Geochemical characteristics of orthopyroxene-bearing metamorphosed lithologies in the northern Chilka lake area, Eastern Ghats belt, India: Insights into the original nature and petrogenesis of precursors

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Abstract

The northern Chilka Lake area stone quarries expose a high-grade terrane comprising a metapelitic suite interbanded with orthopyroxene-bearing gneisses and garnet-biotite leucogneisses (leptynites). This study presents bulk-rock and mineral chemistry data of the orthopyroxene-bearing units to evaluate the true nature of their protoliths and to comment on their petrogenesis. Field, petrographic, and geochemical observations reveal three distinct variants: Enderbitic paragneisses (EP), Garnetiferous Charnockites (G-Ch), and Non-Garnetiferous Charnockites (NG-Ch). Orthopyroxenes in each unit are texturally and compositionally distinct, and highly likely of metamorphic origin. A detailed study of the EPs indicate that their field and chemical characteristics are more consistent with a depositional origin rather than an intrusive one. Based on bulk rock geochemical aspects of non-migmatitic samples, a provenance involving both felsic and mafic-ultramafic rocks is speculated for the unit. Under this interpretation, the ‘ultramafic two-pyroxene granulite xenoliths’ previously reported to be within this unit possibly represent lithic fragments from an ultramafic source, thereby likely reflecting the very low transport distance of some of the additions to the unit. As the G-Ch rocks exhibit a hybrid nature, it is speculated that their precursors likely represent a deep crustal hybrid melt with contributions from both orthogneisses (likely granodioritic) and a suitable metasedimentary source. In contrast, the NG-Ch rocks, with their characteristic HREE+Y depletion, negative Nb, Ti anomalies and high Sr are most likely products of partial melting of predominantly granodioritic component of a lower crustal TTG gneiss. This study therefore highlights the complexity of the metasedimentary units in the Chilka lake area and reaffirms earlier interpretations regarding the presence of a metagreywacke unit, which may be as predominant as the already recognized metapelite unit. Along with that, this study also reports new crustal anatexis events possibly linked to Grenvillian orogeny.

Keywords: orogenic belt; orthopyroxene; protolith; granulite; partial melting; crustal reworking; Chilka lake

1. Introduction

Orogenic belts preserve evidence of crustal processes operating at convergent settings. The Eastern Ghats belt is a Precambrian orogenic belt that was has witnessed intense deformation, high-grade metamorphism, and crustal anatexis. (Dobmeier and Raith, 2003 and references therein). As orogens are prime locations of crustal reworking, it is highly challenging at times to understand the true nature of rock units but attempts in this direction via interpretation of robust geochemical and isotopic data have been shown to be effective. (Rickers et al., 2001) (Bose et al., 2011b) (Bhattacharya et al., 2012) (Sarkar et al., 2015)

Interestingly, as the Chilka lake area in the northern part of the Eastern Ghats belt is widely known for its preservation of deformational imprints, more emphasis has been given to the structural and metamorphic aspects of the terrane (Bhattacharya et al., 1993,1994) (Sen et al., 1995) (Dobmeier and Raith,2000) (Dobmeier and Simmat, 2001) (Bose et al., 2012) (Bose et al., 2016) such that the precise details of the petrogenetic aspects of the units in the region, especially the variants of orthopyroxene-bearing units, have been largely understudied. Apart from the petrogenetic studies conducted on the leucogneisses (Sen and Bhattacharya, 1997), anorthosites (Chakrabarti et al., 2011), mafic granulites (Bose et al., 2011b), metapelitic suite (Bhattacharya et al., 2012) and a migmatitic variant of orthopyroxene-bearing felsic gneisses(OFGs) (Bose et al., 2020) in the Chilka lake area, a detailed study of the petrogenetic aspects of the precursors to all the variants of orthopyroxene-bearing gneisses that exists as bands/lenses within the larger leucogranitoids and metapelitic units is yet to be undertaken.

This study presents a geochemical investigation of orthopyroxene-bearing units in the northern Chilka Lake area to evaluate the petrogenesis of their precursors. The findings are used to better understand the history of the terrane. Since the units have undergone significant metamorphism, behavior of the LILEs has been examined to evaluate the effects of metamorphism and the petrogenetic aspects of the rock suites have been commented upon solely in light of coherent immobile element chemistry.

2. Geological Setting

The Chilka lake area in the northern part of the Eastern Ghats belt hosts a group of intensely metamorphosed rock suites that include a metapelitic suite (comprising mainly khondalites and minor calc-granulites and quartzites), orthopyroxene-bearing units (both massive and localized) of varying compositions, meta-leucogranitoids (leptynites), minor mafic granulites and an anorthosite massif. (Bhattacharya et al., 1993, 1994) (Sen and Bhattacharya, 1997) (Dobmeier and Raith, 2000). According to the domain classification based on mantle model ages, the Chilka Lake area falls under Domain 3, which is interpreted to encompass rocks derived from the reworking of largely homogeneous Proterozoic material, as per studies on khondalites, charnockites, and leptynites (Rickers et al., 2001). The geodynamic history of the area has been interpreted to be predominantly collisional (Rickers et al., 2001; Dasgupta et al., 2013), with the earliest of the major deformation events (D1) understood to have occurred between ~1250 and 1100 Ma (Simmat and Raith, 2008). The peak of subsequent deformation event (D2), which imparted the pervasive S2 foliation to units in the region has been dated ~990Ma (Bose et al., 2016).

The original nature of the metapelitic suite that is identified to preserve the S1 foliation associated with the earliest deformation event was interpreted as a shale-greywacke sedimentary sequence, similar to marine turbidites (Sen and Bhattacharya, 1997; Bhattacharya et al., 2012), and the leptynites were interpreted to have formed by mica dehydration melting of the metapelitic suite during a granulite facies metamorphic event, which is argued to have occurred during D2 (Sen and Bhattacharya, 1997). Both the mafic granulite and anorthosite precursors have been independently interpreted as mantle-derived units based on their respective geochemical characteristics, with their emplacement possibly linked to Rodinia break-up (Bose et al., 2011b; Chakrabarti et al., 2011).

When it comes to the orthopyroxene-bearing units, both of the first two major studies (Bhattacharya et al., 1993, 1994; Dobmeier and Raith, 2000) suggested that they likely represent igneous granitoids that later underwent granulite facies metamorphism. However, the classification of the orthopyroxene-bearing units as ‘charnockites’ by (Bhattacharya et al., 1993, 1994) was later challenged by (Dobmeier and Raith 2000), wherein they argued that the modal volume of plagioclase in these units makes ‘enderbite’ a more appropriate term. More recently, (Bose et al. 2020) suggested that a variant of orthopyroxene-bearing gneisses in the region is plausibly a paragneiss, formed after metamorphism and partial melting of a greywacke-type protolith. In the light of such findings, a comprehensive study of the geochemical characteristics of all the orthopyroxene bearing units in this region is necessary to obtain a complete picture of the compositional variations.

3. Field and petrographic features of orthopyroxene-bearing units

The present study region encompasses active stone quarries exposed in the northern part of Chilka lake area, in a region locally known as Tapang. Thorough field studies were conducted in four stone quarries in the region namely, USIS Chowk Quarry, Greenlake Quarry, Hatia Quarry and Calcutta Quarry (Fig 1b). This study follows the location names originally introduced by (Dobmeier and Raith 2000), with the addition of a new quarry, Greenlake. The field aspects and petrographic studies of representative samples confirmed the presence of the following orthopyroxene bearing entities in the region:

3.1 *Enderbitic paragneisses (EP)*

The most predominant member of the orthopyroxene bearing group in the study region that is characterized by its melanocratic and fine-grained nature. It occurs as centimeter to meter scale bands within khondalites and leptynites (Fig 2a and c). Upon close inspection, its bandings are visible. Petrographic studies did indicate a higher modal proportion of plagioclase feldspar as previously reported by (Dobmeier and Raith, 2000), (Fig 4a). The petrographic study also revealed that a few highly plagioclase rich varieties having no alkali-feldspar or quartz are present in the region. Representative samples of each variant of the unit has been plotted on the QAP plot (Frost and Frost, 2008) (Fig 4a). For reference, the average modal proportions of the unit as reported by (Dobmeier and Raith, 2000) has also been included. The mineral assemblage of this suite in the decreasing order of modal volume is plagioclase+orthopyroxene±quartz±alkali-feldspar+ilmenite±biotite. Biotites are seen to be in contact with orthopyroxene in most cases (Fig3b), but a few euhedral biotites are also prevalent. Structural studies previously undertaken have suggested that S1 schistosity associated with the earliest deformation is preserved in this unit and is subparallel to the compositional layering between the unit and the metapelites (Dobmeier and Raith, 2000).

3.2 *Garnetiferous Charnockites (G-Ch)*

Rock unit characterized by its coarse grained greasy appearance and garnetiferous nature (Fig 2e and f). It is significantly less voluminous than the EP and is seen to range in dimensions from small centimetre scale lenses and pods to metre scale bands and is predominant in the Hatia and Calcutta quarries although one small outcrop was observed in the Greenlake Quarry. Myrmekites are ubiquitous and is seen to occur at the contact between alkali feldspar and plagioclase feldspar (Fig 3h). The mineral assemblage of this unit in decreasing order of modal volume is alkali feldspar+quartz+plagioclase+garnet+orthopyroxene+biotite+ilmenite (Fig 4a). Accessory minerals include zircon, apatite and monazite. Orthopyroxene is mostly seen to be in contact with garnet and biotite (Fig 3f) and is present in lesser modal proportion when compared to the EPs. Foliation associated with the earliest deformational event has been tentatively suggested for this unit (Bhattacharya et al., 1994) but the most pervasive foliation seen in the field is S2 (ENE-WSW) which has clearly overprinted any previous imprints. Microphotographs of sections perpendicular to the foliation plane reveal lenticular quartz, alkali-feldspar, plagioclase and the mafic minerals-orthopyroxene, biotite, garnet and ilmenite to be the foliation defining minerals (Fig 3g). K-feldspars exist as orthoclase.

3.3 *Non-Garnetiferous Charnockites (NG-Ch)*

Greasy green rock units that are coarse grained in nature and devoid of garnets (Fig 2b). These units are predominant in the northern quarries- USIS and Greenlake and are not prevalent in the quarries in the south. The unit also exhibits the pervasive foliation S2 (Fig 2b). This variant was also identified previously by (Bhattacharya et al., 1994) and was put under the broad umbrella of 'charnockitic rocks'. The mineral assemblage of this suite in the decreasing order of modal volume is alkali feldspar+quartz+plagioclase+orthopyroxene+ilmenite+biotite (Fig 4a). Petrographic studies show that most orthopyroxenes are in contact with biotite laths (Fig 3e). Sections prepared perpendicular to the foliation plane exhibits quartz, alkali-feldspar, plagioclase and mafic minerals-orthopyroxene, biotite and ilmenites as the foliation defining minerals while the ones prepared parallel to the foliation plane shows granoblastic texture with characteristic triple junctions (Fig 3d). Myrmekites are ubiquitous in this suite as well and K-feldspars exist as both orthoclase and microcline.

4. Analytical methods

For this study, a total of nine carefully selected samples (five EPs, two G-Ch and two NG-Ch) were processed for whole rock major and trace element analyses. The samples were cautiously collected from the central portions of the exposed lithological bands to avoid contamination from adjacent units. The major oxide data was generated using Bruker S8 Tiger sequential Wavelength Dispersive X ray Fluorescence Spectrometer (WD-XRF) at NCESS, Thiruvananthapuram.

Pressed pellets were prepared using 40 mm aluminium cups filled with Boric acid crystals as binder. Finely powdered sample (-300 mesh) was sprinkled over boric acid (about 2g sample+ 8g boric acid) in the aluminium cup and was pressed in a hydraulic press with 30 ton pressure to produce a circular 40 mm disk. To estimate the accuracy of the readings, USGS rock standards BHVO-2 and G-2 pellets were run as unknowns. Total iron was measured as Fe_2O_3. The accuracy for the major elements were better than 5 %, and P_2O_5 showed accuracy range between 5-7 %. The trace element data including REEs were measured using the quadrupole Inductively Coupled Plasma Mass Spectrometer (ICPMS, Thermo Scientific X-Series II) at Centre for Earth Sciences, Indian Institute of Science, Bengaluru. The digestion and dilution procedure in (Banerjee et al., 2016) was followed for sample solution preparation. USGS international rock standards BHVO-2, AGV-2, BCR-2 and BIR-1a were run along with the samples to keep a check on accuracy. The accuracy for most elements was better than 5% and for the REEs, it was better than 3%. Internal precision based on three repeat measurements was between 2 to 3 % for all elements.

For mineral chemistry analyses, four representative thin sections of the orthopyroxene-bearing rock units (two EPs, one G-Ch and one NG-Ch) were processed for the chemical data of orthopyroxene, biotite, garnet, plagioclase and ilmenite. The concentrations were determined using JEOL JXA-8230 Electron Microprobe (EPMA) at the Advanced Facility for Microscopy and Microanalysis (AFMM) of the Indian Institute of Science, Bengaluru. The acceleration voltage was 15 kV, the probe current was 12 nA, and the probe diameter was set to 3 microns. The calibration was made using natural and synthetic silicates as well as oxides. The data corrections were made by following the standard ZAF correction procedures (Heinrich, 1991).

