

1 Evaluating and Optimizing Raman Analytical Protocols for CO₂ Fluids

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21 Abstract

22 Raman spectroscopy is a key method for determining CO₂ densities in geological fluids, yet
23 acquisition, densimeter calibration, and spectral processing methodologies vary widely between
24 laboratories. However, the precision of this method, and the broader applicability of densimeters
25 generated in a single lab, are still debated. This study describes a series of tests to determine
26 how different instrument and acquisition parameters affect CO₂ densities and errors on any given
27 Raman instrument, allowing users to determine analytical best practices for their specific
28 analytical set up. First, we discuss the effect of spectral non-linearity, demonstrating that shifts in
29 the selected acquisition window and the choice of Ne lines used for drift correction can generate
30 a diversity of densimeters even on a single instrument. However, our repeat measurements of
31 natural fluid inclusion (FI) standards over ~3 years yield a 1σ variation of <0.01 g/cm³, indicating
32 that when consistent methodologies are applied to calibration data and unknowns, Raman
33 analyses are highly reproducible over many years. We investigate the play off between peak fitting
34 error and parameters such as acquisition time, laser power, and sample positioning, rebutting
35 recent suggestions that Raman analyses are associated with very large peak-fitting errors. We
36 show that although high laser powers greatly improve signal:noise ratios (and thus peak fitting
37 errors), natural fluid inclusions hosted in olivine show changes in density outside analytical error
38 with increasing powers (density changing by ~0 to 0.09 g/cm³ per 10 mW of increased power).
39 We show that the amount of heating of the trapped fluid, and thus the change in density, varies
40 drastically based on the absorption coefficient of the host (with melt inclusions and low Fo olivines
41 showing significant more heating than higher Fo olivines). We recommend analyses of fluids with
42 high optical absorption phases at low laser powers (<3-4 mW), and offer practical strategies to
43 optimize signal without raising laser power, noting that small changes in focus and X–Y position
44 can greatly increase signal intensity.

45 1. Introduction

46 Traditionally, the density of CO₂-rich fluids in geological samples have been measured using
47 Microthermometry, a technique in which phase changes are observed during heating and cooling
48 of trapped inclusions. Starting in the 1930s, it was recognised that CO₂ fluid/gas/vapour has a
49 distinct signal when analysed by Raman spectroscopy, with two strong peaks at ~1285 and ~1388
50 cm⁻¹ (Fermi, 1931; McLennan F.R.S. and Smith M.A., 1932). These peaks are collectively called
51 the Fermi diad. For brevity, we refer to the lower wavenumber peak of CO₂ as diad 1 (also ν⁻ or
52 2ν₂), and the higher wavenumber peak as diad 2 (also ν⁺ or ν₁+2ν₂). In the 1970s, the distance
53 between the two peaks – referred to as the Fermi diad separation, splitting, peak splitting or diad
54 splitting – was shown to be related to the density of CO₂ (Wright and Wang, 1973). These findings

55 opened up the possibility of using Raman spectroscopy to quantify the density of CO₂ fluids, an
56 important parameter for many geological research questions (Rosso and Bodnar, 1995).

57 The relationship between diad splitting and CO₂ density on a given Raman instrument is described
58 using a densimeter, which is typically modelled as a polynomial function of Nth degree. A number
59 of groups have developed experimental apparatus to determine densimeters for different
60 instruments (DeVitre et al., 2021; Fall et al., 2011; Kawakami et al., 2003; Lamadrid et al., 2017;
61 Rosso and Bodnar, 1995; Song et al., 2009; Wang et al., 2019, 2011; Yamamoto and Kagi, 2006,
62 see Supporting Table 1). Not only do densimeters vary widely between Raman instruments (e.g.
63 Lamadrid et al. 2017, DeVitre et al. 2021); changes in hardware on a single instrument have also
64 been shown to affect the densimeter. For example, Lamadrid et al. (2017) show that the
65 densimeter on a single instrument can be shifted by changing the laser wavelength (e.g. a ~0.02
66 g/cm³ shift from 514 nm to 632 nm at 2400 g) or the instrument grating (e.g., a 0.02-0.04 g/cm³
67 shift from 600-2400 g at 632 nm). They conclude that the same wavelength and grating must be
68 used for calibration and analysis of unknowns.

69 In contrast, recent work has argued that densimeters may be more universally applicable than
70 previously thought. Remigi et al. (2021) generate a calibration using natural fluid inclusion (FI)
71 standards with densities between 0.37-1.08 g/cm³ and then compare their densimeter to
72 published densimeters using bootstrapping techniques to better account for uncertainty in
73 densimeter regressions. They suggest that published Raman densimeters fall into two main
74 categories – those collected on instruments with higher spectral resolution (~0.4 cm⁻¹/pixel) and
75 those collected on instruments with lower spectral resolution (~1-1.5 cm⁻¹/pixel). For CO₂ densities
76 > 0.3-0.4 g/cm³, they show that published densimeters within the high- or low-resolution category
77 are within error of each other. They suggest that because the spectral resolution is related to the
78 ability of a Raman spectrometer to discriminate overlapping bands, and that a change in the
79 spectral resolution shifts peak center positions in a systematic way. They conclude that
80 instrument-specific densimeters may not be required if a densimeter exists for the spectral
81 resolution of interest.

82 For the remainder of this paper, we refer to the parameter Remigi et al. (2021) term spectral
83 resolution to describe the wavelength range recorded in each pixel as pixel resolution (following
84 Hagiwara et al. 2023). This is not the same as the true *spectral resolution*, which is the smallest
85 width of peak that can be resolved, typically estimated from the full width half maximum (fwhm)
86 of very sharp emission lines such as Ne or Ar, or calcite. These two resolution parameters are not
87 equivalent on many Raman instruments, and there is no simple way to deduce one parameter

88 from the other. For example, using a 1200 grooves/mm grating (g), the pixel resolution reported
89 by Hagiwara et al. (2023, Table 1) is $0.482 \text{ cm}^{-1}/\text{pixel}$ while the spectral resolution estimated from
90 the fwhm of the Ne emission line at 1449 cm^{-1} as 1.59 cm^{-1} . However, using the 1800g, they
91 report a pixel resolution is $0.284 \text{ cm}^{-1}/\text{pixel}$ vs. a spectral resolution of 0.29 cm^{-1} . In Hagiwara et
92 al. (2020), the 1800g yields a pixel resolution of $0.29 \text{ cm}^{-1}/\text{pixel}$ but a spectral resolution of 1.1
93 cm^{-1} estimated from the fwhm of calcite. The vast majority of studies do not report their spectral
94 resolution (Supporting Table 1), which makes it very hard to assess if there truly is a resolution is
95 controlling the densimeter equation.

96 It is also notable that published densimeters differ the most for analyses of gaseous CO_2 at room
97 temperature ($\rho \sim 0\text{-}0.2 \text{ g/cm}^3$; Lamadrid et al. 2017), while Remigi et al. (2021) only consider
98 densities $>0.3 \text{ g/cm}^3$. At low densities, offsets between densimeters can be a factor of 2 or more.
99 This has major implications in igneous petrology, because many melt inclusion vapor bubbles
100 have CO_2 densities below 0.2 g/cm^3 (see Fig. 21 of Wieser et al., 2025). Accurate measurements
101 of these low density vapor bubbles are essential for reconstructing the CO_2 contents of magmas,
102 with implications for determining magma storage depths and the cause of explosive eruptions
103 (Allison et al., 2021; DeVitre et al., 2023b; Moore et al., 2015). However, studies published after
104 the Lamadrid et al., (2017) paper advocating for instrument-specific densimeters still report
105 densities for CO_2 fluids with low densities using calibrations published in other laboratories (e.g.
106 Howe et al., 2025; Venugopal et al., 2020). In this study, we critically evaluate sources of
107 densimeter variability, to evaluate whether generic calibrations based on a simple instrument
108 parameters can replace instrument-specific calibrations.

109 Another recent discussion of the utility of Raman spectroscopy for quantifying CO_2 densities was
110 provided by Bakker (2021). This work assessed the uncertainties associated with various parts of
111 the Raman method (e.g., peak fitting, instrument drift, etc) and suggested that errors in the
112 position of each diad propagate through to very large errors in diad splitting ($\pm 0.26\text{-}0.27 \text{ cm}^{-1}$) and
113 thus CO_2 density ($\sim \pm 0.1 \text{ g/cm}^3$). The large quoted errors of Bakker (2021) are in contrast to the
114 very low theoretical errors for analyses of CO_2 within a high pressure optical cell suggested by
115 Yamamoto and Hagiwara, (2024, 0.0005 cm^{-1} , 0.00015 g/cm^3). Thus, it is imperative to evaluate
116 these contrasting statements, and determine the analytical errors associated with analyses of
117 natural melt and fluid inclusions.

118 Overall, the Bakker (2021) and Remigi et al. (2021) papers demonstrate that while a lot of
119 progress has been made in recent years on optimizing Raman analyses of CO_2 fluids (DeVitre et
120 al., 2023a, 2021; Fall et al., 2011; Lamadrid et al., 2017; Le et al., 2021; Wang et al., 2011), there

121 are still some fundamental disagreements in the community regarding the applicability of
122 densimeters, and the accuracy and precision of the Raman method when applied to natural
123 samples. In this study, we investigate the effect of acquisition and data processing strategies on
124 the precision and accuracy of CO₂ fluid measurements on a WITec alpha300 R in the Department
125 of Earth and Planetary Science, University of California, Berkeley (hereafter, UCB WITec). We
126 describe a series of tests that readers can apply to their own instruments to guide the development
127 of laboratory-specific best practices, with the aim that this leads to a greater understanding in the
128 community of which sensitivities are instrument and hardware dependent, and which are globally
129 applicable.

130 2. Methods

131 We acquire Raman spectra on natural FI and melt inclusions (MI) hosted in olivine using the UCB
132 WITec. The Ultra-High-Throughput-Spectrometer (UHTS600) was calibrated using a Hg-Ar lamp
133 using the WITec control software. The CCD chip (1024X127 pixels) was Peltier cooled to -60°C.
134 We use a green solid-state laser with a wavelength of 532.046 nm, and a Ne lamp in the beam
135 path to correct for instrument drift. For measurements of the CO₂ Fermi diad, the 1800
136 grooves/mm grating yields a pixel resolution ranging from 0.57- 0.51 cm⁻¹/pixel from
137 wavenumbers of 1042-1592 cm⁻¹. The spectral resolution estimated from the width of Ne lines is
138 ~0.75-0.86 cm⁻¹ (Supporting Fig. 1). To convert diad splitting into CO₂ density, we use the
139 densimeter developed for our instrument by DeVitre and Wieser (2024) using a gas cell calibration
140 device (adapted from DeVitre et al. 2021). This calibration line is valid for CO₂ densities spanning
141 0.0096-1.04 g/cm³.

142 The laser was focused on the sample with a variety of objectives depending on the sample size
143 and heating set up; an Olympus 40X objective (0.6 NA, 2.7-4 mm working distance), a Zeiss 50X
144 objective (x0.55NA, 9.1 mm working distance), or a Zeiss 100X objective (0.95 NA, 4 mm working
145 distance). Laser power was controlled using the WITec TruePower module, which measures the
146 laser power in the optical fiber, and allows the user to enter values between 0-50 mW with ±0.1
147 mW precision. We independently verify the laser power after it has left the fibre optic and passed
148 through the optical head using a ThorLabs S121C sensor attached to a PM100D2 optical power
149 meter. There is a strong linear correlation between the measured power through different
150 objectives, and the instrument TruePower, with intercepts close to 1 (Supporting Fig. S2). Placing
151 the power meter at the opening of the objective turret with no objective screwed in yields laser
152 powers that are 85% of the TruePower value. For the 5-100X objective, we measure the power
153 by focusing on the laser power meter in the normal sample position (i.e. after passing through 1

154 cm of air). This yields powers that are 80% (5X), 79% (20X), 77% (40X), 62% (50X) and 43%
155 (100X) of the TruePower value (Supporting Fig. S2). Unlike our previous work examining the
156 effects of laser power which report TruePower values (DeVitre et al., 2023a), in this study all
157 subsequent mention of laser power uses the laser power at the sample, calculated using the
158 correction factors mentioned above:

$$159 \text{ Power}_{\text{sample}} = \text{Power}_{\text{TruePower}} \times \text{loss} \text{ [Eq 1]}$$

160 For example, if we set the TruePower value to 10 mW in the software, we report a power of 4.3
161 mW for 100X (10X0.43). This allows greater comparison between studies performed on different
162 instruments.

163 Various methods were used to heat samples above the critical temperature of CO₂ to ensure a
164 single homogenous fluid was being analysed (DeVitre et al., 2023a). For most analyses, samples
165 were placed on a glass slide in crystalbond and placed on a Peltier heating stage mounted within
166 a custom aluminium holder with a central hole for transmitted light illumination (Supporting Fig.
167 S3). Temperature was monitored using a type K thermocouple taped to the slide with electrical
168 tape. The voltage to the peltier plate was adjusted to yield a sample temperature of 37 °C,
169 matching the gas cell calibration temperature for this instrument (DeVitre and Wieser, 2024).
170 When sample temperature was being directly investigated (Section 3.6), crystals were double
171 polished and placed in a Linkam THMSG600 stage (Supporting Fig. S4). To achieve the long
172 working distance required to see through the stage lid, we initially made measurements using the
173 LWD 40X Olympus objective with a cover slip correction collar. After doing tests with and without
174 the top lid, we found that the stage lid cut out ~50% of the signal, and that there was no difference
175 in sample heating between heating series performed with and without the stage lid (Supporting
176 Fig.S5). Thus, later analyses were performed using the Zeiss 50X objective without the stage lid,
177 but with the small silver cap still placed on top of the sample to help with temperature stability
178 (Supporting Fig. S4).

179 All peak fitting and data processing was performed in the open-source Python3 tool DiadFit
180 (Wieser and DeVitre, 2024) using psuedovoigt peaks, except when investigating peak fit shapes.
181 Versions <v1.0.18 of DiadFit prior to March 2026 use an unweighted least squares routine to fit
182 peaks, resulting in each pixel contributing equal weight to the fit. As part of this work, we revise
183 DiadFit (>v.1.0.19+) to use weighted least squares, with weights assigned to each pixel
184 proportional to the inverse square root of the background-subtracted intensity, consistent with the
185 expected errors from Poisson counting statistics. Uncertainties in peak fit parameters are

186 estimated from the covariance matrix, with the lmfit option scale_covar=True applied to rescale
187 parameter uncertainties such that the reduced chi-squared value is unity. The peak positions do
188 not change within the original uncertainties estimated by v<1.0.18, but the errors on each peak
189 position range from 10 to 80% higher than estimates using unweighted least squares (Supporting
190 Fig. S6-9).

191 3. Results and Discussion

192 3.1 Neon line corrections and spectral non-linearity

193 Instrument drift during a Raman session is commonly assessed using repeated analyses of the
194 spectral emission of Neon generated by a Ne lamp (e.g. DeVitre et al., 2021; Lamadrid et al.,
195 2017. Fig. 1a). Two emission lines bracketing the CO₂ Fermi diad are selected. The theoretical
196 distance between these lines is calculated from the published wavelength of the *observed*
197 emission spectra of Neon in air from the NIST webbook (<https://webbook.nist.gov/>), converted to
198 wavenumber using the wavelength of the laser outputted by the WITec software (a fixed value for
199 each instrument).

$$200 \quad \tilde{\nu}_{line\ 1} = 10^7 \left(\frac{1}{\lambda_o} - \frac{1}{\lambda_{line\ 1}} \right)$$

$$201 \quad \tilde{\nu}_{line\ 2} = 10^7 \left(\frac{1}{\lambda_o} - \frac{1}{\lambda_{line\ 2}} \right)$$

202 Where $\tilde{\nu}_{line\ 1}$ is the wavenumber of Ne line 1, $\lambda_{line\ 1}$ is the Ne emission wavelength of line 1, and
203 λ_o is the laser wavelength. No vacuum to air conversions were applied to the stated laser
204 wavelength. We note that the choice of laser wavelength cancels out, so knowing this value
205 precisely does not affect the correction factor:

$$206 \quad \Delta\tilde{\nu}_{Ne} = \tilde{\nu}_{line\ 2} - \tilde{\nu}_{line\ 1} = 10^7 \left(\frac{1}{\lambda_{line\ 1}} - \frac{1}{\lambda_{line\ 2}} \right)$$

207 The calculated theoretical distance between two Ne lines is then compared to the measured
208 distance in the acquired spectra. For example, if the theoretical distance between Ne lines is
209 330.4776 cm⁻¹, and the measured distance is 331.0587 cm⁻¹, the correction factor is

210 $\frac{\text{Theoretical distance (330.4776)}}{\text{measured distance (331.0587)}} = 0.9982$. If the diad splitting was measured at 103 cm⁻¹, the corrected

211 diad splitting would be 0.9982*103 = 102.8192 cm⁻¹.

212 It has been suggested that the common method of correcting for instrument drift using the
213 distance between two Ne lines (e.g. Lamadrid et al., 2017) does not fully capture the complex
214 non-linearity of Raman spectrometers. Non-linearity arises from the mapping between detector

215 pixel position and wavelength, which is typically 'calibrated' using numerous peaks of known
216 position. On our WITec instrument, this calibration is performed by the WITec control software
217 using a Hg-Ar lamp. Ideally, if spectrometer calibration was perfect, the measured wavelengths
218 of all reference spectrum (e.g. Neon and Argon lamps, benzonitrile, 4-aceamidophenol etc, Liu
219 and Hennelly, 2024) would match the true wavenumbers of these peaks within peak fitting error.
220 However, non-linearity is the norm rather than the exception, and Raman spectra show clear
221 offsets between measured and preferred peak positions, with the offset often varying in linear or
222 sinusoidal patterns as a function of wavenumber (Bakker, 2021; Liu and Hennelly, 2024, Fig. 1b).

223 We investigate spectrometer non-linearity, and the effect of this non-linearity on instrument drift
224 and calculated densities using data collected on the UCB WITec, and two additional alpha 300R
225 spectrometers at Cornell using data from DeVitre et al. (2021). On the UCB instrument, we first
226 perform 50 consecutive repeated acquisitions of the emission spectrum of Neon. We fit
227 pseudo-voigt curves to the 6 major Ne lines within the spectral window for each acquisition using
228 DiadFit. For brevity, we refer to these lines using the calculated wavenumber positions rounded
229 to the nearest integer for the 532.046 nm laser wavelength used in this study (565.66588 nm
230 $\approx 1117 \text{ cm}^{-1}$, 568.98163 nm $\approx 1220 \text{ cm}^{-1}$, 571.92248 nm $\approx 1310 \text{ cm}^{-1}$, 574.82985 nm $\approx 1399 \text{ cm}^{-1}$
231 576.44188 nm $\approx 1447 \text{ cm}^{-1}$, and 580.44496 nm $\approx 1567 \text{ cm}^{-1}$, Fig. 1a, Kramida and Ralchenko,
232 1999). For the line at 1117 cm^{-1} , we iteratively fit two pseudo-voigt functions to resolve the smaller
233 peak on the shoulder. Fig. 1b shows the fitted Ne line position divided by the theoretical position
234 calculated from the laser wavelength (532.046 nm) and the theoretical positions in air. This
235 comparison demonstrates that the spectrometer is indeed non-linear, with the correction factor
236 varying across the acquisition window (Fig. 1b). Spectral non-linearity has been noted as a
237 serious issue when trying to compare spectra collected between instruments, as it persists after
238 performing the wavelength calibration methods built into Raman acquisition software (Liu and
239 Hennelly, 2024).

240 As Bakker (2021) correctly points out, the 'line segment' Ne correction method assumes that the
241 correction factor at any given point between two lines is the same. They suggest that a 'bracketing
242 technique' should be used instead, where a polynomial model is developed to describe how the
243 position of each Ne line deviates from the theoretical position as a function of wavenumber. This
244 polynomial allows a correction factor to be calculated for diad 1 and diad 2. These corrected peak
245 positions are used to calculate the splitting, rather than applying a correction factor to the splitting
246 as in the line-segment method.

247 The polynomial fit through these correction factors is shown for the 'bracketing technique' (grey
248 curve, Fig. 1c) for a single Ne acquisition. Correcting for the position of the CO₂ diad using this
249 curve this would yield a corrected splitting for the diad position shown by blue dotted lines that
250 would be equivalent to a line segment correction factor of 0.9977. The 'line segment' technique
251 for drift correction can be visualized by drawing a straight line between any two datapoints (circles)
252 on this polynomial curve (Fig. 1c). If the gradient of this linear segment is the same as the gradient
253 of a line drawn between where the x positions of the diad intercept the polynomial, the line
254 segment technique and bracketing techniques would give the same result. The line segment
255 method applied to the 1220-1399 cm⁻¹ lines results in a very similar correction factor to the
256 bracketing technique (0.9976 vs. 0.9977). The disadvantage of choosing this pair is that the 1220
257 cm⁻¹ line is relatively small, resulting in larger fitting errors, meaning that correction factors show
258 substantially more variability in each run than using the 1117-1447 cm⁻¹ pair (see error bar on Fig.
259 1c). The 1117-1447 cm⁻¹ pair used by our laboratory yields a reasonably similar correction factor
260 to the bracketing technique (0.9982 vs. 0.9977). However, applying the line segment technique
261 to the 1220-1567 cm⁻¹ lines (red dashed segment) would result in a Ne correction factor that is far
262 smaller than that determined using the bracketing technique (0.9991 vs. 0.9977), yielding a higher
263 apparent diad splitting ($\Delta_{CO_2}^{Corr,brack}=103.450$ cm⁻¹ vs. $\Delta_{CO_2}^{Corr,seg}=103.648$ cm⁻¹).

264 The differences in correction factor mean that the choice of Ne line pair does indeed affect the
265 corrected splitting, and the relationship between CO₂ density and corrected splitting (i.e. the
266 densimeter). To demonstrate this, we take the CO₂ and Neon spectra from the analytical sessions
267 used to calibrate the UCB densimeter, which was originally calibrated from corrected splitting
268 calculated using correction factors from the 1117-1447 cm⁻¹ line pair (DeVitre et al., 2021; DeVitre
269 and Wieser, 2024, black dotted line, Fig. 2a-b). We reprocess the data using 4 new pairs of Ne
270 lines: ~1220-1447 cm⁻¹ (dashed red line, Fig. 2a, c), ~1117-1399 cm⁻¹ (dashed blue line), 1220-
271 1398 cm⁻¹ (solid cyan line), and 1220-1567 cm⁻¹ (solid magenta line, Fig. 2a-b). For each CO₂
272 spectrum collected on a fluid of known density in the calibration cell, this gives us 5 different
273 corrected splitting values (one for each Ne line pair). Using three polynomial fits for low, medium
274 and high densities (following DeVitre et al. 2021), we calibrate coefficients for 5 different
275 densimeter equations using these 5 sets of corrected splitting values (Fig. 2a-b). Despite being
276 collected on a single instrument using the same calibration dataset, the densimeters differ by up
277 to ~0.08 g/cm³ simply based on the Ne line pair used for data reduction. To put the magnitude of
278 these offsets into context, we underlay the densimeters collected on a wide variety of Raman
279 instruments (Fall et al., 2011; Kawakami et al., 2003; Lamadrid et al., 2017; Wang et al., 2011,

280 2011; Yamamoto and Kagi, 2006), colored using the high (grey) and low (salmon) resolution
281 classifications from Remigi et al. (2021). This demonstrates that the variability from a single
282 instrument introduced through the choice of a Neon line pair is as large as the differences between
283 Raman instruments, despite being collected at a single spectral resolution.

284 If a gas cell apparatus is available for calibration of a specific instrument, the offsets between
285 densimeters determined using different Ne line corrections are only problematic if they yield
286 different results when analysing unknowns. We test whether this is the case using data collected
287 on FI from Fogo volcano collected by DeVitre and Wieser (2024), where there are independent
288 constraints on CO₂ density from microthermometry. We process the FI data using each of the 5
289 different Ne line pairs (Fig. 2b), generating 5 values of corrected splitting for each spectrum, and
290 input these values into the densimeter corresponding to each Ne line pair (e.g. data reduced using
291 correction factors for the 1220-1399 pair was fed into the 1220-1399 densimeter). To visualize the
292 differences between these data reduction strategies, we subtract the density calculated using
293 each line pair from that calculated using the 1117-1447 pair, which was shown to correlate very
294 well with Microthermometry (DeVitre and Wieser, 2024). The error bar on each datapoint shows
295 the combination of the peak fitting error, drift correction error, and uncertainty on the densimeter.
296 Densities for the 4 other line combinations are within error of the data reduced using the 1117-
297 1447 cm⁻¹ line pair. We also overlay lines showing the 1σ uncertainty associated with
298 Microthermometry assuming a 0.1 °C and 0.2 °C error in sample temperature (Bakker, 2021). This
299 analysis demonstrates that the uncertainties in the choice of data reduction strategies are small,
300 and comparable to the errors associated with the stage temperature when performing
301 Microthermometry.

302 Repeated acquisitions of Neon lines at UCB in April 2022 and October 2024 shows that the non-
303 linearity response curve shifts up and down with time, but the shape is maintained (Fig. 4a, solid
304 lines). We also tested recalibrating the UCB spectrometer in April 2024 using a Hg-Ar lamp
305 (WITec's default calibration routine), and the same pattern is observed. Thus, while spectrometer
306 non-linearity is clearly significant, and likely introduces diversity in published densimeters, as long
307 as calibration and natural data is processed in the same way, the non-linearity introduces a
308 consistent offset through time, and choice of Ne line will not cause systematic offsets. This stability
309 also means that the bracketing technique is not necessary, and given the additional computation
310 steps, likely results in a larger uncertainty due to the large fitting errors associated with the weaker
311 Ne lines (e.g. 1310 cm⁻¹). However, if a laboratory wishes to change drift correction strategy, they
312 would need to reprocess their calibration data to use the same strategy.

313 This analysis demonstrates that not only does a densimeter need to be generated for each
314 instrument, but that each densimeter is only valid for a specific detector calibration and drift
315 reduction protocol. Supporting Table 1 shows that a wide variety of calibration and drift correction
316 routines are being used by different laboratories publishing densimeters. This can be expected to
317 generate a wide variety of non-linearity responses, and thus we speculate that a lot of the diversity
318 in measured densimeters in the literature is a result of spectral non-linearity, both in its original
319 effect on the diad splitting, and the compounding effect of the use of a line-segment Ne line
320 correction. It is possible that performing a polynomial fit like that suggested by Bakker (2021)
321 would resolve offsets between densimeters calibrated in different laboratories; extensive testing
322 using different Raman instruments for standards with known densities would be a promising
323 avenue for future research.

324 3.1.1 The effect of changing the spectral window

325 The significant non-linearity in the spectrometer response shown in Fig. 1b has other implications
326 for the densimeter equation determined for a given instrument. When setting up a Raman
327 spectrometer, the user must define the wavelength range that is recorded on the CCD. On the
328 three WITec instruments considered here, this is specified by stating the spectral center—the
329 wavenumber at the center of the acquisition window. Along with the grating, this center position
330 determines the range of wavenumbers collected in a single spectral acquisition. Shifting the
331 center means that the photons at a specific wavenumber are collected on a different pixel on the
332 charge-coupled device (CCD) detector within the spectrometer. On other instruments, the window
333 edges rather than center may be the tuneable parameter.

334 Given the strong spectral non-linearity shown in Fig. 1b, it is worth investigating whether shifting
335 the spectral center affects the measured splitting, because each pixel of the diad would shift to a
336 different part of the polynomial in Fig 1b (see also Bakker 2021). We perform repeated
337 acquisitions on an olivine-hosted CO₂-rich FI from Fogo Volcano with a bulk density of ~0.62
338 g/cm³ using different spectral centers. We started at the spectral center used for calibration of the
339 Berkeley densimeter (1325 cm⁻¹), then decreased the spectral center
340 (1325→1320→1275→1250→1225 cm⁻¹, magenta colors, Fig. 3a). To help resolve changes from
341 the spectral center vs instrument drift, we increased (green colors), decreased (yellow colors) and
342 then increased (cyan) the spectral center throughout the day (Fig. 3a). We acquired a Neon line
343 at the same spectral center as each FI measurement to correct each spectra (using the 1117-
344 1447 cm⁻¹ line pair). For each individual acquisition, we calculated the corrected splitting using

345 the Ne line acquired at the same spectral center. Lastly, we calculated density using the UCB
346 densimeter developed using a spectral center of 1325 cm^{-1} .

347 The uncorrected splitting of the diad increases by $\sim 0.22 \text{ cm}^{-1}$ as the spectral center is increased
348 from 1225 to 1350 cm^{-1} ($\sim 0.0017 \text{ cm}^{-1}$ per cm^{-1} of center shift, Fig. 3b). The change in splitting for
349 spectral center shifts of only $\sim 25 \text{ cm}^{-1}$ apart is far larger than the variation at a given spectral
350 center throughout the day (e.g. instrument drift). After correcting each acquisition using the Neon
351 line collected at the same spectral center, the change in splitting is smaller, but still notable (~ 0.15
352 cm^{-1} across the window, 0.0012 cm^{-1} per cm^{-1} of shift, Fig. 3c). Calculated CO_2 densities after a
353 line segment Ne line correction change by 0.08 g/cm^3 across the window (Fig. 3c), which is far
354 outside the $\sim 0.01\text{-}0.02 \text{ g/cm}^3$ uncertainty of the Raman method at the densities of interest (Dayton
355 et al., 2023; DeVitre and Wieser, 2024). Overall, this test indicates that even just shifting the
356 spectral window on a given instrument can generate a new densimeter, and that unknowns must
357 be analysed at the same spectral center/spectral window at which the calibration was performed.

358 To further illustrate the influence of spectral center, we compare the densimeter developed
359 on the UCB WITec alpha300R with densimeters developed in two laboratories at Cornell
360 University (Cornell Center for Materials Research CCMR and Cornell Mass Spectrometry Facility
361 CMaS, DeVitre et al., 2021). These two laboratories were calibrated with a gas cell that is very
362 similar to that at UCB (following the design of DeVitre et al., 2021), and the Raman hardware for
363 all three instruments is very similar: a $\sim 532 \text{ nm}$ laser wavelength, 1800 grating, and Ultra-High-
364 Throughput-Spectrometer UHTS600. All data was reduced with the 565.66588 - 576.44188 nm
365 Neon line pair (corresponding to 1117-1447 for the UCB 532.046 nm laser wavelength).
366 Comparing the fitted neon line positions to the theoretical positions (Fig. 4a) shows that each
367 spectrometer has its own unique signature of non-linearity.

368 Despite CMaS and CCMR having more similar non-linearity responses, the CCMR
369 densimeter shows the biggest offset in densimeter equation relative to UCB and CMaS. We
370 suggest that this is because the UCB densimeter and the CMaS densimeter were both calibrated
371 using a spectral center of 1325 cm^{-1} , while the CCMR densimeter was calibrated with a spectral
372 center of 1340 cm^{-1} . The absolute difference between these densimeters reaches $\sim 0.04 \text{ g/cm}^3$
373 (Fig. 4c). Again, these densimeter differences cannot be explained by differences in resolution as
374 advocated by Remigi et al. (2021), because the pixel resolution of all three of these WITec
375 instruments is identical.

376 3.2 Densimeter stability as a function of time

377 The above tests make it apparent that small changes in instrument hardware can shift the
378 densimeter. It is interesting to assess whether temporal shifts occur when using identical
379 acquisition conditions and spectrometer settings. We evaluate repeated analyses of internal fluid
380 inclusion standards from Fogo volcano (FG04-A1-4, N=280 and FI FG04-101-A, N=96) collected
381 at the start and end of analytical sessions over 2.5 years. The standard deviation calculated from
382 the 280 repeats on A1-4 is 0.008 g/cm^3 (Fig. 5a) and for the 96 repeats of 101A is 0.0070 g/cm^3 .
383 For both inclusions, microthermometry measurements are within 0.01 g/cm^3 of these long-term
384 averages and lie within or extremely close to the 1σ envelope around the mean Raman
385 measurement. Overall, this small deviation and lack of coherent trends with time indicates that at
386 least on the WITec, changes in spectrometer non-linearity does not cause any notable drift in the
387 densimeter, and densimeters are remarkably stable once calibrated. Repeated acquisitions of a
388 natural fluid inclusion in a similar manner to that performed here can be used to assess the point
389 at which the instrument may require a densimeter recalibration.

390 While the Neon method is clearly robust, even with variations in non-linearity, it is
391 worthwhile to determine whether it actually helps correct for drift between and within runs.
392 Yamamoto and Hagiwara (2024) show that over short timescales (minutes) where drift is not
393 controlled by temperature, periodicity is seen in CO_2 peak positions that is not seen in Ne line
394 acquisitions. We reprocess the repeated FI acquisitions using the median Ne correction factor for
395 all the standard analyses, rather than the correction factor calculated by regressing Ne correction
396 factors as a function of time in each analytical run. The standard deviation of both FIs is extremely
397 similar with and without an analysis-specific Ne correction. Interestingly, for FG04 one outlier is
398 generated using a Ne line correction specific to that analysis that goes away using the average
399 correction factor (Fig. 5a), while the opposite is true for standard FG19_101 (one outlier is
400 generated using the median Ne correction). However, it is noteworthy that after the instrument
401 recalibration in April 24, the Raman was moved to a newly renovated laboratory where the HVAC
402 system yields temperature variations of $<\pm 1 \text{ }^\circ\text{C}$ per day, which means temperature variations are
403 minimal, which could explain the small differences between the correction methods. We
404 investigate the ability of line segment Ne line corrections to help minimise drift during larger
405 temperature variations in section 3.3 below.

406 3.3. Densimeter stability within a single analytical session

407 It is well known that changes in temperature can cause considerable drift in peak positions,
408 intensities and peak widths during Raman analyses (e.g., Hagiwara et al., 2018; Inoue et al., 2023;
409 Jakubek and Fries, 2022; Lamadrid et al., 2017; Mestari et al., 1997). For example, Fukura et al.

410 (2006) track peak positions in ruby and naphthalene, showing periodic fluctuations in peak
411 positions of $\sim 0.1 \text{ cm}^{-1}$ with temperature fluctuations of $0.8 \text{ }^\circ\text{C}$. However, they show that the multiple
412 peaks of Naphthalene follow a similar trend, and the difference between two peak positions, which
413 is of most relevance for measurements of CO_2 fluids, is $\sim 8 \text{ X}$ smaller than the absolute peak shift.
414 In contrast, Mestari et al. (1997) perturb room temperature by $1\text{-}2 \text{ }^\circ\text{C}$ and track the relative
415 positions of different Neon lines and show the wavenumber shifts are not identical for different
416 lines and diverge further as the experiment proceeds. This study also shows that amount of
417 relative wavenumber shift between two lines differs drastically as a function of the acquisition
418 conditions.

419 Hagiwara et al. (2018) investigate the effect of both room and sample temperature on CO_2
420 density measurements to isolate the effect of each. They show no notable change in splitting as
421 sample temperature is perturbed by $21\text{-}26 \text{ }^\circ\text{C}$, but an abrupt shift in splitting of $\sim 0.05 \text{ cm}^{-1}$ as the
422 room temperature is raised by 7°C . The splitting then rapidly shifts back to its original value again.
423 As the authors do not plot splitting against room temperature, it is difficult to determine exactly
424 when this perturbation happens relative to the onset of room heating. They suggest the splitting
425 change reflects thermal expansion induced misalignment of the spectrometer, as thermal
426 expansion of the host mineral will be negligible. The Neon line correction routines of Lamadrid et
427 al. (2017) and Lin et al. (2007) do not mitigate the change in splitting during this perturbation, and
428 suggest that repeated measurement a CO_2 standard may provide a better drift correction
429 strategy.

430 Given the differential response to temperature Mestari et al. (1997) report using different
431 instrument settings on a single instrument, and conflicting reports of whether different lines moves
432 together (Mestari et al. 1997 vs. Fukura et al. 2006), we investigate the effect of room temperature
433 variations and whether the Ne line segment technique can minimise the effect on measured
434 splitting on the UCB WITec alpha300R. We turned off the room heating and cooling system the
435 night before a day that was forecast to be particularly warm for the San Francisco Bay area ($\sim 28\text{-}$
436 30°C). We started the test at $\sim 6\text{:}30 \text{ am}$ when the laboratory temperature was $23.5 \text{ }^\circ\text{C}$, 30 mins
437 after the laser and Neon were turned on and the spectrometer set to cool, Two Icelandic fluid
438 inclusions (Kahl et al., 2025) were double polished and placed in the Linkam THM600 stage with
439 a silver lid on top of the sample to keep the sample temperature fixed at $37 \pm 0.1 \text{ }^\circ\text{C}$ throughout the
440 day. We do not use the top lid of the stage, as it cuts out a large proportion of the Raman signal
441 (see Supporting Fig. S4). As the room warmed up, the two FI were analysed repeatedly, with a
442 Ne line acquisition collected after each measurement of the two FI. Room temperature was

443 measured using an Aranet4 sensor located on the table with the Raman and Ne lamp (Fig. 6a).
444 At around 3pm, the room temperature had stabilized at around 26 °C as the outside temperature
445 peaked. To get a larger temperature rise, a temporary heater was switched on. At 5 pm, the
446 temperature began to level out again, so a second heater was turned on, causing the laboratory
447 temperature to rise to 28 °C at ~ 7:30 pm. Two air filters with a combined clean-air delivery rate of
448 500 cubic feet per minute were used to circulate air within the room to help homogenize laboratory
449 temperature. The laser power recorded by the WITec TruePower stays between 5.95 and 6.2 mW
450 with no obvious correlation with room temperature, showing similar variability to analytical days
451 where temperature is stabilized (Supporting Fig. S11).

452 The peak positions of the 1117, 1220, 1400, 1447, and 1567 lines increase by ~0.2-0.3
453 cm⁻¹ over 5°C, which is very similar to the 0.3 cm⁻¹ shift of the peak position of GaN observed
454 over 5°C of room warming by Guo et al. (2020). Interestingly, the proportional change in peak
455 position is extremely similar for all the Ne lines (position at t=i/position at t=0), so the absolute
456 change is largest for the 1567 line and smallest for the 1117 line (Supporting Fig. S12). While the
457 Ne correction factors differ for each Ne line pair, the proportional change in correction factor is
458 within uncertainty for all Ne line pairs (Supporting Fig. S12). This indicates that the magnitude of
459 the drift correction using Ne lines is independent of the choice of line pair.

460 Given the similar response of all Ne lines and their correction factors, for simplicity, we
461 only discuss the effect of the 1117-1447 line pair. The 1117-1447 correction factor shows a strong
462 anticorrelation with temperature (Fig. 6a, note the flipped y axis for temperature). By parametrizing
463 temperature vs. time and Neon correction factor vs. time using high order polynomials (Fig. 6a),
464 a laboratory temperature and correction factor was calculated for each FI measurement. We
465 calculate the corrected splitting using the Ne correction model (blue line, Fig. 6a) and compare
466 this to corrected splitting calculated using the first Ne acquisition at the start of the day (Ne1, Fig.
467 6a). Fitting a linear regression for calculated density vs. laboratory temperature yields a far
468 stronger correlation when using a constant Ne correction factor vs. the Ne correction model (0.57
469 vs. 0.14 for B30, and 0.74 vs. 0.39 for B23). This indicates that during large temperature swings,
470 the line segment Ne line correction does help correct for instrument drift, keeping the rate of
471 change within ~0.01 g/cm³ for 10 °C of laboratory heating, while not using a Ne lamp would
472 generate ~0.03-0.04 g/cm³ of drift.

473 In our experience with our hardware and specific analytical protocol, on a day without such
474 extreme temperature perturbations, measuring a Neon line every 40-60 mins is generally
475 sufficient. However, we have found that our Ne line acquisitions have helped us identify days

476 where the HVAC system was malfunctioning and frequent Ne line acquisitions have allowed us
477 to use the data collected on those days, whereas with less frequent acquisitions we would have
478 had to recollect. Given the differing results reported for Raman temperature sensitivity in the
479 literature, we suggest using room temperature logs to determine typical daily temperature
480 changes in any given laboratory to decide an appropriate interval at which to measure Ne lines.
481 In laboratories with large temperature swings, performing room temperature tests similar to that
482 described here on their instrument can help identify the best drift correction strategy. If other
483 Raman systems behave like ours during temperature perturbations, where very large temperature
484 swings can be corrected using Ne lines so the analytical error only doubles (from an already small
485 value, $\sim 0.01 \text{ g/cm}^3$), this is excellent news for universities who want to conduct Raman
486 measurements but do not have access to temperature-controlled spaces.

487 3.4 Peak fitting errors

488 The above analysis shows that developing a densimeter for a specific analytical conditions and
489 drift correction methods is key to remove the large systematic offsets between different
490 densimeters and generate accurate measurements of CO_2 density. Now, we focus our discussion
491 to optimizing the precision of Raman measurements.

492 Bakker (2021) raised concerns regarding precision, noting the discrete nature of Raman spectra,
493 which comprise x coordinates (wavenumber) collected on the CCD with spacing defined by the
494 grating and other hardware. At each x coordinate, an intensity value is recorded, the magnitude
495 of which is highly dependent on the instrument. Bakker (2021) show that shifting the spectral
496 window one pixel spacing at a time causes a lateral shift in the maximum peak intensity (Fig. 7).
497 They suggest that the error on the peak center is equal to the magnitude of the spectral shift. For
498 two different Raman set ups, they estimate a splitting error of $\sim \pm 0.27 \text{ cm}^{-1}$. A splitting error of
499 $\pm 0.27 \text{ cm}^{-1}$ yields an error of $\pm 0.1 \text{ g/cm}^3$ using the UCB densimeter. For a typical melt inclusion
500 vapour bubbles with 3% bubble volume, this corresponds to an uncertainty of the amount of CO_2
501 held in the bubble of $\pm 1111 \text{ ppm}$. When propagated through to calculate saturation pressure, this
502 would introduce uncertainty of several kbar. Similarly, a FI with a density of $0.3 \pm 0.1 \text{ g/cm}^3$
503 corresponds to entrapment pressures spanning 0.5 to 1.25 kbar (Span and Wagner, 1996). Peak
504 fitting errors of the magnitude suggested by Bakker (2021) would render Raman spectroscopy
505 entirely unsuitable for melt and FI barometry, so it is essential to assess these statements.

506 We obtain the spectra shown in Fig. 18 of Bakker (2021) and fit them with pseudovoigt peaks for
507 the diads and hotbands. These spectra were collected as the spectral window was shifted $\sim 3 \text{ cm}^{-1}$
508 ¹ in steps of 0.137 cm^{-1} on a LabRamHR (600 grating, pixel resolution of $0.62 \text{ cm}^{-1}/\text{pixel}$). To help

509 visualize the shift in the maximum y pixel intensity, we vertically stack the acquisitions collected
510 using different spectral centers after normalizing the intensity for ease of visualization (Fig. 7a).
511 The positions of diad 1 and diad 2 do indeed move by significant amounts as the window is shifted
512 (Fig. 7b). Bakker (2021) state that the total error on the splitting is '*by definition, the sum of the*
513 *uncertainties of the lower and upper band (i.e. peak) of CO₂*', i.e. the amount of shift of each peak.
514 For this dataset, both diad1 and diad2 shift by 0.226 cm⁻¹ as the window is moved (red error bar,
515 Fig. 7c), which would yield a total uncertainty of ~0.45 cm⁻¹ following their logic (and 0.32 cm⁻¹
516 propagating errors in quadrature). However, it is crucial to recognise that diad 1 and diad 2 move
517 in the same direction (Fig. 7b), such that the calculated splitting varies very little and shows no
518 systematic change as the spectral window is shifted (cyan diamonds, Fig. 7c). The overall
519 variation in calculated splitting is well within the estimated error combining the peak fitting errors
520 on diad 1 and diad 2 from Imfit in quadrature within DiadFit (Newville et al., 2016). Thus, the
521 splitting errors suggested by Bakker (2021) are overestimated, because shifting the window
522 causes the position of diad1 and diad2 to move in the same direction, with no notable change in
523 the splitting.

524 Bakker (2021) also suggest that the choice of different probability density functions (PDFs, e.g.
525 Voigt, PseudoVoigt, Gaussian) return different center positions, and that the errors on these
526 parameters are hard to estimate. Fortunately, modern peak fitting packages such as Fityk (Wojdyr,
527 2010), and peak fit packages in python such as sklearn or Imfit (used in DiadFit, Newville et al.,
528 2016) provide an estimate of the uncertainty on the peak center using the covariance matrix
529 obtained during the peak fitting process. We assess the difference in peak position using different
530 fits for Raman data collected on natural FIs from Fogo Volcano, Cabo Verde and Kilauea, Hawaii
531 (DeVitre and Wieser, 2024). We consider Voigt, PseudoVoigt, and Pearson4 functions, as these
532 were identified as the most physically-plausible peak shapes for CO₂ fluids by Yuan and
533 Mayanovic (2017). Peak positions obtained from different fits overlap within the calculated
534 uncertainty on the peak center (Fig. 8, Supporting Fig. S13). This can be shown more formally,
535 by dividing the absolute difference between models by the fitting error combined in quadrature
536 (Fig. 8c-d). Y values <1 on this plot show that the difference between functions is less than the 1
537 σ estimated error. This analysis demonstrates that for a wide range of CO₂ densities and signal
538 intensity, differences between physically plausible peak fit shapes are within the stated error on
539 each model, and that quoted peak fitting errors are realistic.

540 It is notable that the errors on the PseudoVoigt peak position are twice as large for diad 1 as diad
541 2, consistent with its lower intensity and slightly higher full width half maximum (Fig. 9a). Errors

542 are also larger at lower CO₂ densities. However, for all inclusions, peak fitting errors are still
543 significantly smaller than the 0.1-0.3 cm⁻¹ estimated by Bakker (2021). Comparing Raman
544 measurements with Microthermometry measurements provides another opportunity to assess
545 precision (and accuracy, Fig. 9b). We use data from N=58 FI from Fogo volcano from DeVitre and
546 Wieser (2024) supplemented with N=39 new microthermometry and Raman measurements from
547 the Galápagos using the same methods. The Raman error was estimated from Diadfit, by
548 combining the peak fitting errors with the uncertainty from the Neon line correction model and the
549 densimeter. This error was combined in quadrature with the uncertainty in microthermometry
550 homogenization temperature (± 0.2 °C) to create the y error bar shown in Fig. 9b. On Fig. 9c, we
551 perform the same analysis but now show error bars estimated from multiple repeats of FI
552 analyses, rather than the DiadFit peak fitting error. On both plots, the two FI with the largest offset
553 between Raman and Microthermometry have CO₂ densities very close to critical, where
554 Microthermometry errors are hard to quantify. The majority of inclusions lie within the ~ 0.01 - 0.02
555 g/cm³ error commonly quoted for the Raman method (red bars, Dayton et al., 2023; DeVitre and
556 Wieser, 2024; Kobayashi et al., 2012; Remigi et al., 2021). In contrast, it is clear that when
557 propagated to density, the splitting errors of ~ 0.2 - 0.3 cm⁻¹ (grey dotted and dashed lines) from
558 Bakker, (2021) are not consistent with this dataset.

