accuracy greater than 70%. An example of this quantitative summary for one of the days is shown in Figure 5. Finally, we corroborate our results with the simple idea of using SNR itself as a proxy for source directionality. Azimuthal polar plots of SNR are obtained in a manner similar to Figure 2, with the exception that, here, distance does not play an explicit role – SNR values for different interstation distances within an azimuth bin are averaged. Yang & Ritzwoller (2008) used this kind of analysis on long-range correlations to infer noise-source directions at continental scales. In this study, we find that subtle variations – on daily time scales – in the dominant wave-propagation directions (primarily between ENE and NNE) inferred from SNR polar plots, correlate well with variations in the azimuthal noise-source distribution obtained by our inverse method (Figure 6). Also, there is qualitative agreement between the two methods, on the apparent strength of noise sources on different days – SNR values are highest on days 3-5, as are source amplitudes obtained by inversion.

5 Discussion and conclusions

We have introduced a new method of determining ambient noise-source directionality by inverting NCF amplitudes. In its present form, the method is suitable for local-scale studies, where the effects of Earth's sphericity are negligible. The main attraction of the technique compared to existing approaches is the combination of computational ease and rigorous waveform-fitting. Although the method owes its economy to strong simplifying approximations, these do not preclude its application. In particular, the use of amplitudes ensures that Earth structure has a diminished impact on the measurements. We applied the technique to a study region with known strong lateral variations in geological structure, and achieved waveform fits comparable with typical seismic tomography applications.

The real data application systematically accounts for observational uncertainties, and demonstrates the suitability of the method for studying non-stationarity in ambient-noise-source distributions over short timescales (e.g. hours or days for reservoir monitoring). This is possible because the technique may be applied to ambient seismic noise of shorter duration than is typically used in most applications. In this respect, it contributes to a growing body of seismological research (e.g. Roux & Ben-Zion, 2017; Fichtner et al., 2017) that treats NCFs as fundamental seismic observables in their own right, without the requirement of convergence to an interstation Green's function.

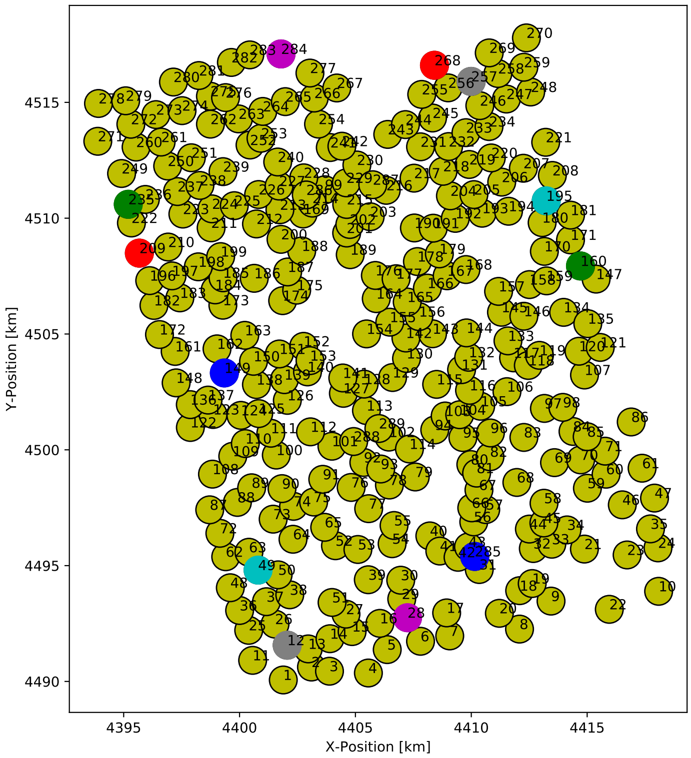

Figure 1. Map view of network of recording geophones, numbered from 1 to 289. Colours other than yellow indicate six receiver pairs selected for later reference in Figure 5.