5. Results

5.1 Whole rock Geochemistry

5.1a Major and minor elements

The major and minor element chemistry of the orthopyroxene-bearing suites are given in Table 1 and has been illustrated through the plots given in Fig 4. The TAS diagram (Le Bas et al., 1986) shows the mafic to slightly intermediate nature of the EPs (with one isolated sample plotting on the ultramafic region) and felsic nature of the G-Ch and NG-Ch rocks (Fig 4b). The EPs shows metaluminous character in the A/CNK vs A/NK plot (Shand,1943) with A/NK values ranging from 3.60 to 7.75 while the G-Ch and NG-Ch rocks show slightly metaluminous character with A/NK values ranging from 1.59 to 1.61 and 1.40 to 1.42 respectively (Fig 4c). The AFM diagram (Irvine and Baragar, 1971) reveals the tholeiitic affinity of the EPs and the calc-alkaline affinities of the G-Ch and NG-Ch rocks (Fig 4d). The $FeO_t/(FeO_t+MgO)$ wt% vs SiO_2 wt% diagram (Frost and Frost, 2008) shows that the EPs exhibit magnesian to ferroan character while the G-Ch rocks show ferroan and NG-Ch rocks, slightly magnesian character (Fig 4f). When plotted on the Na_2O+K_2O-CaO wt% vs SiO_2 wt% diagram (Frost et al., 2001), the EPs straddle between the calcic and calc-alkalic fields while the G-Ch and NG-Ch samples plot on the calc-alkalic and alkali-calcic fields respectively (Fig 4g). In the normative Ab-An-Or diagram (Barker,1979), the EPs plot very close to An vertex while the G-Ch and NG-Ch rocks plot in the granite region (Fig 4h). As seen in Table 1, overall the high TiO_2, Al_2O_3, CaO, MgO, Fe_2O_3 and low SiO_2 and K_2O of the EPs are in sharp contrast with the composition ranges of these oxides for the G-Ch and NG-Ch rocks. The slightly higher CaO and lower Na_2O contents of the G-Ch rocks compared to the NG-Ch rocks are also notable.

5.1b Trace elements

The trace element data of orthopyroxene-bearing units is reported in Table 2. EPs exhibit lower and more variable Rb and Ba concentrations when compared to G-Ch rocks and NG-Ch rocks.

The Sr concentrations in EPs range narrowly (103.58-113.20 ppm), with one outlier at 296.20 ppm, while G-Ch rocks show similar levels (111.50-172.30 ppm) and NG-Ch rocks exhibit significantly higher values (326.85-340.60 ppm). EPs exhibit Nb and Y concentrations in the range 8.60-33.23 ppm and 28.11-66.90 ppm respectively while G-Ch rocks show narrower ranges ([18.82-19.70 ppm] and [42.84-43.08 ppm]). Notably, NG-Ch rocks have low concentrations of Nb and Y (6.75-10.80 ppm and 6.60-8.16 ppm respectively). Zr abundance of the EPs are in the range 115-229 ppm. In comparison, the G-Ch rocks show a higher Zr abundance (163-224 ppm) while the NG-Ch samples, although high, exhibits a larger variation (26-254ppm).

In the chondrite normalized diagram, the EPs show a flat to slightly enriched LREE pattern [$(La/Sm)_n$=1.30-2.13], with no discernible HREE depletion [$(Gd/Yb)_n$ = 1.40-2.16]. The $(La/Yb)_n$ values range from 2.19-5.47 indicating a slightly fractionated REE pattern. An isolated sample shows a slightly positive Eu anomaly (1.14). The G-Ch rocks exhibit slightly higher LREE values [$(La/Sm)_n$=2.89-3.65], slightly higher HREE depletion [$(Gd/Yb)_n$=2.29-2.87] and a more fractionated overall REE pattern [$(La/Yb)_n$=9.10-11.05]. They also show a negative to slightly positive Eu/Eu* [0.64-1.08]. The REE pattern of the NG-Ch rocks diverge markedly from the EPs and G-Ch rocks as it shows high LREE enrichment [$(La/Sm)_n$=4.90-8.94], significant HREE depletion [$(Gd/Yb)_n$=5.76-11.67], and therefore a highly fractionated REE pattern [$(La/Yb)_n$=26.46-80.46] and variable Eu anomaly [Eu/Eu*=0.85-1.55] (Fig 5a and b)

In the primitive mantle normalized multi-element diagram, the EPs show high variability in the Rb values (around 10 times to more than 100 times the primitive mantle values) while the Ba values show a comparatively narrower range. There is no notable Nb anomaly and the values are around 10 times the primitive mantle values. A slightly positive K anomaly is observed for three samples (around 100 times the PM value) and a negative Sr anomaly is present, except for one sample. The lack of any notable Ti or Y anomalies is also conspicuous in the diagram. On the other hand, the G-Ch and NG-Ch rocks show more elevated Rb and Ba values. The presence of negative Ti anomaly and strong positive K anomaly is also evident in both rock units. One NG-Ch sample shows a negative Th anomaly. A negative Zr anomaly for one NG-Ch sample is also evident in the spider plot. The G-Ch rocks show a negative Sr anomaly while the NG-Ch rocks show flat to slightly positive Sr anomaly. Other notable contrasting feature includes the presence of a negative Y, Nb anomaly in the NG-Ch rocks. (Fig 6a and b)

5.2 Mineral Chemistry

Representative mineral chemistry data of orthopyroxenes and biotites from each unit are provided in Table 3. Orthopyroxenes in EPs range from ~En_{51} to ~En_{40}. They also have higher TiO_2 (~0.1-0.2%) and CaO (~0.8%) but lower Al_2O_3 (~0.7-0.8%) compared to other units. In contrast, orthopyroxenes in the G-Ch are ~En_{34} and exhibit lower TiO_2 (~0.1%) and CaO (~0.5%) but higher Al_2O_3 (~1.3%). Orthopyroxenes in NG-Ch are ~En_{40} and show similar X_{Mg} to those found in one variant of EP but have higher Al_2O_3 (~1.0%) and lower CaO (~0.5%) and TiO_2 (<0.1%), distinguishing them from EP orthopyroxenes. Biotites in EPs have high X_{Mg} (~0.70) relative to those in G-Ch (~0.41) while biotites in NG-Ch have intermediate X_{Mg} (~0.55). TiO_2 contents are high in biotites from both EP (4.5%) and G-Ch (5.7-5.9%), while NG-Ch biotites contain slightly lower TiO_2 (3.8-4.0%). Feldspar compositions are listed in Table 4. EP plagioclase is highly calcic, with An content ranging from ~An_{82} to ~An_{66}. Plagioclase in NG-Ch and G-Ch is more sodic but distinct in composition (~An_{32} and ~An_{52}, respectively). Garnet and opaque (oxides) compositions are shown in Table 5. Garnets in G-Ch are predominantly Almandine (~$Alm_{70}Pyr_{13}Grs_{15}Sps_2$). Opaques across all orthopyroxene-bearing units are ilmenites, with FeO_t of ~48% in EP, ~52% in NG-Ch, and ~50% in G-Ch.

6. Discussion

The major, trace, and mineral chemistry data clearly indicate that the three orthopyroxene-bearing units are distinct entities and cannot be grouped under the same umbrella. Their characteristics suggest that each unit may have originated from independent sources, so the nature of the source regions must be evaluated using robust proxies.

6.1 Element mobility

Since all the rock units have undergone significant metamorphism, it is crucial to examine the mobility of elements under these settings. The mobility of LILEs under such settings have been well documented, but the REEs and other HFSEs are generally regarded to be largely unaffected by metamorphic fluid activity (Bau, 1991). The plots shown in Fig 7 a through c can be used to examine the mobility of elements. The K/Rb vs K(wt%) plot (Rudnick and Presper,1990) shows that the G-Ch , NG-Ch rocks plot in the field of common igneous and sedimentary rocks with a higher K_2O value, while EG appears perturbed (Fig 7a). The Rb vs Sr (log-log) plot also shows that the overall trend of EG is incoherent (Fig 7b). The Th/U vs La/Th plot is an effective way to examine the depletion of U and Th (Rudnick and Presper, 1990). It shows that only a NG-Ch sample plots on the U loss side (Fig 7c). Overall, the trend of the REEs and other HFSEs composition within each orthopyroxene bearing unit (Fig 5, Fig 6) suggests that they are robust and can be used to comment on the characteristics of their precursors. The HREE+Y depletion observed in the NG-Ch rocks can be considered intrinsic to their precursor lithologies because if it were the activity of metamorphic fluids, it would have been prevalent in other units in the terrane as well, but this depletion is exclusive to the NG-Ch rocks. The coherent trend observed in Sr abundances within each unit suggests that it could also be used to make some broad conclusions about the true nature of the unit. The higher Sr contents of the NG-Ch rocks when compared to the other two rock units would therefore be reflective of the characteristics of their precursor lithologies. Overall, it is the case that only the EP elements appear to be incoherent, thus these variations could be intrinsic to the precursor.

6.2 Key insights from mineral chemistry

The mineral chemistry data reaffirm the distinctness of each orthopyroxene-bearing unit, as established by the bulk rock data (Fig 8). Petrographic associations suggest that the orthopyroxenes are likely of metamorphic origin, with compositional variations possibly linked to the nature of the reactant ferromagnesian minerals. In the EPs and NG-Ch rocks, the close association between biotite and orthopyroxene indicates that biotite is the main ferromagnesian phase dehydrating to form orthopyroxene (Fig. 8c and d). This potentially explains the higher modal volume of orthopyroxenes in the EPs as it could be a reflection of the initially high modal volume of biotite in the unit. The scarcity of biotite in one EP sample may therefore reflect its near-complete consumption (Fig 3c). In contrast, the G-Ch rocks show a close association between garnet, biotite, and orthopyroxene, suggesting that both biotite and garnet are involved in orthopyroxene formation (Fig. 8e). Although distinguishing metamorphic from igneous orthopyroxenes petrographically can be challenging in high-grade rocks (Touret & Huizenga, 2012), mineral chemistry provides a broader basis for distinction. Rajesh and Santosh (2011) compared X_{Mg} (molar Mg/(Mg+Fe)) and X_{Al} (=Al/2, where Al is the cation proportion for six oxygens) in orthopyroxenes from metamorphic zones and unmetamorphosed igneous charnockites. Their diagram showed that orthopyroxenes of igneous origin tend to exhibit lower and narrower X_{Al} values than those from metamorphic zones. Although they did not recommend the plot as a strict discrimination tool, it can still be used for broad classification, keeping in mind that it does not account for bulk rock compositional variations, as the reference dataset included only charnockites (sensu stricto). X_{Mg} and X_{Al} values of orthopyroxenes from the present study are plotted on the diagram adapted from (Meshram et al., 2021) (Fig. 8b). G-Ch orthopyroxenes fall within the dehydration field, EG orthopyroxenes plot mainly within the igneous field, and NG-Ch orthopyroxenes straddle both fields. However, as mentioned before, this diagram alone cannot be used to reach a conclusion regarding the EP orthopyroxenes as the plot does not appear to be useful in informing us about the composition range of orthopyroxenes in mafic-intermediate units. In this context, it is necessary to consider the composition of the reactant phases involved while interpreting the origin of orthopyroxenes. The higher X_{Mg} and lower X_{Al} in EP orthopyroxenes likely reflect the the higher X_{Mg} of associated biotites, which are compositionally phlogopitic (Fig. 9a). Their lower X_{Al} content also corresponds with the lower Al content of the reactant biotite (Table 3). Similarly, within the G-Ch rocks, garnets in contact with orthopyroxene show lower X_{Mg} (~0.15) compared to other G-Ch garnets (X_{Mg} ~0.17), supporting garnet breakdown during orthopyroxene formation. When it comes to the plagioclases, the ones found in EPs are clearly distinct from those in NG-Ch and G-Ch rocks. The higher An content (~82) is highly unlike the composition observed in regular granitoids. (Fig 9b). The plagioclase compositions of the G-Ch and NG-Ch shows that the range is within what is generally observed in granitoids, however, G-Ch plagioclase possess a slightly higher An content.

6.3 Implications of geochemistry for tectonic setting

Since tectonic discrimination diagrams employing mobile elements may not reliably represent the rock units in this study, they were excluded. Although the Ta vs Yb plot (Pearce et al., 1984) shows that all rock units plot largely within the volcanic arc field, only the NG-Ch rocks can be confidently categorized as exhibiting a proper VAG like signature (Fig 7f). As Sr abundances are coherent, as discussed in section 6.1, the Sr/Y vs Y diagram (Martin, 1986; Hansen et al., 2002; Defant and Drummond, 1990) was used, showing that the NG-Ch rocks plot fairly within the adakite field, with Sr/Y values ranging between 41 and 47 (Fig 7e). Given that the study area is a complex terrane, the actual scenario is likely more nuanced than what is suggested solely by discrimination diagrams, as inputs from multiple sources and partial melting events of various source units, depending on their fertility and heat sources, may occur in orogenic belts. The geochemical data appears to be in-line with potential processes in a convergent tectonic environment. However, whether they are more indicative of an accretionary or collisional setting requires further evaluation through more detailed studies and comparison with rock suites from well-studied regions. These aspects are undertaken and further discussed in section 6.3.