559 Although the errors estimated from comparing Microthermometry and Raman (Fig. 9), and
560 repeated measurements of standards (Fig. 5) are far smaller than that estimated by Bakker,
561 (2021), it is noteworthy that the standard deviation in density obtained from averaging multiple
562 repeats is about twice as large as the error estimated from DiadFit (using weighted least squares
563 or unweighted least squares, Fig. 9). Similarly, the individual fitting errors on repeated
564 measurements of FI standards (Fig. 5) are about half the total standard deviation in repeated
565 measurements. This indicates that the DiadFit error accounting for uncertainty in peak position,
566 Neon correction models, and the densimeter itself is not the only contribution to the density error.
567 In Section 3.6, we suggest that a reasonable fraction of the unaccounted-for uncertainty comes
568 from the temperature of the sample, and other sources of instrumental variability.

569 3.5. Optimizing Signals

570 While the above discussion indicates that peak fitting errors are far smaller than those indicated
571 by Bakker (2021), it is still worthwhile investigating methods to reduce uncertainties as much as
572 possible. This is particularly important for low density CO₂ fluids (i.e. melt inclusion vapour
573 bubbles), where peak fitting errors can approach 50-100% near the instrument detection limit
574 (~ 0.01 - 0.02 g/cm³, Bearden et al., 2025). Peak fitting errors are highly dependent on both the

575 instrument hardware (e.g., spectrometer efficiency, spectral resolution, read and dark noise from
576 the CCD), analytical conditions (e.g., laser power, acquisition time, number of accumulations
577 averaged), and sample characteristics (e.g., density of CO₂, size of inclusion).

578 On most instruments, users can adjust the acquisition time (time spent collecting a single
579 spectrum) and the number of accumulations (number of individual spectra that are collected and
580 averaged). We investigate the interplay between these two parameters using a variety of samples
581 with low CO₂ densities (which have the largest peak fitting errors). We collect N consecutive
582 individual acquisitions of variable duration using the 'slow series' option in the WITec acquisition
583 software. This mode keeps the laser on continuously, and collects and outputs spectra every Xs.
584 This is analogous to a single accumulation when run in the regular mode. To simulate the software
585 collecting two accumulations of time Xs, we average the 1st and 2nd collected spectra to get one
586 averaged spectra and then fit that spectra. We then average the 2-3rd acquisition, the 3-4th and
587 so forth. For N=14 acquisitions collected, this yields 13 averaged spectra that we fit (shown as
588 individual dots on Fig. 10a). We also calculate the average and standard deviation of each peak
589 parameter for the 13 averaged spectra. To simulate 3 accumulations, we average the 1-2-3rd
590 spectra, 2-3-4th, and for 4 accumulations, the 1-2-3-4th spectra, 2-3-4-5th spectra etc. Tests were
591 performed for different integration times (30s, 45s, 60s, 90s, 120s and 200s).

592 Fig. 10 shows a CO₂ capillary tube with a CO₂ density of 60 MPa, corresponding to $\rho=0.011\pm 0.011$
593 g/cm³ (Le et al., 2021). Supporting Fig. S16-S17 shows the same plot for a low density melt
594 inclusion vapour bubble from Mauna Ulu, Kīlauea (Wieser et al., 2019). Each 'averaged spectrum'
595 is shown as a semi-transparent dot on panel a-b, and the overall average for combinations
596 yielding N accumulations is shown as a cross with an error bar. For a single spectra, the errors
597 are smaller for 90s compared with 45s, as expected given the more intense spectra, and higher
598 signal to noise ratio (Fig. 10c vs 11d). The peak fit error for diad1 is about twice as large as for
599 diad 2. The error drops quickly as the number of accumulations is increased, and then the trend
600 flattens out (Fig. 10a-b). In terms of assessing ways to optimize instrument time, it is useful to
601 compare the errors in the context of the total acquisition time (Fig. 10e). There are no clear
602 differences in error outside of uncertainty the same total time for acquisitions 45s or greater. For
603 the two melt inclusion vapour bubbles, 30s acquisitions to return errors that are slightly higher
604 than all other times.

605 These observations can be understood when considering the factors that influence the uncertainty
606 in the intensity of each pixel collected on the CCD (Hagiwara et al. 2023). The error for each pixel
607 is a function of the shot noise (i.e., the Raman photons being delivered to the detector), dark noise

608 (arising from thermally produced electrons in the detector) and read noise (introduced during
 609 conversion of charge to voltage during readout). Shot noise (σ_{shot}^2) is proportional to counts, dark
 610 noise (σ_{dark}^2) is proportional to exposure time, and read noise is proportional to the number of
 611 times the detector is read (i.e., the number of accumulations). Increasing the time spent collecting
 612 each spectrum will decrease σ_{shot} , but increase σ_{dark} , while performing more integrations will
 613 increase the total contribution from read noise (σ_{read}).

614 We collect 20 accumulations of 2s, 5s, 10s, 30s and 60s with the laser and room lights off using
 615 the WITec slow series option. For each pixel (wavenumber) we calculate the variance after
 616 rejection of cosmic rays and then calculate the median variance for all pixels. Plotting this variance
 617 (σ_{signal}^2) against exposure time yields a straight line (Supporting Fig. S19). We assume that the
 618 noise in the signal is given by $\sigma_{signal}^2 = \sigma_{read}^2 + Dt$, where D is the dark current rate in counts²
 619 per second. We estimate these parameters from the intercept and gradient of a linear regression.
 620 For our conditions, $\sigma_{read}^2 \sim 5.8$ cts² per read, and $D = 0.014$ cts² per second (Supporting Fig.
 621 S19). Our detector has a gain factor (K) of 2.55, meaning that a measure signal of one intensity
 622 unit corresponds to 2.55 electrons. Shot noise can be expressed as $\sigma_{shot}^2 = N_{counts}/K$. As shot
 623 noise scales with \sqrt{counts} (the noise), and Raman signals increase proportionally with counts and
 624 thus t (the signal), the signal to noise ratio from shot noise improves with time following \sqrt{t} , and
 625 errors decreases proportional to $1/\sqrt{t}$. This means for a theoretically perfect detector with no read
 626 or dark noise, the total acquisition time will control the signal to noise ratio, which will be the same
 627 for 10s X 6 accumulations as 1X60s. The relatively low read and dark noise from our detector
 628 means that for strong signals, this is approximately true. For a single 120s accumulation, with a
 629 signal that is 1000 counts above background:

$$630 \quad \sigma_{Total}^2 = \frac{N_{counts}}{K} + \sigma_{read}^2 + Dt = 1000/2.55 + 5.8 + 0.014*120 = 399.6 \text{ cts}^2.$$

631 This yields a signal to noise ratio of 50 ($1000/\sqrt{399.6}$). Shot noise accounts for 98.1% of the error,
 632 dark noise 0.4% and read noise 1.5%. This is an expected result on modern, cooled CCDs where
 633 shot noise dominates (Liu et al., 2018). In contrast, for the same total acquisition time but with n
 634 accumulations, if the spectrum was collected for 30s X 4, the signal would be 4* weaker (250 cts),
 635 yielding the following errors:

$$636 \quad \sigma_{Total}^2 = \frac{N_{counts}}{nK} + \frac{\sigma_{read}^2}{n} + \frac{Dt}{n} = 250/(4*2.55) + 5.8/4 + 0.014*30/4 = 26 \text{ cts}^2$$

637 This yields a signal to noise ratio of 49. Shot noise is 94% of the error, dark noise 0.4% and read
638 noise 5.6%. The small increase in read noise accompanying more accumulations is relatively
639 negligible for strong signals, explaining why the # of accumulations has relatively little influence
640 on the signal to noise ratio, and thus the overall peak fitting error in Fig. 10. However, if we
641 consider now a scenario where the 120s accumulation has a signal only 100 cts above
642 background (e.g. Supporting Fig. S17), a single 120sX1 accumulation has a signal to noise ratio
643 of 14.6 (83% shot noise, 3.6% dark noise and 12.4% read noise) while 30sx4 accumulations with
644 a signal of only 25 cts above background have a signal to noise ratio of 12.5 (61% shot noise,
645 2.6% dark noise and 36% read noise). If there was no read or dark noise, these signal to noise
646 ratios would be the same. However, for very weak signals, the penalty of multiple accumulations
647 and the associated increase in read noise become significant.

648 Overall, these equations explain why peak fitting errors declines with $1/\sqrt{t}$ for strong signals, and
649 there is little difference between different combinations of accumulations and acquisition times for
650 the same total analysis time. However, for the very weakest signals, acquisition times of 30s
651 generate noticeably larger errors (Supporting Fig. S16-17). On our system, because of the decline
652 in error following $1/\sqrt{t}$, we find that analysis times beyond about 4 minutes results in rapidly
653 diminishing returns. Longer acquisition times also run the risk of real drift due to changes in
654 laboratory temperature, Raman laser wavelength, and changing sample and focus positioning.
655 As errors will vary as a function of the Raman hardware and characteristics of the samples being
656 targeted (e.g. density, fluid size etc), we suggest that performing these tests on each specific
657 Raman instrument can help inform an analytical strategy and total acquisition time that reduce
658 peak fitting error while remaining cost and time-effective.

659 3.6 Sample Temperature

660 It is important to investigate the effect of sample temperature on calculated densities. This has
661 implications for how the sample should be heated during Raman analyses. Heating samples
662 above the critical point of CO₂ (31.1 °C) ensures that a homogenous fluid can be analysed, which
663 removes issues associated with analysis of a mixture of gaseous and liquid phases (see DeVitre
664 et al. 2023). We use a low cost heating apparatus consisting of a \$60 DC power supply, a \$41
665 Peltier plate with a hole in it to maintain the ability to use transmitted light for sample navigation,
666 and a \$24 type K thermocouple with a precision of ± 2 °C attached to the surface of the sample
667 with electrical tape (Supporting Fig. S3). However, the limited precision of type K thermocouples
668 means that the sample may be at 35 °C or 39 °C, even if the display reads 37 °C. Determining the
669 error in splitting and density that results from this uncertainty in sample temperature is important

670 to determine whether higher precision thermocouples or a more sophisticated heating apparatus
671 is needed (i.e. a Linkam stage with a precision of ± 0.1 °C).

672 A number of studies have investigated how sample (and thus fluid) temperature affects diad
673 splitting. Kawakami et al. (2003) measured CO₂-rich FIs with densities of ~ 1 g/cm³ in mantle
674 xenoliths and found that splitting stayed remarkably constant between 20-200 °C (see their Fig.
675 10). Thus, they deduce that splitting is sensitive only to density, and is insensitive of temperature.
676 Wang et al. (2011) investigate splitting vs. temperature for a CO₂-rich FI (0.72 g/cm³) and a H₂O-
677 CO₂ FI (see their Fig. 11). Both these FI exhibit a drop in splitting of ~ 0.2 cm⁻¹ between ~ 40 and
678 200 °C, corresponding to a drop in calculated density of 0.08 g/cm³. As well as making
679 measurements on natural FIs with specific densities, Wang et al. (2011) also investigated
680 temperature sensitivity using a high-pressure optical apparatus where the analyte CO₂ gas was
681 held at a controlled pressure and temperature. Their acquisitions at 21, 25, 50, 100, 150 and 200
682 °C indicated that Δ decreases for a given CO₂ density as temperature increases (their experiments
683 span densities between ~ 0.2 and 0.7 g/cm³).

684 More recently, Wang et al. (2019) use a high pressure optical cell to investigate the relationship
685 between splitting and density for a pure CO₂ fluid between -20 and 360°C, and pressures of 0.51
686 to 50 MPa. Their parameterization of how temperature affects splitting is shown in Fig. 11a. We
687 can also calculate the change in density that would occur if the splitting measured on a fluid
688 heated at higher temperatures was converted into density using the 37°C densimeter (Fig. 11d).
689 This is a scenario which may occur if a natural sample is analysed at a different temperature to
690 the densimeter, because of a lack of heating apparatus, an issue with the thermocouple on a
691 heating apparatus, or the heating of the fluid by the laser itself (see section 3.4). The Wang et al.
692 (2019) model predicts that the effect of temperature is largest for densities of ~ 0.5 -0.6 g/cm³ (Fig.
693 11d).

694 Sublett et al. (2021) also performed analyses on pure CO₂ in a high-pressure optical apparatus
695 at 22 °C, 50 °C, 100 °C, 200 °C and 350 °C, for pressures spanning 10-500 bars. Their
696 parameterization (Fig. 11b) is broadly similar to that of Wang et al. (2019), except it shows a larger
697 sensitivity to temperature at the lowest densities, and the change in calculated density vs. true
698 density is more linear with increasing density (Fig. 11c). Finally, Hagiwara et al. (2021, hereafter
699 H2021) investigate the sensitivity of higher density fluids, using pressures from 7-250 MPa for 23
700 °C-200 °C, yielding densities of ~ 0.65 -1.2 g/cm³ (Fig. 11c). This confirms the downward trend in
701 temperature sensitivity towards higher densities indicated by Wang et al. (2019).

702 To assess these model predictions (Fig. 11c), we perform heating tests on natural olivine-hosted
703 FIs with densities spanning 0.4-0.9 g/cm³ (Fig. 11 d-i). Double polished wafers were placed within
704 a LinkamTHMS600 stage and spectra were acquired with the TruePower meter set to 6 mW for
705 temperatures between 37 °C and 70 °C. Over this temperature window, thermal expansion of the
706 olivine induces negligible volume and thus density changes (<0.1 % following the approach of
707 Hagiwara et al., 2018; see also Yamamoto and Kagi, 2008). To account for the possibility of
708 instrument drift, analyses were run with increasing temperature and then decreasing temperature
709 (2 repeats made at each temperature). We fit a linear regression to temperature vs. apparent
710 density calculated using the UCB densimeter, and extrapolate this regression to calculate the
711 density at 37 °C (Fig. 12). For each acquisition, we then subtract this calculated 37 °C density from
712 the density calculated from the measured splitting using the 37 °C densimeter, and compare this
713 to the prediction of the three models (Fig. 11d, Fig. 12).

714 As predicted by all three models, heating causes the 37 °C densimeter to underestimate the true
715 density, and the density change is largest for FI with densities of ~0.4-0.7 g/cm³, and smallest for
716 ones at 0.9 g/cm³. The Wang et al. (2019) model does a very good job of predicting the density
717 change for 6 fluid inclusions (Fig. 11d, Fig. 12).

718 These temperature-density gradients can be used to assess the uncertainty introduced by
719 measuring the sample temperature. The ±2 °C error introduced from the thermocouple plus offsets
720 of 1-2 °C we observed based on the type of tape used to stick the thermocouple down and the
721 exact positioning of the thermocouple contributes an additional error of ~±0.0016-0.0048 g/cm³,
722 (depending on the density of the fluid being analysed). For FG04-A1-A4, the average analytical
723 error calculated in DiadFit is ~0.004 g/cm³, while the standard deviation of hundreds of
724 acquisitions over many years is 0.0085 g/cm³. Assuming the worst-case scenario of temperature
725 contributing the full 0.0048 g/cm³, this leaves a residual error of 0.0057 g/cm³ unaccounted for.
726 The remaining variability may reflect subtle variations in laser wavelength with time, laser
727 positioning within the sample, and temperature drift which is not fully accounted for by Ne line
728 corrections (e.g. Fig. 6). However, to a first order estimate, peak fitting errors are a good estimate
729 of the uncertainty of Raman analyses within a factor of two, and are certainly not several orders
730 of magnitude too small (c.f. Bakker, 2021).

731 3.7. Laser Power

732 These heating tests give us a framework to assess the effect of laser power on CO₂ density. This
733 has been a topic of extensive debate in the literature. Fall et al. (2011) suggest that the laser
734 doesn't heat the inclusion more than a few 10ths of a degree, based on observations that many

735 inclusions analysed in their laboratory which homogenized during microthermometry at
736 temperatures very close to laboratory temperature still showed distinctive liquid and gaseous
737 phases when analysed by Raman spectroscopy. However, Hagiwara et al. (2021b) simulate laser
738 heating with a numerical model, concluding that even laser powers of 10 mW can cause the
739 temperature to rise tens or hundreds of degrees. Specifically, they show that the absorption
740 coefficient of the host mineral is the most important parameter for determining the amount of
741 heating, because the CO₂ inclusion itself has very little absorption at the wavelengths of the laser.
742 They perform Raman analyses at varying laser powers, and use a thermometer based on hot
743 bands to determine the laser heating coefficient B (°C per mW of laser power) for FIs in olivine,
744 orthopyroxene, clinopyroxene, spinel and quartz. It is notable how much lower their B values are
745 for quartz (0.010-0.014 °C/mW) vs. mafic minerals such as olivine (0.02-1 °C/mW). In their tests,
746 as laser power is changed from 4-14 mW, the olivine FI shows a ~0.02 g/cm³ decrease in density
747 while the quartz inclusion shows no change outside uncertainty. Thus, while Fall et al. (2011) do
748 not note the host mineralogy of the inclusions they draw their conclusions based on, the inclusion
749 shown in their figure that they reference as evidence for a lack of laser heating is hosted in quartz.
750 If the Fall et al. (2011) observations are from quartz inclusions, they are not in conflict with the
751 results of Hagiwara et al. (2021b) who predict minimal heating for quartz. DeVitre et al. (2023)
752 also demonstrate the effect of laser heating on calculated CO₂ densities through the analyses of
753 olivine-hosted melt inclusion vapour bubbles with densities of ~0.3, 0.5 and 0.6 g/cm³. They show
754 a clear drop in splitting (and thus density) with increasing power (~0.1 g/cm³ from 0-20 mW using
755 TruePower readings).

756 It is vital to determine the effect of laser heating in different samples and bring some consensus
757 to the literature, because it is very tempting to use high laser powers for analysis to obtain strong
758 signals and small peak fitting errors. For example, Supporting Figure. S20 shows the peak fitting
759 error associated with the analysis of FIs and vapour bubbles using laser powers ranging from ~5-
760 35 mW. A laser power of ~5 mW yields a peak fitting error 2-3X higher than a laser power of 35
761 mW. This reduction in error is highly beneficial for low density samples like melt inclusion vapour
762 bubbles (often 0.02-0.04 g/cm³), where peak fitting errors can introduce an uncertainty of 20-50%
763 in density, but this is only true if the higher laser powers do not change calculated density.

764 We perform analyses using TruePower values of ~5 to 50 mW (scaling to <0-30 mW at sample
765 depending on the objective) for 4 olivine-hosted melt inclusion vapour bubbles from Mauna Ulu,
766 Kīlauea (Wieser et al., 2019) and the Cascades (Wieser et al., 2023), and 29 olivine-hosted FIs
767 from Iceland (Kahl et al., 2025), Fogo volcano (DeVitre et al., 2023b), and American Samoa (kindly

768 supplied by Drew Downs). We also analyse 3 pyroxene-hosted FI from the Galápagos (Gleeson
769 et al., 2025). These MI and FI span a wide range of CO₂ densities (Fig. 13). Microthermometry
770 measurements and broad wavelength scans performed by the studies cited above demonstrate
771 that the vast majority of these inclusions are pure CO₂ (e.g. Fogo and Iceland, DeVitre and Wieser,
772 2024; Kahl et al., 2025). Some of the Galápagos fluid inclusions show anomalously low freezing
773 temperatures but no other Raman-active peaks, so are inferred to contain small amounts of
774 Raman-inactive noble gases (Gleeson et al., 2025). We also perform power series on the FI and
775 melt inclusion vapour bubbles that were used for heating series (Fig. 14). As for temperature, we
776 analyse each FI with increasing then decreasing laser power.

777 We also conduct a series of tests to determine if a steady state is reached, given that Hagiwara
778 et al. (2021b) show that the time it takes for an inclusion to reach 90% of its max temperature
779 upon laser heating is 18.5s, and that this time is heavily dependent on the boundary condition.
780 Using the WITec slow series function where the laser remains on during the entire interval, we
781 collect 15 spectra spaced between 5s and 30s, and fit each of these spectra individually. There
782 are no significant and consistent decreases in CO₂ density or changes with increasing # of
783 acquisitions (Supporting Fig. S21). The power series show the area ratio of the sum of the two
784 hotbands relative to the sum of the two diads increases with increasing laser power, providing
785 evidence that the laser is increasing the internal temperature of the inclusion (Supporting Fig.
786 S22). There is no notable change in the hotband to diad ratio during the slow series acquisitions,
787 also indicating that a steady state is reached quickly, and doesn't affect our conclusions regarding
788 heating rates using our typical 45s x5 accumulations.

789 The vast majority of fluid inclusions show a notable drop in CO₂ density as the laser power is
790 increased (Figure, 13, Supporting Fig. S23-26). We calculate the gradient of a linear regression
791 between the laser power at the sample and CO₂ density, and find highly variable gradients for
792 different samples (~0 to -0.006 g/cm³ per mW, Fig. 13a). In general, FIs show weaker negative
793 gradients than melt inclusions from the same samples (see also Supporting Fig. S27). However,
794 FIs show great diversity in gradients. One fluid inclusion (FI 101C from Fogo volcano) shows a
795 very large change in density (~0.95 to 0.8 g/cm³, Fig. 13d). This corresponds to a change in
796 calculate depth of 8 km in the crust. The three Icelandic FI shown in Fig. 12c-e show density
797 change varies from -0.08 g/cm³ to -0.018 g/cm³ for a 30 mW change in laser power. The drop in
798 density for B22_1_FIA corresponds to a change in calculated depth of ~ 3 km (Fig. 12c, RH axis).
799 These changes in density are far larger than the uncertainty on each Raman measurement (see
800 error bars, Fig. 12c-e).

801 Hagiwara et al., (2021b) suggest that the absorption coefficient of the host phase is the most
802 important parameter affecting laser heating. A wide variety of absorption coefficients for olivine
803 have been reported in the literature, making it hard to quantitatively model the heating process.
804 For their laser heating simulations, Hagiwara et al. (2021) use an olivine absorption coefficient of
805 1.14 cm^{-1} based on measurements of a hydrothermal olivine from the Cola Peninsula, Russia
806 (Fo_{95} Taran and Matsyuk, 2013). In contrast, Chassé et al. (2015) report an olivine absorption
807 coefficient at wavenumbers corresponding to a 532 nm green laser ($\sim 18800 \text{ cm}^{-1}$) of 11 cm^{-1} (the
808 Fo content is not given). Other papers have shown that the absorption coefficient of olivine is
809 sensitive to the Fo content, and the crystallographic orientation of the host (Hazen et al., 1977;
810 Taran and Matsyuk, 2013; Ullrich et al., 2002). For example, at 532 nm, a Fo_{90} crystal has an
811 absorption coefficient of $\sim 6 \text{ cm}^{-1}$ parallel to a vs. 3 cm^{-1} parallel to b (Ullrich et al. 2002). For Fo_{10} ,
812 the coefficient parallel to the b axis is $\sim 97 \text{ cm}^{-1}$ and 76 cm^{-1} parallel to c. For comparison, Chassé
813 et al. (2015) show absorption coefficients for basaltic glass and melt inclusions of $\sim 100 \text{ cm}^{-1}$ at
814 532 nm. The order of magnitude higher absorption coefficient of basaltic glass vs. typical igneous
815 olivine forsterite contents (70-90) account for the higher rates of heating seen in melt inclusions
816 vs. fluid inclusions.

817 We investigate the relationship between heating rates and host forsterite contents for a subset of
818 Icelandic and American Samoa crystals that have had the host chemistry analysed, given the
819 expected increase in absorption coefficient as the olivine becomes more Fayalitic. Our data shows
820 a strong correlation between the measured gradient and the Fo content (Fig. 13b), indicating that
821 the absorption coefficient of the host is strongly controlling the amount of heating. However, olivine
822 chemistry cannot explain all the variability in laser heating rates. Crystal 101 from Fogo volcano
823 contains three FI in the same secondary trail within $\sim 100 \mu\text{m}$ of each other. FIC experiences a
824 very dramatic drop in density, from $\sim 0.93 \text{ g/cm}^3$ to $\sim 0.75 \text{ g/cm}^3$ over 30 mW of heating (gradient
825 of -0.0053 g/cm^3 per mW). In contrast, FIB experiences no change in density outside of uncertainty
826 (gradient of -0.0001 g/cm^3 per mW), and FIA experiences a minor change, with a gradient of $-$
827 0.001 g/cm^3 per mW (Supporting Fig. S26). Hagiwara et al. (2021b) show that in addition to
828 absorption coefficient, the host crystal thickness and radius can also affect heating rates, with
829 higher heating coefficients for larger crystals, and lower heating rates for larger host radius.
830 However, these parameters are also the same for these three fluid inclusions. We initially
831 speculated that the extreme heating experienced by FIC may result from its proximity to the
832 surface ($\sim 6 \mu\text{m}$ depth). However, when the crystal was flipped over and a power series was
833 performed from the other side ($\sim 68 \mu\text{m}$ depth), the gradient was almost identical (Supporting Fig.
834 S30). However, 3D mapping reveals that 101C contains carbonate (Supporting Fig. S31). It has

835 been suggested that microparticles which are not visible optically can enhance the laser heating
836 effect (e.g. Burke, 2001), so it is possible that carbonate is playing this role in this inclusion. The
837 two Icelandic samples that show higher heating than expected based on their Fo contents (Fig.
838 12b) also contain notable carbonate in 3D maps. However, of the three FI from American Samoa
839 (pink edges, Fig. 12b), the two which follow the gradient-Fo trend defined by the Icelandic samples
840 contain abundant carbonate, while the one with more negative gradients contains no carbonate.
841 The three Mg# 79 pyroxenes also show very variable heating amounts (Supporting Fig. S32),
842 despite the fact these samples are very homogenous and there are no apparent color differences
843 between crystals (Gleeson et al., 2025).

844 As predicted from the modelling of Hagiwara et al., (2021b), we also note that the gradient
845 depends on the heating apparatus used, which controls the boundary condition by which the
846 crystal loses heat. When FI 101C is heated when embedded in crystalbond on a glass slide, it
847 experiences an even greater heating rate (gradient of $\sim 0.008 \text{ g/cm}^3/\text{mW}$) than when placed on
848 the Linkam stage (Supporting Fig. S31). Overall, these power series demonstrate that while laser
849 heating can be partially predicted based on the absorption coefficient of the host, the extreme
850 change in density experienced by very specific FI, replicated across different analytical sessions
851 in our lab, is worrying. Given that many melt and FIs are analysed prior to knowing the Fo content
852 (as that is often the final step of measurement) and that we cannot diagnose why a few FI show
853 such extreme heating, this work demonstrates that it is absolutely vital to use low laser powers,
854 as the heating response is currently not entirely predictable. Thus, we recommend each
855 laboratory investigate laser heating using similar tests to those performed here for new sample
856 suites to determine the correct laser power to use, as the heating response is currently not
857 predictable.

858 Performing laser heating and laser power tests on the same FI within the Linkam stage allows us
859 to estimate the heating coefficient, B ($^{\circ}\text{C}/\text{mW}$ of laser power) for this specific apparatus; B values
860 will likely vary depending on the exact mounting mechanism used and the boundary conditions
861 they induce. For each FI, we subsample the power series to be within the range of densities seen
862 in the heating series, and calculate the gradient of density vs. temperature, and density vs. laser
863 power. On Fig. 14 (and Supporting Fig. S34), we adjust the power and temperature scale such
864 that the gradients of the two datasets align. We calculate B as:

$$865 \quad B \text{ (}^{\circ}\text{C}/\text{mW}) = \frac{\text{Grad} \left(\frac{\text{g}}{\text{cm}^3} / \text{mW} \right)}{\text{Grad} \left(\frac{\text{g}}{\text{cm}^3} / ^{\circ}\text{C} \right)} \text{ [Eq 2]}$$

866 B value errors were calculated from the error on the gradient of each linear regression calculated
867 from ordinary least squares regression from scipy stats linregress propagated in quadrature. B
868 values mostly span ~0.3-0.7 °C/mW, although a few exceed 1 (Supporting Fig. S34). This
869 compares to 0.02-1°C/mW from Hagiwara et al. (2021), and ~0.36 °C/mW from Wang et al. (2024).

870 3.5. Other strategies to optimize signal.

871 Having to use low laser powers to minimise heating is a real issue in terms of getting enough
872 signal to minimise peak fitting errors. In addition to optimizing the acquisition times (Fig. 10), it is
873 worth thinking about other ways to optimize signals. Our repeated measurements of FI standards
874 (Fig. 5) not only allow us to track the stability of the densimeter, but also the intensity of signals
875 with time. If the FI standard returns lower peak heights than normal, we analyse a Si chip which
876 yields ~2000 cts for a 10s acquisition when the system is well aligned. We use the oscilloscope
877 feature ('live spectra' on HORIBA) to perform micro adjustments on the pinhole position to
878 enhance laser transmission through the fibre to achieve these counts. We find that these
879 adjustments can easily increase signals by a factor of 2-3. We typically adjust the system every
880 few weeks, although sometimes it goes months with no decline in signal, and other times we have
881 to readjust after a few days. Typically, after HVAC failures where the temperature of the laboratory
882 has fluctuated, adjustments are needed, suggesting this adjustment is required based on thermal
883 expansion/contraction of components. On a mirror-based instrument, mirror alignments will have
884 a similar effect.

885 Once the system's optics are optimized, it is worth considering the analytical geometry. The depth
886 of the FI within the sample has been discussed in a variety of papers. Frezzotti et al. (2012) note
887 that the depth resolution of Raman spectroscopy degrades with depth in the sample, so it is best
888 to analyse FI within 30 µm of the surface. Burke (2001) states that '*usually one strives to analyse
889 inclusions as near the surface as possible' inclusions deeper than 50 µm start to give problems,
890 especially if their density is not too high*'. To assess the influence of FI depth, we individually
891 ground 3 olivine-hosted FI down, performing analyses after each round of grinding and polishing.
892 We measure the depth by focusing on the surface of the sample in RL, and then focusing down
893 to the melt inclusion. We correct the movement of the Z stage to true depth accounting for the
894 refractive index of olivine ($N_{\text{olivine}} \sim 1.7$, 1.64-1.67 for Fo, 1.84-1.89 for Fa).

895 True depth = $N_{\text{olivine}} Z_{\text{stage movement}}$ [Eq 3]

896 We use the WITec autofocus function at each depth. Surprisingly, we find no coherent trends in
897 signal intensity (quantified here using the area of diad 2) or the error on calculated density with

898 increasing depth in the sample (Fig. 15a, Supporting Fig. 35-37). However, we do find big changes
899 with minor changes in focus at a single sample depth. To demonstrate this, we perform
900 acquisitions at different focuses when the FI was 91.8 μm deep in the sample. We first find the
901 optical focus yielding a sharp inclusion edge (yellow box and spectra, Fig. 15). We then use the
902 WITec autofocus feature at our maximum laser power. This collects spectra for ~ 3 s at various
903 depths to find the maximum intensity in user-selected spectral window. The z position was ~ 2 μm
904 close to the surface than the best optical focus and yielded very similar signal strengths (black
905 spectra and box). We then used the oscilloscope function to manually scroll through Z positions
906 to find the best signal (pink box and spectra). The 6 focus positions between -14.1 μm and 0 μm
907 all yield optically reasonable images, and yet the signal strength varies by a factor of 3.
908 Interestingly, the blue focus position ($+2.3$ μm) yields an optically poor image, but still a better
909 focus than the oscilloscope feature found. This comparison clearly demonstrates that focus has
910 far more effect on spectral strength than changing the depth in the sample by ~ 250 μm (Fig. 15
911 c-d).

912 This result was replicated for several inclusions (Supporting Fig. S35-S37). We also find that the
913 signal can vary quite considerably spatially, and small shifts in x-y position can also result in
914 stronger signals. This becomes very apparent when taking 2D transects through and across FI
915 (Supporting Fig. S38-39). We suggest that when signals are weak, the autofocus or oscilloscope
916 feature should be used to investigate different x-y-z positions around the optimal focus. We
917 perform these tests at high laser power (49 mW) for speed, and then drop the power to perform
918 the analyses at the optimal Z position. Variability in signal strength can also be seen by performing
919 2D and 3D maps of signal intensity.

920 Conclusions

921 In summary, we investigate the effect of several instrument parameters and data reduction
922 methods on the accuracy and precision of Raman analyses. Our findings can be summarized as
923 follows:

- 924 1. Spectral non-linearity means that the spectral window/center and drift correction method
925 (e.g. choice of Ne and Ar lines) used for densimeter calibration must be replicated when
926 analysing unknowns in future sessions. Spectral resolution alone cannot account for the
927 diversity in published densimeters, shown by the fact even a single instrument can
928 generate a wide variety of densimeter relationships. We instead speculate that densimeter
929 diversity in the literature is likely related to non-linearity, causing the splitting to vary
930 depending on which CCD pixels the CO_2 diad is projected onto.

- 931 2. Estimates of peak fitting error from pseudovoigt curves, combined with uncertainty in the
932 Ne line correction model, account for approximately half the scatter seen in repeated
933 Raman analyses of the same inclusion over ~ 2.5 years (~0.004 g/cm³ for FI with densities
934 of >0.6 g/cm³). These low uncertainties make suggestions of very large peak fitting errors
935 untenable. The additional 0.004 g/cm³ of scatter seen in repeated standard acquisitions
936 likely results from uncertainty in the temperature of the sample and room, combined with
937 variations in laser wavelength and other parameters with time.
- 938 3. We recommend tracking the density of a single standard during every run to ensure there
939 has been no drift from the laboratory densimeter relationship. These standards also help
940 to track the efficiency of the instrument (e.g. checking the peak heights are as high as
941 previous days), so the user knows when re-alignment is required.
- 942 4. In a temperature-controlled laboratory, drift during and between analytical sessions is
943 minimal using an WITec alpha300R Raman. However, when the laboratory temperature
944 is perturbed, the Ne line correction helps to reduce instrument drift by a factor of ~3. Thus,
945 while it is better to have a Raman in a temperature-controlled laboratory, it is not essential.
946 Large temperature swings of +5 °C result in an additional uncertainty of <0.01 g/cm³ if Ne
947 lines are used for drift correction.
- 948 5. Laser heating is a significant issue, particularly for materials with higher optical absorption
949 coefficients (e.g., silicate glass, lower Fo olivine). However, variable absorption
950 coefficients alone are not enough to explain the large variability in heating response seen.
951 <4-8 mW of laser power at the sample seems a good way to minimise heating for all but
952 the most strongly laser-heated samples, while still providing sufficient signal to obtain good
953 peak fits.
- 954 6. To help counteract the relatively weak signals achieved using low laser powers, the analyst
955 should investigate the play off between uncertainty and acquisition time for their specific
956 instrument. They should experiment with different focuses using an autofocus/live spectra
957 function of the oscilloscope at high power to find the optimum signal within the fluid. Focus
958 has a far larger effect than depth to the fluid within the sample (even at depths of ~250
959 μm).

960 We suggest that prior to performing density measurements in natural samples with a Raman
961 spectroscopy set up, users conduct the various tests here to optimize analytical conditions:

- 962 1) Choose a spectral window over which to collect data, and if for any reason you want to
963 change this after calibration, investigate the impact of this change on the density of a
964 natural or synthetic FI
- 965 2) Use a temperature logger to explore variations in laboratory temperature. If the
966 temperature fluctuates by several degrees, perform repeated analysis of a natural or
967 synthetic fluid inclusion, and test whether the selected drift correction method (Ne lines or
968 repeated standard measurements) minimises this drift or not.
- 969 3) Investigate the relationship between laser power and signal using the available power
970 options, and between analytical errors and acquisition time and accumulations
- 971 4) For the samples of interest (with their different absorption coefficients), investigate the
972 effects of laser heating to choose a laser power that gives sufficient signal without inducing
973 density changes.

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983 greatly improved the manuscript.

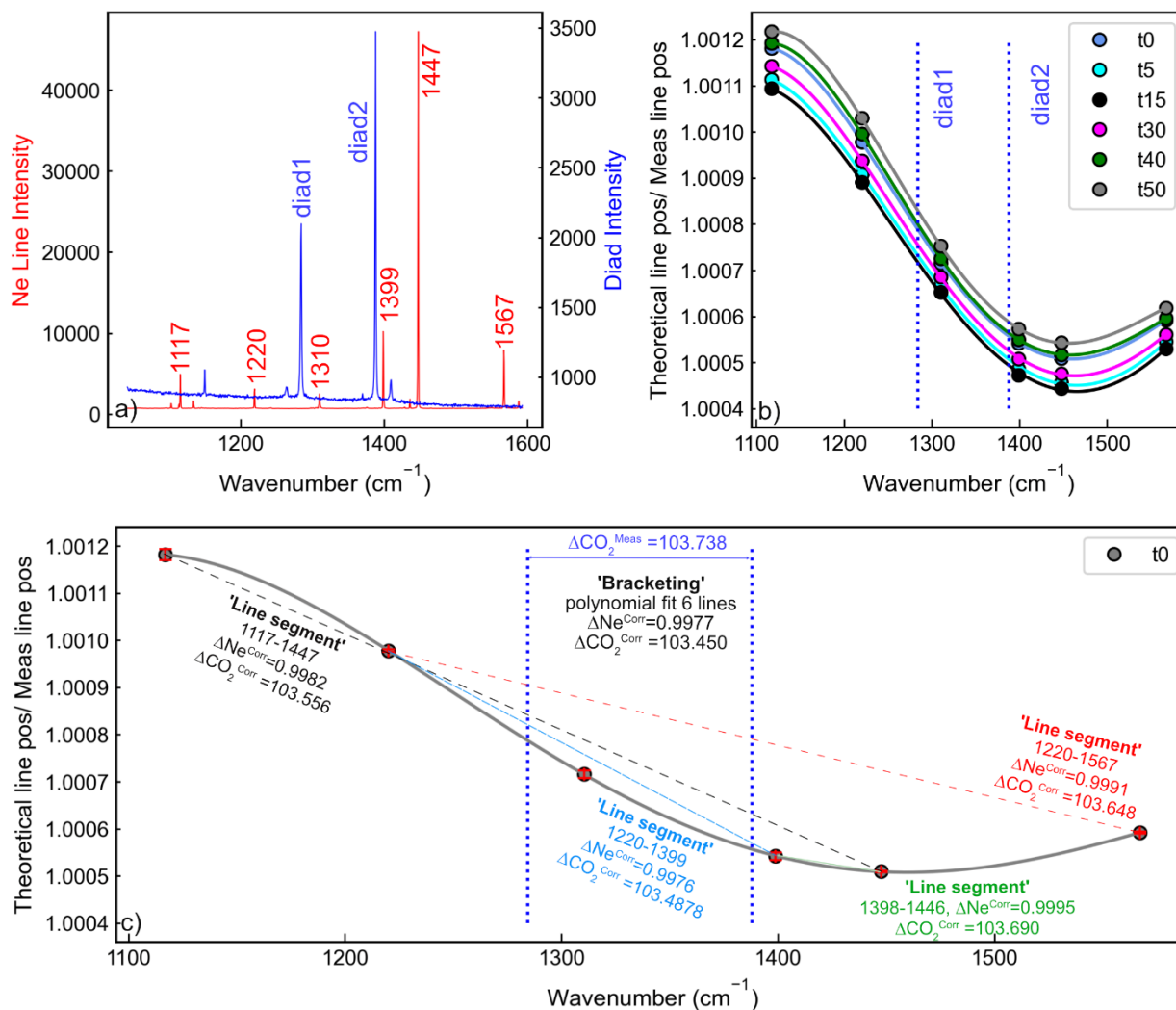
984 Supporting Information

985 All data and Jupyter Notebooks needed to recreate the figures in this study are provided in the
986 Github repository. We also supply a supporting spreadsheet containing processed data for each
987 figure (e.g. peak fits etc).

988 [PennyWieser/Raman_Best_Practices: Code for Raman best practices paper by Wieser et al.](#)

989 Author Contributions

990 PW and CLD conceived the project. PW, CLD and IS collected the Raman data and performed
991 peak fitting. PW wrote the manuscript. All authors contributed to the discussion of ideas and the
992 editing of the manuscript.



994

995 Figure 1 – Examining the difference between the Line segment method and the bracketing method

996 when correcting instrument drift using Ne lines. a) An example Ne line acquisition compared to a

997 diad. Note that the intensity of the Ne (left hand axis) is almost 10X stronger than the diad (right

998 hand axis). This is different from Fig. 8 of Bakker 2021 where Ne lines show very low intensities.

999 b) For repeated acquisitions of Ne lines through the day, the actual peak position is compared to

1000 the theoretical peak position. For each line acquisition, a polynomial has been fitted through the

1001 data. c) For the first acquisition (t0), we show the polynomial fit through the 6 Ne peaks. The

1002 'Bracketing technique' would correct each diad position based on the value for the polynomial at

1003 that wavenumber (where the blue dotted lines intersect the grey line). The 'Line segment' method

1004 will give similar results if the gradient drawn between two points is parallel to the intersection of

1005 the dotted blue and grey lines.

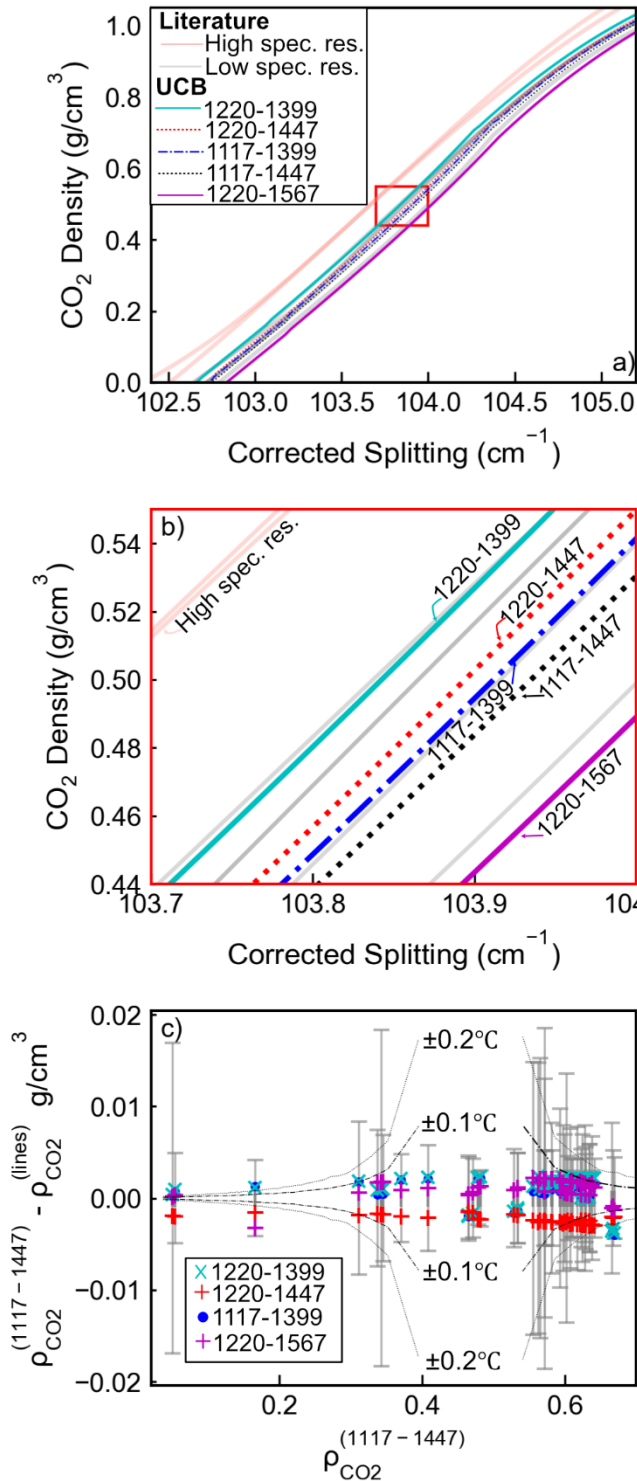


Figure 2 – Assessing the effect of different Ne line reduction routines on densimeters and densities calculated for natural FIs. a) The relationship between corrected splitting and CO_2 density (i.e. the densimeter) was determined using 5 different pairs of Neon lines for the same UCB calibration dataset. This variation is compared to other literature densimeters (grey = high precision, salmon – low precision based on the definitions of Remigi et al. 2021). b) Inset of area in red box in a). c) Natural FI data from DeVitre and Wieser, (2024) was reduced using these different Neon correction routines and densimeters. The y axis shows the difference in density for the default 1117-1447 line option compared with other data reduction strategies. The grey error bar shows the uncertainty on the 1117-1447 density. This demonstrates densities are within uncertainty as long as spectra from unknowns and calibration data are drift-corrected in the same way. For reference, we overlay the uncertainty associated with Microthermometry for ± 0.1 and 0.2°C in sample temperature.

1035

1036

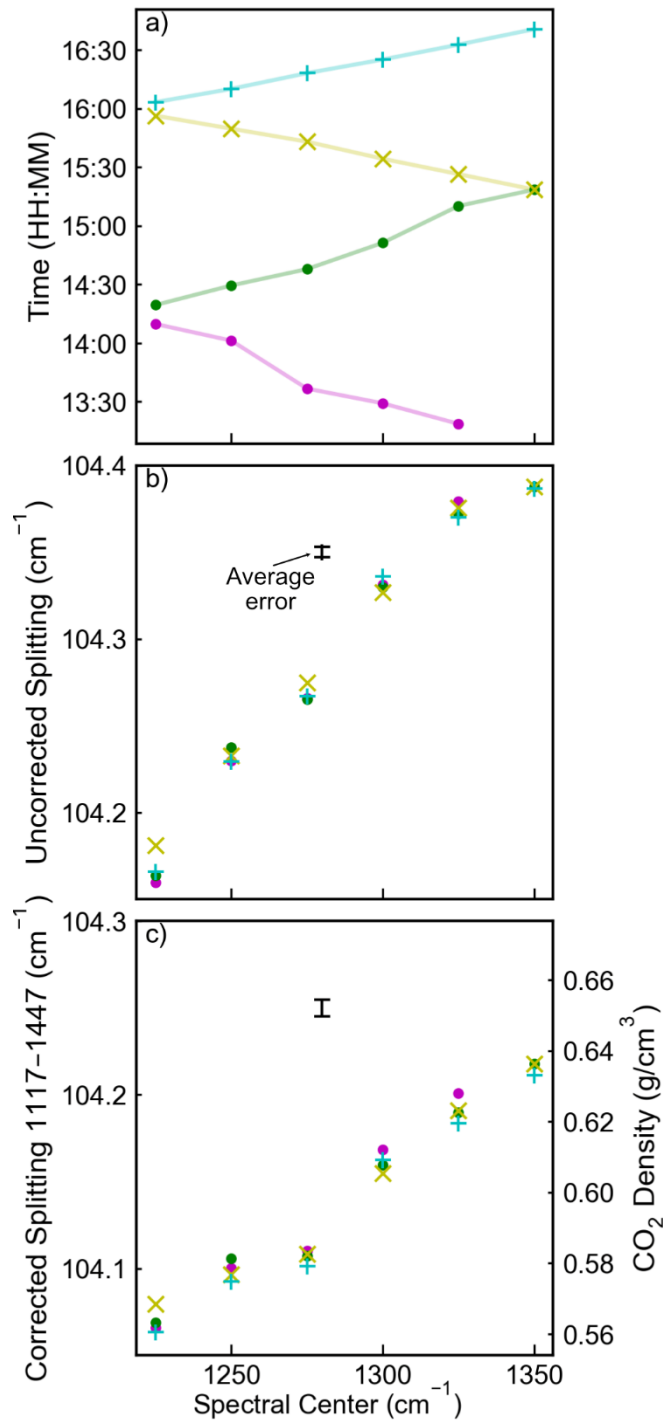


Figure 3 – Change in splitting and calculated density for a single FI as a function of the spectral center. a) The four symbol colors show ‘up’ and ‘down’ cycles changing the spectral center as a function of time through the day. b) Uncorrected splitting vs. spectral center and c) corrected splitting, using Ne lines collected at the same spectral center as the diad. The black error bars show a representative error for the splitting (b) and corrected splitting (c) respectively (offset from trend for clarity). The density axis in c is calculated using the densimeter developed for a spectral center of 1325 cm^{-1} .