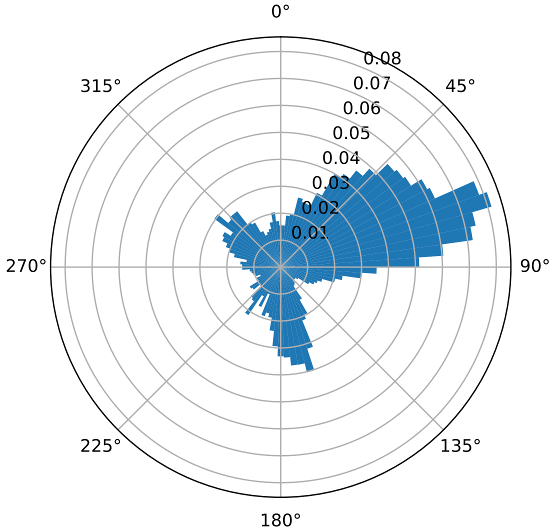

Figure 2. Single-day example (day 3) of azimuthal variation in the departure of the observed cross-correlation energies from a $1/r$ energy decay regime, where r is interstation distance. Plotted values are the χ^2 misfit obtained in each azimuth bin – refer to text and Figure A.2.

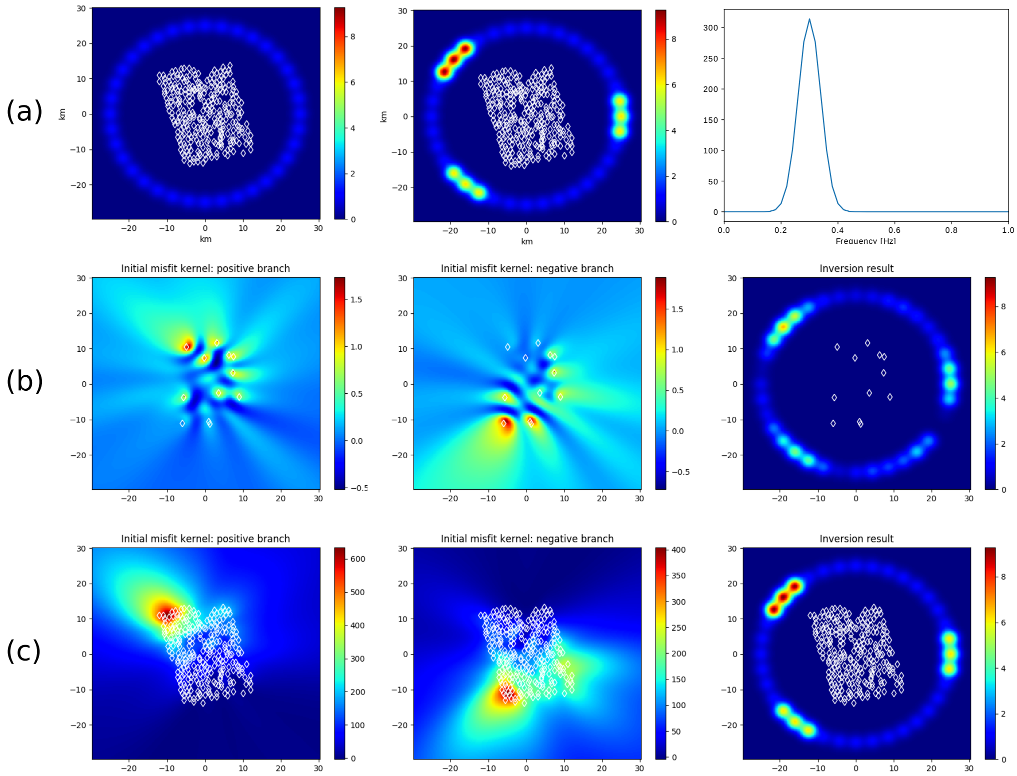

Figure 3. Type 1 synthetic test of the inverse method. (a) Left: Reference model with an azimuthally homogeneous $s(\mathbf{x})$ shown in colour. All basis functions have unit amplitude in this model and it serves as the starting model for inversions. Middle: The "true model" for the synthetic test, comprising three high-amplitude regions relative to the reference model. Right: Assumed $P(\omega)$ for the sources, centred at 0.3 Hz. (b) Inversion performed using only 12 receivers: starting misfit kernels for the positive- (left panel) and negative-branch (middle panel) measurements, and the inversion result after 8 iterations (right panel). (c) Same as (b) but with 256 receivers. Note that all the model plots have the same colour scale and in all plots, the white diamonds mark the receiver locations.