6.3 Original nature of precursors

6.3a EP

Although enderbitic units are generally interpreted to represent metamorphosed tonalites, this rock suite appears to be different and not fit that line of interpretation. The bulk rock compositions are mostly mafic, with composition of plagioclase ranging from An_{82} to An_{66}, which is not the range generally expected if it is a tonalite (sensu stricto) (Le Maitre, 1976, 2002). Although the modal proportions of quartz is equal to or slightly above 20% in selected samples, most of the varieties do not possess the characteristic modal volume observed in tonalites, with some even plotting in the diorite/gabbro/anorthosite region (Fig 4a). Previous studies attributed this to igneous grading, but the variations in the modal volume and bulk rock compositions suggest that an alternative process might be the reason behind this nature of the unit. As the normative feldspar diagram reveals (Fig 4h), the samples are not plotting in the tonalite region but instead very close to the anorthite vertex. This is reflective of the highly calcic composition of the plagioclases. Interestingly, a recent study (Bose et al., 2020) has argued for the presence of a metagreywacke unit in the field, and after the comparison of field images and photomicrographs, it is clear that the OFGs they present and the EPs discussed in this study are the same entities. Clearly, a depositional origin can better explain the trends observed in the bulk rock chemistry. Under this interpretation of the unit, the variations in the Eu/Eu*(0.39-1.14) can be suggestive of a multi-component provenance, and based on such signatures along with extremely high An composition observed in non-migmatitic variants of EP and the overall high modal volume of such grains, this study speculates the presence of mafic components in the provenance for the unit, along with the felsic components already suggested based on the study of detrital zircons from this unit (Bose et al., 2020). Another interesting implication of this interpretation is that the ultramafic xenoliths previously suggested to be within this unit are possibly lithic fragments of an ultramafic component in the provenance. This essentially suggests a short transport distance and mineralogical immaturity, suggesting rapid erosion of nearby mafic-ultramafic sources. This is in alignment with the interpretation that the unit likely represents arc-derived turbidity sequences (Bose et al., 2020) Overall the bulk rock and mineral chemistry data are more in-line with the conclusions drawn regarding the metagreywacke unit in the region (Bose et al., 2020) and it highlights the complexity of the depositional history of the terrane. A detailed provenance study of this unit can provide much insights on the earliest history of the Chilka lake area.

6.3b G-Ch

The G-Ch suite appears to represent a largely homogenous entity. Although the G-Ch suite differs from the leptynites in macroscopic appearance, the two are similar petrographically, with the presence of orthopyroxene and a more coarser grained nature in the former being the only distinguishing features under thin section. The average composition of Post-Archean Indian shale (Ray and Paul, 2021) and the composition of a leptynite sample from the Hatia quarry has been plotted on a spider diagram for reference (Fig 10). Since it is established that the precursors to the leptynites arose via the mica dehydration melting of the previously existing metapelitic suite in the region (Sen&Bhattacharya, 1997), this study was compelled to test the validity of such a mode of petrogenesis for the G-Ch rocks. However, unlike the leptynites, the G-Ch rocks are coarse grained and mesocratic to melanocratic in appearance, as previously mentioned. They also appear to contain a higher modal volume of ilmenite and biotite. The most contrasting aspect is the metaluminous nature of the G-Ch, when compared to the largely peraluminous nature of the metaleucogneisses. Partial melting studies conducted on metasedimentary units emphasize that only peraluminous leucogranitoids may be regarded as representative of liquid melts, and that other granitoids either involve some form of restite influence or are hybrid entities with mafic inputs (Montel and Vielzeuf, 1997). In light of such findings, it is highly unlikely that the G-Ch is akin to a peraluminous S-type granitoid. In addition, a partial melting based origin with a restite dominated pattern from a metapelitic source also doesn’t seem to align with the characteristics of the G-Ch rocks as such entities reported from within the leucogneissses were also peraluminous. (Sen and Bhattcharya, 1997). The partial melting of a greywacke entity in the field also cannot explain the unit, as the melts would still be peraluminous, based on leucosome studies of a greywacke within the metasedimentary suite (Sen and Bhattacharya, 1997), and these entities exhibit significant melt retention (Bose et al., 2020) such that melts cannot migrate from the source efficiently either.

Given these aspects, an interesting possibility is a deep crustal hybrid melting. Although isotopic studies would be needed to establish such a mode of petrogenesis, as the G-Ch represents a high-

K metaluminous unit with high Rb, Ba, moderate CaO and low Sr, along with variations in Eu abundance and a lack of any Nb depletions, it is showing a hybrid nature, therefore, it is speculated that a largely well-equilibriated hybrid melt of a suitable deep metasedimentary unit and possibly granodioritic gneisses could perhaps explain the G-Ch composition. As inferences based on major and trace element chemistry alone cannot be conclusive, this mode of petrogenesis is speculative at this stage. It is suspected that the melt was largely homogenized before crystallization. This is observable in the trace element chemistry, as the two samples collected from two distant quarries (see Fig. 1b) exhibit nearly comparable compositions. The orthopyroxenes, as previously mentioned are likely metamorphic in origin, with the mineral forming reaction involving the breakdown of both garnet and biotite. Finally, ~An_{50} plagioclase in the sample also suggests a non-pelitic influence, most likely from a Ca-rich crustal component, consistent with a hybrid origin.

As the region has undergone significant quarrying, the true extent of this unit could not be ascertained, but based on the limited occurrences, it is suspected that the unit is not a voluminous one. As this is a high-grade terrane, the true textures of protoliths are lost and given that textural indicators are typically subtle in such felsic–felsic mixing, no definitive textural evidence could be identified. Nevertheless, this study is tempted to propose such an origin for the rock unit based on its chemical characteristics. However, this petrogenetic model is speculative and needs to be verified using a larger dataset and isotope systematics.

6.3c NG-Ch

The NG-Ch rocks as previously discussed is a distinct entity characterized by its highly fractionated REE pattern, low Nb, Ti, Y and higher Sr concentration. Based on the volume of the unit as observed in the field, and the absence of any mafic-intermediate unit with a similar geochemical signature in the area, the NG-Ch precursors are more likely representative of products of partial melting rather than products of fractionation. Coming to a possible source region, the most likely explanation, in light of the current understanding of the geodynamic history of the region, appears to be melting of a crustal source which already possess the aforementioned geochemical signatures. As the HREE+Y depletion is suggestive of residual garnet, melting of older lower crustal TTG gneisses (possibly formed in an arc setting during a past event) which can generate sufficient volume of melt to eventually rise and emplace into the upper crustal metasedimentary suite appears to be a plausible model. This way, the negative Nb, Ti anomaly, high $(La/Yb)_n$ and high Sr/Y ratios could be explained as signatures directly inherited from the source region.

However, fluid absent melting experiments on sodic TTGs suggest that the resulting melts are usually low in K_2O (Watkins et al., 2007), thereby indicating that significant volumes of high K_2O granites cannot be explained as melt products of TTGs. Interestingly, taking into consideration this 'low K_2O' problem, a granite suite identified in the Fiskenæsset region having a 'TTG' like signature just like the NG-Ch was interpreted to be a product of partial melting of predominantly granodioritic component of a lower crustal TTG, thereby explaining its higher K_2O abundances ((Huang et al., 2016). This interpretation was further corroborated by the presence of negative Eu anomalies in representative granites. The compositional range of the Fiskenæsset region granites (excluding altered samples and cumulates) is shown in Fig 11. Interestingly, the NG-Ch rocks exhibit a similar trend, but with notable differences in the Eu anomalies and degree of HREE depletion. Although the current dataset of NG-Ch rocks is limited, the trends in the Eu anomaly and ΣREE are suggestive of melt evolution, with the sample with the least ΣREE (~70ppm) showing a positive Eu anomaly (~1.55) and a more evolved sample (ΣREE= ~285ppm) showing a slightly negative Eu anomaly (~0.85). A comparison of the composition range of Fiskenæsset region granites with the trend of the NG-Ch suite suggests that the latter is likely more enriched in incompatible elements overall. This, by extension, reflects a slightly more enriched source for the NG-Ch. However, irrespective of these nuances in the nature of source TTG, the fractionated pattern is common to both, and the signature anomalies in the spider plots also align, implying a similar mode of petrogenesis. In light of these aspects, this study suggests that the NG-Ch rocks were likely formed via dehydration melting of a lower crustal TTG gneiss, with melts derived predominantly from the granodioritic component. Though the NG-Ch rocks are voluminous enough to be mappable, they are not of batholith scale, therefore such a mode of petrogenesis appears consistent. Regarding the mechanism that could have triggered this melting, crustal thickening associated with the peak of the orogenic event in the region (D2) appears to be the most plausible. The presence of this entity, therefore, signifies the complex crustal anatexis history of the Chilka lake area.

6.4 Evolution of terrane

Based on the obtained understanding of the origin of the orthopyroxene-bearing entities in the field, it is important to properly place them in the timescale of the evolution of the Chilka lake area. If we go by the argued similar deformational histories of the metagreywackes and metapelites, then it seems to be the case that both these sedimentary units were likely established before the first deformation event (D1), thereby making them the oldest entities in the region. On the basis of EPMA dating of monazites from khondalites, the timing of the D1 event has been estimated to be between 1250 and 1100 Ma (Simmat and Raith, 2008). The leucogranitoids are interpreted to have been formed during the foliation-defining and crustal anatexis phase (D2), which is identified to have occurred between 980 and 940 Ma (Simmat and Raith, 2008), thereby making them a later addition. However, although the G-Ch could be explained via D2 event, previous studies have argued for its pre-F2 status (Bhattacharya et al., 1994), therefore, based on relative dating alone, it is difficult to place this event, only geochronological studies could reveal the true timing of its origin. As the origin of NG-Ch rock precursors appears to involve crustal anatexis, it is likely that they also formed during the D2 stage, as they too preserve the pervasive S2 foliation. However, radiometric dating is necessary to verify this. The precursors to mafic granulites and anorthosites identified further south of the Chilka lake area are potentially linked to Rodinia breakup (Bose et al., 2011b; Chakrabarti et al., 2011), which suggests that they are possibly younger than the aforementioned entities. This study therefore highlights that crustal reworking was the dominant process during the Rodinia assembly, with notable juvenile additions occurring only later, during Rodinia breakup.

7. Conclusions

- The geochemical study of orthopyroxene-bearing entities in the northern Chilka Lake area suggests that each identified unit, i.e., EPs, G-Ch, and NG-Ch, represents a distinct lithology originating from an independent source.
- The field aspects, bulk rock and mineral chemistry data of the EPs are more consistent with a depositional origin, with a provenance likely involving, felsic, mafic and minor ultramafic rocks, thereby reaffirming their previous characterization as akin to arc derived turbidity sequences.
- The G-Ch rocks are speculated to represent a deep crustal hybrid melt from a suitable metasedimentary source and an orthogneiss (possibly granodioritic).

- NG-Ch rocks are likely products of partial melting of predominantly granodioritic component of a lower crustal TTG, with their characteristic signature likely inherited from the source region.
- This study emphasizes the intricate depositional history and complex crustal anatexis events of the Chilka lake area. It also highlights the overall complexity of granulite terranes, which can alter the original nature of rock units to such an extent that their present characteristics can be misleading.

8. Future Prospects

As discussed earlier, high-resolution geochronological studies are necessary to better constrain the timing of origin for the identified units. In addition, isotopic proxies should be employed to evaluate the validity of the petrogenetic models proposed in this study. It appears to be the case that each distinct member of the metasedimentary suite and the NG-Ch rocks were likely not represented in previous comprehensive isotopic studies of Domain 3 (Rickers et al., 2001). The provenance studies of the prevalent metagreywacke unit will also shed light on the true nature of the source of sediments, therefore readers with expertise in that domain are encouraged to take up such research.

