1   Best practices for the analyses of CO$_2$ fluids by Raman Spectroscopy

2   Penny E. Wieser[1*], Charlotte L DeVitre[1,2], Isabelle Susman[1]

3   [1] Earth and Planetary Sciences, University of California, Berkeley, CA 94270, USA

4   [2] Now at Department of Earth and Environmental Sciences, University of Ottawa, Ottawa, ON
5   K1N 6N5, Canada

6   *penny_wieser@berkeley.edu

7

11

12

13

14

15

16

17

18

19

20

21

22

23

24

25

26

27

Abstract

Raman spectroscopy is a key method for determining CO_2 densities in geological fluids, yet acquisition, calibration, and processing methodologies vary widely between laboratories. This study evaluates how these parameters affect precision and accuracy. We show that spectral non-linearity can cause a single instrument to show variable relationships between CO_2 density and spectral parameters as the spectral center and drift correction routine is varied, explaining some of the diversity in published densimeters. We show that when the calibration methodology is applied to unknowns, the Raman method is robust. For example, 240 repeat measurements of a natural fluid inclusion (FI) over ~2 years gives a 1σ variation of only ~0.008 g/cm³. We also assess the influence of variations in sample and room temperature on measurements to help inform analytical strategies. We show that high laser powers in fluid and melt inclusions causes heating of the CO_2 fluid, and unacceptable drops in CO_2 density. We offer practical strategies to optimize signal without raising laser power, noting that small changes in focus and X–Y position greatly affect intensity (but not depth in sample).

1. Introduction

Traditionally, the density of CO_2-rich fluids in geological samples have been measured using Microthermometry, a technique in which phase changes are observed during heating and cooling of trapped inclusions. Starting in the 1930s, it was recognised that CO_2 fluid/gas/vapour has a distinct signal when analysed by Raman spectroscopy, with two strong peaks at ~1285 and ~1388 cm^{-1} (Fermi, 1931; McLennan F.R.S. and Smith M.A., 1932). These peaks are collectively called the Fermi diad. In the 1970s, the distance between the two peaks – referred to as the Fermi diad separation, splitting, peak splitting or diad splitting – was shown to be related to the density of CO_2 (Wright and Wang, 1973). These findings opened up the possibility of using Raman spectroscopy to quantify the density of CO_2 fluids, an important parameter in many geological research questions.

The relationship between diad splitting and CO_2 density on a given Raman instrument is described using a densimeter, which is typically modelled as a polynomial function of Nth degree. A number of groups have developed experimental apparatus to determine densimeters for different instruments (DeVitre et al., 2021; Fall et al., 2011; Kawakami et al., 2003; Lamadrid et al., 2017;

57 Rosso and Bodnar, 1995; Song et al., 2009; Wang et al., 2019, 2011; Yamamoto and Kagi, 2006).

58 Not only do densimeters vary widely between Raman instruments (e.g. Lamadrid et al. 2017,

59 DeVitre et al. 2021); changes in hardware on a single instrument have also been shown to affect

60 the densimeter. For example, Lamadrid et al. (2017) show that the densimeter on a single

61 instrument can be shifted by changing the laser wavelength (e.g. a ~0.02 g/cm^3 shift from 514 nm

62 to 632 nm at 2400 g) or the instrument grating (e.g., a 0.02-0.04 g/cm^3 shift from 600-2400 g at

63 632 nm). They conclude that the same wavelength and grating must be used during analysis and

64 calibration.

65 In contrast, Remigi et al. (2021) investigated offsets between densimeters by generating a

66 calibration using natural fluid inclusion (FI) standards with densities between 0.37-1.08 g/cm^3.

67 They compare their densimeter to published densimeters using bootstrapping techniques to better

68 account for uncertainty in densimeter regressions. They conclude that published Raman

69 densimeters fall into two main categories – those collected on instruments with higher spectral

70 resolution (~0.4 cm^{-1}) and those collected on instruments with lower spectral resolution (~1-1.5

71 cm^{-1}). For CO_2 densities > 0.3-0.4 g/cm^3, they show that published densimeters within the high-

72 or low-resolution category are within error of each other. Thus, they suggest that published

73 densimeters have wider applicability than previously thought, and instrument-specific calibrations

74 may not be needed.

75 However, it is worth noting that all the high resolution densimeters they consider were performed

76 on a HORIBA HR instruments, so do not encompass the true variability of Raman hardware.

77 Additionally, they do not consider offsets between densimeters at lower densities (<0.3 g/cm^3).

78 When considering densimeter offsets as a percentage of the CO_2 density, published densimeters

79 differ the most for analyses of gaseous CO_2 at room temperature (ρ~ 0-0.2 g/cm^3; Lamadrid et al.

80 2017). At these densities offsets can be a factor of 2 or more. This has major implications in

81 volcanology, because many melt inclusion vapor bubbles have CO_2 densities below 0.2 g/cm^3

82 (see Fig. 21 of Wieser et al., 2025). Accurate measurements of vapor bubble CO_2 contents are

83 essential for reconstructing the CO_2 contents of magmas, with implications for determining

84 magma storage depths and the cause of explosive eruptions (Allison et al., 2021; DeVitre et al.,

85 2023b; Moore et al., 2015). Despite this, many recent studies are still measuring CO_2 fluids with

86 low densities using calibrations published in other laboratories (e.g. Howe et al., 2025).

87 Another recent discussion of the utility of Raman spectroscopy for quantifying CO_2 densities was

88 provided by Bakker (2021). This work assessed the uncertainties associated with various parts of

89 the Raman method (e.g., peak fitting, instrument drift etc) and suggested that errors in the position

90 of each diad propagate through to very large errors in diad splitting (±0.26-0.27 cm^{-1}) and thus

91 CO_2 density (~±0.1 g/cm^3). This is in contrast to the very low theoretical errors suggested by

92 (Yamamoto and Hagiwara, 2024). Bakker (2021) also suggests that a common method of

93 correcting for instrument drift using the distance between two Ne lines (e.g. Lamadrid et al., 2017)

94 does not fully capture the complex nonlinearity of Raman spectrometers.

95 Overall, the Bakker (2021) and Remigi et al. (2021) papers demonstrate that while a lot of

96 progress has been made in recent years on optimizing Raman analyses of CO_2 fluids (DeVitre et

97 al., 2023a, 2021; Fall et al., 2011; Lamadrid et al., 2017; Le et al., 2021; Wang et al., 2011), there

98 are still some fundamental questions regarding the factors affecting the densimeter relationship

99 for a single instrument, and the accuracy and precision of the Raman method when applied to

100 natural samples. In this study, we investigate the effect of acquisition and data processing

101 strategies on the precision and accuracy of CO_2 fluid measurements, with the goal of establishing

102 coherent best practices for the broader community.

103 ## 2. Methods

104 We acquire Raman spectra on natural FI and melt inclusions (MI) hosted in olivine using a WITec

105 Alpha 300R Raman spectrometer in the Department of Earth and Planetary Science, University

106 of California, Berkeley (hereafter, UCB WITec). We use a green solid-state laser with a wavelength

107 of 532.046 nm, and a Ne lamp in the beam path to correct for instrument drift. For measurements

108 of the CO_2 Fermi diad, the 1800 grooves/mm grating was used, yielding a spectral resolution of

109 0.57 cm^{-1}. The laser was focused on the sample with a Zeiss 50X objective (x0.55NA, 9.1 mm

110 working distance) or a Zeiss 100X (0.95 Na, 4 mm working distance). This Raman system is

111 equipped with a TruPower module, allowing control of the laser power with an accuracy of ±0.1

112 mW.

113 Various methods were used to heat samples above the critical temperature of CO_2 to ensure a

114 single homogenous fluid was being analysed (DeVitre et al., 2023a). For most analyses, samples

115 were placed on a glass slide in crystalbond and placed on a Peltier heating stage with a central

116 hole for transmitted light illumination (Supporting Fig. 1). The target sample temperature was

117 37 °C, matching the gas cell calibration temperature for this instrument (DeVitre and Wieser,

118 2024). Temperature was monitored using a type K thermocouple taped to the slide with electrical

119 tape. When tighter temperature control was needed, crystals were double polished and placed in

120 a Linkam THMSG600 stage. To achieve the long working distance required to see through the

121 window, we initially made measurements using a LWD 40X objective with a cover slip correction

122 collar. After doing tests with and without the window, for the wide range of sample temperatures

123   investigated here (37-70℃), there was no difference in heating series. Thus, later analyses
124   performed them without the window using the Zeiss 50X objective to help increase the strength
125   of signals. We covered the sample with a silver lid with a top hole to enhance uniform heating. All
126   data processing was performed in the open-source Python3 tool DiadFit (Wieser and DeVitre,
127   2024).

128   ## 3. Results and Discussion

129   ### 3.1 Neon lines and spectral nonlinearity

130   Instrument drift during a Raman session is commonly assessed using repeated analyses of the
131   spectral emission of Neon generated by a Ne lamp (e.g. DeVitre et al., 2021; Lamadrid et al.,
132   2017. Fig. 1a). Two emission lines bracketing the diad are selected. The theoretical distance
133   between these lines is calculated from the published wavelength of the emission spectra of Neon
134   in air from the NIST webbook, converted to wavenumber using the wavelength of the laser used
135   for analysis. This theoretical distance is then compared to the measured distance in the acquired
136   spectra. For example, if the theoretical distance between Ne lines is 330.4776 cm^{-1}, and the
137   measured distance is 331.0587 cm^{-1}, the correction factor is $\frac{\text{Theoretical distance (330.4776)}}{\text{measured distance (331.0587)}} = 0.9982$.
138   If the diad splitting was measured at 103 cm^{-1}, the corrected diad splitting would be 0.9982*103 =
139   102.8192 cm^{-1}.

140   Bakker (2021) criticized this method (which they refer to as the 'line segment' technique) because
141   it assumes that the correction factor at any given point between two lines is the same, which is
142   not true if there is any nonlinearity in the spectrometer. Nonlinearity causes different line pairs to
143   yield different correction factors (as noted by Wieser et al. 2021). Bakker (2021) suggest that a
144   'bracketing technique' should be used instead, where a polynomial model is developed to describe
145   how the position of each Ne line deviates from the theoretical position as a as a function of
146   wavenumber. For brevity, we refer to the lower wavenumber peak of CO_2 as diad 1 (also v- or
147   2v2), the higher wavenumber peak as diad 2 (also v+ or v_1+2v_2) and the Fermi diad separation
148   as diad splitting. This polynomial allows a correction factor to be calculated for diad 1 and diad 2.
149   These corrected peak positions are used to calculate the splitting, rather than applying a
150   correction factor to the splitting as in the line-segment method.

151

152   We investigate spectrometer nonlinearity using three different WITec instruments, and determine
153   the effect of this nonlinearity on instrument drift corrections and calculated CO_2 densities. First,
154   we perform 50 consecutive repeated acquisitions of the emission spectrum of Neon. We fit

155 pseudovoigt curves to the 6 major Ne lines within the spectral window for each acquisition using

156 DiadFit (565.66588 nm ≈1117 cm^{-1}, 568.98163 nm ≈1220 cm^{-1}, 571.92248 nm≈1310 cm^{-1},

157 574.82985 nm≈1399 cm^{-1} 576.44188 nm≈1447 cm^{-1}, and 580.44496 nm≈1567 cm^{-1} for the

158 532.046 nm laser wavelength used here, Fig. 1a, Kramida and Ralchenko, 1999). For the line at

159 1117 cm^{-1}, we iteratively fit two pseudovoigt functions to resolve the smaller peak on the shoulder.

160 Fig. 1b shows the fitted Ne line position divided by the theoretical position calculated from the

161 laser wavelength and the theoretical positions in air. This comparison demonstrates that the

162 spectrometer is indeed non-linear, with the correction factor varying across the acquisition window

163 (Fig. 1b).

164 The polynomial fit through these correction factors is shown for the 'bracketing technique' (grey

165 curve, Fig. 1c) for a single Ne acquisition. Correcting for the position of the CO_2 diad using this

166 curve this would yield a corrected splitting for the diad position shown by blue dotted lines that

167 would be equivalent to a line segment correction factor of 0.9977. The 'line segment' technique

168 for drift correction can be visualized by drawing a straight line between any two datapoints (circles)

169 on this polynomial curve (Fig. 1c). If the gradient of this linear segment is the same as the gradient

170 of a line drawn between where the x positions of the diad intercept the polynomial, the line

171 segment technique and bracketing techniques would give the same result. The line segment

172 method applied to the 1220-1399 cm^{-1} lines results in a very similar correction factor to the

173 bracketing technique (0.9976 vs. 0.9977). The disadvantage of choosing this pair is that the 1220

174 cm^{-1} line is relatively small, resulting in larger fitting errors, meaning that correction factors show

175 substantially more variability in a given run than using the 1117-1447 cm^{-1} pair (see error bar on

176 Fig. 1c). The 1117-1447 cm^{-1} pair used by our laboratory yields a reasonably similar correction

177 factor to the bracketing technique (0.9982 vs. 0.9977). However, applying the line segment

178 technique to the 1220-1567 cm^{-1} lines (red dashed segment) would result in a Ne correction factor

179 that is far smaller than that determined using the bracketing technique (0.9991 vs. 0.9977),

180 yielding a higher apparent diad splitting ($\Delta_{CO2}^{Corr,brack}$=103.450 cm^{-1} vs. $\Delta_{CO2}^{Corr,seg}$=103.648 cm^{-1}).

181 The differences in correction factor mean that the choice of Ne line pair affect the corrected

182 splitting, and thus the relationship between CO_2 density and corrected splitting (i.e. the

183 densimeter). To demonstrate this, we take the CO_2 and Neon spectra from the analytical sessions

184 used to calibrate the UCB densimeter. The published UC densimeter was calibrated from

185 corrected splittings calculated using correction factors from the 1117-1447 cm^{-1} Ne line pair

186 (DeVitre et al., 2021; DeVitre and Wieser, 2024, black dotted line, Fig. 2a-b). We reprocess the

187 data using 4 new pairs of Ne lines: ~1220-1447 cm^{-1} (dashed red line, Fig. 2a, c), ~1117-1399 cm^{-}

188   ^{1} (dashed blue line), 1220-1398 cm^{-1} (solid cyan line), and 1220-1567 cm^{-1} (solid magenta line,

189   Fig. 2a-b). Thus, for each CO$_2$ spectrum, we have one value for the known density of the fluid

190   from the calibration cell, and 5 different corrected splitting values – one for each Ne line pair.

191   Using three polynomial fits, one for low, medium and high densities (following DeVitre et al. 2021),

192   we generate 5 different densimeter equations (Fig. 2a-b). Despite being collected on a single

193   instrument using the same calibration dataset, the densimeters differ by up to ~0.08 g/cm^3 simply

194   based on Ne line pair used for data reduction. To put the magnitude of these offsets into context,

195   we underlay the densimeters collected on a wide variety of Raman instruments (Fall et al., 2011;

196   Kawakami et al., 2003; Lamadrid et al., 2017; Wang et al., 2011, 2011; Yamamoto and Kagi,

197   2006), colored using the high (grey) and low (salmon) resolution classifications from Remigi et al.

198   (2021). The variability from a single instrument introduced through the choice of a Neon line pair

199   is as large as the differences between Raman instruments. It is also noteworthy that despite our

200   instrument having a 'high resolution' spectrometer, all our possible densimeters lie with the low

201   resolution densimeters of Remigi et al. (2021). Table 1 shows that different laboratories use a

202   wide variety of methods to calibrate their spectrometers and subsequently correct them for

203   instrument drift. Thus, we suggest that the choice of drift correction method is likely a major

204   contributor to the wide diversity of densimeters published by different laboratories (see also

205   Bakker, 2021).

206   Table 1- Compilation of studies publishing densimeters, and the methods used to calibrate the

207   spectrometer (which sets the nonlinearity response seen in Fig. 1b) and correct data for drift

208   during a run. n.d. indicates the publication did not say what method was used.

Paper	Spectrometer Calibration method	Drift correction used
DeVitre et al., 2021; DeVitre and Wieser, 2024	Hg-Ar lamp (atomic emission line)	Ne lamp (lines at 1117 cm^{-1} and 1447 cm^{-1} at 532.046 nm)
Wieser et al., (2021)	Si chip	Repeated FI measurement
Lamadrid et al. (2017)	Laser line and Ne line at 1458.58 cm^{-1}	Ne lamp (lines at 1031.42 cm^{-1} and 1458.58 cm^{-1} for a 514 nm laser)
Fall et al. (2011)	Silicon wafer	n.d. say methods similar to Lin et al. (2007) who use Ne lines

Wang et al. (2019, 2011)	Benzonitrile - (Raman shift standard)	Diamond
Yamamoto and Kagi, (2006)	Napthalene (Raman shift standard)	n.d.
Song et al. (2009)	Diamond	n.d.

209

210 These offsets between these densimeters are only problematic if they yield different results when
211 analysing unknowns. We test whether this is the case using data collected on FI from Fogo
212 volcano collected by DeVitre and Wieser (2024), where there are independent constraints on CO_2
213 density from microthermometry. We process the FI data using each of the 5 different Ne line pairs
214 (Fig. 2b), generating 5 values of corrected splitting for each spectrum, and input these values into
215 the densimeter corresponding to each Ne line pair (e.g. data reduced using correction factors for
216 the 1220-1399 pair was fed into the 1220-1399 densimeter). To visualize the differences between
217 these data reduction strategies, we subtract the density calculated using each line pair from that
218 calculated using the 1117-1447 pair, which was shown to correlate very well with
219 Microthermometry (DeVitre and Wieser, 2024). The error bar on each datapoint shows the
220 combination of the peak fitting error, drift correction error, and uncertainty on the densimeter
221 (errors propagated using DiadFit, (Wieser and DeVitre, 2023). Densities for the 4 other line
222 combinations are within error of the data reduced using the 1117-1447 cm^{-1} line pair. We also
223 overlay lines showing the 1σ uncertainty associated with Microthermometry assuming a 0.1℃ and
224 0.2℃ error in sample temperature (Bakker, 2021). This analysis demonstrates that the
225 uncertainties in the choice of data reduction strategies are small, and comparable to the errors
226 associated with the stage temperature when performing Microthermometry.

227 Repeated acquisitions of Neon lines at UCB in April 22 and October 2024 shows that the shape
228 of the nonlinearity curve shifts up and down with time, but the shape is maintained (Fig. 4a, solid
229 lines). We also tested recalibrating the UCB spectrometer in April 2024 using a Hg-Ar lamp
230 (WITec's default calibration routine). The nonlinearity curve is shifted down (dotted vs. solid lines,
231 Fig. 4a), but the shape is preserved. Thus, while spectrometer nonlinearity is clearly significant,
232 as long as calibration and natural data is processed in the same way, the nonlinearity introduces
233 a consistent offset through time, and choice of Ne line is not a major source of uncertainty. This
234 stability also means that the bracketing technique is not necessary, given the additional
235 computation steps, and the relatively large uncertainty associated with fitting a polynomial through

236   some weaker Ne lines (e.g. 1310 cm^{-1}). However, if a laboratory wishes to change drift correction

237   strategy, they would need to reprocess their calibration data to use the same strategy.

3.1.1 The effect of changing the Spectral Center

239   The significant nonlinearity in the spectrometer response shown in Fig. 1b has other implications

240   for the densimeter equation determined for a given instrument. When setting up a Raman

241   analysis, the user must choose the spectral center—the wavenumber at the center of the

242   acquisition window. Along with the grating, this center position determines the range of

243   wavenumbers collected in a single spectral acquisition. Shifting the center means that the photons

244   at a specific wavenumber are collected on a different pixel on the charge-coupled device (CCD)

245   within the spectrometer.

246   Given the strong spectral nonlinearity shown in Fig. 1b, it is worth investigating whether shifting

247   the spectral center affects the measured splitting, because each pixel of the diad would shift to a

248   different part of the polynomial in Fig 1b (see also Bakker 2021). We perform repeated

249   acquisitions on an olivine-hosted CO_2-rich FI from Fogo Volcano with a bulk density of ~0.62

250   g/cm^3 using different spectral centers. We started at the spectral center used for calibration of the

251   Berkeley densimeter (1325 cm^{-1}), then decreased the spectral center

252   (1325→1320→1275→1250→1225 cm^{-1}, magenta colors, Fig. 3a). To help resolve changes from

253   the spectral center vs instrument drift, we increased (green colors), decreased (yellow colors) and

254   then increased (cyan) the spectral center throughout the day (Fig. 3a). We acquire a Neon line at

255   the same spectral center as each FI measurement to correct each spectra (using the 1117-1447

256   cm^{-1} line pair). For each individual acquisition, we calculate the corrected splitting using the Ne

257   line acquired at the same spectral center. Lastly, we calculate density using the UCB densimeter

258   developed using a spectral center of 1325 cm^{-1}.

259   The uncorrected splitting of the diad increases by ~0.22 cm^{-1} as the spectral center is increased

260   from 1225 to 1350 cm^{-1} (~0.0017 cm^{-1} per cm^{-1} of center shift, Fig. 3b). The change in splitting for

261   spectral center shifts of only ~ 25 cm^{-1} apart is far larger than the variation at a given spectral

262   center throughout the day (e.g. instrument drift). After correcting each acquisition using the Neon

263   line collected at the same spectral center, the change in splitting is smaller, but still notable (~0.15

264   cm^{-1} across the window, 0.0012 cm^{-1} per cm^{-1} of shift, Fig. 3c). Calculated CO_2 density changes

265   by 0.08 g/cm^3 across the window (Fig. 3c), which is far outside the ~0.01-0.02 g/cm^3 uncertainty

266   of the Raman method at the densities of interest (Dayton et al., 2023; DeVitre and Wieser, 2024).

267   Interestingly, this 0.08 g/cm^3 offset is also comparable to the offset between all the 'high resolution'

268   densimeters identified by Remigi et al. (2021; Fig. 3c). Overall, this test indicates that unknowns

269   must be analysed at the same spectral center at which the calibration was performed, else
270   calculated densities will be incorrect.

271   To further illustrate the influence of spectral center, we compare the densimeter developed
272   on the UCB WITec alpha300R with densimeters developed in two laboratories at Cornell
273   University (Cornell Center for Materials Research CCMR and Cornell Mass Spectrometry Facility
274   CMaS, DeVitre et al., 2021). These two laboratories were calibrated with a gas cell that is very
275   similar to that at UCB (following the design of DeVitre et al., 2021), and the Raman hardware for
276   all three instruments is very similar: ~532 nm laser wavelength, 1800 grating, Ultra-High-
277   Throughput-Spectrometer UHTS600 for VIS. All data was reduced with the 565.66588 -
278   576.44188 nm Neon line pair (corresponding to 1117-1447 for the UCB 532.046 nm laser
279   wavelength). Comparing the fitted neon line positions to the theoretical positions (Fig. 4a) shows
280   that each spectrometer has its own unique signature of nonlinearity.

281   Despite CMaS and CCMR having more similar nonlinearity responses, the CCMR
282   densimeter shows the biggest offset in densimeter equation relative to UCB and CMaS. We
283   suggest that this is because the UCB densimeter and the CMaS densimeter were both calibrated
284   using a spectral center of 1325 cm^{-1}, while the CCMR densimeter was calibrated with a spectral
285   center of 1340 cm^{-1}. The absolute difference between these densimeters reaches ~0.04 g/cm^3
286   (Fig. 4c). Notably, these densimeter differences cannot be explained by differences in resolution
287   as advocated by Remigi et al. (2021), because the resolution of all three of these WITec
288   instruments is identical.

289   We can also track the stability of our densimeter (and thus spectral non-linearity) with time
290   using repeated acquisitions of FIs from Fogo volcano (FG04-A1-4 and FI FG04-101-A) collected
291   during different analytical sessions. The standard deviation calculated from the 240 repeats on
292   A1-4 collected by PIs, postdocs, graduate students and undergrads spanning almost two years
293   is 0.008 g/cm^3 (Fig. 5a), and is very similar for the N=82 acquisitions of FG04-101-A (0.0078
294   g/cm^3, Supporting Fig. 2). For both inclusions, microthermometry measurements are within, or
295   extremely close to the Raman analyses considering the 1σ values. Overall, this small deviation
296   and lack of coherent trends with time indicates that at least on the WITec, changes in spectrometer
297   nonlinearity does not cause any notable drift in the densimeter, and densimeters are remarkably
298   stable once calibrated.

299   While the Neon method is clearly robust, even with variations in non-linearity, it is
300   worthwhile to determine whether it actually helps correct for drift between and within runs.

301 Yamamoto and Hagiwara, (2024) show that over short-timescales (minutes) where drift is not
302 controlled by temperature, the periodicity seen in CO_2 peak positions is not seen in Ne line
303 acquisitions. We reprocess the repeated FI acquisitions using the median Ne correction factor for
304 all the standard analyses after the instrument recalibration, rather than the correction factor
305 calculated by regressing Ne correction factors as a function of time in each analytical run. The
306 standard deviation of both FIs is extremely similar with and without an analysis-specific Ne
307 correction. Interestingly, for FG04 one outlier is generated using a Ne line correction specific to
308 that analysis that goes away using the average correction factor (Fig. 5a), while the opposite is
309 true for standard FG19_101 (Supporting Fig. 2), one outlier is generated using the median Ne
310 correction. It is noteworthy that after the instrument recalibration in April 24, the Raman was
311 located in a laboratory with an excellent HVAC system, yielding temperature variations of <±1℃
312 per day, which could explain the small difference between these two correction methods.

313 To test the efficiency of the Neon correction under more extreme temperature fluctuations,
314 we turned off the heating and cooling system in the Raman laboratory the night before a forecast
315 for a hot day in Berkeley. We start the test at ~ 6:30 am after 30 mins of instrument warm up time,
316 when the laboratory temperature was 23.5℃. Two Icelandic inclusions (Kahl et al. in review) were
317 double polished and mounted in a Linkam THM600 stage to keep the sample temperature at
318 37±0.1 ℃. As the room warmed up, these two FI were analysed repeatedly. After each acquisition
319 of the two FI, a Ne line was collected. Room temperature was measured using an Aranet4 sensor
320 located on the table with the Raman and Ne lamp (Fig. 6a). At around 3pm, the room temperature
321 had stabilized at around 26℃ as the external temperature peaked. To continue heating the lab, a
322 temporary heater was switched on. At 5 pm, the temperature began to level out again, so a second
323 heater was turned on, causing the laboratory temperature to rise to 28℃ at ~ 7:30 pm. Two air
324 filters with a combined clean-air delivery rate of 500 cubic feet per minute were used to circulate
325 air within the room to ensure even temperatures with these different heat sources.

326 The Ne correction factor shows a strong anticorrelation with temperature (Fig. 6a, note the
327 flipped y axis for temperature). By parametrizing temperature vs. time and Neon correction factor
328 vs. time using high order polynomials (Fig. 6a), a laboratory temperature and correction factor
329 was calculated for each FI measurement. We calculate the corrected splitting using the Ne
330 correction model (blue line, Fig. 6a) and compare this to corrected splitting calculated using the
331 first Ne acquisition at the start of the day (Ne1, Fig. 6a). Fitting a linear regression for calculated
332 density vs. laboratory temperature yields a far stronger correlation when using a constant Ne
333 correction factor vs. the Ne correction model (0.57 vs. 0.14 for B30, and 0.74 vs. 0.39 for B23).

334 This indicates that during large temperature swings, the Ne line correction does help correct for
335 instrument drift, keeping the rate of change within ~0.01 g/cm^3 for 10℃ of laboratory heating. Not
336 using a Neon lamp could generate ~0.03-0.04 g/cm^3 of drift for 10℃. This test also shows that
337 while it is better to have the Raman within a temperature-controlled lab, if Ne lines are used,
338 having a suboptimal room only doubles the typical analysis error. This is excellent news for
339 institutions and departments without the resources to build a temperature-controlled laboratory.

340   Overall, these tests demonstrates that spectral nonlinearity can generate significant
341 diversity in densimeters on a single instrument based on the choice of spectra center and neon
342 line correction routine, and that variability in these parameters can explain some of the offsets
343 between instruments with nearly identical hardware. While it may be tempting to shift the spectral
344 center to capture secondary peaks that appear in specific samples, or swap to a new Neon line
345 data processing strategy, this would require a new densimeter relationship to be developed. We
346 suggest that extreme caution must be used when applying a densimeter from another instrument,
347 which likely has a different nonlinearity characteristic and different acquisition parameters (e.g.
348 spectral center). Additionally, this analysis shows that just because two instruments have the
349 same spectral resolution, does not mean they have the same densimeter (c.f. Remigi et al., 2021).
350 In a temperature-stable laboratory, Ne line corrections are useful to tie the analysis of unknowns
351 to the densimeter, but at least for the WITec used here, the system is stable enough there is no
352 appreciable drift in the spectral nonlinearity over years. We suggest that measuring a Neon line
353 every 40-60 mins is generally sufficient and possibly too frequent if the instrument is housed in a
354 temperature-controlled laboratory environment. In contrast, if the laboratory undergoes large
355 temperature swings, more frequent Neon acquisitions are usefull to correct for drift (Fig. 6).
356 However, we have found that our Ne line acquisitions help us to identify days where the HVAC
357 system was malfunctioning and correct the data for those days, given their strong relationship
358 with laboratory temperature.