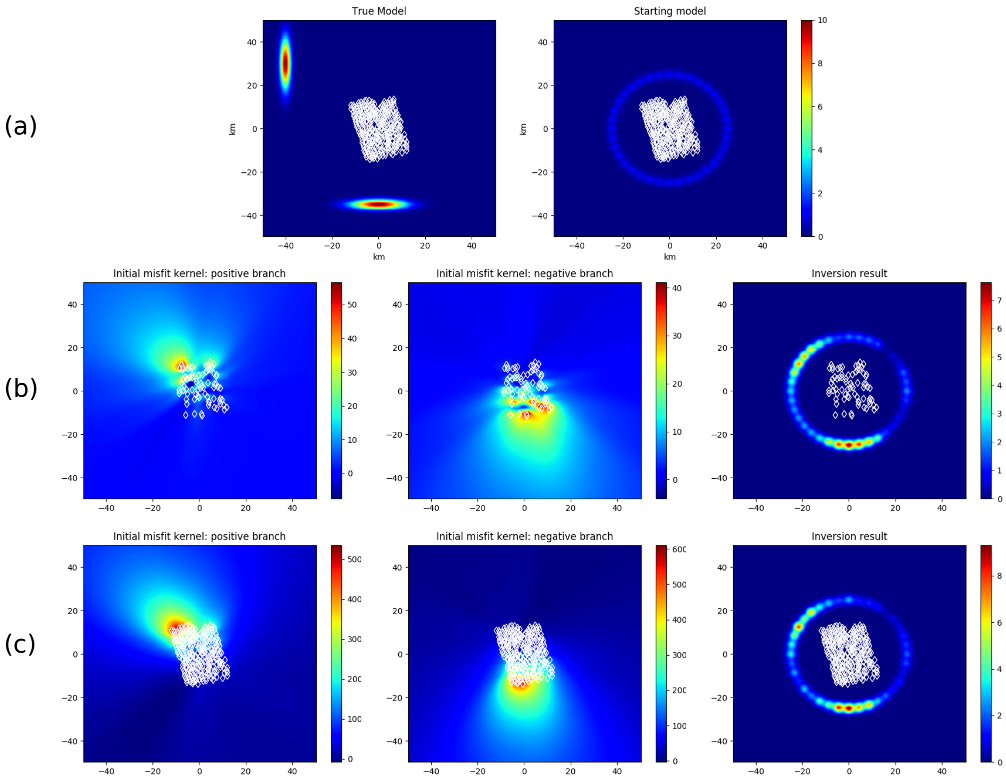

Figure 4. Type 2 synthetic test. (a) True model (left) and starting model for inversions (right, same as in Figure 3). In the true model, the high amplitude regions to the North West and South of the receiver network are at distances of 50 km and 35 km respectively, from the network centre. (b)-(c) same as in Figure 3 except that (b) uses 72 receivers. Note that both inversions predict higher source amplitudes South of the receiver network than North West of it – commensurate with the sources in the true model being closer to the network in the South than in the North West.