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Tables

Table 1: Major-oxide compositions of orthopyroxene-bearing entities and one leptynite

Sample	GRL02	CAL04	HAT04	CAL02	USIS02	HAT01	CAL01	USIS01	GRL03	HATLP
Rock Type	EP	EP	EP	EP	EP	G-Ch	G-Ch	NG-Ch	NG-Ch	Leptynite
Major oxides (wt%)										
SiO_2	42.70	49.44	51.36	52.48	51.42	66.54	68.66	67.29	68.03	71.01
TiO_2	1.96	1.31	1.50	1.35	2.41	0.72	0.63	0.35	0.48	0.35
Al_2O_3	19.86	17.37	17.29	15.83	17.63	15.20	14.69	14.47	15.28	14.13
Fe_2O_3(t)	15.27	11.90	11.83	12.19	13.42	4.21	3.77	4.88	3.28	2.65
MnO	0.16	0.16	0.15	0.14	0.19	0.05	0.05	0.04	0.04	0.04
MgO	5.81	6.38	5.06	5.95	2.77	0.69	0.57	1.45	0.85	0.21
CaO	11.67	9.83	8.32	7.87	8.06	3.66	3.26	2.7	2.62	2.69
Na_2O	1.11	1.61	1.65	1.24	1.92	1.60	1.98	2.53	2.65	2.08
K_2O	0.68	1.52	1.91	2.17	0.59	6.27	5.53	5.53	6.07	6.00
P_2O_5	0.31	0.17	0.38	0.31	0.84	0.35	0.31	0.25	0.18	0.30
Total	99.53	99.69	99.45	99.53	99.25	99.29	99.45	99.49	99.48	99.46
Mg#	43.00	51.53	45.89	49.18	29.04	24.52	23.06	37.07	33.94	13.59
K_2O/Na_2O	0.61	0.94	1.16	1.75	0.31	3.92	2.79	2.18	2.29	2.88

Mg# = molar $Mg/(Mg+Fe^{2+})$

Table 2: Trace element compositions of orthopyroxene-bearing entities and one leptynite

Sample	GRL02	CAL04	HAT04	CAL02	USIS02	HAT01	CAL01	USIS01	GRL03	HATLP
Rock Type	EG	EG	EG	EG	EG	G-Ch	G-Ch	NG-Ch	NG-Ch	Leptynite
Trace elements (ppm)										
Rb	3.86	72.15	12.58	134.70	4.23	219.20	275.60	167.10	163.30	202.70
Ba	84.38	118.60	58.47	238.30	66.78	1618.00	875.80	1280.00	1722.00	601.40
Sr	296.20	104.70	103.58	113.20	111.00	172.30	111.50	313.10	340.60	82.70
Pb	2.44	8.52	9.65	15.51	7.54	32.08	38.51	27.82	36.42	35.21
Th	0.55	0.09	0.14	4.07	1.15	4.36	6.52	0.26	26.08	3.88
U	0.13	0.05	0.17	1.86	0.28	0.28	0.46	0.29	0.68	0.32
Zr#	135	115	154	124	229	224	163	26	254	272
Ta	0.83	0.68	0.85	0.84	1.81	1.09	1.11	0.20	0.37	0.84
Y	28.11	38.42	40.47	41.41	66.90	42.84	43.08	6.60	8.15	40.28
Nb	9.65	8.60	15.32	11.61	33.23	19.70	18.82	6.76	10.80	13.39
Sc	36.13	38.73	36.25	39.38	44.55	8.29	5.84	4.88	2.80	8.40
Cr	80.30	301.70	254.00	216.10	85.48	13.54	10.91	187.50	14.42	5.39
Ni	72.02	47.63	35.77	34.84	39.32	11.56	11.22	67.41	9.63	3.29
Co	81.56	69.91	55.80	69.50	71.48	55.69	59.69	47.35	38.92	28.45
V	267.10	239.10	207.39	223.80	193.50	50.32	39.61	75.14	45.06	14.42
Zn	23.51	18.85	10.11	32.24	26.92	15.00	15.36	12.34	10.20	44.40
Cu	61.43	47.96	29.96	49.50	103.20	29.91	21.83	48.98	14.34	8.73
La	8.05	15.99	23.19	30.21	38.63	51.42	50.92	16.6	76.05	45.61
Ce	20.73	36.38	58.93	71.92	92.66	101.8	106.90	29.29	131.30	90.61
Pr	3.09	4.86	6.69	9.38	12.30	11.54	12.42	3.16	13.40	10.35
Nd	14.47	19.79	28.04	38.25	53.66	45.52	50.65	11.52	42.81	39.48
Sm	4.00	5.11	7.02	9.15	13.79	9.09	11.35	2.19	5.52	8.40
Eu	1.59	1.17	1.28	1.37	1.85	3.59	2.43	1.33	2.04	1.76
Gd	4.94	6.05	7.22	10.36	15.25	11.26	11.51	3.14	9.57	9.34
Tb	0.76	0.95	1.18	1.34	2.19	1.38	1.62	0.29	0.61	1.33
Dy	4.99	6.23	7.49	7.58	12.27	7.43	8.29	1.26	1.80	7.73
Ho	1.04	1.32	1.53	1.51	2.44	1.54	1.52	0.23	0.32	1.49
Er	2.87	3.74	4.30	4.14	6.61	4.29	3.55	0.56	0.80	4.01
Tm	0.43	0.58	0.63	0.62	0.96	0.61	0.50	0.07	0.10	0.57
Yb	2.64	3.59	3.91	3.96	6.15	4.05	3.30	0.45	0.68	3.53
Lu	0.38	0.51	0.58	0.57	0.88	0.57	0.46	0.06	0.10	0.51
ΣREE	69.53	106.31	152.01	190.37	259.65	254.09	265.44	70.15	285.10	224.72
Eu/Eu*	1.14	0.64	0.55	0.43	0.39	1.08	0.64	1.55	0.85	0.61
$(Gd/Yb)_{CN}$	1.41	1.39	1.52	2.16	2.05	2.30	2.88	5.76	11.67	2.19
$(La/Yb)_{CN}$	2.19	3.20	4.25	5.47	4.51	9.10	11.04	26.46	80.45	9.28

CN- Chondrite normalized, Eu/Eu*= $(Eu)_{CN}/[(Sm)_{CN} \times (Gd)_{CN}]^{0.5}$, #data from XRF

Table 3: Microprobe data and cation proportions of Orthopyroxene and Biotite

Sample	GRL02	USIS02	HAT01	GRL03	GRL02	HAT01	HAT01	GRL03	GRL03
Mineral	Opx	Opx	Opx	Opx	Bt	Bt	Bt	Bt	Bt
SiO_2	51.76	50.05	48.62	49.68	38.41	35.35	35.75	37.89	38.01
TiO_2	0.17	0.10	0.07	0.07	4.54	5.91	5.71	3.80	3.98
Al_2O_3	0.79	0.74	1.30	1.03	12.37	13.06	13.08	13.56	13.60
Cr_2O_3	0.00	0.00	0.06	0.02	0.17	0.07	0.04	0.05	0.00
FeO_t	28.88	34.33	38.11	34.73	12.63	22.60	22.73	18.01	18.14
MnO	0.49	0.82	0.37	0.50	0.01	0.02	0.04	0.04	0.00
MgO	17.39	13.32	11.54	13.48	16.36	8.75	8.64	12.60	12.99
CaO	0.77	0.80	0.50	0.52	0.01	0.05	0.00	0.04	0.04
Na_2O	0.03	0.07	0.00	0.05	0.06	0.00	0.00	0.09	0.10
K_2O	0.00	0.00	0.00	0.00	10.35	9.66	9.99	10.15	10.12
F	0.00	0.00	0.00	0.01	3.16	0.92	1.12	-	-
Cl	0.01	0.02	0.01	0.01	0.25	0.07	0.07	-	-
H_2O_{calc}	-	-	-	-	2.59	3.43	3.36	3.98	4.01
O=F+Cl	-	-	-	-	-1.39	-0.41	-0.49	-	-
Total	100.29	100.25	100.59	100.09	99.52	99.48	100.04	100.19	100.99
O basis	6	6	6	6	22	22	22	22	22
Si	1.98	1.98	1.95	1.97	5.55	5.46	5.49	5.71	5.68
Ti	0.00	0.00	0.00	0.00	0.49	0.69	0.66	0.43	0.48
Al	0.04	0.03	0.06	0.05	2.10	2.38	2.37	2.41	2.40
Cr	0.00	0.00	0.00	0.00	0.02	0.01	0.01	0.01	0.00
Fe^{2+}	0.93	1.14	1.28	1.15	1.53	2.92	2.92	2.27	2.27
Mn	0.02	0.03	0.01	0.02	0.00	0.00	0.01	0.01	0.00
Mg	0.99	0.79	0.69	0.80	3.52	2.02	1.98	2.83	2.90
Ca	0.03	0.03	0.02	0.02	0.00	0.01	0.00	0.01	0.01
Na	0.00	0.01	0.00	0.00	0.02	0.00	0.00	0.03	0.03
K	0.00	0.00	0.00	0.00	1.91	1.90	1.96	1.95	1.93
F	0.00	0.00	0.00	0.00	1.44	0.45	0.54	-	-
Cl	0.00	0.00	0.00	0.00	0.06	0.02	0.02	-	-
OH	-	-	-	-	2.50	3.53	3.44	4.00	4.00
X_{Mg}	0.52	0.41	0.35	0.41	0.70	0.41	0.40	0.55	0.56

XMg = Mg/(Mg+Fe^{2+}), calc- H_2O contents of Bt calculated by stoichiometry

Table 4: Microprobe data and cation proportions of feldspars

Sample	GRL02	GRL02	USIS02	USIS02	GRL03	GRL03	HAT01	HAT01
Mineral	Pl(core)	Pl(rim)	Pl(core)	Pl(rim)	Pl(core)	Pl(rim)	Pl(core)	Pl(rim)
SiO_2	48.00	47.96	51.45	51.32	60.27	60.33	54.17	53.27
TiO_2	0.00	0.00	0.03	0.04	0.02	0.04	0.00	0.00
Al_2O_3	33.28	33.64	30.80	31.02	24.68	24.88	28.74	29.33
Cr_2O_3	0.00	0.00	0.00	0.01	0.00	0.01	0.00	0.03
FeO_t	0.10	0.19	0.16	0.21	0.08	0.19	0.05	0.15
MnO	0.00	0.04	0.00	0.02	0.00	0.00	0.00	0.00
MgO	0.00	0.02	0.01	0.00	0.00	0.00	0.00	0.00
CaO	16.69	16.76	13.93	13.98	6.92	6.88	11.04	12.01
Na_2O	2.04	2.14	3.69	3.53	7.90	7.82	5.44	4.81
K_2O	0.04	0.05	0.18	0.14	0.24	0.20	0.24	0.21
Total	100.15	100.81	100.25	100.28	100.11	100.34	99.68	99.83
O basis	8	8	8	8	8	8	8	8
Si	2.20	2.18	2.34	2.33	2.69	2.68	2.46	2.42
Ti	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	1.79	1.81	1.65	1.66	1.30	1.30	1.54	1.57
Cr	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe^{2+}	0.00	0.01	0.01	0.01	0.00	0.01	0.00	0.01
Mn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Mg	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ca	0.82	0.82	0.68	0.68	0.33	0.33	0.54	0.58
Na	0.18	0.19	0.32	0.31	0.68	0.67	0.48	0.42
K	0.00	0.00	0.01	0.01	0.01	0.01	0.01	0.01
An	81.70	81.00	66.91	68.08	32.17	32.34	52.14	57.31
Ab	18.07	18.70	32.07	31.11	66.47	66.53	46.50	41.51
Or	0.23	0.30	1.01	0.81	1.35	1.13	1.36	1.18

Table 5: Microprobe data and cation proportions of Garnet and Ilmenite

Sample Mineral	HAT01 Grt	HAT01 Grt	GRL02 Ilm	USIS02 Ilm	GRL03 Ilm	HAT01 Ilm
SiO_2	37.65	37.73	0.00	0.00	0.00	0.00
TiO_2	0.00	0.03	52.67	50.50	47.50	50.14
Al_2O_3	20.69	20.92	0.03	0.02	0.02	0.01
Cr_2O_3	0.00	0.00	0.20	0.09	0.12	0.04
FeO_t	32.19	32.08	47.07	48.58	52.19	50.07
MnO	0.87	0.85	0.49	0.49	0.25	0.20
MgO	3.16	3.64	0.25	0.44	0.21	0.24
CaO	5.35	5.23	0.00	0.00	0.00	0.00
Na_2O	0.00	0.00	0.04	0.02	0.03	0.00
K_2O	0.00	0.00	0.01	0.00	0.00	0.00
Total	99.92	100.48	100.74	100.14	100.31	100.73
O basis	12	12	3	3	3	3
Si	3.01	3.00	0.00	0.00	0.00	0.00
Ti	0.00	0.00	0.99	0.97	0.93	0.96
Al	1.95	1.96	0.00	0.00	0.00	0.00
Cr	0.00	0.00	0.00	0.00	0.00	0.00
Fe^{2+}	2.15	2.13	0.99	1.03	1.13	1.07
Mn	0.06	0.06	0.01	0.01	0.01	0.00
Mg	0.38	0.43	0.01	0.02	0.01	0.01
Ca	0.46	0.44	0.00	0.00	0.00	0.00
Na	0.00	0.00	0.00	0.00	0.00	0.00
K	0.00	0.00	0.00	0.00	0.00	0.00
Alm	70.49	69.61	-	-	-	-
Pyr	12.46	14.05	-	-	-	-
Grs	15.08	14.38	-	-	-	-
Sps	1.97	1.96	-	-	-	-

Figure Captions

Fig 1: Simplified Geological map of the study region. (a) Modified after Geological and Mineral Map of Orissa (2011) - Geological Survey of India. (b) Lithological map of the studied stone quarries showing sample locations.