359   ### 3.2 Peak fitting errors
360 The above analysis shows that developing a densimeter for a specific analytical conditions and
361 drift correction methods is key to remove the large systematic offsets between different
362 densimeters, and generate accurate measurements of CO_2 density. Now, we focus our discussion
363 to optimizing the precision of Raman measurements. Bakker (2021) raised concerns regarding
364 precision, noting the discrete nature of Raman spectra, which comprise x coordinates
365 (wavenumber) collected on the CCD with spacing defined by the grating and other hardware. At
366 each x coordinate, an intensity value is recorded, the magnitude of which is highly dependent on

367 the instrument. Bakker (2021) show that shifting the spectral window one pixel at a time when
368 collecting a spectrum causes a lateral shift in the maximum peak intensity (Fig. 7). They suggest
369 that the error on the peak center is equal to the magnitude of the spectral shift. For two different
370 Raman set ups, they estimate a splitting error of ~ ±0.27 cm^{-1}. They emphasize that '*these*
371 *numbers must be regarded as the best estimated uncertainties of Raman band peak positions'*.
372 A splitting error of ±0.27 cm^{-1} yields an error of ±0.1 g/cm3 using the UCB densimeter. For a typical
373 melt inclusion vapour bubbles with 3% bubble volume, this corresponds to an uncertainty of the
374 amount of CO_2 held in the bubble of ±1111 ppm, which would introduce several kbar of uncertainty
375 into melt inclusion saturation pressures. Similarly, a FI with a density of 0.3±0.1 g/cm^3 corresponds
376 to pressures spanning 0.5-1.25 kbar (Span and Wagner, 1996). Overall, peak fitting errors of this
377 magnitude would make using Raman spectroscopy for melt and FI barometry untenable, so these
378 statements must be investigated.

379 We use DiadFit (Wieser and DeVitre, 2023) to refit the spectra of Bakker (2021) shown on their
380 Fig. 18 in DiadFit (Wieser and DeVitre, 2023) using pseudovoigt peaks. Their spectra were
381 collected as the spectral window was shifted ~3 cm^{-1} in steps of 0.137 cm^{-1} on a LabRamHR (600
382 grating, spectral resolution of 0.62 cm^{-1}). To help visualize the shift in the maximum y pixel
383 intensity, we vertically stack the acquisitions collected using different spectral centers after
384 normalizing the intensity for ease of visualization (Fig. 7a). The positions of diad 1 and diad 2 do
385 indeed move by significant amounts as the window is shifted (Fig. 7b). Bakker (2021) state that
386 the total error on the splitting is '*by definition, the sum of the uncertainties of the lower and upper*
387 *band (i.e. peak) of* CO_2', i.e. the amount of shift – yielding 0.274 cm^{-1} in this scenario (red error
388 bar, Fig. 7c). We also combine these errors in quadrature (purple error bar, Fig. 7c, total error=
389 $\sqrt{diad\ 1\ shift^2 + diad\ 2\ shift^2}$, Hughes and Hase, 2010). However, it is crucial to recognise that
390 diad 1 and diad 2 move in the same direction (Fig. 7b), such that the calculated splitting varies
391 very little (cyan diamonds, Fig. 7c). In fact, the overall variation in calculated splitting is smaller
392 than the estimated error from combining the peak fitting errors on diad 1 and diad 2 from lmfit in
393 quadrature within DiadFit (Newville et al., 2016). Clearly, the errors suggested by Bakker (2021)
394 are drastic overestimates, based on neglecting the similar effect shifting the window has on diad
395 1 and diad 2.

396 Bakker (2021) also suggest that the choice of different probability density functions (PDFs, e.g.
397 Voigt, PseudoVoigt, Gaussian) return different center positions, and that '*these best fit*
398 *reproductions provide peak position at maximum intensity, but do not consider uncertainty of each*
399 *measurement'.* Fortunately, modern peak fitting packages in python such as sklearn, or lmfit (used

400 in DiadFit, Newville et al., 2016) provide an estimate of the uncertainty on the peak center using
401 the covariance matrix obtained during the peak fitting process. To assess how much difference
402 the choice of PDF makes, we fit Raman data collected on natural FIs from Fogo Volcano, Cabo
403 Verde (DeVitre and Wieser, 2024), and Kīlauea Volcano, Hawai'i (DeVitre et al., 2024) using Voigt,
404 PseudoVoigt, and Pearson4 functions. These were identified as the most physically-plausible
405 peak shapes for CO_2 fluids by Yuan and Mayanovic (2017). Peak positions obtained from different
406 fits overlap within the calculated uncertainty on the peak center (Fig. 8, Supporting Fig. 3). This
407 can be shown more formally, by dividing the absolute difference between models by the fitting
408 error combined in quadrature (Fig. 8c-d). Y values <1 on this plot show that the difference between
409 functions is less than the 1 σ estimated error. This analysis demonstrates that for a wide range of
410 CO_2 densities and signal intensity, differences between physically plausible peak fit shapes are
411 within the stated error on each model, and that quoted peak fitting errors are realistic.

412 It is notable that the errors on the PseudoVoigt peak position are twice as large for diad 1 as diad
413 2, as a result of its lower intensity (Fig. 9a). Errors are also larger at lower CO_2 densities. However,
414 for all inclusions, peak fitting errors are still significantly smaller than the 0.1-0.3 cm^{-1} estimated
415 by Bakker (2021). Comparing Raman measurements with Microthermometry measurements for
416 specific fluid inclusions provide another opportunity to assess uncertainty (Fig. 9b). We use data
417 from N=58 FI from Fogo volcano from DeVitre and Wieser (2024) supplemented with N=39 new
418 measurements from Galápagos made in this study using the same technique. The Raman error
419 was estimated from Diadfit, by combining the peak fitting errors with the uncertainty from the Neon
420 line correction model and the densimeter. This error was combined in quadrature with the
421 uncertainty in microthermometry homogenization temperature (±0.2 ℃) to create the y error bar
422 shown in Fig. 9b. On Fig. 9c, we perform the same analysis, but now show error bars estimated
423 from multiple repeats of FI analyses, rather than the DiadFit peak fitting error. On both plots, the
424 two FI with the largest offset between Raman and Microthermometry have CO_2 densities very
425 close to critical, where Microthermometry errors are hard to quantify. The majority of inclusions lie
426 within the ~0.01-0.02 g/cm^3 error commonly quoted for the Raman method (red bars, Dayton et
427 al., 2023; DeVitre and Wieser, 2024; Kobayashi et al., 2012; Remigi et al., 2021). In contrast, it is
428 clear that a splitting error of ~0.2-0.3 cm-1 (grey dotted and dashed lines) from Bakker, (2021) is
429 a drastic overestimate (see also Yamamoto and Hagiwara, 2024, who suggest a best-case
430 scenario splitting error of 0.0005 cm^{-1} for a high pressure optical cell).

431 Although the errors estimated from comparing Microthermometry and Raman (Fig. 9), and
432 repeated measurements of standards (Fig. 5) are far smaller than that estimated by Bakker,

433 (2021), it is noteworthy that the standard deviation in density obtained from averaging multiple
434 repeats is about twice as large as the error estimated from DiadFit. This DiadFit error accounts
435 for uncertainty in peak position, Neon correction models, and the densimeter itself. In Section 3.3,
436 we suggest that a large fraction of this unaccounted-for uncertainty comes from the temperature
437 of the sample. Overall, once this temperature effect is accounted for, peak fitting errors estimated
438 from PDF functions, Ne line corrections and densimeters provide a good estimate of the
439 uncertainty of Raman analyses.

3.1.1. Optimizing Signals

441 While the above discussion indicates that peak fitting errors are far smaller than those indicated
442 by Bakker (2021), it is still worthwhile investigating methods to reduce uncertainties as much as
443 possible. This is particularly important for low density CO_2 fluids (i.e. melt inclusion vapour
444 bubbles), where relative errors can approach 50-100% near the instrument detection limit (~0.01-
445 0.02 g/cm^3, Bearden et al., 2025). Peak fitting errors are highly dependent on both the instrument
446 hardware (e.g., spectrometer efficiency, spectral resolution), analytical conditions (e.g., laser
447 power, acquisition time, number of accumulations averaged), and sample characteristics (e.g.,
448 density of CO_2, size of inclusion). Most instruments can adjust the acquisition time, which is the
449 time spent collecting a single spectrum, and the number of accumulations, which is the number
450 of individual acquisitions that are collected and averaged for the final spectrum. In general,
451 increasing the acquisition time results in higher counts, and more accumulations help to average
452 out noise in the signal. Increasing both parameters will result in better spectra, but also longer
453 analysis times.

454 We investigate the interplay between acquisition time and accumulations using a variety of
455 samples with low CO_2 densities, which have the largest peak fitting errors. To investigate the effect
456 of changing the number of accumulations (spectra averaged), we collect 14 consecutive individual
457 acquisitions of 45s duration using the 'slow series' option in the WITec acquisition software. We
458 fit each acquisition individually (i.e. a single accumulation). To simulate two accumulations, we
459 average the 1st and 2nd collected spectra, the 2-3rd, 3-4th and so forth, and then fit that averaged
460 spectra. To simulate 3 accumulations, we average the 1-2-3rd spectra, 2-3-4th, and for 4
461 accumulations, the 1-2-3-4th spectra, 2-3-4-5th spectra etc. Fig. 10 shows a CO_2 capillary tube
462 with a CO_2 density of 60 MPa, corresponding to ρ=0.011±0.011 g/cm^3 (Le et al., 2021). Supporting
463 Fig. 4 shows the same plot for a low density melt inclusion vapour bubble from Mauna Loa
464 (ρ=0.038±0.005 g/cm^3). Each 'averaged spectrum' is shown as a semi-transparent dot on panel
465 a-b, and the overall average for combinations yielding N accumulations is shown as a cross with

466   an error bar. These tests were performed for different integration times (45s, 60s, 90s, and 120s).

467   The errors are smaller for 90s compared with 45s, as expected given the more intense spectra,

468   and higher signal to noise ratio (Fig. 10c vs 11d), although for both, the error for diad1 is about

469   twice that as for diad 2. The error drops quickly as the number of accumulations is increased, and

470   then the trend flattens out (Fig. 10a-b). In terms of assessing ways to optimize instrument time, it

471   is useful to compare the errors in the context of the total acquisition time (Fig. 10e). Interestingly,

472   there are no clear differences outside of error; while 90 s acquisitions have slightly smaller errors

473   for a given total analysis time than 45 s acquisition for this capillary, for the vapour bubble, the

474   errors for 120 s are smaller (Supporting Fig. 4). For both series, increasing total analysis times

475   beyond about 4 minutes results in rapidly diminishing returns. We suggest that performing these

476   tests on each specific Raman instrument can help inform an analytical strategy that is cost-

477   effective and reduces peak fitting error as much as possible.

478   ### 3.3 Sample Temperature

479   It is important to investigate the effect of sample temperature on calculated densities. This has

480   implications for how the sample should be heated during Raman analyses. Heating samples

481   above the critical point of CO_2 (31.1℃) ensures that a homogenous fluid can be analysed, which

482   removes issues associated with analysis of a mixture of gaseous and liquid phases (see DeVitre

483   et al. 2023). We use a low cost heating apparatus consisting of a $60 DC power supply, a $41

484   Peltier plate with a hole in it to maintain the ability to use transmitted light for sample navigation,

485   and a $24 type K thermocouple with a precision of ±2℃ attached to the surface of the sample with

486   electrical tape (Supporting Fig. 1). However, the precision of type K thermocouples means that

487   the sample may be at 35℃ or 39℃, even if the display reads 37℃. Determining the error in splitting

488   and density that results from this uncertainty in sample temperature is important to determine

489   whether higher precision thermocouples or a more sophisticated heating apparatus is needed

490   (i.e. a Linkam stage with a precision of ±0.1℃).

491   A number of studies have investigated how sample (and thus fluid) temperature affects diad

492   splitting. Kawakami et al. (2003) measured CO_2-rich FIs with densities of ~1 g/cm^3 in mantle

493   xenoliths and found that splitting stayed remarkably constant between 20-200℃ (see their Fig.

494   10). Thus, they deduce that splitting is sensitive only to density, and is insensitive of temperature.

495   Wang et al. (2011) investigate splitting vs. temperature for a CO_2-rich FI (0.72 g/cm^3) and a H_2O-

496   CO_2 FI (see their Fig. 11). Both these FI exhibit a drop in splitting of ~0.2 cm^{-1} between ~40 and

497   200℃, corresponding to a drop in calculated density of 0.08 g/cm^3. As well as making

498   measurements on natural FIs with specific densities, Wang et al. (2011) also investigated

499   temperature sensitivity using a high-pressure optical apparatus where the analyte CO_2 gas was

500   held at a controlled pressure and temperature. Their acquisitions at 21, 25, 50, 100, 150 and

501   200℃ indicated that Δ decreases for a given CO_2 density as temperature increases (their

502   experiments span densities between ~0.2 and 0.7 g/cm^3).

503   Sublett et al. (2021) performed analyses on pure CO_2 in a high-pressure optical apparatus at

504   22℃, 50℃, 100℃, 200℃ and 350℃. They provide a parameterization for the relationship between

505   CO_2 density and splitting as a function of temperature (for 22-250℃ and 10-500 bars). We use

506   this equation to generate densimeters for 4 different temperatures (Fig. 11a). The strong increase

507   in pressure required to achieve a certain CO_2 content means that the densimeters for higher

508   temperatures terminate at progressively lower densities as the pressure limit of the apparatus

509   was reached. The temperature effect on splitting is largest at higher densities, and smaller at

510   lower densities, although all densities show some sensitivity to temperature. Hagiwara et al.

511   (2021, hereafter H2021) extended the results of Sublett et al. (2021, hereafter S2021) to higher

512   CO_2 densities, comparing densimeters for 23℃-200℃ for pressures between ~70-2500 bars,

513   yielding densities of ~0.65-1.2 g/cm^3 (Fig. 11b). For both H2021 and S2021, we calculate the

514   expected offset in density if an analysis is performed at X℃, but splitting is converted to density

515   using the densimeter obtained at 37 ℃ (Fig. 11c). This would reflect the sample being overheated

516   due to issues with a thermocouple and/or heating stage, or heating of the fluid by the laser itself

517   (see section 3.4). These models predict that CO_2 densities of ~0.6-0.8 g/cm^3 are the most

518   temperature sensitive, with the change in calculated density for 1 g/cm^3 being half that for 0.7

519   g/cm^3.

520   To assess these model predictions (Fig. 11c), we perform heating tests on natural olivine-hosted

521   FIs with densities spanning 0.4-0.9 g/cm^3 (Fig. 11 d-i). Double polished wafers were placed within

522   a LinkamTHMS600 stage and spectra were acquired at 6 mW laser power at temperatures

523   between 37℃ and 70℃. To account for the possibility of instrument drift, analyses were run with

524   increasing temperature and then decreasing temperature (2 repeats made at each temperature).

525   We fit a linear regression to plots of temperature vs. density, and use this to define a density at

526   37℃. For each acquisition, we then subtract this 37 ℃ from the calculated density. The change in

527   density is largest for FI with densities of ~0.4-0.7 g/cm^3, and smallest for ones at 0.9 g/cm^3. We

528   compare the change in density to that predicted from H2021 and S2021 (Fig. 11c, d-i). The

529   extrapolation of the H2021 model for the 0.4 g/cm^3 FI overpredicts the change in density with

530   heating, while S2021 (within its calibration range) underpredicts it. Both models, but particularly

531   H2021, do a good job of predicting the change in density for the inclusions at ~0.7 g/cm^3 (Fig.

532   11c, e-g). Both models drastically overpredict the density change for the inclusions at ~0.9 g/cm^3

533   (Fig. 11h-i).

534   These temperature-density gradients can be used to assess the uncertainty introduced by

535   measuring the sample temperature. The gradients (Fig. 11e-i) mean that the ±2℃ error introduced

536   from the thermocouple causes an additional error of ±0.0008-0.0028 g/cm^3. Combined with offsets

537   we observe based on the type of tape used to stick the thermocouple down (1-2℃), we believe

538   sample temperature variation accounts for the factor of 2 offset determined from repeated

539   analyses of standards vs. the average error estimated from the uncertainty in spectral and Ne

540   fitting, and the densimeter alone (Fig. 5). However, this is still a small uncertainty, and the

541   additional analytical time and instrument cost required to heat samples in a Linkam stage is not

542   justified.

3.4. Laser Power

544   These heating tests give us a framework to assess the effect of laser power on CO_2 density. This

545   has been a topic of extensive debate in the literature. Fall et al. (2011) suggest that the laser

546   doesn't heat the inclusion more than a few 10ths of a degree, based on observations that many

547   inclusions analysed in their laboratory which homogenized during microthermometry at

548   temperatures very close to laboratory temperature still showed distinctive liquid and gaseous

549   phases when analysed by Raman spectroscopy. However, Hagiwara et al. (2021b) simulate laser

550   heating with a numerical model, concluding that the absorption coefficient of the host mineral is

551   the most important parameter for determining the amount of heating, because the CO_2 inclusion

552   itself has very little absorption at the wavelengths of the laser. They perform Raman analyses at

553   varying laser powers, and use a thermometer based on hot bands to determine the laser heating

554   coefficient B (℃ per mW of laser power) for FIs in olivine, orthopyroxene, clinopyroxene, spinel

555   and quartz. It is notable how much lower their B values are for quartz (0.010-0.014℃/mW) vs.

556   mafic minerals such as olivine (0.02-1℃/mW). In their tests, as laser power is changed from 4-14

557   mW, the olivine FI shows a ~0.02 g/cm^3 decrease in density while the quartz inclusion shows no

558   change outside uncertainty. Thus, while Fall et al. (2011) do not note the host mineralogy of the

559   inclusions they draw their conclusions based on, the inclusion shown in their figure that they

560   reference as evidence for a lack of laser heating is hosted in quartz. If the Fall et al. (2011)

561   observations are from quartz inclusions, they are not in conflict with the results of Hagiwara et al.

562   (2021b) who predict minimal heating for quartz.. DeVitre et al. (2023) also demonstrate the effect

563   of laser heating on calculated CO_2 densities through the analyses of olivine-hosted melt inclusion

564   vapour bubbles with densities of ~0.3, 0.5 and 0.6 g/cm^3. They show a clear drop in splitting (and

565   thus density) with increasing power (~0.1 g/cm^3 from 0-20 mW).

566   It is vital to determine the effect of laser heating in different samples and bring some consensus

567   to the literature, because it is very tempting to use high laser powers for analysis to obtain strong

568   signals and small peak fitting errors. For example, Supporting Figure. 5 shows the peak fitting

569   error associated with the analysis of FIs and vapour bubbles using laser powers ranging from ~5-

570   35 mW. A laser power of ~5 mW yields a peak fitting error 2-3X higher than a laser power of 35

571   mW. This reduction in error seems highly beneficial for low density samples like melt inclusion

572   vapour bubbles (often 0.02-0.04 g/cm^3), where peak fitting errors can introduce an uncertainty of

573   20-50% in density.

574   We perform analyses at laser powers from ~5 to 50 mW on 4 olivine-hosted melt inclusion vapour

575   bubbles from Mauna Loa and the Cascades (Bearden et al., 2025; Wieser et al., 2023), 29 olivine-

576   hosted FIs from Iceland (Kahl et al. in review), Fogo volcano (DeVitre et al., 2023b), and American

577   Samoa, and 3 pyroxene-hosted FI from the Galápagos (Gleeson et al., 2025). These span a wide

578   range of CO_2 densities (Fig. 12, Supporting Figs 6-10). We also perform power series on the FI

579   and melt inclusion vapour bubbles that were used for heating series (Fig. 13, Supporting Fig. 6-

580   7). As for temperature, we analyse each FI with increasing then decreasing laser power. We

581   calculate the gradient of a linear regression between laser power and CO_2 density. We find that

582   this gradient varies dramatically between different fluids from ~0 to -0.005 g/cm^3 per mW (Fig.

583   12a, Supporting Figure 4-5). In general, FIs show weaker negative gradients than melt inclusions

584   from the same samples (see Supporting Fig. 10). However, FIs show great diversity in gradients.

585   The largest change in density is for FI 101C from Fogo volcano. Over 50 mW, the calculated

586   depth would change by 8 km (Fig. 13d). The three Icelandic FI shown in Fig. 12c-e show density

587   change varies from -0.08 g/cm^3 to -0.018 g/cm^3 for a 50 mW change in laser power. The drop in

588   density for B22_1_FIA between 0-50 mW corresponds to a change in calculated depth of ~ 3km

589   (Fig. 12c, RH axis). These changes in density are far larger than the uncertainty on each Raman

590   measurement (see error bars, Fig. 12c-e).

591   For a subset of fluid inclusions, EDS measurements were conducted to determine the forsterite

592   (Fo) content of the host olivine. We observe a strong correlation between the measured gradient

593   and the Fo content (Fig. 12b). This relationship is consistent with the known dependence of olivine

594   color on Fo content, which will affect the absorption coefficient. A wide variety of absorption

595   coefficients for olivine have been reported in the literature. For their laser heating simulations,

596   Hagiwara et al. (2021) use an olivine absorption coefficient of 1.14 cm^{-1} based on measurements

597   of a hydrothermal olivine from the Cola Peninsula, Russia (Fo_{95} Taran and Matsyuk, 2013). In

598   contrast, Chassé et al. (2015) report an olivine absorption coefficient at wavenumbers

599   corresponding to a 532 nm green laser (~18800 cm^{-1}) of 11 cm^{-1} (the Fo content is not given).

600   Other papers have shown that the absorption coefficient of olivine is sensitive to the Fo content,

601   and the crystallographic orientation of the host (Hazen et al., 1977; Taran and Matsyuk, 2013;

602   Ullrich et al., 2002). For example, at 532 nm, a Fo_{90} crystal has an absorption coefficient of ~6

603   cm^{-1} parallel to a vs. 3 cm^{-1} parallel to b (Ullrich et al. 2002). For Fo_{10}, the coefficient parallel to

604   the b axis is ~97 cm^{-1} and 76 cm^{-1} parallel to c. This large difference in absorption coefficient as

605   a function of olivine Fo content likely accounts for the stronger laser heating experienced by lower

606   Fo olivine hosts (Fig. 12b).

607   However, olivine chemistry cannot explain all the variability in laser heating rates. Crystal 101

608   from Fogo volcano contains three FI in the same secondary trail within ~100 µm of each other

609   (Supporting Fig. 12). FIC experiences a very dramatic drop in density, from ~0.93 g/cm^3 to ~0.75

610   g/cm^3 over 50 mW of heating (gradient of -0.033 g/cm^3 per mW). In contrast, FIB experiences no

611   change in density outside of uncertainty (gradient of -0.0001 g/cm^3 per mW), and FIA experiences

612   a minor drop with a gradient of -0.0006 g/cm^3 per mW. Initially, we thought the extreme heating

613   experienced by FIC may result from its proximity to the surface (~6 µm depth). However, when

614   the crystal was flipped over and a power series was performed from the other side (~68 µm depth),

615   the gradient was almost identical (Supporting Fig. 13). 3D mapping reveals that 101C contains

616   carbonate (Supporting Fig. 14). It has been suggested that microparticles which are not visible

617   optically can enhance the laser heating effect (e.g. Burke, 2001). It is possible that carbonate is

618   playing this role. The two Icelandic samples that lie off the trend in Fig. 12b also contain notable

619   carbonate in 3D maps. However, interestingly, of the three FI from American Samoa (pink edges,

620   Fig. 12b), the two which follow the gradient-Fo trend defined by the Icelandic samples contain

621   abundant carbonate, while the one with more negative gradients contains no carbonate.

622   Performing laser heating and laser power tests on the same FI allows us to estimate the heating

623   coefficient, B (℃/mW of laser power). For each FI, we calculate the gradient of density vs.

624   temperature, and density vs. laser power. On Fig. 13, we adjust the power and temperature scale

625   such that the gradients of the two datasets align. We calculate B as:

626    $B\ (℃/mW) = \dfrac{Grad\ (\frac{g}{cm3}\ /\ mW)}{Grad\ (\frac{g}{cm3}\ /\ ℃)}$

627   B values range from 0.22 to 5.3 ℃/mW, compared to 0.02-1℃/mW from Hagiwara et al. (2021).
628   Overall, these power series demonstrate that while laser heating can be partially predicted based
629   on the absorption coefficient of the host, the extreme change in density experienced by very
630   specific FI, replicated across different analytical sessions in our lab, is worrying. Given that many
631   melt and FIs are analysed prior to knowing the Fo content (as that is often the final step of
632   measurement) and that we cannot diagnose why a few FI show such extreme heating, this work
633   demonstrates that it is absolutely vital to use low laser powers, as the heating response is
634   currently not entirely predictable.

635   ### 3.5. Other strategies to optimize signal.
636   Having to use low laser powers is a real issue in terms of getting enough signal to minimise peak
637   fitting errors. In addition to optimizing the acquisition times (Fig. 10), it is worth thinking about
638   other ways to optimize signals. Our repeated measurements of FI standards (Fig. 5) not only allow
639   us to track the stability of the densimeter, but also the intensity of signals with time. If the FI
640   standard returns lower peak heights than normal, we analyse a Si chip. We use the oscilloscope
641   feature ('live spectra' on HORIBA) to perform micro adjustments on the pinhole position to
642   enhance laser transmission through the fiber. We find that this can easily increase signals by a
643   factor of 2-3. We typically adjust this every few weeks, although sometimes it goes months with
644   no decline in signal, and other times we have to readjust after a few days. Typically, after HVAC
645   or power failures, adjustments are needed, suggesting this adjustment is required based on
646   thermal expansion/contraction of components. On a mirror-based instrument, mirror alignments
647   will have a similar effect.

648   Once the system's optics are optimized, it is worth considering the analytical geometry. The depth
649   of the FI within the sample has been discussed in a variety of papers. Frezzotti et al. (2012) note
650   that the depth resolution of Raman spectroscopy degrades with depth in the sample, so it is best
651   to analyse FI within 30 μm of the surface. Burke, (2001) states that 'usually one strives to analyse
652   inclusions as near the surface as possible' inclusions deeper than 50 μm start to give problems,
653   especially if their density is not too high'. To assess the influence of FI depth, we individually
654   ground 3 olivine-hosted FI down, performing analyses after each round of grinding and polishing.
655   We measure the depth by focusing on the surface of the sample in RL, and then focusing down
656   to the melt inclusion. We correct the movement of the Z stage to true depth accounting for the
657   refractive index of olivine ($N_{olivine}$ ~1.7, 1.64-1.67 for Fo, 1.84-1.89 for Fa).

658   True depth = $N_{olivine} \, Z_{stage\ movement}$

659 We use the WITec autofocus function at each depth. Surprisingly, we find no coherent trends in
660 signal intensity (quantified here using the area of diad 2) or the error on calculated density with
661 increasing depth in the sample (Fig. 14a, Supporting Fig. 17-19). However, we do find big changes
662 with minor changes in focus at a single sample depth. To demonstrate this, we perform
663 acquisitions at different focuses when the FI was 91.8 µm deep in the sample. We first find the
664 optical focus yielding a sharp inclusion edge (yellow box and spectra, Fig. 14). We then use the
665 WITec autofocus feature at 50 mW. This collects spectra for ~3 s at various depths to find the
666 maximum intensity in user-selected spectral window. The z position was ~ 2 µm close to the
667 surface than the best optical focus and yielded very similar signal strengths (black spectra and
668 box). We then used the oscilloscope function to manually scroll through Z positions to find the
669 best signal (pink box and spectra). The 6 focus positions between -14.1 µm and 0 µm all yield
670 optically reasonable images, and yet the signal strength varies by a factor of 3. Interestingly, the
671 blue focus position (+2.3 µm) yields an optically poor image, but still a better focus than the
672 oscilloscope feature found. This comparison clearly demonstrates that focus has far more effect
673 on spectral strength than changing the depth in the sample by ~250 µm (Fig. 14 c-d).

674 This result was replicated for several inclusions (Supporting Fig. 17-21). We also find that the
675 signal can vary quite considerably spatially, and small shifts in x-y position can also result in
676 stronger signals. This becomes very apparent when taking 2D transects through and across FI
677 (Supporting Fig. 22-23). We suggest that when signals are weak, the autofocus or oscilloscope
678 feature should be used to investigate different x-y-z positions around the optimal focus. We
679 perform these tests at high laser power (49 mW) for speed, and the drop the power to perform
680 the analyses at the optimal Z position. Variability in signal strength can also be seen by performing
681 2D and 3D maps of signal intensity.

682 Conclusions
683 In summary, we investigate the effect of several instrument and data reduction methods on the
684 accuracy and precision of Raman analyses. Our findings can be summarized as follows:

685 1. The spectra center and drift correction method (e.g. choice of Ne and Ar lines) used for
686   densimeter calibration must be replicated when analysing unknowns in future sessions.
687   Spectral resolution alone cannot account for the diversity in published densimeters, and
688   even for a single spectral resolution, a single instrument can generate a wide variety of
689   densimeter relationships as the spectral center or drift correction method is changed
690   because of the nonlinearity of the spectrometer.

2. Estimates of peak fitting error from pseudovoigt curves, combined with uncertainty in the Ne line correction model, account for approximately half the scatter seen in repeated Raman analyses of the same inclusion over ~ 2 years (~0.004 g/cm^3 for FI with densities of >0.6 g/cm^3). These low uncertainties make prior suggestions of very large peak fitting errors untenable. The additional 0.004 g/cm^3 of scatter seen in repeated standard acquisitions likely results from uncertainty in the temperature of the sample. However, , this small uncertainty does not justify a more sophisticated heating apparatus for most geological applications.