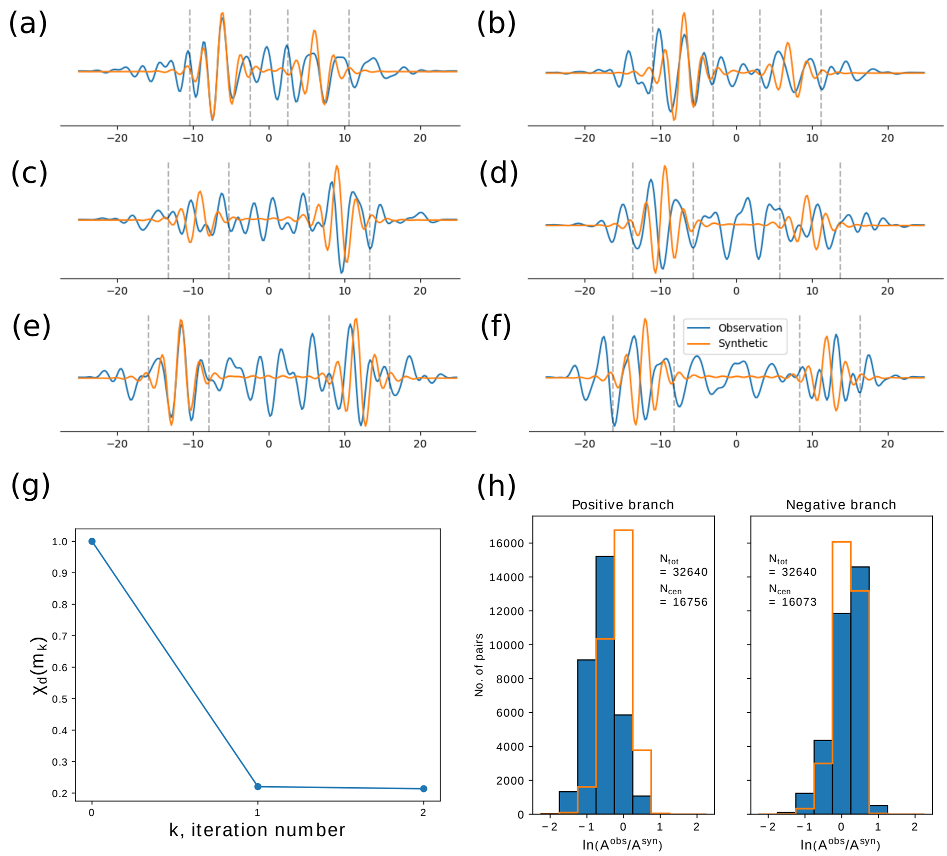

Figure 5. Summary of inversion for day 3. (a)-(f) Six examples of waveform fits after inversion; these correspond to receiver pairs highlighted in blue (285-149), red (268-209), green (235-160), cyan (195-49), magenta (284-28) and grey (257-12) respectively, in Figure 1. All plots share the same legend and the vertical black dashed lines represent the 8-s measurement window in each case. Note that (e) has the greatest symmetry in amplitude between the positive and negative branches, and it corresponds to a receiver pair that is oriented approximately transverse to the dominant incoming energy direction (see Figure 6). (d) Evolution of the total data misfit χ_d (normalized for plotting) through the inversion, showing a misfit reduction of nearly 80% (e) Histograms of Δd values from all interstation pairs in the starting model (blue filled bars) and in the final inverted model (orange step bars). Note that data errors are included in χ_d but not in Δd. Text on the plots indicates the total number of measurements (N_{tot}) and those that lie in the central histogram bin after inversion (N_{cen}). The central bin corresponds to $|\Delta d| <= 0.25$, or a maximum waveform discrepancy (as defined by the measurement) of about 28%. In this case, 32829 measurements (out of the total 65280) lie in the central bin after inversion, meaning that $\sim 50.3\%$ of total measurements are fit to an accuracy of at least 72%.