Fig 2: Field images from the study area (a) Folded EP layer within metatexite metapelite. (b) Non-garnetiferous charnockite with prominent S2 foliation. (c) Dismembered blocks of EPs within the lepytnite. (d) A close-up of EP and leptynite. (e) Garnetiferous charnockite with prominent S2 foliation. (f) Garnetiferous charnockite displaying pervasive S2 foliation, as observed in the southernmost quarry.

Fig 3: Photomicrographs of orthopyroxene-bearing units. (a) Representative photomicrograph of EP exhibiting the characteristic granoblastic texture. (b) Biotite in contact with orthopyroxene in EP. (c) Orthopyroxene and ilmenite in EP with a paucity in biotite. (d) Representative photomicrograph of non-garnetiferous charnockite showing the typical granoblastic texture. (e) Orthopyroxene in contact with biotite and ilmenite in the non garnetiferous charnockite. (f) Orthopyroxene in contact with garnet, ilmenite, and biotite in the garnetiferous charnockite. (g) Elongate orthopyroxene in contact with garnet, defining the S2 foliation in garnetiferous charnockite. (h) Close-up of the ubiquitous myrmekitic texture along the boundary between alkali and plagioclase feldspars in garnetiferous charnockite. Mineral abbreviations follow IMA conventions

Fig 4: Classification diagram for orthopyroxene bearing units (a) QAP plot (After Frost and Frost, 2008) (b) TAS diagram (After Le Bas et al. 1986) (c) A/CNK [$Al_2O3/(CaO+Na_2O+K_2O)$] vs. A/NK [$Al_2O_3/(Na_2O+K_2O)$] diagram (After Shand, 1943) (d) AFM diagram (After Irvine&Baragar, 1971) (e) FeO_t/MgO vs SiO_2 discrimination diagram (After Miyashiro, 1974) (f) $FeO_t/(FeO_t+MgO)$ wt% vs SiO_2 (After Frost and Frost,2008) (g) Na_2O+K_2O-CaO (MALI) vs. SiO_2 diagram (Frost et al., 2001) (h) Normative feldspar diagram (After Barker 1979). Legend same as Fig 1b

Fig 5: Chondrite normalized REE plots for (a) EP, (b) G-Ch and NG-Ch. Normalization values from (Sun and McDonough, 1989). Average composition of UCC from (Taylor and McLennan, 1995) and PAAS from (McLennan, 1989)

Fig 6: Primitive mantle normalized diagram for (a) EP, (b) G-Ch and NG-Ch. Normalization values from (Sun and McDonough, 1989). Average composition of UCC from (Taylor and Mclennan, 1995)

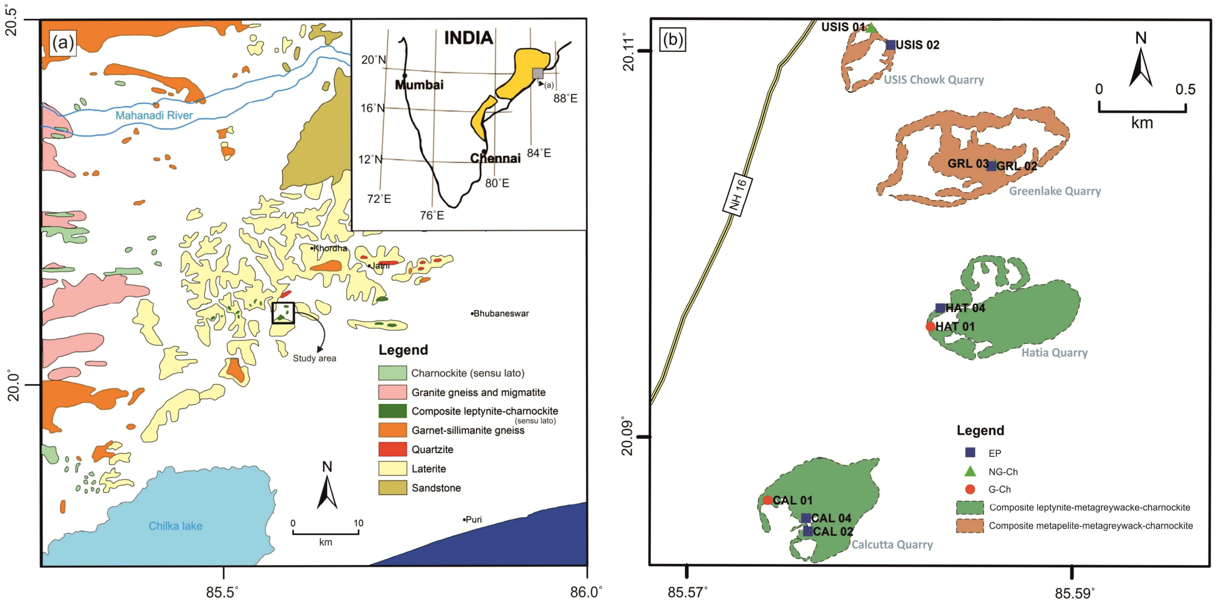
Fig 7: Geochemical plots to assess element mobility and plausible tectonic setting (a) K/Rb vs K(wt%) (Rudnick and Presper, 1990) (b) Rb vs Sr. UCC and mantle values from (Taylor and McLennan, 1995) and (Sun and McDonough, 1989) respectively. Th/U vs La/Th (Rudnick and Presper, 1990) (d) Eu/Eu* vs ΣREE (e) Sr/Y vs Y (Martin, 1986; Hansen et al., 2002; Defant and Drummond, 1990) (f) Ta vs Yb (Pearce et al., 1984) See main text for discussion.

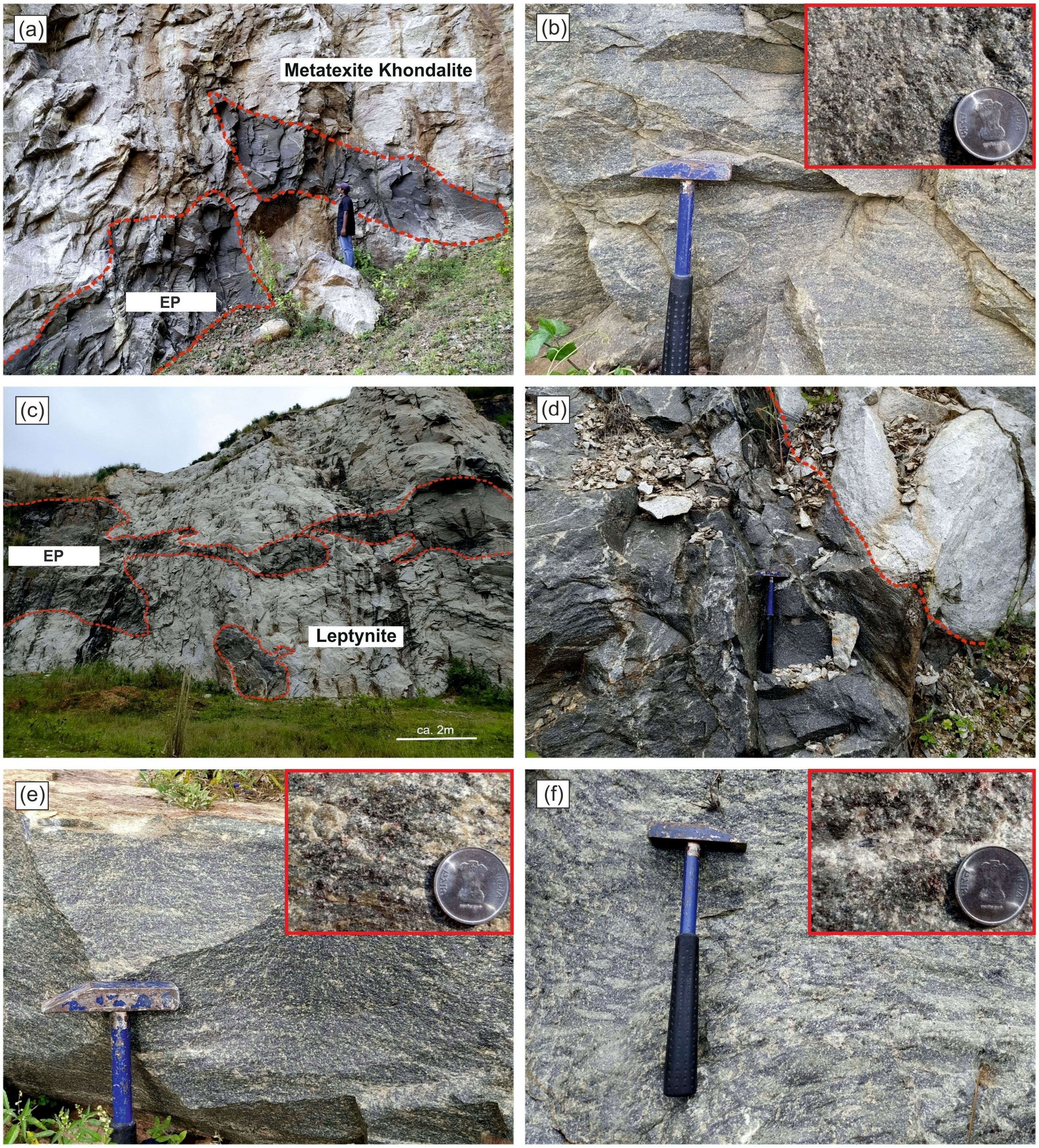
Fig 8: Classification diagrams and BSE images of orthopyroxenes (a) Ternary classification diagram (Wo-En-Fs) (Morimoto et al., 1988) (b)XAl vs XMg plot (Rajesh and Santosh, 2011) (Meshram et al., 2021). See main text for discussion. (c)Biotite in contact with orthopyroxene in EP. (d)Biotite within orthopyroxene in NG-Ch. (e) Orthopyroxene in contact with both garnet and biotite in G-Ch. Mineral abbreviations follow IMA conventions

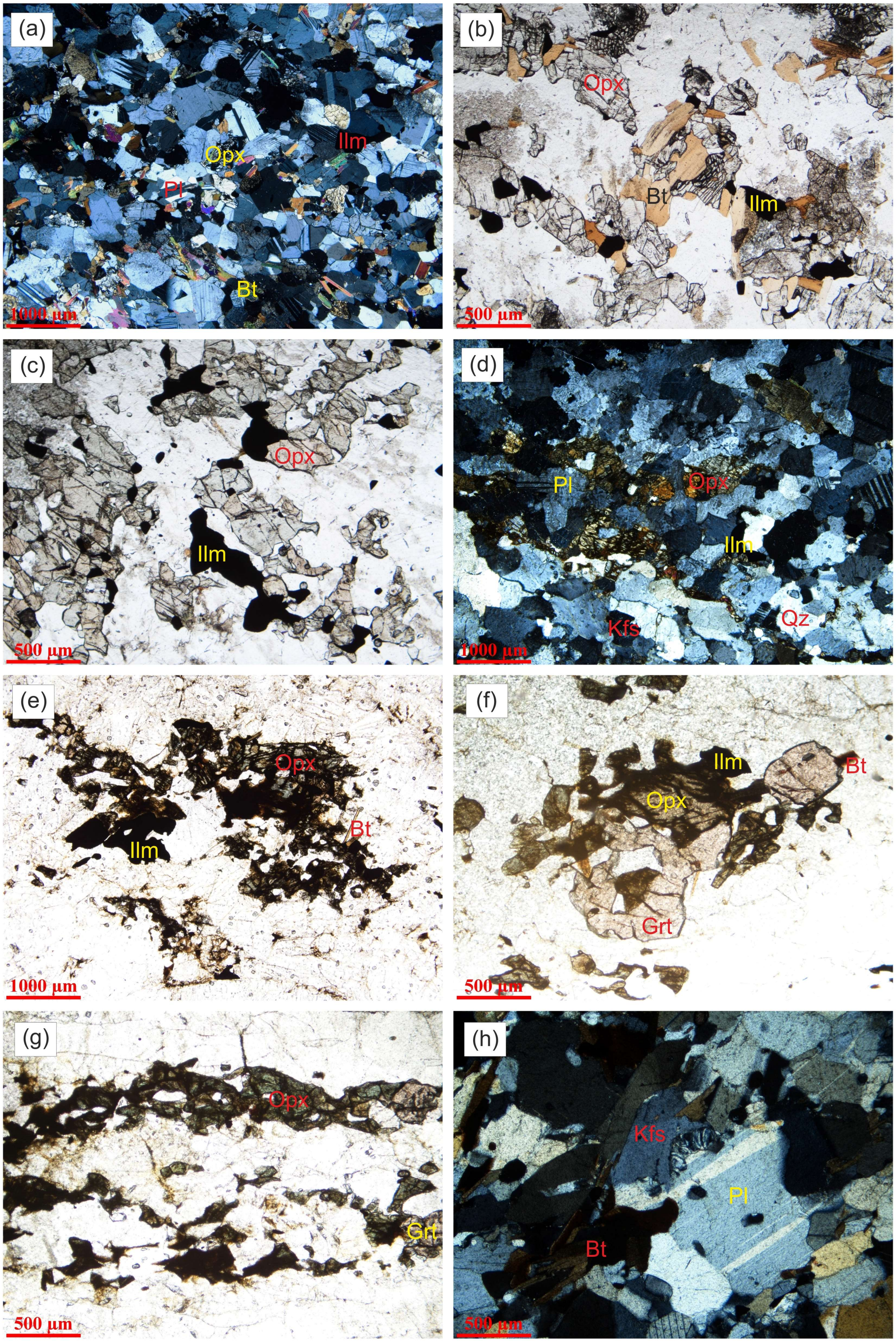
Fig 9: Classification diagrams for biotite and feldspars (a) Biotite classification plot using Al^{IV} vs Fe/ (Fe + Mg) (Deer et al., 1992). (b) Ternary classification diagram Ab-An-Or for feldspars