3. We recommend tracking the density of a single standard during every run to ensure there has been no drift from the laboratory densimeter relationship. These standards also help to track the efficiency of the instrument (e.g. checking the peak heights are as high as previous days), so the user knows when re-alignment is required.

4. In a temperature-controlled laboratory, drift during and between analytical sessions is minimal using an WITec alpha300R Raman. However, when the laboratory temperature is perturbed, the Ne line correction helps to reduce instrument drift by a factor of ~3. Thus, while it is better to have a Raman in a temperature-controlled laboratory, it is not essential. Large temperature swings of +5℃ result in an additional uncertainty of <0.01 g/cm^3 if Ne lines are used for drift correction.

5. Laser heating is a significant issue, particularly for materials with higher optical absorption coefficients (e.g., silicate glass, lower Fo olivine). However, absorption coefficients alone are not enough to explain the large variability in heating response seen. 6mW laser power seems a good way to minimise heating for all but the most strongly laser-heated samples, while still providing sufficient signal to obtain good peak fits.

6. To help counteract the relatively weak signals achieved using low laser powers, the analyst should investigate the play off between uncertainty and acquisition time for their specific instrument. They should experiment with different focuses using an autofocus/live spectra function of the oscilloscope at high power to find the optimum signal within the fluid. Focus has a far larger effect than depth to the fluid within the sample (even at depths of ~250 µm).

samples, Drew Downs for the American Samoa samples, and Matthew Gleeson for the Galápagos samples. We are grateful to Alex Bearden for his early contributions to investigating laser heating in some Mauna Loa vapour bubbles.

Author Contributions

PW and CJD conceived the project. PW, CJD and IS collected the Raman data and performed peak fitting. PW wrote the manuscript. All authors contributed to the discussion of ideas and the editing of the manuscript.
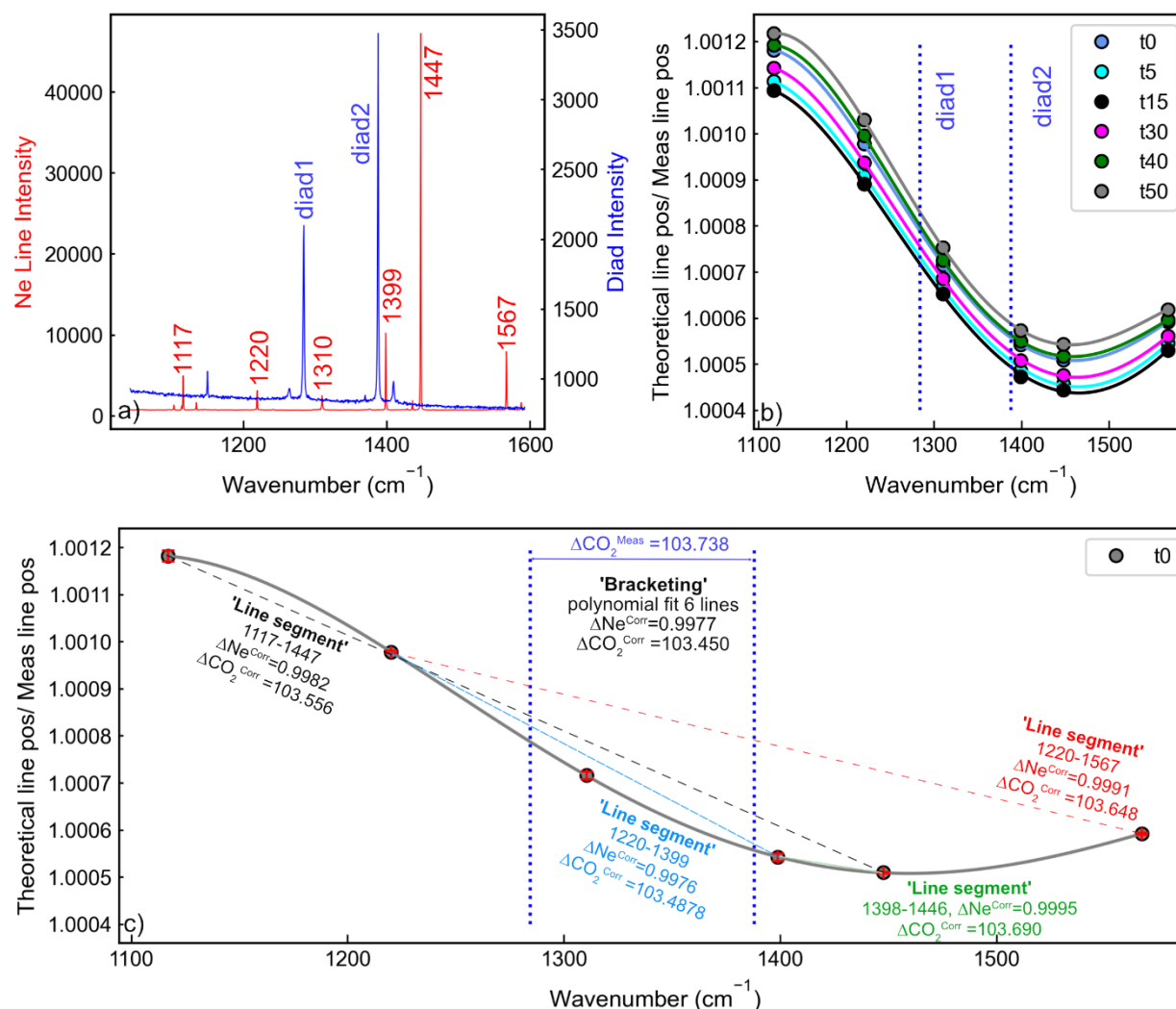
Figure 1 – Fitting Neon lines to demonstrate differences between the Line segment method and the bracketing technique. a) An example Ne line acquisition compared to a diad. Note that the intensity of the Ne (left hand axis) is almost 10X stronger than the diad (right hand axis). This is different from Fig. 8 of Bakker 2021 where Ne lines show very low intensities. b) For repeated acquisitions of Ne lines through the day, the actual peak position is compared to the theoretical peak position. For each line acquisition, a polynomial has been fitted through the data. c) For the first acquisition (t0), we show the polynomial fit through the 6 Ne peaks. The 'Bracketing technique' would correct each diad position based on the value for the polynomial at that wavenumber (where the blue dotted lines intersect the grey line). The 'Line segment' method will give similar results if the gradient drawn between two points is parallel to the intersection of the dotted blue and grey lines.
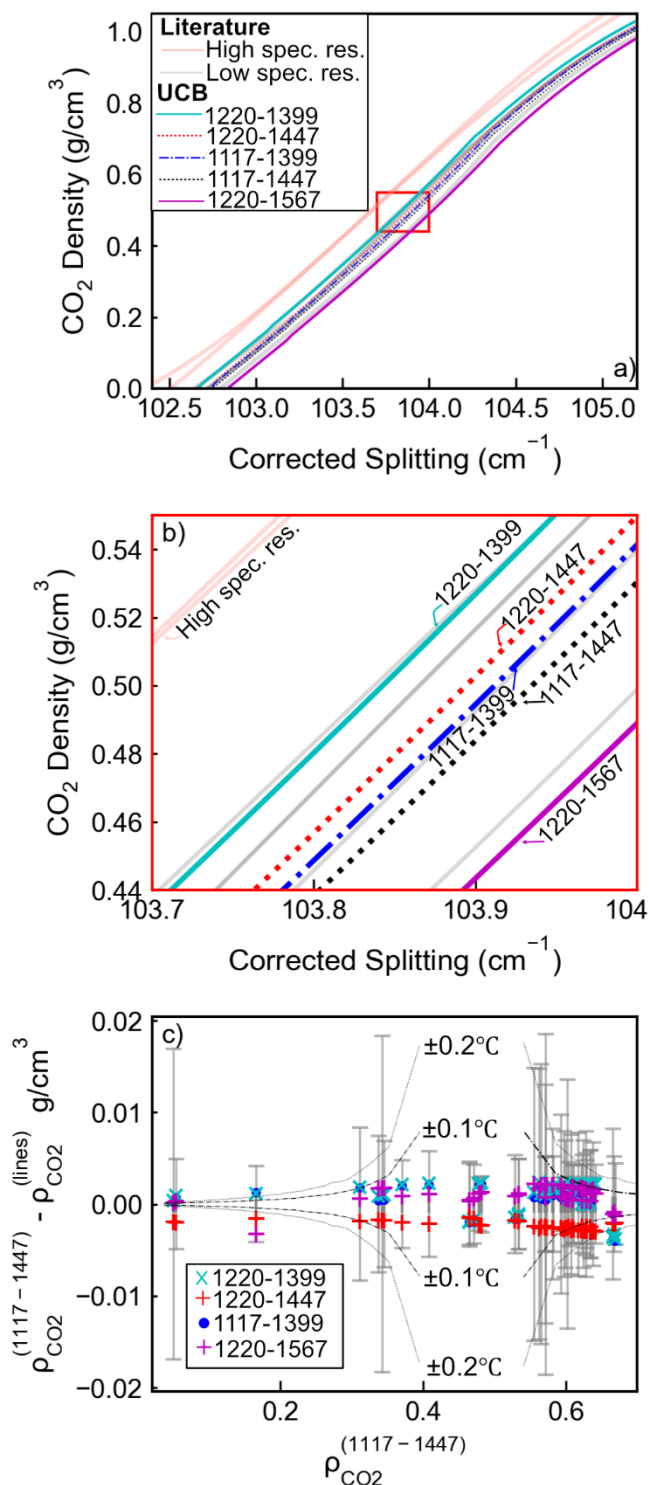
Figure 2 – Assessing the effect of different Ne line reduction routines on densimeters and densities calculated for natural FIs. a) The relationship between corrected splitting and CO_2 density (i.e. the densimeter) was determined using 5 different pairs of Neon lines for the same UCB calibration dataset. This variation is compared to other literature densimeters (grey = high precision, salmon – low precision based on the definitions of Remigi et al. 2021). b) Inset of area in red box in a). c) Natural FI data from DeVitre and Wieser, (2024) was reduced using these different Neon correction routines and densimeters. The y axis shows the difference in density for the default 1117-1447 line option compared with other data reduction strategies. The grey error bar shows the uncertainty on the 1117-1447 density. This demonstrates densities are within uncertainty as long as spectra from uknowns and calibration data are reduced in the same way. For reference, we overlay the uncertainty associated with Microthermometry for ±0.1 and 0.2℃ in sample temperature.

779

780
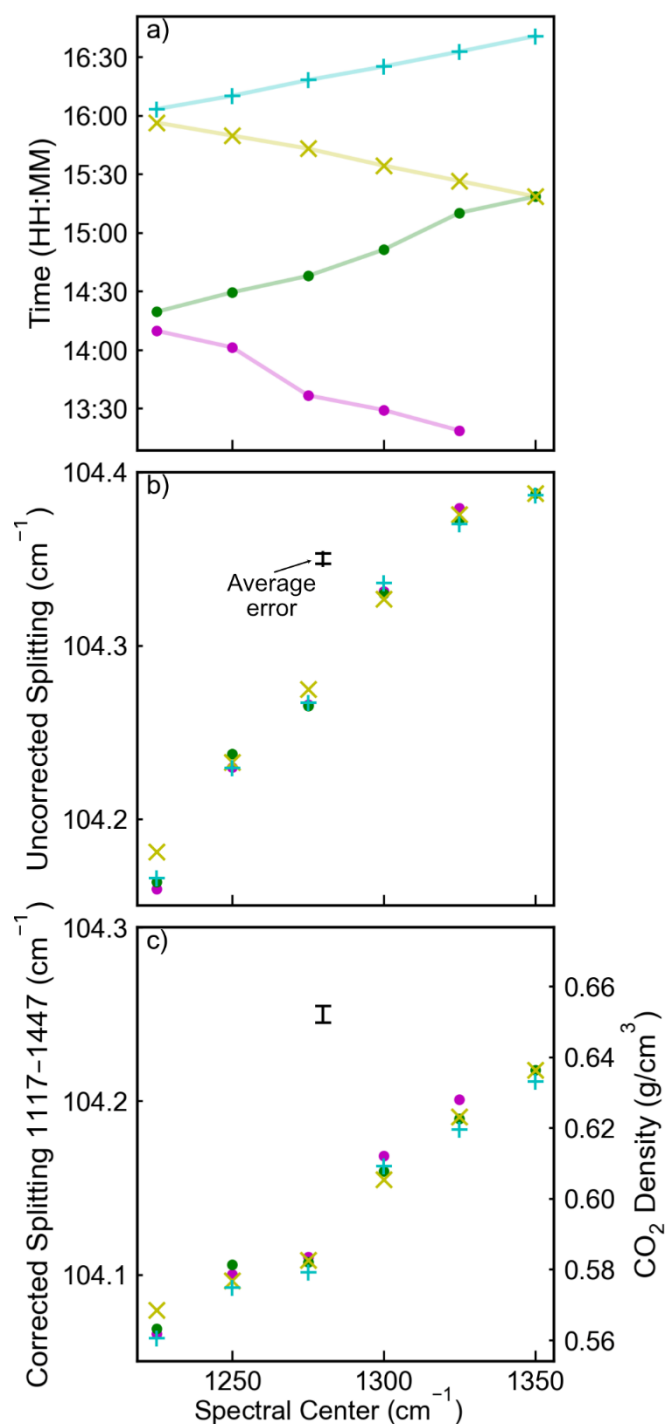
Figure 3 – Change in splitting and calculated density for a single FI as a function of the spectral center. a) The four symbol colors show 'up' and 'down' cycles changing the spectral center as a function of time through the day. b) Uncorrected splitting vs. spectral center and c) corrected splitting, using Ne lines collected at the same spectral center as the diad. The black error bars show a representative error for the splitting (b) and corrected splitting (c) respectively (offset from trend for clarity). The density axis in c is calculated using the densimeter developed for a spectral center of 1325 cm-1.
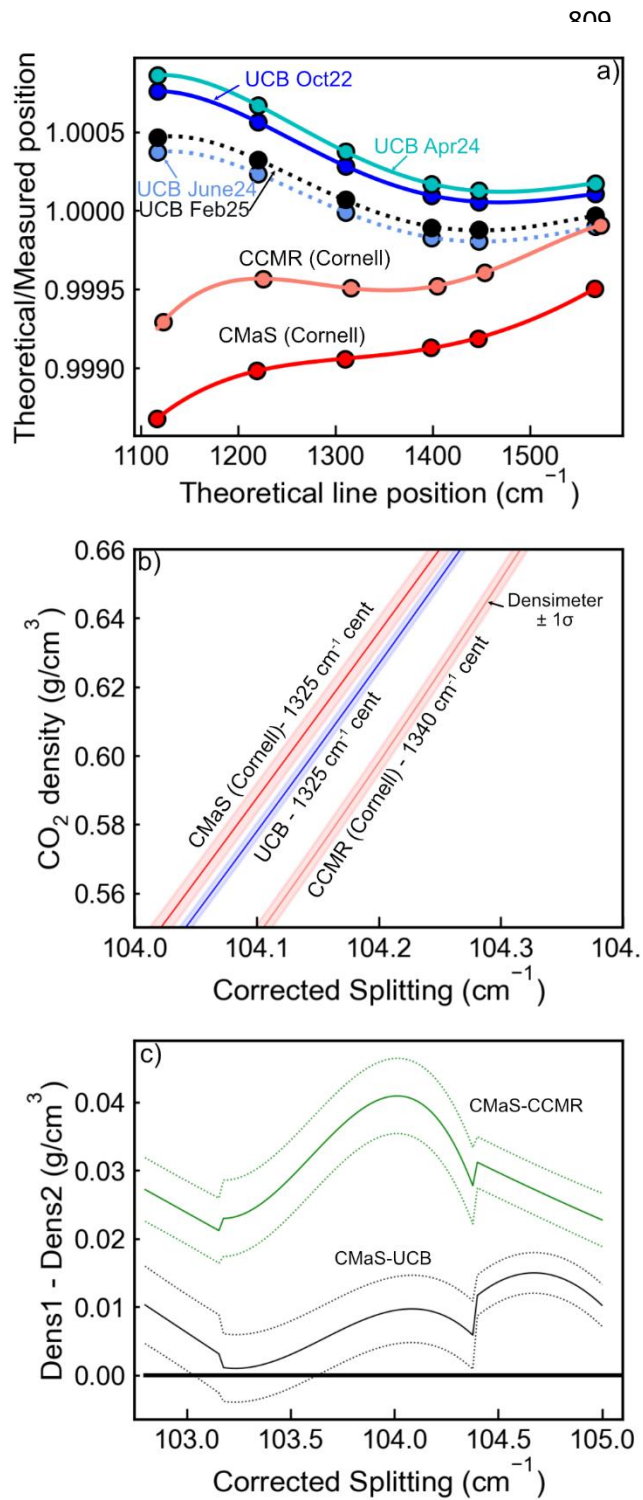
Figure 4 – a) Comparison of spectral nonlinearity determined from Ne line acquisitions for three different WITec alpha300R instruments. b) Comparison of the densimeters across a smaller density range showing that UCB and CMaS are more similar than CCMR and CMaS, despite CCMR-CMaS showing much more similar spectral non-linearities (a). c) Difference between densimeters as a function of splitting. The filled lines on b) show the 1σ errors on the densimeter fit, combined in quadrature and shown as dashed lines in c).

824
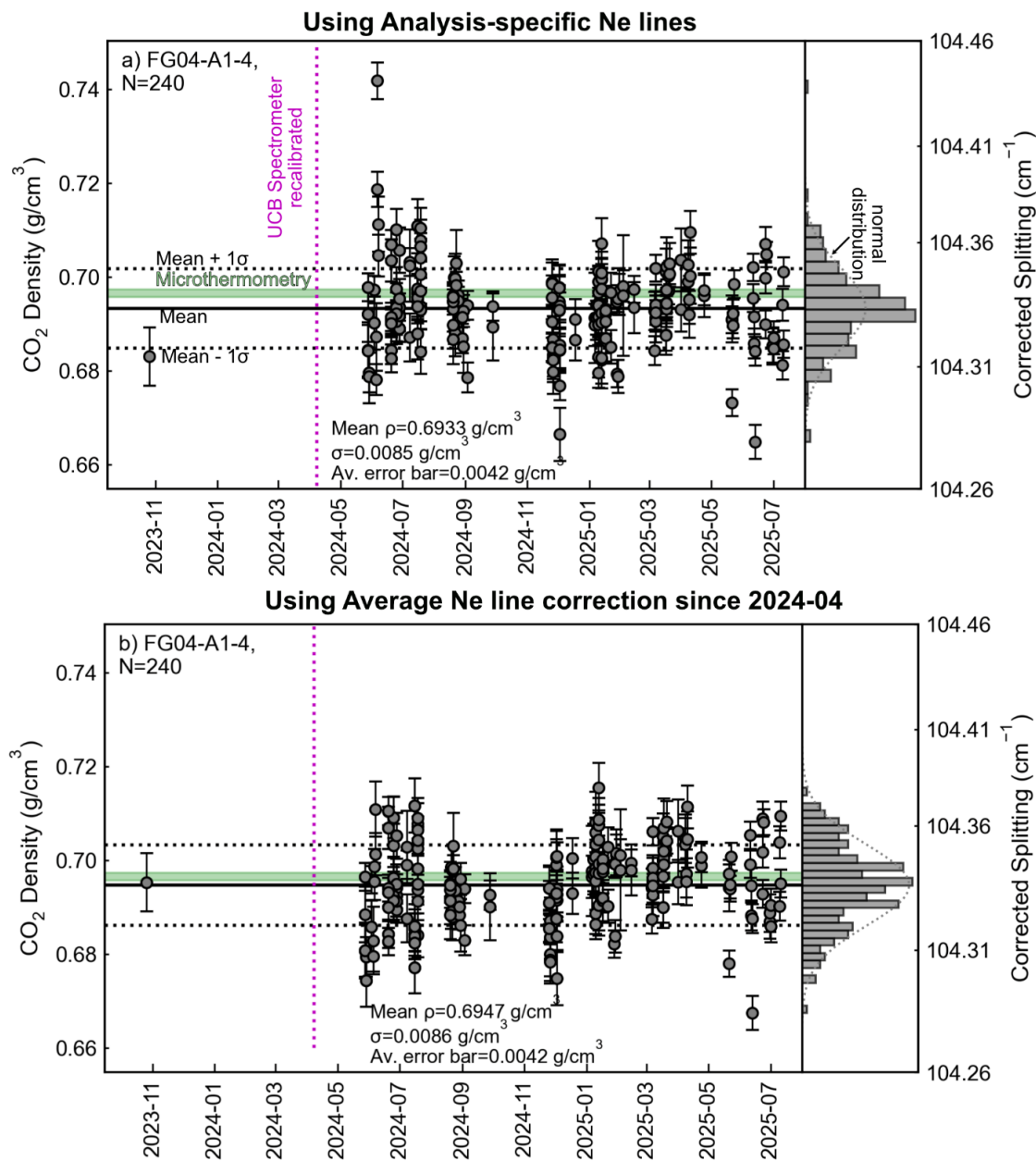
Figure 5- Long-term variability in standard value for FI FG04-A1-4 collected at UCB. a) Data reduced using a Ne drift correction model for each analytical session. b) Data reduced using the average Ne line correction factor for all standard acquisitions after the spectrometer recalibration (marked with a magenta line). The green bar shows the density determined by Microthermometry, accounting for the uncertainty in the temperature reading of the stage. The black solid line shows the mean Raman value for the standard, and the dotted black lines ±1 σ. The distribution of densities is shown as a histogram, compared to a normal distribution based on the mean and standard deviation. Supporting Fig. 2 shows a second standard run less frequently at UCB.

833
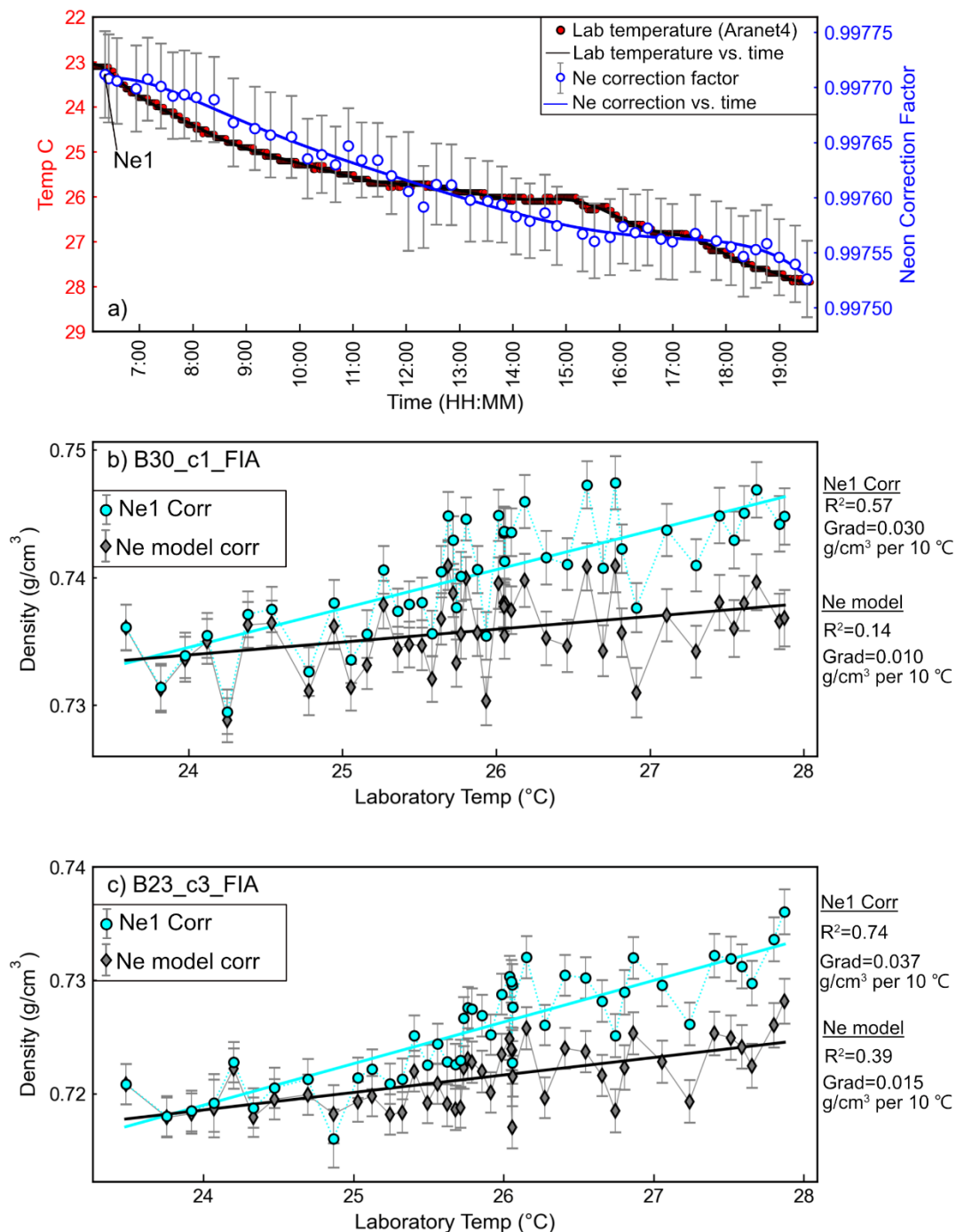
Figure 6 – Testing the sensitivity of Raman acquisitions to large fluctuations in room temperature. a) Temperature log from an Aranet4 logger (red dots) against time. The white circles with blue dots show the Ne correction factor calculated for each Ne line acquisition. These series were fitted with a regression (black and blue lines) to calculate temperature and Ne correction factor for each FI acquisition. b) Change in density of FI B30_c1_FIA as a function of laboratory temperature.

The grey diamonds are corrected based on the Ne line correction model (blue line in a), while the
840 cyan circles use the correction factor for the first Ne acquisition (to replicate not applying a drift
841 correction through time). The straight lines represent linear regressions, with the statistics shown
842 on the right. c) Same for B23_c3_FIA.
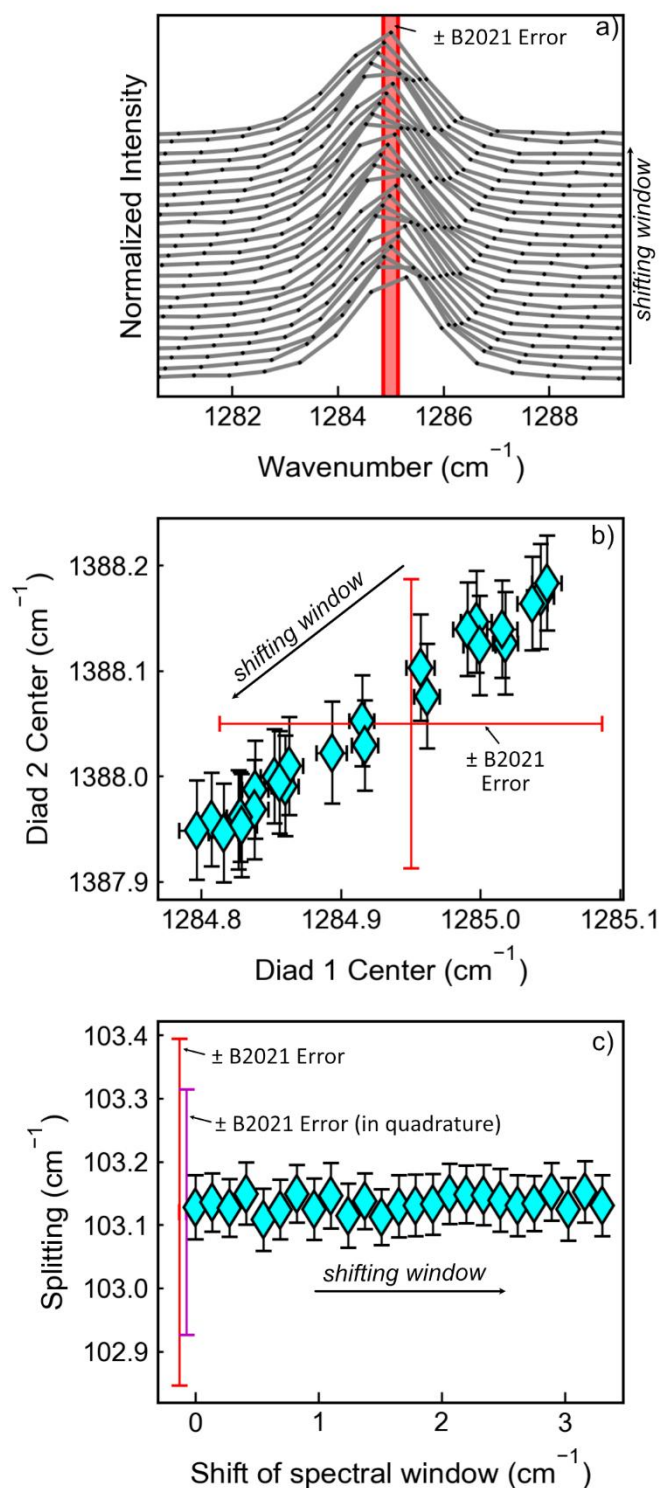
Figure 7. a) Stacked spectra collected as the spectral window was progressively shifted by Bakker (2021). The red band shows their stated error on the peak position, based on the wavenumber shift in the window. b) Plot showing the changing position of diad1 and diad2 for each acquisition shown in a). The error bar shows the 1 sigma error estimated by lmfit from the peak fit c) Change in splitting as a function of the shift of the window. Errors estimated by Bakker (2021) are shown in red, and purple shows their stated error on each peak combined in quadrature. The black error bar shows the uncertainty based on the lmfit peak positions propagated in quadrature.