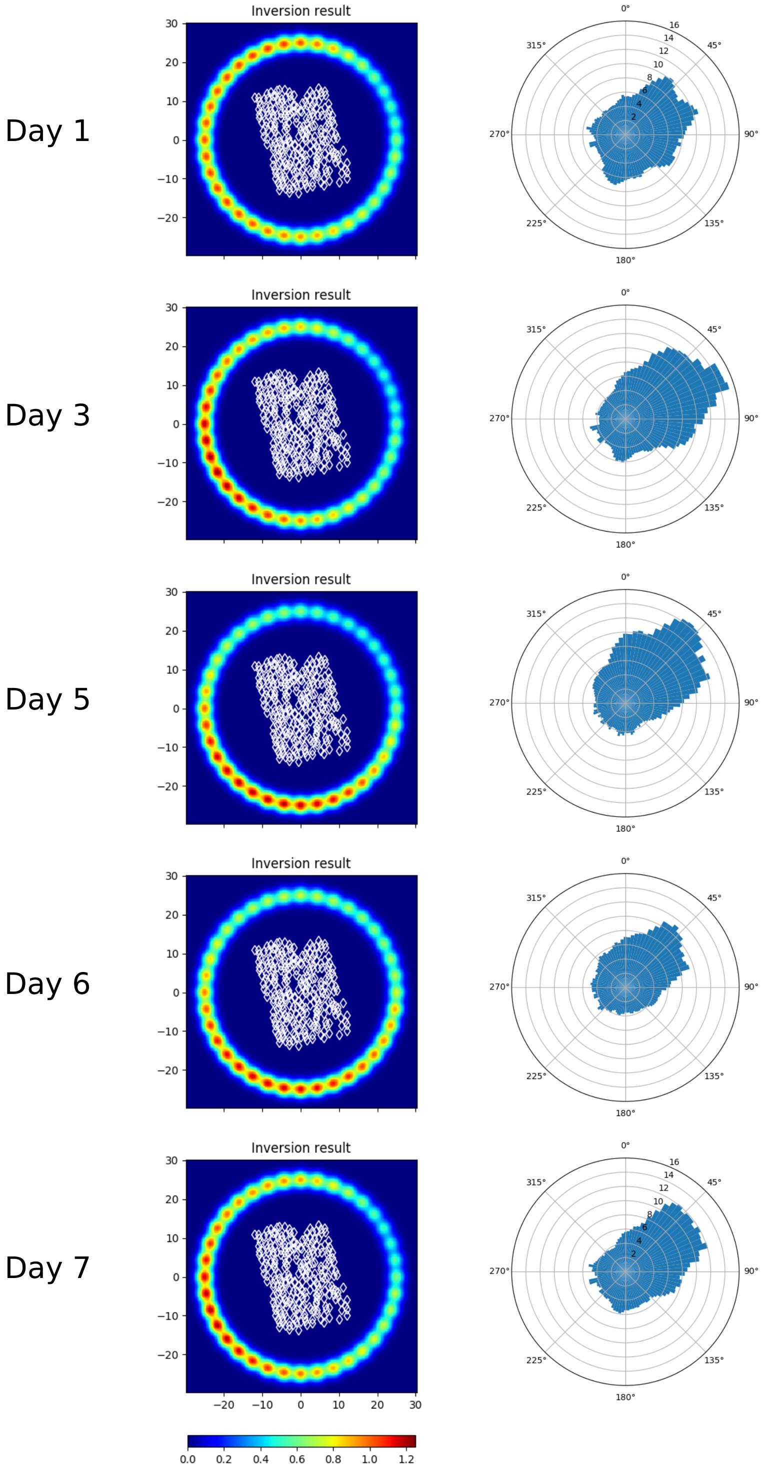

Figure 6. Results corresponding to different days of data from our method (left column) and from SNR analysis (right column). In the SNR plots, each bar points in the direction of wave propagation, i.e. it points away from the source location.