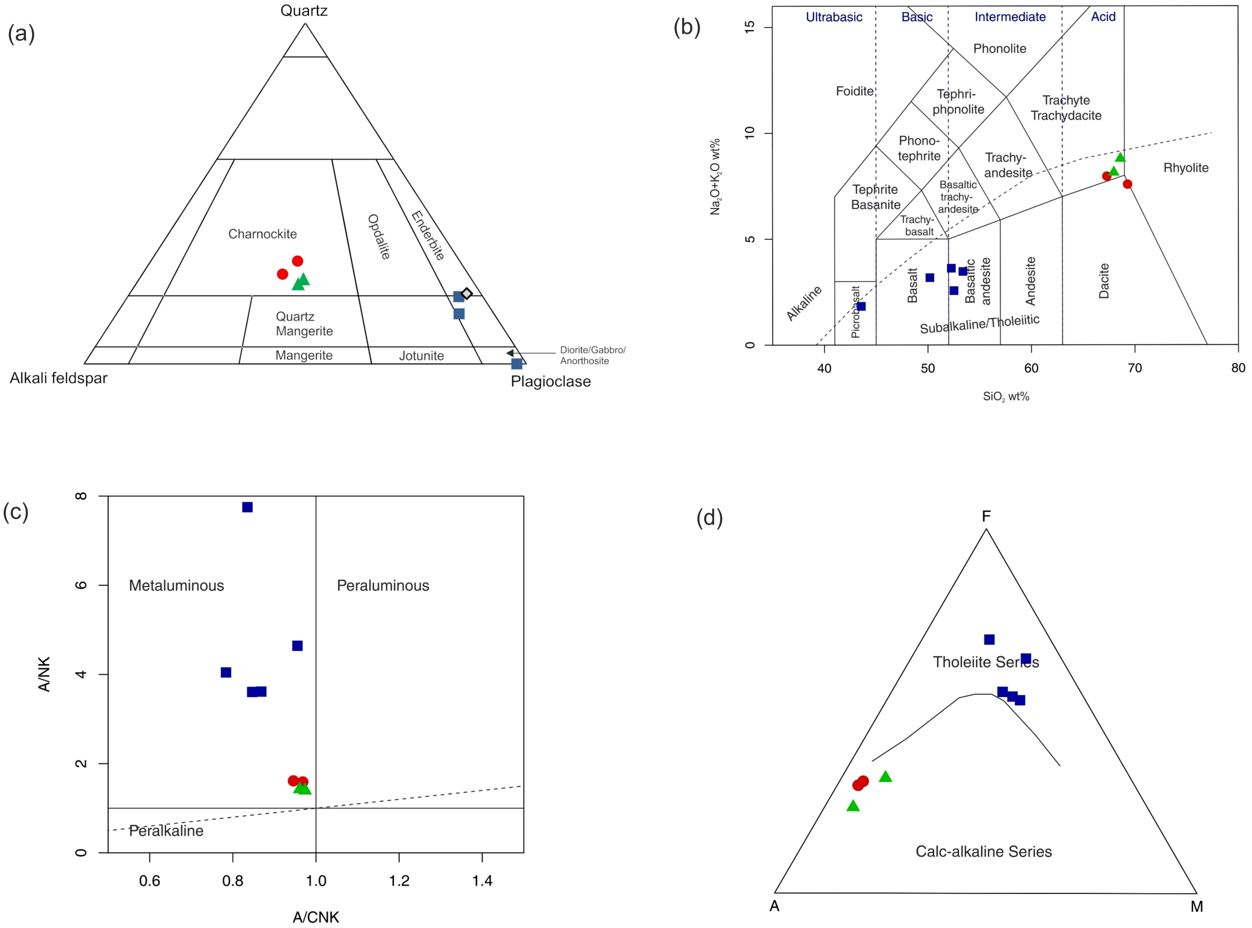
Fig 10: Spider plot comparing the G-Ch rocks with leptynite; the average composition of Post-Archean Indian Shale (Ray and Paul, 2021) is also plotted for reference. See main text for discussion.

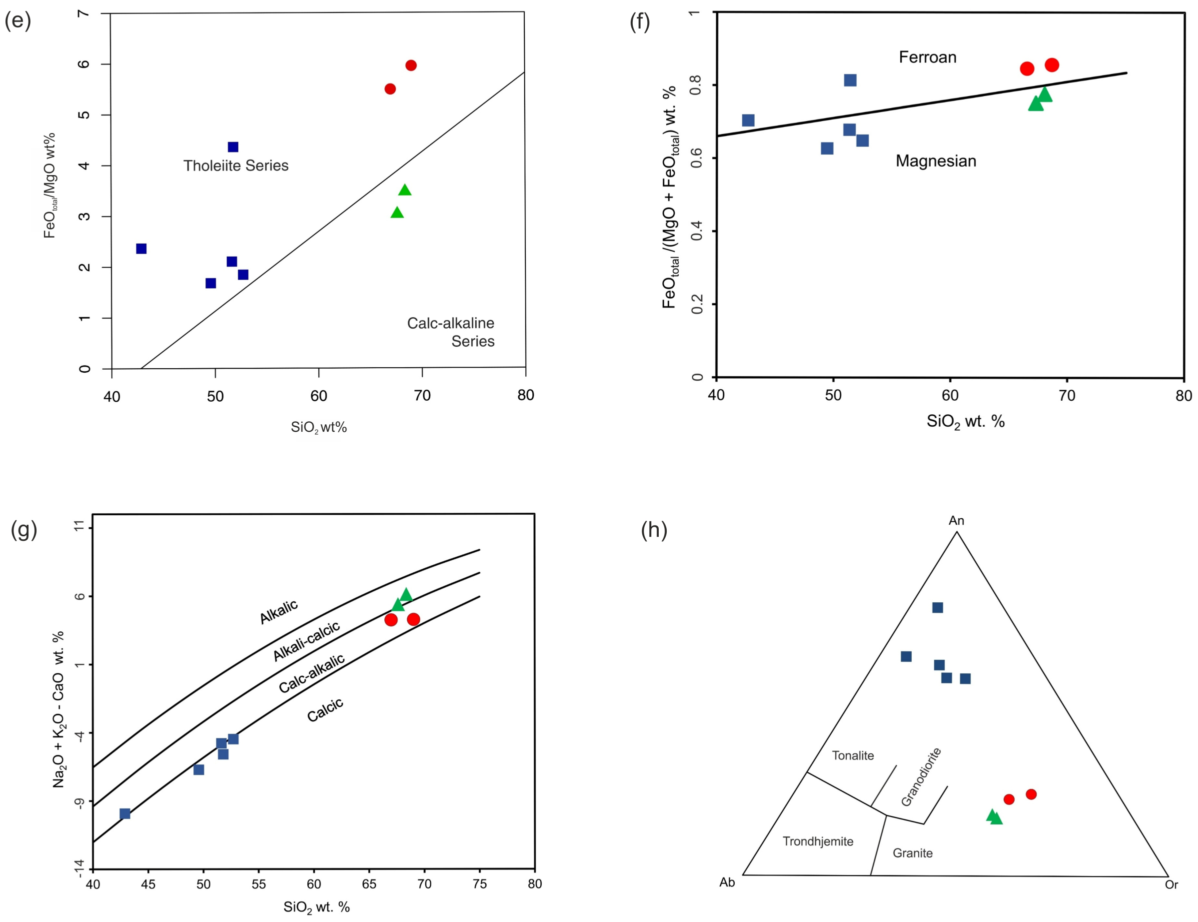
Fig 11: Spider plot comparing the NG-Ch unit with the composition range of Fiskenæsset region granites (Huang et al., 2016) (excluding altered samples and cumulates). See main text for discussion.

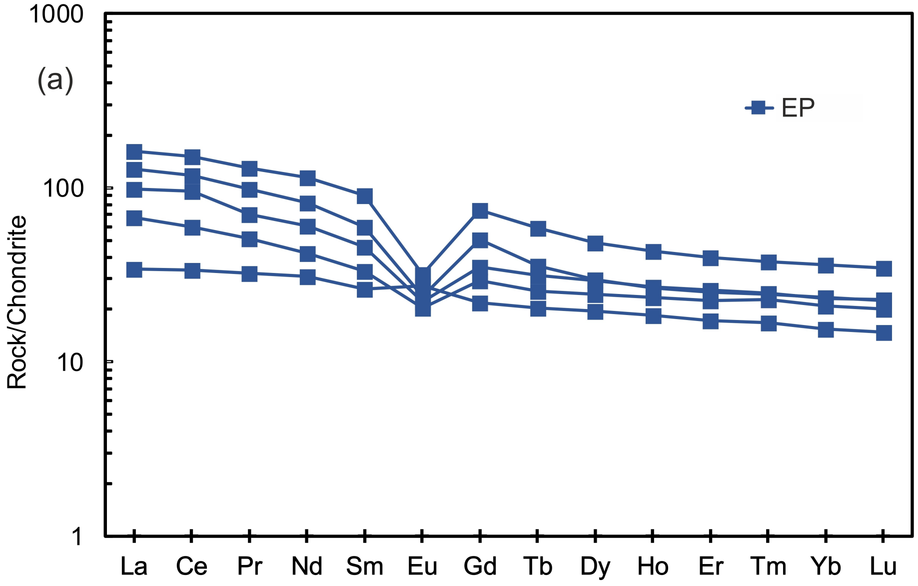

(a)
20.5˚
20.0˚
85.5˚
86.0˚
INDIA
20˚N
16˚N
12˚N
72˚E
76˚E
80˚E
84˚E
88˚E
Mumbai
Chennai
(a)
Mahanadi River
Khordha
Jatni
Bhubaneswar
Study area
Chilka lake
Puri
N
0
10
km
Legend
Charnockite (sensu lato)
Granite gneiss and migmatite
Composite leptynite-charnockite (sensu lato)
Garnet-sillimanite gneiss
Quartzite
Laterite
Sandstone
(b)
20.11˚
20.09˚
85.57˚
85.59˚
NH 16
USIS 01
USIS 02
USIS Chowk Quarry
GRL 03
GRL 02
Greenlake Quarry
HAT 04
HAT 01
Hatia Quarry
CAL 01
CAL 04
CAL 02
Calcutta Quarry
N
0
0.5
km
Legend
EP
NG-Ch
G-Ch
Composite leptynite-metagreywacke-charnockite
Composite metapelite-metagreywack-charnockite

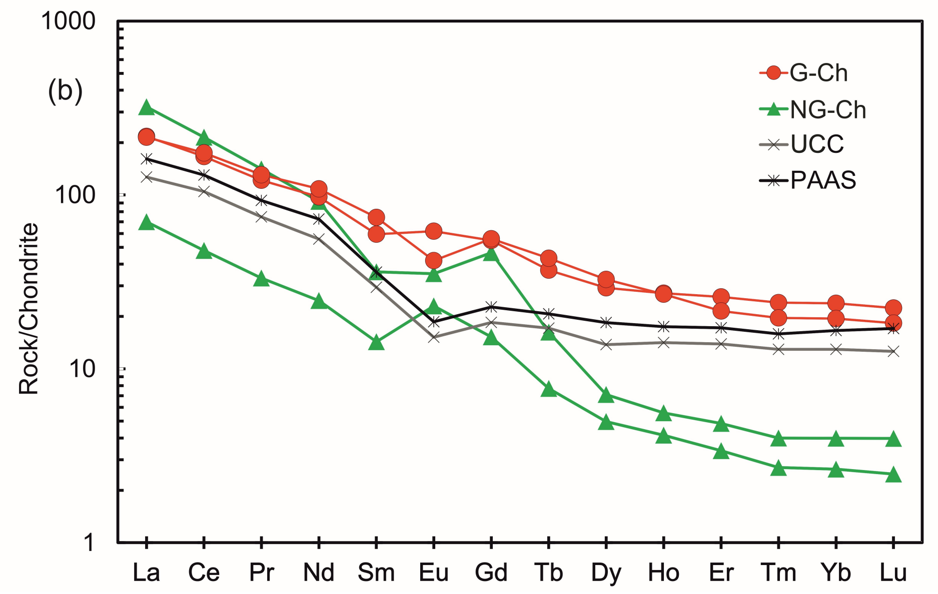

(a)
Metatexite Khondalite
EP
(b)
(c)
EP
Leptynite
ca. 2m
(d)
(e)
(f)