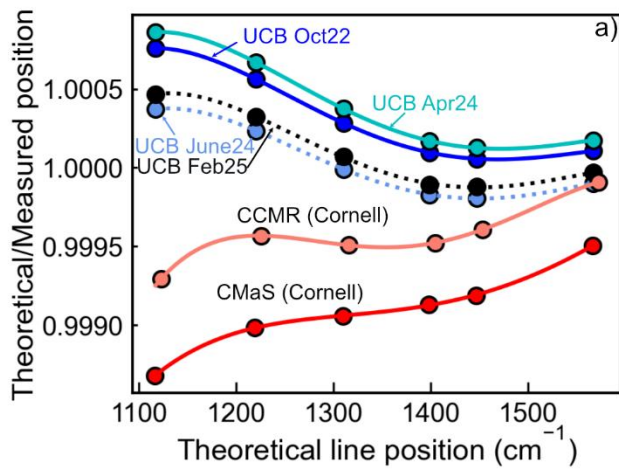
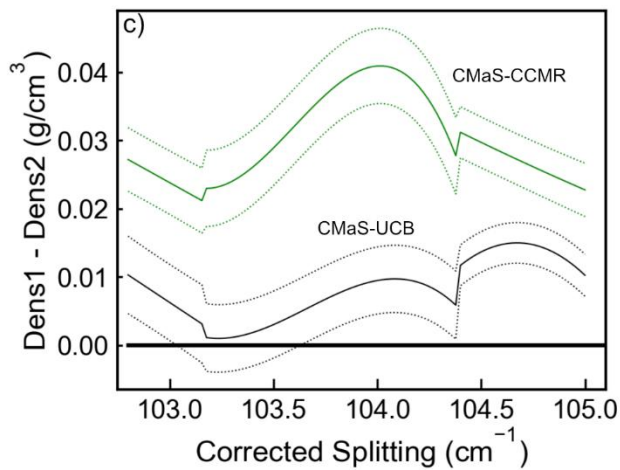
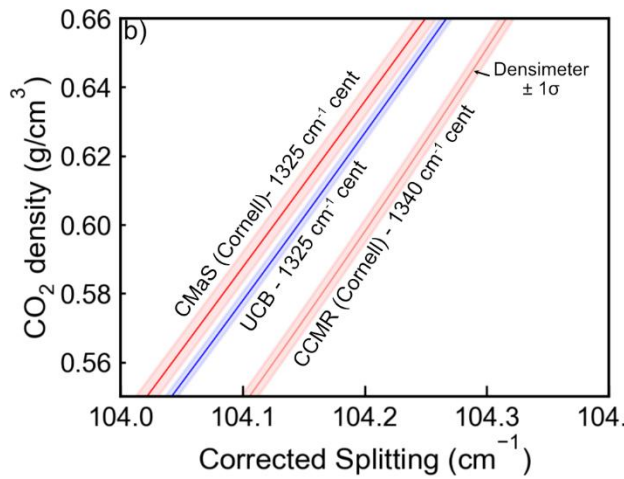
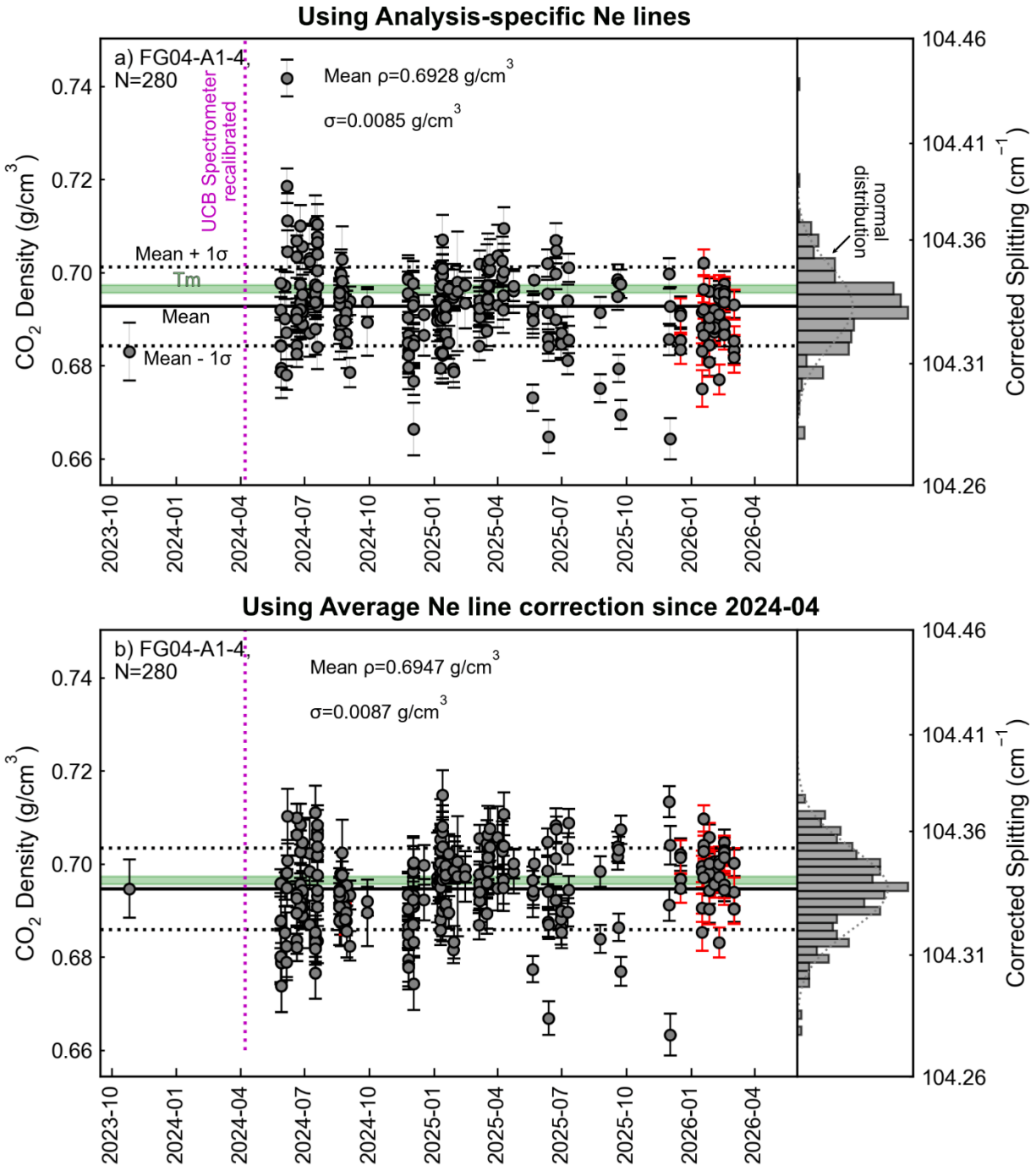


Figure 4 – a) Comparison of spectral non-linearity determined from Ne line acquisitions for three different WITec alpha300R instruments. b) Comparison of the densimeters across a smaller density range showing that UCB and CMaS are more similar than CCMR and CMaS, despite CCMR-CMaS showing much more similar spectral non-linearities (a). c) Difference between densimeters as a function of splitting. The filled lines on b) show the 1 σ errors on the densimeter fit, combined in quadrature and shown as dashed lines in c).

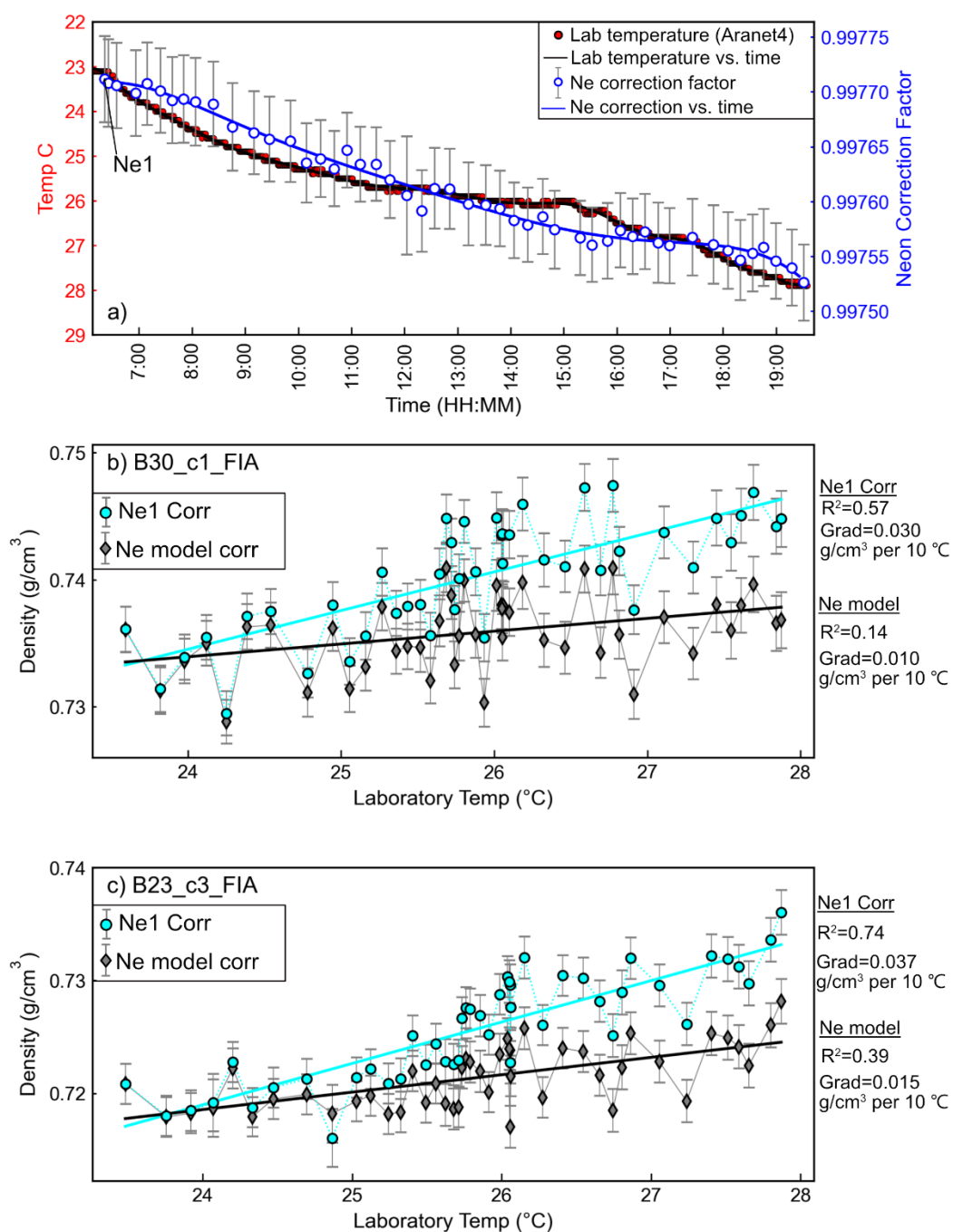




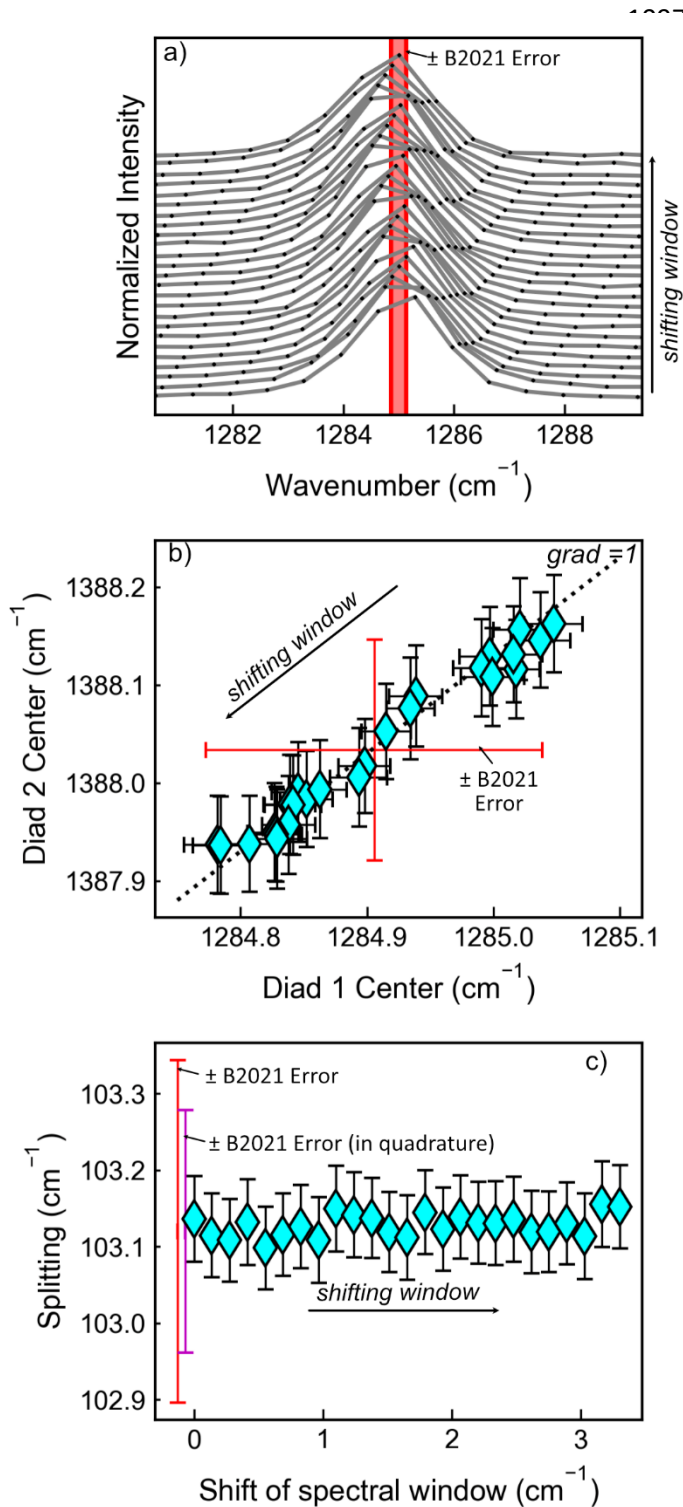
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1082 Figure 5- Long-term variability in standard value for FI FG04-A1-4 collected at UCB. a) Data
 1083 reduced using a Ne drift correction model for each analytical session. b) Data reduced using the
 1084 average Ne line correction factor for all standard acquisitions after the spectrometer recalibration
 1085 (marked with a magenta line). The green bar shows the density determined by Microthermometry,
 1086 accounting for the uncertainty in the temperature reading of the stage. The black solid line shows

1087 the mean Raman value for the standard, and the dotted black lines $\pm 1 \sigma$. The distribution of
 1088 densities is shown as a histogram, compared to a normal distribution based on the mean and
 1089 standard deviation. Supporting Fig. S10 shows the same plot for a second standard run less
 1090 frequently at UCB. Grey error bars use unweighted least squares, red error bars used weighted
 1091 least squares routines from newer versions of DiadFit.



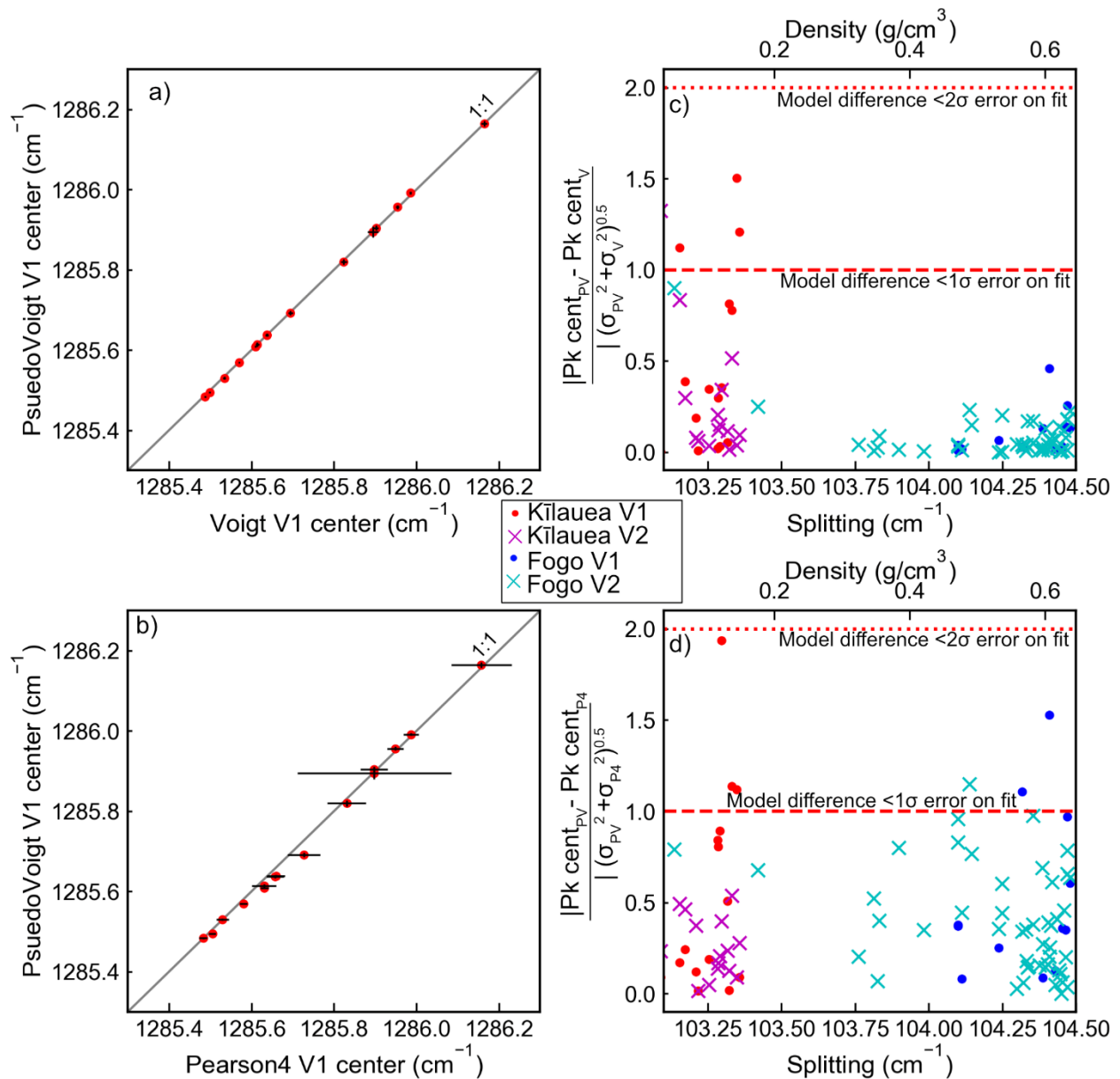
1093 Figure 6 – Testing the sensitivity of Raman acquisitions to large fluctuations in room temperature.
 1094 a) Temperature log from an Aranet4 logger (red dots) against time. The white circles with blue
 1095 dots show the Ne correction factor calculated for each Ne line acquisition. These series were fitted
 1096 with a regression (black and blue lines) to calculate temperature and Ne correction factor for each



FI acquisition. b) Change in density of FI B30_c1_FIA as a function of laboratory temperature. The grey diamonds are corrected based on the Ne line correction model (blue line in a), while the cyan circles use the correction factor for the first Ne acquisition (to replicate not applying a drift correction through time). The straight lines represent linear regressions, with the statistics shown on the right. c) Same for B23_c3_FIA. Error bars from weighted least squares fits.

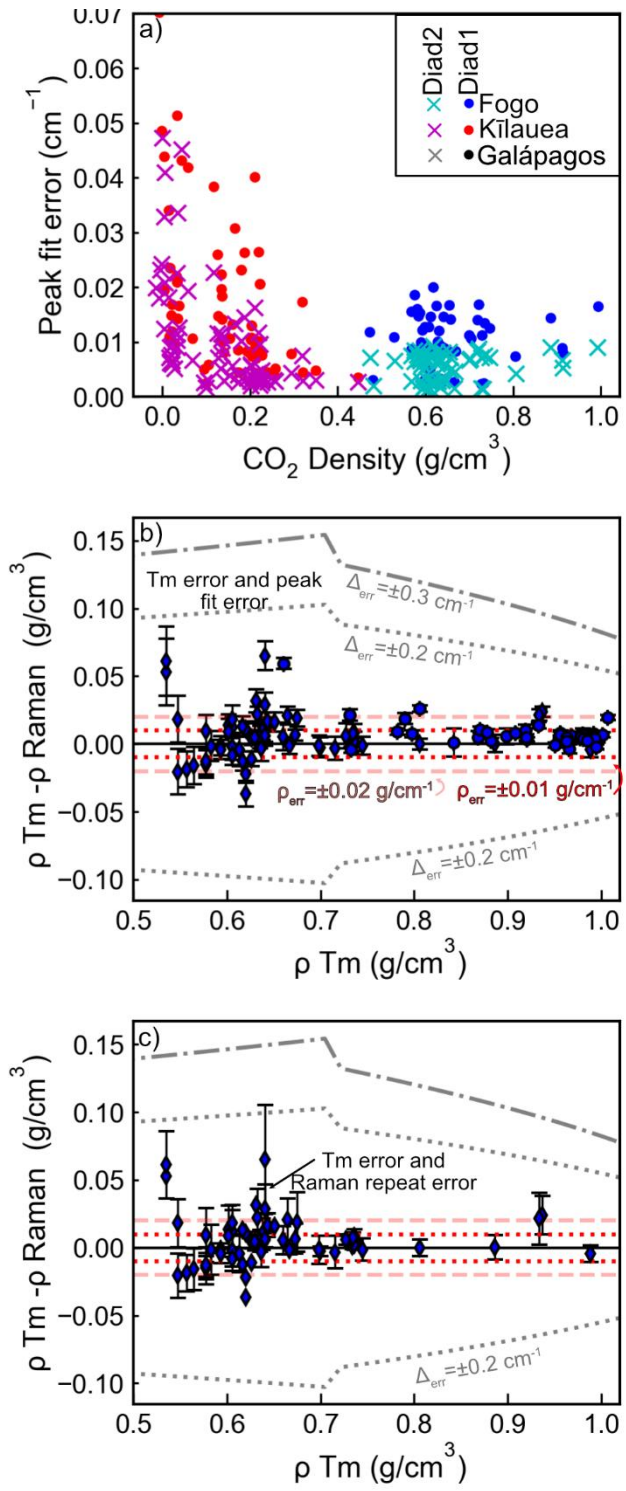
Figure 7 a) Stacked spectra collected as the spectral window was progressively shifted by Bakker (2021). The red band shows their stated error on the peak position, based on the wavenumber shift in the window. b) Plot showing the changing position of diad1 and diad2 for each acquisition shown in a). The black dashed line has a gradient of 1 to emphasize that diad1 and diad2 shift by equal amounts. The error bar shows the 1 sigma error estimated by lmfit from the peak fit c) Change in splitting as a function of the shift of the window. Errors estimated by Bakker (2021) are shown in red, and purple shows their stated error on each peak combined in quadrature. The black error bar

1126 shows the uncertainty based on the Imfit peak positions propagated in quadrature using a
 1127 weighted least squares routine.



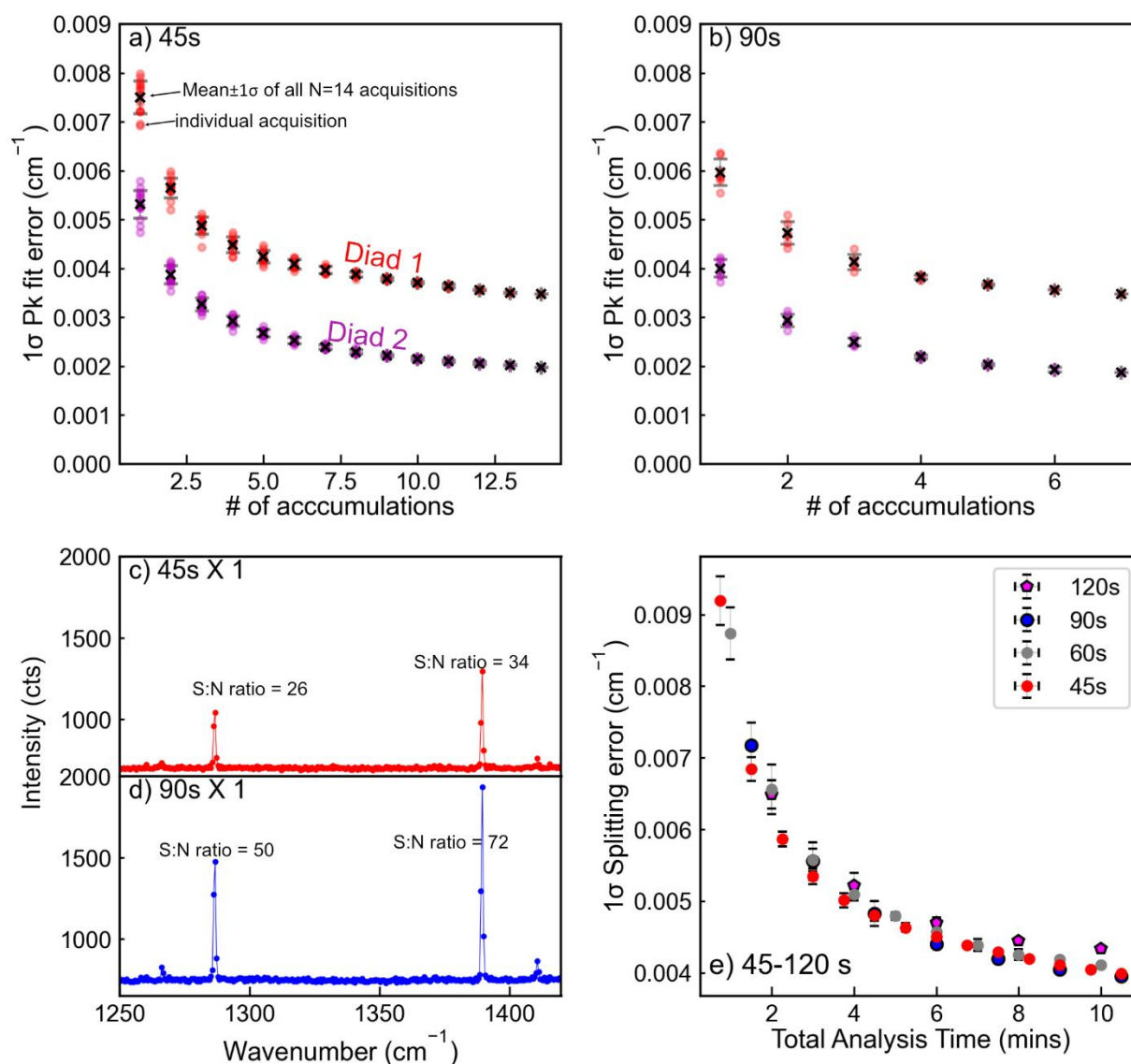
1128
 1129 Figure 8 - Comparison of peak centers obtained with Voigt (V), Psuedovoigt (PV) and Pearson4
 1130 (P4) peak fits for Diad 1 (V1) and Diad2 (V2) for natural FIs from Fogo and Kīlauea Volcano. a-b)
 1131 Comparison of Diad 1 position for Kīlauea FI, with errorbars showing the 1σ error on each peak
 1132 fit from Imfit. c-d) The absolute difference between the peak center for each model is divided by
 1133 the error on each model combined in quadrature. For datapoints lying below 1, the absolute

1134 difference between the models is smaller than the 1σ on each model. All differences are $<2\sigma$.
 1135 Comparison of PDF fits for the Fogo inclusions like in a-b) are shown in Supporting Fig. S13.

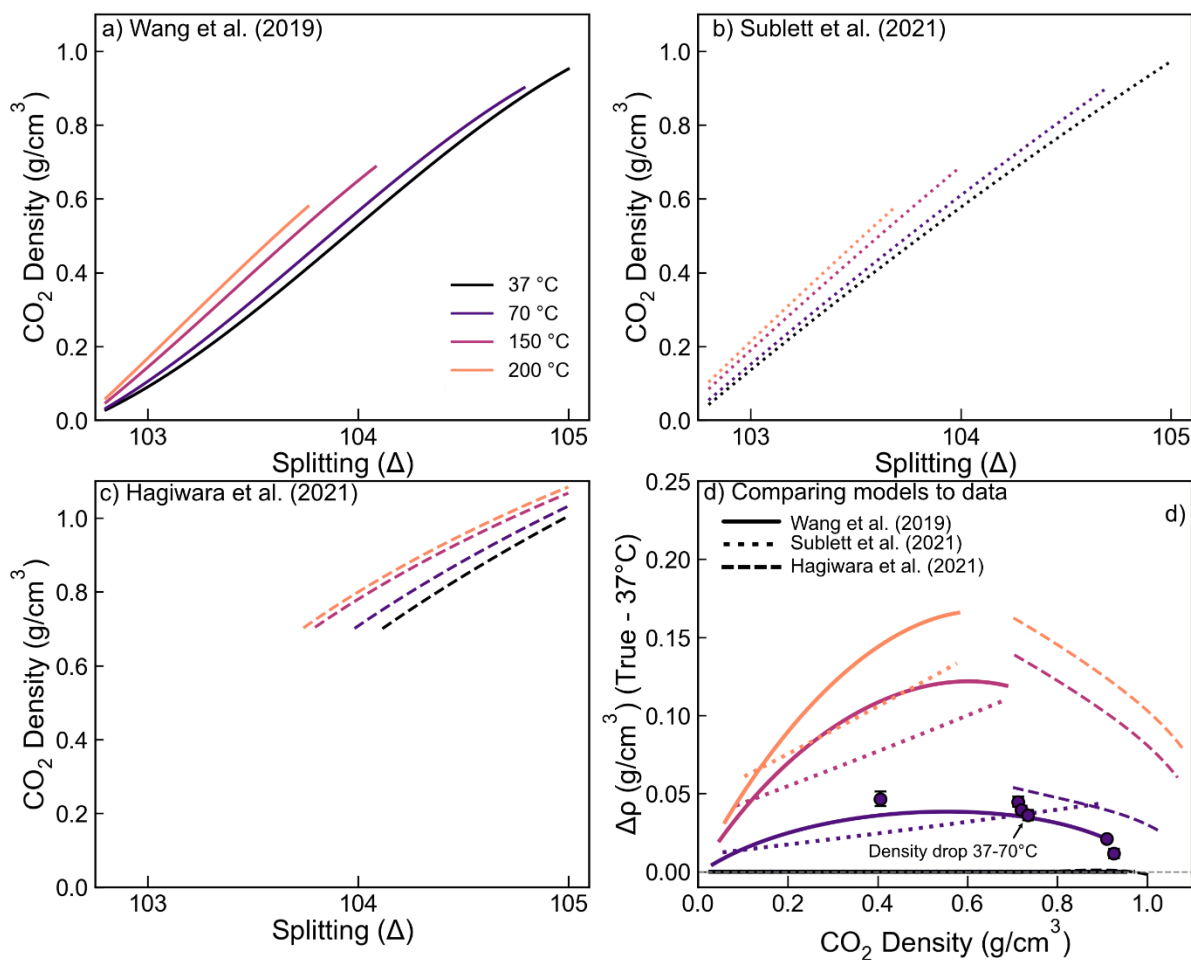


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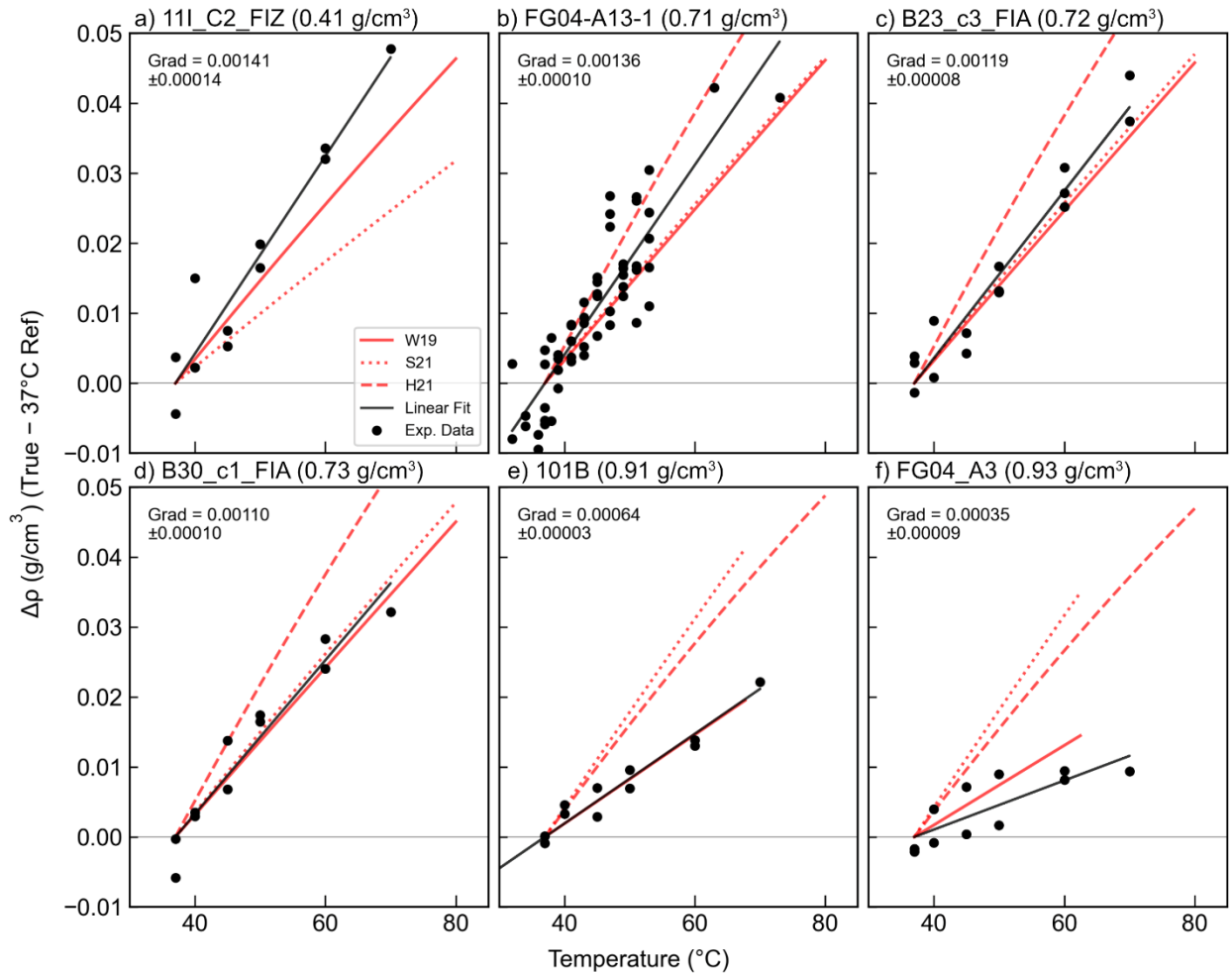
1137 Figure 9 –Assessment of Raman errors using natural FIs with independent density constraints
 1138 from microthermometry a) Errors on each peak position obtained from Imfit for low and high
 1139 density FIs from DeVitre and Wieser (2024) from Fogo and Wieser et al. (2025) from Mauna Loa.
 1140 b) Comparison of densities obtained from Raman spectroscopy and Microthermometry. The error
 1141 bars show 1 sigma errors estimated from peak fitting, analytical drift and densimeter uncertainty
 1142 for Raman spectroscopy and Microthermometry combined in quadrature. c) Same comparison
 1143 but using standard errors calculated from repeated Raman measurements (only FI with repeated
 1144 analyses are shown). On b-c), we overlay the $\pm 1 \sigma$ error bounds that would result from the splitting
 1145 errors suggested by Bakker (2021). We also overlay red lines showing the ± 0.01 - 0.02 g/cm^3 errors
 1146 that numerous studies comparing the two methods have stated.



1148 Figure 10 – Investigating peak fitting error as a function of integration time and accumulations
 1149 using a low density capillary tube ($\sim 0.01 \text{ g/cm}^3$, Le et al., 2021). a-b) Error on the peak position
 1150 of diad1 and diad2 for 45 and 90s acquisitions. The x axis shows the number of spectra that were
 1151 averaged prior to fitting the peaks (analogous to accumulation). The dots show an individual fitted
 1152 spectra (e.g. spectra 1 to 14 for accumulations =1, average of spectra 1-2, 2-3 3-4 for n=2
 1153 accumulations etc.), and the cross and error bar show the mean and standard deviation of
 1154 averaged spectra for each number of accumulations (e.g. 14 individually fitted spectra for n=1
 1155 accumulations). c-d) Example spectra for 45s x 1 and 90s x 1. The signal to noise (S:N) ratio is
 1156 calculated as the height of the highest pixel divided by the standard deviation of the region
 1157 between the two diads. e) Error on peak fitting as a function of total acquisition time. For example,
 1158 five 120s acquisitions were averaged for the 10 minute total acquisition time. The same figure for
 1159 a low density melt inclusion vapour bubble is shown in Supporting Fig. 4.



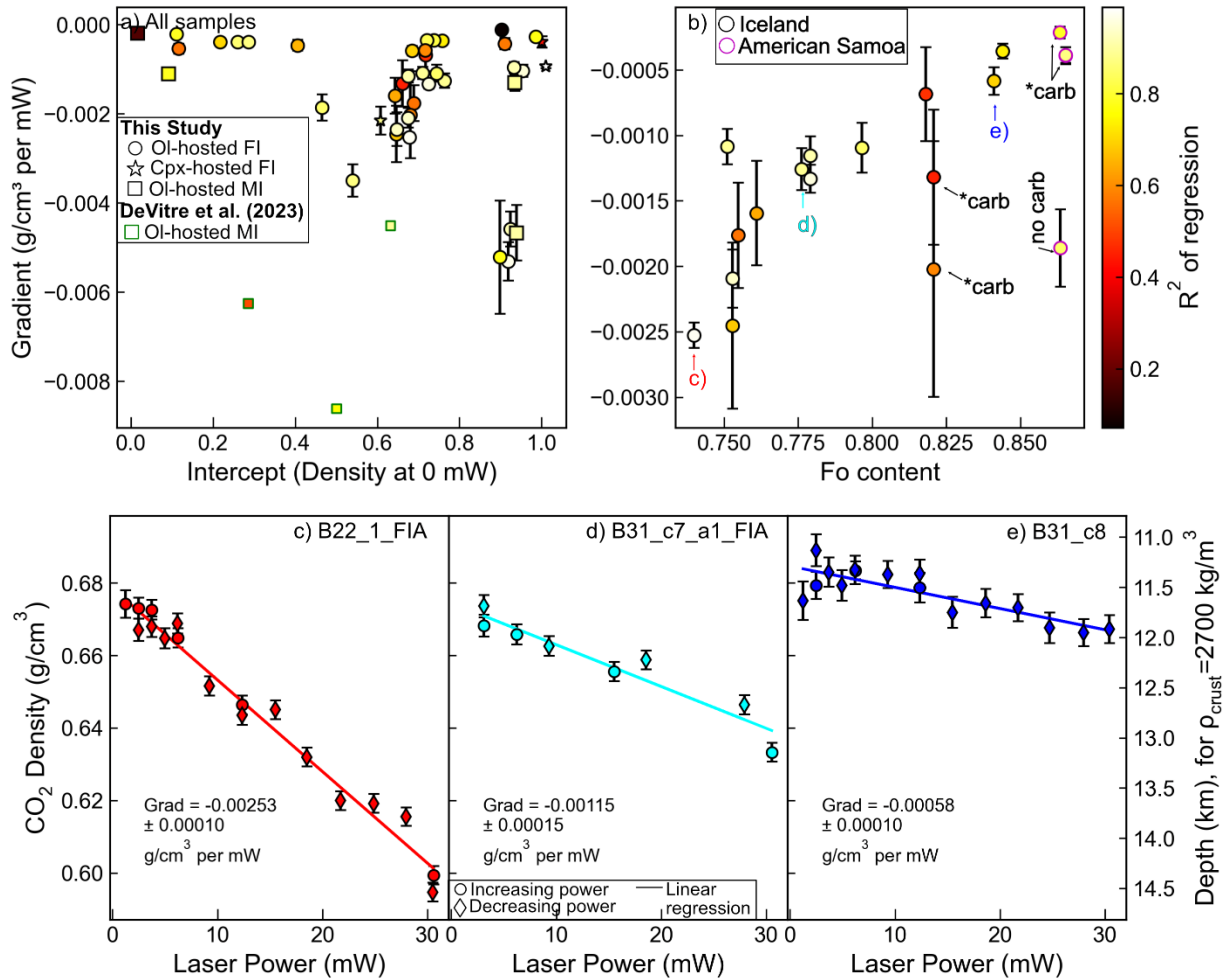
1161 Figure 11 – Models describing the temperature sensitivity of the densimeter. a-c) Relationship
 1162 between splitting and CO₂ for 4 temperatures calculated using equation 3 of Wang et al., (2019,
 1163 W19, panel a) equation 8 of Sublett et al. (2021, S2021, panel b) and equation 2 of Hagiwara et
 1164 al. (2021a, H2021, panel c). Lines are terminated outside the calibration range of each study. d)
 1165 Difference between true density of a FI and the density calculated using the 37 °C when analysed
 1166 at T=X °C. We overlay the data from fluid inclusions we analyze at a variety of temperatures (See
 1167 Figure 12).



1168

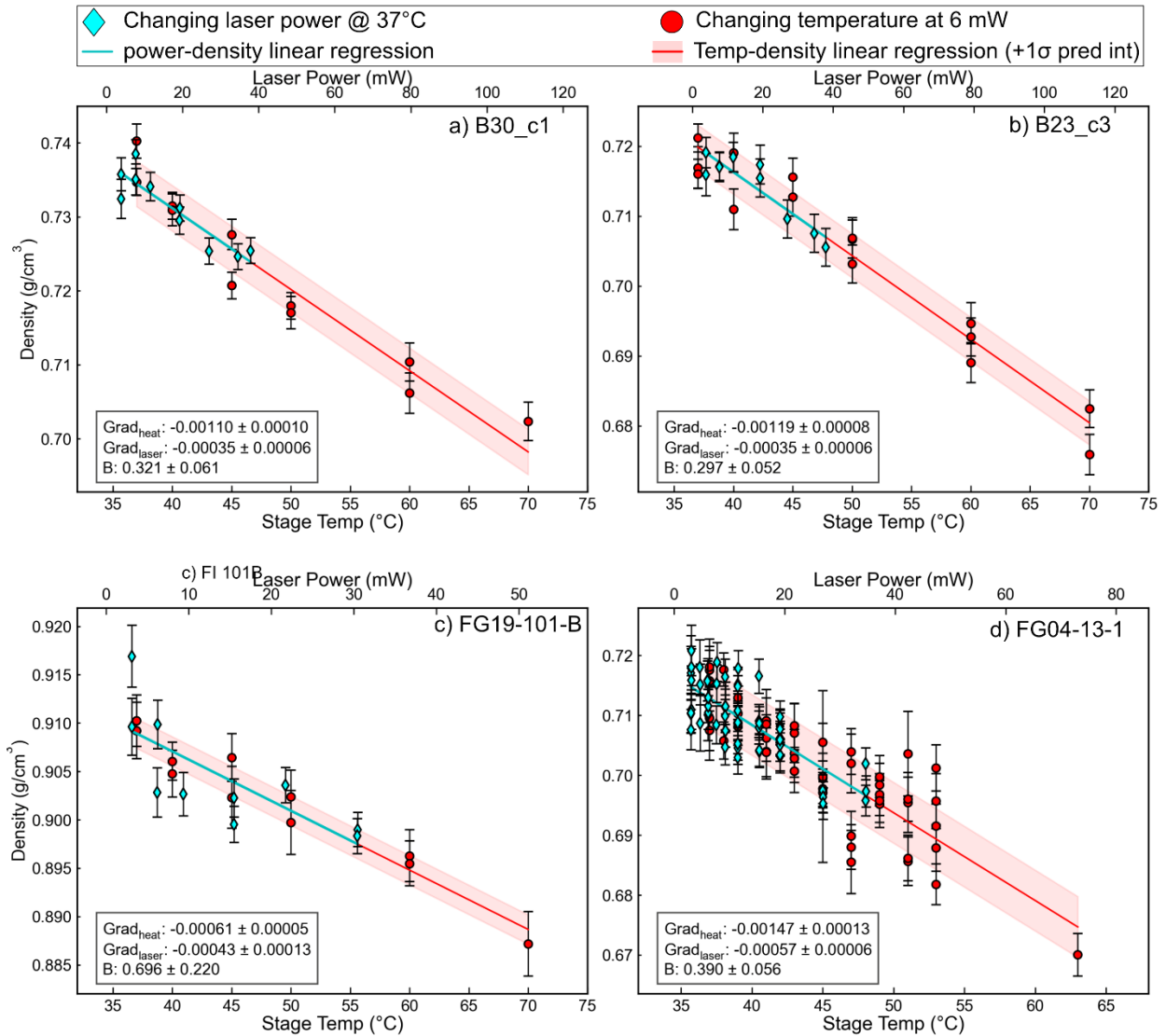
1169 Figure 12 – Changes in apparent density calculated using the 37 °C densimeter for fluid inclusions
 1170 heated in a linkam stage. The ‘True’ density of the inclusion was determined from a regression
 1171 between density and temperature, taking the value at 37 °C (the temperature of our gas cell
 1172 calibration). This was then compared to the calculated density at T=X°C using the 37 °C
 1173 densimeter (i.e. panel a, the density at 70 °C was 0.046 g/cm³ lower than the intercept at 37 °C).

1174 The solid, dotted and dashed lines show the predicted change in density using the models of
 1175 W19, S21 and H21 respectively. The W19 model provides an excellent fit to all fluid inclusions.