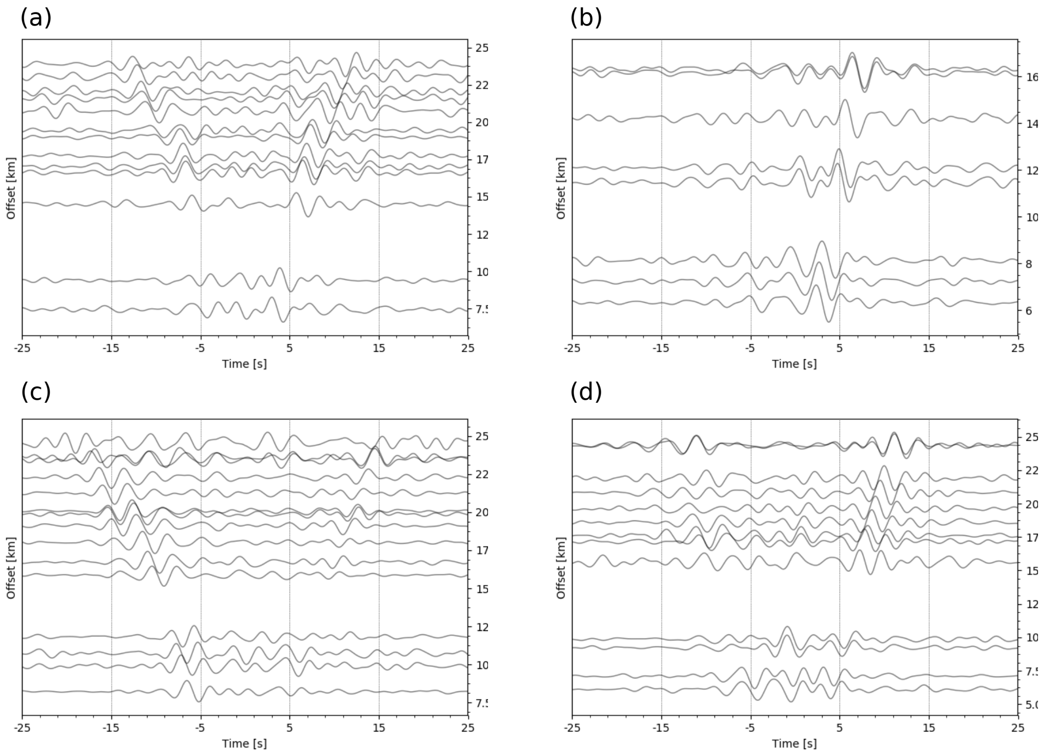

Figure A.1. Examples of record sections of NCFs obtained from one day of data, corresponding to a single base receiver (number 29) and azimuths of (a) 10° (b) 30° (c) 332° and (d) 344° with respect to North. Note that relative amplitudes are not represented in these plots because each plot contains independent normalization for each waveform.