858
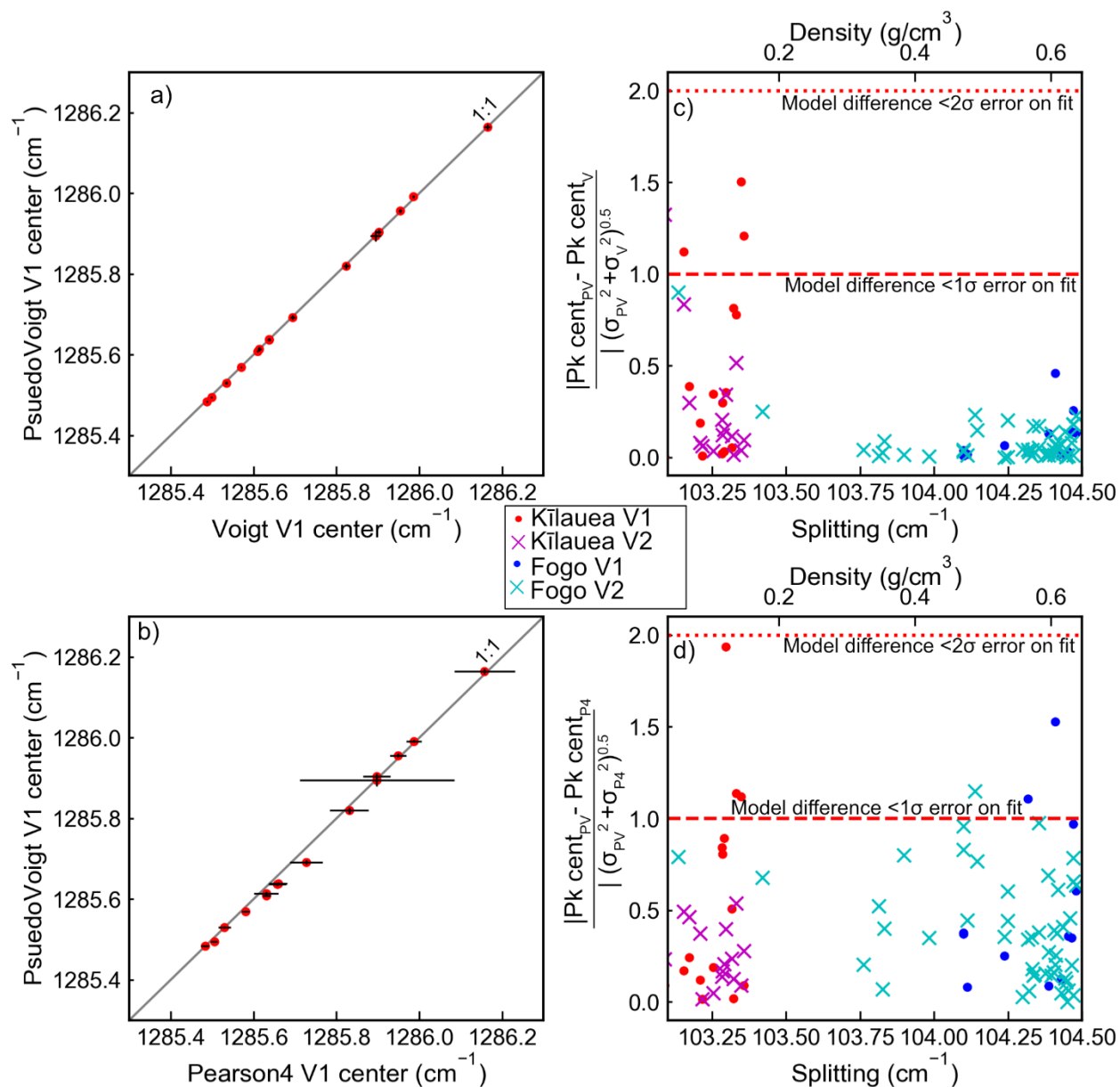
Figure 8 - Comparison of peak centers obtained with Voigt (V), Psuedovoigt (PV) and Pearson4 (P4) peak fits for Diad 1 (V1) and Diad2 (V2) for natural FIs from Fogo and Kīlauea Volcano. a-b) Comparison of Diad 1 position for Kīlauea FI, with errorbars showing the 1σ error on each peak fit from lmfit. c-d) The absolute difference between the peak center for each model is divided by the error on each model combined in quadrature. For datapoints lying below 1, the absolute difference between the models is smaller than the 1σ on each model. All differences are <2 σ. Comparison of PDF fits for the Fogo inclusions like in a-b) are shown in Supporting Fig. 3.
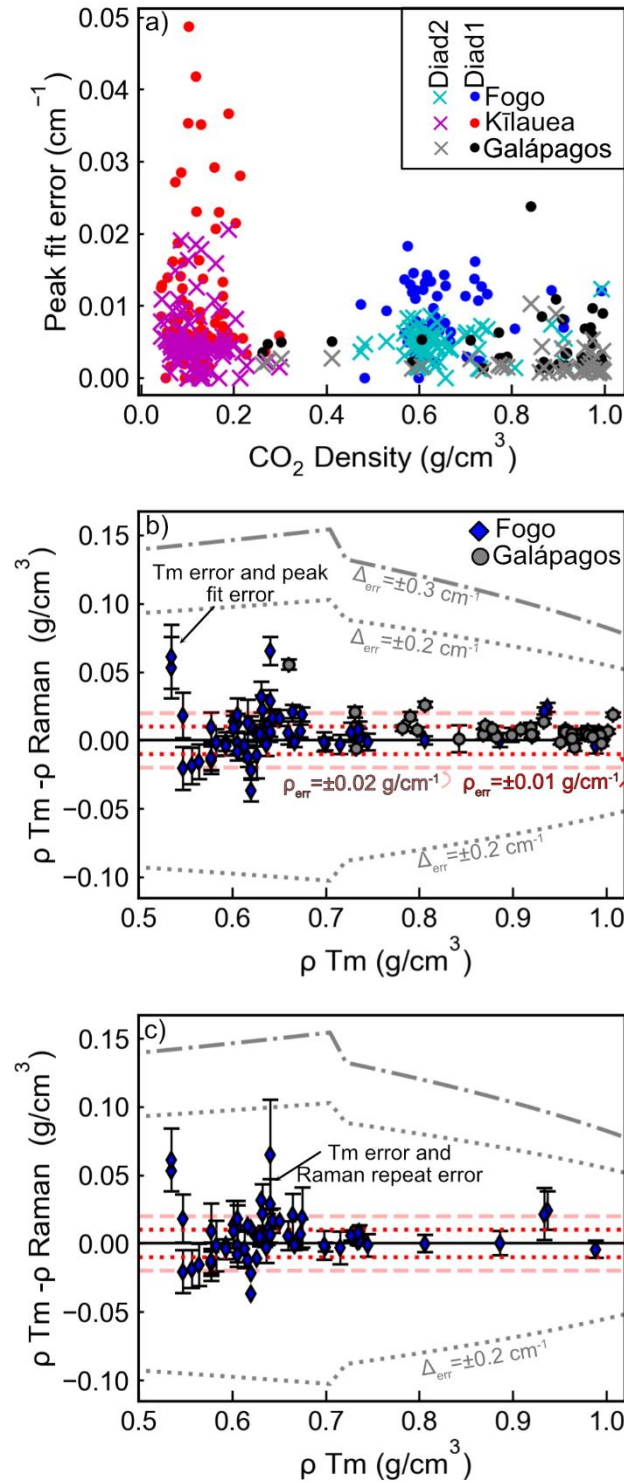
Figure 9 –Assessment of Raman errors using natural FIs with independent density constraints from microthermometry a) Errors on each peak position obtained from lmfit for low and high density FIs from DeVitre and Wieser (2024) from Kīlauea and Fogo, and from the Galápagos (this study). b) Comparison of densities obtained from Raman spectroscopy and Microthermometry. The error bars show 1 sigma errors estimated from peak fitting, analytical drift and densimeter uncertainty for Raman spectroscopy and Microthermometry combined in quadrature. c) Same comparison but using standard errors calculated from repeated Raman measurements (only FI with repeated analyses are shown). On b-c), we overlay the ±1 σ error bounds that would result from the splitting errors suggested by Bakker (2021). We also overlay red lines showing the errors determined from numerous studies comparing Raman and Microthermometry (±0.01-0.02 g/cm^3).
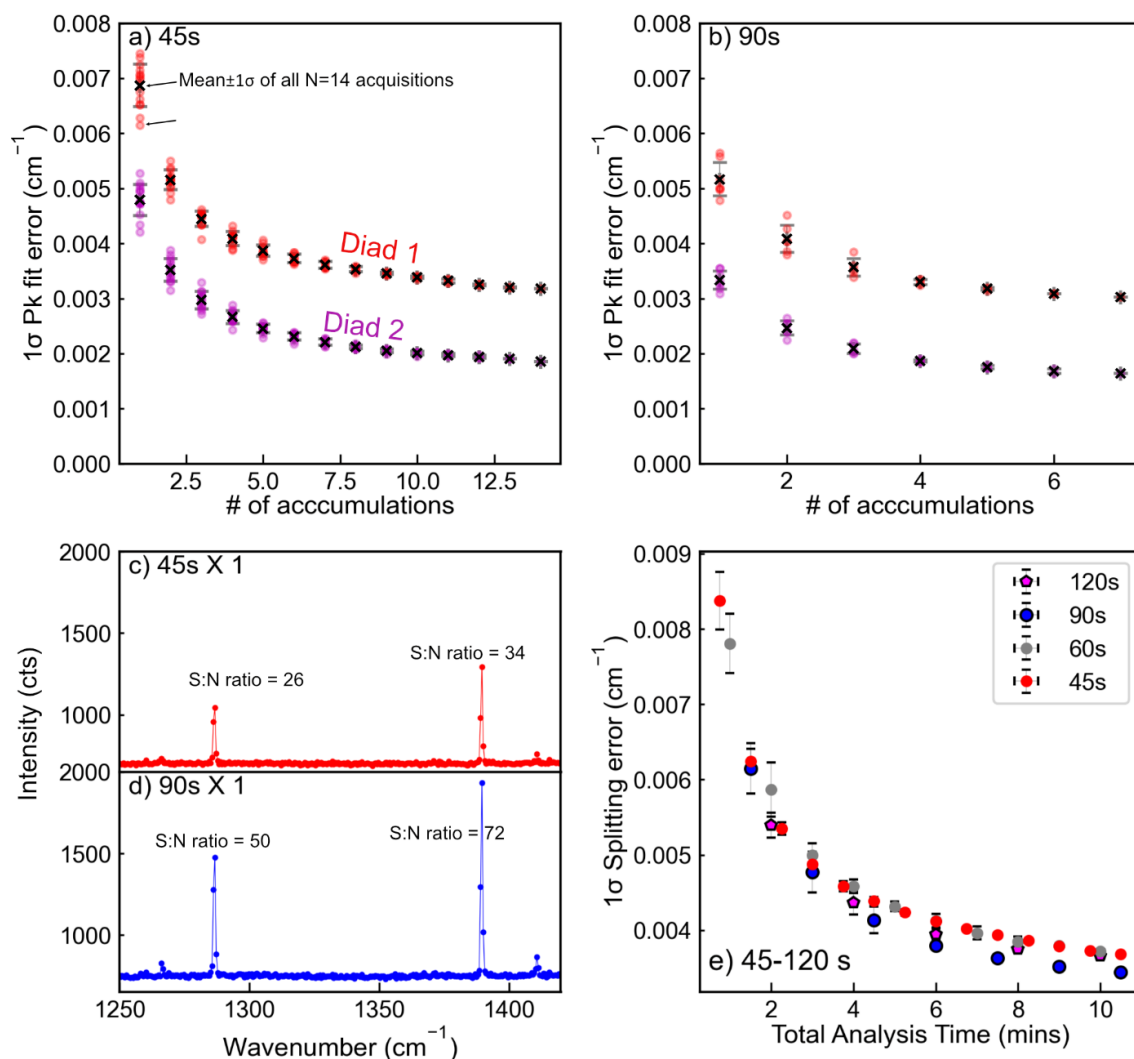
Figure 10 – Investigating peak fitting error as a function of integration time and accumulations using a low density capillary tube (~0.01 g/cm^3, Le et al., 2021). a-b) Error on the peak position of diad1 and diad2 for 45 and 90s acquisitions. The x axis shows the number of spectra that were averaged prior to fitting the peaks (analogous to accumulation). The dots show an individual fitted spectra (e.g. spectra 1 to 14 for accumulations =1, average of spectra 1-2, 2-3 3-4 for n=2 accumulations etc.), and the cross and error bar show the mean and standard deviation of averaged spectra for each number of accumulations (e.g. 14 individually fitted spectra for n=1 accumulations). c-d) Example spectra for 45s x 1 and 90s x 1. The signal to noise (S:N) ratio is calculated as the height of the highest pixel divided by the standard deviation of the region between the two diads. e) Error on peak fitting as a function of total acquisition time. For example, five 120s acquisitions were averaged for the 10 minute total acquisition time. The same figure for a low density melt inclusion vapour bubble is shown in Supporting Fig. 4.

903
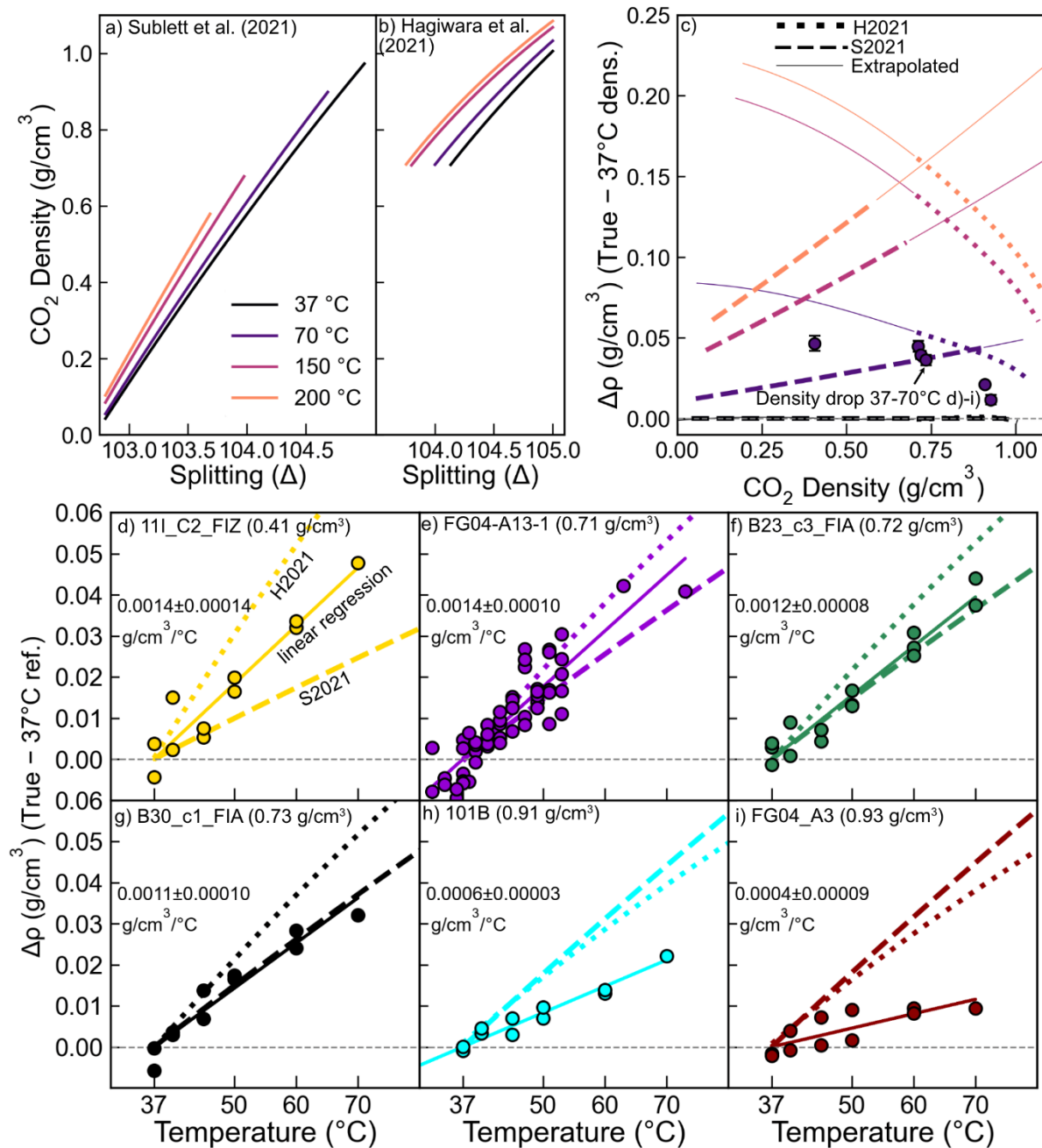
Figure 11 – Temperature sensitivity of splitting. a) Relationship between splitting and CO_2 for 4 temperatures calculated using equation 8 of Sublett et al. (2021, hereafter S2021). b) Same using equation 2 of Hagiwara et al. (2021a, hereafter H2021). We only show the lines within the calibration range of each study. c) Difference between true density of a FI and the density calculated using the 37℃ when analysed at T=X℃. Thin straight lines represent the equations of S2021 and H2021 when extrapolated outside their calibration range. d-i) For each inclusion, a linear regression was produced between calculated density using the 37℃ densimeter and the

911 sample temperature (Supporting Fig. 4). The 'True' density of the inclusion was determined from
912 the best fit line at 37℃. This was then compared to the calculated density at T=X using the 37℃
913 densimeter (i.e. panel d, the density at 70℃ was 0.05 g/cm3 lower than the intercept at 37 ℃).
914 The dotted and dashed lines show the predicted change in density using the models of H2021
915 and S2021 respectively.

916
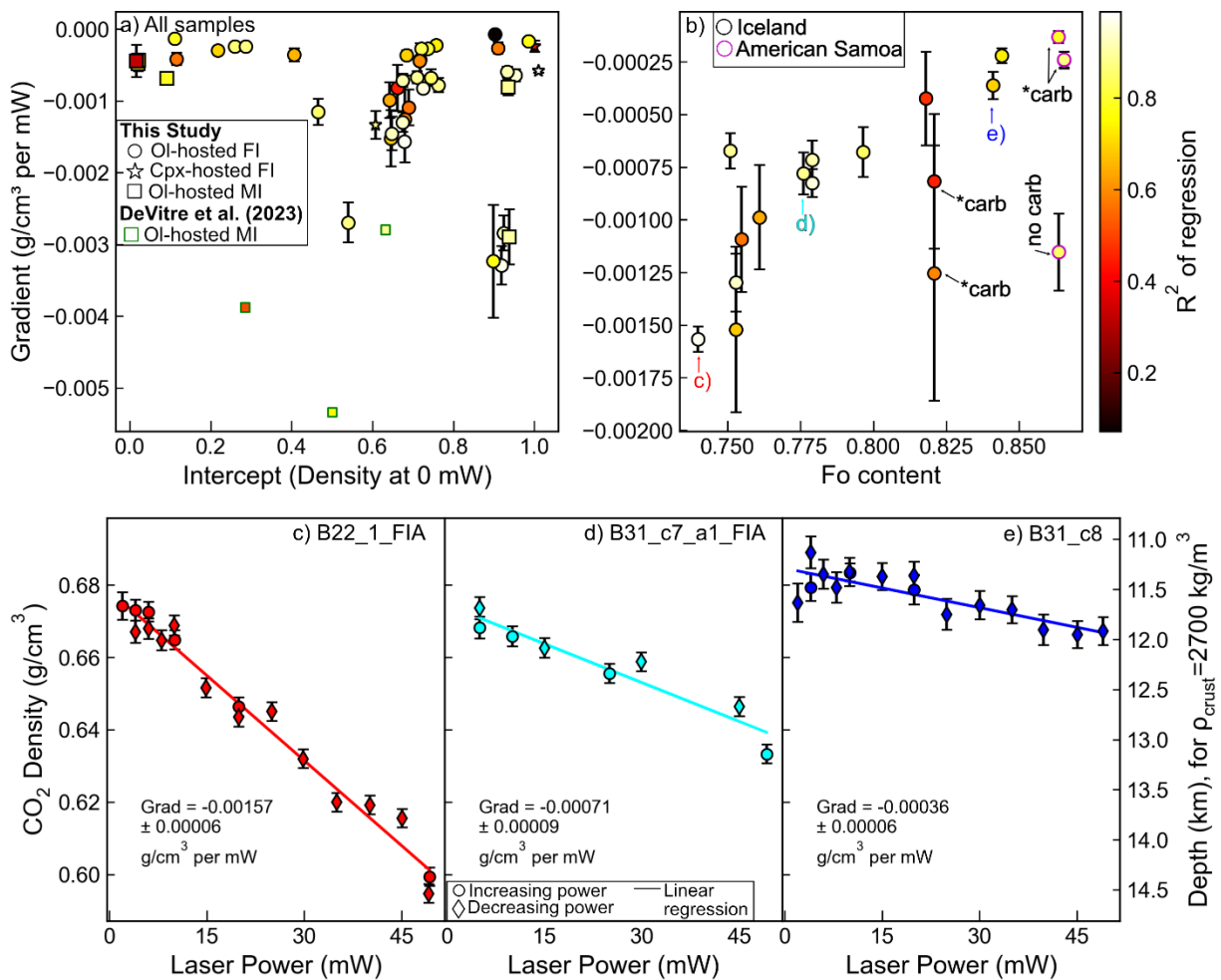
917 Figure 12 – Relationship between density and laser power. a) Gradient of laser power vs. density
918 for all the heating tests conducted here (individual regressions shown in Supporting Figures).
919 Symbol shapes show the host mineral and whether the sample is a fluid inclusion or melt
920 inclusion. Squares with green edges show the three melt inclusions from DeVitre et al. (2023) b)
921 Gradient of each power series vs. the forsterite content of the host crystal. FI with a carbonate
922 are indicated with arrows. c-e) Laser power series on three Icelandic FI indicated by arrows in b).

923   The circle symbols show data collected at increasing laser power, and the diamonds at decreasing

924   laser power. The gradients for the best-fit linear regression are shown.
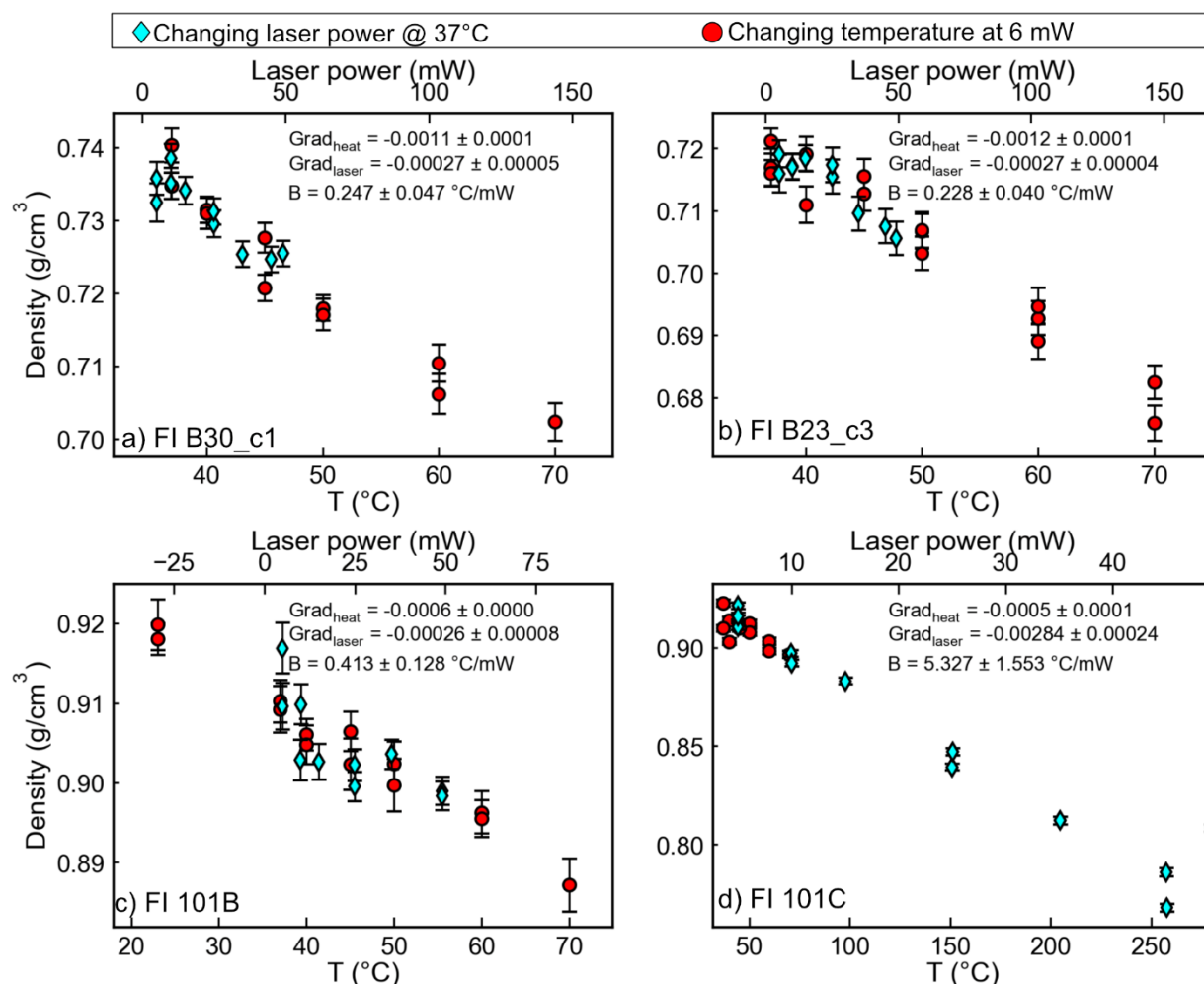
925

926   Figure 13 – Comparison of heating and laser power series. For each FI, we perform acquisitions

927   at temperatures from 37-70℃ (red circles) at 6 mW, and at laser powers from 6 mW to ~50 mW

928   with the sample temperature controlled at 37℃ (cyan diamonds). We adjust the limits of the two

929   x axes such that the gradients of these two data series are identical. The gradient of each series

930   is shown, along with the B value. The most noteworthy comparisons are 101B and 101C, which

931   are located within 100 μm of each other (see Supporting Fig. 12).
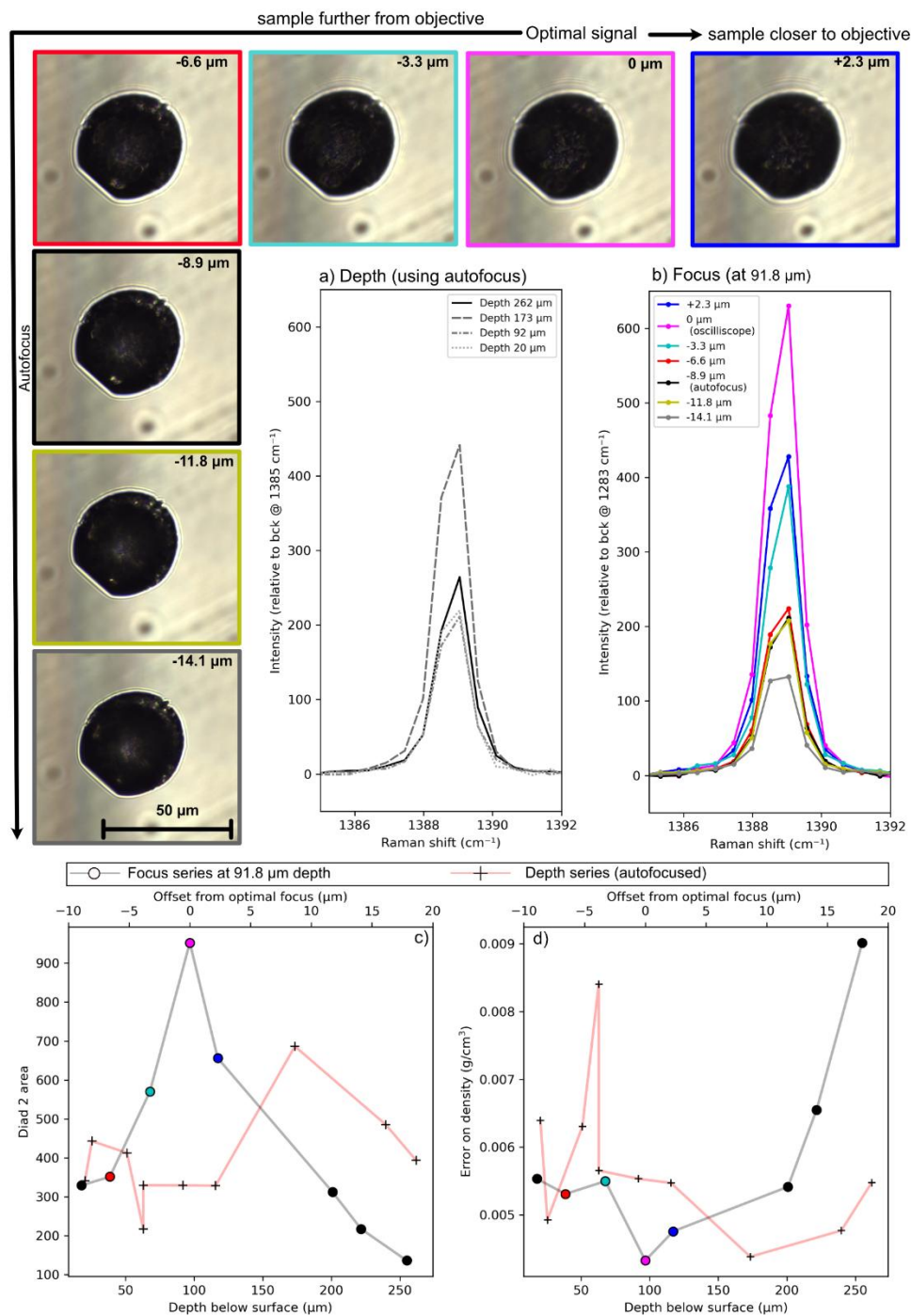
Figure 14 – Investigation of how signal strengths are affected by focus vs. depth in the sample.
a) Spectra near Diad 2, showing the signal strength for different depths of the FI in the sample
during serial sectioning. After each polishing round, the autofocus function was used to find the
optimum depth. b) When the FI was at 91.8 μm depth in the sample, we acquire spectra at 7
different focus positions. The black spectra and dot represents the focus position found by the

938 autofocus, and the pink spectra and dot the best focus position using the oscilloscope feature.
939 The images on the left-hand side show what the FI looked like at each focus. c) Change in diad
940 2 intensity and d) density error as a function of depth in the sample (circles) and focus (crosses).