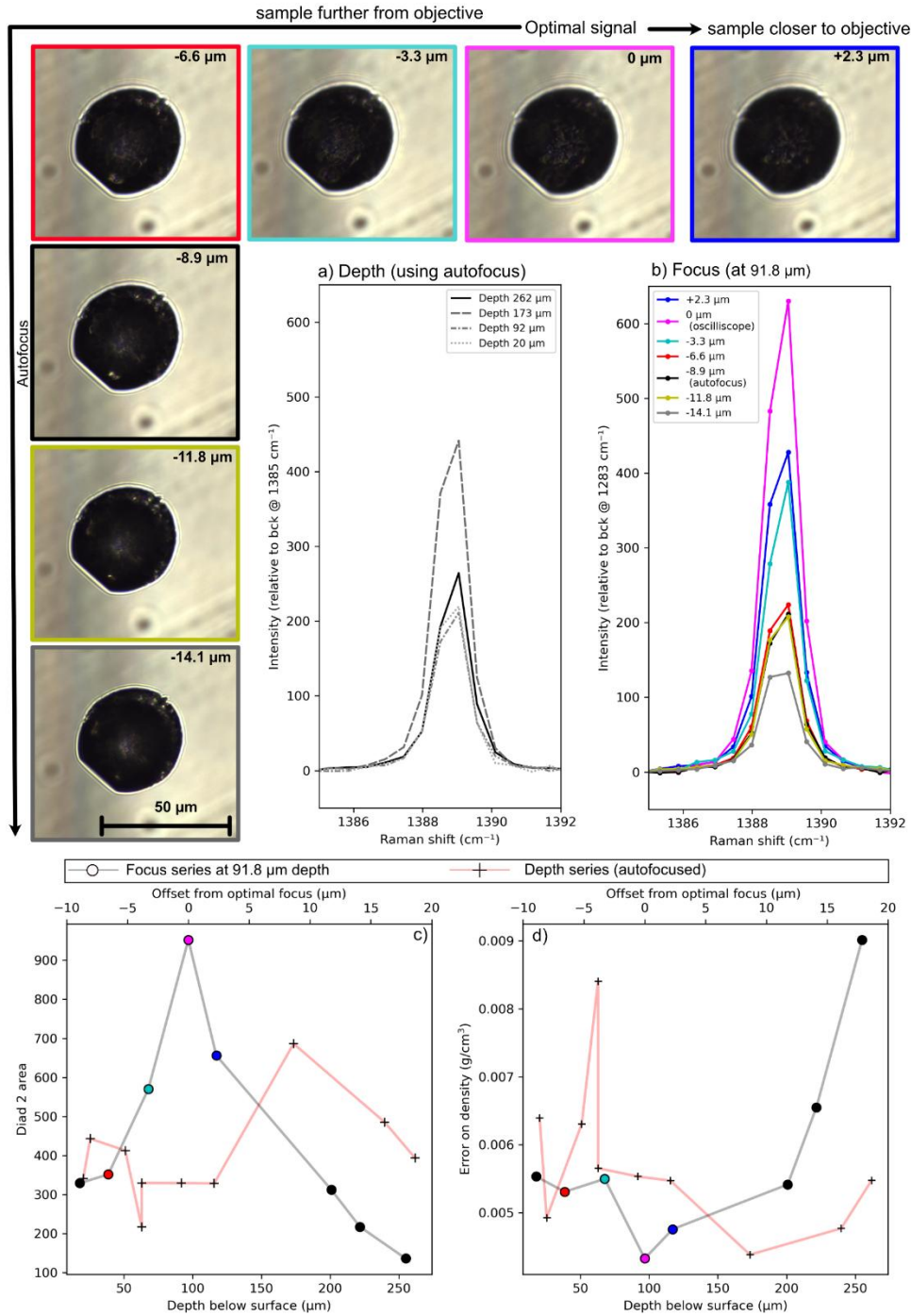
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1177 Figure 13 – Relationship between density and laser power at the sample. a) Gradient of laser
 1178 power vs. density for all the heating tests conducted here (individual regressions shown in
 1179 Supporting Figures). Symbol shapes show the host mineral and whether the sample is a fluid
 1180 inclusion or melt inclusion. Squares with green edges show the three melt inclusions from DeVitre
 1181 et al. (2023) b) Gradient of each power series vs. the forsterite content of the host crystal. FI with
 1182 a carbonate are indicated with arrows. c-e) Laser power series on three Icelandic FI indicated by
 1183 arrows in b). The circle symbols show data collected at increasing laser power, and the diamonds
 1184 at decreasing laser power. The gradients for the best-fit linear regression are shown.



1185

1186 Figure 14 – Comparison of heating and laser power series. For each FI, we perform acquisitions
 1187 at temperatures from 37-70 °C (red circles) at 6 mW, and at laser powers from 6 mW to ~50 mW
 1188 with the sample temperature controlled at 37 °C (cyan diamonds). We adjust the limits of the two
 1189 x axes such that the gradients of these two data series are identical. The gradient of each series
 1190 is shown, along with the B value.



1191

1192 Figure 15 – Investigation of how signal strengths are affected by focus and depth in the sample.
 1193 a) Spectra near Diad 2, showing the signal strength for different depths of the FI in the sample
 1194 during serial sectioning. After each polishing round, the autofocus function was used to find the
 1195 optimum depth. b) When the FI was at 91.8 μm depth in the sample, we acquire spectra at 7
 1196 different focus positions. The black spectra and dot represents the focus position found by the

1197 autofocus, and the pink spectra and dot the best focus position using the oscilloscope feature.
1198 The images on the left-hand side show what the FI looked like at each focus. c) Change in diad
1199 2 intensity and d) density error as a function of depth in the sample (circles) and focus (crosses).

1200

1201 [References](#)

- 1202 Allison, C.M., Roggensack, K., Clarke, A.B., 2021. Highly explosive basaltic eruptions driven by CO₂
1203 exsolution. *Nat. Commun.* 12, 217. <https://doi.org/10.1038/s41467-020-20354-2>
- 1204 Bakker, R.J., 2021. The perfection of Raman spectroscopic gas densimeters. *J. Raman Spectrosc.* 52,
1205 1923–1948. <https://doi.org/10.1002/jrs.6245>
- 1206 Bearden, A.T., Wieser, P.E., Rangel, B., DeVitre, C.L., Gleeson, M.L.M., Monteleone, B.D., Lynn, K.J., 2025.
1207 Volatile systematics and magma storage depths at Mauna Loa: a melt inclusion perspective., in:
1208 In Chapman Conference on Caldera-Forming Eruptions at Basaltic Volcanoes: Insights and
1209 Puzzles from Kīlauea and Beyond. AGU.
- 1210 Burke, E.A.J., 2001. Raman microspectrometry of fluid inclusions. *Lithos* 55, 139–158.
1211 [https://doi.org/10.1016/S0024-4937\(00\)00043-8](https://doi.org/10.1016/S0024-4937(00)00043-8)
- 1212 Chassé, M., Lelong, G., Van Nijnatten, P., Schoofs, I., De Wolf, J., Galois, L., Calas, G., 2015. Optical
1213 Absorption Microspectroscopy (μ -OAS) Based on Schwarzschild-Type Cassegrain Optics. *Appl.*
1214 *Spectrosc.* 69, 457–463. <https://doi.org/10.1366/14-07628>
- 1215 Dayton, K., Gazel, E., Wieser, P., Troll, V.R., Carracedo, J.C., La Madrid, H., Roman, D.C., Ward, J., Aulinas,
1216 M., Geiger, H., Deegan, F.M., Gisbert, G., Perez-Torrado, F.J., 2023. Deep magma storage during
1217 the 2021 La Palma eruption. *Sci. Adv.* 9, eade7641. <https://doi.org/10.1126/sciadv.ade7641>
- 1218 DeVitre, C.L., Allison, C.M., Gazel, E., 2021. A high-precision CO₂ densimeter for Raman spectroscopy
1219 using a Fluid Density Calibration Apparatus. *Chem. Geol.* 584, 120522.
1220 <https://doi.org/10.1016/j.chemgeo.2021.120522>
- 1221 DeVitre, C.L., Dayton, K., Gazel, E., Pamukçu, A., Gaetani, G., Wieser, P.E., 2023a. Laser heating effect on
1222 Raman analysis of CO₂ co-existing as liquid and vapor in olivine-hosted melt inclusion bubbles.
1223 *Volcanica* 6, 201–219. <https://doi.org/10.30909/vol.06.02.201219>
- 1224 DeVitre, C.L., Gazel, E., Ramalho, R.S., Venugopal, S., Steele-MacInnis, M., Hua, J., Allison, C.M., Moore,
1225 L.R., Carracedo, J.C., Monteleone, B., 2023b. Oceanic intraplate explosive eruptions fed directly
1226 from the mantle. *Proc. Natl. Acad. Sci.* 120, e2302093120.
1227 <https://doi.org/10.1073/pnas.2302093120>
- 1228 DeVitre, C.L., Wieser, P.E., 2024. Reliability of Raman analyses of CO₂-rich fluid inclusions as a
1229 geobarometer at Kīlauea. *Geochem. Perspect. Lett.* 29, 1–8.
1230 <https://doi.org/10.7185/geochemlet.2404>
- 1231 Fall, A., Tattitch, B., Bodnar, R.J., 2011. Combined microthermometric and Raman spectroscopic
1232 technique to determine the salinity of H₂O–CO₂–NaCl fluid inclusions based on clathrate
1233 melting. *Geochim. Cosmochim. Acta* 75, 951–964. <https://doi.org/10.1016/j.gca.2010.11.021>
- 1234 Fermi, E., 1931. Über den ramaneffekt des kohlendioxyds. *Z. Für Phys.* 71(3–4).
- 1235 Frezzotti, M.L., Tecce, F., Casagli, A., 2012. Raman spectroscopy for fluid inclusion analysis. *J. Geochem.*
1236 *Explor.* 112, 1–20. <https://doi.org/10.1016/j.gexplo.2011.09.009>
- 1237 Fukura, S., Mizukami, T., Odake, S., Kagi, H., 2006. Factors Determining the Stability, Resolution, and
1238 Precision of a Conventional Raman Spectrometer. *Appl. Spectrosc.* 60, 946–950.
1239 <https://doi.org/10.1366/000370206778062165>

1240 Gleeson, M., Wieser, P.E., DeVitre, C.L., Shi, S.C., Millet, M.-A., Muir, D.D., Stock, M.J., Lissenberg, J., 2025.
1241 Persistent High-Pressure Magma Storage beneath a Near-Ridge Ocean Island Volcano (Isla
1242 Floreana, Galápagos). *J. Petrol.* 66, egaf031. [https://doi.org/10.1093/](https://doi.org/10.1093/petrology/egaf031)
1243 Guo, X., Mughal, A., Dunphy, D., Stone, G., Miller, D., Huh, S., 2020. Ambient Temperature–Corrected
1244 Mechanical Stress Mapping of Gallium Nitride and Aluminum Indium Gallium Phosphide Films by
1245 Raman Scattering Spectroscopy for Characterization of Light-Emitting Diodes. *Phys. Status Solidi*
1246 *A* 217, 1900776. <https://doi.org/10.1002/pssa.201900776>
1247 Hagiwara, Y., Kawano, T., Takahata, K., Torimoto, J., Yamamoto, J., 2021a. Temperature dependence of a
1248 Raman CO₂ densimeter from 23°C to 200°C and 7.2 to 248.7 MPa: Evaluation of density
1249 underestimation by laser heating. *J. Raman Spectrosc.* 52, 1744–1757.
1250 <https://doi.org/10.1002/jrs.6188>
1251 Hagiwara, Y., Sogo, Y., Takahata, K., Yamamoto, J., 2018. Temperature dependence of CO₂ densimetry
1252 using micro-Raman Spectrometry at laboratory conditions. *Geochem. J.* 52, 379–383.
1253 <https://doi.org/10.2343/geochemj.2.0523>
1254 Hagiwara, Y., Torimoto, J., Yamamoto, J., 2020. Pressure measurement and detection of small H₂O
1255 amounts in high-pressure H₂O–CO₂ fluid up to 141 MPa using Fermi diad splits and bandwidths
1256 of CO₂. *J. Raman Spectrosc.* 51, 1003–1018. <https://doi.org/10.1002/jrs.5865>
1257 Hagiwara, Y., Yokokura, L., Yamamoto, J., 2023. Unlocking ultimate precision of intensity and area ratio
1258 measurements in Raman spectroscopy: Insights from simulation, experimentation, and theory
1259 and implications for isotope ratio analysis. *J. Raman Spectrosc.* 54, 1440–1464.
1260 <https://doi.org/10.1002/jrs.6594>
1261 Hagiwara, Y., Yoshida, K., Yoneda, A., Torimoto, J., Yamamoto, J., 2021b. Experimental variable effects on
1262 laser heating of inclusions during Raman spectroscopic analysis. *Chem. Geol.* 559, 119928.
1263 <https://doi.org/10.1016/j.chemgeo.2020.119928>
1264 Hazen, R., Mao, H., Bell, P.M., 1977. Effects of compositional variation on absorption spectra of lunar
1265 olivine. *Proc Lunar Sci Conf 8th* 1081–1090.
1266 Howe, T.A., Christopher, T.E., Moune, S., Tuffen, H., Schiavi, F., 2025. Melt inclusion bubbles provide new
1267 insights into crystallisation depths and CO₂ systematics at Soufrière Hills Volcano, Montserrat.
1268 *Front. Earth Sci.* 12, 1509409. <https://doi.org/10.3389/feart.2024.1509409>
1269 Inoue, Y., Okiyama, R., Hagiwara, Y., Yamamoto, J., 2023. Raman spectroscopic evaluation of precision of
1270 oxygen isotope ratio of carbon dioxide. *Geochem. J.* 57, 92–99.
1271 <https://doi.org/10.2343/geochemj.GJ23007>
1272 Jakubek, R.S., Fries, M.D., 2022. Calibration of the temporal drift in absolute and relative Raman
1273 intensities in large Raman images using a mercury–argon lamp. *J. Raman Spectrosc.* 53, 137–147.
1274 <https://doi.org/10.1002/jrs.6259>
1275 Kahl, M., Morgan, D.J., Wieser, P.E., Bali, E., Guðfinnsson, G.H., Neave, D.A., Walshaw, R., 2025. Crystal-
1276 mush remobilization timescales and magma storage depth in the Snæfellsnes Volcanic Zone (W-
1277 Iceland): insights from olivine Fe-Mg diffusion chronometry and fluid inclusion barometry. *Bull.*
1278 *Volcanol.* 87, 118. <https://doi.org/10.1007/s00445-025-01892-3>
1279 Kawakami, Y., Yamamoto, J., Kagi, H., 2003. Micro-Raman Densimeter for CO₂ Inclusions in Mantle-
1280 Derived Minerals. *Appl. Spectrosc.* 57, 1333–1339.
1281 <https://doi.org/10.1366/000370203322554473>
1282 Kobayashi, T., Yamamoto, J., Hirajima, T., Ishibashi, H., Hirano, N., Lai, Y., Prikhod'ko, V.S., Arai, S., 2012.
1283 Conformity and precision of CO₂ densimetry in CO₂ inclusions: microthermometry *versus*
1284 Raman microspectroscopic densimetry: Conformity and precision of CO₂ densimetry in CO₂
1285 inclusions. *J. Raman Spectrosc.* 43, 1126–1133. <https://doi.org/10.1002/jrs.3134>
1286 Kramida, A., Ralchenko, Y., 1999. NIST Atomic Spectra Database, NIST Standard Reference Database 78.
1287 <https://doi.org/10.18434/T4W30F>

1288 Lamadrid, H.M., Moore, L.R., Moncada, D., Rimstidt, J.D., Burruss, R.C., Bodnar, R.J., 2017. Reassessment
1289 of the Raman CO₂ densimeter. *Chem. Geol.* 450, 210–222.
1290 <https://doi.org/10.1016/j.chemgeo.2016.12.034>
1291 Le, V.-H., Caumon, M.-C., Tarantola, A., 2021. FRAnCIs calculation program with universal Raman
1292 calibration data for the determination of PVX properties of CO₂–CH₄–N₂ and CH₄–H₂O–NaCl
1293 systems and their uncertainties. *Comput. Geosci.* 156, 104896.
1294 <https://doi.org/10.1016/j.cageo.2021.104896>
1295 Lin, F., Bodnar, R.J., Becker, S.P., 2007. Experimental determination of the Raman CH₄ symmetric
1296 stretching (ν₁) band position from 1–650bar and 0.3–22°C: Application to fluid inclusion studies.
1297 *Geochim. Cosmochim. Acta* 71, 3746–3756. <https://doi.org/10.1016/j.gca.2007.05.016>
1298 Liu, D., Hennelly, B., O’Neill, L., Byrne, H.J., 2018. Investigation of wavenumber calibration for Raman
1299 spectroscopy using a polymer standard, in: Berghmans, F., Mignani, A.G. (Eds.), *Optical Sensing
1300 and Detection V. Presented at the Optical Sensing and Detection, SPIE, Strasbourg, France*, p. 79.
1301 <https://doi.org/10.1117/12.2307574>
1302 Liu, D., Hennelly, B.M., 2024. Wavenumber Calibration Protocol for Raman Spectrometers Using Physical
1303 Modelling and a Fast Search Algorithm. *Appl. Spectrosc.* 78, 790–805.
1304 <https://doi.org/10.1177/00037028241254847>
1305 McLennan F.R.S., J.C., Smith M.A., H.D., 1932. Raman effects with liquid and solid carbon dioxide. *Can. J.*
1306 *Res.* 7, 551–555. <https://doi.org/10.1139/cjr32-104>
1307 Mestari, A., Gaufres, R., Huguet, P., 1997. Behaviour of the calibration of a Raman spectrometer with
1308 temperature changes. *J. Raman Spectrosc.* 28, 785–789. [https://doi.org/10.1002/\(SICI\)1097-
1309 4555\(199710\)28:10%253C785::AID-JRS148%253E3.0.CO;2-D](https://doi.org/10.1002/(SICI)1097-4555(199710)28:10%253C785::AID-JRS148%253E3.0.CO;2-D)
1310 Moore, L.R., Gazel, E., Tuohy, R., Lloyd, A.S., Esposito, R., Steele-MacInnis, M., Hauri, E.H., Wallace, P.J.,
1311 Plank, T., Bodnar, R.J., 2015. Bubbles matter: An assessment of the contribution of vapor bubbles
1312 to melt inclusion volatile budgets. *Am. Mineral.* 100, 806–823. [https://doi.org/10.2138/am-
1313 2015-5036](https://doi.org/10.2138/am-2015-5036)
1314 Newville, M., Stensitzki, T., Allen, D.B., Rawlik, M., Ingargiola, A., Nelson, A., 2016. LMFIT: Non-linear
1315 least-square minimization and curve-fitting for Python. *Astrophys. Source Code Libr.* ascl-1606.
1316 Remigi, S., Mancini, T., Ferrando, S., Frezzotti, M.L., 2021. Interlaboratory Application of Raman CO₂
1317 Densimeter Equations: Experimental Procedure and Statistical Analysis Using Bootstrapped
1318 Confidence Intervals. *Appl. Spectrosc.* 75, 867–881. <https://doi.org/10.1177/0003702820987601>
1319 Rosso, K.M., Bodnar, R.J., 1995. Microthermometric and Raman spectroscopic detection limits of CO₂ in
1320 fluid inclusions and the Raman spectroscopic characterization of CO₂. *Geochim. Cosmochim.*
1321 *Acta* 59, 3961–3975. [https://doi.org/10.1016/0016-7037\(95\)94441-H](https://doi.org/10.1016/0016-7037(95)94441-H)
1322 Song, Y., Chou, I., Hu, W., Robert, B., Lu, W., 2009. CO₂ Density-Raman Shift Relation Derived from
1323 Synthetic Inclusions in Fused Silica Capillaries and Its Application. *Acta Geol. Sin. - Engl. Ed.* 83,
1324 932–938. <https://doi.org/10.1111/j.1755-6724.2009.00090.x>
1325 Span, R., Wagner, W., 1996. A New Equation of State for Carbon Dioxide Covering the Fluid Region from
1326 the Triple-Point Temperature to 1100 K at Pressures up to 800 MPa. *J. Phys. Chem. Ref. Data* 25,
1327 1509–1596. <https://doi.org/10.1063/1.555991>
1328 Sublett, D.M., Sendula, E., Lamadrid, H.M., Steele-MacInnis, M., Spiekermann, G., Bodnar, R.J., 2021.
1329 Raman spectral behavior of N₂, CO₂, and CH₄ in N₂–CO₂–CH₄ gas mixtures from 22°C to
1330 200°C and 10 to 500 bars, with application to other gas mixtures. *J. Raman Spectrosc.* 52, 750–
1331 769. <https://doi.org/10.1002/jrs.6033>
1332 Taran, M.N., Matsyuk, S.S., 2013. Fe²⁺, Mg-distribution among non-equivalent structural sites M1 and
1333 M2 in natural olivines: an optical spectroscopy study. *Phys. Chem. Miner.* 40, 309–318.
1334 <https://doi.org/10.1007/s00269-013-0572-x>

1335 Ullrich, K., Langer, K., Becker, K.D., 2002. Temperature dependence of the polarized electronic absorption
1336 spectra of olivines. Part I - fayalite. *Phys. Chem. Miner.* 29, 409–419.
1337 <https://doi.org/10.1007/s00269-002-0248-4>

1338 Venugopal, S., Schiavi, F., Moune, S., Bolfan-Casanova, N., Druitt, T., Williams-Jones, G., 2020. Melt
1339 inclusion vapour bubbles: the hidden reservoir for major and volatile elements. *Sci. Rep.* 10,
1340 9034. <https://doi.org/10.1038/s41598-020-65226-3>

1341 Wang, H., Lu, W., Wang, W., Liu, Q., Yamamoto, J., 2024. High-precision analysis of carbon isotopic
1342 composition for individual CO₂ inclusions via Raman spectroscopy: Addressing issues arising
1343 from the laser-heating effects. *Chem. Geol.* 651, 122014.
1344 <https://doi.org/10.1016/j.chemgeo.2024.122014>

1345 Wang, W., Caumon, M.-C., Tarantola, A., Pironon, J., Lu, W., Huang, Y., 2019. Raman spectroscopic
1346 densimeter for pure CO₂ and CO₂-H₂O-NaCl fluid systems over a wide P-T range up to 360 °C
1347 and 50 MPa. *Chem. Geol.* 528, 119281. <https://doi.org/10.1016/j.chemgeo.2019.119281>

1348 Wang, X., Chou, I.-M., Hu, W., Burruss, R.C., Sun, Q., Song, Y., 2011. Raman spectroscopic measurements
1349 of CO₂ density: Experimental calibration with high-pressure optical cell (HPOC) and fused silica
1350 capillary capsule (FSCC) with application to fluid inclusion observations. *Geochim. Cosmochim.*
1351 *Acta* 75, 4080–4093. <https://doi.org/10.1016/j.gca.2011.04.028>

1352 Wieser, P., Kent, A.J.R., DeVitre, C.L., Gazel, E., Till, C.B., Johnson, Emily, Wallace, Paul, Couperthwaite,
1353 F.K., 2023. Bubble Trouble: Raman measurements of Cascade melt inclusion vapour bubbles
1354 indicate substantial underestimation of magma storage depths in legacy data. Presented at the
1355 AGU Fall Meeting 2023. Poster No. 0203, id. V23D-0203.

1356 Wieser, P.E., Edmonds, M., Maclennan, J., Jenner, F.E., Kunz, B.E., 2019. Crystal Scavenging from Mush
1357 Piles Recorded by Melt Inclusions. *Nat. Commun.*

1358 Wieser, P.E., Lamadrid, H., Maclennan, J., Edmonds, M., Matthews, S., Iacovino, K., Jenner, F.E., Gansecki,
1359 C., Trusdell, F., Lee, R.L., Ilyinskaya, E., 2021. Reconstructing Magma Storage Depths for the 2018
1360 Kīlauean Eruption From Melt Inclusion CO₂ Contents: The Importance of Vapor Bubbles.
1361 *Geochem. Geophys. Geosystems* 22. <https://doi.org/10.1029/2020GC009364>

1362 Wieser, P.E., Shi, S.C., Gleeson, M.L.M., Rangel, B., DeVitre, C.L., Bearden, A.T., Lynn, K.J., Camille-
1363 Caumon, M., 2025. Fluid inclusion constraints on the geometry of the magmatic plumbing
1364 system beneath Mauna Loa — Part 1: lavas and tephros. *Bull. Volcanol.* 87, 89.
1365 <https://doi.org/10.1007/s00445-025-01874-5>

1366 Wojdyr, M., 2010. *Fityk* : a general-purpose peak fitting program. *J. Appl. Crystallogr.* 43, 1126–1128.
1367 <https://doi.org/10.1107/S0021889810030499>

1368 Wright, R.B., Wang, C.H., 1973. Density effect on the Fermi resonance in gaseous CO₂ by Raman
1369 scattering. *J. Chem. Phys.* 58, 2893–2895. <https://doi.org/10.1063/1.1679594>

1370 Yamamoto, J., Hagiwara, Y., 2024. Precision evaluation of Raman densimetry for carbon dioxide:
1371 improvement by correction for the drift effect. *Appl. Opt.* 63, 1402.
1372 <https://doi.org/10.1364/AO.507939>

1373 Yamamoto, J., Kagi, H., 2008. Application of densimetry using micro-Raman spectroscopy for CO₂ fluid
1374 inclusions: a probe for elastic strengths of mantle minerals. *Eur. J. Mineral.* 20, 529–535.
1375 <https://doi.org/10.1127/0935-1221/2008/0020-1825>

1376 Yamamoto, J., Kagi, H., 2006. Extended Micro-Raman Densimeter for CO₂ Applicable to Mantle-
1377 originated Fluid Inclusions. *Chem. Lett.* 35, 610–611. <https://doi.org/10.1246/cl.2006.610>

1378 Yuan, X., Mayanovic, R.A., 2017. An Empirical Study on Raman Peak Fitting and Its Application to Raman
1379 Quantitative Research. *Appl. Spectrosc.* 71, 2325–2338.
1380 <https://doi.org/10.1177/0003702817721527>

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1 **Supporting Information for:** Evaluating and Optimizing Raman Analytical Protocols for CO₂
 2 Fluids

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 6 K1N 6N5, Canada

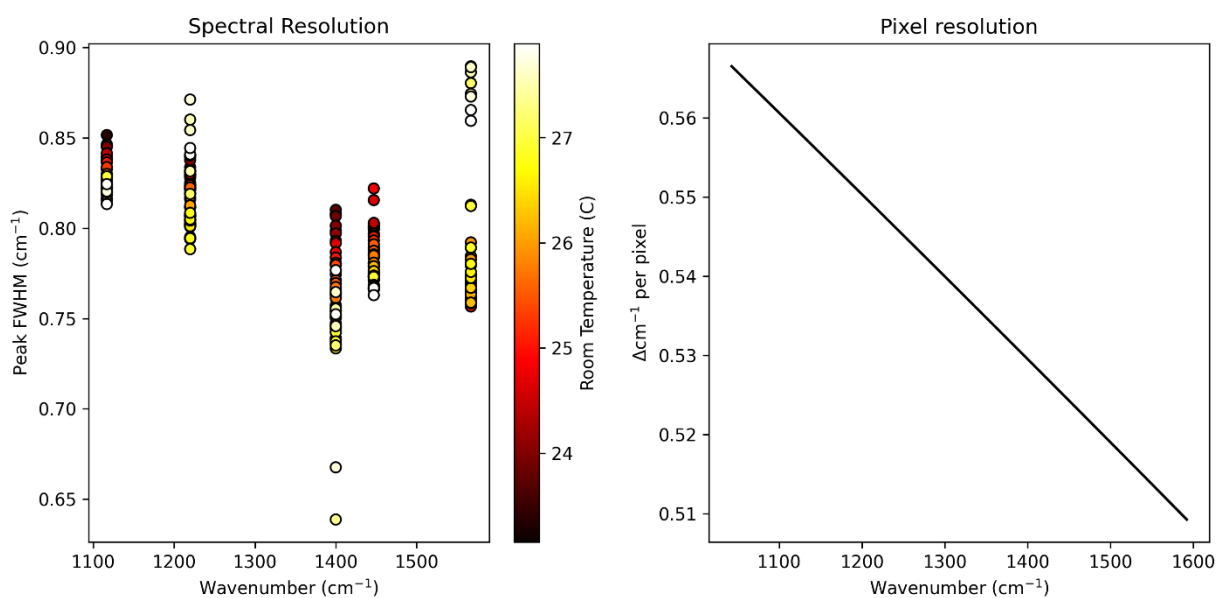
7 **Supporting Table 1-** Compilation of studies publishing densimeters, and the methods used to
 8 calibrate the spectrometer (which sets the non-linearity response seen in Fig. 1b) and correct data
 9 for drift during a run. n.d. indicates the publication did not say what method was used. It is not
 10 always clear if the Raman shift standard was used alone, or in addition to some form of instrument
 11 software calibration.

| Reference | Spectrometer Calibration method | Drift correction used | Pixel resolution | Spectral resolution |
|--|--|---|---|---------------------|
| Rosso and Bodnar, (1995) | Spectrometer position calibrated between analyses using Neon lines and a mercury line. | | 0.84 cm ⁻¹ /pixel | n.d. |
| Kawakami et al. (2003) | Raman shift calibrated with Naphthalene | n.d. | 1.5 cm ⁻¹ /pixel | n.d. |
| Yamamoto and Kagi, (2006) | Raman shift calibrated with Naphthalene | n.d. | 1.5 cm ⁻¹ /pixel | n.d. |
| Song et al. (2009) | Diamond (Raman shift standard) | n.d. | State spectral resolution of 1 cm ⁻¹ , (doesn't clarify if resolution per pixel or the fwhm of a narrow line) | |
| Lamadrid et al. (2017) LabRam HR (800 mm) | Establish linearity of the spectrometer using the Laser line and Ne line at 1458.58 cm ⁻¹ | Ne lamp (lines at 1031.42 cm ⁻¹ and 1458.58 cm ⁻¹ for a 514 nm laser) | No specific data, state spectral resolution of different instruments ranges from 1.4-8 cm ⁻¹ (doesn't clarify if | |

| | | | | |
|-------------------------|--|---|--|--|
| | | | resolution per pixel or the fwhm of a narrow line) | |
| Wang et al. (2011) | Co-collected diamond were used to calibrate the individual positions of each CO ₂ peak. Diad splits were corrected using the 1192.6 and 1598.9 cm ⁻¹ peaks of benzonitrile collected every 15 minutes. | | State spectral resolution of 1 cm ⁻¹ (doesn't clarify if resolution per pixel or the fwhm of a narrow line) | |
| Fall et al. (2011) | n.d. | n.d. | n.d. | |
| Wang et al. (2019) | Neon Lamp | Neon Lamp using the L17 correction for the 1447 and 1613 cm ⁻¹ lines (532.06 nm laser) | | |
| Le et al. (2019) | Say no need for absolute wavelength calibration of the spectrometer | n.d. | n.d. | 1.67 cm ⁻¹ using fwhm of Ne peak at 2348 cm ⁻¹ |
| Le et al., (2020) | n.d. | n.d. | n.d. | 1.6 cm ⁻¹ don't say method |
| Sublett et al. (2020) | No calibration method given ,but say they systematically subtract a X cm ⁻¹ offset for their peak positions of gases relative to other studies, | Ne peaks collected simultaneously to correct for spectrometer drift. | n.d. | n.d. |
| (Hagiwara et al., 2020) | Real-time wavelength calibration using 1449 and 1222 cm ⁻¹ Ne lines following Lamadrid and Lin et al. | | 0.29 cm ⁻¹ /pixel | 1.1 cm ⁻¹ |

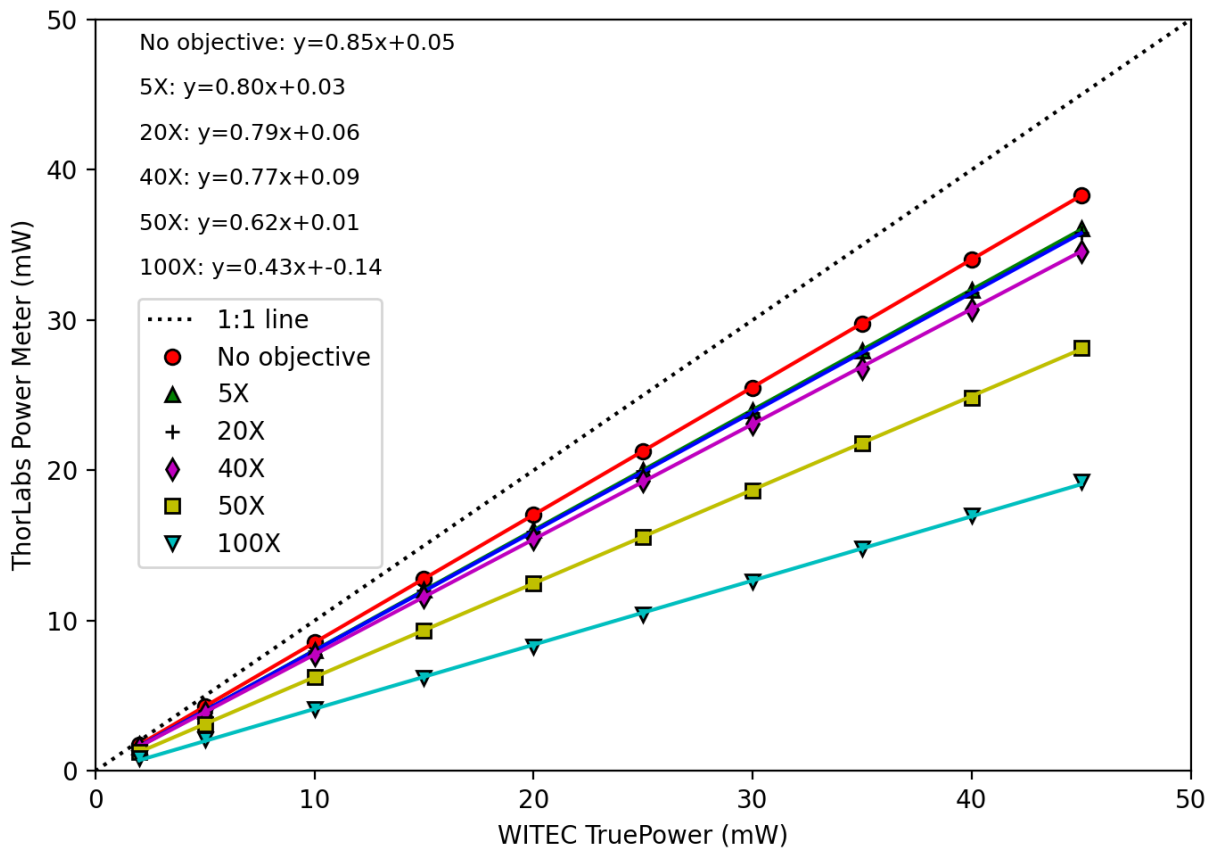
| | | | | |
|--------------------------|--|--|------------------------------|------|
| | n.d. | Ne lines at 1221.73 and 1449.19 cm ⁻¹ . | 0.28 cm ⁻¹ /pixel | n.d. |
| Wieser et al., (2021) | Si chip | Repeated synthetic FI measurement | n.d. | n.d. |
| DeVitre et al. (2021) | Hg-Ar lamp (atomic emission line) | Neon lamp using the L17 correction for the 1116 cm ⁻¹ and 1447 cm ⁻¹ lines (532.067 nm laser) | n.d. | n.d. |
| Allison et al. (2021) | Raman shift calibrated using three calibration standards (cyclohexane, naphthalene, 1,4 bis (2-methylstyryl) benzene. Unclear if analysed repeatedly to correct for drift. | | n.d. | n.d. |
| DeVitre and Wieser, 2024 | Hg-Ar lamp (atomic emission line) | Neon lamp using the L17 correction for the lines at 1117 cm ⁻¹ and 1447 cm ⁻¹ (532.046 nm laser) | 0.54 cm ⁻¹ /pixel | n.d. |

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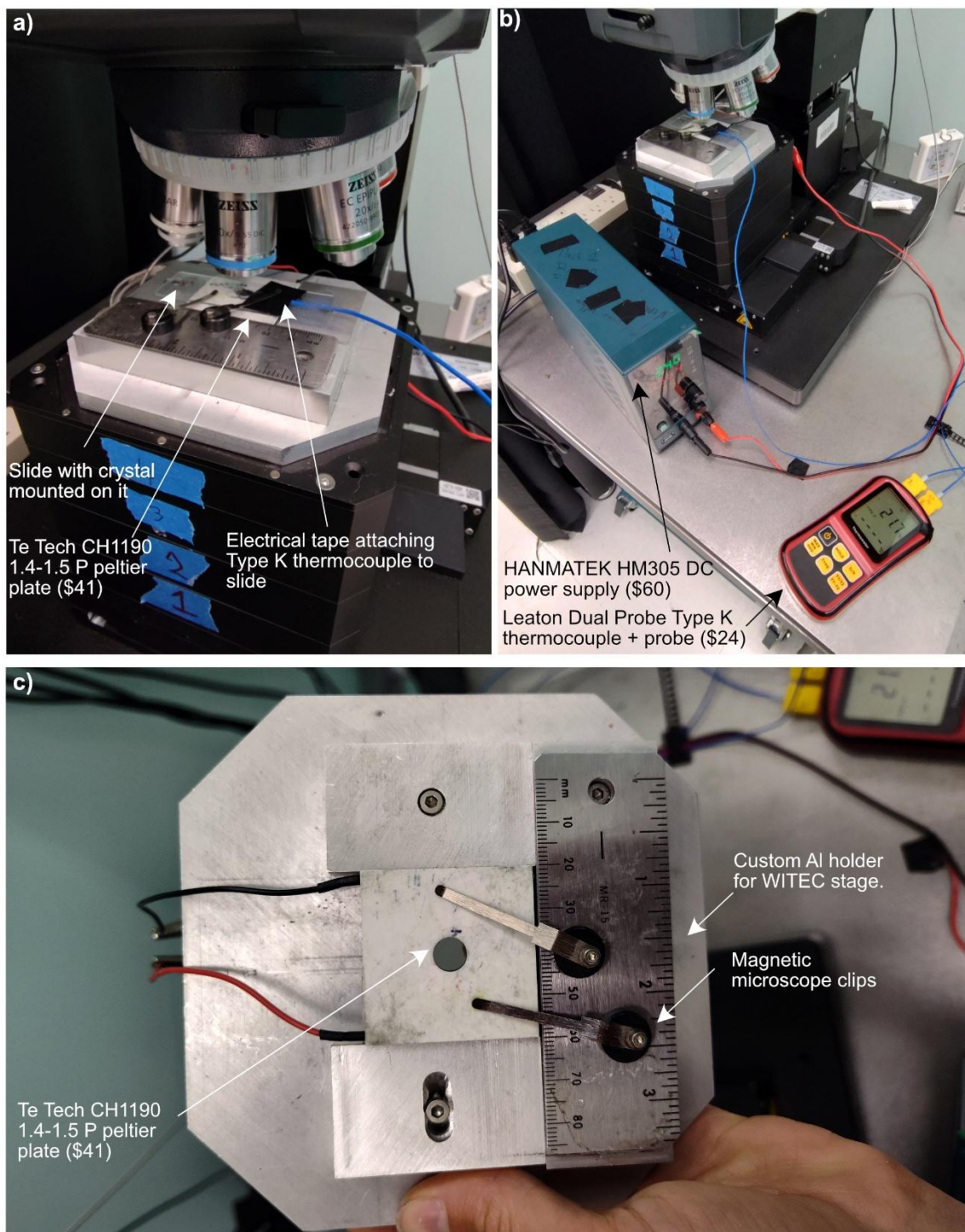
14 **Supporting Figure S1** – We estimate the spectral resolution using the full width half
 15 maximum(FWHM) of Ne lines following Hagiwara et al. (2023). Spectra were collected during the
 16 increase in room temperature, and show a decrease in peak FWHM as room T increases. This
 17 assumes that the Ne lines are infinitely narrow, and the measured FWHM is a function of the
 18 spectral resolution of the spectrometer. There are complex changes in FWHM with maximum –
 19 with not all lines showing the same pattern. We compare this to the pixel resolution, calculated by
 20 subtracting the x coordinate of the n+1 pixel from the nth pixel.



21

22 **Supporting Figure S2** – Comparison of power from the WITec TruePower read out with the power
 23 measured by a ThorLabs power meter. For the ‘No objective’ measurement, the S121C sensor
 24 was pushed into the objective turret hole after unscrewing the objective (it fits perfectly). For the
 25 5-100X measurements, the S121C sensor was set on the sample stage and the z height was
 26 varied until the maximum power reading was achieved on the sensor (this was very close to the
 27 optimal optical focus using reflected light). We found that x-y positioning of the laser on the sensor
 28 affected the laser power by roughly 2-3%.

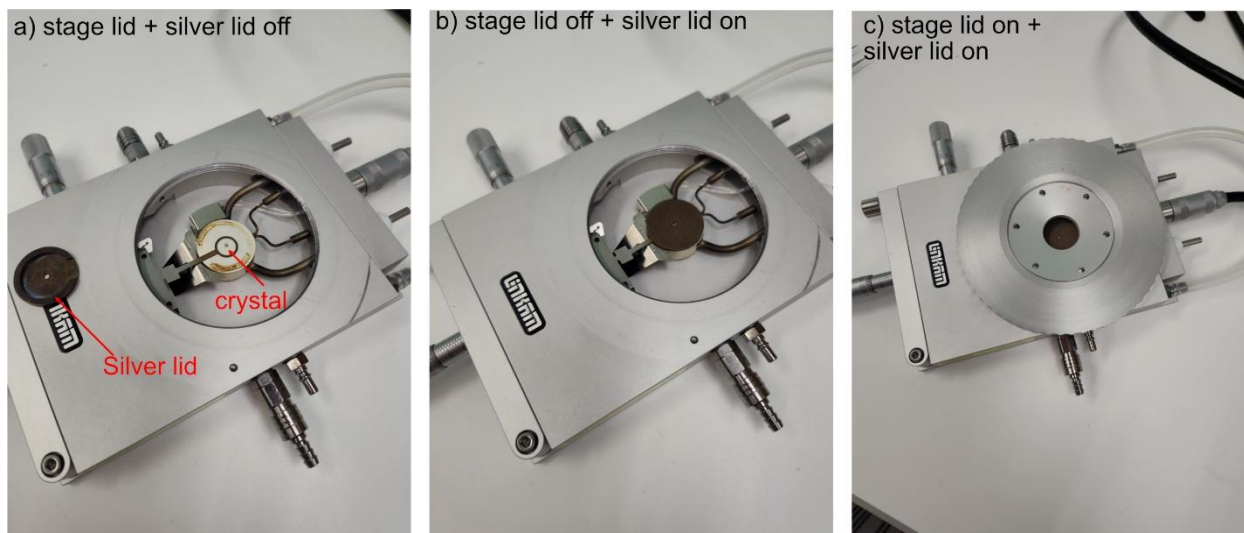
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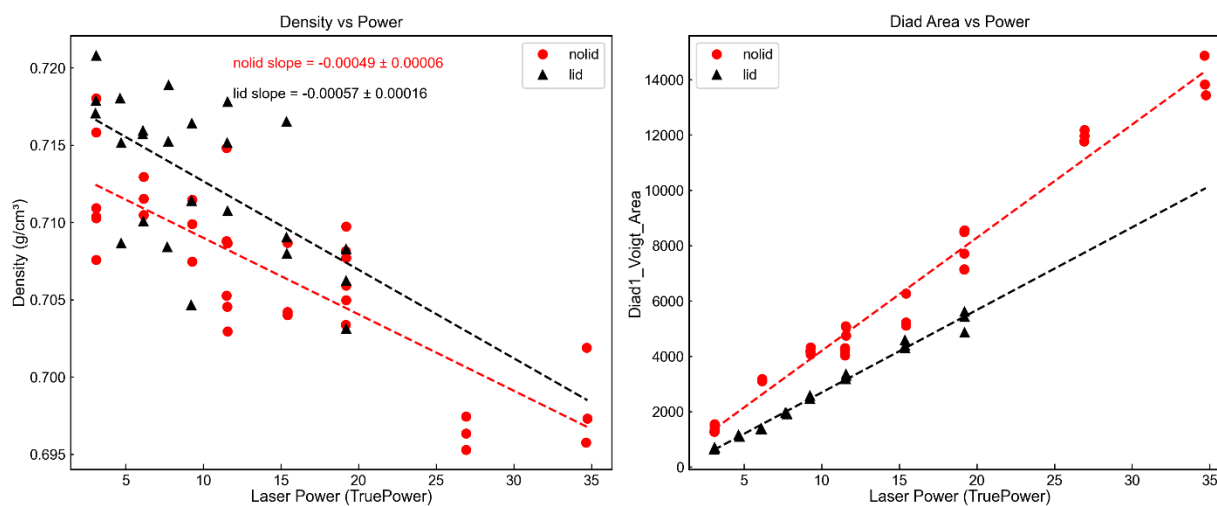
31 **Supporting Figure S3** – Apparatus used for heating samples to 37°C. a) Picture showing a glass
 32 slide with a FI mounted in crystalbond under a 50X Zeiss objective. b) Picture showing the entire
 33 WITec system and the sample holder. A Type K thermocouple is taped on using electrical tape.
 34 This picture was taken while the slide was still warming up (hence the 21.7 °C temperature). c)
 35 Close up of the custom Aluminum holder that has the bottom machined out to slot onto the shape
 36 of the WITec base. There are 4 screws that attach to the magnets on the WITec stage (because
 37 Al is not magnetic). The ruler is glued to the top to provide a magnetic surface to attach a set of

38 magnetic microscope clips to help keep the sample in place. This holder was manufactured by
 39 our in-house machine shop (labour and material cost, \$900). This holder keeps the peltier plate
 40 and sample still, making it possible to analyze very small inclusions without drift (1-2 μm).



41
 42 **Supporting Figure S4** – Linkam apparatus which is mounted on the Raman for heating tests. a)
 43 The stage is entirely open, and the crystal of interest is placed on the glass circle. b) The silver
 44 lid is then placed on top to help with temperature stability, but without sacrificing the Raman signal
 45 (as it has a hole in the top). In c), in addition to the silver lid, the stage lid is screwed on, leading
 46 to a drop in Raman signal.

47



48
 49 **Supporting Figure S5** – Comparison of heating series for FI FG04_a13_12 performed with and
 50 without the stage lid. The gradient is within uncertainty when density is plotted against laser power.
 51 However, the intensity of Diad 1 (peak area) is significantly lower with the stage lid in place.

52 **Peak fitting errors**

53 DiadFit versions $\leq 1.0.18$ used a least squares cost function during peak fitting, where the
 54 deviation between each pixel intensity and the best fit model was given equal weight. This

55 simplification neglects the fact that Raman noise is heteroscedastic, where the error increases as
56 a function of the intensity.