A Supporting figures

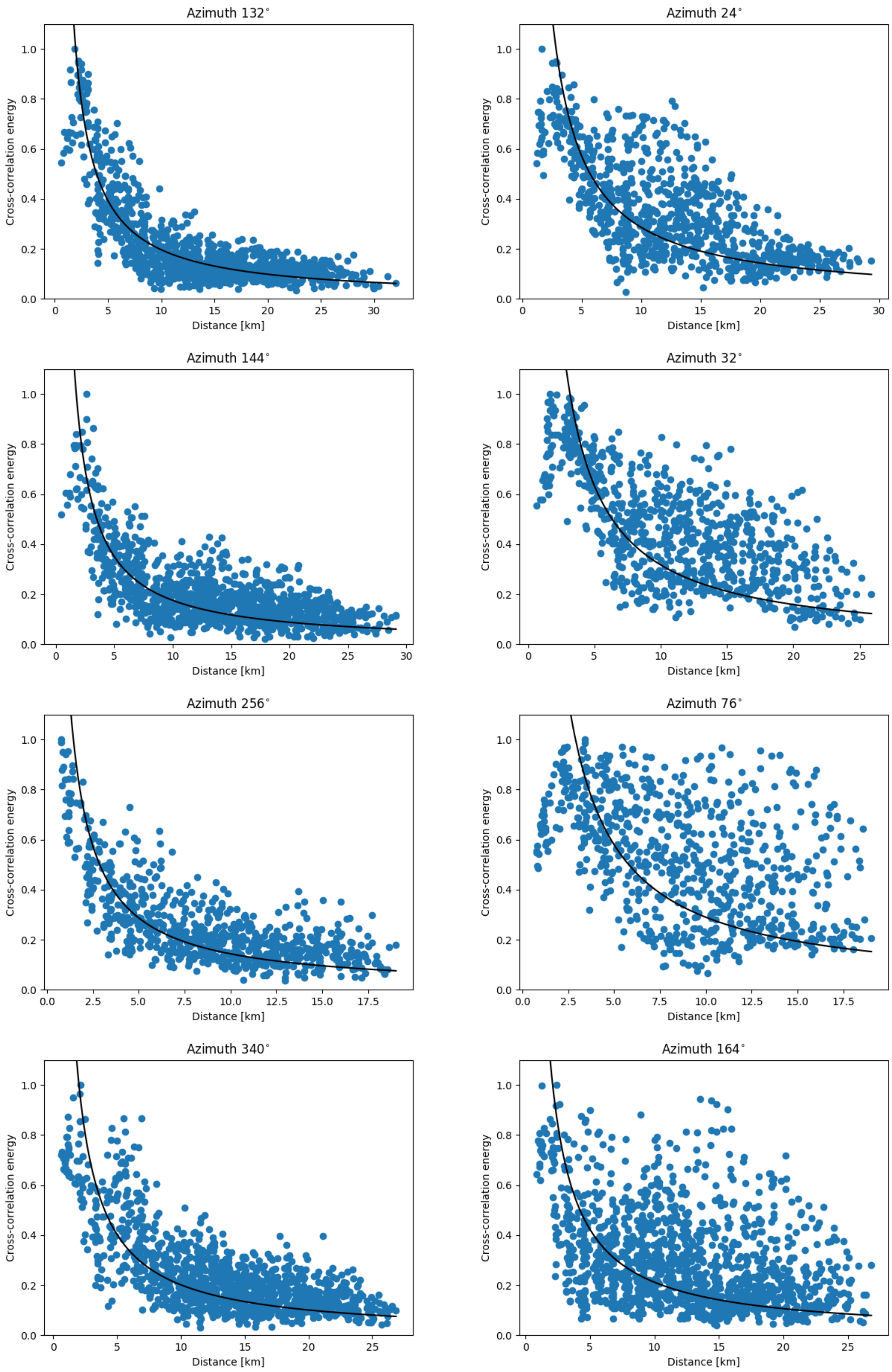

Figure A.2. Normalized observed NCF energies (blue circles) and the best-fit $1/r$ curve, as a function of interstation distance r. Shown are plots for 8 randomly selected azimuth bins, qualitatively separated into “good” (left) and “bad” (right) fits. Corresponds to Figure 2 of main paper.

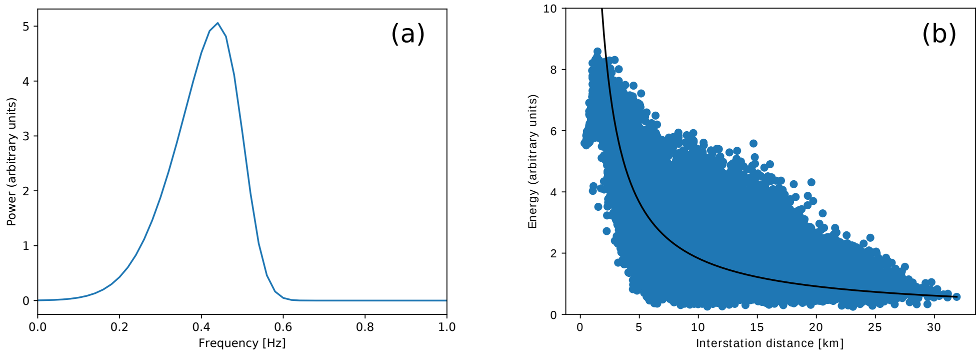

Figure A.3. (a) Shape of $P(\omega)$ used for real data inversions. (b) Azimuth-independent, complete-NCF energy fitting used to initialize said inversions. The starting synthetics for inversion have energies given by the black curve. The curve and symbols have the same meaning as in Figure A.2.

B Model parameterization and inversion

Equation (28) of Hanasoge (2013) may be written as:

$$\delta\chi = -\int K(\mathbf{x})\delta s(\mathbf{x})d^2\mathbf{x} \tag{B.1}$$

where $s(\mathbf{x})$ is the spatial source distribution, $K(\mathbf{x})$ the misfit kernel and χ the misfit functional defined as:

$$\begin{aligned}\chi &= \frac{1}{2}\sum_i \left(ln\frac{A_i^{obs}}{A_i^{syn}}\right)^2 \\ &= \frac{1}{2}[\mathbf{d}_{obs} - \mathbf{d}_{syn}]^T[\mathbf{d}_{obs} - \mathbf{d}_{syn}]\end{aligned} \tag{B.2}$$

with index i denoting the i^{th} receiver-receiver pair. Expanding the model perturbation into the model-space basis functions B_j, we have

$$\delta s(\mathbf{x}) = \sum_{j=1}^{M} \delta m_j B_j(\mathbf{x}) \tag{B.3}$$

Using (B.3) in (B.1) gives

$$\delta\chi = \sum_{j=1}^{M}\left[-\int K(\mathbf{x})B_j(\mathbf{x})d^2\mathbf{x}\right]\delta m_j \tag{B.4}$$

from which it is clear that the quantity in square brackets represents the gradient ($\mathbf{g}$) of χ:

$$\frac{\partial\chi}{\partial m_j} = -\int K(\mathbf{x})B_j(\mathbf{x})d^2\mathbf{x} = g_j \tag{B.5}$$

To express the gradient in terms of the customary $\mathbf{G}$ matrix that linearizes the inverse problem, we need only invoke the fact that the misfit kernel is a weighted sum of source kernels for individual receiver pairs (Hanasoge, 2013, eq. 30):

$$K = \sum_i ln\left(\frac{A_i^{obs}}{A_i^{\text{syn}}}\right)K_i(\mathbf{x}) \tag{B.6}$$