941

942 References

943 Allison, C.M., Roggensack, K., Clarke, A.B., 2021. Highly explosive basaltic eruptions driven by
944   CO_2 exsolution. Nat. Commun. 12, 217. https://doi.org/10.1038/s41467-020-20354-2

945 Bakker, R.J., 2021. The perfection of Raman spectroscopic gas densimeters. J. Raman
946   Spectrosc. 52, 1923–1948. https://doi.org/10.1002/jrs.6245

947 Bearden, A.T., Wieser, P.E., Rangel, B., DeVitre, C.L., Gleeson, M.L.M., Monteleone, B.D.,
948   Lynn, K.J., 2025. Volatile systematics and magma storage depths at Mauna Loa: a melt
949   inclusion perspective., in: In Chapman Conference on Caldera-Forming Eruptions at
950   Basaltic Volcanoes: Insights and Puzzles from Kīlauea and Beyond. AGU.

951 Burke, E.A.J., 2001. Raman microspectrometry of fluid inclusions. Lithos 55, 139–158.
952   https://doi.org/10.1016/S0024-4937(00)00043-8

953 Chassé, M., Lelong, G., Van Nijnatten, P., Schoofs, I., De Wolf, J., Galoisy, L., Calas, G., 2015.
954   Optical Absorption Microspectroscopy (μ-OAS) Based on Schwarzschild-Type
955   Cassegrain Optics. Appl. Spectrosc. 69, 457–463. https://doi.org/10.1366/14-07628

956 Dayton, K., Gazel, E., Wieser, P., Troll, V.R., Carracedo, J.C., La Madrid, H., Roman, D.C.,
957   Ward, J., Aulinas, M., Geiger, H., Deegan, F.M., Gisbert, G., Perez-Torrado, F.J., 2023.
958   Deep magma storage during the 2021 La Palma eruption. Sci. Adv. 9, eade7641.
959   https://doi.org/10.1126/sciadv.ade7641

960 DeVitre, C.L., Allison, C.M., Gazel, E., 2021. A high-precision CO_2 densimeter for Raman
961   spectroscopy using a Fluid Density Calibration Apparatus. Chem. Geol. 584, 120522.
962   https://doi.org/10.1016/j.chemgeo.2021.120522

963 DeVitre, C.L., Dayton, K., Gazel, E., Pamukçu, A., Gaetani, G., Wieser, P.E., 2023a. Laser
964   heating effect on Raman analysis of CO_2 co-existing as liquid and vapor in olivine-
965   hosted melt inclusion bubbles. Volcanica 6, 201–219.
966   https://doi.org/10.30909/vol.06.02.201219

967 DeVitre, C.L., Gazel, E., Ramalho, R.S., Venugopal, S., Steele-MacInnis, M., Hua, J., Allison,
968   C.M., Moore, L.R., Carracedo, J.C., Monteleone, B., 2023b. Oceanic intraplate explosive
969   eruptions fed directly from the mantle. Proc. Natl. Acad. Sci. 120, e2302093120.
970   https://doi.org/10.1073/pnas.2302093120

971 DeVitre, C.L., Wieser, P.E., 2024. Reliability of Raman analyses of CO2-rich fluid inclusions as a
972   geobarometer at Kīlauea. Geochem. Perspect. Lett. 29, 1–8.
973   https://doi.org/10.7185/geochemlet.2404
974 DeVitre, C.L., Wieser, P.E., Bearden, A.T., Richie, A., Rangel, B., Gleeson, M.L.M., Grimsich, J.,
975   Lynn, K.J., Downs, D.T., Deligne, N.I., Mulliken, K.M., 2024. Depths in a day — A new
976   era of rapid-response Raman-based barometry using fluid inclusions. J. Petrol. egae119.
977   https://doi.org/10.1093/petrology/egae119
978 Fall, A., Tattitch, B., Bodnar, R.J., 2011. Combined microthermometric and Raman spectroscopic
979   technique to determine the salinity of H2O–CO2–NaCl fluid inclusions based on
980   clathrate melting. Geochim. Cosmochim. Acta 75, 951–964.
981   https://doi.org/10.1016/j.gca.2010.11.021
982 Fermi, E., 1931. Über den ramaneffekt des kohlendioxyds. Z. Für Phys. 71(3–4).
983 Frezzotti, M.L., Tecce, F., Casagli, A., 2012. Raman spectroscopy for fluid inclusion analysis. J.
984   Geochem. Explor. 112, 1–20. https://doi.org/10.1016/j.gexplo.2011.09.009
985 Gleeson, M., Wieser, P.E., DeVitre, C.L., Shi, S.C., Millet, M.-A., Muir, D.D., Stock, M.J.,
986   Lissenberg, J., 2025. Persistent High-Pressure Magma Storage beneath a Near-Ridge
987   Ocean Island Volcano (Isla Floreana, Galápagos). J. Petrol. 66, egaf031.
988   https://doi.org/10.1093/petrology/egaf031
989 Hagiwara, Y., Kawano, T., Takahata, K., Torimoto, J., Yamamoto, J., 2021a. Temperature
990   dependence of a Raman CO_2 densimeter from 23°C to 200°C and 7.2 to 248.7 MPa:
991   Evaluation of density underestimation by laser heating. J. Raman Spectrosc. 52, 1744–
992   1757. https://doi.org/10.1002/jrs.6188
993 Hagiwara, Y., Yoshida, K., Yoneda, A., Torimoto, J., Yamamoto, J., 2021b. Experimental variable
994   effects on laser heating of inclusions during Raman spectroscopic analysis. Chem. Geol.
995   559, 119928. https://doi.org/10.1016/j.chemgeo.2020.119928
996 Hazen, R., Mao, H., Bell, P.M., 1977. Effects of compositional variatoin on absorption spectra of
997   lunar olivine. Proc Lunar Sci Conf 8th 1081–1090.
998 Howe, T.A., Christopher, T.E., Moune, S., Tuffen, H., Schiavi, F., 2025. Melt inclusion bubbles
999   provide new insights into crystallisation depths and CO2 systematics at Soufrière Hills
1000   Volcano, Montserrat. Front. Earth Sci. 12, 1509409.
1001   https://doi.org/10.3389/feart.2024.1509409
1002 Hughes, I., Hase, T., 2010. Measurements and their uncertainties: a practical guide to modern
1003   error analysis. OUP.

1004 Kawakami, Y., Yamamoto, J., Kagi, H., 2003. Micro-Raman Densimeter for CO_2 Inclusions in
1005   Mantle-Derived Minerals. Appl. Spectrosc. 57, 1333–1339.
1006   https://doi.org/10.1366/000370203322554473

1007 Kobayashi, T., Yamamoto, J., Hirajima, T., Ishibashi, H., Hirano, N., Lai, Y., Prikhod'ko, V.S.,
1008   Arai, S., 2012. Conformity and precision of CO_2 densimetry in CO_2 inclusions:
1009   microthermometry *versus* Raman microspectroscopic densimetry: Conformity and
1010   precision of CO_2 densimetry in CO_2 inclusions. J. Raman Spectrosc. 43, 1126–1133.
1011   https://doi.org/10.1002/jrs.3134

1012 Kramida, A., Ralchenko, Y., 1999. NIST Atomic Spectra Database, NIST Standard Reference
1013   Database 78. https://doi.org/10.18434/T4W30F

1014 Lamadrid, H.M., Moore, L.R., Moncada, D., Rimstidt, J.D., Burruss, R.C., Bodnar, R.J., 2017.
1015   Reassessment of the Raman CO_2 densimeter. Chem. Geol. 450, 210–222.
1016   https://doi.org/10.1016/j.chemgeo.2016.12.034

1017 Le, V.-H., Caumon, M.-C., Tarantola, A., 2021. FRAnCIs calculation program with universal
1018   Raman calibration data for the determination of PVX properties of CO2–CH4–N2 and
1019   CH4–H2O–NaCl systems and their uncertainties. Comput. Geosci. 156, 104896.
1020   https://doi.org/10.1016/j.cageo.2021.104896

1021 McLennan F.R.S., J.C., Smith M.A., H.D., 1932. RAMAN EFFECTS WITH LIQUID AND SOLID
1022   CARBON DIOXIDE. Can. J. Res. 7, 551–555. https://doi.org/10.1139/cjr32-104

1023 Moore, L.R., Gazel, E., Tuohy, R., Lloyd, A.S., Esposito, R., Steele-MacInnis, M., Hauri, E.H.,
1024   Wallace, P.J., Plank, T., Bodnar, R.J., 2015. Bubbles matter: An assessment of the
1025   contribution of vapor bubbles to melt inclusion volatile budgets. Am. Mineral. 100, 806–
1026   823. https://doi.org/10.2138/am-2015-5036

1027 Newville, M., Stensitzki, T., Allen, D.B., Rawlik, M., Ingargiola, A., Nelson, A., 2016. LMFIT: Non-
1028   linear least-square minimization and curve-fitting for Python. Astrophys. Source Code
1029   Libr. ascl-1606.

1030 Remigi, S., Mancini, T., Ferrando, S., Frezzotti, M.L., 2021. Interlaboratory Application of Raman
1031 CO_2 Densimeter Equations: Experimental Procedure and Statistical Analysis Using
1032   Bootstrapped Confidence Intervals. Appl. Spectrosc. 75, 867–881.
1033   https://doi.org/10.1177/0003702820987601

1034 Rosso, K.M., Bodnar, R.J., 1995. Microthermometric and Raman spectroscopic detection limits
1035   of C02 in fluid inclusions and the Raman spectroscopic characterization of C02.
1036   Geochim. Cosmochim. Acta 59, 3961–3975. https://doi.org/10.1016/0016-
1037   7037(95)94441-H

1038 Song, Y., Chou, I., Hu, W., Robert, B., Lu, W., 2009. CO2 Density-Raman Shift Relation Derived
1039   from Synthetic Inclusions in Fused Silica Capillaries and Its Application. Acta Geol. Sin. -
1040   Engl. Ed. 83, 932–938. https://doi.org/10.1111/j.1755-6724.2009.00090.x

1041 Span, R., Wagner, W., 1996. A New Equation of State for Carbon Dioxide Covering the Fluid
1042   Region from the Triple-Point Temperature to 1100 K at Pressures up to 800 MPa. J.
1043   Phys. Chem. Ref. Data 25, 1509–1596. https://doi.org/10.1063/1.555991

1044 Sublett, D.M., Sendula, E., Lamadrid, H.M., Steele-MacInnis, M., Spiekermann, G., Bodnar,
1045   R.J., 2021. Raman spectral behavior of N_2, CO_2, and CH_4 in N_2–CO_2–CH_4 gas
1046   mixtures from 22°C to 200°C and 10 to 500 bars, with application to other gas mixtures.
1047   J. Raman Spectrosc. 52, 750–769. https://doi.org/10.1002/jrs.6033

1048 Taran, M.N., Matsyuk, S.S., 2013. Fe2+, Mg-distribution among non-equivalent structural sites
1049   M1 and M2 in natural olivines: an optical spectroscopy study. Phys. Chem. Miner. 40,
1050   309–318. https://doi.org/10.1007/s00269-013-0572-x

1051 Ullrich, K., Langer, K., Becker, K.D., 2002. Temperature dependence of the polarized electronic
1052   absorption spectra of olivines. Part I - fayalite. Phys. Chem. Miner. 29, 409–419.
1053   https://doi.org/10.1007/s00269-002-0248-4

1054 Wang, W., Caumon, M.-C., Tarantola, A., Pironon, J., Lu, W., Huang, Y., 2019. Raman
1055   spectroscopic densimeter for pure CO2 and CO2-H2O-NaCl fluid systems over a wide
1056   P-T range up to 360 °C and 50 MPa. Chem. Geol. 528, 119281.
1057   https://doi.org/10.1016/j.chemgeo.2019.119281

1058 Wang, X., Chou, I.-M., Hu, W., Burruss, R.C., Sun, Q., Song, Y., 2011. Raman spectroscopic
1059   measurements of CO2 density: Experimental calibration with high-pressure optical cell
1060   (HPOC) and fused silica capillary capsule (FSCC) with application to fluid inclusion
1061   observations. Geochim. Cosmochim. Acta 75, 4080–4093.
1062   https://doi.org/10.1016/j.gca.2011.04.028

1063 Wieser, P., DeVitre, C., 2023. DiadFit: An Open-SourcePython3 Tool for Peak fitting of Raman
1064   Data from silicate melts and CO2 fluids (preprint). Earth Sciences.
1065   https://doi.org/10.31223/X5CQ1F

1066 Wieser, P., Kent, A.J.R., DeVitre, C.L., Gazel, E., Till, C.B., Johnson, Emily, Wallace, Paul,
1067   Couperthwaite, F.K., 2023. Bubble Trouble: Raman measurements of Cascade melt
1068   inclusion vapour bubbles indicate substantial underestimation of magma storage depths
1069   in legacy data. Presented at the AGU Fall Meeting 2023. Poster No. 0203, id. V23D-
1070   0203.

1071   Wieser, P.E., Lamadrid, H., Maclennan, J., Edmonds, M., Matthews, S., Iacovino, K., Jenner,
1072   F.E., Gansecki, C., Trusdell, F., Lee, R.L., Ilyinskaya, E., 2021. Reconstructing Magma
1073   Storage Depths for the 2018 Kīlauean Eruption From Melt Inclusion CO_2 Contents: The
1074   Importance of Vapor Bubbles. Geochem. Geophys. Geosystems 22.
1075   https://doi.org/10.1029/2020GC009364
1076   Wright, R.B., Wang, C.H., 1973. Density effect on the Fermi resonance in gaseous CO2 by
1077   Raman scattering. J. Chem. Phys. 58, 2893–2895. https://doi.org/10.1063/1.1679594
1078   Yamamoto, J., Hagiwara, Y., 2024. Precision evaluation of Raman densimetry for carbon
1079   dioxide: improvement by correction for the drift effect. Appl. Opt. 63, 1402.
1080   https://doi.org/10.1364/AO.507939
1081   Yamamoto, J., Kagi, H., 2006. Extended Micro-Raman Densimeter for CO_2 Applicable to
1082   Mantle-originated Fluid Inclusions. Chem. Lett. 35, 610–611.
1083   https://doi.org/10.1246/cl.2006.610
1084   Yuan, X., Mayanovic, R.A., 2017. An Empirical Study on Raman Peak Fitting and Its Application
1085   to Raman Quantitative Research. Appl. Spectrosc. 71, 2325–2338.
1086   https://doi.org/10.1177/0003702817721527
1087

1088

1089

1090

1091
1092

Supporting Information for: Best practices for the analyses of CO₂ fluids by Raman Spectroscopy

Penny E. Wieser[1*], Charlotte L DeVitre[1, 2], Isabelle Susman[1]

[1] Earth and Planetary Sciences, University of California, Berkeley, CA 94270, USA

[2] Now at Department of Earth and Environmental Sciences, University of Ottawa, Ottawa, ON K1N 6N5, Canada
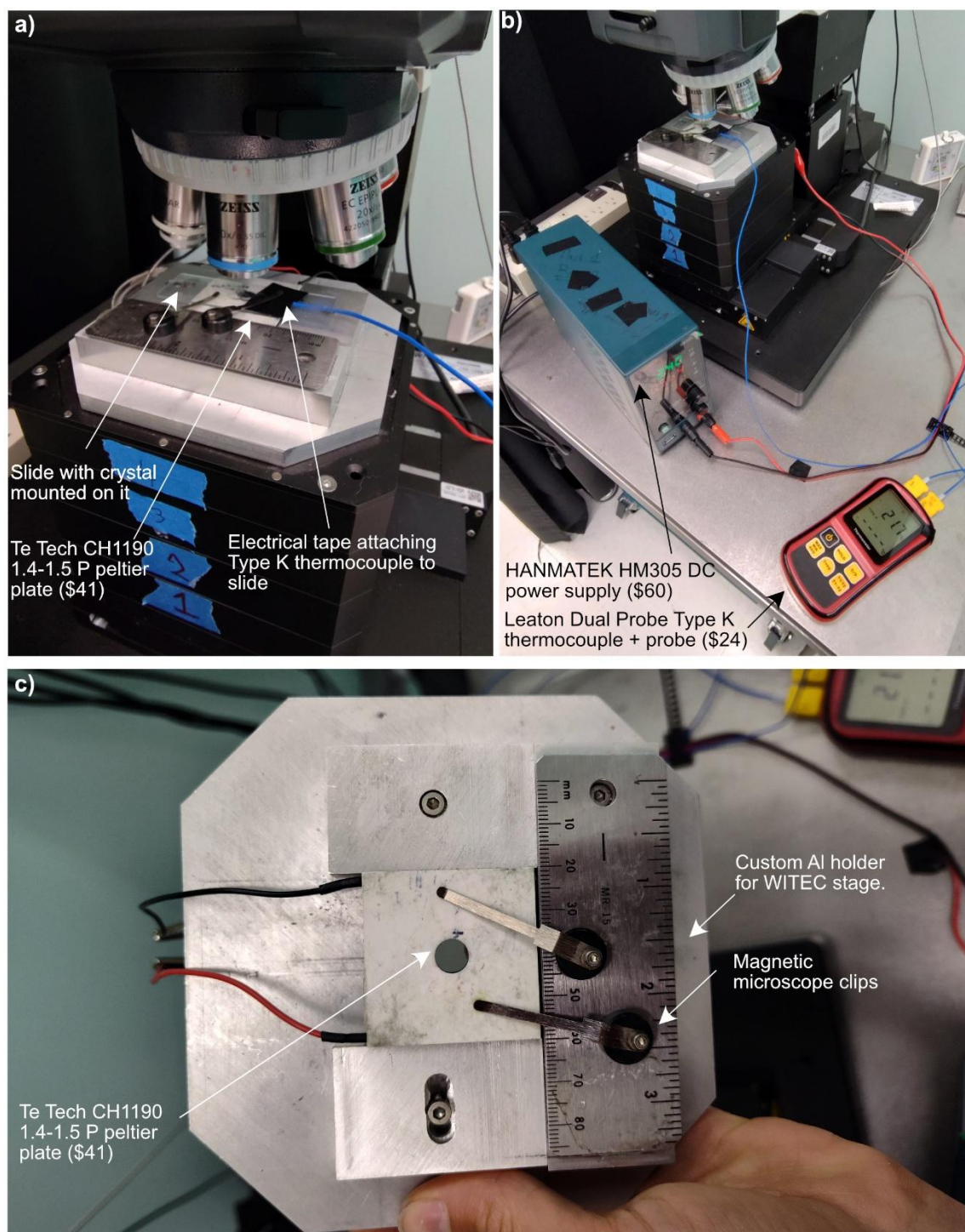
Supporting Figure 1 – Apparatus used for heating samples to 37℃. a) Picture showing a glass slide with a FI mounted in crystalbond under a 50X Zeiss objective. b) Picture showing the entire WITec system and the sample holder. A Type K thermocouple is taped on using electrical tape. This picture was taken while the slide was still warming up (hence the 21.7 ℃ temperature). c) Close up of the custom Aluminum holder that has the bottom machined out to slot onto the shape of the WITec base. There are 4 screws that attach to the magnets on the WITec stage (because Al is not magnetic). The ruler is glued to the top to provide a magnetic surface to attach a set of

magnetic microscope clips to help keep the sample in place. This holder was manufactured by our in-house machine shop (labour and material cost, $900), and can be ordered by contacting Wieser. This holder keeps the peltier plate and sample still, making it possible to analyze very small inclusions without drift (1-2 μm).
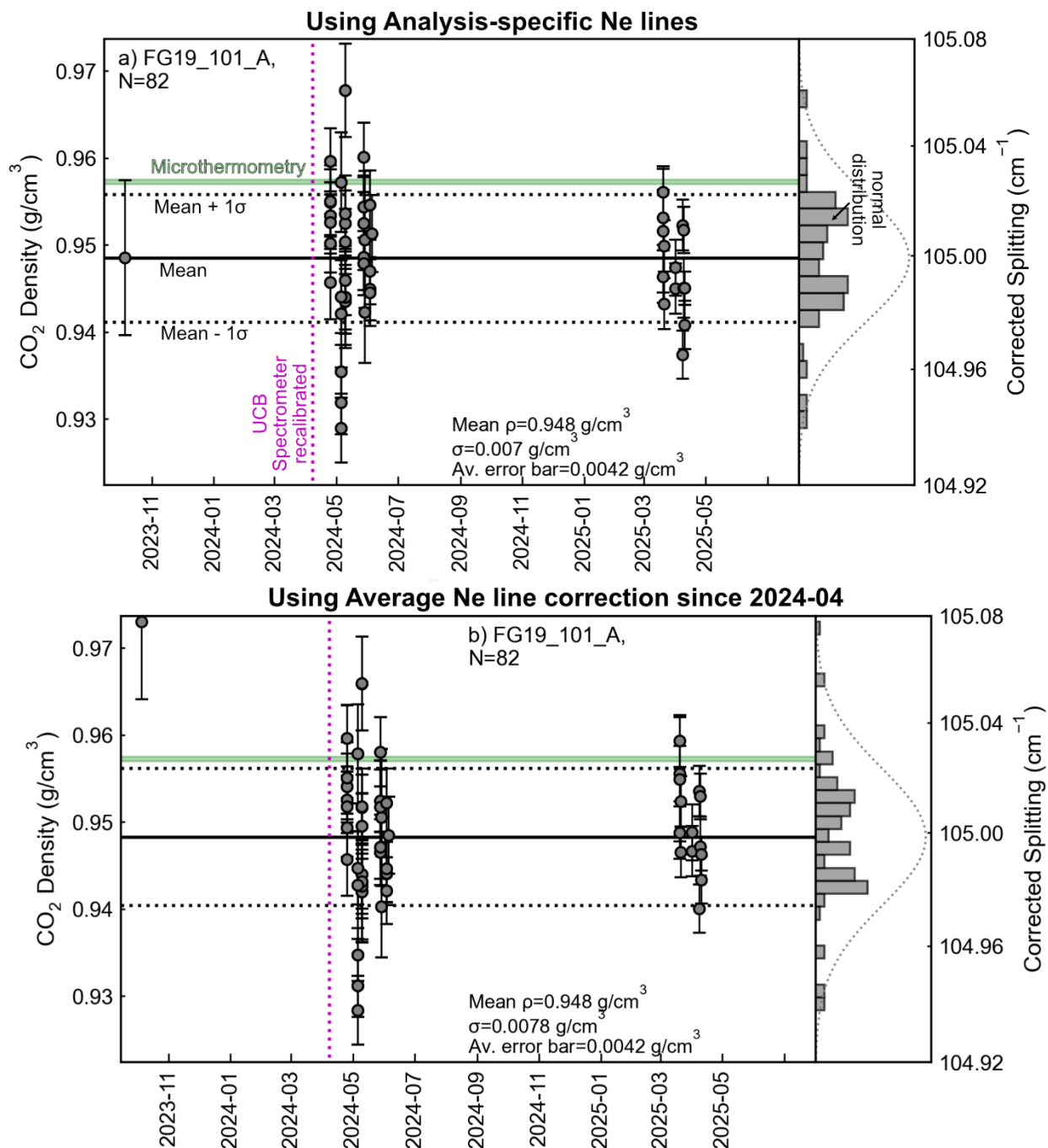
Supporting Figure 2 - Long-term variability in standard values collected at UCB. a) Data reduced using a drift correction model for each analytical session. b) Data reduced using the average Neon line correction factor for all standard acquisitions after the instrument recalibration (marked with a magenta line). The green bar shows the density determined by Microthermometry, accounting