57 From v>= 1.0.19, DiadFit fit functions have a new input argument so that the weighting strategy
58 can be changed. For reproducibility, we preserve an option to fit peaks with the old least squares
59 approach (setting the parameter 'minimise' to 'least squares'), but by default the code now uses
60 a weighted least square approach (minimise='least_squares'). According to the documentation,
61 Lmfit defines the weights to use to calculate the fit residual as follows:

62 $\text{weights}*(\text{data}-\text{fit})$

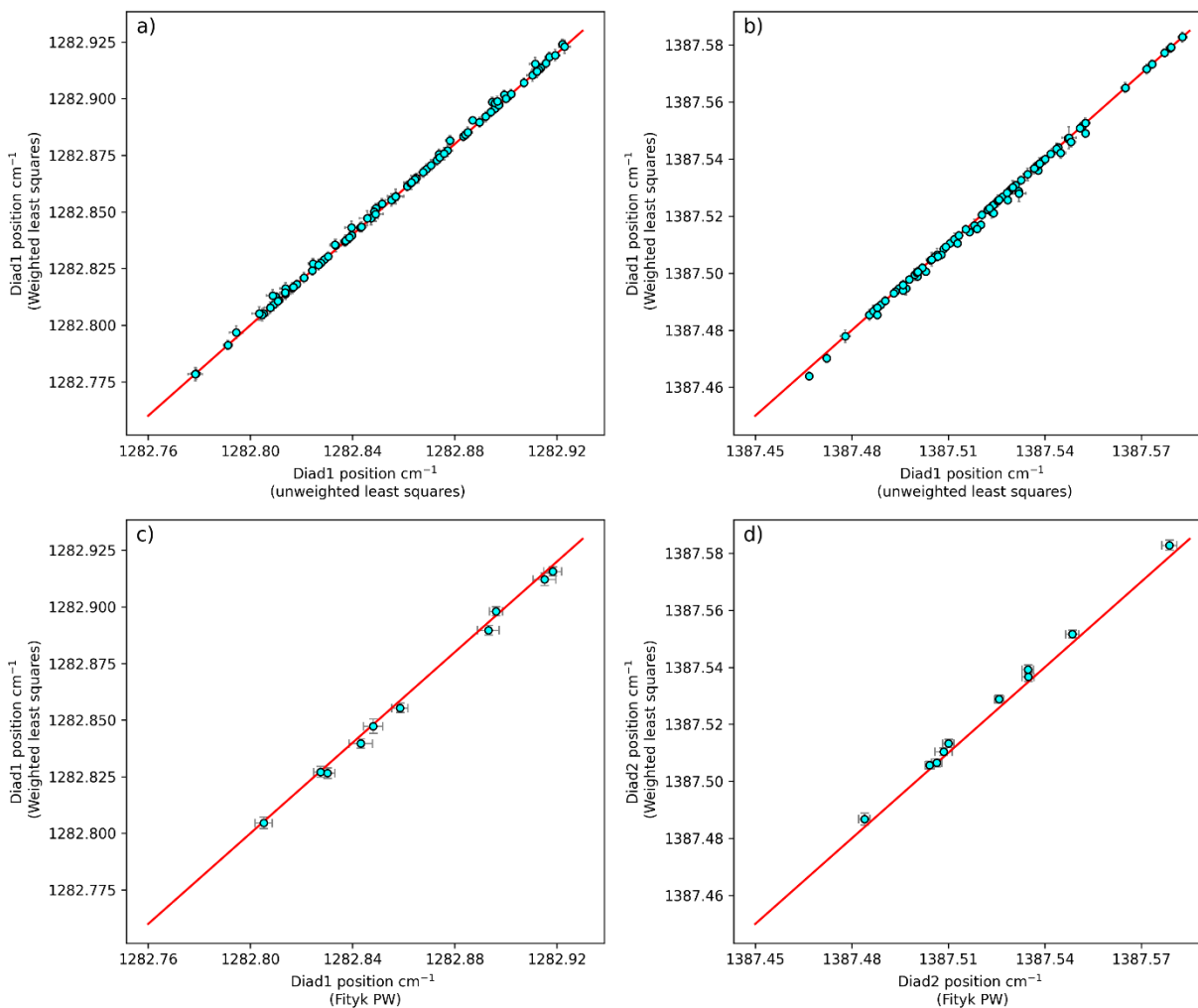
63 Following poisson statistics, we assume that $\sigma = \sqrt{N}$, where N is total counts. In lmfit, the weights
64 are inputted as 1/error ([python 3.x - Taking experimental errors into account in lmfit - Stack
65 Overflow](#)). Our code assigns a weight to each pixel as follows:

66 $\text{Weight} = \frac{1}{\sqrt{N}}$

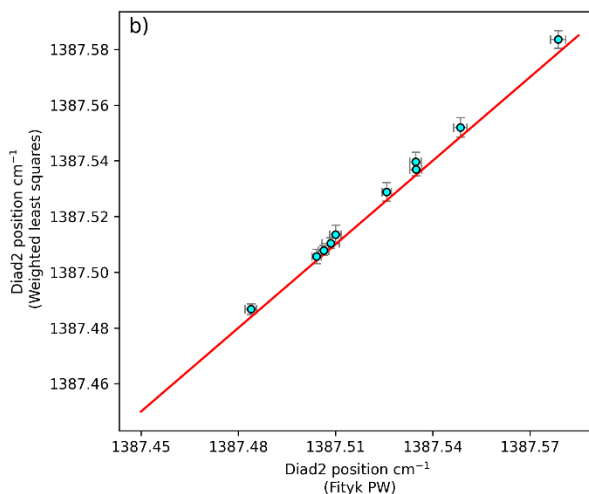
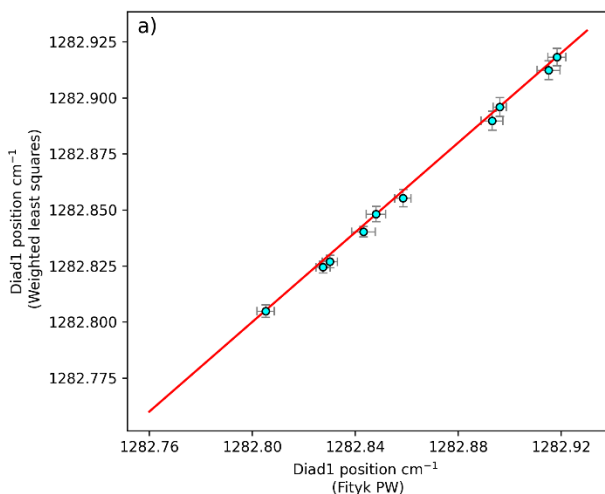
67 When extracting errors from the covariance matrix, lmfit has two options: `scale_covar=True`, and
68 `scale_covar=False`. The lmfit documentation suggests that in cases where the absolute value of
69 the uncertainty is unknown on each point, but the peak fit is good, `scale_covar` should be set to
70 True, which rescales the uncertainty in the data such that the reduced chi-square is 1.

71 First, we assess how changing the minimization regime from least-squares to weighted least
72 squares affects the peak positions using the room heating test data shown in Fig. 6. The shift in
73 peak position is very small, and within the uncertainty of the two fit routines (Supporting Fig. S6).
74 We also fit a subset of 10 spectra in Fityk. Because of the effect of user-choice related to manual
75 fitting in the Fityk GUI, both C. DeVitre and P. Wieser independently conducted the fits using
76 pseudoVoigt peaks. Where possible, a Gaussian was added under the peaks that were elevated
77 above the background, but we found it very challenging to get a good fit in Fityk (where clicks are
78 required), so for many spectra a Gaussian could not be reliably added as it resulted in a visually
79 worse fit. The peak positions obtained from DiadFit are also extremely similar to Fityk, but more
80 so for Diad1 where the Gaussian is not required, and less so for Diad2 where a gaussian is
81 required (Fig. S7-S8). We turn off the Gaussian argument in DiadFit to better compare, which
82 causes the diad2 peak positions to show a much better alignment (although the fits are now
83 noticeably worse to the data, Supporting Figure S8).

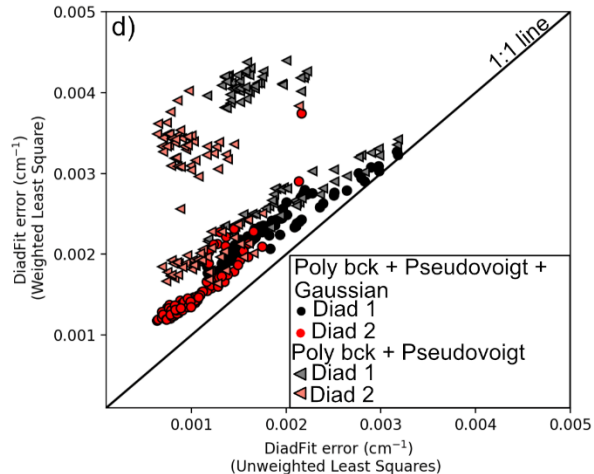
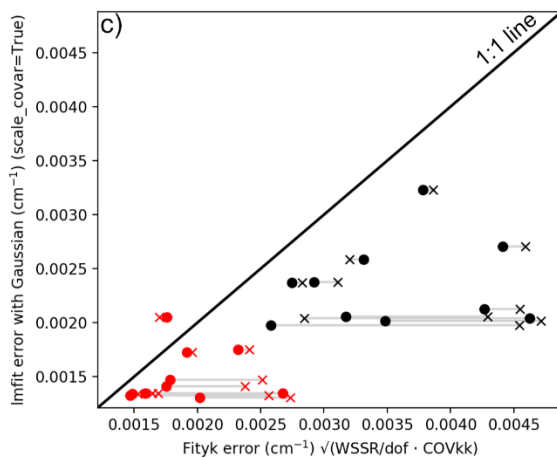
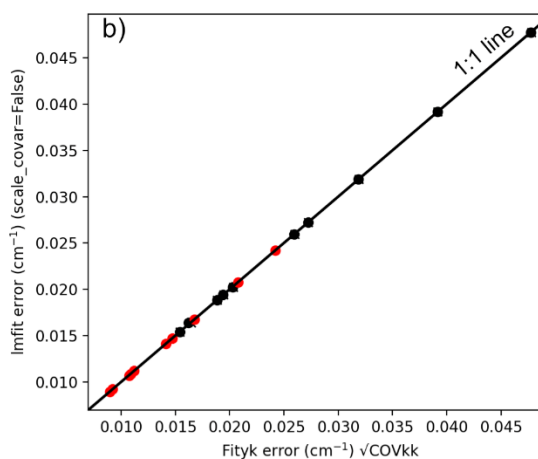
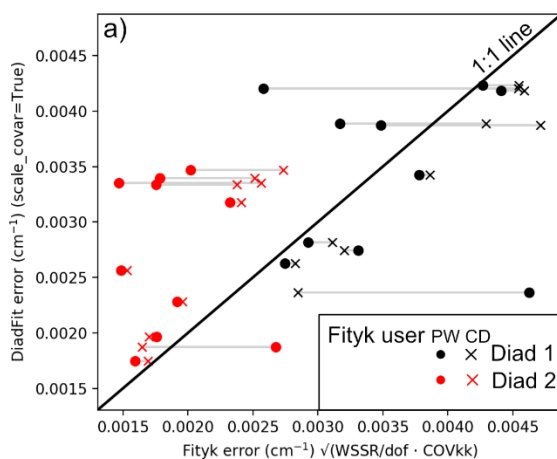
84 We also compare our errors to those from Fityk to benchmark the new fitting routine. Like lmfit,
85 Fityk can return two types of errors. The default output is labelled '*±standard errors:*
86 *sqr(WSSR/DoF COV_kk)*' which are equivalent to `scale_covar=True` in lmfit, and there is also an
87 option to select errors labelled '*sqr(COV_kk)*', which are equivalent to `scale_covar=False` in lmfit
88 (Supporting Figure S8). The unscaled errors are completely implausible given agreement
89 between lmfit and Fityk, and the long- and short-term laboratory stability we observe (Figure 5
90 and 6), but agree well between these two bits of software.



91
 92 **Supporting Figure S6** – Comparing different peak fitting routines. a-b) comparing peak positions
 93 determined from unweighted least squares (DiadFit v<1.0.18) and weighted least squares
 94 (DiadFit v≥1.0.19). The red line shows a 1:1 line, and the two methods do not deviation from the
 95 1:1 line in a meaningful way considering the peak fitting errors. For a subset of 10 spectra, we
 96 also fit the data using Fityk.



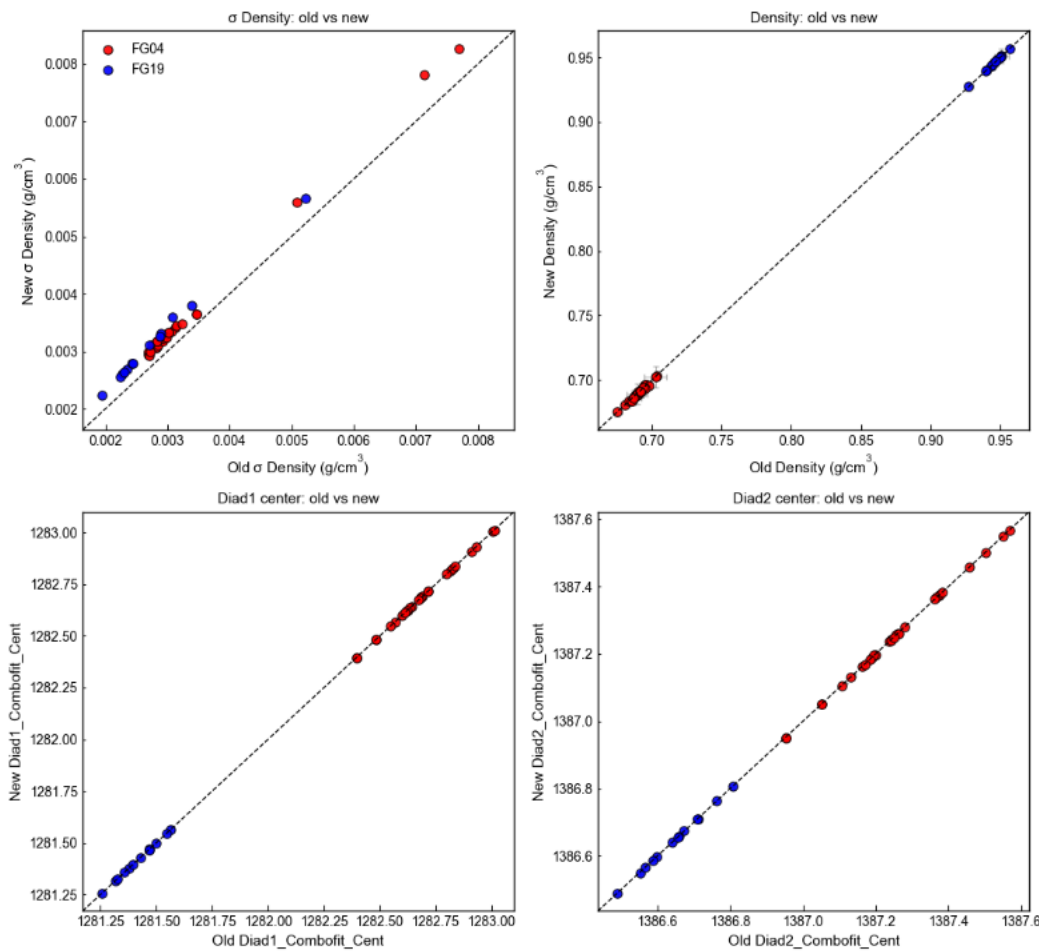
97
 98 **Supporting Figure S7-** Comparing Fityk to DiadFit where no Gaussian was used to improve the
 99 fit to stronger peaks.



100
 101 **Supporting Figure S8 –** Comparison of errors determined from different methods. a) scaled
 102 errors from DiadFit (excluding the Gaussian) and Fityk. For each spectra, errors determined from

103 P. Wieser (PW) and C. DeVitre (CD) are shown (dot and x joined by a line). It is clear that small
104 changes in clicked parameters such as user backgrounds, peak positions have a large influence
105 on the quality of the fit, and thus scaled errors. b) Unscaled errors are identical for Fityk and Imfit,
106 regardless of user (because these errors are insensitive to the quality of the fit). c) Same
107 comparison as in a), but using fits with the Gaussian function enabled in DiadFit. This reduces
108 the errors. d) Comparison of errors from the unweighted least squares and weighted least squares
109 approaches in Diadfit, with and without Gaussian peaks enabled.

110

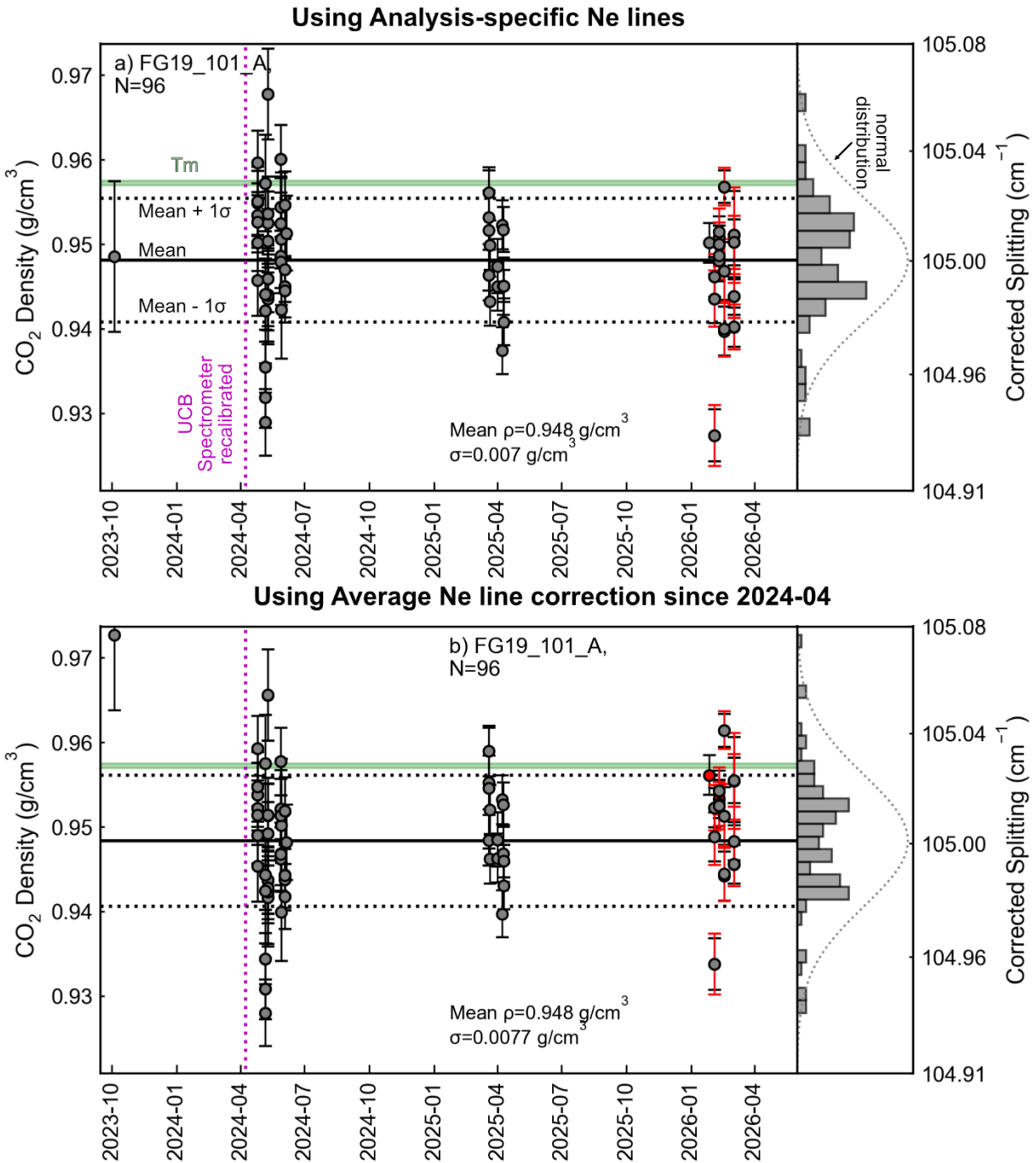


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112 **Supporting Figure S9** – Comparison of peak fitting positions and errors using least squares (old)
113 and weighted least squares (new) for standard analyses from 13 analytical separate analytical
114 session.

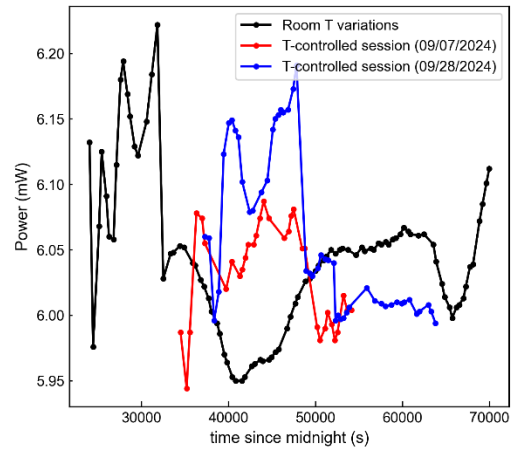
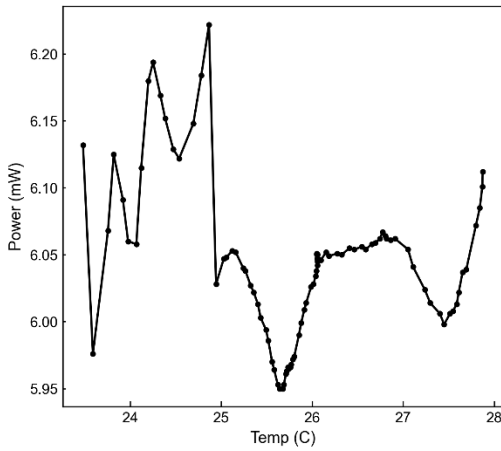
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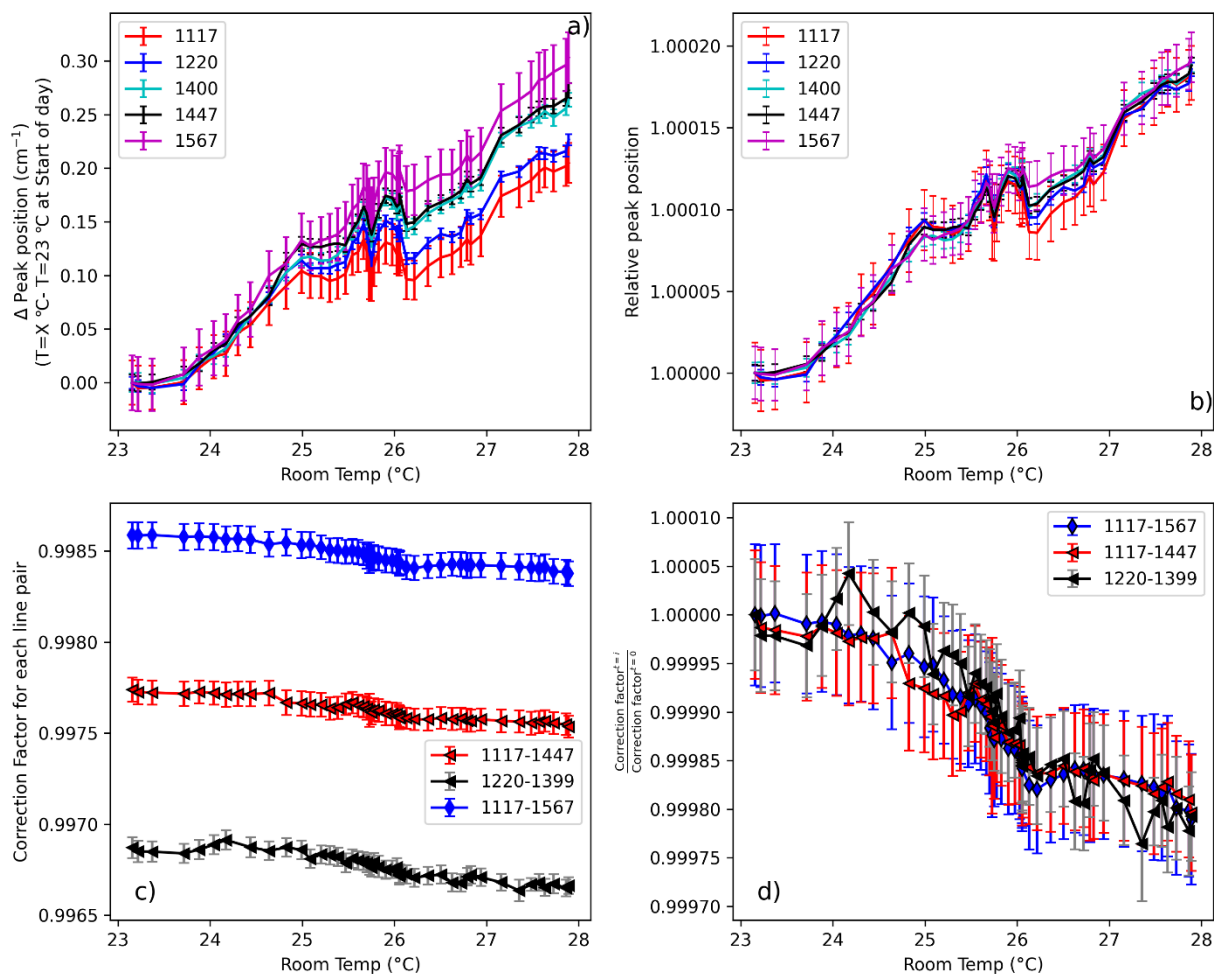
118 **Supporting Figure S10** - Long-term variability in standard values collected at UCB for standard
 119 FG19_101_A. a) Data reduced using a drift correction model for each analytical session. b) Data
 120 reduced using the average Neon line correction factor for all standard acquisitions after the
 121 instrument recalibration (marked with a magenta line). The green bar shows the density
 122 determined by Microthermometry, accounting for the uncertainty in the temperature reading of the
 123 stage. The black solid line shows the mean Raman value for the standard, and the dotted black
 124 lines $\pm 1 \sigma$. The distribution of densities is shown as a histogram, compared to a normal
 125 distribution based on the mean and standard deviation.



126

127 **Supporting Fig. S11** – Plot of laser power measured with the TruePower meter against
128 Temperature during the room temperature heating test to show there are no coherent variations
129 outside of the ± 0.1 mW read out precision of the TruePower meter. The variability seen is of similar
130 magnitude to that seen during two analytical sessions with the HVAC system regulating
131 temperature (red and blue lines from Kahl et al. 2025).

132



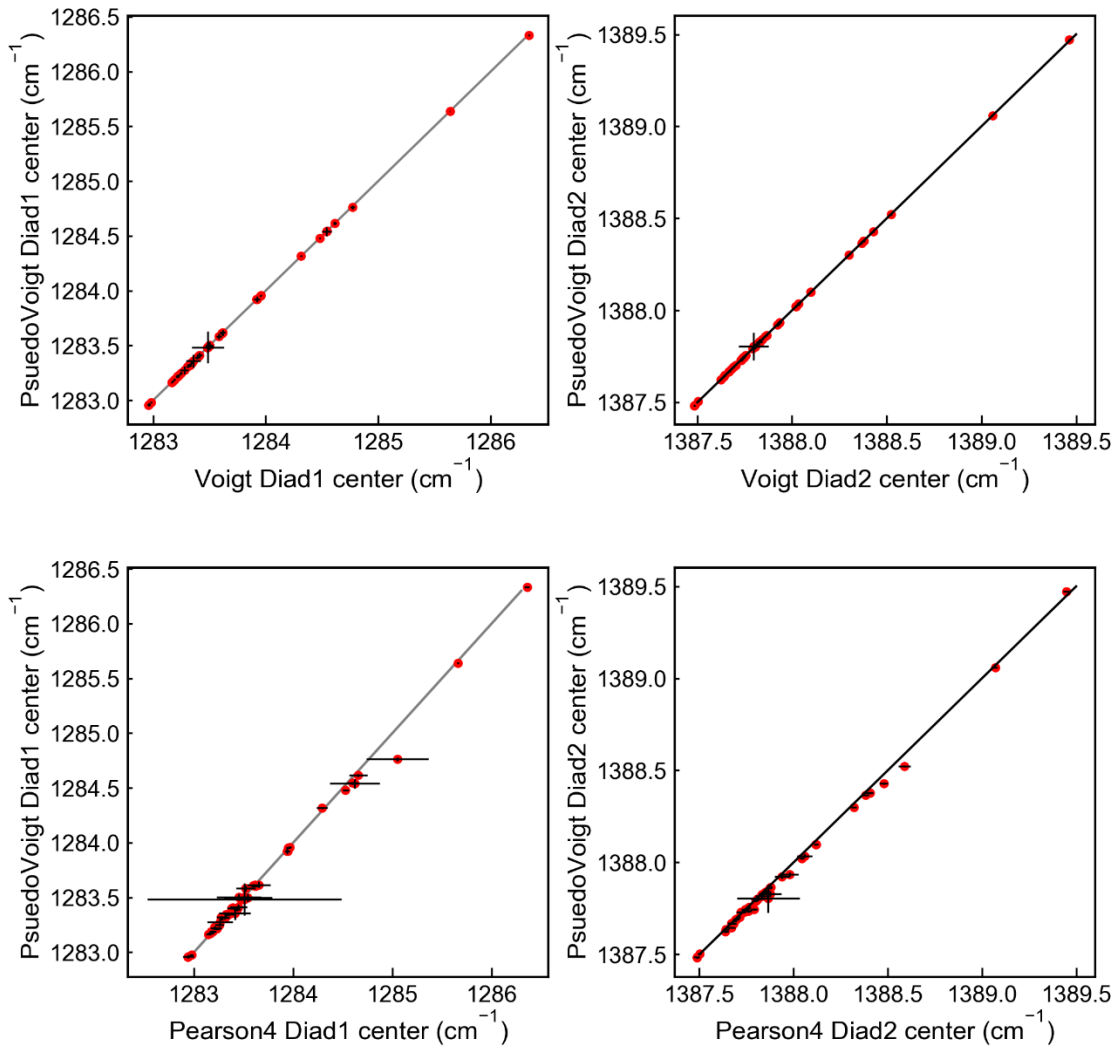
133

134 **Supporting Figure S12** – Change in Neon line position, and correction factors calculated for
 135 different Neon line pairs during perturbations in the room temperature of 5 °C. a) Absolute change
 136 in peak position, calculated as the peak position at $T=X$ °C minus the peak position in the first
 137 Neon Acquisition b) Proportional change in peak position, calculated as the peak position at $T=X$
 138 °C divided by the peak position in the first Neon Acquisition. b) Absolute correction factor,
 139 calculated as the difference between the measured position of the two Neon lines and the
 140 theoretical distance. c) Change in correction factor, defined as the correction factor at each step
 141 divided by the correction factor in the first acquisition. All Neon lines follow the same trend within
 142 uncertainty so would have the same effect on removing instrument drift due to temperature
 143 fluctuations.

144

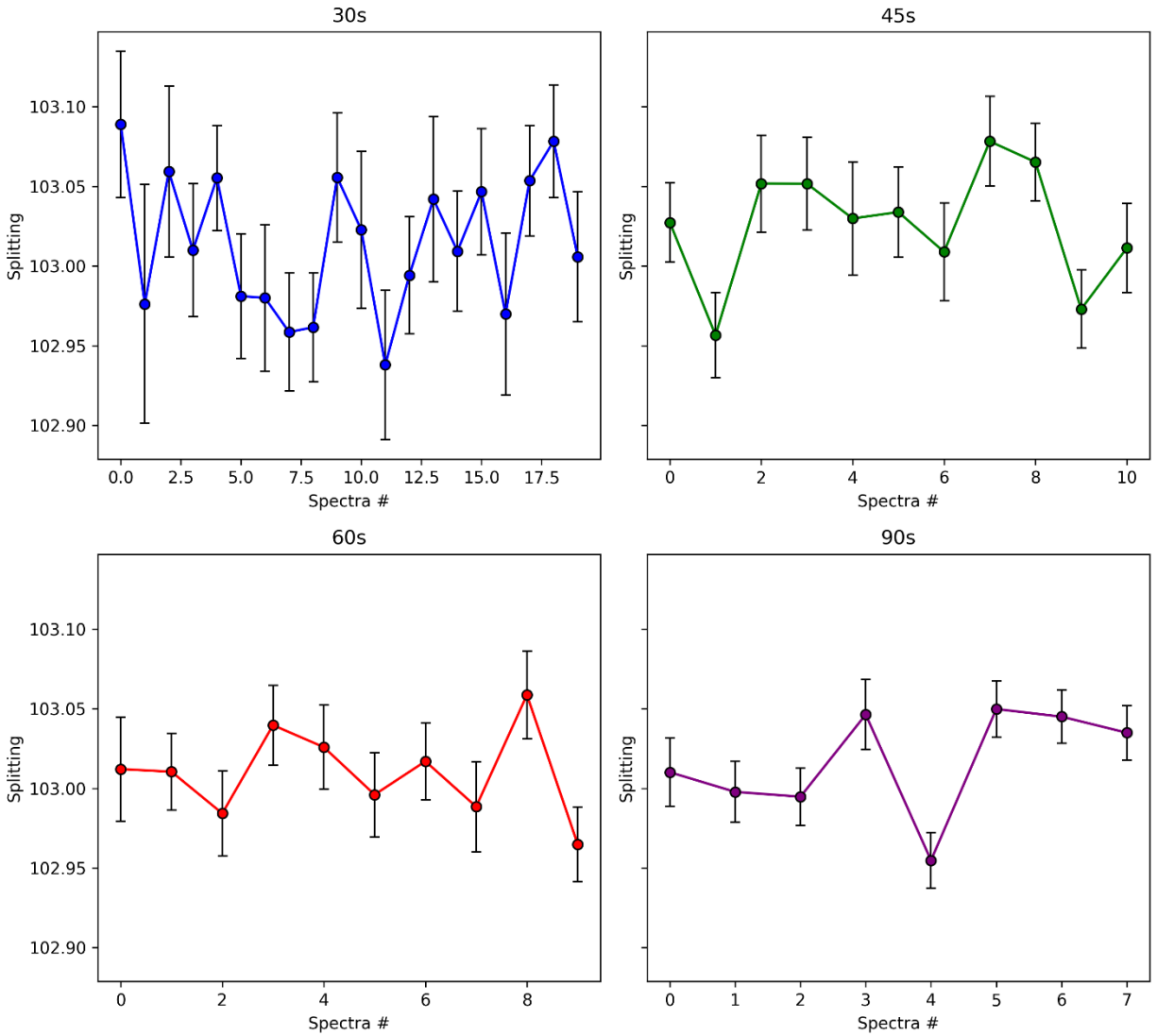
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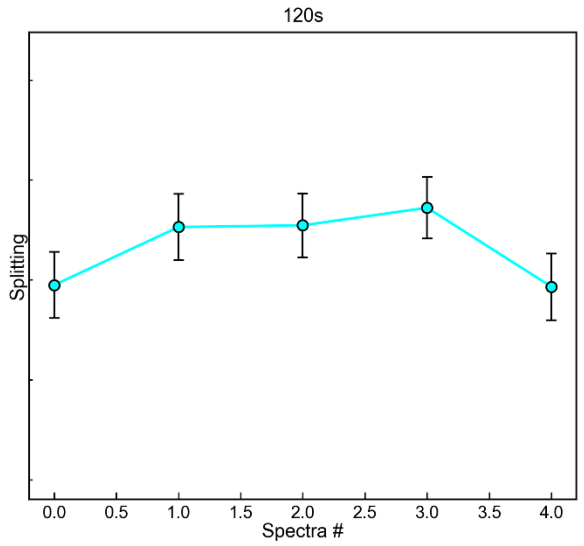
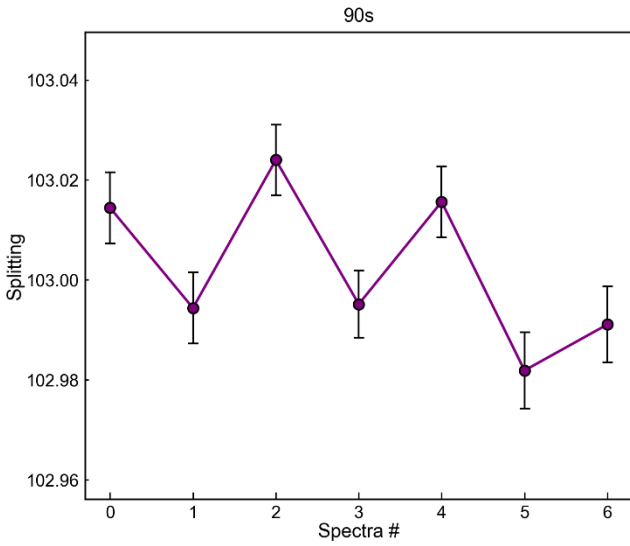
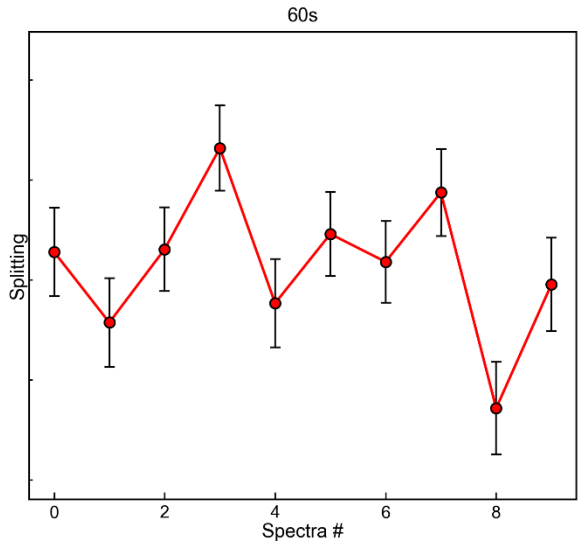
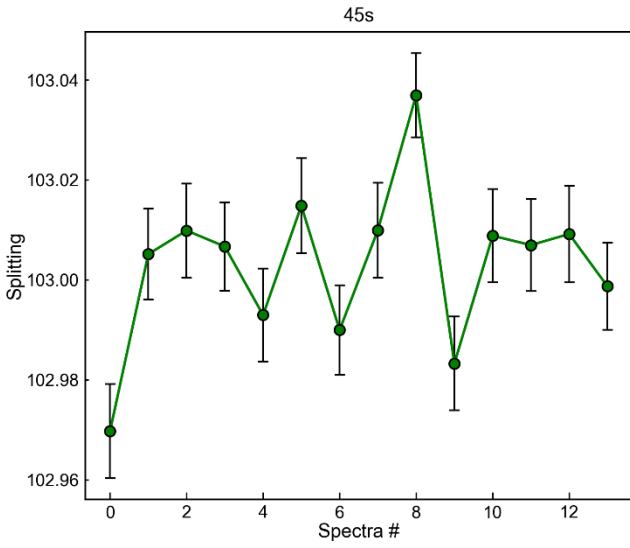
148 **Supporting Figure S13** – Comparison of different PDF fits for Fogo fluid inclusions (as for Fig 8
 149 in the main text).



150

151 **Supporting Figure S14** – Variability in splitting from each spectra for N=1 accumulation for a melt
 152 inclusion vapour bubbles from Mauna Ulu ($MI1 \sim 0.015 \text{ g/cm}^3$, Wieser et al., 2019). The errors bars
 153 are 1 sigma errors from DiadFit.

154

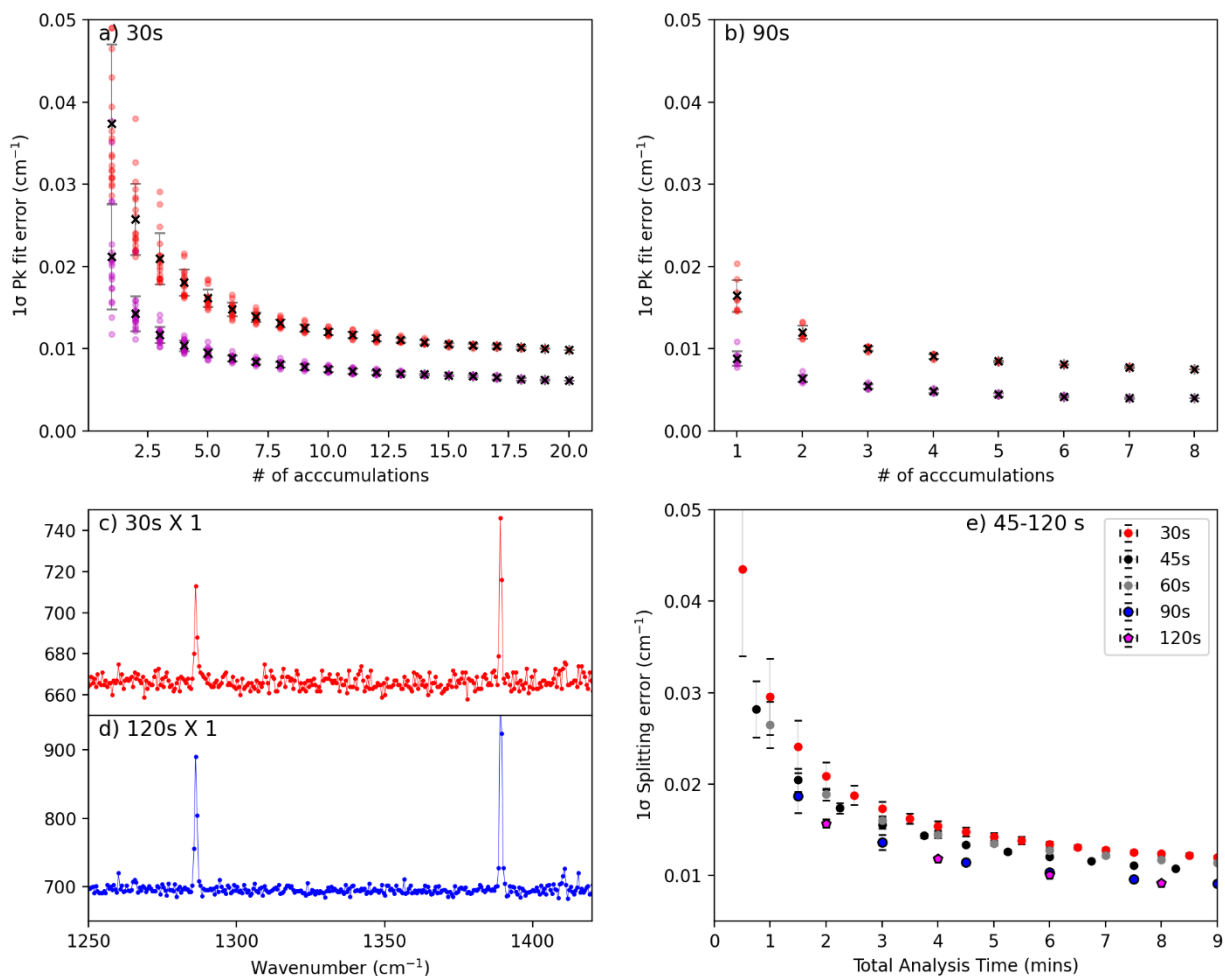


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156 **Supporting Fig. S15-** Same for the Capillary capsule.

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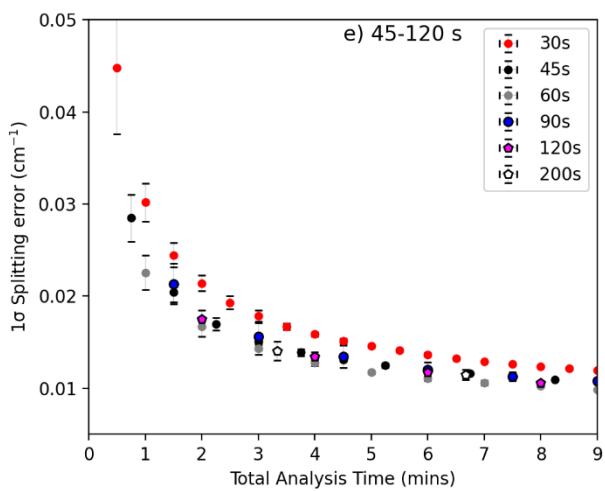
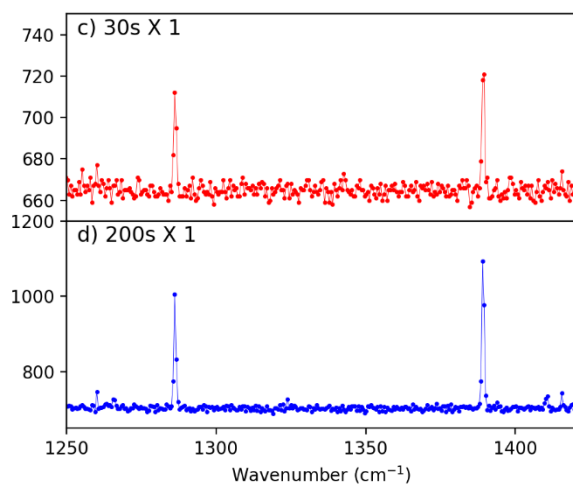
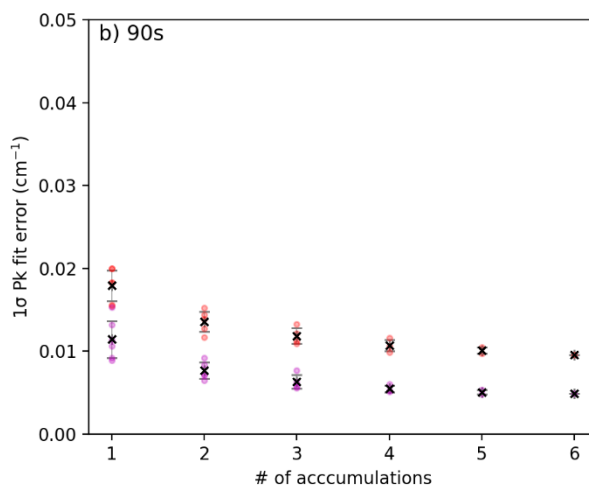
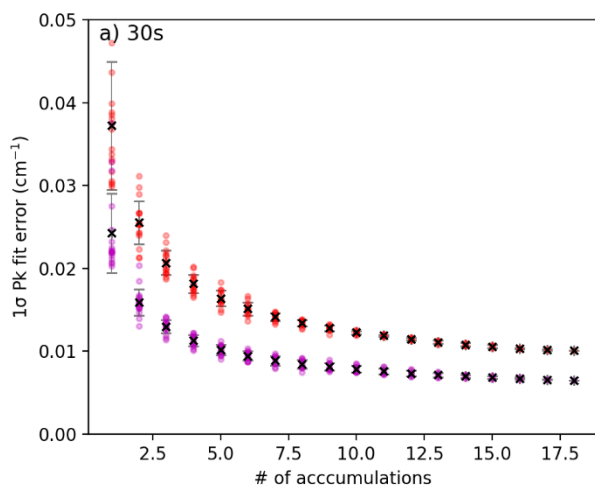
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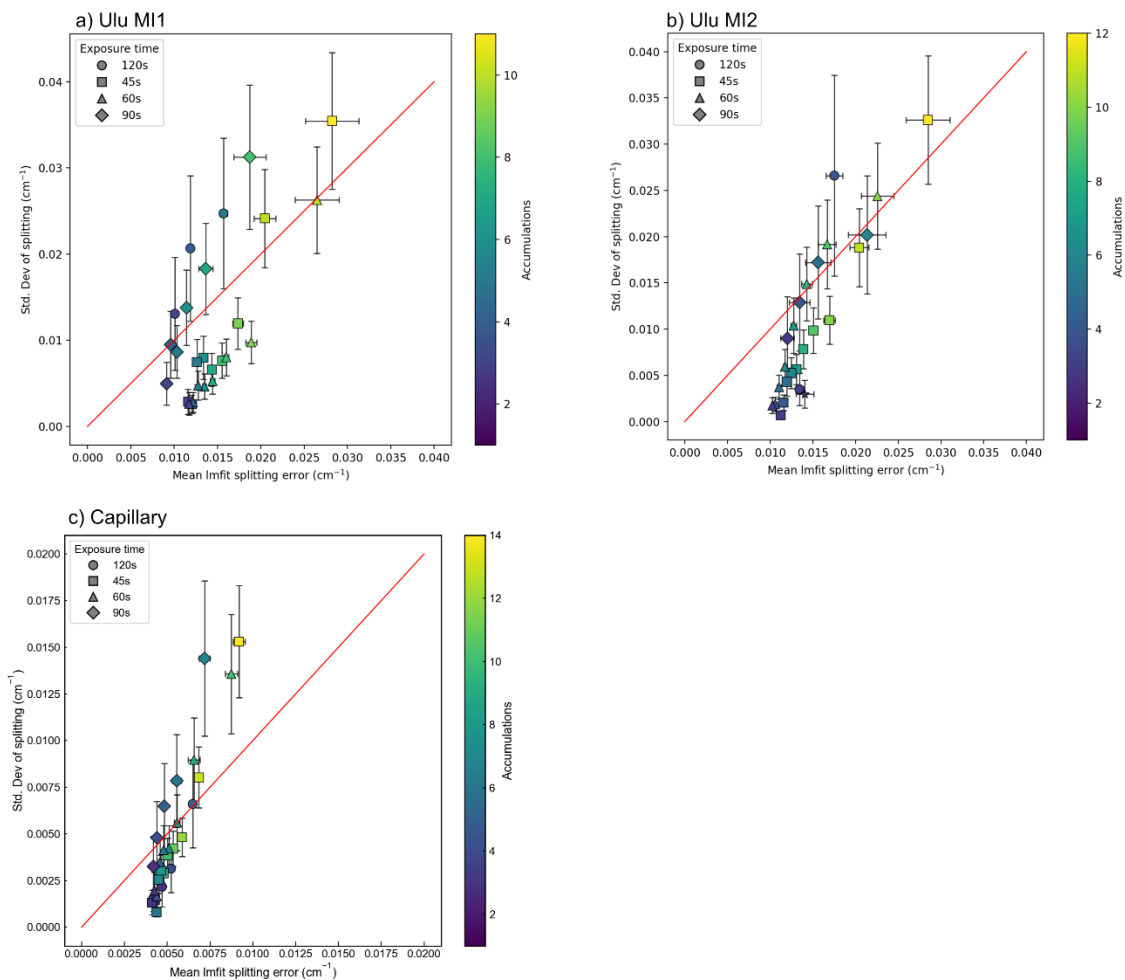
160

161 **Supporting Figure S16** – Investigating peak fitting error as a function of integration time and
 162 accumulations using a melt inclusion vapour bubbles from Mauna Ulu (MI1~0.015 g/cm³, Wieser
 163 et al., 2019). a-b) fitting error estimated from Imfit error for the peak position of diad1 and diad2
 164 for 30 and 90s acquisitions. The x axis shows the number of spectra that were averaged prior to
 165 fitting the peaks (analogous to accumulation). The dots show an individual fitted spectra (e.g.
 166 spectra 1 to 10 for accumulations =1, average of spectra 1-2, 2-3 3-4 for n=2 accumulations etc.).
 167 The cross shows the mean Imfit error, and the errorbar shows the std. dev of the Imfit error for
 168 each number of accumulations (e.g. 14 individually fitted spectra for n=1 accumulations). c-d)
 169 Example spectra for 45s x 1 and 90s x 1. e) Error on calculated splitting after combining different
 170 numbers of individual acquisitions to give different acquisition times. For example, 5 different 60s
 171 acquisitions were averaged for the 5 minute total acquisition time.