(a)
Ilm
Opx
Pl
Bt
1000 µm
(b)
Opx
Bt
Ilm
500 µm
(c)
Opx
Ilm
500 µm
(d)
Pl
Opx
Ilm
Kfs
Qz
1000 µm
(e)
Opx
Bt
Ilm
1000 µm
(f)
Ilm
Bt
Opx
Grt
500 µm
(g)
Opx
Grt
500 µm
(h)
Kfs
Pl
Bt
500 µm


(a)
Quartz
Charnockite
Opdalite
Enderbite
Quartz Mangerite
Mangerite
Jotunite
Diorite/Gabbro/ Anorthosite
Alkali feldspar
Plagioclase
(b)
Ultrabasic
Basic
Intermediate
Acid
Phonolite
Foidite
Tephri-phonolite
Trachyte Trachydacite
Phono-tephrite
Trachy-andesite
Rhyolite
Tephrite Basanite
Basaltic trachy-andesite
Trachy-basalt
Alkaline
Picrobasalt
Basalt
Basaltic andesite
Andesite
Dacite
Subalkaline/Tholeiitic
Na_2O+K_2O wt%
SiO_2 wt%
0
5
10
15
40
50
60
70
80
(c)
Metaluminous
Peraluminous
Peralkaline
A/NK
A/CNK
0
2
4
6
8
0.6
0.8
1.0
1.2
1.4
(d)
F
Tholeiite Series
Calc-alkaline Series
A
M


(e)
FeO$_{total}$/MgO wt%
0
1
2
3
4
5
6
7
Tholeiite Series
Calc-alkaline Series
40
50
60
70
80
SiO$_2$ wt%
(f)
FeO$_{total}$/(MgO + FeO$_{total}$) wt. %
0
0.2
0.4
0.6
0.8
1
Ferroan
Magnesian
40
50
60
70
80
SiO$_2$ wt. %
(g)
Na$_2$O + K$_2$O - CaO wt. %
-14
-9
-4
1
6
11
Alkalic
Alkali-calcic
Calc-alkalic
Calcic
40
45
50
55
60
65
70
75
80
SiO$_2$ wt. %
(h)
An
Tonalite
Granodiorite
Trondhjemite
Granite
Ab
Or


(a)
EP
Rock/Chondrite
1000
100
10
1
La
Ce
Pr
Nd
Sm
Eu
Gd
Tb
Dy
Ho
Er
Tm
Yb
Lu


(b)
G-Ch
NG-Ch
UCC
PAAS
Rock/Chondrite
1000
100
10
1
La
Ce
Pr
Nd
Sm
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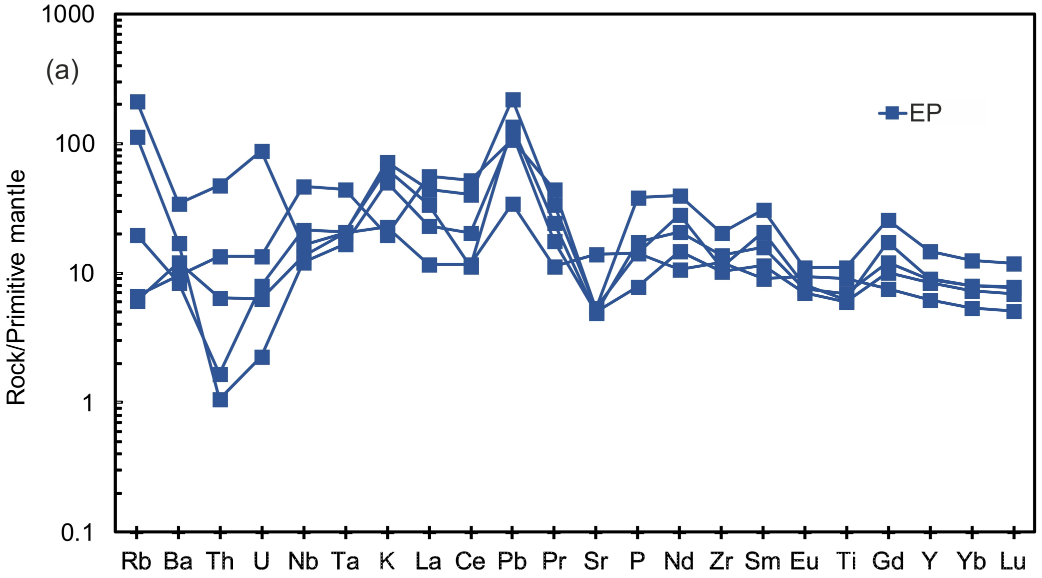
Tb
Dy
Ho
Er
Tm
Yb
Lu


1000
(a)
100
10
1
0.1
Rock/Primitive mantle
EP
Rb
Ba
Th
U
Nb
Ta
K
La
Ce
Pb
Pr
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P
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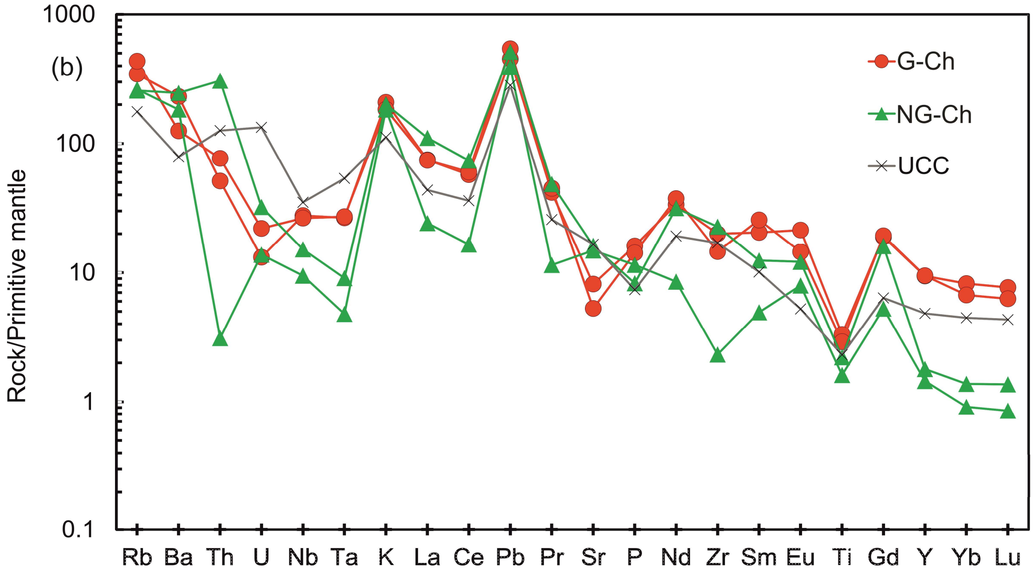
Eu
Ti
Gd
Y
Yb
Lu