for the uncertainty in the temperature reading of the stage. The black solid line shows the mean Raman value for the standard, and the dotted black lines ±1 σ. The distribution of densities is shown as a histogram, compared to a normal distribution based on the mean and standard deviation
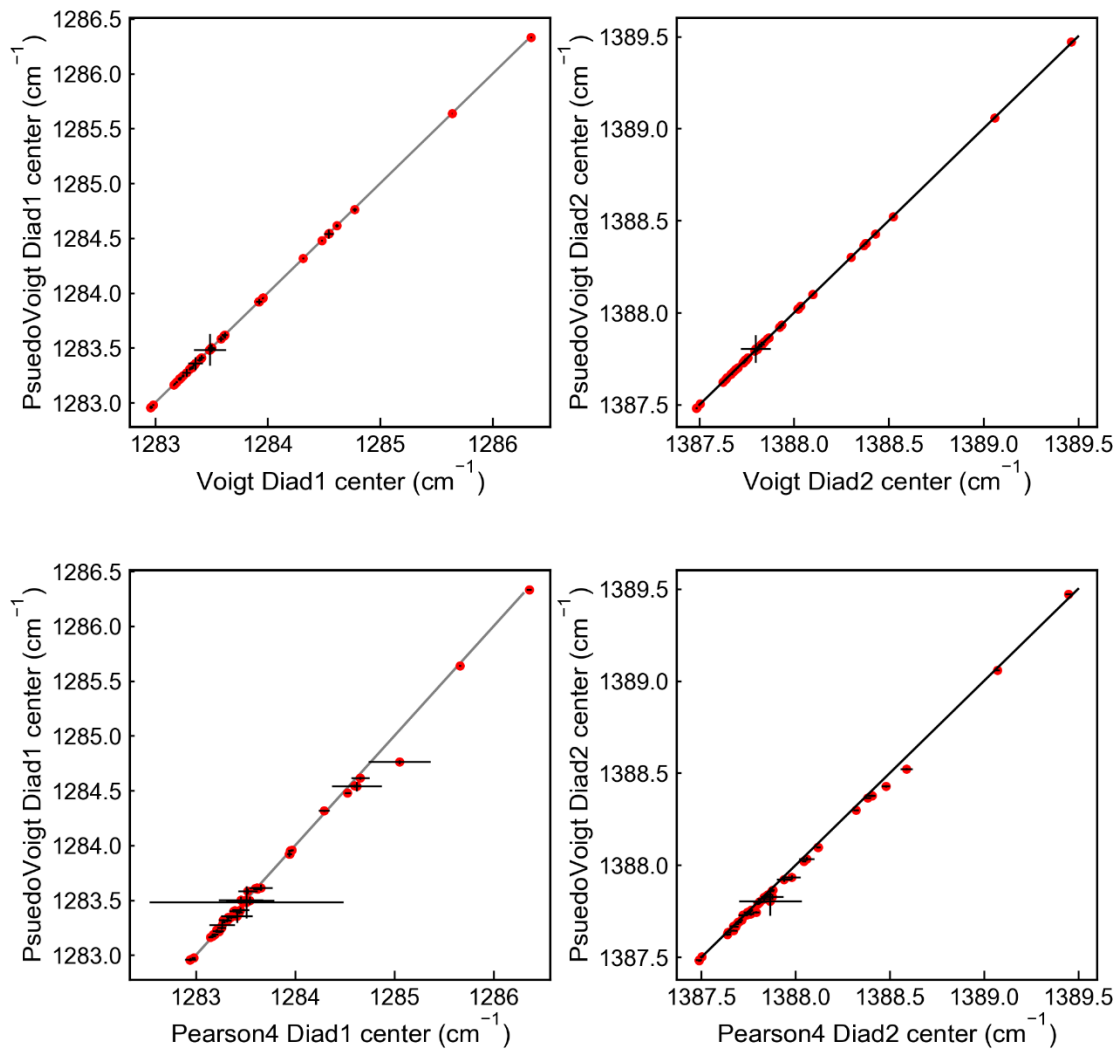
Supporting Figure 3 – Comparison of different PDF fits for Fogo fluid inclusions (as for Fig 8 in the main text).
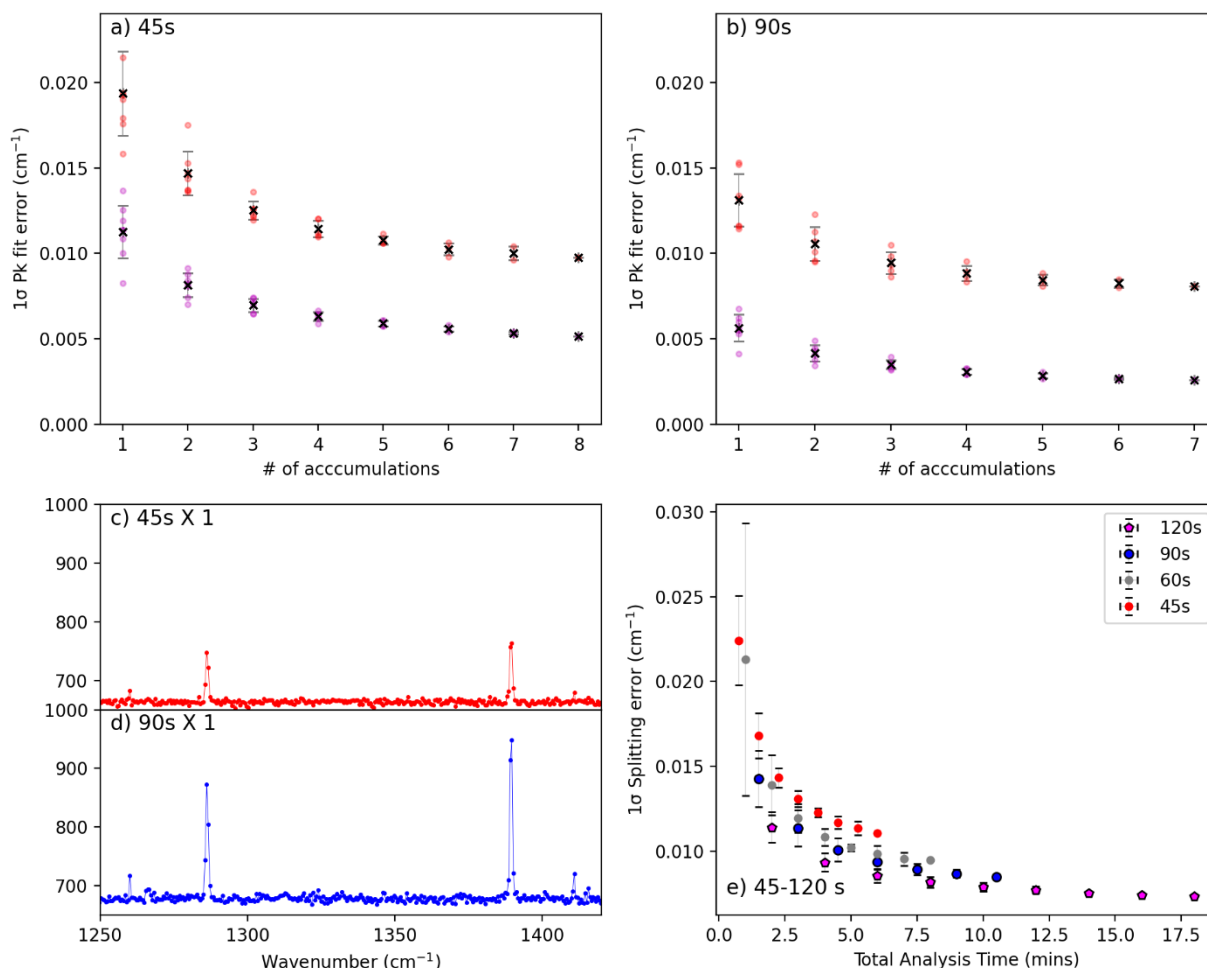
Supporting Figure 4 – Investigating peak fitting error as a function of integration time and accumulations using a melt inclusion vapour bubble from Mauna Loa (~0.01 g/cm^3, (Bearden et al., 2025). a-b) Error on the peak position of diad1 and diad2 for 45 and 90s acquisitions. The x axis shows the number of spectra that were averaged prior to fitting the peaks (analogous to accumulation). The dots show an individual fitted spectra (e.g. spectra 1 to 14 for accumulations =1, average of spectra 1-2, 2-3 3-4 for n=2 accumulations etc.), and the cross and error bar show the mean and standard deviation of averaged spectra for each number of accumulations (e.g. 14 individually fitted spectra for n=1 accumulations). c-d) Example spectra for 45s x 1 and 90s x 1. e) Error on calculated splitting after combining different numbers of individual acquisitions to give different acquisition times. For example, 5 different 120s acquisitions were averaged for the 10 minute total acquisition time.
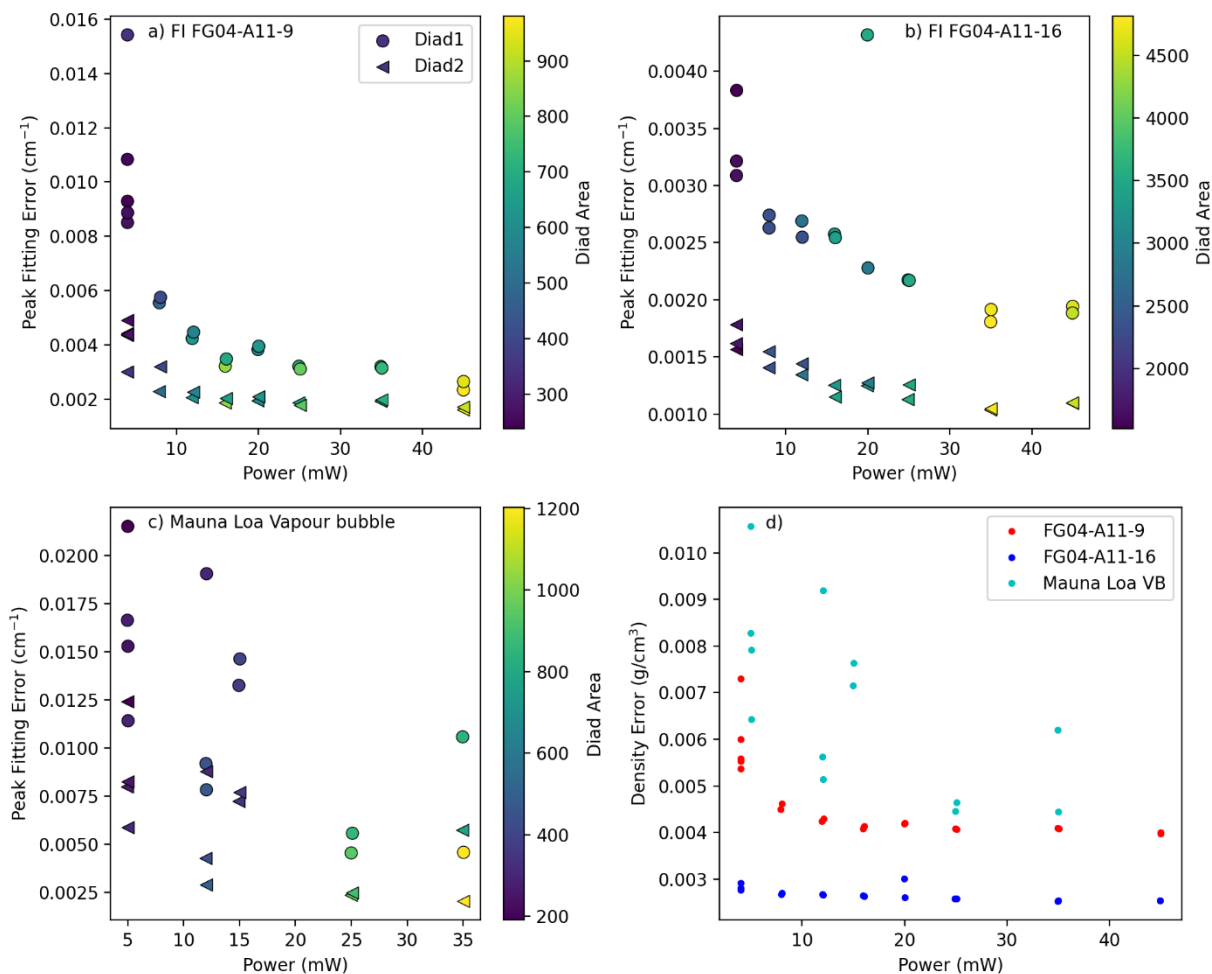
Supporting Figure 5– Change in peak fitting errors with increasing laser power. a-b) There is a clear drop in the peak fitting error for diad1 and diad2 as laser power is increased for Fogo fluid inclusion (FG04-A11-9), and for a melt inclusion vapour bubble from Mauna Loa (c). d) Same datasets but showing the error in calculated density.

a) B30_c1_power
Slope = -0.00028 ± 0.00004

b) B23_c3_power
Slope = -0.00028 ± 0.00004

c) 11i_c2_power
Slope = -0.00034 ± 0.00009

d) FG04-13-1_power
Slope = -0.00044 ± 0.00005
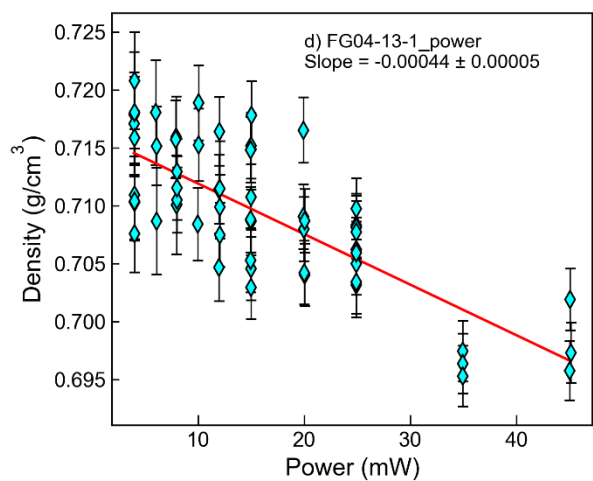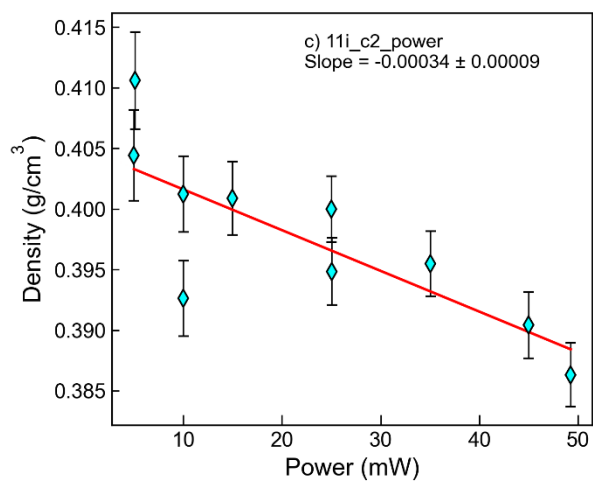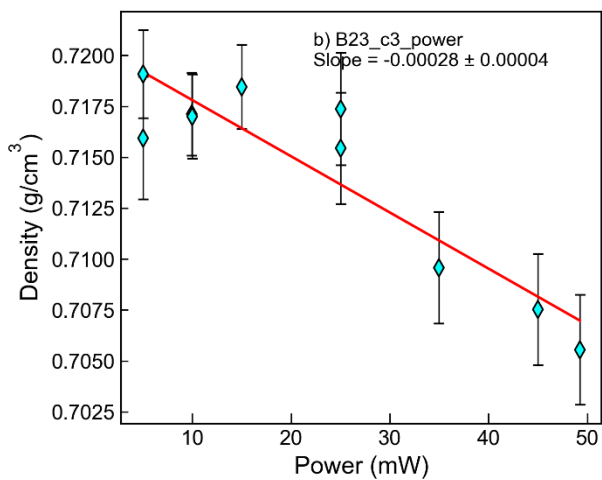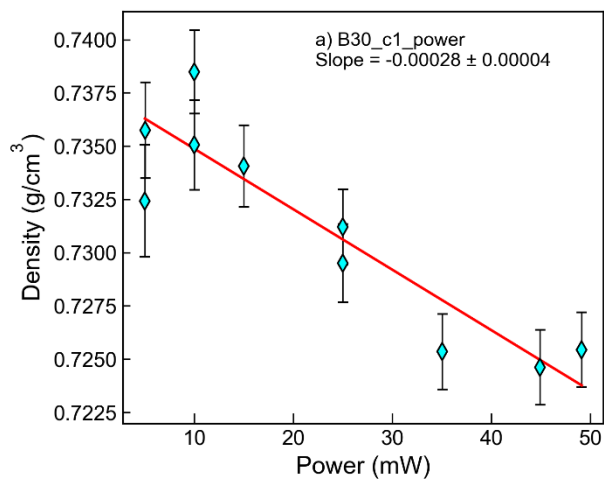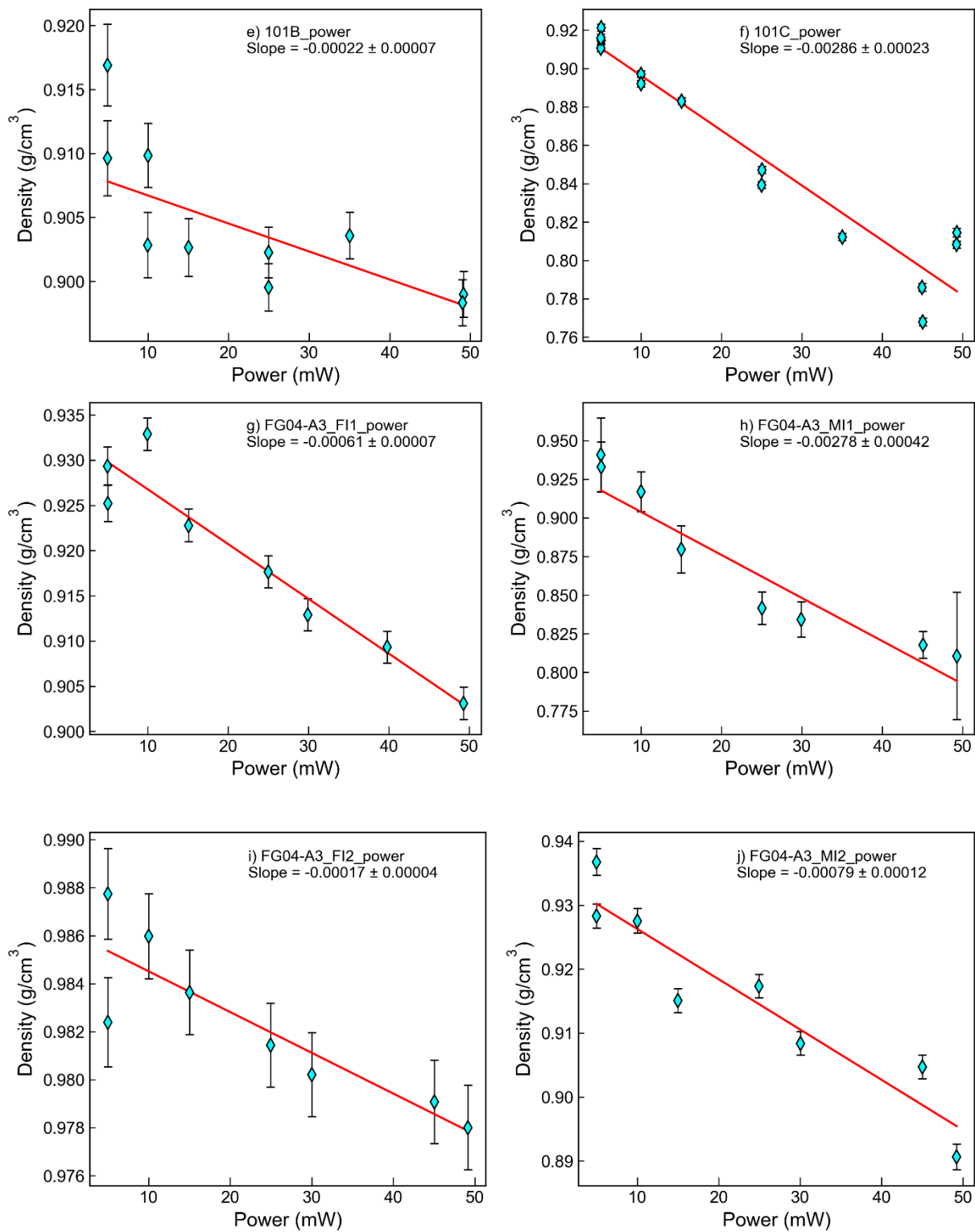
Supporting Figure 6 – Power series for 10 FI/MI, 4 of which are shown in Fig. 13 of the main text. The gradient and error on the slope is shown.
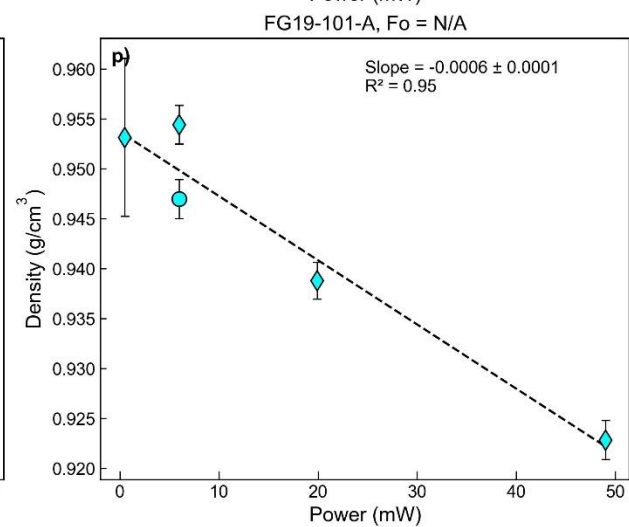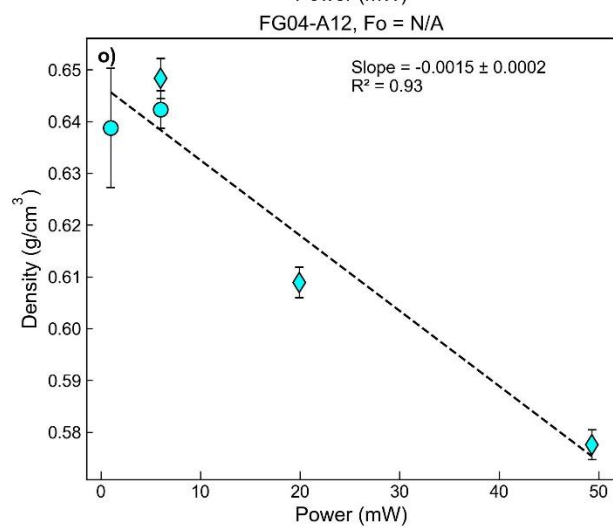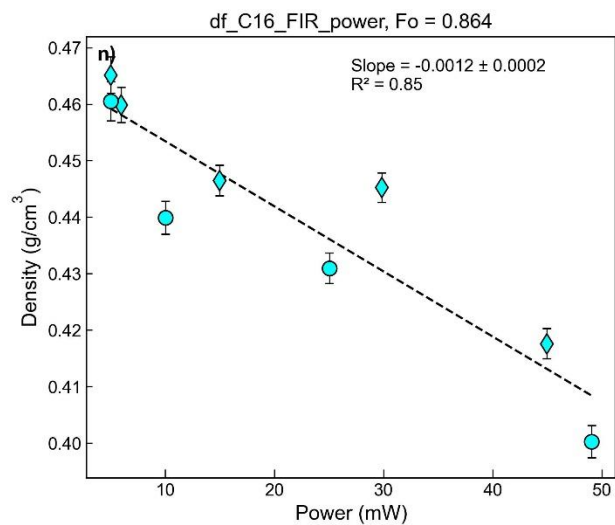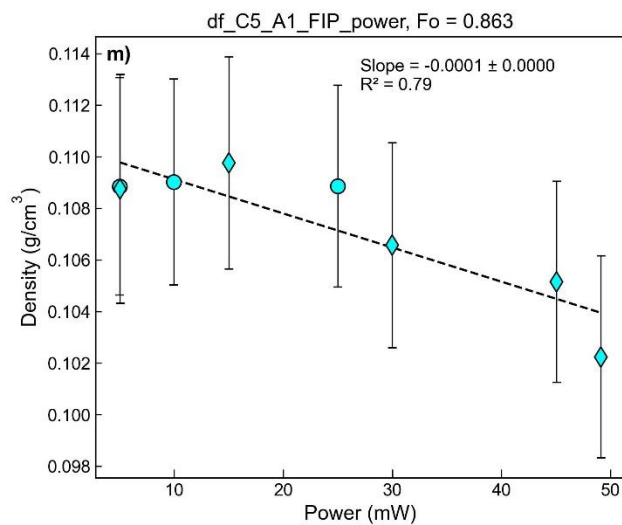
a) B30_c1_heat
Slope = -0.001 ± 0.000

b) B23_c3_heat
Slope = -0.001 ± 0.000

c) 11i_c2_heat
Slope = -0.001 ± 0.000

d) FG04-13-1_heat
Slope = -0.001 ± 0.000
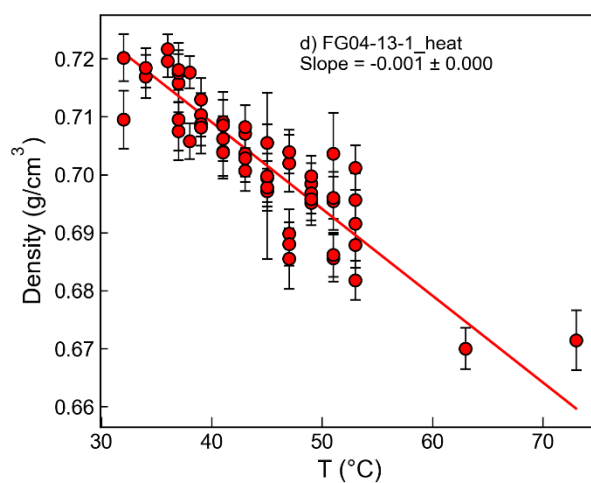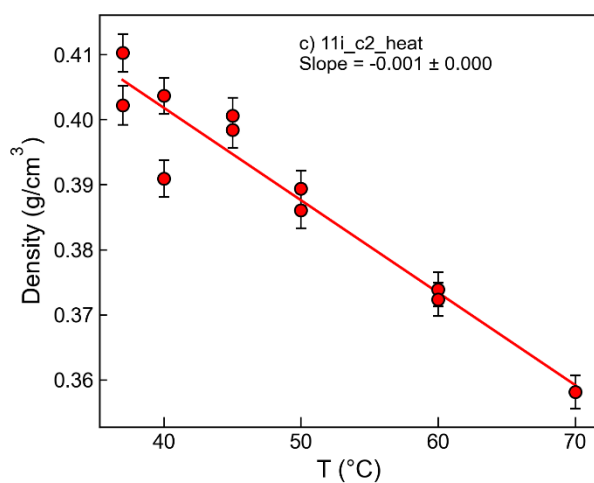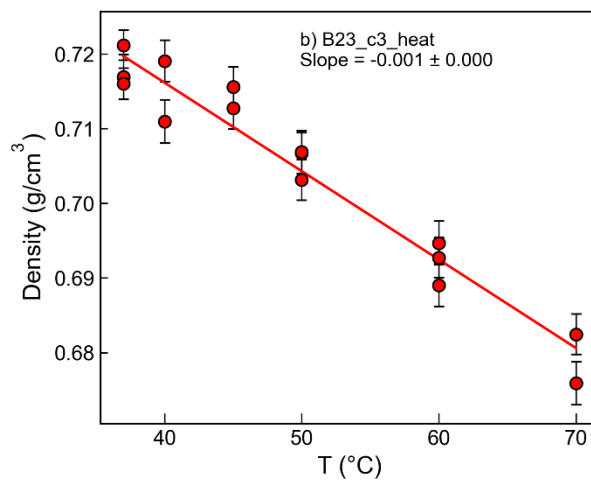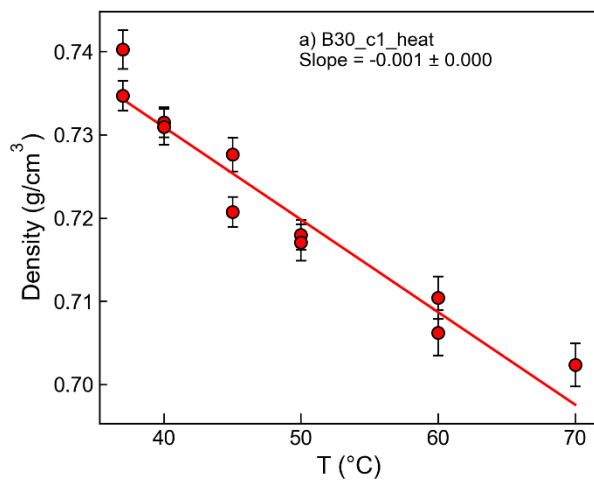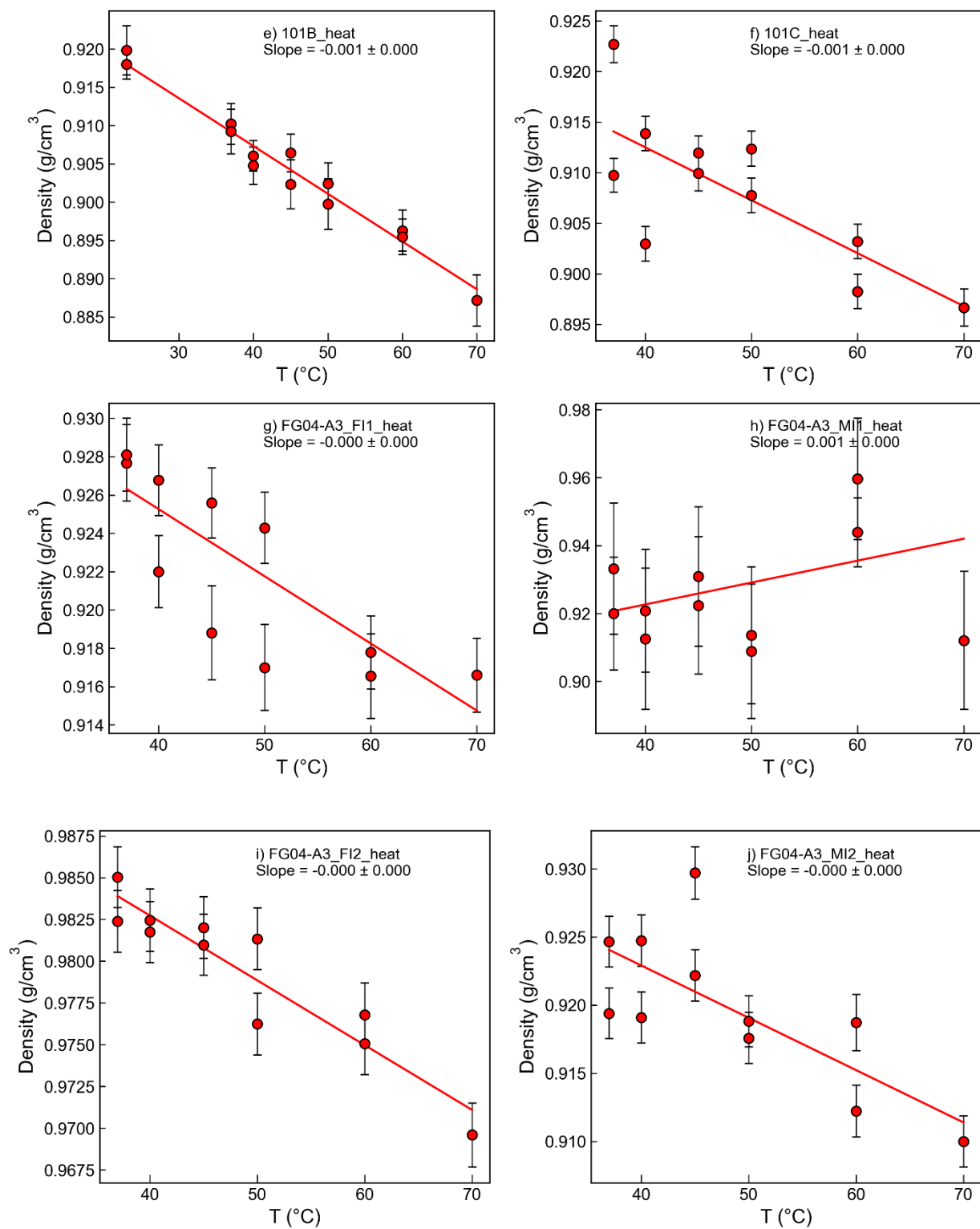
Supporting Figure 7 – Heating series for 10 FI/MI, 4 of which are shown in Fig. 13 of the main text. The gradient and error on the slope is shown.