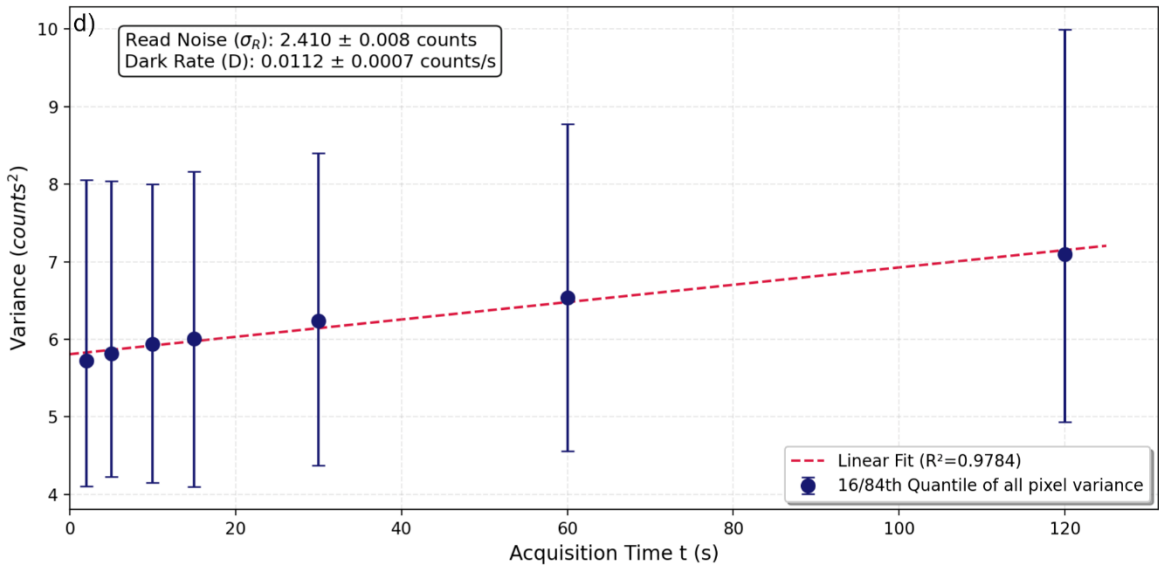
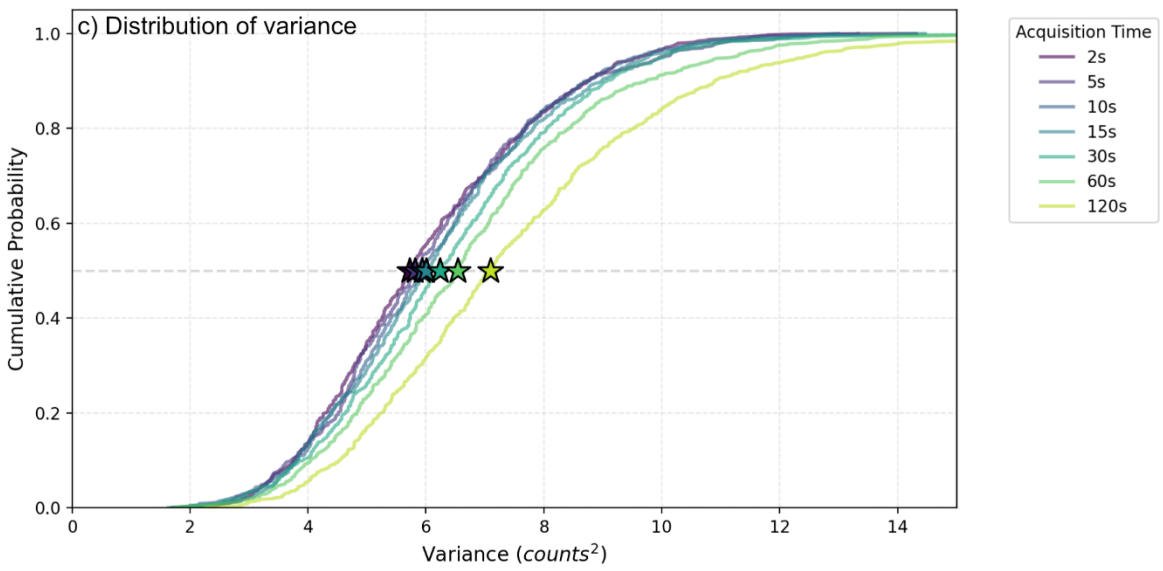
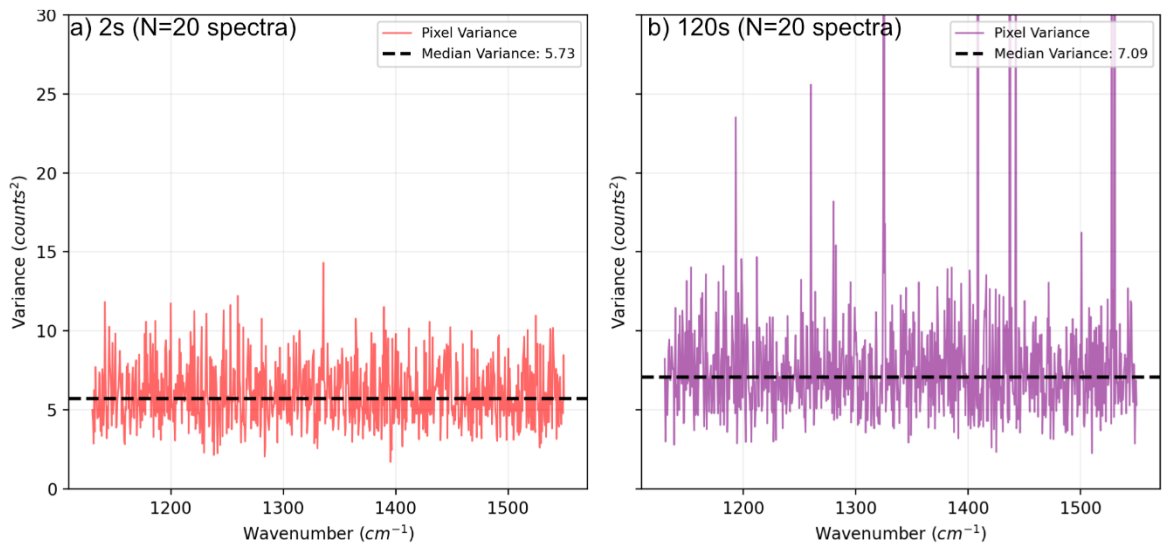
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173 **Supporting Figure S17** – Same as above but for MI2, including a 200s acquisition.



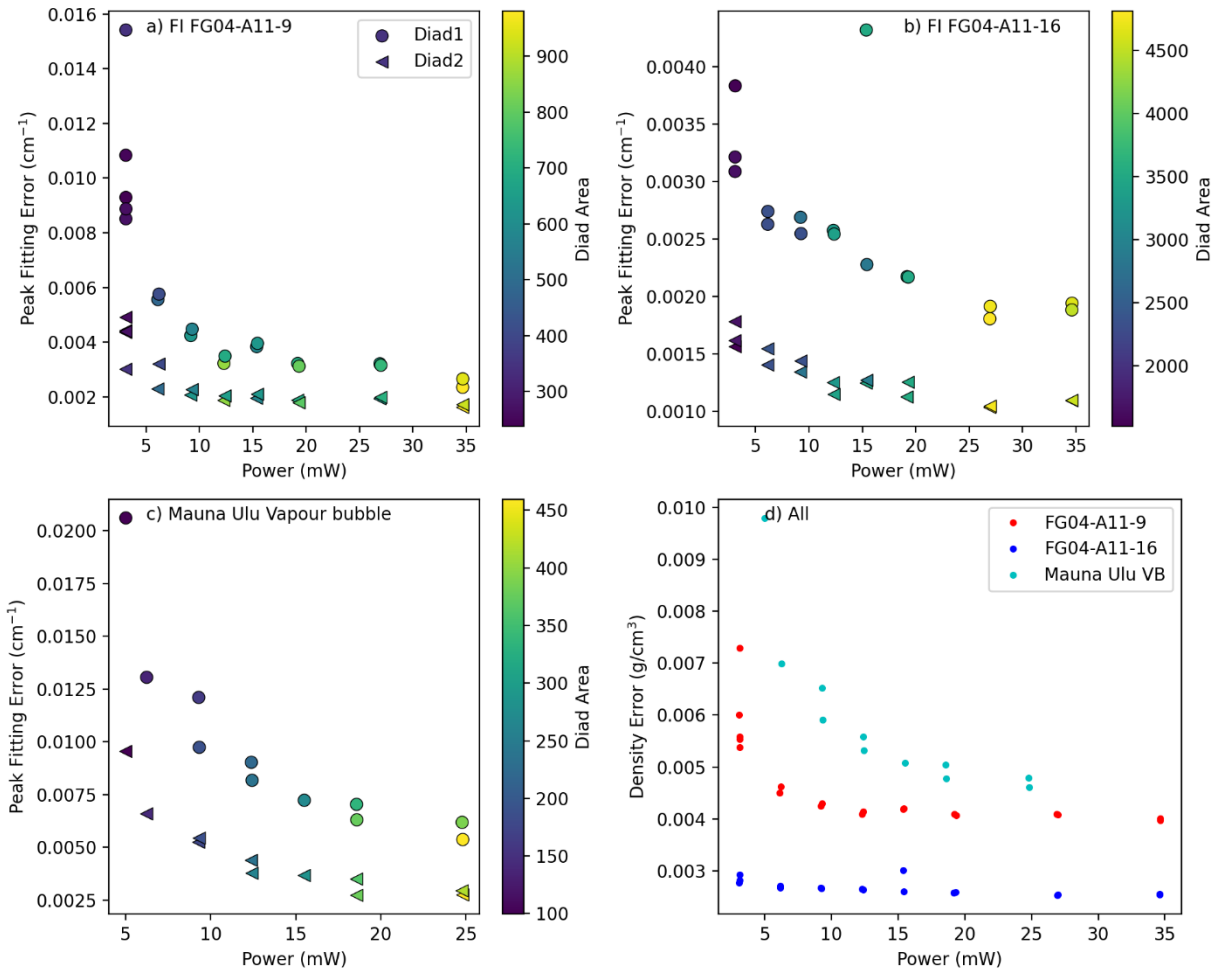
174

175 **Supporting Figure S18** – Comparison of mean Imfit splitting error (e.g. average error from N=5
 176 spectra), and the standard deviation of the splitting determined from the 5 fits. The red line shows
 177 a 1:1 line. The x error bar represents the standard deviation on the calculated Imfit error, and the
 178 y error is calculated from the standard deviation based on the number of samples averaged.



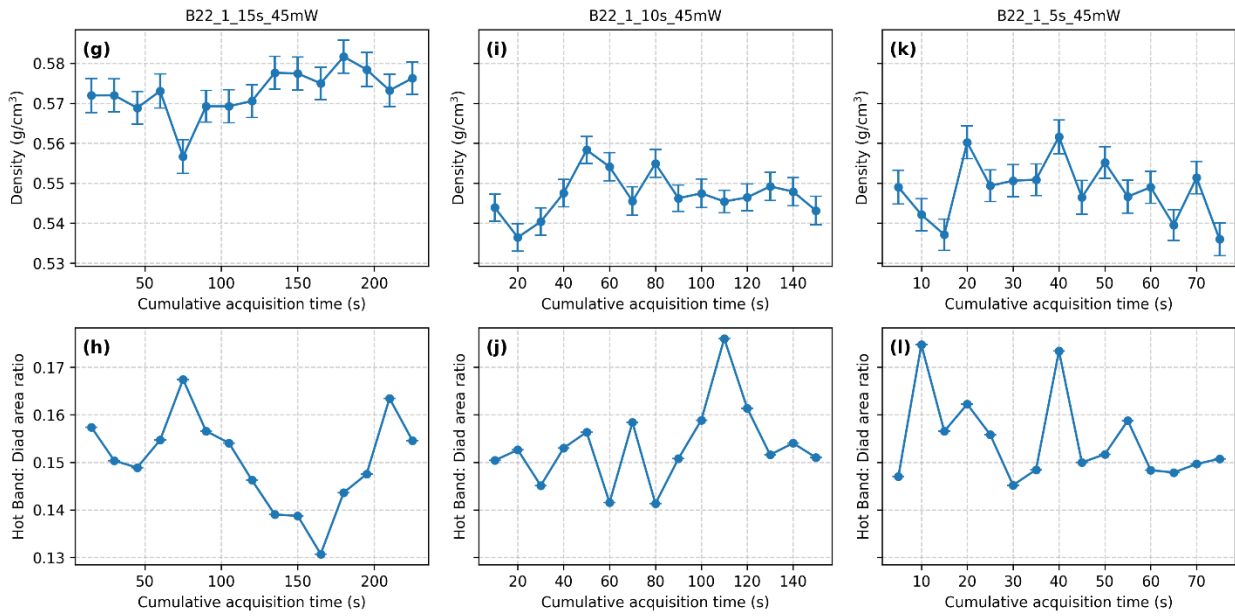
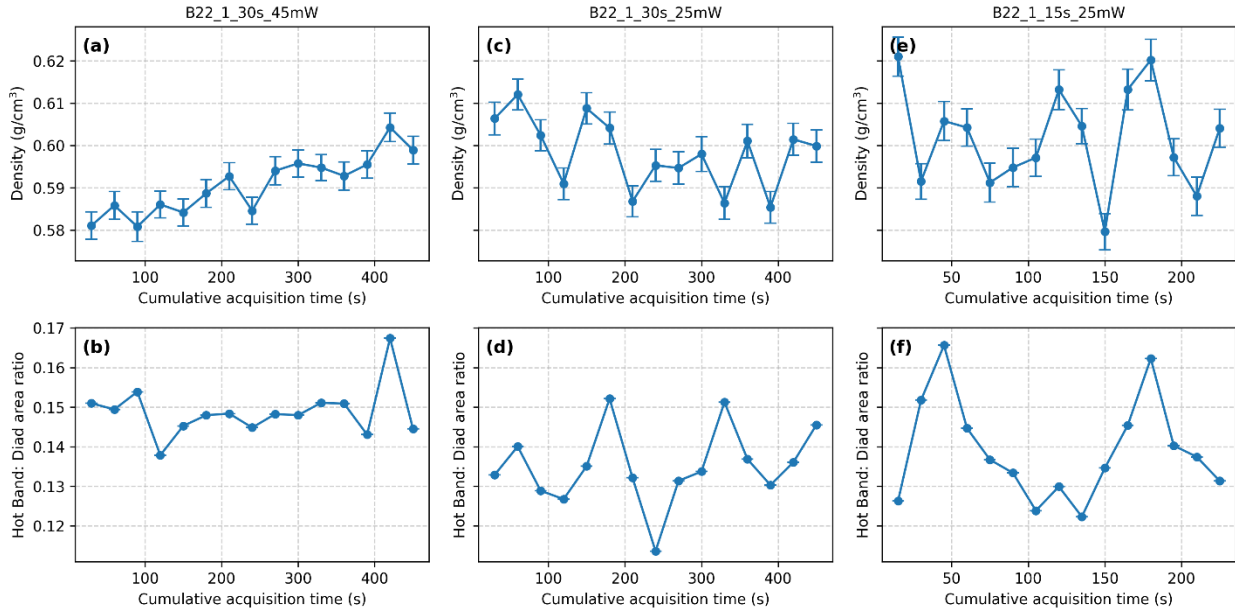
180 **Supporting Fig. S19** – Estimating dark and read noise from pixel variance with the laser and
 181 room lights off. a-b) Variance of each pixel across 20 repeated acquisitions at 2s and 120s
 182 acquisition times. The larger number of high variance pixels at 120s reflects cosmic rays (large
 183 ones were removed, but small ones remain). c) Distribution of variance as a function of acquisition
 184 time. The median clearly increases, even though the distributions overlap. d) Estimating read
 185 noise and dark noise from the increase in median variance with acquisition time. The fact that the
 186 trend is so subtle relative to the spread of each acquisition time indicates that dark noise is very
 187 small on this instrument.

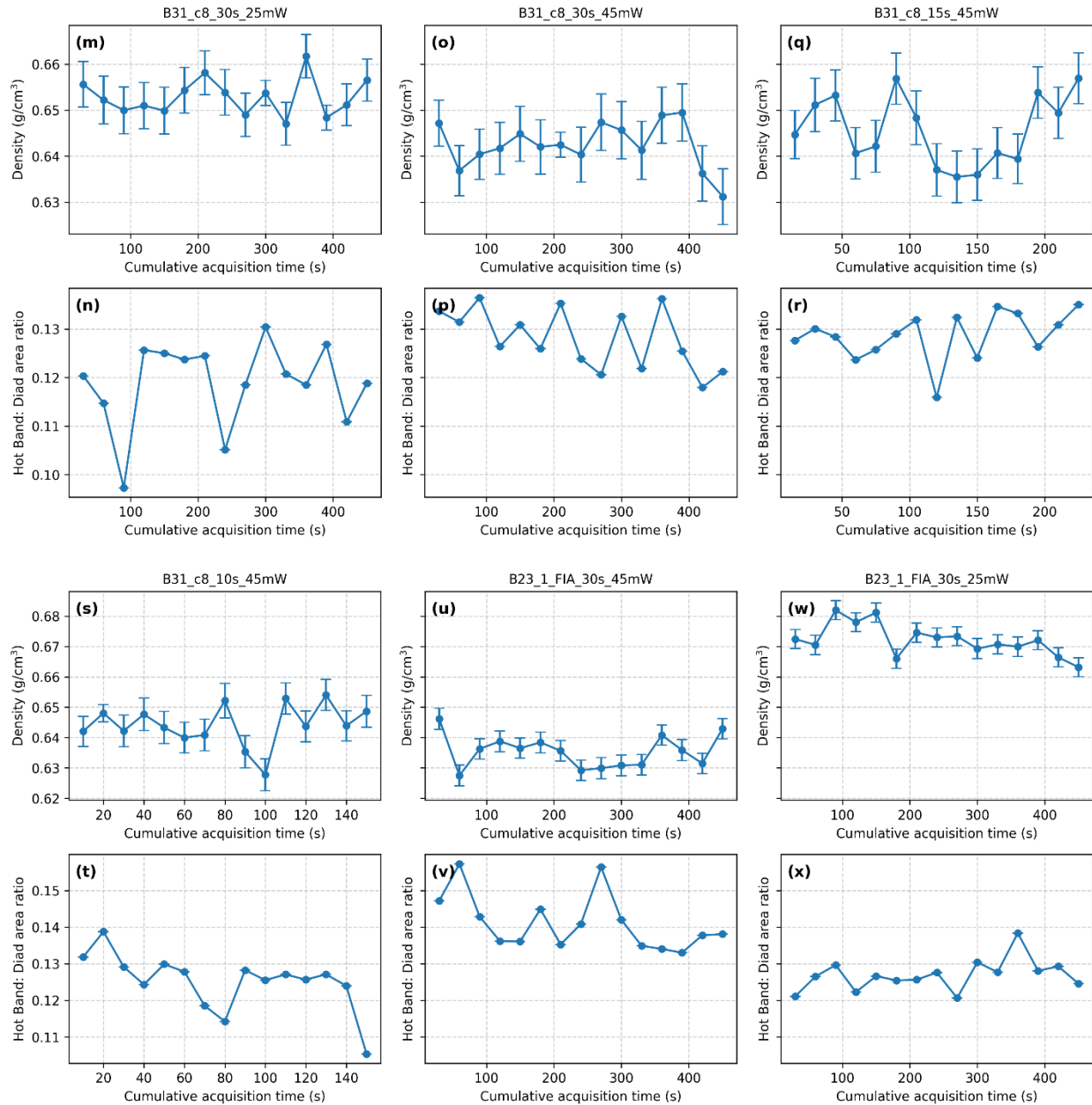
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 189



190

191 **Supporting Figure S20**– Change in peak fitting errors with increasing laser power. a-b) There is
 192 a clear drop in the peak fitting error for diad1 and diad2 as laser power is increased for Fogo fluid
 193 inclusion (FG04-A11-9), and for a melt inclusion vapour bubble from Mauna Ulu, Kīlauea (Wieser
 194 et al., 2019). (c. d) Same datasets but showing the error in calculated density. Laser power is at
 195 the sample, calculated from the TruePower reading multiplied by the correction factors for each
 196 objective measured using the laser power meter. This is true for all subsequent plots in this
 197 supplement.



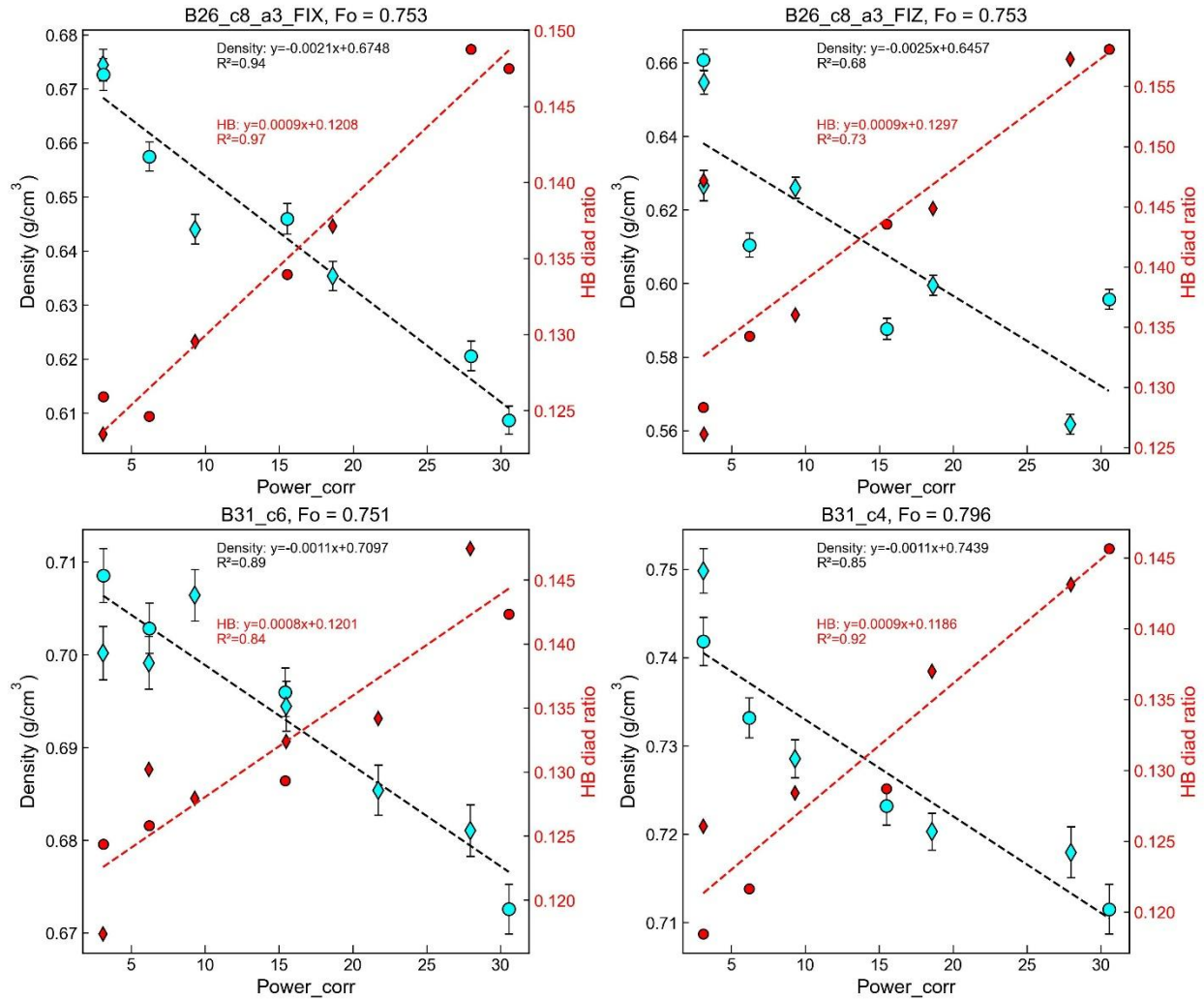


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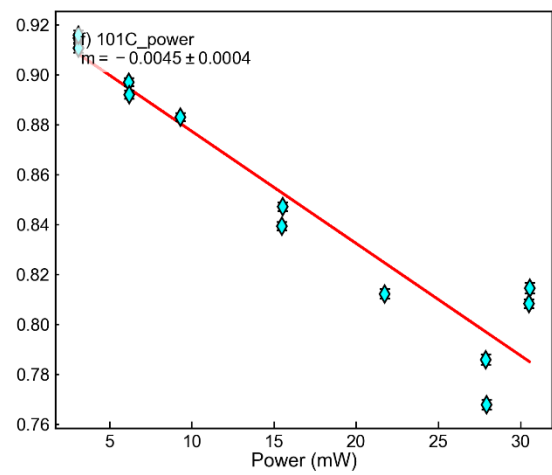
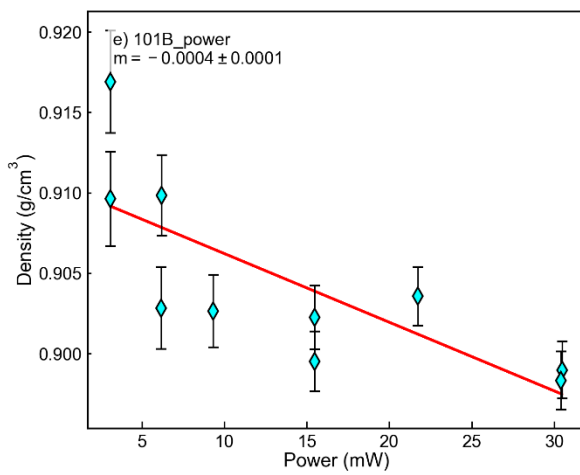
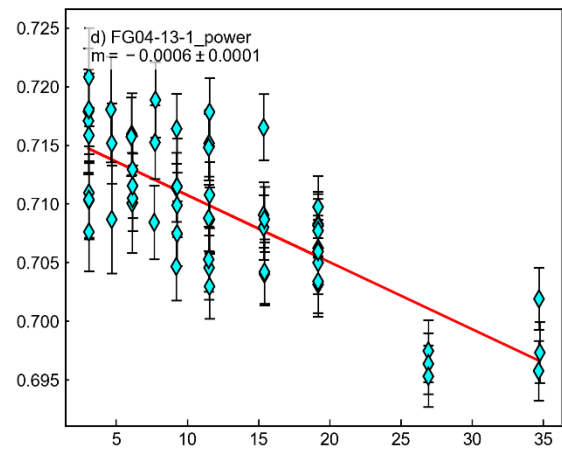
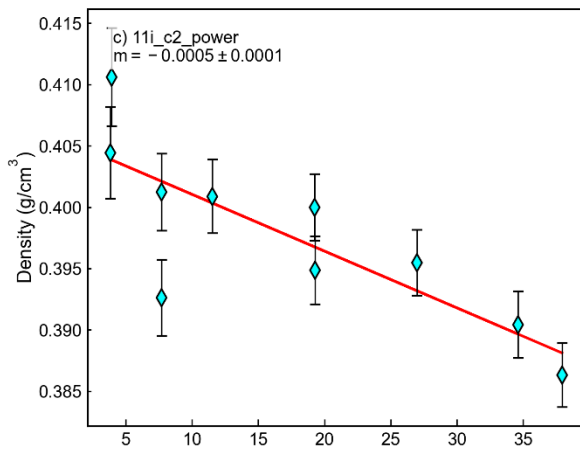
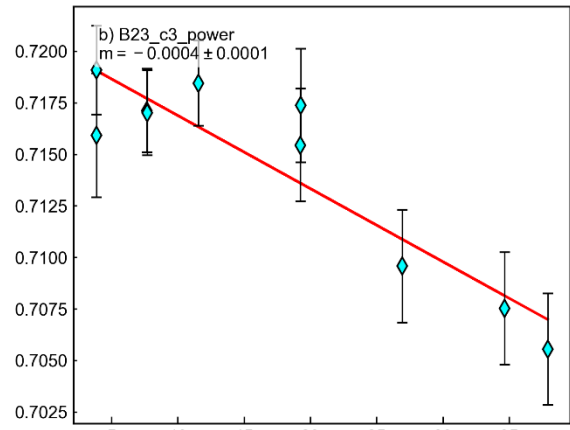
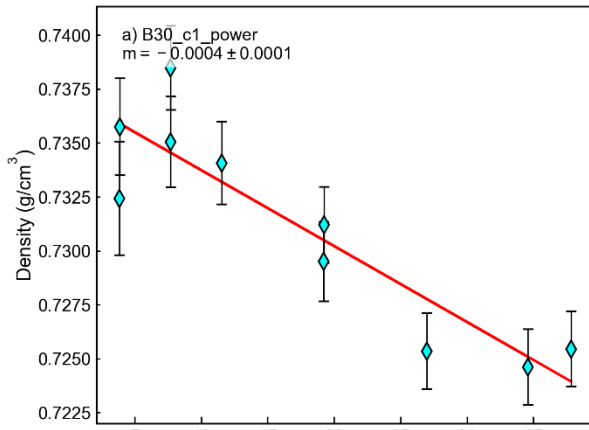
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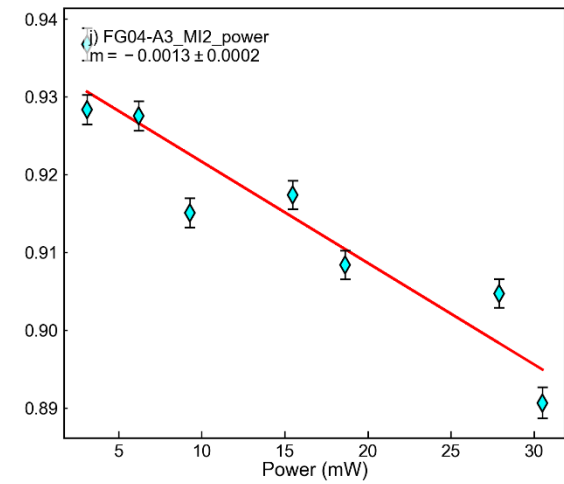
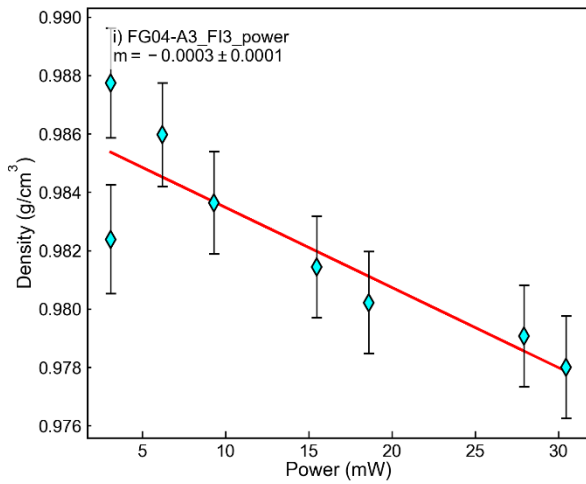
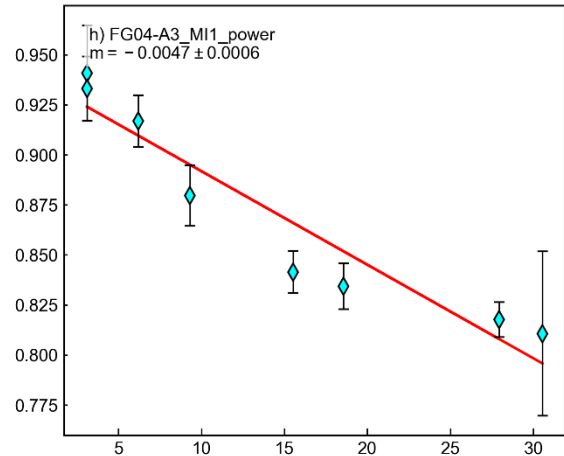
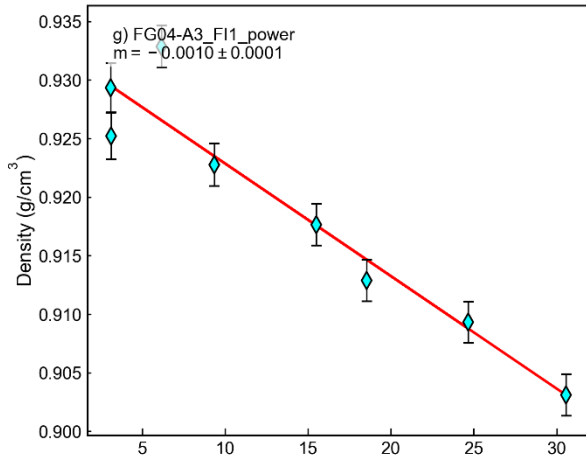
203 **Supporting Fig. S21** – Change in density and HB:diad area ratio from a WITEC slow series for
 204 the acquisition times and laser powers shown in the title to investigate how long it takes a sample
 205 to reach steady state. For example, the panel labelled B22_1_30s_45mW shows acquisitions
 206 every 30s with the power set to 45 mW in the software using the TruePower meter. The laser is
 207 continuously on during the acquisition. There is no clear change in these trends, indicating the
 208 sample reaches a stable temperature very quickly after being exposed to the laser.



209

210 **Supporting Figure S22**–Comparison of Density vs. power, and HB:diad ratio vs power (defined
 211 as (Area HB1 + Area HB2)/(Area Diad1 + Area of Diad2). The anticorrelation shows strong
 212 evidence for laser heating.

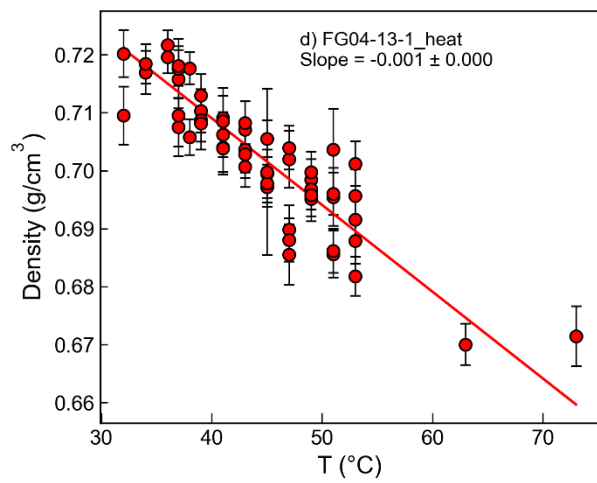
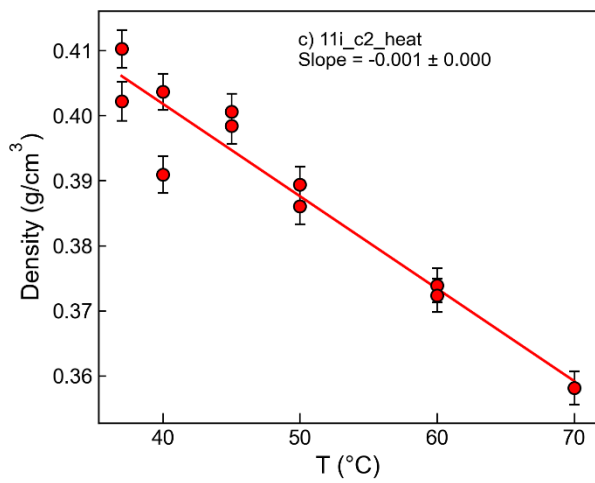
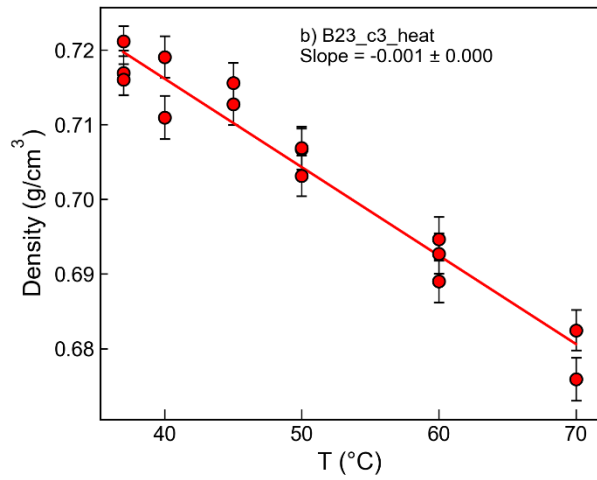
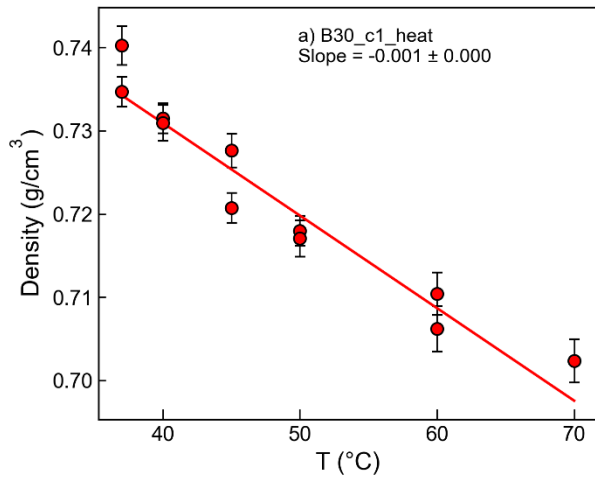


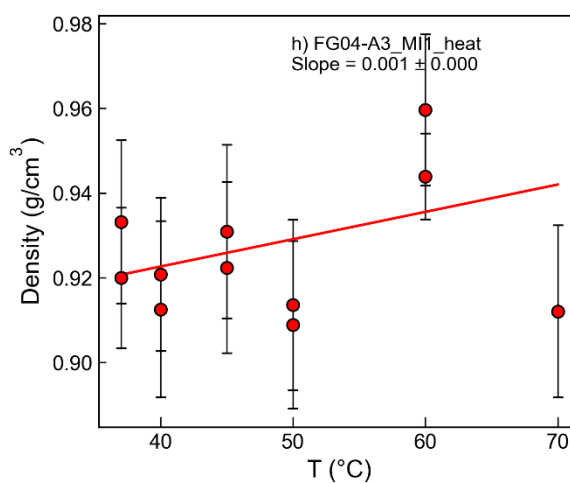
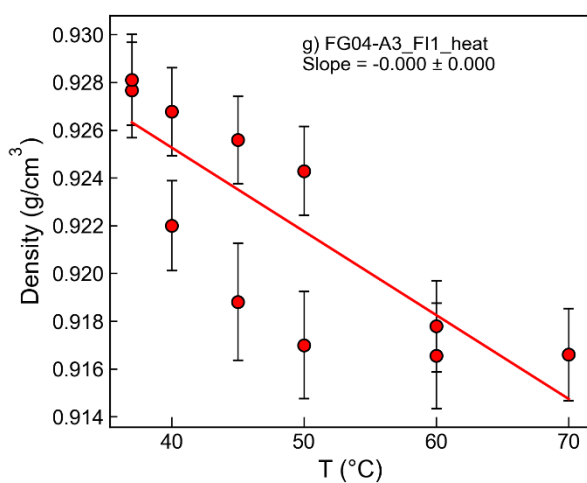
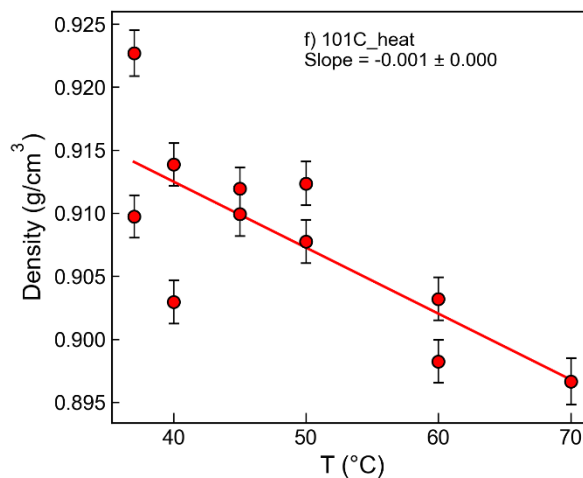
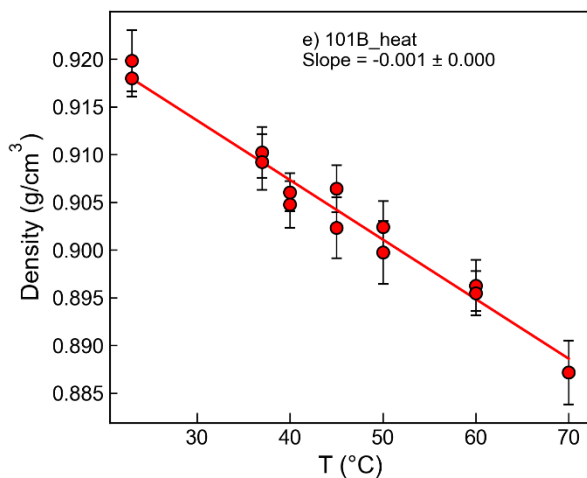


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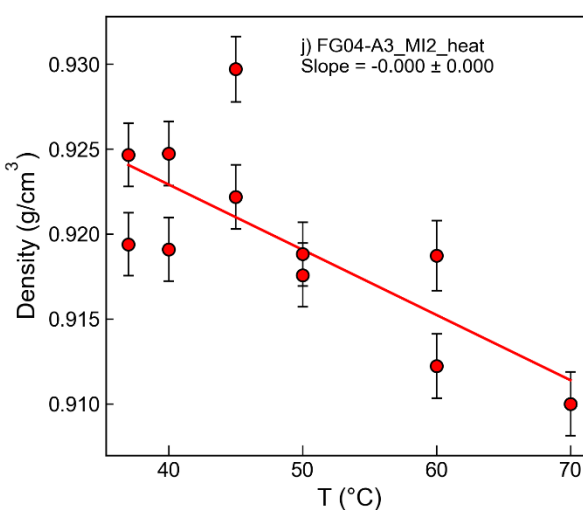
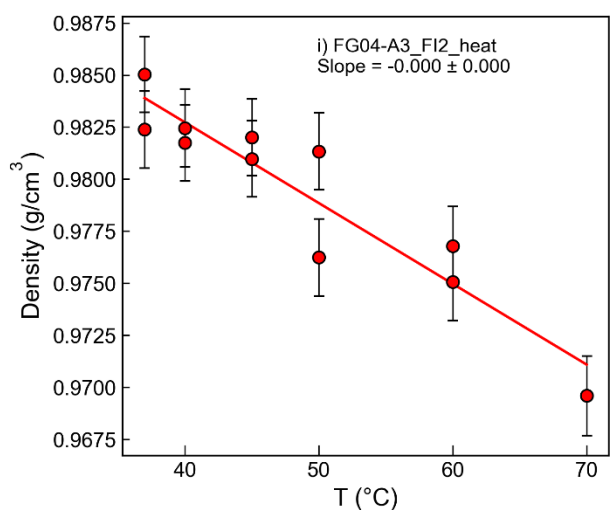
216 **Supporting Figure S23** – Power series for 10 FI/MI, 4 of which are shown in Fig. 14 of the main
 217 text. The gradient and error on the slope is shown.