1000
(b)
100
10
1
0.1
Rock/Primitive mantle
G-Ch
NG-Ch
UCC
Rb
Ba
Th
U
Nb
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La
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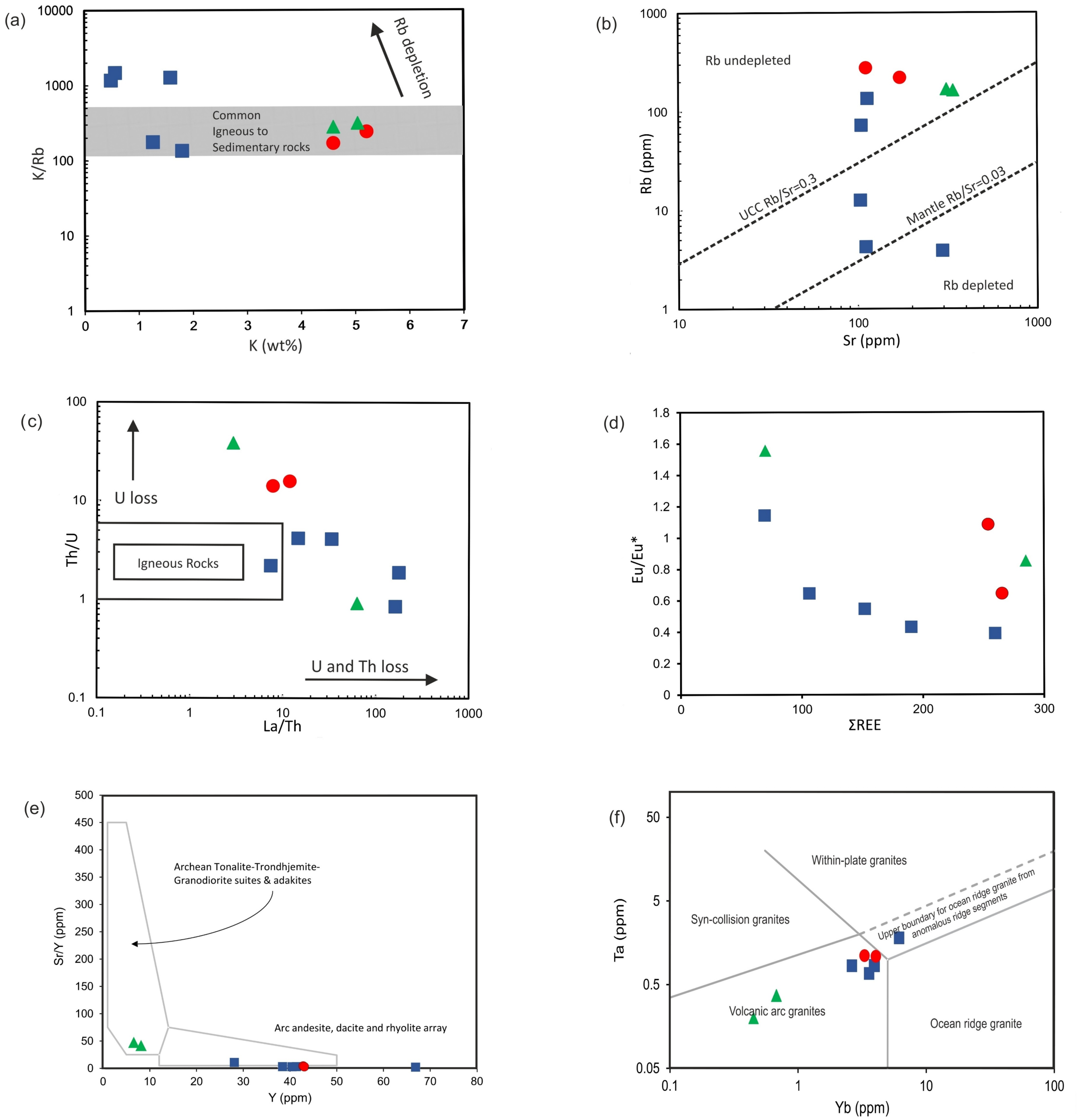
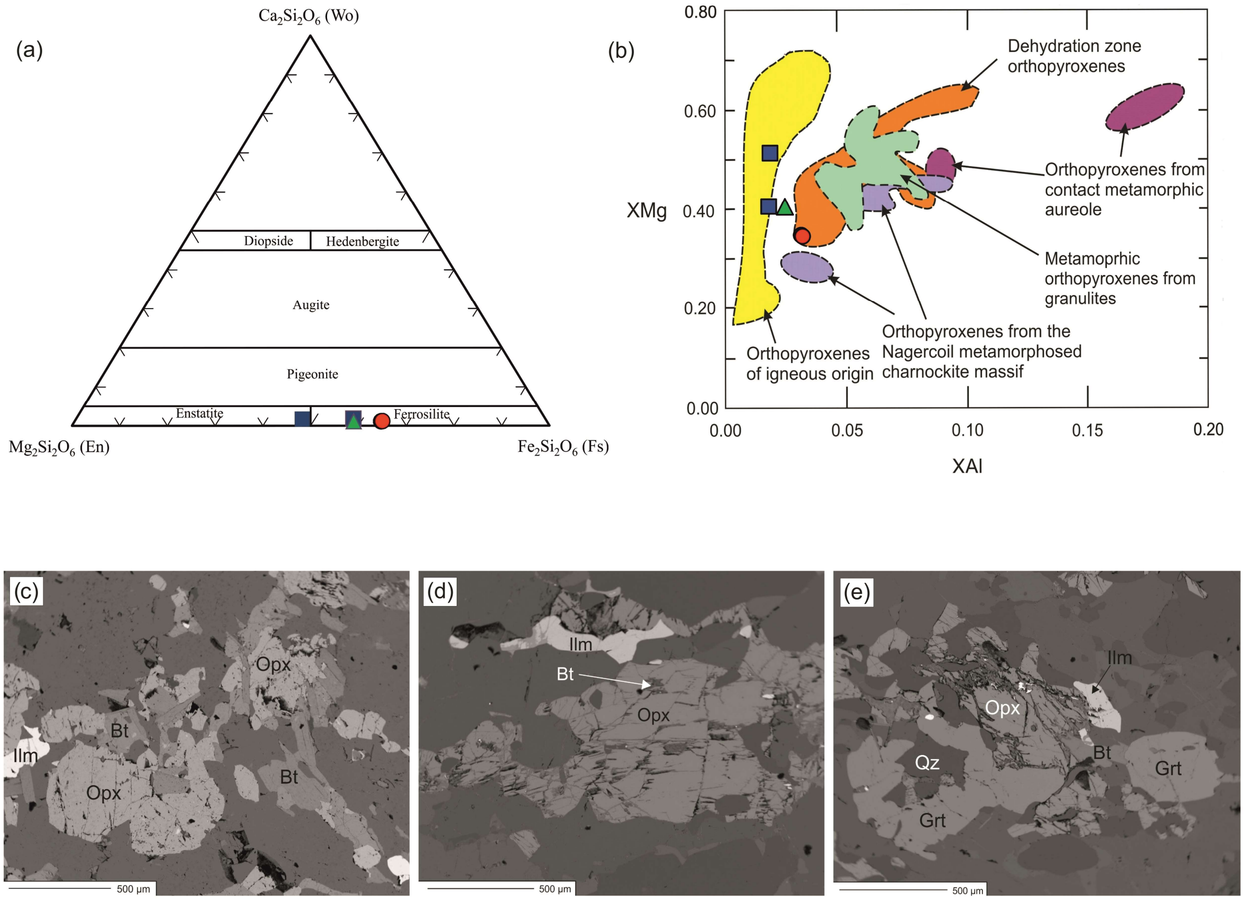
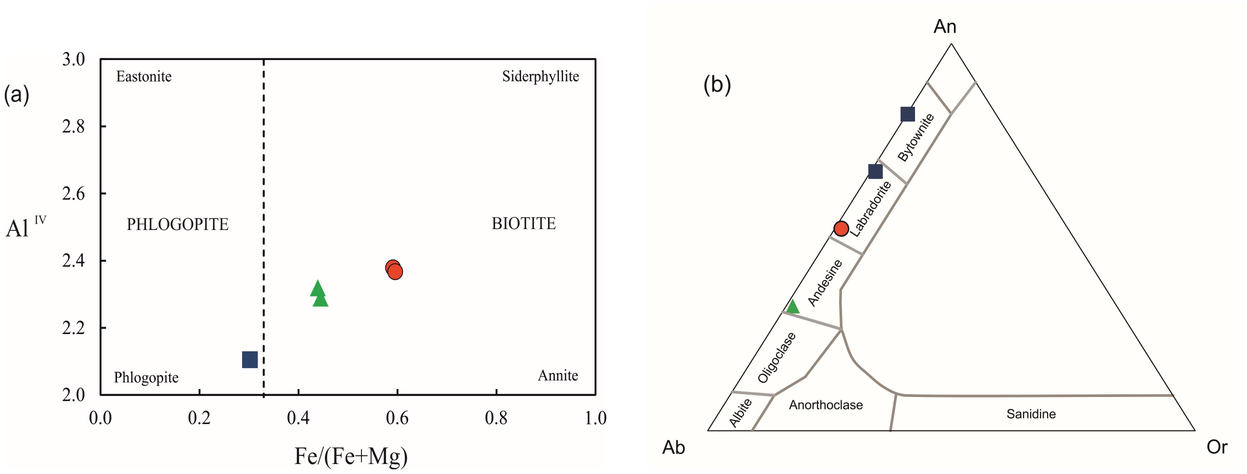
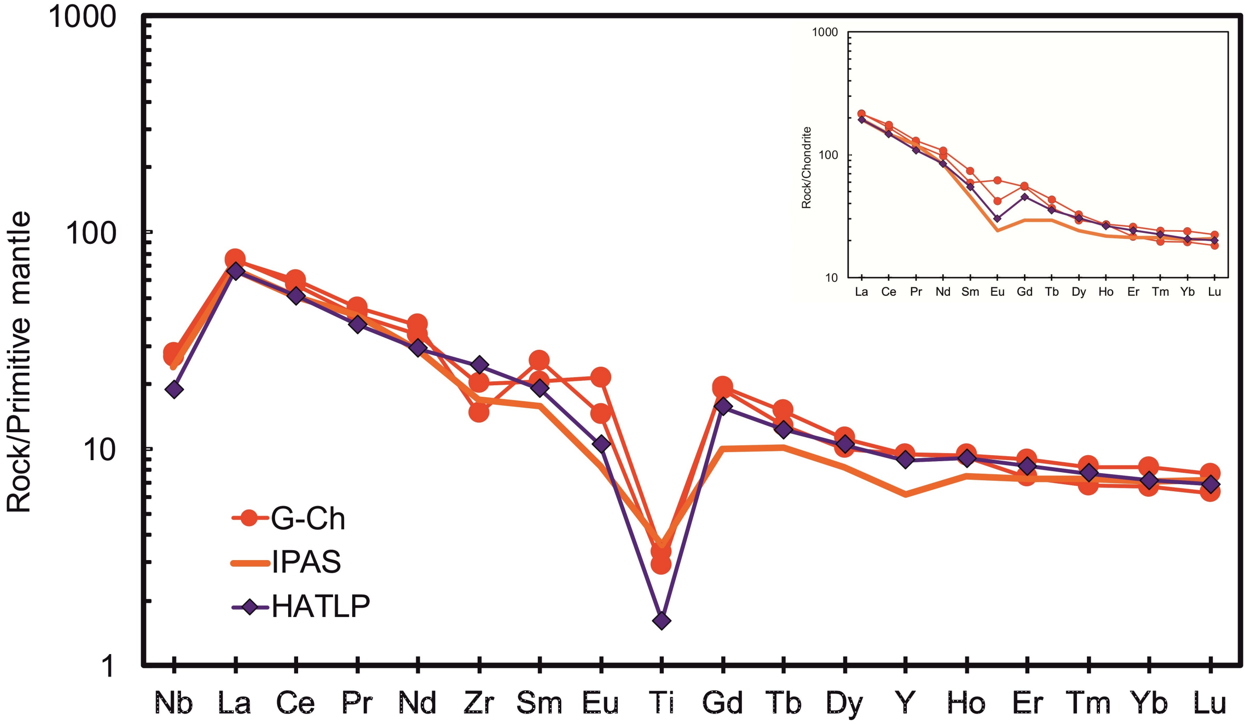
Eu
Ti
Gd
Y
Yb
Lu

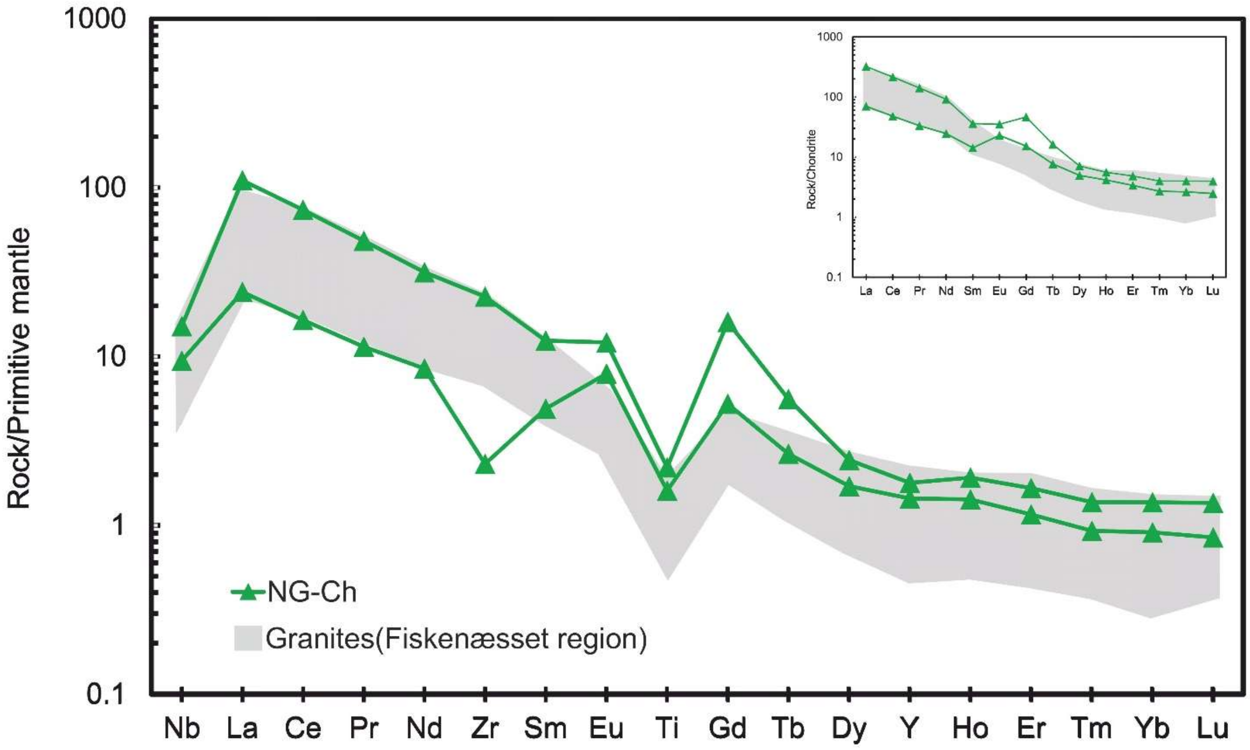

(a)
K/Rb
K (wt%)
Rb depletion
Common
Igneous to
Sedimentary rocks
10000
1000
100
10
1
0
1
2
3
4
5
6
7
(b)
Rb (ppm)
Sr (ppm)
Rb undepleted
UCC Rb/Sr=0.3
Mantle Rb/Sr=0.03
Rb depleted
1000
100
10
1
10
100
1000
(c)
Th/U
La/Th
U loss
Igneous Rocks
U and Th loss
100
10
1
0.1
0.1
1
10
100
1000
(d)
Eu/Eu*
ΣREE
1.8
1.6
1.4
1.2
1
0.8
0.6
0.4
0.2
0
0
100
200
300
(e)
Sr/Y (ppm)
Y (ppm)
Archean Tonalite-Trondhjemite-
Granodiorite suites & adakites
Arc andesite, dacite and rhyolite array
500
450
400
350
300
250
200
150
100
50
0
0
10
20
30
40
50
60
70
80
(f)
Ta (ppm)
Yb (ppm)
Within-plate granites
Syn-collision granites
Upper boundary for ocean ridge granite from
anomalous ridge segments
Volcanic arc granites
Ocean ridge granite
50
5
0.5
0.05
0.1
1
10
100


(a)
Ca2Si2O6 (Wo)
Diopside
Hedenbergite
Augite
Pigeonite
Enstatite
Ferrosilite
Mg2Si2O6 (En)
Fe2Si2O6 (Fs)
(b)
0.80
0.60
XMg 0.40
0.20
0.00
0.00
0.05
0.10
0.15
0.20
XAl
Dehydration zone orthopyroxenes
Orthopyroxenes from contact metamorphic aureole
Metamoprhic orthopyroxenes from granulites
Orthopyroxenes from the Nagercoil metamorphosed charnockite massif
Orthopyroxenes of igneous origin
(c)
Opx
Bt
Ilm
Opx
Bt
500 µm
(d)
Ilm
Bt
Opx
500 µm
(e)
Ilm
Opx
Qz
Bt
Grt
Grt
500 µm


(a)
Al IV
3.0
2.8
2.6
2.4
2.2
2.0
0.0
0.2
0.4
0.6
0.8
1.0
Fe/(Fe+Mg)
Eastonite
Siderphyllite
PHLOGOPITE
BIOTITE
Phlogopite
Annite
(b)
An
Ab
Or
Bytownite
Labradorite
Andesine
Oligoclase
Albite
Anorthoclase
Sanidine



1000
100
10
1
Rock/Primitive mantle
Nb
La
Ce
Pr
Nd
Zr
Sm
Eu
Ti
Gd
Tb
Dy
Y
Ho
Er
Tm
Yb
Lu
G-Ch
IPAS
HATLP
1000
100
10
Rock/Chondrite
La
Ce
Pr
Nd
Sm
Eu
Gd
Tb
Dy
Ho
Er
Tm
Yb
Lu



1000
100
10
1
0.1
Rock/Primitive mantle
Nb
La
Ce
Pr
Nd
Zr
Sm
Eu
Ti
Gd
Tb
Dy
Y
Ho
Er
Tm
Yb
Lu
NG-Ch
Granites(Fiskenæsset region)
1000
100
10
1
0.1
Rock/Chondrite
La
Ce
Pr
Nd
Sm
Eu
Gd
Tb
Dy
Ho
Er
Tm
Yb
Lu