a) B22_1, Fo = 0.740
Slope = -0.0016 ± 0.0001
R² = 0.98

b) B31_c8, Fo = 0.841
Slope = -0.0004 ± 0.0001
R² = 0.69

c) B27_c4, Fo = 0.844
Slope = -0.0002 ± 0.0000
R² = 0.73

d) B27_c19, Fo = 0.755
Slope = -0.0011 ± 0.0002
R² = 0.56
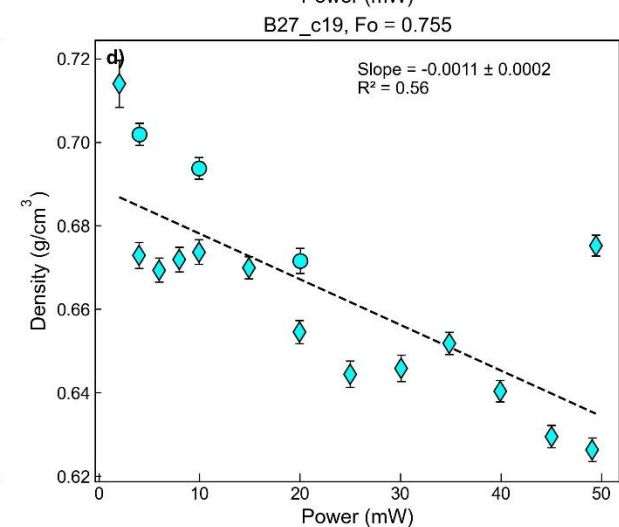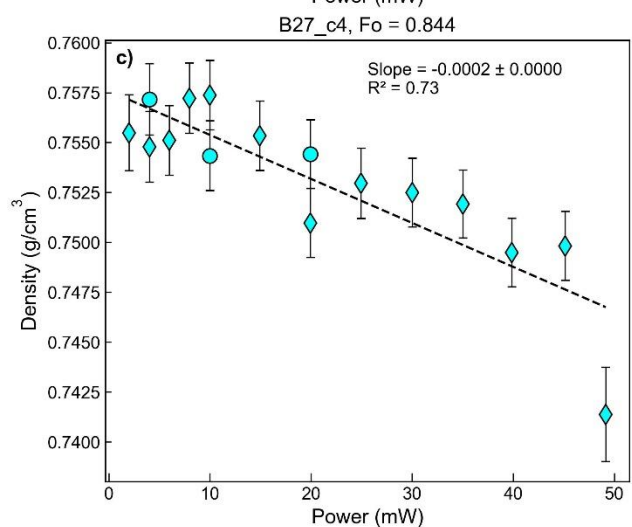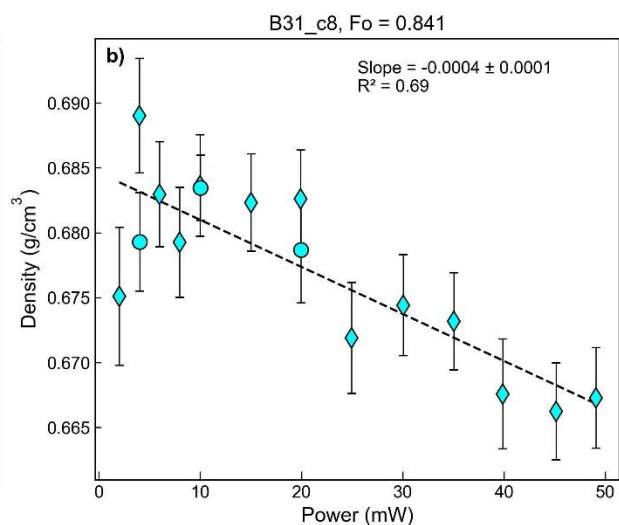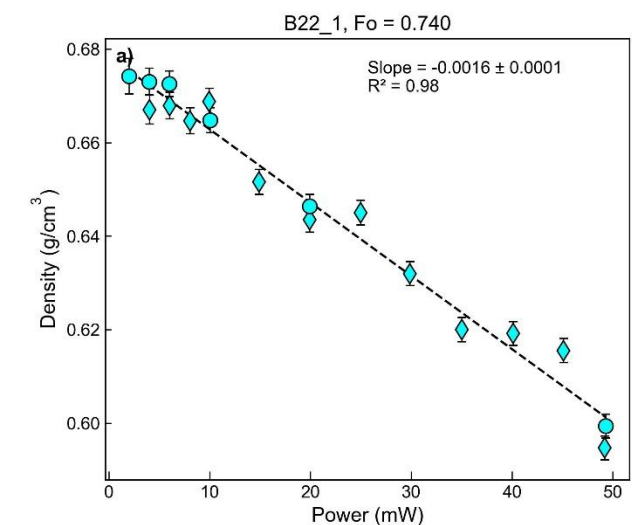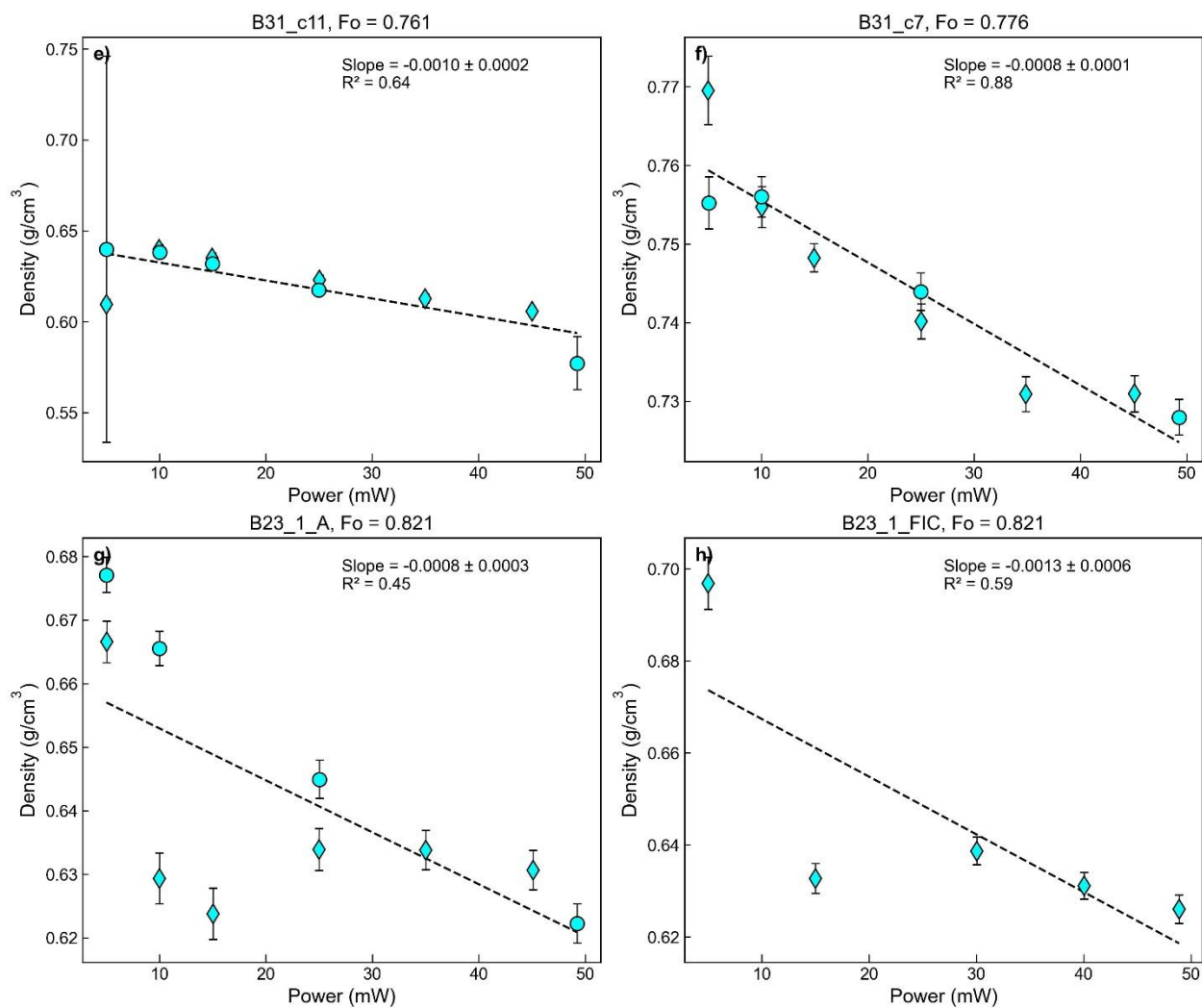
Supporting Figure 8 – Power series for Icelandic fluid inclusions with known Fo contents. The gradient, Fo content, and R^2 value is shown.

a) ML22_77_power, Fo = N/A

Slope = -0.0005 ± 0.0001
R² = 0.80

b) MP2_204_power, Fo = N/A

Slope = -0.0004 ± 0.0001
R² = 0.70

c) ML22_77_4_MIA_power, Fo = N/A

Slope = -0.0004 ± 0.0002
R² = 0.32

d) FG04-A11-9_power, Fo = N/A

Slope = -0.0004 ± 0.0001
R² = 0.56

e) FG04-A11-16_power, Fo = N/A

Slope = -0.0003 ± 0.0001
R² = 0.70

f) FG04-A11-4_power, Fo = N/A

Slope = -0.0027 ± 0.0003
R² = 0.87

g) FG19_101_FIB_power, Fo = N/A

Slope = -0.0001 ± 0.0000
R² = 0.07

h) Fo1_power, Fo = N/A

Slope = -0.0007 ± 0.0001
R² = 0.78

m) df_C5_A1_FIP_power, Fo = 0.863

Slope = -0.0001 ± 0.0000
R² = 0.79

n) df_C16_FIR_power, Fo = 0.864

Slope = -0.0012 ± 0.0002
R² = 0.85

o) FG04-A12, Fo = N/A

Slope = -0.0015 ± 0.0002
R² = 0.93

p) FG19-101-A, Fo = N/A

Slope = -0.0006 ± 0.0001
R² = 0.95

Supporting Figure 9 – Power series for all other fluid inclusions, and melt inclusions (a, b, c, h) shown in Fig. 12a of the main text.
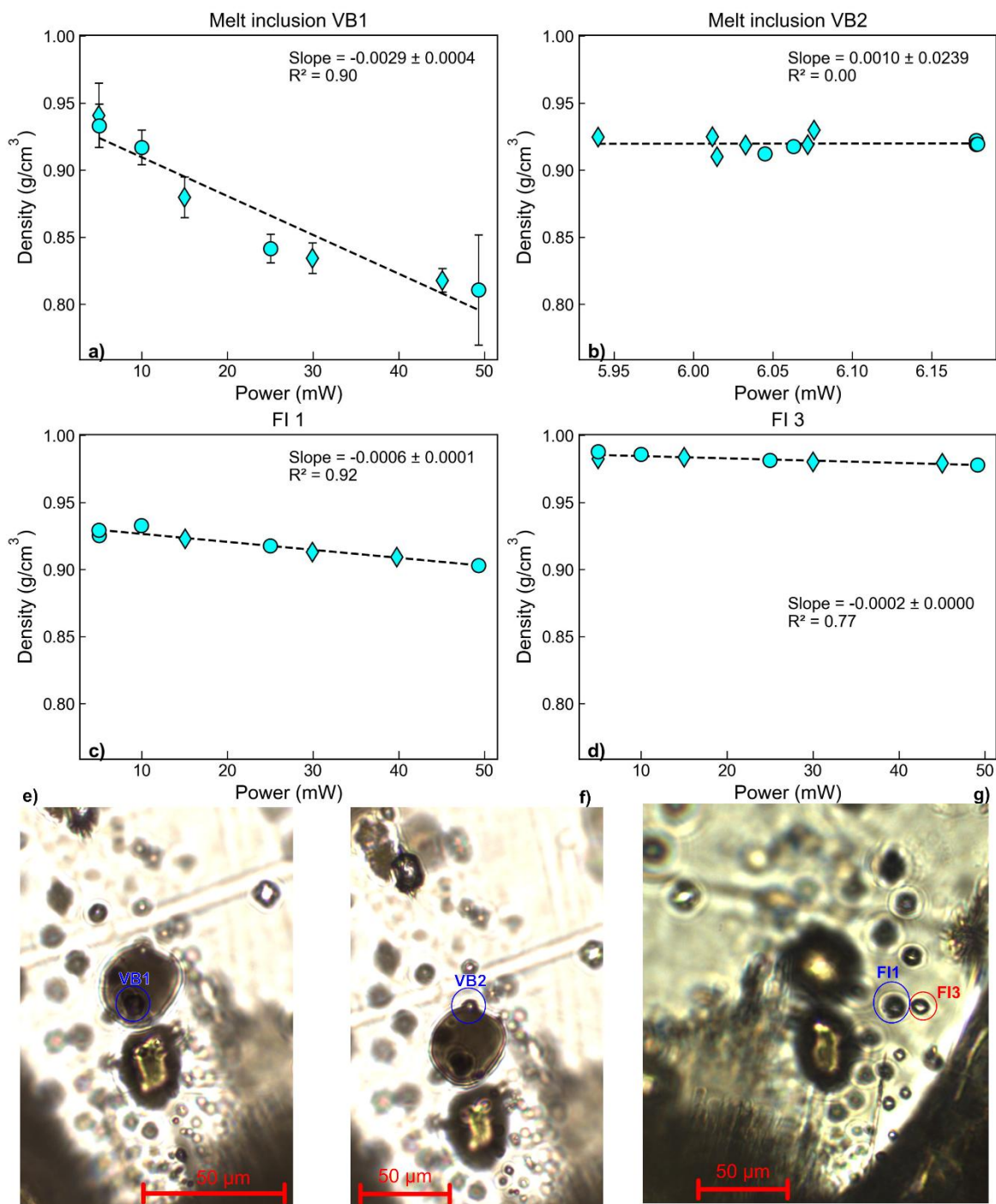
Supporting Figure 10 – Comparison of change in density as a function of laser power for two melt inclusion vapour bubbles and two fluid inclusions within the same crystal. VB1 is surrounded by melt (e) while VB2 is partially touching melt and partially touching olivine (f, see Supporting Fig. 11). The fact that VB2 is surrounded by olivine and melt accounts for its lower heating rate compared to VB1. It is noteworthy that the two FI, despite being in close proximity, show gradients that differ by a factor of 3.
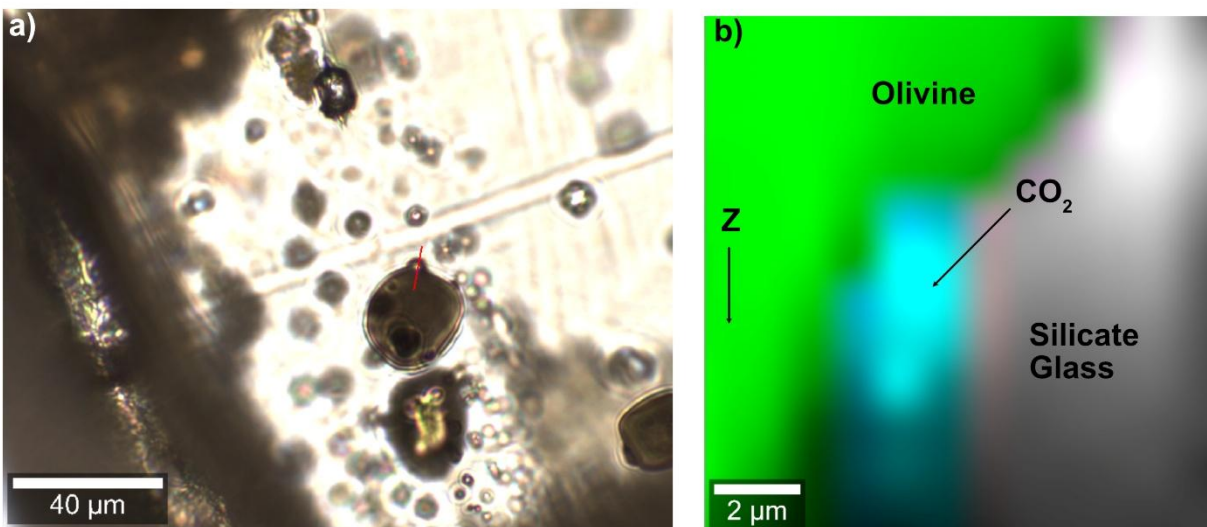
Supporting Figure 11 – a) Transmitted light image of Bub2, with the red line showing the location of a vertical Raman map. b) Processed Raman map, showing that the bubble is partially touching the silicate glass of the melt inclusion (grey colors), partially touching the olivine (green). Map processed using the WITec software.
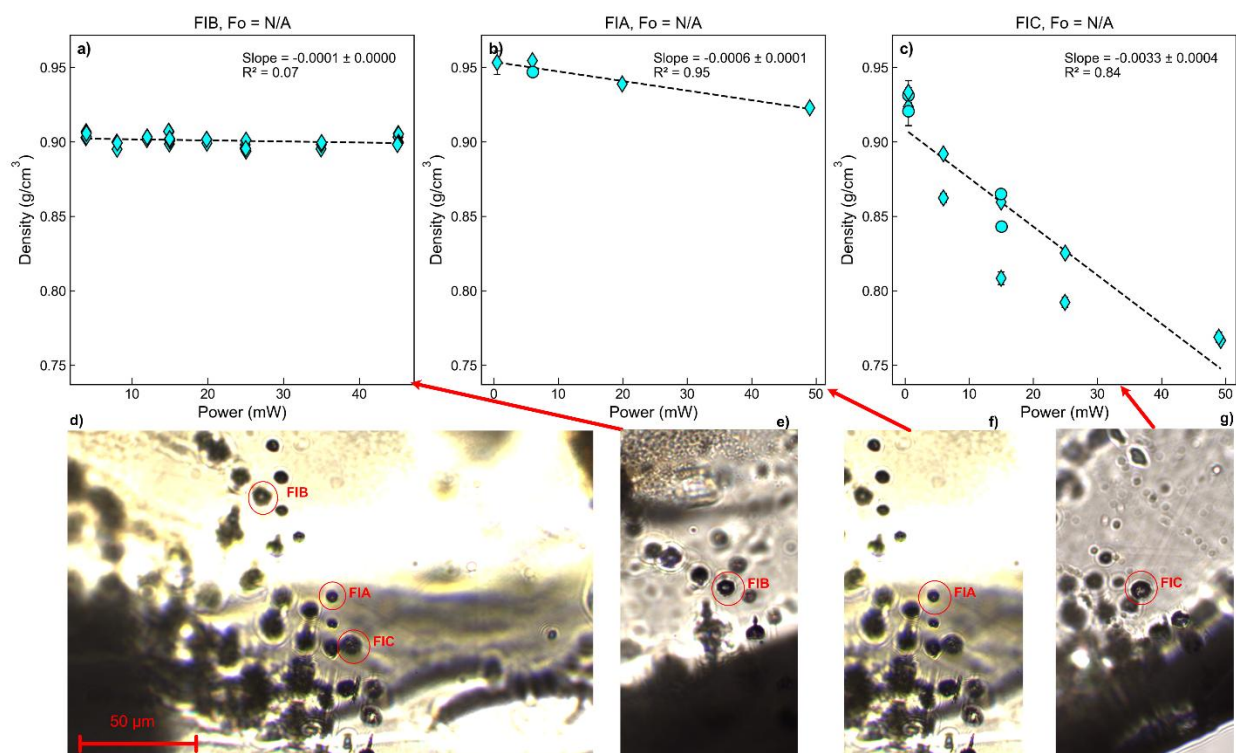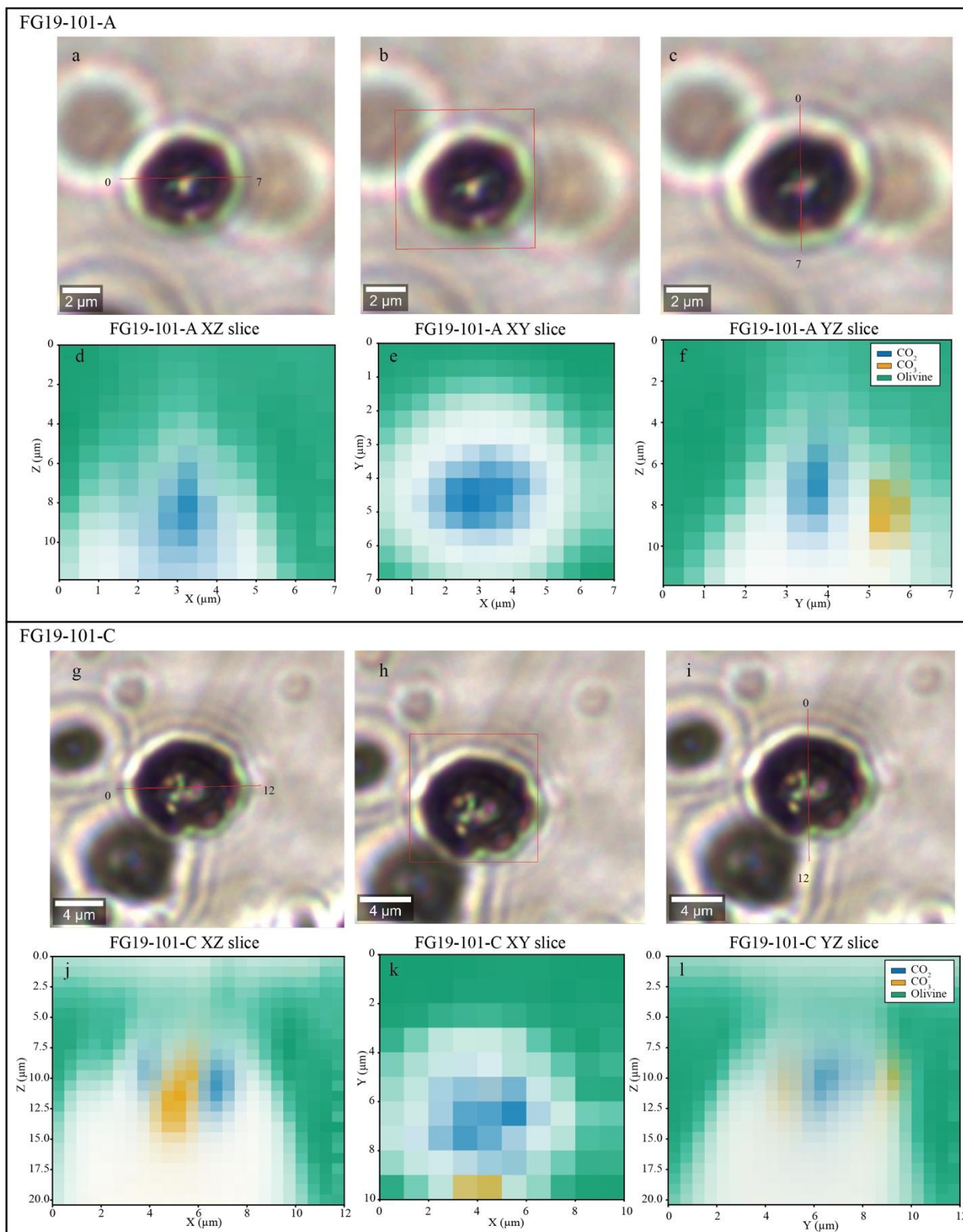

Supporting Figure 12. Comparison of change in density with increasing laser power (a-c) for 3 FI hosted within a single olivine crystal (FG19_101). d) Transmittted light image of the fluid

inclusion trail and the location of each FI. e-g) shows a focused picture of each FI. The scale is the same for all pictures. Raman maps of A and C are shown in Supporting Figure 14.
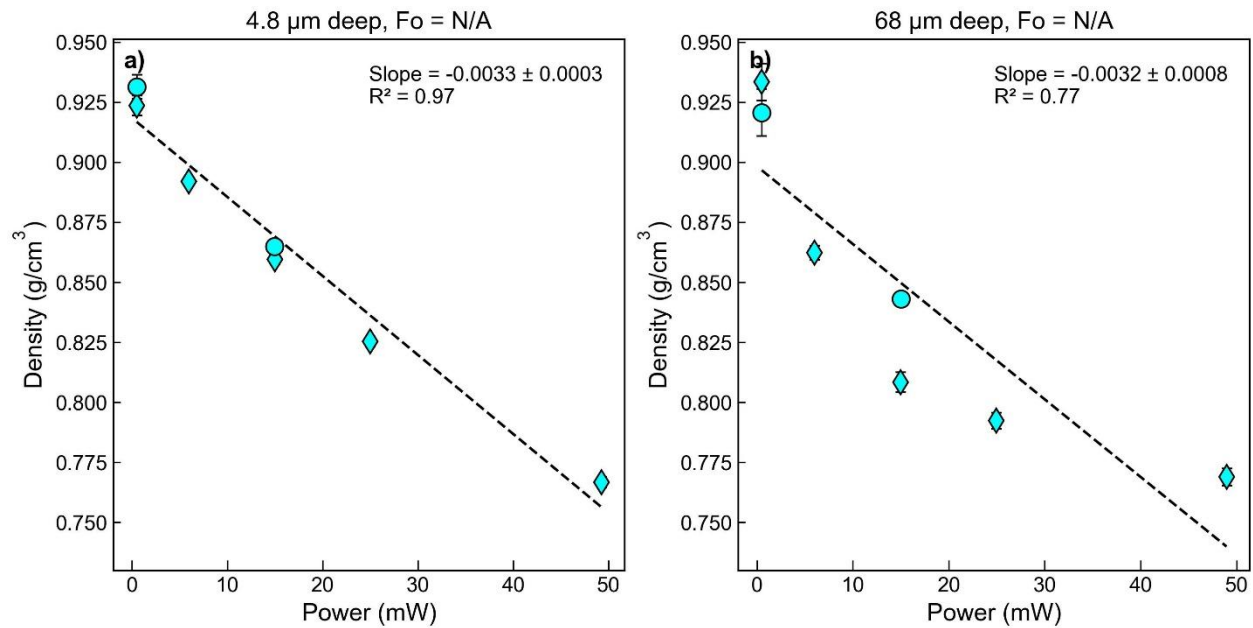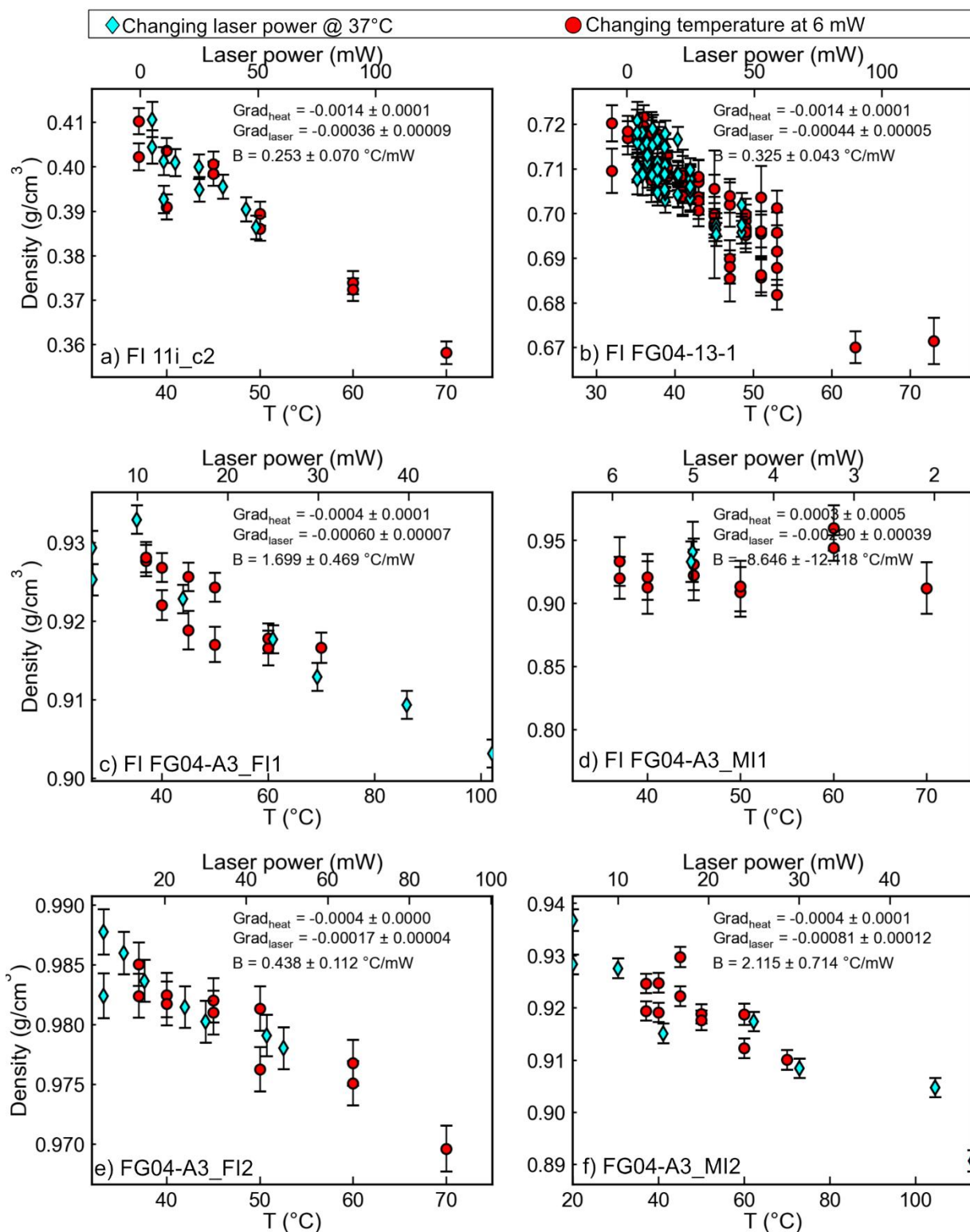

Supporting Figure 13 – Comparison of change in density with increasing laser power for FI FG19-101C at two different depths below the surface a) Collected when the FI was ~4.8 μm below the surface. The crystal was then flipped over, and a second power series conducted (The FI was ~68 μm below the top surface). The gradients of these series are within error of each other. This supports the modelling work of Hagiwara et al. (2021) suggesting that depth in the sample doesn't strongly affect heating rate.