218



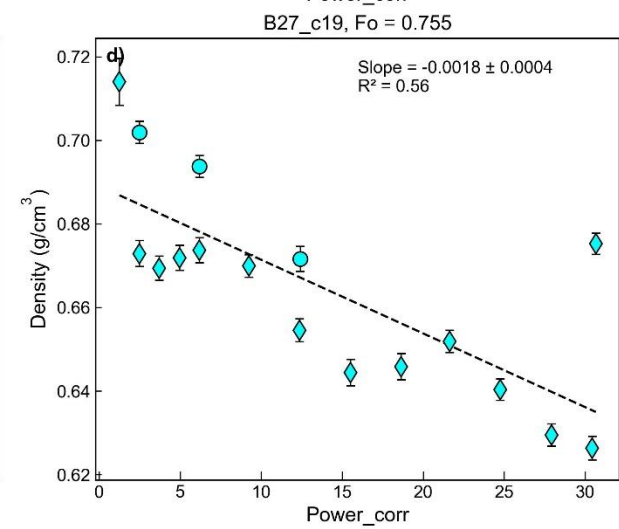
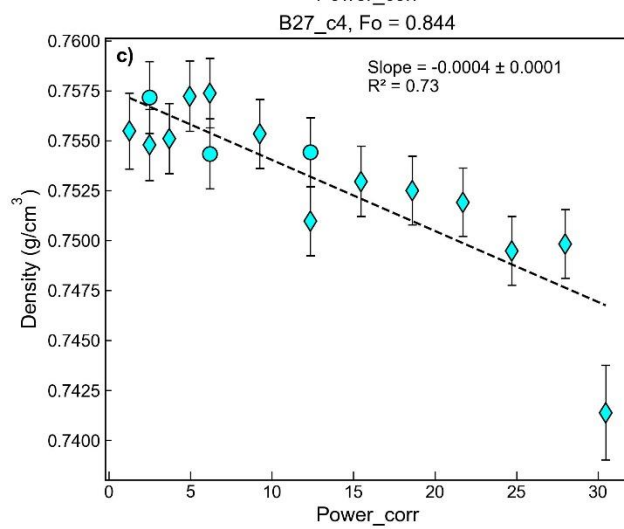
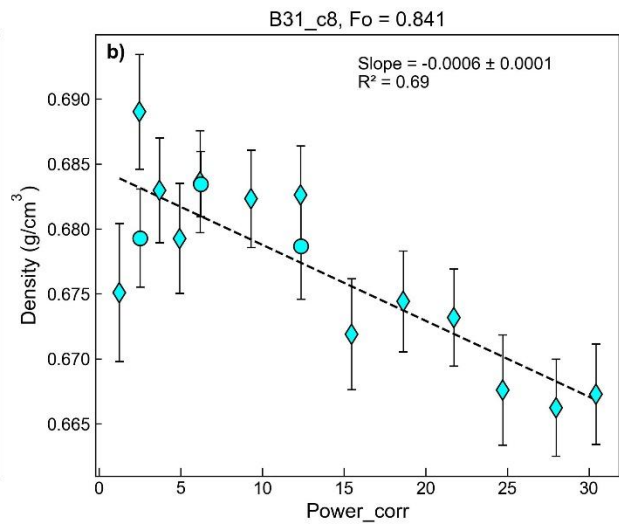
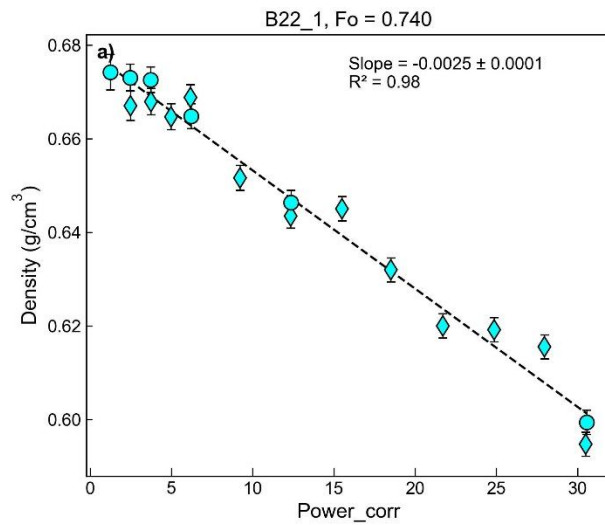


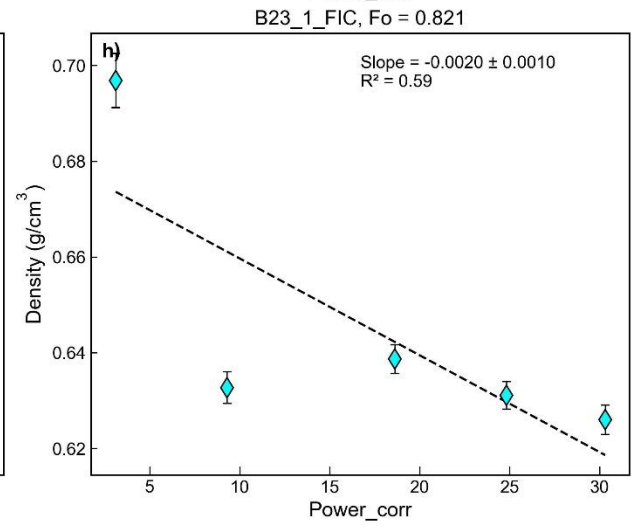
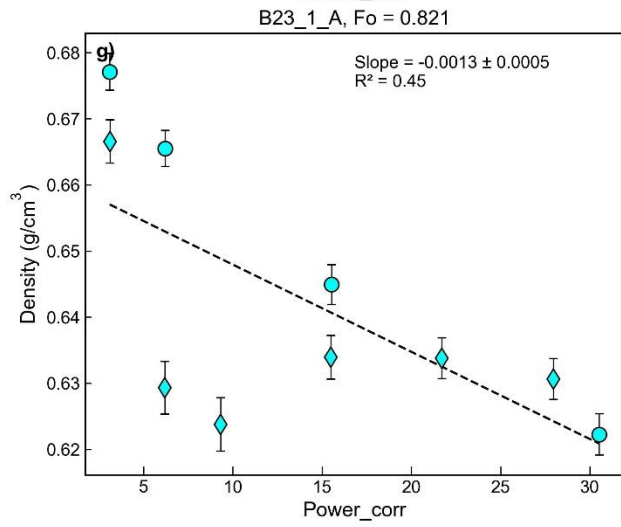
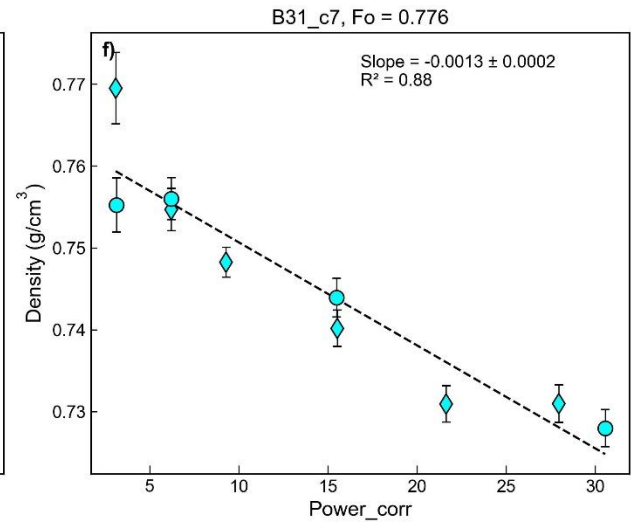
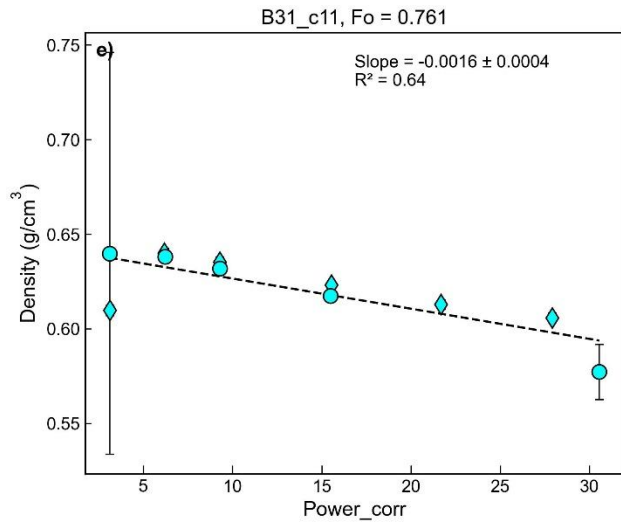
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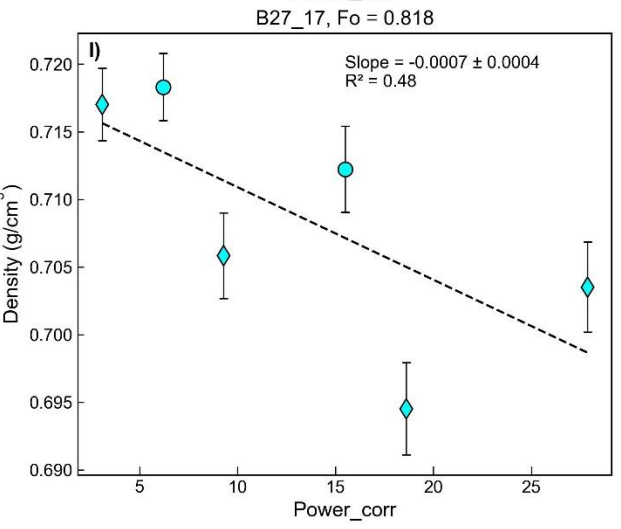
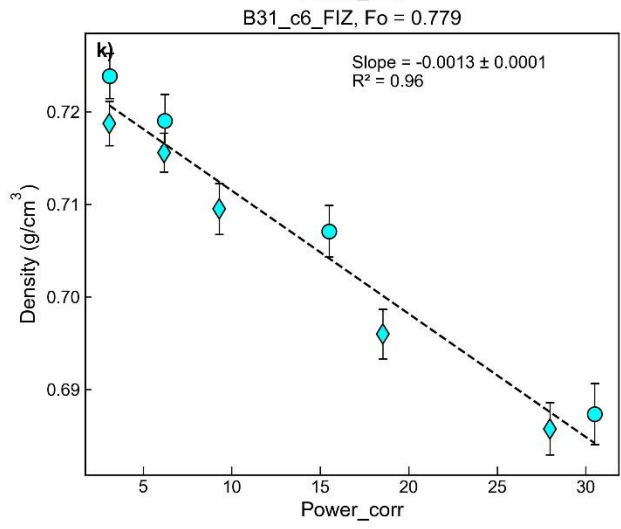
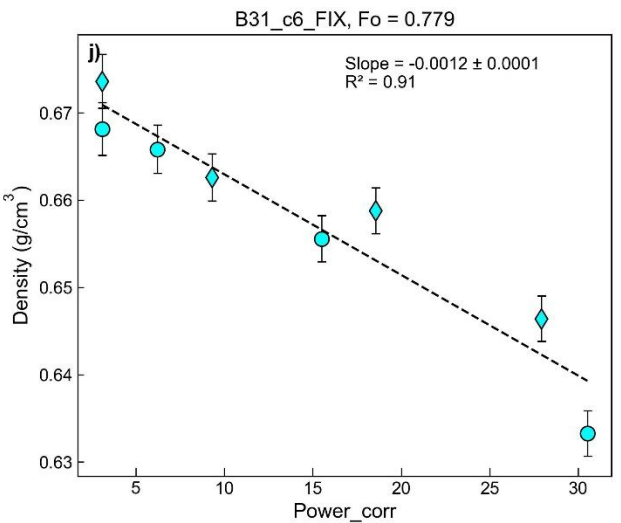
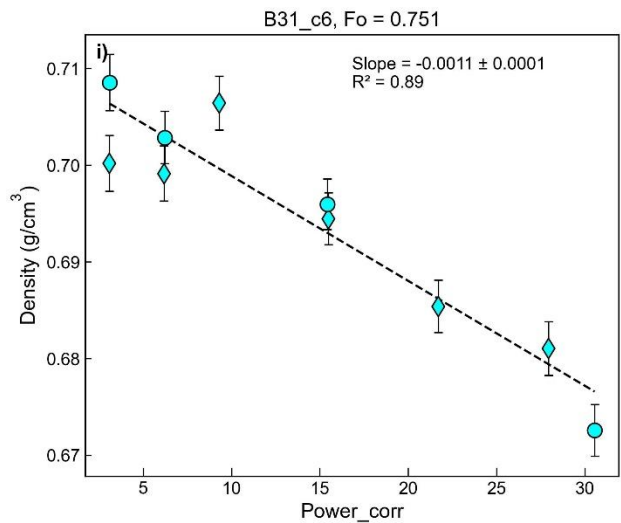


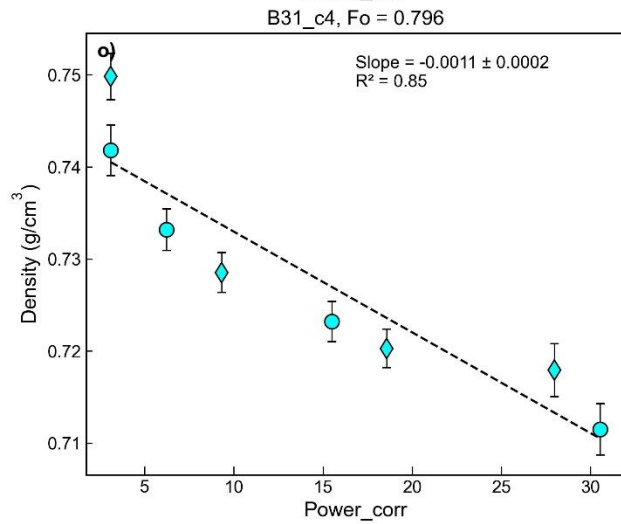
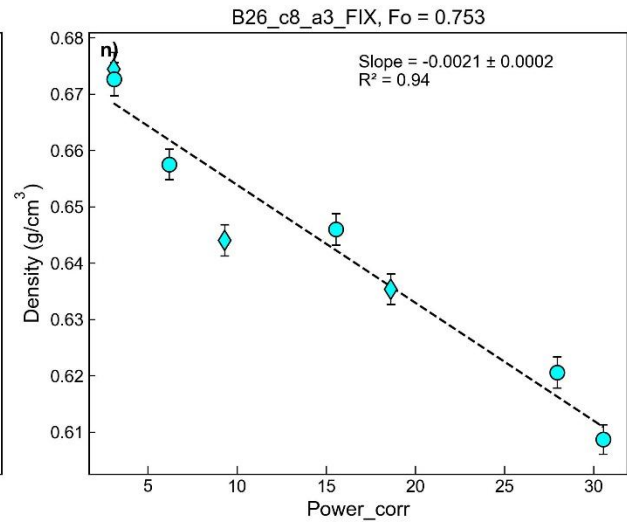
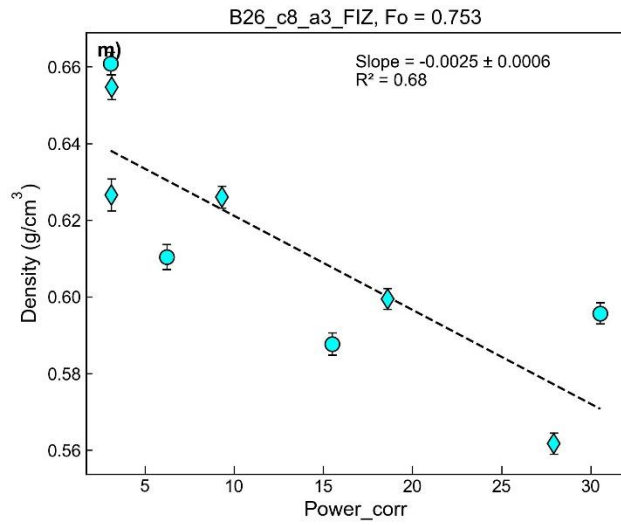
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222 **Supporting Figure S24** – Heating series for 10 FI/MI, 4 of which are shown in Fig. 14 of the main
 223 text. The gradient and error on the slope is shown.





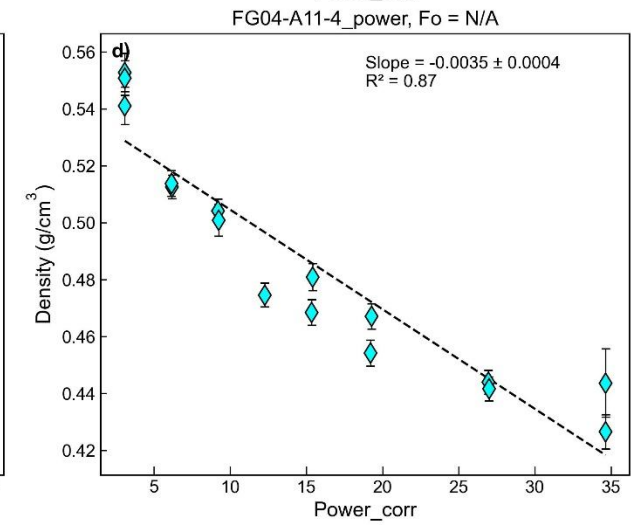
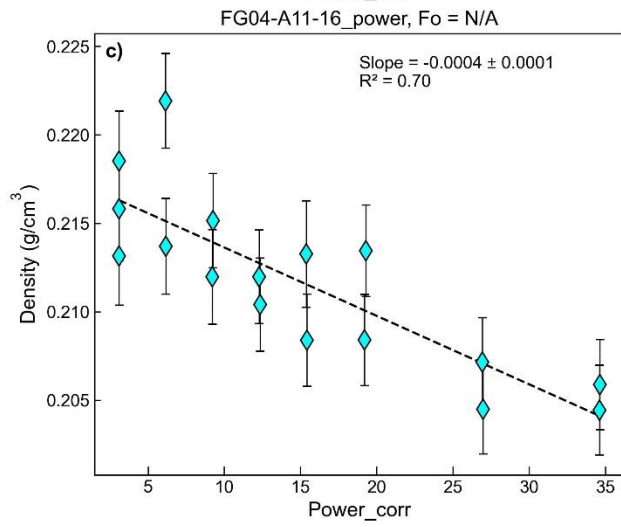
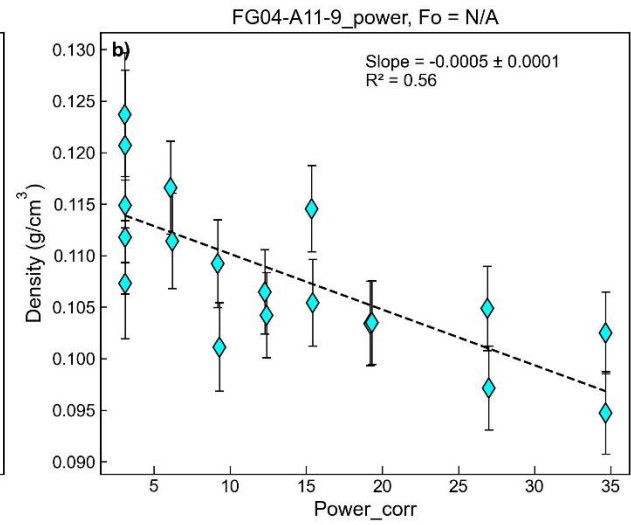
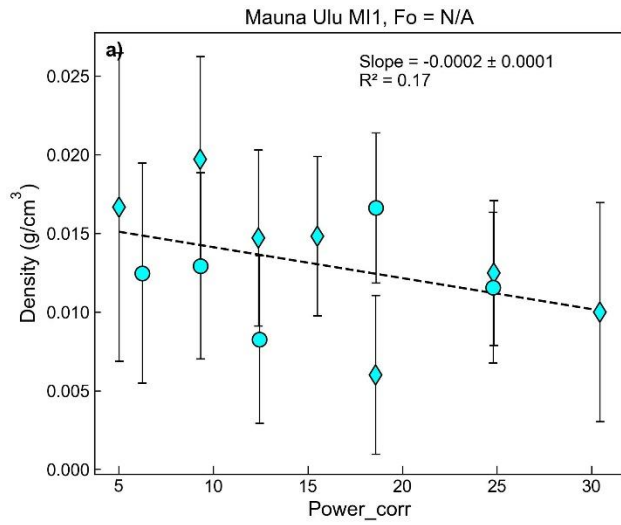




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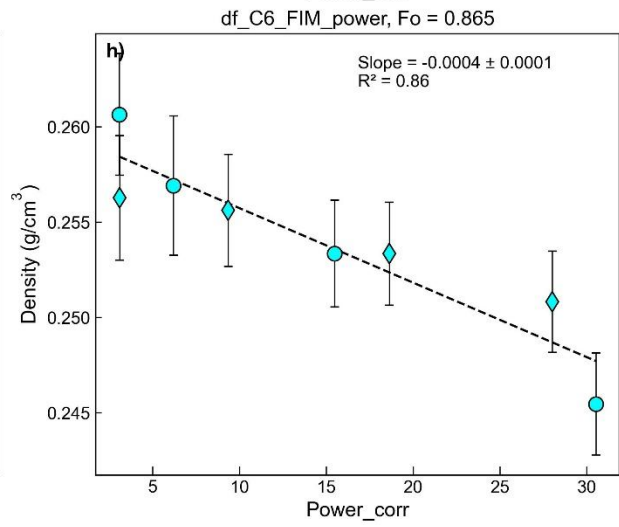
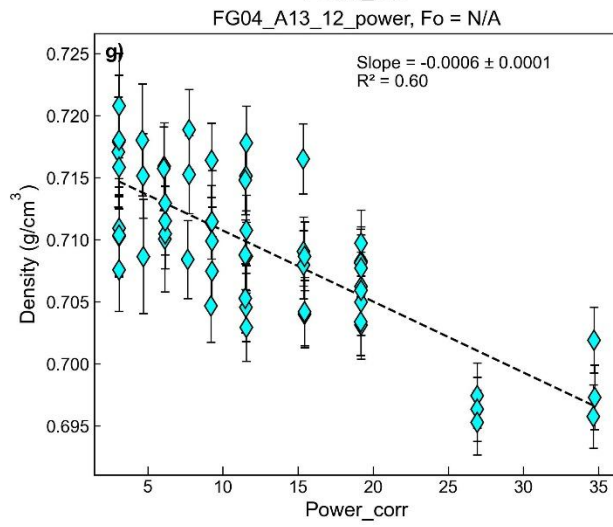
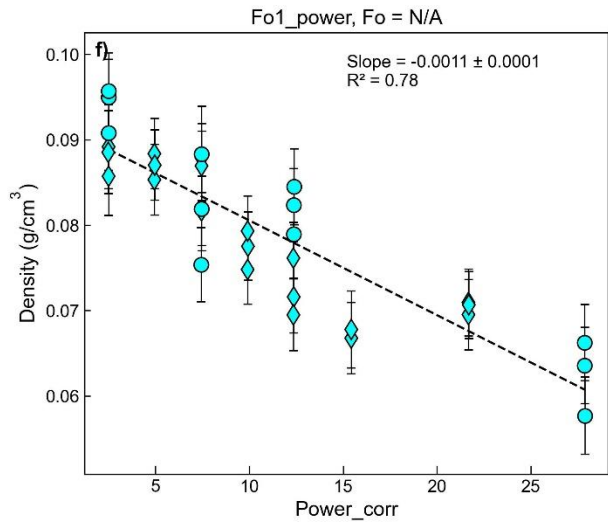
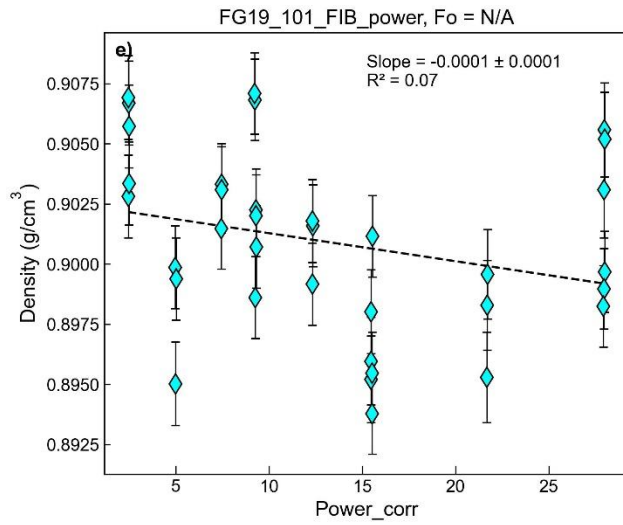
228 **Supporting Figure S25** – Power series for Icelandic fluid inclusions with known Fo contents. The
 229 gradient, Fo content, and R^2 value is shown.

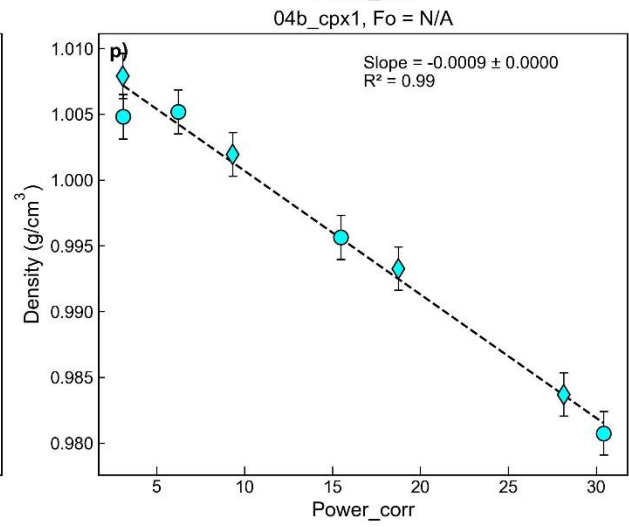
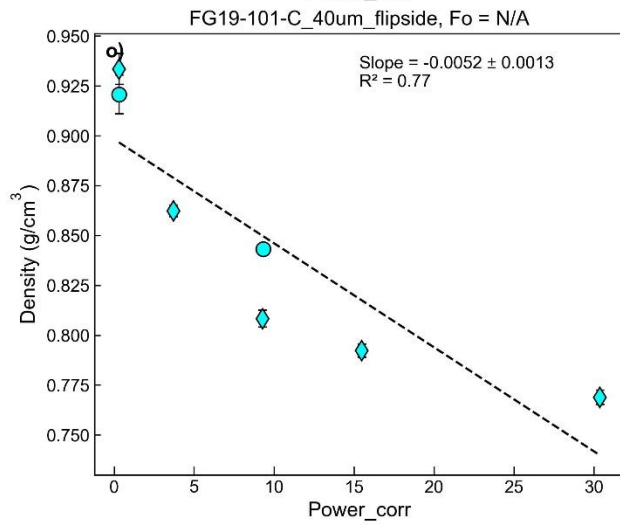
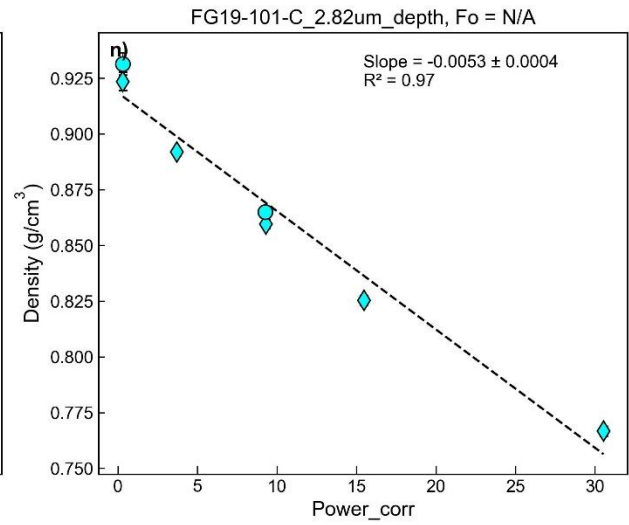
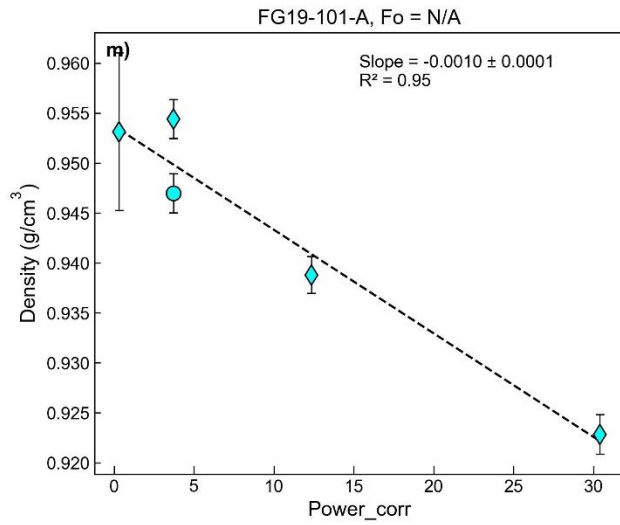
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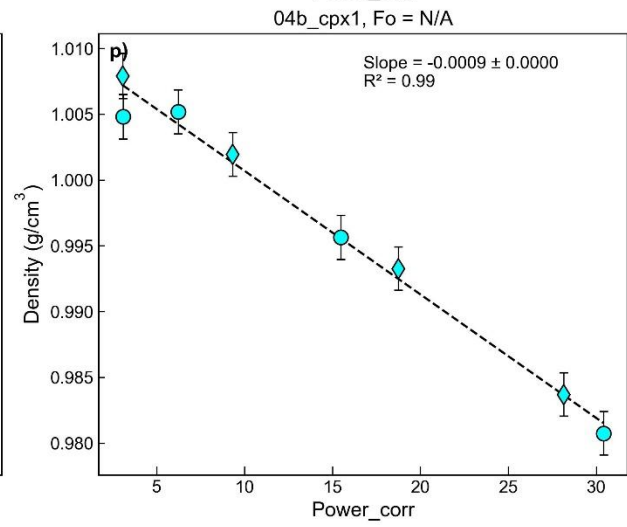
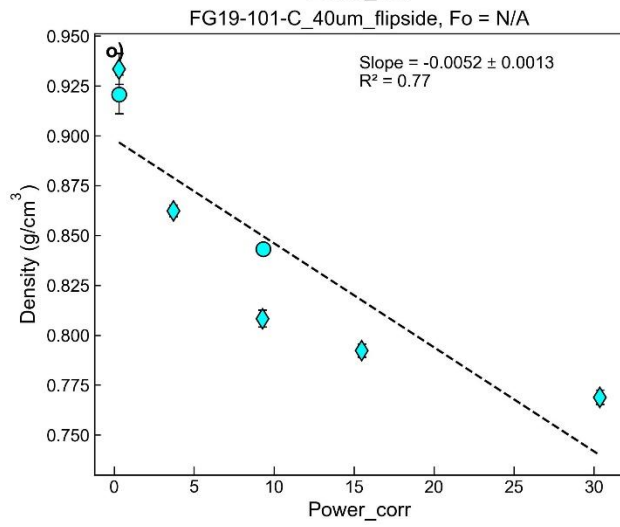
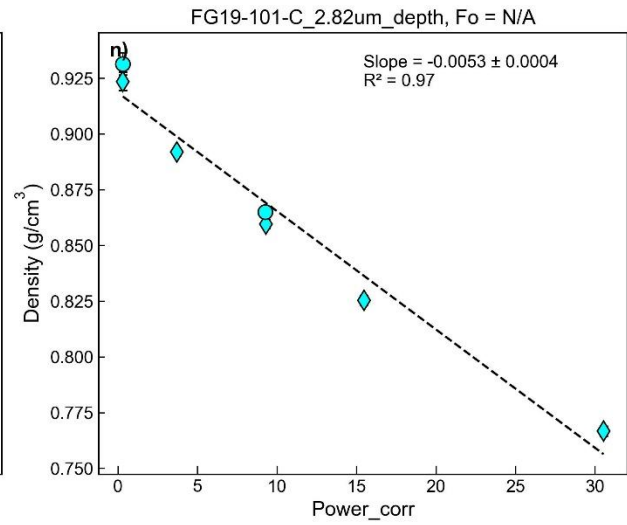
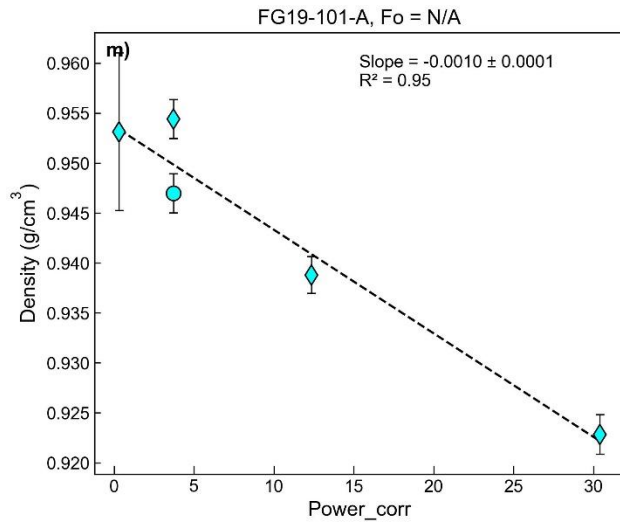
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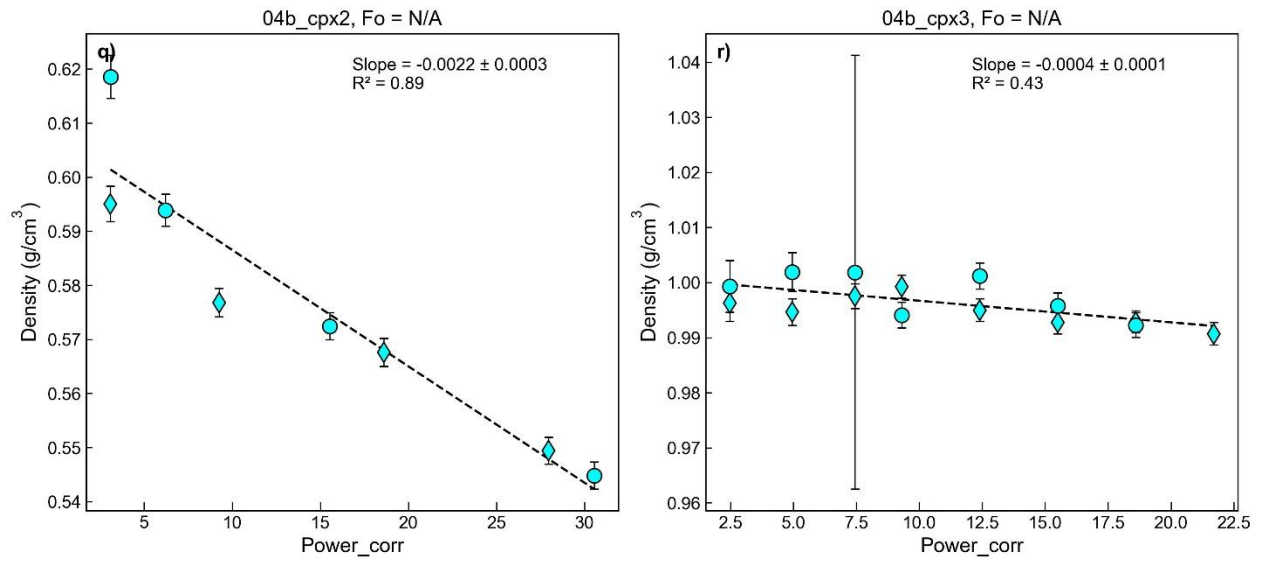


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241 **Supporting Figure S26** – Power series for all other fluid inclusions, and melt inclusions (a, f)
 242 shown in Fig. 12a of the main text.

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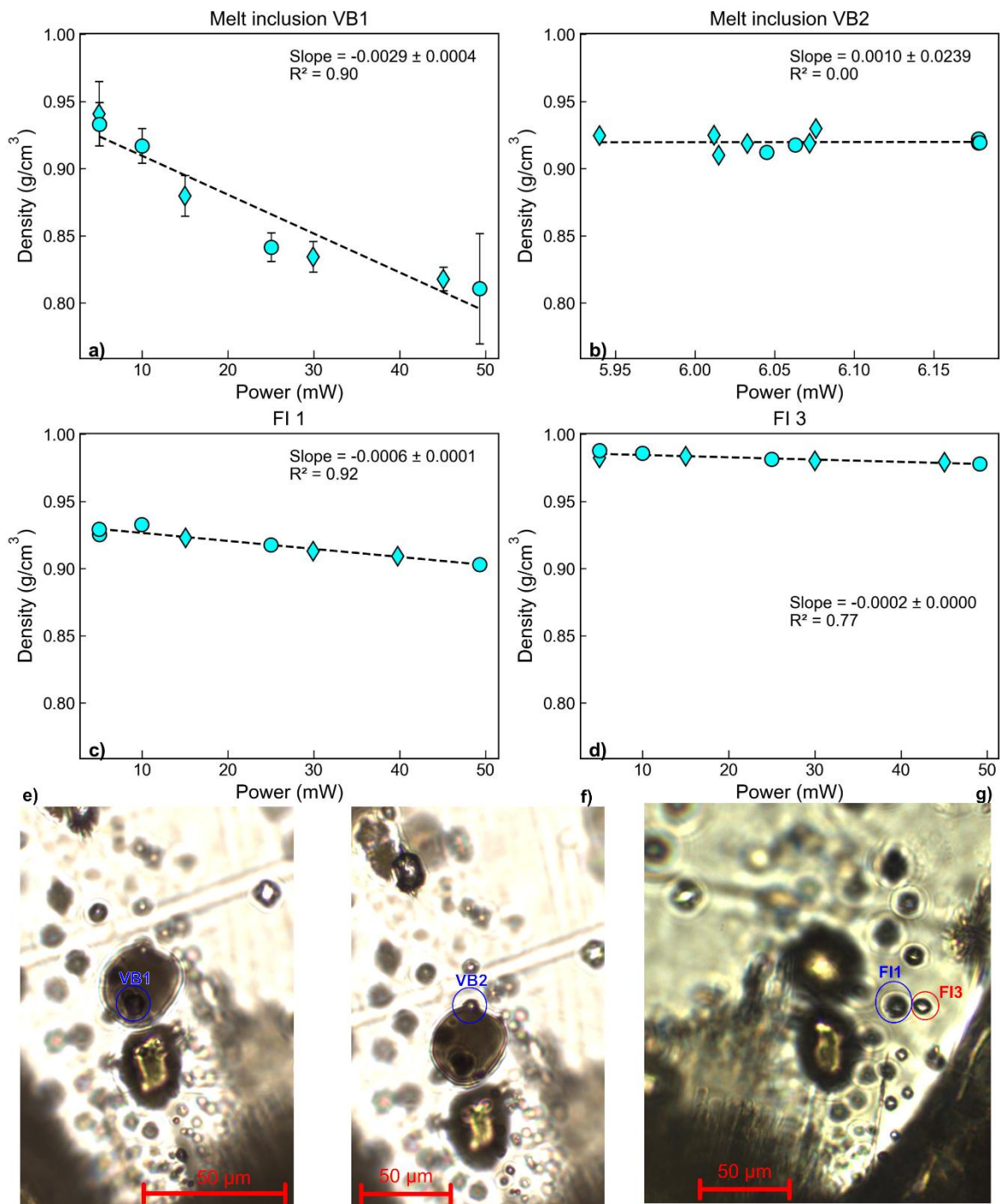
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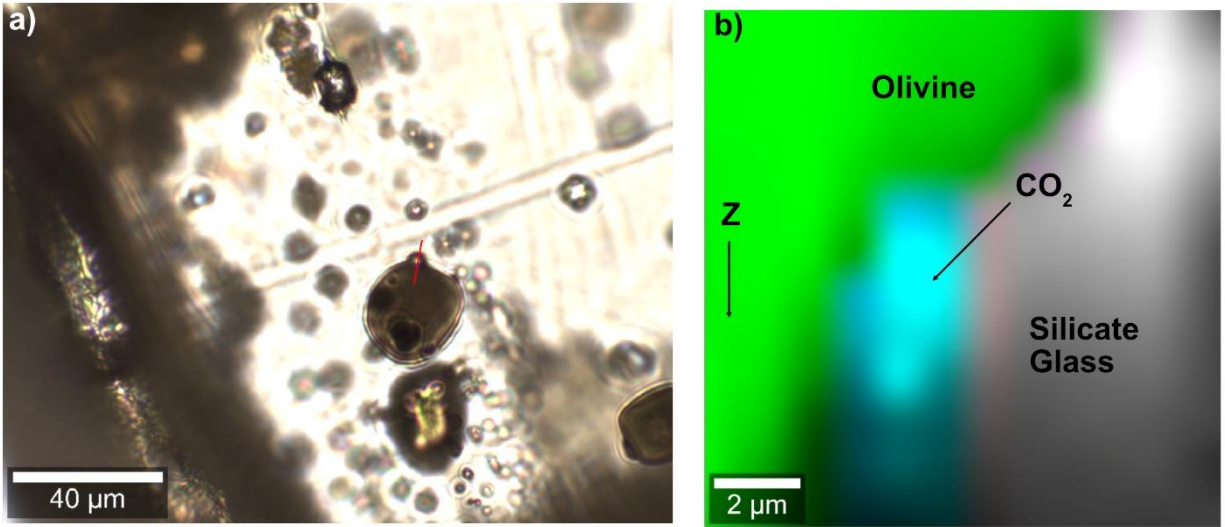
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250 **Supporting Figure S27** – Comparison of change in density as a function of laser power for two
 251 melt inclusion vapour bubbles and two fluid inclusions within the same crystal. VB1 is surrounded
 252 by melt (e) while VB2 is partially touching melt and partially touching olivine (f, see Supporting
 253 Fig. 11). The fact that VB2 is surrounded by olivine and melt accounts for its lower heating rate
 254 compared to VB1. It is noteworthy that the two FI, despite being in close proximity, show gradients
 255 that differ by a factor of 3.



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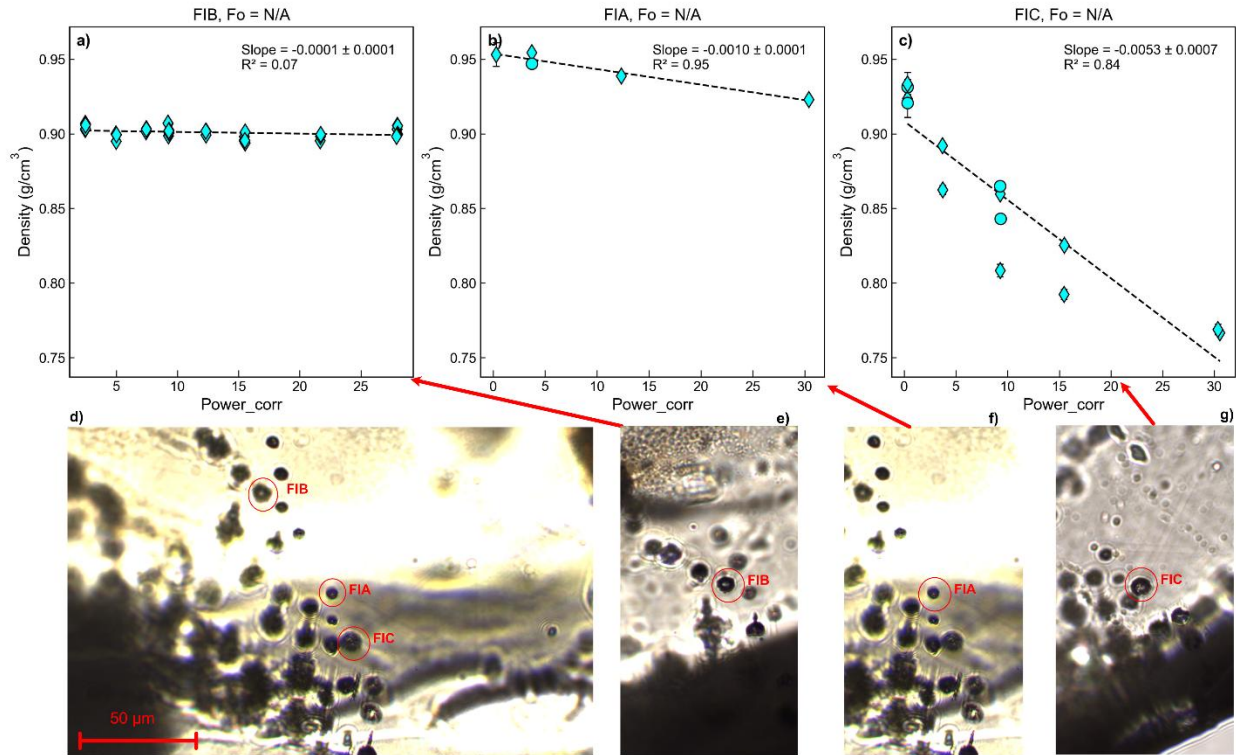
258 **Supporting FigureS28**– a) Transmitted light image of Bub2, with the red line showing the location
259 of a vertical Raman map. b) Processed Raman map, showing that the bubble is partially touching
260 the silicate glass of the melt inclusion (grey colors), partially touching the olivine (green). Map
261 processed using the WITec software.

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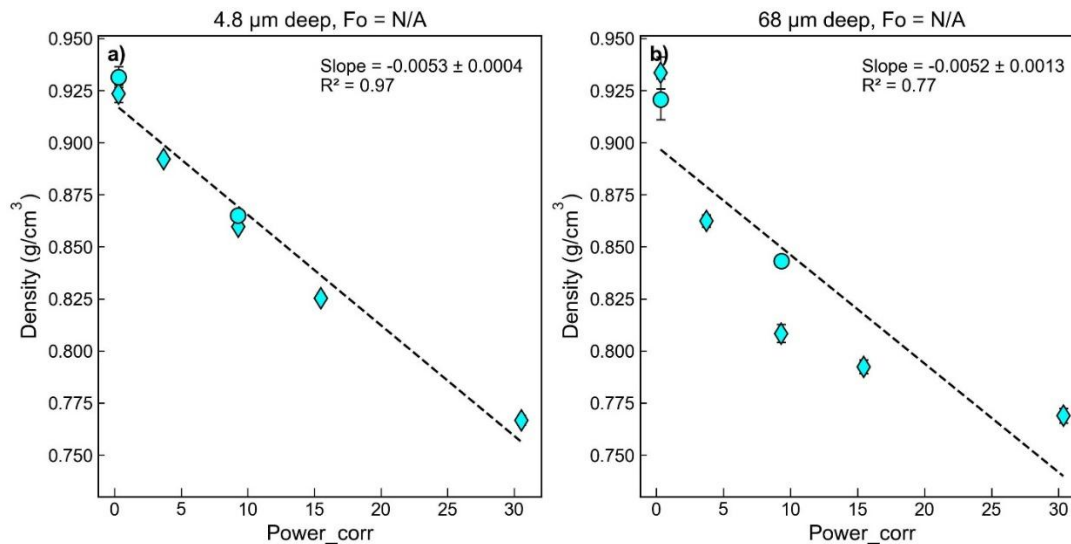
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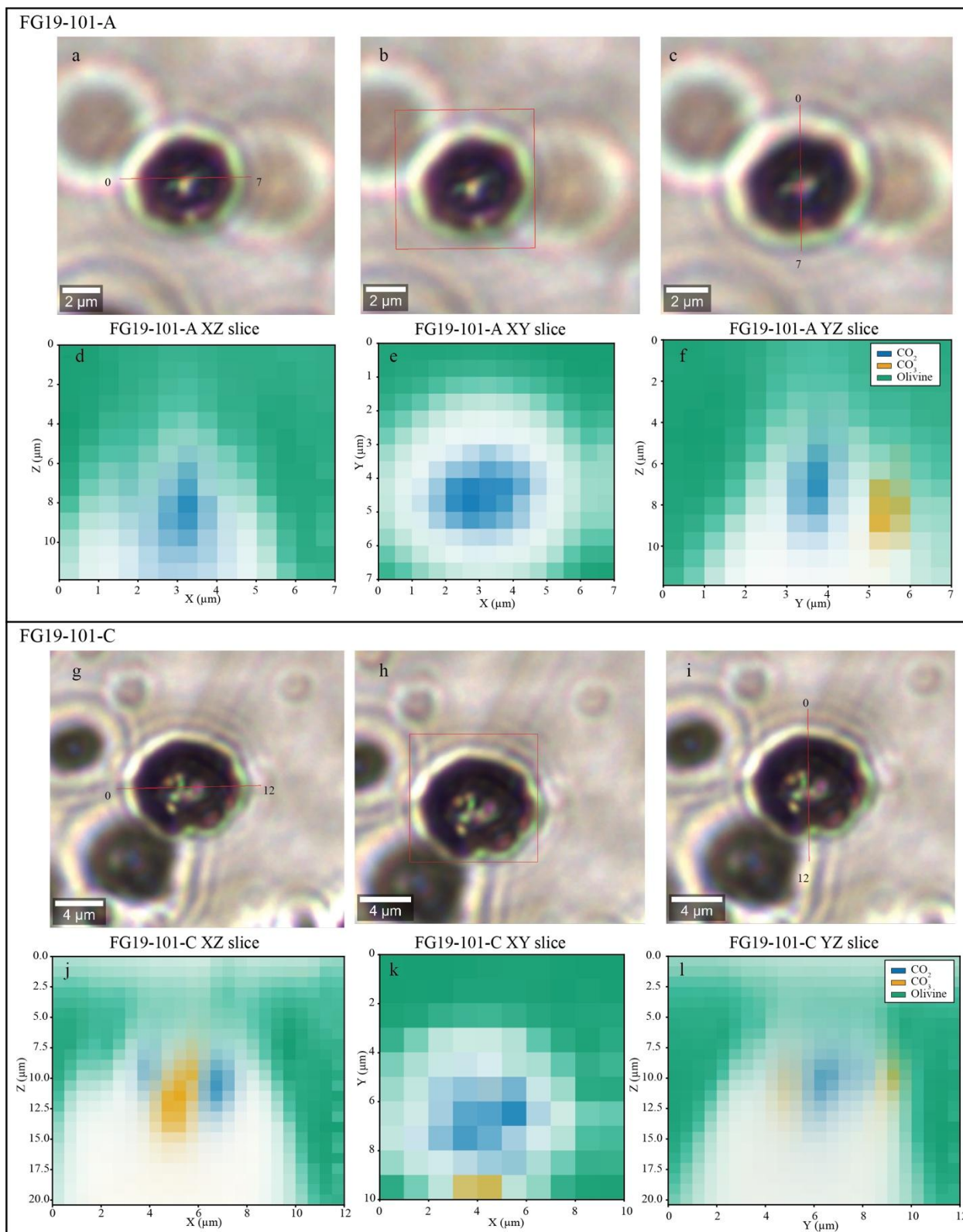
266

267 **Supporting Figure S29.** Comparison of change in density with increasing laser power (a-c) for 3
 268 FI hosted within a single olivine crystal (FG19_101). d) Transmitted light image of the fluid
 269 inclusion trail and the location of each FI. e-g) shows a focused picture of each FI. The scale is
 270 the same for all pictures. Raman maps of A and C are shown in Supporting Figure 13.