Using (B.6) in (B.5), we have

$$\begin{aligned}g_j &= -\sum_i ln\left(\frac{A_i^{\text{obs}}}{A_i^{\text{syn}}}\right)\int K_i(\mathbf{x})B_j(\mathbf{x})d^2\mathbf{x} \\ &= -G_{ij}\Delta d_i\end{aligned} \tag{B.7}$$

or

$$\mathbf{g} = -\mathbf{G}^T\Delta\mathbf{d} \tag{B.8}$$

where the $\mathbf{G}$ matrix and $\Delta\mathbf{d}$ vector are given by:

$$\begin{aligned} G_{ij} &= \int K_i(\mathbf{x}) B_j(\mathbf{x}) d^2\mathbf{x} \\ \Delta d_i &= ln\left(\frac{A_i^{\mathrm{obs}}}{A_i^{\mathrm{syn}}}\right) \end{aligned} \tag{B.9}$$

Using $\mathbf{G}$, the approximate Hessian $\mathbf{H}$ for a Gauss-Newton inversion is easily computed:

$$\mathbf{H} = \mathbf{G}^T\mathbf{G} \tag{B.10}$$

In practice, the canonical expressions (B.8) and (B.10) are modified, because the objective function itself is modified by damping (required to overcome solution underdeterminacy) and weighting by model or data (co-)variances (required to control the extent of damping or to incorporate observation errors, as in Section 4). The inverse problem we actually work with is (e.g. Tarantola, 2005):

$$\begin{aligned} \tilde{\chi} &= \frac{1}{2}[\mathbf{d}_{obs} - \mathbf{d}_{syn}]^T\mathbf{C}_D^{-1}[\mathbf{d}_{obs} - \mathbf{d}_{syn}] + \frac{1}{2}[\mathbf{m} - \mathbf{m}_{prior}]^T\mathbf{C}_M^{-1}[\mathbf{m} - \mathbf{m}_{prior}] \\ &= \chi_d + \chi_m \\ \tilde{\mathbf{g}} &= -\mathbf{G}^T\mathbf{C}_D^{-1}\Delta\mathbf{d} - \mathbf{C}_M^{-1}(\mathbf{m} - \mathbf{m}_{prior}) \\ \tilde{\mathbf{H}} &= \mathbf{G}^T\mathbf{C}_D^{-1}\mathbf{G} + \mathbf{C}_M^{-1} \end{aligned} \tag{B.11}$$

Here $\mathbf{C}_D$ and $\mathbf{C}_M$ are the model and data covariance matrices. We use $\mathbf{C}_D = \mathbf{I}$ and $\mathbf{C}_D(i,j) = \sigma_{ij}^2\delta_{ij}$ in sections 3 and 4 respectively. Correspondingly, for $\mathbf{C}_M$ we use $\mathbf{C}_M = \frac{1}{\gamma^2}\mathbf{I}$ and $\mathbf{C}_M(i,j) = \frac{1}{\gamma^2}exp\left(\frac{-(i-j)^2}{2L^2}\right)$ respectively. Here γ is a damping parameter and L is a correlation length that controls the smoothness of the model; they are both determined empirically. These choices for $\mathbf{C}_D$ and $\mathbf{C}_M$ imply that the synthetic test inversions are carried out with no data errors and no smoothing, whereas the real data inversions are done with both errors and smoothing included.

Choosing $\mathbf{m}_{prior} = \mathbf{m}_0$ and applying the Gauss-Newton method $\tilde{\mathbf{H}}\Delta\mathbf{m} = -\tilde{\mathbf{g}}$ to (B.11) leads to the following iterative solution:

$$\mathbf{m}_{k+1} = \mathbf{m}_k + \left[\mathbf{G}_k^T\mathbf{C}_D^{-1}\mathbf{G}_k + \mathbf{C}_M^{-1}\right]^{-1}\left[\mathbf{G}_k^T\mathbf{C}_D^{-1}\Delta\mathbf{d}_k + \mathbf{C}_M^{-1}(\mathbf{m}_k - \mathbf{m}_0)\right] \tag{B.12}$$

where index k denotes the iteration number.

Acknowledgments

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