FG19-101-A

a

b

c

FG19-101-A XZ slice   FG19-101-A XY slice   FG19-101-A YZ slice

d   e   f

CO$_2$
CO$_3^{2-}$
Olivine

FG19-101-C

g

h

i

FG19-101-C XZ slice   FG19-101-C XY slice   FG19-101-C YZ slice

j   k   l

CO$_2$
CO$_3^{2-}$
Olivine

Supporting Fig. 14. Two dimensional Raman maps of FIA and FIc show the presence of CO$_2$ and carbonate a-c) Transmitted light images of FG19-101-A. The red lines and rectangles show the plane along which Raman mapping was conducted (shown in panels d-f). g-i) Transmitted

light images of FG19-101-A and corresponding 2D slice (red line) . It is evident that carbonate is much more abundant in 101-C than 101-A, though they are only ~20 um apart from each other (Supporting Fig. 11). Both sets of Raman maps were collected using 50X objective, with a laser power of 49 mW, an integration time of 0.5 s, a 300 grooves/mm grating, and a 2350 cm-1 spectral center. Z dimension was recalculated accounting for the refraction index of olivine (1.7). In the Z slices, the bottom under the inclusions is white (very low intensity) due to loss of confocality at depth. Other white pixels are areas of highly mixed spectra which are difficult to reliably extract components from. Maps were processed using Ramanspy 0.2.10 (Georgiev et al., 2024). To produce the maps, we first denoise using savitzky golay filter with a window length of 4, and a polynomial of 3rd order, then apply ASLS baseline routine in Ramanspy and normalize each spectrum to using Min-Max normalization. We apply the whole preprocessing routine to all the maps. We extract "pure" endmember spectra for CO_2, CO_3 and Olivine from the map spectra and use Non-negative Constrained Least Squares (NNLS) abundance method in RamanSpy to extract abundance maps for each component at every pixel. These maps are not quantitative, but an approximation of the distribution of each component
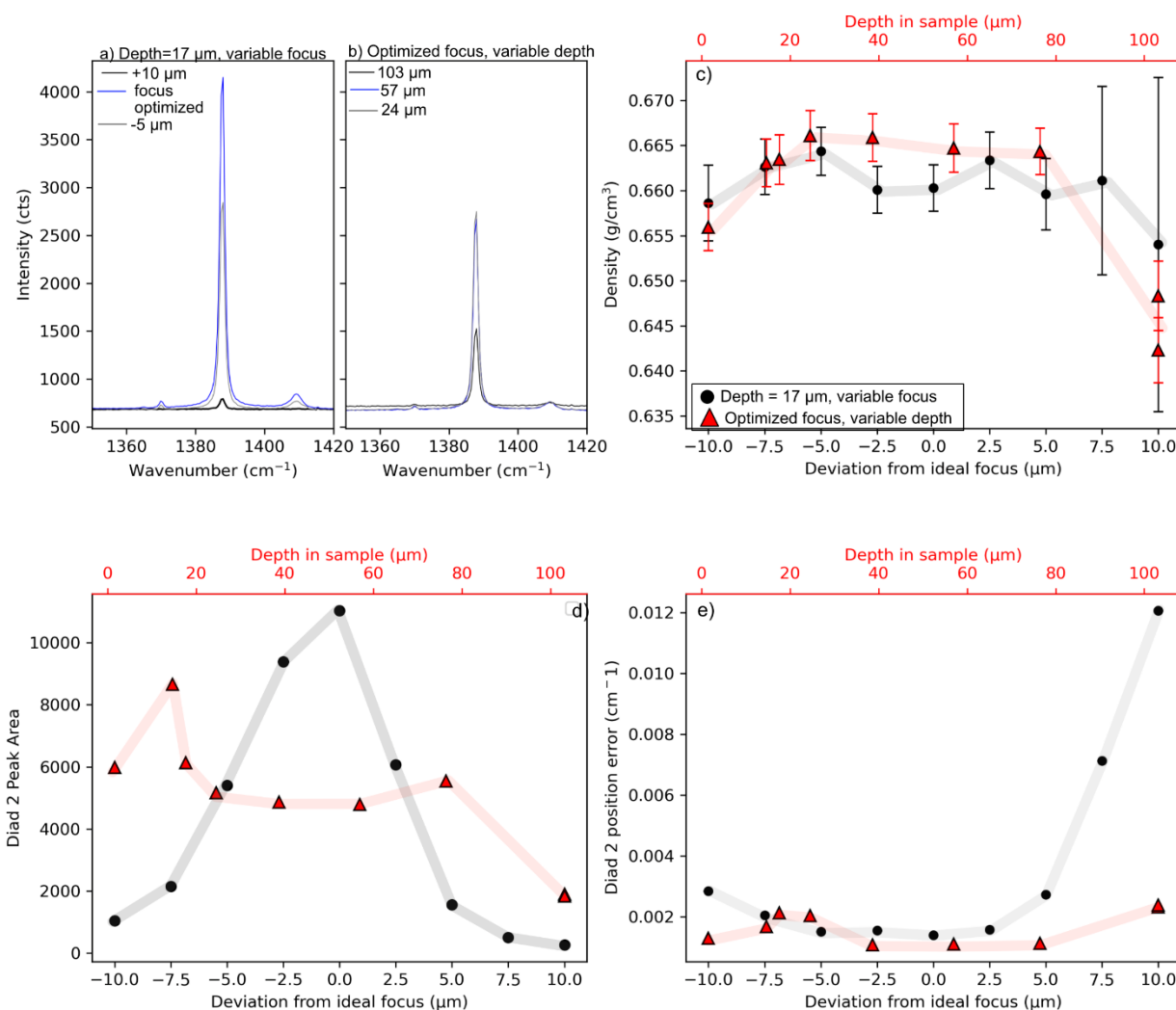
Supporting Figure 15 – Comparison of heating and power series for additional samples to those shown in Fig. 13 of the main text.
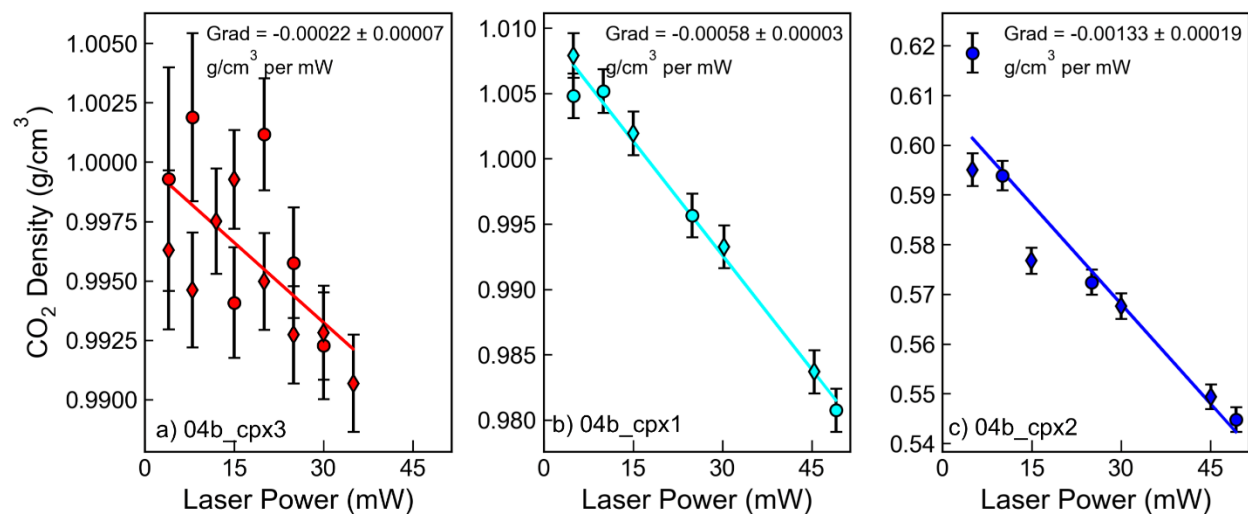
Supporting Figure 16 – Laser power series for ~Mg$_{79}$# clinopyroxenes for fluid inclusions from Gleeson et al. (2025). The gradient varies by a factor of ~3.8, despite the fact these samples are very chemically homogenous and there are no obvious color differences between the different crystals.

Supporting Figure 17 – Comparison of peak intensity and peak fitting errors with increasing depth in sample vs. a change in focus for FG04_A12. a) Spectra centered around diad2 showing the difference in intensity for the optimized focus (blue) vs. -5 μm deeper in the sample (grey) and 10 μm shallower (black). The fluid inclusion has a radius of ~7 μm. b) Change in intensity with depth of the fluid inclusion optimal focus below the sample. At 24 μm and 57 μm depth, the intensities are very similar, although they drop to above half for 103 μm depth. c) Change in density as a function of depth and ideal focus. Red triangles show variation in depth in sample (top red axis) and black circles show deviation from the ideal focus (black). Apart from the very worst focus and deepest depth, there is no change outside of uncertainty. Points are joined up with lines to help separate the two series. d) Peak area of diad 2 (a measure of intensity) as a function of depth and focus. e) Error on the position of diad 2.
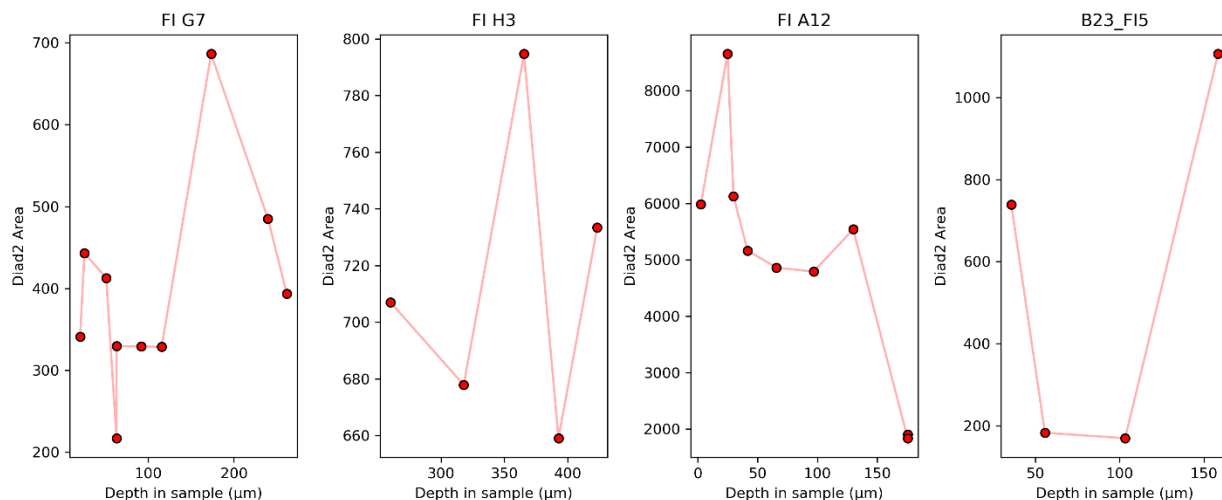
Supporting Figure 18 – Changes in signal intensity as a function of depth of the FI below the sample for 4 different olivine-hosted FI. FI H3 was lost during serial sectioning, hence the shorter range of depths.
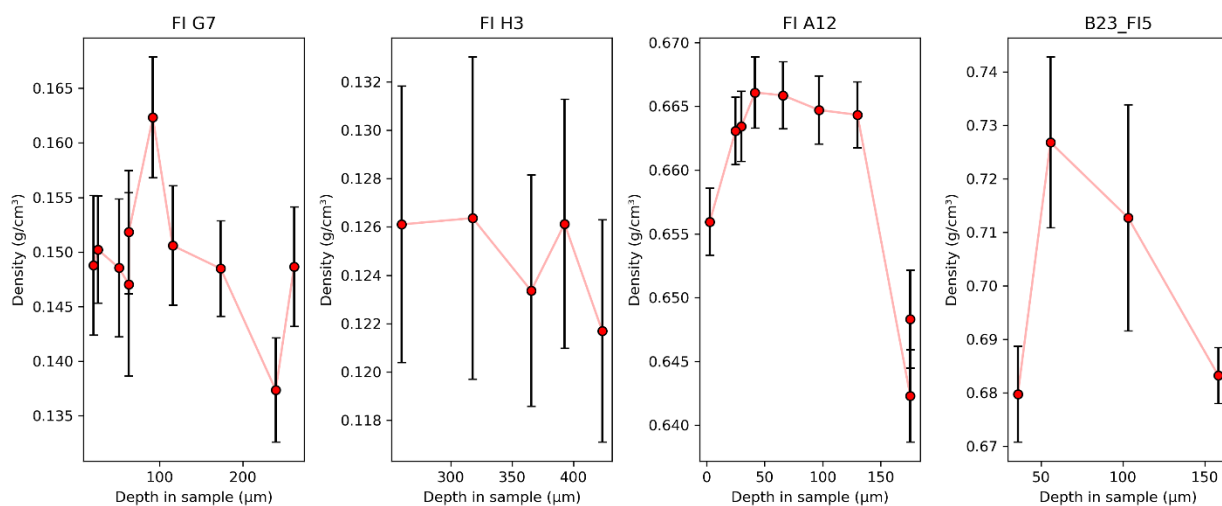

Supporting Figure 19- Density (and error) as a function of depth. The highly variable densities in B23_FI5 likely result from the very poor surface polish, resulting in backgrounds which were tricky to fit.
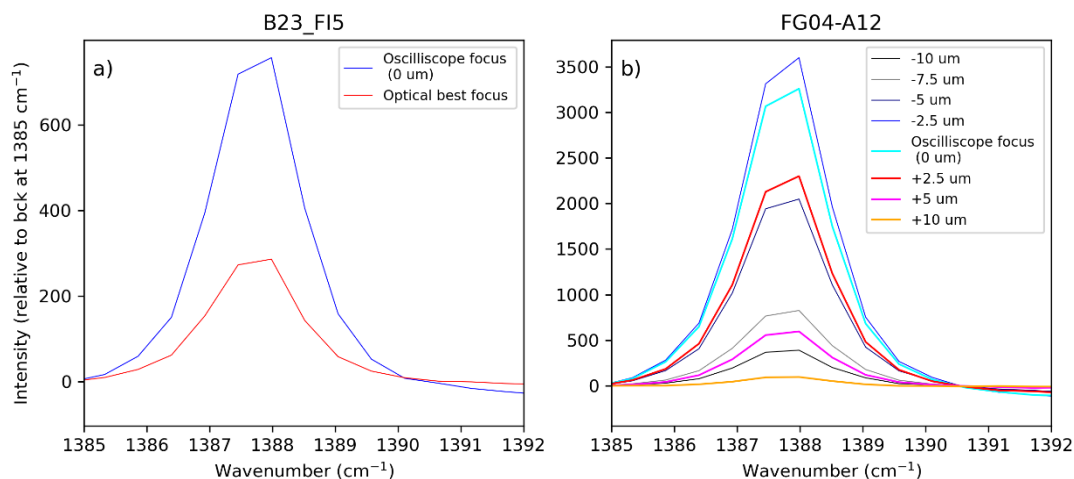
Supporting Figure 20 – Example of the effect of small focus changes on the intensity of diad 2. a) Compares the user-selected optical focus based on sharpness of the FI edges to the optimal focus position found by the oscilloscope. b) Compares the optimal focus position found by the oscilloscope to small perturbations from this position.
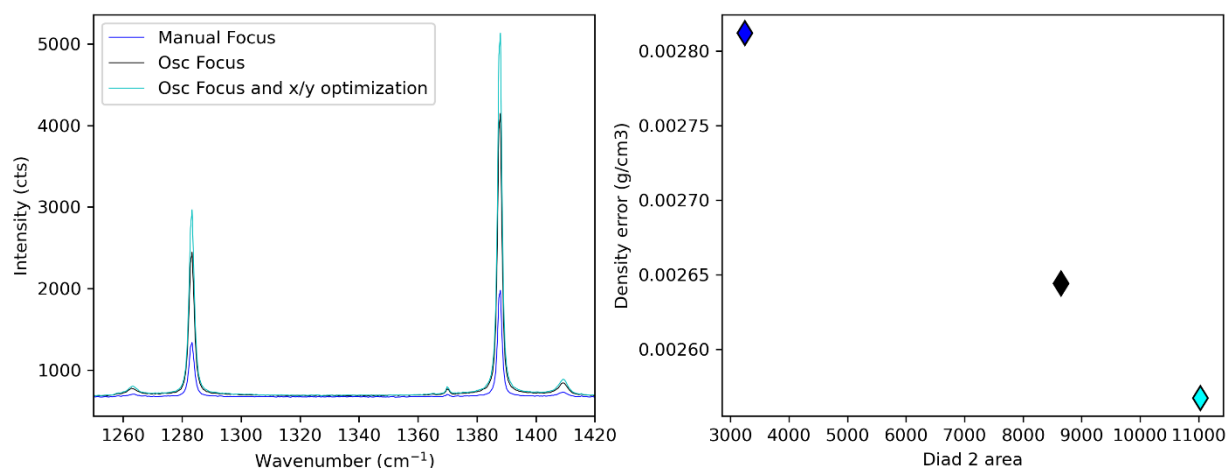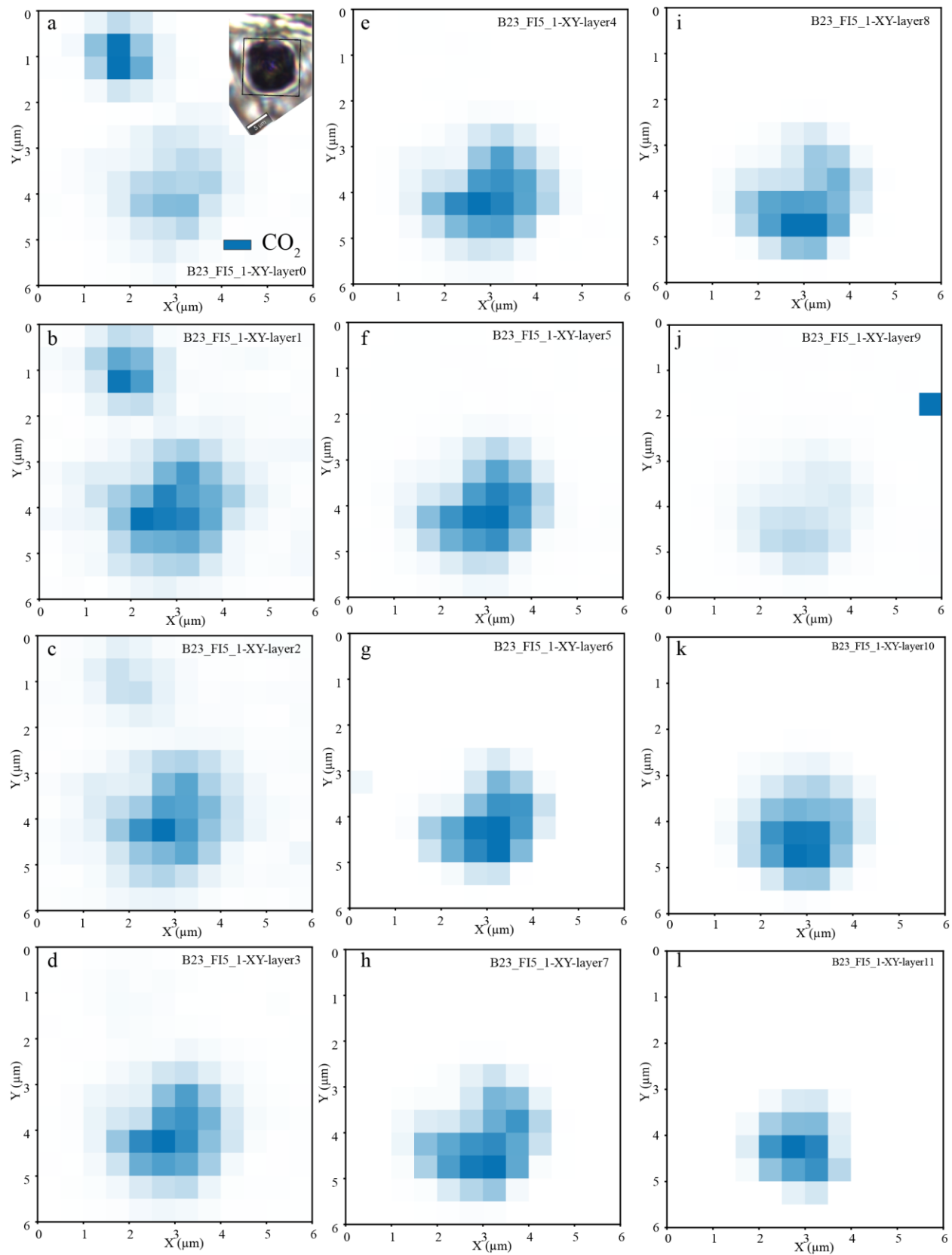

Supporting Fig. 21– Compares the user-selected optical focus based on sharpness of the FI edges to the optimal focus position found by the oscilloscope (dark blue), to the optimum Z position found by the oscilloscope (black), and the optimum Z and x-y position determined by Raman mapping for FG04-A12.
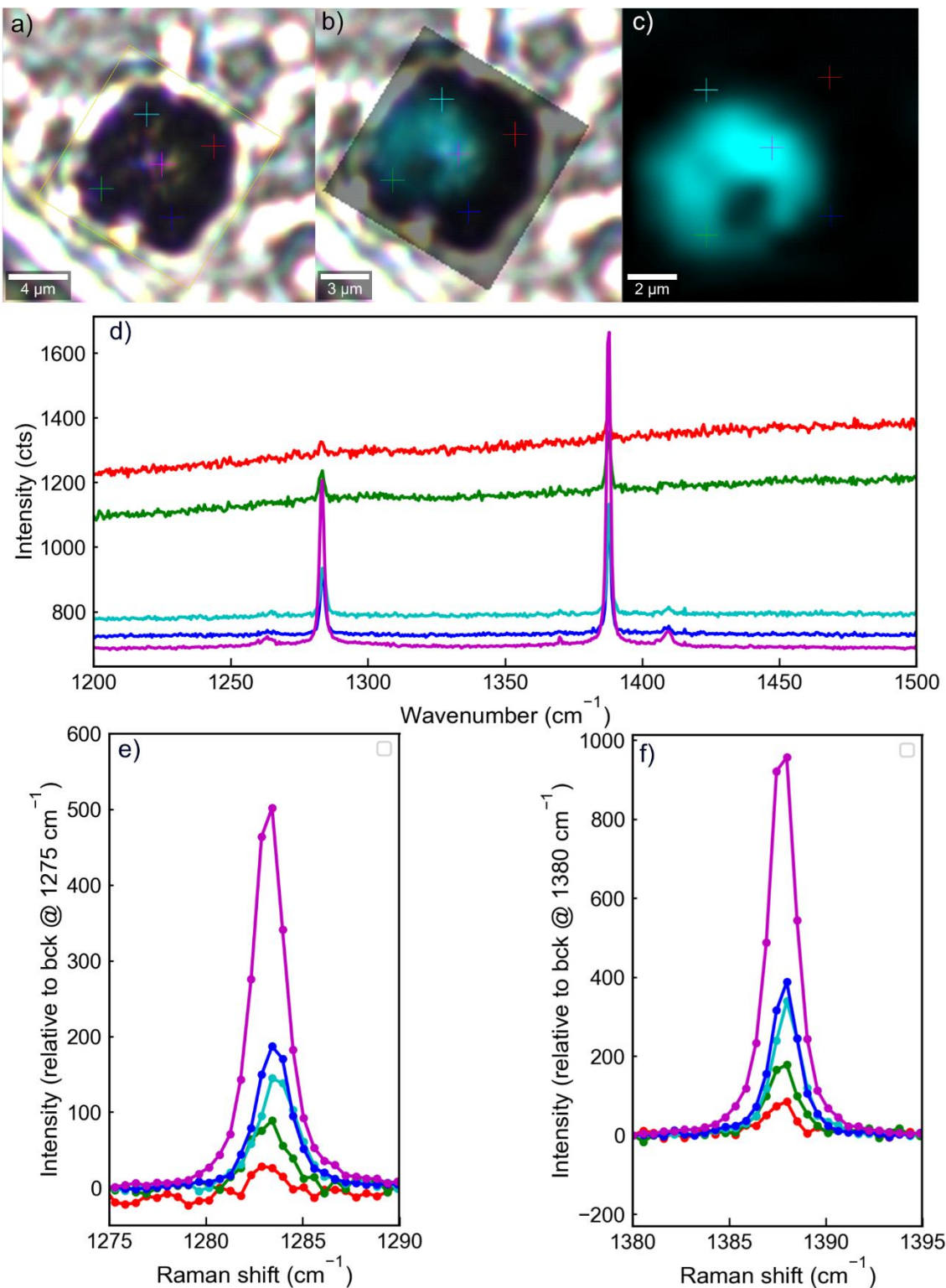
Supporting Figure 22 – Variation in signal intensity as a function of X-Y position within the fluid inclusion. a) Transmitted light image of fluid inclusion B23_FI5. The 5 colored crosses show the position of different spectral acquisitions shown in d-f. b) An 1800 g grating Raman map with the

intensity of the CO_2 signal shown as cyan colors overlain on this transmitted light image. c) shows the map with no underlay. d-f) Comparing the intensity of different acquisitions. While it is not surprising that the strongest signal is obtained from the center of the inclusion, it is surprising that despite the cyan, red, green and blue cross being equidistance from the center, they yield vastly different intensities.

a B23_FI5_1-XY-layer0

CO$_2$

b B23_FI5_1-XY-layer1

c B23_FI5_1-XY-layer2

d B23_FI5_1-XY-layer3

e B23_FI5_1-XY-layer4

f B23_FI5_1-XY-layer5

g B23_FI5_1-XY-layer6

h B23_FI5_1-XY-layer7

i B23_FI5_1-XY-layer8

j B23_FI5_1-XY-layer9

k B23_FI5_1-XY-layer10

l B23_FI5_1-XY-layer11

Supporting Figure 23 – Raman map of CO_2 abundance in B23_FI5_1. The inset in a) shows the geometry of the map relative to the fluid inclusion. Each X-Y graph shows a different Z slice through the inclusion, with the stronger blue colors showing stronger CO_2 intensities. It is clear that the x-y positoin of max intensity varies with depth. The map geometry in the WITec software was set as 12x12x12 um, so each slice was separated by 1 um of stage movement. With the refractive index of olivine being 1.7 and CO_2 being 1, there is some vertical distortion. Maps were collected using 50X objective, with a laser power of 49 mW, an integration time of 1 s, the 1800 grooves/mm grating (so olivine is not visible), and a 1325 cm-1 spectral center. Processing was performed using Ramanspy 0.2.10 (Georgiev et al., 2024). To produce the maps, we first denoise using savitzky golay filter with a window length of 4, and a polynomial of 3rd order, then apply ASLS baseline routine and normalize each spectrum to using Min-Max normalization. We apply the whole preprocessing routine to all the maps. We extract "pure" endmember spectra for CO_2 from the map spectra and use Non-negative Constrained Least Squares (NNLS) abundance method in RamanSpy to extract abundance maps for CO_2 at every pixel. These maps are not quantitative, but an approximation of the distribution of CO_2. Each individual map is normalized to the strongest pixel within that slice.

References

Bearden, A.T., Wieser, P.E., Rangel, B., DeVitre, C.L., Gleeson, M.L.M., Monteleone, B.D., Lynn, K.J., 2025. Volatile systematics and magma storage depths at Mauna Loa: a melt inclusion perspective., in: In Chapman Conference on Caldera-Forming Eruptions at Basaltic Volcanoes: Insights and Puzzles from Kīlauea and Beyond. AGU.

Georgiev, D., Pedersen, S.V., Xie, R., Fernández-Galiana, Á., Stevens, M.M., Barahona, M., 2024. RamanSPy: An Open-Source Python Package for Integrative Raman Spectroscopy Data Analysis. Anal. Chem. 96, 8492–8500. https://doi.org/10.1021/acs.analchem.4c00383

Gleeson, M., Wieser, P.E., DeVitre, C.L., Shi, S.C., Millet, M.-A., Muir, D.D., Stock, M.J., Lissenberg, J., 2025. Persistent High-Pressure Magma Storage beneath a Near-Ridge Ocean Island Volcano (Isla Floreana, Galápagos). Journal of Petrology 66, egaf031. https://doi.org/10.1093/petrology/egaf031

Hagiwara, Y., Yoshida, K., Yoneda, A., Torimoto, J., Yamamoto, J., 2021. Experimental variable effects on laser heating of inclusions during Raman spectroscopic analysis. Chemical Geology 559, 119928. https://doi.org/10.1016/j.chemgeo.2020.119928