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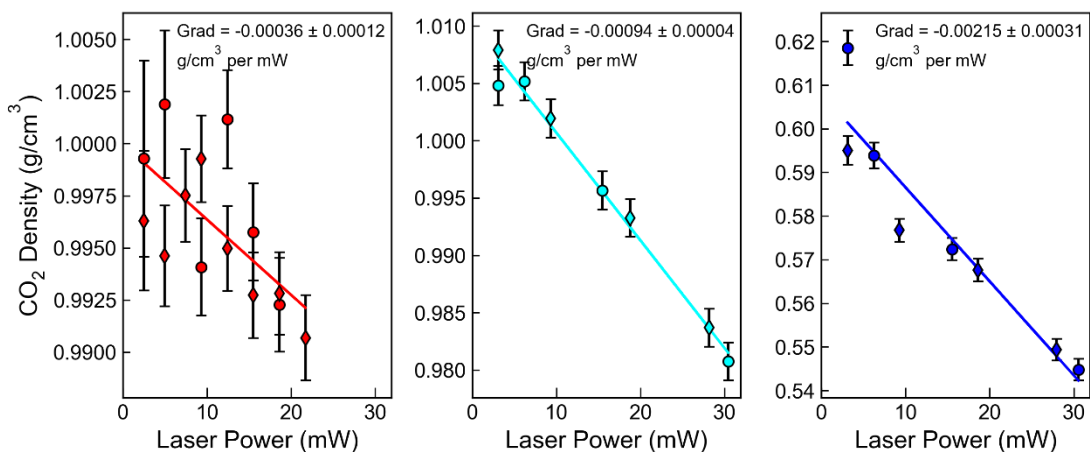
272 **Supporting Figure S30** – Comparison of heating rate for FI FG19-101C as a function of depth
 273 below the surface. a) shows the heating rate when the FI was ~4.8 μm below the surface. The
 274 crystal was then flipped over, and then a power series was conducted from the other side, such
 275 that the FI was ~68 μm below the surface. The gradients of these series is within error of each
 276 other. This supports the modelling work of Hagiwara et al. (2021) suggesting that depth in the
 277 sample doesn't strongly affect heating rate.



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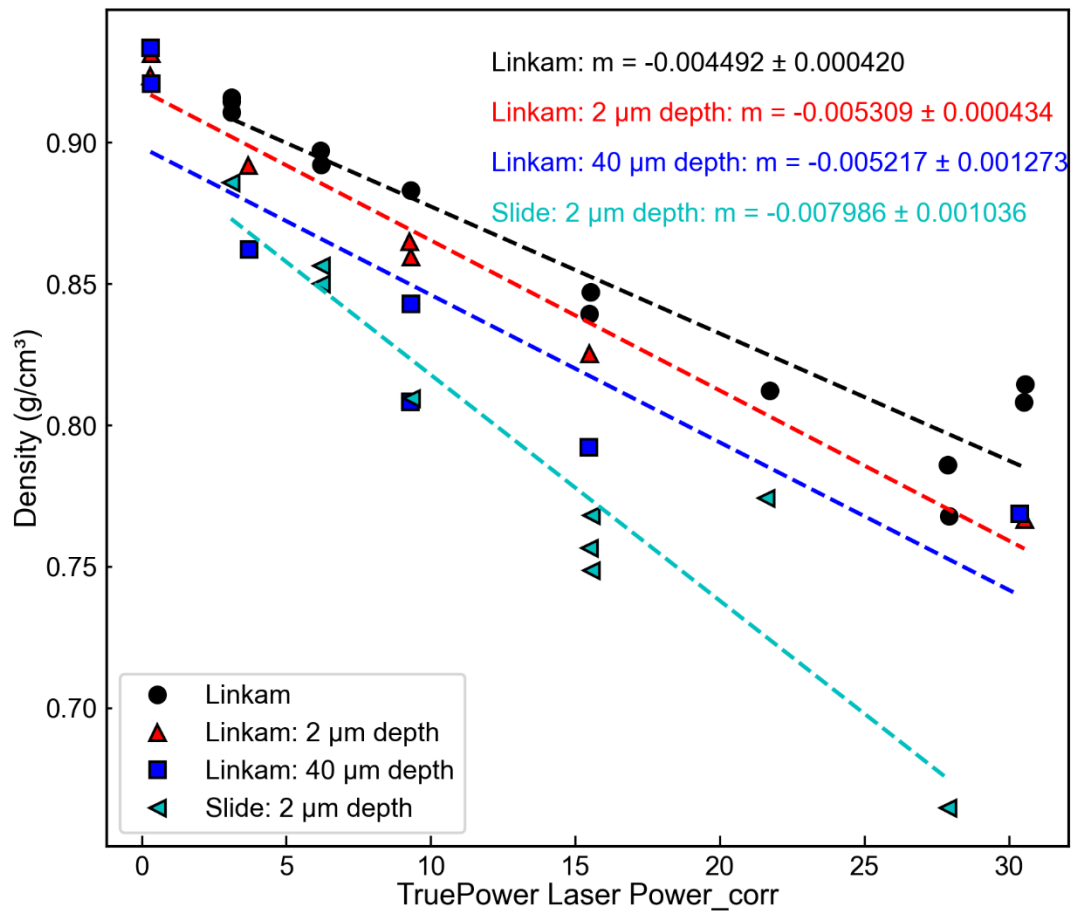
279 **Supporting Fig. S31.** Two dimensional Raman maps of FI FG19-101-A showing the presence
 280 of CO_2 and carbonates in the FI. a-c) Transmitted light images of FG19-101-A. The red lines and
 281 rectangles show the plane along which Raman mapping was conducted (shown in panels d-f). g-

282 i) Transmitted light images of FG19-101-A and corresponding 2D slice (red line). Both sets of
 283 Raman maps were collected using 50X objective, with a laser power of 49 mW, a integration time
 284 of 0.5 s, a 300 grooves/mm grating, and a 2350 cm⁻¹ spectral center. Z dimension is recalculated
 285 accounting for the refraction index of olivine (1.7). In the Z slices, the bottom under the inclusions
 286 is white (very low intensity) due to loss of confocality at depth. Other white pixels are areas of
 287 highly mixed spectra which are difficult to reliably extract components from. Maps were processed
 288 using Ramanspy 0.2.10 (Georgiev et al., 2024). To produce the maps, we first denoise using
 289 savitzky golay filter with a window length of 4, and a polynomial of 3rd order, then apply ASLS
 290 baseline routine in Ramanspy and normalize each spectrum to using Min-Max normalization. We
 291 apply the whole preprocessing routine to all the maps. We extract “pure” endmember spectra for
 292 CO₂, CO₃ and Olivine from the map spectra and use Non-negative Constrained Least Squares
 293 (NNLS) abundance method in Ramanspy to extract abundance maps for each component at
 294 every pixel. These maps are not quantitative, but an approximation of the distribution of each
 295 component. It is evident that carbonate is much more abundant in 101-C than 101-A, though they
 296 are only ~20 μm apart from each other (Supporting Fig. 11).



297
 298 **Supporting Figure S32** – Laser power series for ~Mg₇₉# clinopyroxenes for fluid inclusions from
 299 Gleeson et al. (2025). The gradient varies by a factor of ~3.8, despite the fact these samples are
 300 very chemically homogenous and there are no obvious color differences between the different
 301 crystals.

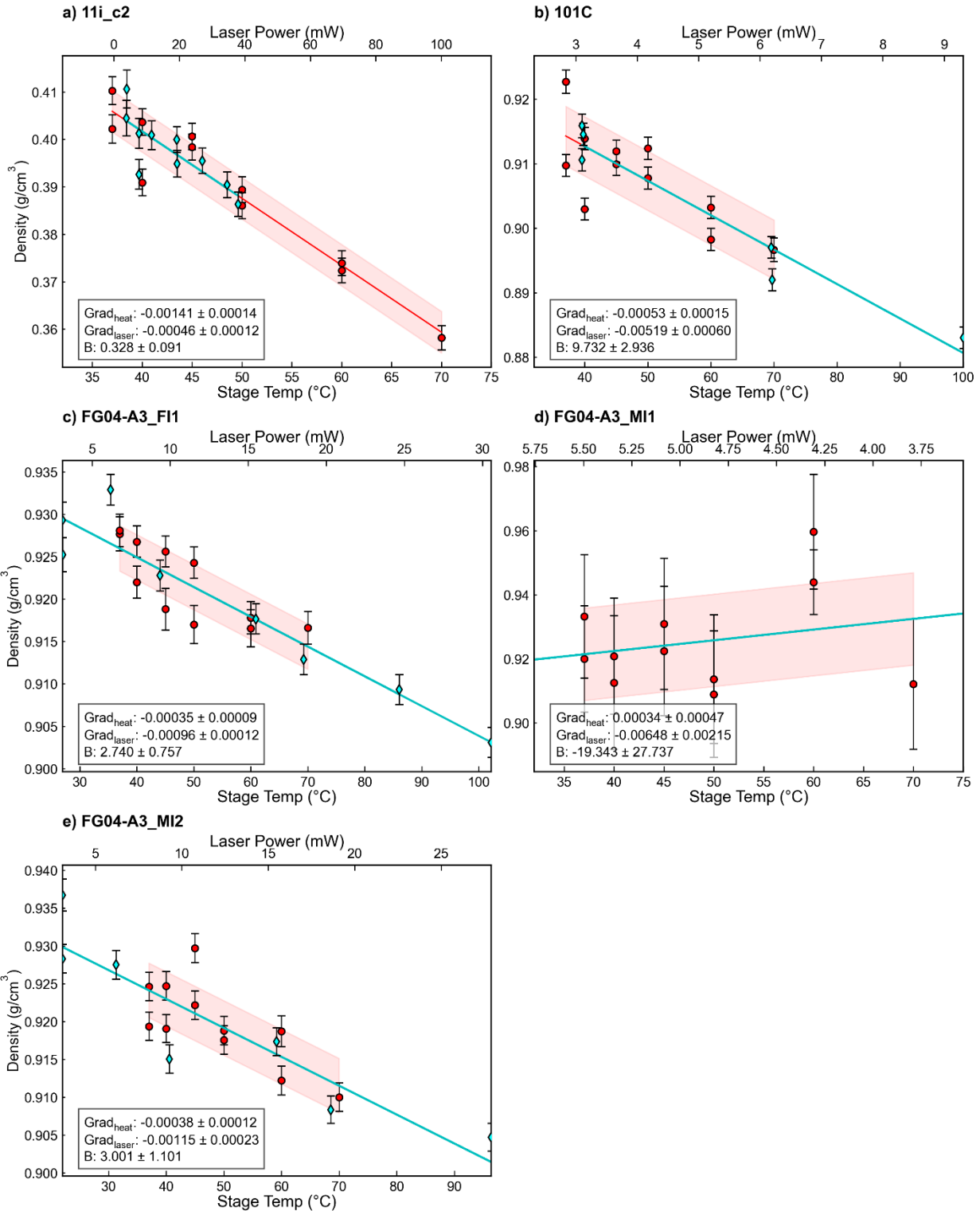
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305 **Supporting Fig. S33** – Comparing heating rates for FG19_101C when the crystal is mounted in
 306 the linkam stage (not surrounded by resin) at varying depths (black, red and blue), vs. when the
 307 crystal is mounted on a glass slide in crystalbond resin (cyan colors).

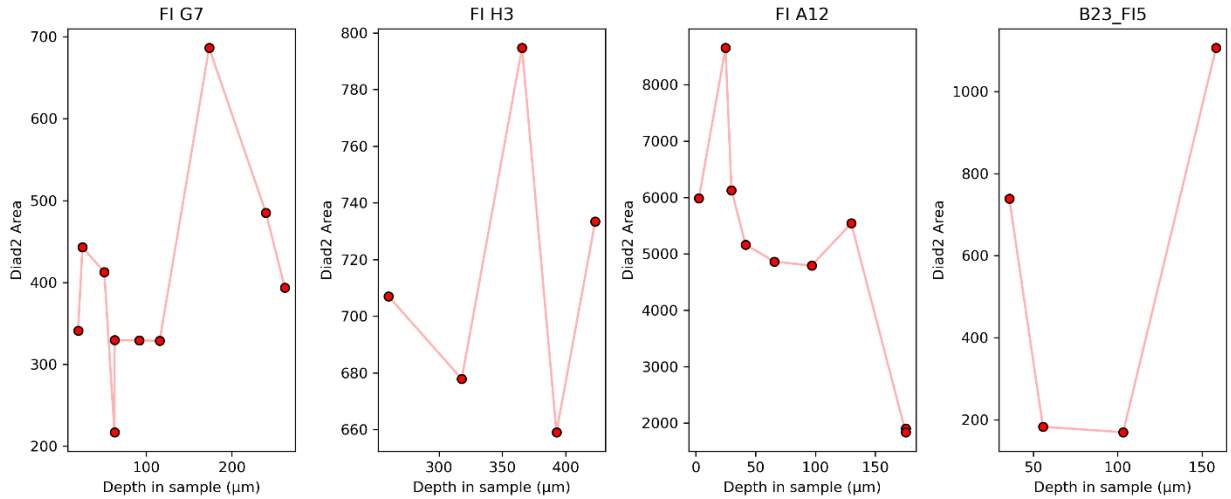
308



309

310 **Supporting Figure S34** – Comparison of heating and power series as in Fig. 14 of the main text.

311

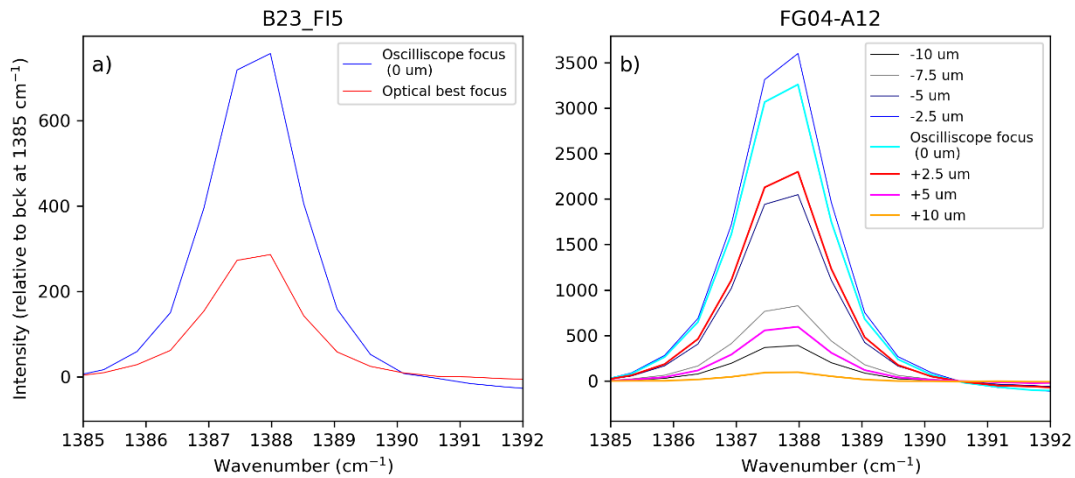


312

313 **Supporting Figure S35**– Changes in signal intensity as a function of depth in sample for 4
 314 different olivine-hosted FI.

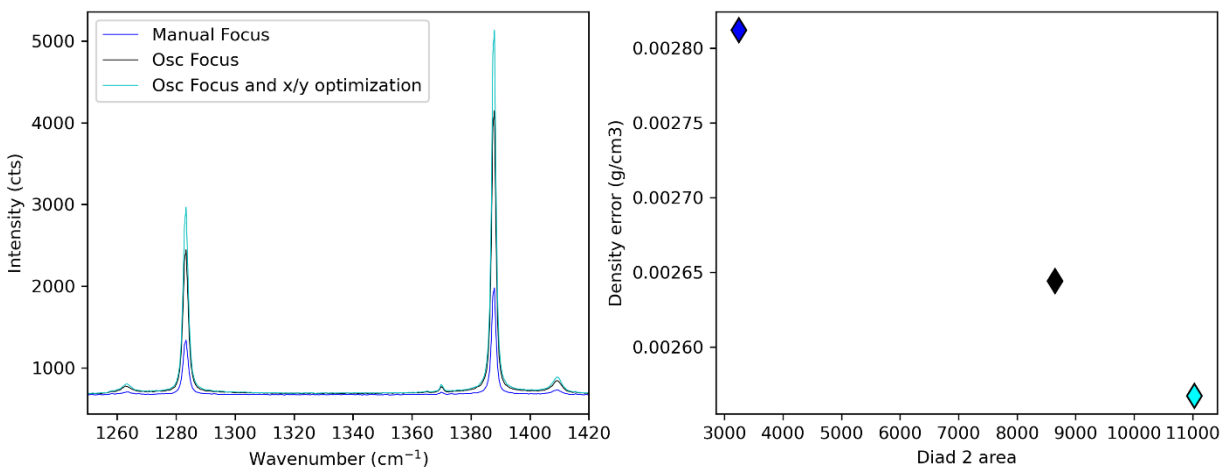
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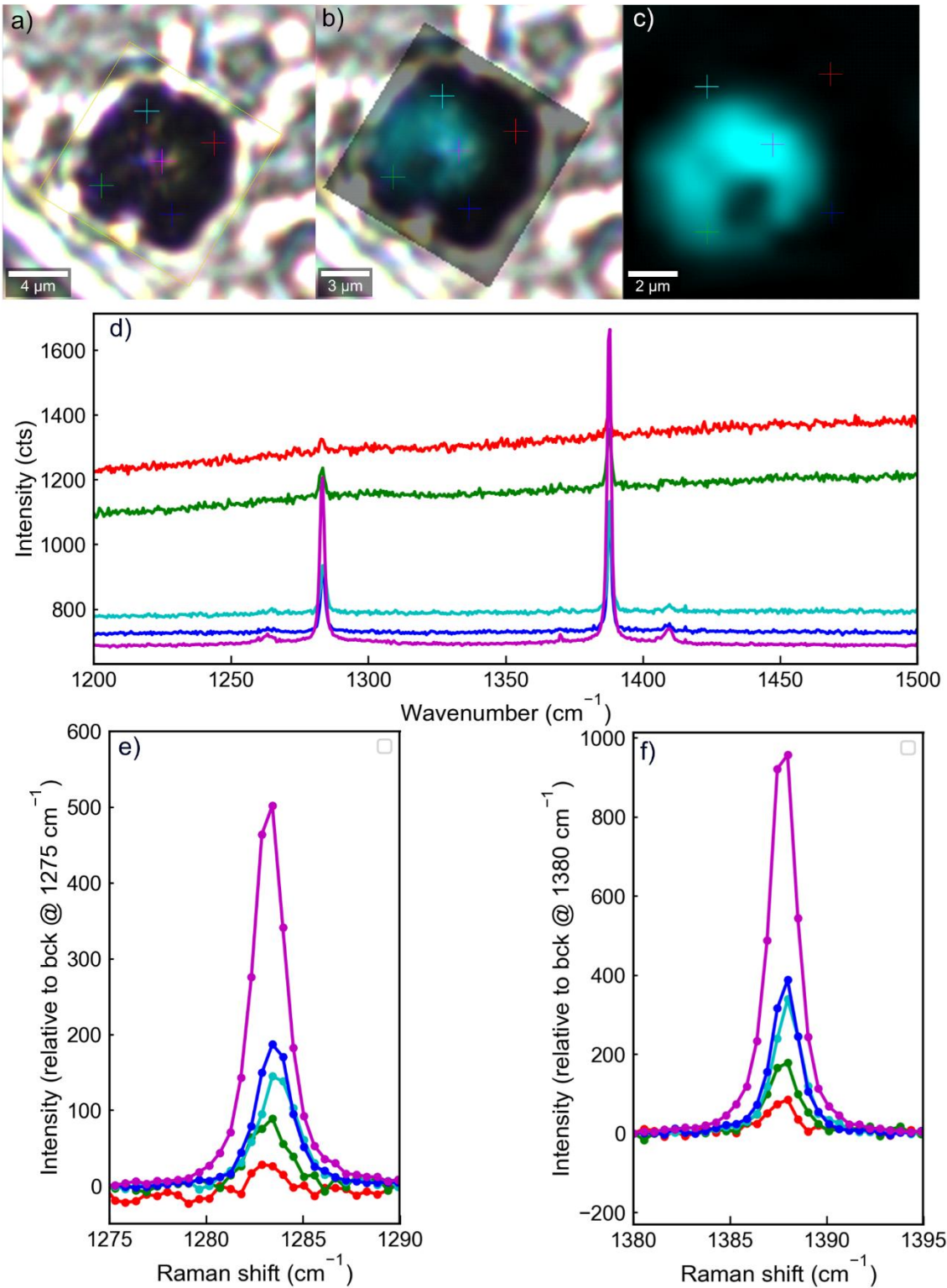
317

318 **Supporting Figure S36** – Example of the effect of small focus changes on the intensity of diad
 319 2. a) Compares the user-selected optical focus based on sharpness of the FI edges to the optimal
 320 focus position found by the oscilloscope. b) Compares the optimal focus position found by the
 321 oscilloscope to small perturbations from this position.



322

323 **Supporting Fig. S37**– Compares the user-selected optical focus based on sharpness of the FI
 324 edges to the optimal focus position found by the oscilloscope (dark blue), to the optimum Z
 325 position found by the oscilloscope (black), and the optimum Z and x-y position determined by
 326 Raman mapping for FG04-A12.

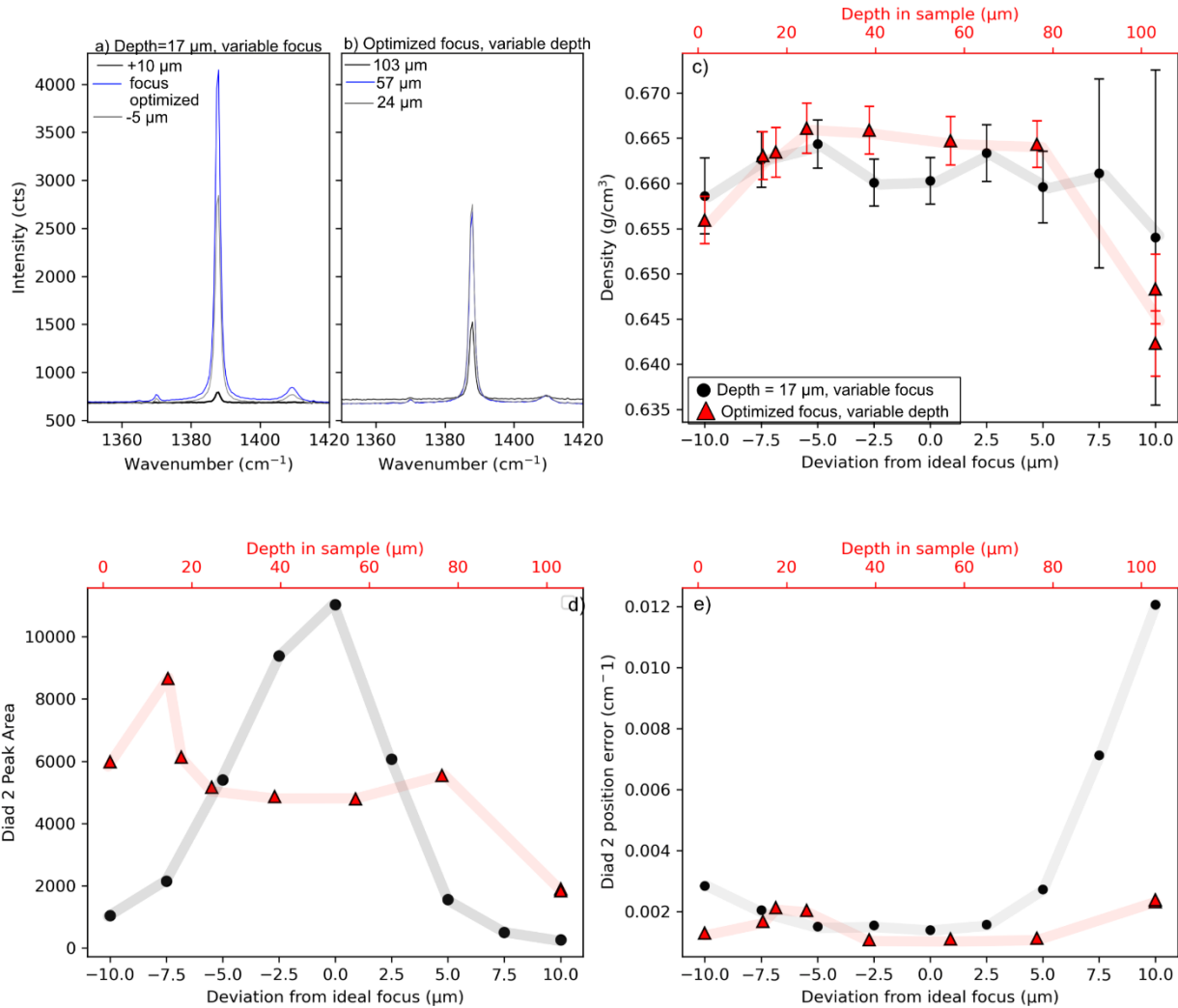


327

328 **Supporting Figure S38** – Variation in signal intensity as a function of X-Y position within the fluid
 329 inclusion. a) Transmitted light image of fluid inclusion B23_FI5. The 5 colored crosses show the
 330 position of different spectral acquisitions shown in d-f. b) An 1800 g grating Raman map with the

331 intensity of the CO₂ signal shown as cyan colors overlain on this transmitted light image. c) shows
 332 the map with no underlay. d-f) Comparing the intensity of different acquisitions. While it is not
 333 surprising that the strongest signal is obtained from the center of the inclusion, it is surprising
 334 that despite the cyan, red, green and blue cross being equidistance from the center, they yield
 335 vastly different intensities.

336

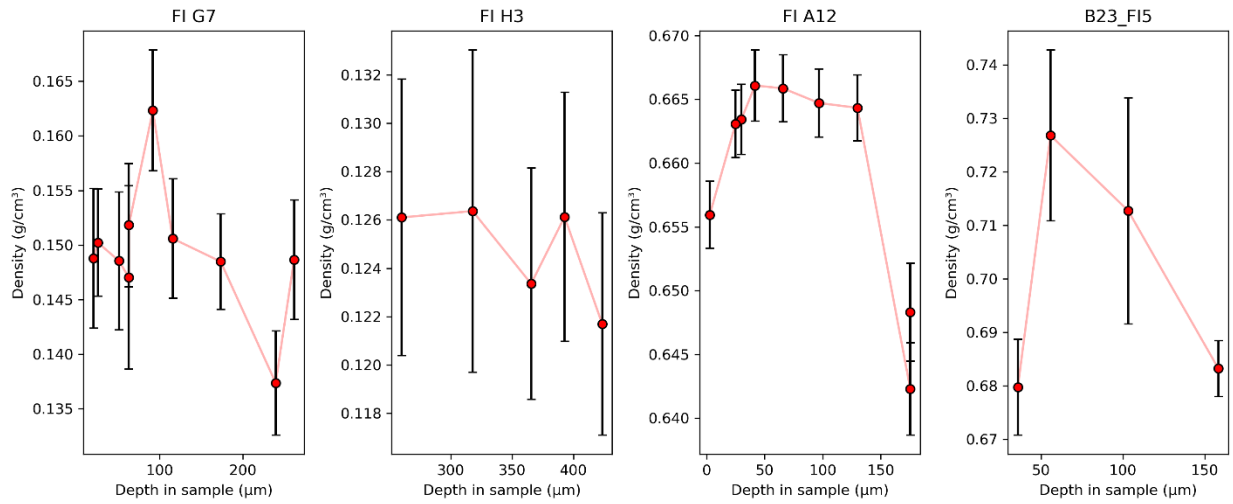


337 **Supporting Figure S39** – Comparison of peak intensity and peak fitting errors with increasing
 338 depth in sample vs. a change in focus for FG04_A12. a) Spectra centered around diad2 showing
 339 the difference in intensity for the optimized focus (blue) vs. -5 μm deeper in the sample (grey) and
 340 10 μm shallower (black). The fluid inclusion has a radius of ~7 μm. b) Change in intensity with
 341 depth of the fluid inclusion optimal focus below the sample. At 24 μm and 57 μm depth, the
 342 intensities are very similar, although they drop to above half for 103 μm depth. c) Change in
 343 density as a function of depth and ideal focus. Red triangles show variation in depth in sample
 344 (top red axis) and black circles show deviation from the ideal focus (black). Apart from the very
 345 worst focus and deepest depth, there is no change outside of uncertainty. Points are joined up
 346

347 with lines to help separate the two series. d) Peak area of diad 2 (a measure of intensity) as a
348 function of depth and focus. e) Error on the position of diad 2.

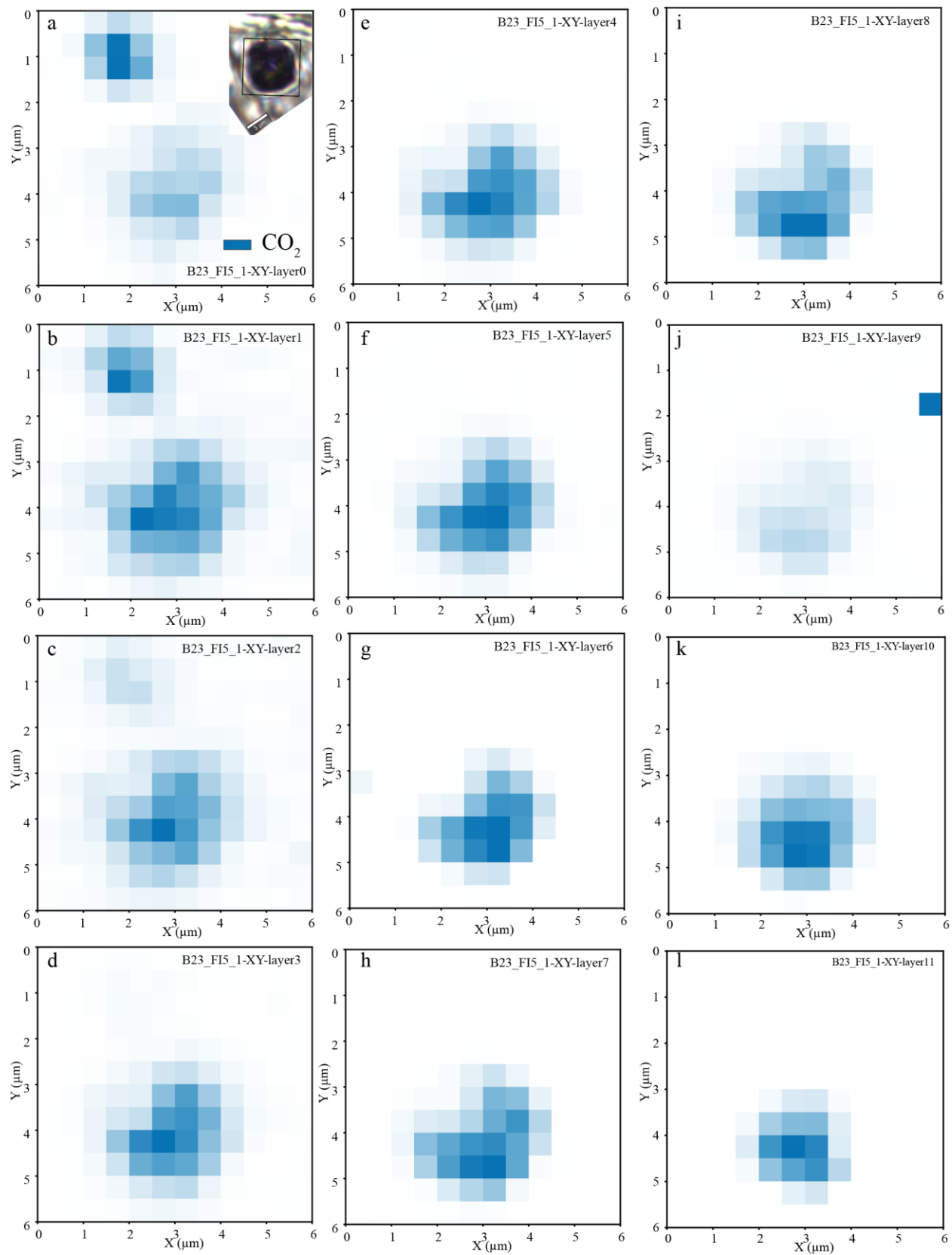
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352 **Supporting Figure S40-** Density (and error) as a function of depth. The highly variable densities
353 in B23_FI5 likely result from the very poor surface polish, resulting in backgrounds which were
354 tricky to fit.

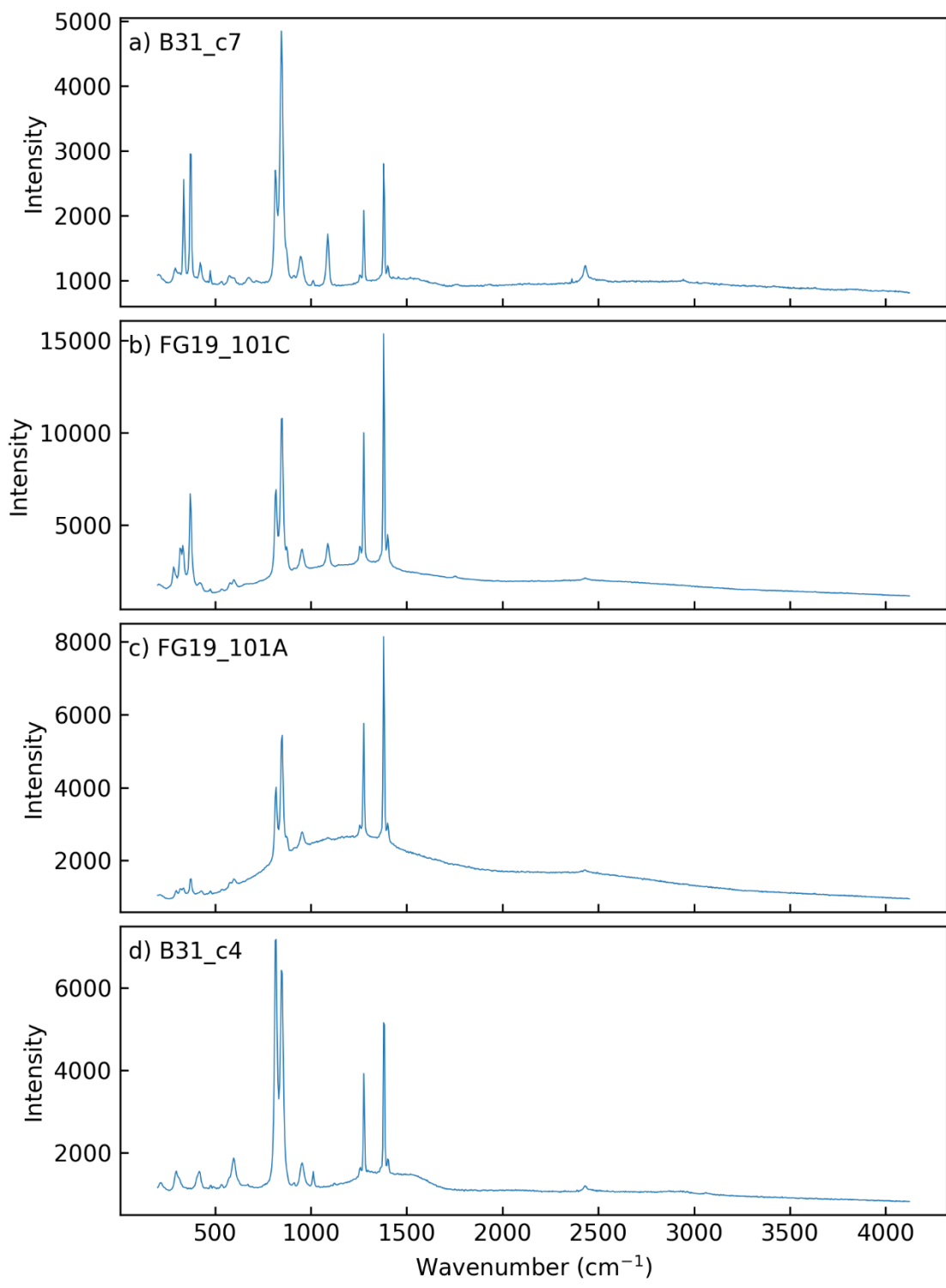


356 **Supporting Figure S41** – Raman map of CO₂ abundance in B23_F15_1. The inset in a) shows
357 the geometry of the map relative to the fluid inclusion. Each X-Y graph shows a different Z slice
358 through the inclusion, with the stronger blue colors showing stronger CO₂ intensities. It is clear
359 that the x-y distribution of max intensity varies with depth. The map geometry in the WITec
360 software was set as 12x12x12 um, so each slice was separated by 1 um of stage movement.
361 With the refractive index of olivine being 1.7 and CO₂ being 1, there is some vertical distortion.
362 Maps were collected using 50X objective, with a laser power of 49 mW, an integration time of 1
363 s, the 1800 grooves/mm grating (so olivine is not visible), and a 1325 cm⁻¹ spectral center.
364 Processing was performed using Ramanspy 0.2.10 (Georgiev et al., 2024). To produce the maps,
365 we first denoise using savitzky golay filter with a window length of 4, and a polynomial of 3rd order,
366 then apply ASLS baseline routine and normalize each spectrum to using Min-Max normalization.
367 We apply the whole preprocessing routine to all the maps. We extract “pure” endmember spectra
368 for CO₂ from the map spectra and use Non-negative Constrained Least Squares (NNLS)
369 abundance method in Ramanspy to extract abundance maps for CO₂ at every pixel. These maps
370 are not quantitative, but an approximation of the distribution of CO₂.

371

372

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374

375 **Supporting Fig. S42**– 300 g wavelength scans of 4 inclusions targeted for laser heating tests.

376 There is not significant amounts of gas species other than CO₂.

377

378 **References**

- 379 Allison, C.M., Roggensack, K., Clarke, A.B., 2021. Highly explosive basaltic eruptions driven by
380 CO₂ exsolution. *Nat Commun* 12, 217. <https://doi.org/10.1038/s41467-020-20354-2>
- 381 Dayton, K., Gazel, E., Wieser, P., Troll, V.R., Carracedo, J.C., La Madrid, H., Roman, D.C., Ward,
382 J., Aulinas, M., Geiger, H., Deegan, F.M., Gisbert, G., Perez-Torrado, F.J., 2023. Deep
383 magma storage during the 2021 La Palma eruption. *Sci. Adv.* 9, eade7641.
384 <https://doi.org/10.1126/sciadv.ade7641>
- 385 DeVitre, C.L., Allison, C.M., Gazel, E., 2021. A high-precision CO₂ densimeter for Raman
386 spectroscopy using a Fluid Density Calibration Apparatus. *Chemical Geology* 584,
387 120522. <https://doi.org/10.1016/j.chemgeo.2021.120522>
- 388 DeVitre, C.L., Wieser, P.E., 2024. Reliability of Raman analyses of CO₂-rich fluid inclusions as a
389 geobarometer at Kīlauea. *Geochem. Persp. Lett.* 29, 1–8.
390 <https://doi.org/10.7185/geochemlet.2404>
- 391 Fall, A., Tattitch, B., Bodnar, R.J., 2011. Combined microthermometric and Raman spectroscopic
392 technique to determine the salinity of H₂O–CO₂–NaCl fluid inclusions based on clathrate
393 melting. *Geochimica et Cosmochimica Acta* 75, 951–964.
394 <https://doi.org/10.1016/j.gca.2010.11.021>
- 395 Georgiev, D., Pedersen, S.V., Xie, R., Fernández-Galiana, Á., Stevens, M.M., Barahona, M.,
396 2024. RamanSPy: An Open-Source Python Package for Integrative Raman Spectroscopy
397 Data Analysis. *Anal. Chem.* 96, 8492–8500.
398 <https://doi.org/10.1021/acs.analchem.4c00383>
- 399 Gleeson, M., Wieser, P.E., DeVitre, C.L., Shi, S.C., Millet, M.-A., Muir, D.D., Stock, M.J.,
400 Lissenberg, J., 2025. Persistent High-Pressure Magma Storage beneath a Near-Ridge
401 Ocean Island Volcano (Isla Floreana, Galápagos). *Journal of Petrology* 66, egaf031.
402 <https://doi.org/10.1093/petrology/egaf031>
- 403 Hagiwara, Y., Torimoto, J., Yamamoto, J., 2020. Pressure measurement and detection of small H₂
404 O amounts in high-pressure H₂O–CO₂ fluid up to 141 MPa using Fermi diad splits and
405 bandwidths of CO₂. *J Raman Spectroscopy* 51, 1003–1018.
406 <https://doi.org/10.1002/jrs.5865>
- 407 Hagiwara, Y., Yokokura, L., Yamamoto, J., 2023. Unlocking ultimate precision of intensity and area
408 ratio measurements in Raman spectroscopy: Insights from simulation, experimentation,
409 and theory and implications for isotope ratio analysis. *J Raman Spectroscopy* 54, 1440–
410 1464. <https://doi.org/10.1002/jrs.6594>
- 411 Hagiwara, Y., Yoshida, K., Yoneda, A., Torimoto, J., Yamamoto, J., 2021. Experimental variable
412 effects on laser heating of inclusions during Raman spectroscopic analysis. *Chemical*
413 *Geology* 559, 119928. <https://doi.org/10.1016/j.chemgeo.2020.119928>
- 414 Kawakami, Y., Yamamoto, J., Kagi, H., 2003. Micro-Raman Densimeter for CO₂ Inclusions in
415 Mantle-Derived Minerals. *Appl Spectrosc* 57, 1333–1339.
416 <https://doi.org/10.1366/000370203322554473>
- 417 Lamadrid, H.M., Moore, L.R., Moncada, D., Rimstidt, J.D., Burruss, R.C., Bodnar, R.J., 2017.
418 Reassessment of the Raman CO₂ densimeter. *Chemical Geology* 450, 210–222.
419 <https://doi.org/10.1016/j.chemgeo.2016.12.034>
- 420 Le, V.-H., Caumon, M.-C., Tarantola, A., Randi, A., Robert, P., Mullis, J., 2020. Calibration data
421 for simultaneous determination of P-V-X properties of binary and ternary CO₂ - CH₄ - N₂
422 gas mixtures by Raman spectroscopy over 5–600 bar: Application to natural fluid
423 inclusions. *Chemical Geology* 552, 119783.
424 <https://doi.org/10.1016/j.chemgeo.2020.119783>
- 425 Le, V.-H., Caumon, M.-C., Tarantola, A., Randi, A., Robert, P., Mullis, J., 2019. Quantitative
426 Measurements of Composition, Pressure, and Density of Microvolumes of CO₂ –N₂ Gas

427 Mixtures by Raman Spectroscopy. *Anal. Chem.* 91, 14359–14367.
428 <https://doi.org/10.1021/acs.analchem.9b02803>

429 Rosso, K.M., Bodnar, R.J., 1995. Microthermometric and Raman spectroscopic detection limits
430 of CO₂ in fluid inclusions and the Raman spectroscopic characterization of CO₂.
431 *Geochimica et Cosmochimica Acta* 59, 3961–3975. [https://doi.org/10.1016/0016-](https://doi.org/10.1016/0016-7037(95)94441-H)
432 [7037\(95\)94441-H](https://doi.org/10.1016/0016-7037(95)94441-H)

433 Song, Y., Chou, I., Hu, W., Robert, B., Lu, W., 2009. CO₂ Density-Raman Shift Relation Derived
434 from Synthetic Inclusions in Fused Silica Capillaries and Its Application. *Acta Geologica*
435 *Sinica - English Edition* 83, 932–938. <https://doi.org/10.1111/j.1755-6724.2009.00090.x>

436 Sublett, D.M., Sendula, E., Lamadrid, H., Steele-MacInnis, M., Spiekermann, G., Burruss, R.C.,
437 Bodnar, R.J., 2020. Shift in the Raman symmetric stretching band of N₂, CO₂, and CH₄
438 as a function of temperature, pressure, and density. *J Raman Spectroscopy* 51, 555–568.
439 <https://doi.org/10.1002/jrs.5805>

440 Wang, W., Caumon, M.-C., Tarantola, A., Pironon, J., Lu, W., Huang, Y., 2019. Raman
441 spectroscopic densimeter for pure CO₂ and CO₂-H₂O-NaCl fluid systems over a wide P-
442 T range up to 360 °C and 50 MPa. *Chemical Geology* 528, 119281.
443 <https://doi.org/10.1016/j.chemgeo.2019.119281>

444 Wang, X., Chou, I.-M., Hu, W., Burruss, R.C., Sun, Q., Song, Y., 2011. Raman spectroscopic
445 measurements of CO₂ density: Experimental calibration with high-pressure optical cell
446 (HPOC) and fused silica capillary capsule (FSCC) with application to fluid inclusion
447 observations. *Geochimica et Cosmochimica Acta* 75, 4080–4093.
448 <https://doi.org/10.1016/j.gca.2011.04.028>

449 Wieser, P.E., Edmonds, M., MacLennan, J., Jenner, F.E., Kunz, B.E., 2019. Crystal Scavenging
450 from Mush Piles Recorded by Melt Inclusions. *Nature Communications*.

451 Wieser, P.E., Lamadrid, H., MacLennan, J., Edmonds, M., Matthews, S., Iacovino, K., Jenner, F.E.,
452 Gansecki, C., Trusdell, F., Lee, R.L., Ilyinskaya, E., 2021. Reconstructing Magma Storage
453 Depths for the 2018 Kīlauean Eruption From Melt Inclusion CO₂ Contents: The
454 Importance of Vapor Bubbles. *Geochem Geophys Geosyst* 22.
455 <https://doi.org/10.1029/2020GC009364>

456 Yamamoto, J., Kagi, H., 2006. Extended Micro-Raman Densimeter for CO₂ Applicable to Mantle-
457 originated Fluid Inclusions. *Chem. Lett.* 35, 610–611. <https://doi.org/10.1246/cl.2006.610>

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